American Mineralogist, Volume 89, pages 1597–1609, 2004

## A Mössbauer study of the oxidation state of Fe in silicate melts

## KASTHURI D. JAYASURIYA,<sup>1,\*</sup> HUGH ST.C. O'NEILL,<sup>2,†</sup> ANDREW J. BERRY,<sup>2</sup> AND STEWART J. CAMPBELL<sup>1</sup>

<sup>1</sup>School of Physical, Environmental and Mathematical Sciences, University of New South Wales, Australian Defence Force Academy, Canberra, ACT 2600, Australia

<sup>2</sup>Research School of Earth Sciences, Australian National University, Canberra, ACT 0200, Australia

## ABSTRACT

 $Fe^{3+}/\Sigma Fe$  ratios were determined from Mössbauer spectra recorded for a series of 17 anorthitediopside eutectic glasses containing 1 wt% <sup>57</sup>Fe<sub>2</sub>O<sub>3</sub> quenched from melts equilibrated over a range of oxygen fugacities from  $f_{O_2} \sim 10^5$  bars (Fe<sup>3+</sup>/ $\Sigma$ Fe = 1) to  $10^{-13}$  bars (Fe<sup>3+</sup>/ $\Sigma$ Fe = 0) at 1682 K. Fe<sup>3+</sup>/Fe<sup>2+</sup> was found to be proportional to  $f_{0_2}$  to the power of 0.245 ± 0.004, in excellent agreement with the theoretical value of 0.25 expected from the stoichiometry of the reaction  $Fe^{2+}O + 0.25 O_2 = Fe^{3+}O_{1,5}$ . The uncertainty in the Fe<sup>3+</sup>/ $\Sigma$ Fe ratios determined by Mössbauer spectroscopy was estimated as  $\pm$ 0.01 (1 $\sigma$ ) from the fit of the data to the theoretical expression, which is significantly less than that quoted for previous measurements on silicate glasses; this results from fitting the spectra of a large number of systematically varying samples, which allows many of the ambiguities associated with the fitting procedure to be minimized.  $Fe^{3+}/\Sigma Fe$  ratios were then determined for samples of the anorthitediopside eutectic composition equilibrated at selected values of  $f_{0.3}$ , to which up to 30 wt% Fe<sub>2</sub>O<sub>3</sub> had been added. Fe<sup>3+</sup>/ $\Sigma$ Fe was found to vary with  $\Sigma$ Fe (or FeO<sub>T</sub>), but both the 1 wt% and high FeO<sub>T</sub> data could be satisfactorily fit assuming the ideal stoichiometry (i.e.,  $Fe^{3+}/Fe^{2+} \propto f_{O_2}^{1/4}$ ) by the inclusion of a Margules term describing Fe<sup>2+</sup>-Fe<sup>3+</sup> interactions. The large negative value of this term indicates a tendency toward the formation of Fe<sup>2+</sup>-Fe<sup>3+</sup> complexes in the melt. The resulting expression, using the ideal exponent of 0.25, gave a fit to 289 Fe<sup>3+</sup>/ $\Sigma$ Fe values, compiled from various literature sources, of similar quality as previous empirical models which found an exponent of  $\sim 0.20$ . Although the empirical models reproduce  $Fe^{3+}/\Sigma Fe$  values of glasses with high  $FeO_T$  reasonably well, they describe the data for 1 wt% FeO<sub>T</sub> poorly. The non-ideal values of the exponent describing the dependence of  $Fe^{3+}/\Sigma Fe$  on  $f_{O_2}$  at high FeO<sub>T</sub> are an artifact of models that did not include a term explicitly to describe the  $Fe^{2+}-Fe^{3+}$  interactions. An alternative model in which Fe in the silicate melt is described in terms of three species,  $Fe^{2+}O$ ,  $Fe^{3+}O_{1,5}$ , and the non-integral valence species  $Fe^{2.6+}O_{1,3}$ , was also tested with promising results. However, at present there is no model that fits the data within the assessed accuracy of the experimental measurements.