REVIEW

Pathways for nitrogen cycling in Earth's crust and upper mantle: A review and new results for microporous beryl and cordierite

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ABSTRACT



Earth's atmosphere contains 27-30% of the planet's nitrogen and recent estimates are that about one-half that amount (11-16%) is located in the continental and oceanic crust combined. The percentage of N in the mantle is more difficult to estimate, but it is thought to be near 60%, at very low concentrations. Knowledge of the behavior of N in various fluid-melt-rock settings is key to understanding pathways for its transfer among the major solid Earth reservoirs.

Nitrogen initially bound into various organic materials is transferred into silicate minerals during burial and metamorphism, often

as NH⁴ substituting for K⁺ in layer silicates (clays and micas) and feldspars. Low-grade metamorphic rocks appear to retain much of this initial organic N signature, in both concentrations and isotopic compositions, thus in some cases providing a relatively un- or little-modified record of ancient biogeochemical cycling. Devolatilization can release significant fractions of the N initially fixed in crustal rocks through organic diagenesis, during progressive metamorphism at temperatures of ~350–550 °C (depending on pressure). Loss of fractionated N during devolatilization can impart an appreciable isotopic signature on the residual rocks, producing shifts in δ^{15} N values mostly in the range of +2 to +5‰. These rocks then retain large fractions of the remaining N largely as NH⁴₄, despite further heating and ultimately partial melting, with little additional change in δ^{15} N. This retention leads to the storage of relatively large amounts of N, largely as NH⁴₄, in the continental crust. Nitrogen can serve as a tracer of the mobility of organic-sedimentary components into and within the upper mantle.

This contribution focuses on our growing, but still fragmentary, knowledge of the N pathways into shallow to deep continental crustal settings and the upper mantle. We discuss the factors controlling the return of deeply subducted N to shallower reservoirs, including the atmosphere, via metamorphic devolatilization and arc magmatism. We discuss observations from natural rock suites providing tests of calculated mineral-fluid fractionation factors for N. Building on our discussion of N behavior in continental crust, we present new measurements on the N concentrations and isotopic compositions of microporous beryl and cordierite from medium- and high-grade metamorphic rocks and pegmatites, both phases containing molecular N_2 , and NH_4^4 -bearing micas coexisting with them. We suggest some avenues of investigation that could be particularly fruitful toward obtaining a better understanding of the key N reservoirs and the more important pathways for N cycling in the solid Earth.

Keywords: Nitrogen cycling, nitrogen isotopes, ammonium, microporous silicate, isotope fractionation, layer silicates, cordierite, Review article, Invited Centennial article

	Beryl	Beryl	Muscovite	Muscovite		Beryl	Beryl
Sample	$\delta^{15}N_{air}$	N (ppm)	$\delta^{15}N_{air}$	N (ppm)	$\Delta^{15}N_{mica-beryl}$	$\delta^{15}C_{VPDB}$	C (ppm)
80192	7.9	17	9.4	80	1.5		31
23215	5.9	11	8.1	273	2.2	-8.47	48
40597	4.0	25	8.3	305	4.3		
1	3.8	18	6.7	41	2.9	28	-
80145	5.1	39	8.7	632	3.6		
	Cordierite	Cordierite	Biotite	Biotite		Cordierite	Cordierite
Sample	δ ¹⁵ N _{air}	N (ppm)	δ ¹⁵ N _{air}	N (ppm)	$\Delta^{\rm 15} N_{\rm mica-crd}$	δ ¹⁵ C _{VPDB}	C (ppm)
	2 · ·aii		Pegmatite/pa				- (- [)
88593	12.0	5	8.6	134	-3.4		80
80537	10.5	6	11.0	103	0.5		31
G-155a	5.1	17				-10.5	16
C006	9.0	29				-6.9	936
TUB-1	4.3	33				-13.6	195
26230	7.5	38	1.8	70	-5.7	1510	
C004	11.4	60	110		517	-8.8	600
84264	7.4	67				-14.0	327
04204	7.4	07	Mid and a second	· · · · · · · · · · · · · · · · · · ·		14.0	527
Wards	16.0	0	Mid-grade me	tamorphics		6.5	245
25 Geco Mine	16.9 10.4	8 19				-6.5	345
WYO-2	5.1	30				-10.8	590
	6.5	30 95					590 694
118171	0.5	95				-8.4	694
			Granulite	facies			
X-1	4.7	41				-4.0	277
42/IA	0.9	55				-12.0	543
CL-177-1	30.0	71				-36.4	1200
TA-5	7.0	101				-13.0	221
129875	10.4	104					
13	4.8	162	7.8	116	3.0		
26539	9.9	232	0.5	65	-9.4	-9.4	1099
VS-1	5.9	273				-6.7	1039
7114	2.9	634				-22.3	991
S. India 1	3.1	923				-16.3	408
89 V	3.6	1342					
NE86A-24b	8.6	4525				-16.5	445
			Uncatego	orized			
CTSiM	9.0	27				-6.7	976
10398	5.5	56	5.2	48	-0.3		
43090	8.3	89	2.7	86	-5.5	-8.3	820
33294	7.5	154	4.4	25	-3.1	-9.3	614
H06	2.0	446		20	5	2.0	
106886	3.9	1457				-6.1	623

TABLE 2. Isotopic data for beryl and cordierite (and coexisting micas)

Nitrogen partitioning behavior between coexisting cordierite and biotite

Biotite in pegmatites has N concentrations of 70 to 134 ppm higher than those of the coexisting cordierites, but to varying degrees (see Fig. 9). In two of the three experiments, the δ^{15} N values (+1.8 to +11‰) of the biotites are lower than the values for cordierite from the same rock sample (Fig. 9). In the third case, the δ^{15} N values for the coexisting minerals are similar. For two granulite facies rocks, biotite has N concentrations of 65 and 116 ppm and they are lower than those of the coexisting cordierites with concentrations of 162 and 232 ppm N, respectively. For three rock samples whose origins are uncertain ("uncategorized"), cordierites have N concentrations and δ^{15} N values higher than those of coexisting biotite. Thus, the partitioning and isotopic fractionation results for N in coexisting biotite and cordierite do not show any consistent behavior and, thus, factors other than temperature must be considered.

The wide range of measured N concentrations and δ^{15} N values in cordierite could be attributed to differences in the temperature of crystallization, compositional heterogeneity in the protoliths (see the compilation of data for sediments in Kerrich et al. 2006), and differing magnitudes of positive isotopic shifts

in δ^{15} N resulting from lower-grade devolatilization (see Bebout and Fogel 1992; Jia 2006; Palya et al. 2011). Three additional effects that could explain the range of $\delta^{15}N$ values are: (1) the presence of some NH4 in cordierite, which could result in decreased or no fractionation of N with NH⁴ in coexisting biotite; (2) chemical disequilibrium between biotite and cordierite that could affect the N concentrations and isotopic compositions; or (3) post-crystallization or retrograde modification of the $\delta^{15}N$ values due to differential diffusive loss (i.e., preferential loss of ¹⁴N). Possibility 3 must be addressed in future studies of microporous silicates and fluid-rock processes (see discussion of diffusive loss for H₂O and CO₂ in cordierite by Vry et al. 1990). Regarding the first effect, we consider the presence of significant amounts of NH₄⁺ in cordierite as unlikely for crystal chemical reasons (the concentration of K, with very few exceptions, in cordierite is very low).

Nitrogen partitioning behavior between cordierite and its rock matrix for a medium-grade schist

Nitrogen concentrations and $\delta^{15}N$ values were measured for a gem-quality cordierite and its muscovite-rich matrix for one medium-grade metasedimentary schist from Connecticut, U.S.A. (Fig. 8b). Here, quite interestingly, cordierite showed no measurable N within the experimental detection limits, whereas the muscovite-rich matrix contained 350 ppm N. Essentially all N in the schist resides as NH⁴ in the micaceous matrix. This result is consistent with the observation that N concentrations in cordierite are highest at the highest metamorphic grades where muscovite is not present. The interpretation is that N residing in the muscovite, after its breakdown with increasing temperature, is taken up in K-feldspar and cordierite (see Palya et al. 2011).

Nitrogen partitioning behavior between coexisting beryl and muscovite

Figure 10 shows the N concentrations and δ^{15} N values for five beryl-muscovite pairs taken from four pegmatites and one metasedimentary schist (see the data in Table 2). These results show that muscovite always contains far greater amounts of N than coexisting beryl. A mean $\Delta^{15}N_{musc-beryl}$ ($\delta^{15}N_{musc} - \delta^{15}N_{beryl}$) value of +2.9 ($1\sigma = 1.1\%$) was obtained for these pairs. The direction and magnitude of this isotopic fractionation are similar to those measured for bioite-fluid inclusion pairs in a vein in metasedimentary rocks from Bastogne, Belgium (see discussion above) and predicted by the fractionation factors calculated using spectroscopic data (see Fig. 2a). Apparently, based on the limited data, the partitioning behavior of N and its isotopes between beryl and muscovite, compared to the case for cordierite-biotite pairs, is more systematic (Fig. 9). This could, in part, reflect (1) the presence of N₂ as the single N species in the beryl (i.e., not also NH₄⁺) or (2) the more rapid cooling of pegmatites compared to most granulites. Rapid cooling would allow greater retention of N₂ incorporated during peak crystallization conditions and thus better preservation of the peak-temperature partitioning behavior.

Carbon concentrations and δ^{13} C values of cordierite and beryl

Figure 11a shows a plot of N and C concentrations for various cordierites and a single beryl sample (see also Table 2). The two elements show a rough positive correlation. Figure 11b shows a plot of C concentrations and $\delta^{13}C_{\text{VPDB}}$ values. The various samples have roughly similar $\delta^{13}C$ values with a mean = -9.8‰ ($1\sigma = 3.5\%$), with the exception of two cordierites outliers from granulites having $\delta^{13}C$ values of -36.4 and -22.3‰. For the range of rocks types studied here, pegmatite cordierites tend to have lower C concentrations. Cordierites from granulites and





Muscovite

 Beryl

FIGURE 10. Nitrogen isotopic compositions and concentrations of coexisting beryl and muscovite from pegmatites and metamorphic schist (Table 2 and Appendix A¹; Lazzeri 2012). The inset photograph of beryl in muscovite is used courtesy of Rob Lavinsky, www.iRocks.com (for photo and specimer; horizontal dimension ~8 cm). (Color online.)

FIGURE 11. C and N concentrations and δ^{13} C values for the cordierites and beryl analyzed in this study. (**a**) N vs. C concentrations. Cordierites from the granulite facies tend to show the highest N concentrations and those from pegmatites the lowest, as is also the case for beryl. (**b**) Carbon concentration vs. δ^{13} C showing the relatively narrow range of isotopic compositions and the wide range in C concentrations.

medium-grade metapelites (and the uncategorized cordierites) have, in general, similar concentrations of C, but the most Crich samples (i.e., 990 to 1200 ppm C) come from granulites (there is considerable overlap among the data for medium grade, granulite, and uncategorized cordierites). These observations for C concentrations and δ^{13} C values are consistent with those made by Vry et al. (1990) in their isotopic investigation of cordierite. Beryl, unlike cordierite, contains very little C, with only one sample containing amounts sufficient for an analysis of δ^{13} C.

How important are microporous silicates and tourmaline for storage of nitrogen in continental crust?

Cordierite can be a volumetrically significant rock-forming mineral in metapelitic rocks. It is stable over a wide range of temperatures and at low to moderate pressures corresponding to upper to middle levels of the continental crust (e.g., Schreyer 1965; Kalt et al. 1998; White et al. 2003; Palva et al. 2011). Tourmaline, which may also contain appreciable amounts of N (e.g., Wunder et al. 2015), occurs in metapelitic rocks over a broader range of P and T, even in UHP rocks that experienced pressures of up to ~3.0 GPa (Bebout and Nakamura 2003; see the summary by Marschall et al. 2009; van Hinsberg et al. 2011). However, its modal abundance is limited in both metamorphic and igneous rocks, therefore it is unlikely to play a role as a major sink in the crust and for the deep-Earth N cycle. Tourmaline and beryl in larger amounts can occur in pegmatites (London and Evensen 2002), but pegmatites are unlikely to act as significant N reservoirs due to their relative scarcity. Summarizing, cordierite could be a notable sink for N in shallow- to mid-levels of the continental crust.

CONCLUSIONS AND OUTLOOK

Fluxes of N among the oceanic and continental crust, mantle, oceans, and atmosphere largely determine the abundance and the isotopic composition of N in all of these reservoirs. Models of modern and ancient volatile cycling on Earth are highly dependent on understanding the nature of these fluxes (Javoy 1997; Tolstikihn and Marty 1998; Zhang and Zindler 1993). Biological processes play a key role in affecting the concentrations and behavior of N in the solid Earth. Nitrogen in the oceans and atmosphere can be incorporated (via biological processes) into mineral phases, some of which are carried into the deep Earth (via burial and subduction). Nitrogen can thus be an effective tracer in the study of the transfer of sedimentary and organic components into and within the crust and upper mantle. In this article, we present some important observations regarding this hydrosphere-crust-upper mantle transfer, based on the studies to date, and we suggest several areas needing attention.

 NH⁴₄ can replace K⁺ via a solid solution mechanism in some K-bearing rock-forming silicate minerals, especially layer silicates (clays and micas) and feldspar. This substitution is so prevalent that an estimate of N subduction-input fluxes can be based on knowledge of the rates of K subduction (see Busigny et al. 2003, 2011; Busigny and Bebout 2013). Recent research suggests that cordierite and tournaline can also serve as reservoirs for N, with concentrations roughly similar to those of coexisting micas. Cordierite could be a significant phase for the storage of molecular N₂ in shallowto mid-levels of the continental crust.

- Low-temperature devolatilization of organic matter, and the concomitant crystallization of clay minerals such as illite, permit the retention of this initially organic N as NH[‡], apparently with little isotopic fractionation. Although whole-rock C/N ratios of very low-grade metamorphosed sediments can retain biogeochemical information, kerogen itself (the reduced C reservoir) contains very little N. Most of the initially organic N is transferred into and housed in clays and low-grade metamorphic micas (and in some cases, authigenic feldspars; Svensen et al. 2008).
- Considerable loss of N from minerals to fluids can occur at low to medium metamorphic grades, depending upon the prograde *P*-*T* path the rocks experience (see Bebout and Fogel 1992; Busigny et al. 2013; Li and Keppler 2014). In many cases, the isotopic shifts associated with this loss point to a N₂-NH⁺₄ exchange mechanism. Li et al. (2009) proposed, however, the possibility of kinetically controlled NH₃-NH⁺₄ exchange. In relatively "cool" subduction zone settings, sedimentary rocks can retain a large fraction of their original N contents to great depths, perhaps even the depths beneath volcanic fronts (see Busigny et al. 2003; Bebout et al. 2013a). Experimental phase equilibrium studies document the stability of mica (e.g., phengite), the key N mineral reservoir, to great depths in most subduction zones (Schmidt and Poli 2014).
- Further investigation of the concentrations, isotope compositions, and fluxes of N into and within the continental crust is badly needed (see the discussion by Johnson and Goldblatt 2015). The estimated average concentration of 56 ppm for this reservoir is based on a very small number of analyses (see Wedepohl 1995; Bach et al. 1999; Palya et al. 2011; Rudnick and Gao 2014). It goes without saying that further work on the concentrations and isotopic compositions of N in the mantle is needed, as the mantle could contain ~60% of the Earth's N (Table 1; see the discussions by Cartigny and Marty 2013; cf. Johnson and Goldblatt 2015).
- The rates and mechanisms by which N diffuses in key minerals such as the micas, alkali feldspars, clinopyroxenes, and microporous silicates are poorly understood (see Watson and Cherniak 2014). Closure temperatures for the retention of N in these phases are not known, complicating the assessment of N isotope behavior at high geologic temperatures.
- All research done thus far on N in silicate systems has been on either whole-rock samples or mineral separates. It will be important to develop microanalytical methods for analyzing N concentrations and δ^{15} N at scales allowing consideration of intramineral heterogeneity (see the first analyses on cordierite using the ion microprobe by Hervig et al. 2014).
- Finally, we stress the need for experimental calibration of N isotope fractionation in silicate fluid-mineral systems. This is required to understand and resolve the differences among the various calculated fractionation factors.

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