## Biotite dehydration, partial melting, and fluid composition: Experiments in the system KAlO<sub>2</sub>-FeO-MgO-SiO<sub>2</sub>-H<sub>2</sub>O-CO<sub>2</sub>

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## ABSTRACT

Biotite solid-solutions are a significant H<sub>2</sub>O reservoir in the lithosphere, and the assemblage Bt+Opx+Kfs+Qtz is commonly used to estimate  $a_{H_0}$  in high-grade metamorphic and magmatic rocks. Here we report experimental constraints on subsolidus mineral equilibria involving biotite and orthopyroxene in the system KAIO<sub>2</sub>-MgO-FeO-SiO<sub>2</sub>-H<sub>2</sub>O-CO<sub>2</sub>. Our experiments address the question of stability of biotite of a given  $X_{\rm Fe}$  in the assemblage Bt+Qtz±Sa, or the stability of the assemblage Opx+Sa±Qtz. Clemens (1993) and Clemens et al. (1997) concluded that  $CO_2$  has no effect other than to lower  $a_{HO}$  and thereby raise the solidus T. Our data at  $X_{H,O}^{\text{H}} < 1$  extend these conclusions to encompass Fe-bearing systems more similar to natural rocks. From a comparison of experimental data and calculated isopleths of biotite composition in the divariant assemblage Bt+Opx+Kfs+Qtz+fluid, it appears that phlogopite-annite solid-solutions must be significantly non-ideal (at T < T800 °C) or that enstatite-ferrosilite solid-solutions must have negative values for their Margules-type parameters. Ignoring these factors would result in any calculated  $a_{H,O}$  values being too low. Although various models allow us to estimate  $X_{H,O}$  in H<sub>2</sub>O-CO<sub>2</sub> fluids, we are still unable to use biotite equilibria to estimate  $a_{H,O}$  accurately during high-grade metamorphism and magma crystallization. We also consider qualitatively the effects Fe-Mg biotite solid-solution on partial melting equilibria in fluid-poor (rock-dominated) systems in which hydration-dehydration reactions control the fluid composition.