

Crystal-chemical controls on trace element partitioning between garnet and anhydrous silicate melt

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ABSTRACT

We performed experiments at 3.0 GPa and 1530–1565 °C to investigate the effects of crystal composition on trace element partitioning between garnet and anhydrous silicate melt. Bulk compositions along the pyrope (Py: Mg₃Al₂Si₃O₁₂)-grossular (Gr: Ca₃Al₂Si₃O₁₂) join, doped with a suite of trace elements (Li, B, K, Sc, Ti, Sr, Y, Zr, Nb, Cd, In, REE, Hf, Ta, Th, and U) produced homogeneous garnets, ranging in composition from Py₈₄Gr₁₆ to Py₉Gr₉₁, in equilibrium with melt.

Trace element partition coefficients (*D*-values), measured by SIMS, depend greatly on the Mg/(Mg + Ca) of garnet. For example, from Py₈₄ to Py₉, *D*_{La} increases from 0.004 to 0.2, whereas *D*_U increases from 0.029 to 0.42. These variations can be explained by the lattice strain model of Blundy and Wood (1994), which describes trace element partitioning of an element *i* in terms of the ionic radius of *i* (*r_i*), the size of the lattice site on which *i* partitions (*r₀*), the Young's modulus of the site (*E*), and the (theoretical) partition coefficient *D*₀ for an ion of radius *r₀*.

For trivalent cations substituting in the garnet X-site (Y, REE, Sc, and In), apparent values of *r₀* fitted to our data vary systematically from 0.935 ± 0.004 Å (Py₈₄) to 0.99 ± 0.01 Å (Py₉), a trend consistent with variations in the size of the X-site. Values of *D*₀ show an increase from Py₉ (*D*₀ = 2.8 ± 0.1) to Py₈₄ (4.8 ± 0.1) and Young's modulus *E* varies from 257 ± 20 GPa for Py₆₀ to 590 ± 40 GPa for Py₈₄. These results allow a quantitative assessment of the influence of crystal chemistry on garnet-melt *D*-values, thereby forming the basis for a predictive model similar to that recently developed for clinopyroxene-melt partitioning by Wood and Blundy (1997). Our new data emphasize the importance of taking into account crystal composition when modeling trace element behavior in natural systems.