Ca-Sr fractionation between zoisite, lawsonite, and aqueous fluids: An experimental study at 2.0 and 4.0 GPa/400 to 800 °C

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ABSTRACT

The Ca-Sr fractionation between zoisite and, respectively, lawsonite and an aqueous fluid has been determined by synthesis experiments in the presence of a 1 M (Ca,Sr)Cl₂ aqueous fluid at 2.0 GPa/550, 600, and 700 °C and 4.0 GPa/800 °C for zoisite and 2.0 GPa/400 °C and 4.0 GPa/600 °C for lawsonite. Solid run products were characterized by EMP. SEM. and XRD with Rietveld refinement and fluids were analyzed by ICP-OES. Zoisite exhibits notable intracrystalline Ca-Sr fractionation between the A1 and A2 sites and calculated intracrystalline exchange coefficients $K_{D(Srca)}^{A1-A2} = 1.5$ to 26 show strong preference of Sr over Ca for the slightly larger A2 site. Calculated individual site-dependent zoisite/ aqueous fluid (af, in superscripts)-exchange coefficients for the studied 1 M (Ca,Sr)Cl₂ aqueous fluids are $K_{(\text{Sr-Ca})}^{\text{zo A1-af}} = 3.38$ to 41.08 for the A1 site and $K_{(\text{Sr-Ca})}^{\text{zo A2-af}} = 0.45$ to 6.51 for the A2 site. Assuming $\gamma_{\text{Ca}}^{\text{af}} = \gamma_{\text{Sr}}^{\text{af}}$ and a symmetric mixing model, the thermodynamic evaluation of the site-dependent exchange reactions $Ca^{2+(af)} + Sr^{A1}(M^{2+})^{A2}Al_3[Si_3O_{11}(O/OH)] = Sr^{2+(af)} + Ca^{A1}(M^{2+})^{A2}Al_3[Si_3O_{11}(O/OH)]$ and $Ca^{2+(af)} + Ca^{A1}(M^{2+})^{A2}Al_3[Si_3O_{11}(O/OH)]$ $(M^{2+})^{A_1}Sr^{A_2}Al_3[Si_3O_{11}(O/OH)] = Sr^{2+(af)} + (M^{2+})^{A_1}Ca^{A_2}Al_3[Si_3O_{11}(O/OH)]$ yields $\Delta\mu^0 = -29$ kJ/mol and $W_{\text{SeCa}}^{zo} = 5.5 \text{ kJ/mol for the A1 site and } \Delta \mu^0 = -1.1 \text{ kJ/mol and } W_{\text{SeCa}}^{zo} = 0 \text{ kJ/mol for the A2 site at } P$ and T of the experiments. The data indicates ideal Ca-Sr substitution on the A2 site. Lawsonite formed in both the orthorhombic *Cmcm* and the monoclinic $P2_1/m$ form. Calculated lawsonite-aqueous fluidexchange coefficients indicate overall preference of Ca over Sr in the solid and are $K_{DavSr(ca)}^{DavSr(ca)} = 1.12$ to 11.32 for orthorhombic and $K_{D(Sr-Ca)}^{law P21m-af} = 1.67$ to 4.34 for monoclinic lawsonite. Thermodynamic evaluation of the exchange reaction $Ca^{2+(af)} + SrAl_2Si_2O_7(OH)_2 \cdot H_2O = Sr^{2+(af)} + CaAl_2Si_2O_7(OH)_2 \cdot H_2O$ assuming $\gamma_{ca}^{a} = \gamma_{sr}^{a}$ and a symmetric mixing model yields similar values of $\Delta \mu^0 = -9$ kJ/mol and W_{sr-ca}^{law} $C_{mcm} = 10 \text{ kJ/mol}$ for orthorhombic and $\Delta \mu^0 = -10 \text{ kJ/mol}$ and $W_{1avP21/m}^{1avP21/m} = 11 \text{ kJ/mol}$ for monoclinic lawsonite. Calculated Nernst distribution coefficients for the studied 1 M (Ca,Sr)Cl₂ aqueous fluids are $D_{\text{Sr}}^{\text{so-af}} = 2.8 \pm 0.7$ for zoisite at 2 GPa/600 °C and $D_{\text{Sr}}^{\text{law Cmcm-af}} = 0.6 \pm 0.2$ for orthorhombic lawsonite at 4 GPa/600 °C and show Sr to be compatible in zoisite but incompatible in lawsonite. This opposite mineral-aqueous fluid-fractionation behavior of Sr with respect to zoisite and lawsonite on the one hand and the ideal Ca-Sr substitution on the zoisite A2 site in combination with the strong intracrystalline Ca-Sr fractionation in zoisite on the other hand, make Sr a potential tracer for fluid-rock interactions in zoisite- and lawsonite-bearing rocks. For low Sr-concentrations, x_{Sr}^{so} directly reflects x_{sf}^{sf} and allows us to calculate Sr-concentrations in a metamorphic aqueous fluid. During high-pressure aqueous fluid-rock interactions in subduction zone settings the opposite mineral-aqueous fluid-fractionation behavior of Sr results in different aqueous fluid characteristics for lawsonite- vs. zoisite-bearing rocks. Ultimately, subduction zone magmas may trace these different aqueous fluid characteristics and allow distinguishing between cold, lawsonite-bearing vs. warm, zoisite-bearing thermal regimes of the underlying subduction zone.

Keywords: Ca-Sr fractionation, zoisite, lawsonite, high-pressure fluids, experimental petrology, subduction zone