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3 **The use of boron nitride to impose reduced redox conditions in experimental petrology**

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12
13 **Abstract**

14 BN is a commonly used pressure transmitting material in experimental petrology. It is often
15 considered to be as inert as MgO or Al₂O₃, and its redox potential is seldomly discussed. It is
16 generally implied that, when used as a capsule sleeve, BN may impose relatively reduced
17 conditions, similar to the effect of the fayalite-magnetite-quartz (FMQ) buffer. However,
18 sediment melting experiments performed at 1050°C and 3 GPa with BN as the capsule sleeve,
19 produced a hydrous rhyolitic melt with dissolved H₂S and CH₄ (Li et al. 2021). The resulting *f*O₂
20 estimate is significantly more reduced than that for the magnetite-wustite (MW)-buffered
21 experiment where H₂S and CH₄ were undetected (Li et al. 2021), possibly to the extent of the
22 quartz-iron-fayalite (QIF) buffered conditions produced when BN is used as a capsule or crucible
23 (Wendlandt et al. 1982). In order to establish an explanation for such a discrepancy, we have

24 conducted further investigation to better constrain the fO_2 imposed by BN, when used as a
25 capsule sleeve. Here we report results on analyses of Fe content in Au capsules, a comparative
26 experiment using a QIF buffer and an experiment with an Fe-(Mg,Fe)O sensor for direct analysis
27 of fO_2 . The calibration of the equilibrium between FeO in melt and Fe in the Au capsule, from
28 Ratajeski and Sisson (1999) appears to be inadequate in constraining fO_2 for our experiments.
29 However, we were able to obtain Fe diffusion coefficients in Au from the Fe diffusion profiles
30 observed in the capsule of the Fe-(Mg,Fe)O sensor experiment, and both the inner and outer
31 capsules of the MW-buffered experiment, with resulting values of $1 \times 10^{-13} \text{ m}^2/\text{s}$, $3 \times 10^{-14} \text{ m}^2/\text{s}$,
32 and $5 \times 10^{-14} \text{ m}^2/\text{s}$, respectively. The QIF-buffered and Fe-(Mg,Fe)O sensor experiments provide
33 several lines of evidence supporting the observation that BN imposes QIF-like redox conditions.
34 Firstly, the Fe-(Mg,Fe)O sensor returned an fO_2 value of QIF. Secondly, the “apparent” partition
35 coefficients between FeO content in melt and Fe in the Au capsules are similar between the BN
36 experiment and the QIF-buffered experiment. Thirdly, we observe CH_4 and H_2O peaks with
37 similar intensities in the Raman spectra of melts from these two experiments, suggesting similar
38 H_2 and thus O_2 fugacity. As our experiments were performed on a cubic press with the
39 experimental assembly encased in a pyrophyllite cube, we interpret that the significantly reduced
40 conditions imposed by BN are likely due to high H_2O activity maintained by dehydration of
41 pyrophyllite, which can be explained using the reaction $2BN + 3H_2O = B_2O_3 + N_2 + 3H_2$. Lower H_2O
42 activity will reduce or inhibit the oxidation of BN and its fO_2 buffering ability. If heat-treated,
43 BN acts as a highly efficient H_2 barrier, as shown by Truckenbrodt et al. (1997). Through our
44 efforts to determine the fO_2 imposed by using BN as a capsule sleeve in our experimental
45 assembly, we are able to demonstrate the reducing ability of BN as an assembly component, and
46 furthermore shed light on the process by which BN imposes such reduced fO_2 . We hereby

47 present what we have learnt during the course of this investigation, in the hope that the effect of
48 BN on fO_2 control is both recognized and further exploited in future experimental studies.

49

50 **Keywords:**

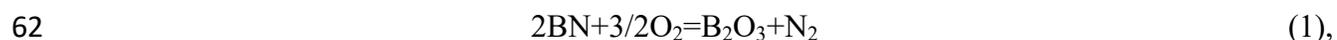
51 BN, oxygen fugacity control, Fe in Au calibration, Fe-(Mg,Fe)O sensor, experimental
52 petrology

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Introduction

55 Boron Nitride (BN) is an often-used pressure transmitting material in experimental petrology;
56 however, it has been considered to be as inert as MgO or Al₂O₃, and its redox potential has
57 seldomly been discussed. The use of BN as a capsule material or crucible for silicate systems has
58 been previously reported (Mysen 1979; Mysen and Popp 1980), but did not gain wide popularity
59 due to the fact that BN imposes highly reducing conditions; and furthermore that B₂O₃, produced
60 through oxidation of BN, acts as a fluxing agent for melting (Wendlandt et al. 1982). Based on
61 the BN oxidation reaction



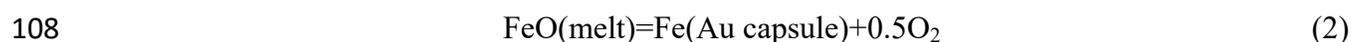
63 Wendlandt et al. (1982) estimated that the fO_2 imposed by BN capsules is ~3 log units below the
64 quartz-iron-fayalite (QIF) buffer at 1027°C and 1 bar for pure N₂ conditions; while for 1027°C
65 and 3 GPa conditions, the estimated fO_2 is close to that of the QIF buffer. Note that such a
66 difference arises from the different N₂ fugacity between 1 bar and 3 GPa. The reducing ability of
67 BN was verified by Wendlandt et al. (1982) with experiments performed using BN capsules
68 containing various starting materials. They showed that FeO and NiO were reduced to Fe and Ni

69 metal respectively; Fe metal was also produced in experiments with either olivines of various
70 compositions or a garnet lherzolite as starting materials.

71 The redox imposing character of BN initially came to our attention after a sediment melting
72 experiment performed in a previous study (Li et al. 2021; exp. LMD670) at 3 GPa, 1050°C using
73 BN as a capsule sleeve and spacer, produced a hydrous rhyolitic melt with dissolved H₂S and
74 CH₄, as shown by the Raman spectra of the quenched melt (Fig. 1). Dissolution of CH₄ in silicate
75 melt has been reported mainly in experiments performed at reduced conditions where a metal
76 alloy phase is stable (i.e., $fO_2 < IW$, IW stands for iron-wustite buffer) (Kadik et al. 2006, 2015,
77 2017; Mysen et al. 2009; Ardia et al. 2013; Dasgupta et al. 2013; Armstrong et al. 2015; Li et al.
78 2015, 2016; Dalou et al. 2019; Grewal et al. 2020). Moreover, H₂S and CH₄ were undetected in
79 the melt from the magnetite-wustite (MW)-buffered experiment in the same series with identical
80 starting composition and *PT* conditions (Fig. 1). We therefore interpret that the intrinsic fO_2
81 imposed by the BN capsule sleeve and spacer is lower than that of the MW buffer. In addition,
82 graphite was a stable phase in the MW-buffered experiment, but absent in the BN experiment,
83 also suggesting that BN imposes more reduced conditions than the MW buffer. The absence of
84 graphite in the BN experiment (LMD670) provides an upper-bound constraint on the fO_2 of the
85 experiment. The fO_2 limit for graphite stability can be calculated based on the reaction
86 $C + 2H_2O = CH_4 + O_2$, with the equilibrium constant calculated according to Ohmoto and Kerrick
87 (1977). Considering that the gas phase in equilibrium with the experimental charge should be
88 dominated by CH₄, we can assume fCH_4 as being equal to that of pure CH₄ at *PT* conditions of 3
89 GPa, 1050°C (calculated using the CHO program). In combination with the H₂O fugacity
90 estimated based on H₂O content in melt (Li et al. 2021), the calculated fO_2 is ~FMQ-5.5. During
91 the course of our previous study (Li et al. 2021), no independent assessment was attempted to

92 directly determine the fO_2 for the BN experiment. We thus assumed that the fO_2 imposed by BN
93 would be similar to that of the QIF buffer, as calculated for the BN capsule at 3 GPa, 1050°C.
94 This appeared to be a reasonable assumption considering that the data point for the BN
95 experiment fits the trend of variation of melt S solubility with fO_2 , and also agrees with the
96 model prediction of Clemente et al. (2004). On a separate note, the presence of pyrrhotite, rather
97 than Fe metal in exp. LMD670 at such reduced conditions can be attributed to its high $\log fS_2$,
98 estimated to be -2.2 (Li et al. 2021), higher than the calculated $\log fS_2$ (-2.6 at 3 GPa, 1050°C) for
99 Fe metal and FeS coexistence based on the reaction $Fe + 1/2S_2 = FeS$.

100 A search of the literature revealed that the reducing effect of BN when used as pressure
101 transmitting material has been indirectly observed before. For example, Fe loss to Au capsules
102 was reported by Ratajeski and Sisson (1999). Their experiments investigated melting of natural
103 hornblende quartz gabbro at 8 kbar, 800-975°C with an estimated bulk H₂O content of ~1.3 wt%.
104 In their NaCl-graphite experimental assembly, BN powder was packed around the Au capsule to
105 minimize oxidation of the sample. Based on the analyses of Fe content in Au capsules and FeO
106 content in melt, they obtained an expression for the apparent equilibrium constant (K') of the
107 reaction



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$$\ln K' = -22.953[1000/T(\text{K})] + 3.421 \quad (r^2 = 0.99) \quad (3)$$

110 with K' defined as

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$$K' = \frac{X_{Fe}^{\text{alloy}} \times fO_2^{0.5}}{X_{FeO}^{\text{melt}}} \quad (4)$$

112 which is in effect representing the distribution coefficient since the activity coefficients for Fe in
113 Au ($\gamma_{Fe}^{\text{alloy}}$) and FeO in melt ($\gamma_{FeO}^{\text{melt}}$) were not calculated. The fO_2 value used in the calibration
114 was assumed to be FMQ-1, which was based on the fO_2 estimation for similar experiments from

115 Patiño Douce and Beard (1994, 1995), although the latter experiments were performed without
116 BN in their experimental assembly. When Fe content in Au capsules and FeO content in melt are
117 known, fO_2 can in turn be derived from calculated K' values:

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$$\log fO_2 = 2 \times \frac{\ln K' - \ln X_{Fe}^{alloy} + \ln X_{FeO}^{melt}}{\ln(10)} \quad (5)$$

119 Sisson et al. (2005) performed hydrous (1.7-2.3 wt% H₂O) basalt melting experiments at 700
120 MPa with either BN or NaCl as capsule sleeves, and reported that sulfide was stable in
121 experiments using BN, while titanomagnetite was stable in experiments using NaCl. The fO_2 for
122 the former case was estimated to be similar to FMQ-buffered experiments, while the fO_2 for the
123 latter was estimated to be between FMQ and (Re-ReO₂)-buffered conditions. Most recently,
124 Pelletier et al. (2021) reported S contents at sulfide saturation (SCSS) in hydrous rhyolitic melts,
125 produced by melting of Ca-poor pelite at 3 GPa, 750-1000°C using a Talc-pyrex-graphite
126 assembly, but with either MgO or BN capsule sleeves. Higher SCSS values were obtained for
127 experiments using MgO than those using BN as capsule sleeves. For example, the SCSS value
128 was reported to be ~150 ppm for the experiment performed at 800°C with a bulk S content of 2.1
129 wt% when using a BN capsule sleeve, but reached ~830 ppm for the 750° C experiment where a
130 MgO capsule sleeve was used. Such a difference in melt S content reflects the higher fO_2
131 conditions imposed by MgO with respect to BN. In summary, it is clear that the use of BN in
132 these experimental studies imposed reduced conditions (\leq FMQ). However, the estimated fO_2
133 conditions were not as reduced as those we have empirically observed in our own study ($<$ MW,
134 likely equivalent to QIF).

135 To better constrain the intrinsic fO_2 of our experimental assembly using BN as a capsule
136 sleeve, new experiments have been conducted and assessed in combination with further analyses
137 of experiments from our previous work. Inspired by the work of Ratajeski and Sisson (1999), we

138 have analyzed the post-run Fe content in Au capsules from the experimental series. We also
139 report an additional experiment with identical starting composition and *PT* conditions to those
140 previously reported, but with a QIF buffer, which serves as a comparison to the BN experiment.
141 An experiment employing an Fe-(Mg,Fe)O sensor as the starting material and BN as the capsule
142 sleeve, has been performed for direct analysis of the experimental fO_2 . The application of BN in
143 fO_2 control will also be discussed.

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Methods

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Experimental methods

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A detailed description of both the experimental design and the method for fO_2 estimation are presented in Li et al. (2021). Here we briefly describe the key points. All of the experiments, both new and previously reported, have been conducted using a CS-IV 6 ×14 MN cubic press at the High Pressure and High Temperature laboratory, Peking University. The experimental assembly (BJC-11) has a cylindrical configuration, which is comprised of a BN outer sleeve, graphite heater, and a BN/MgO spacer and capsule sleeve, similar to the common assembly used in piston-cylinder experiments, the primary difference being that the cylindrical unit is further encased in a pyrophyllite cube with an edge length of 32.5 mm (Fig. 2). For experiments with buffer assemblies, a double-capsule configuration is employed with an inner capsule (2.5 mm O.D.) containing ~10 mg starting material and an outer capsule (5 mm O.D.) containing 50-60 mg of buffer material together with 4-5 mg of deionized H₂O. Both inner and outer capsules are fabricated from Au tubing. Au is the preferred inert medium when the starting composition contains Fe and S, when compared to other potential capsule materials, e.g., Pb, Pd, etc. As the experimental series regarding sediment melting was intended to study sulfur solubility in

161 sediment melt as a function of oxygen fugacity, the starting composition used (EPSM-S1) has a
162 major element composition similar to that of the average “global subducting sediment” (GLOSS)
163 (Plank and Langmuir 1998), but with a sulfur content of ~1.9 wt%. Due to the low H₂O activity
164 (~0.01) in the inner capsule estimated based on H₂O content in melt, the fO_2 in the inner capsule
165 is generally ~2 log units lower than that imposed by the buffer.

166 Both the QIF-buffered exp. LMD776 and the Fe-(Mg,Fe)O sensor exp. LMD783 were
167 performed at 3 GPa, 1050°C; identical to the *PT* conditions of the previous experimental series
168 (Table 1). The QIF buffer assemblage has a mixing ratio of 1:1:1 between SiO₂, Fe metal powder
169 and fayalite. For the Fe-(Mg,Fe)O sensor experiment, the starting composition was a mix of Fe
170 metal powder and Mg(OH)₂ in order to achieve H₂O-saturated conditions. The resultant Fe:MgO
171 mixing ratio was 3:1.

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173 **Analytical methods**

174 The measurement of Fe in Au capsules was accomplished using an electron microprobe
175 (JEOL JXA 8230) at Peking University. By employing beam conditions of 20 kV and 50 nA and
176 a counting time of 60 s, a detection limit of ~55 ppm can be achieved. The diffusion profiles of
177 Fe in Au capsules have also been obtained across both the inner and outer capsule walls, with a
178 step size of 20 μm.

179 The analysis of S content in melt was performed on the same electron microprobe, employing
180 a PETH analyzing crystal, in combination with a defocused beam of 15 kV and 50 nA, and 100 s
181 counting time. The resultant detection limit was ~20 ppm. ZnS was used as the calibration
182 standard.

183 The Raman spectra of melts were recorded with a Horiba LabRam HR Evolution
184 spectrometer at Peking University, employing a Nd:YAG laser with a wavelength of 532 nm and
185 an emission power of 80 mW. The recording conditions are as follows: 100× objective lens, 100%
186 laser power, 600 grating and 500 μm pinhole. Acquisition time was 10s for each window,
187 accumulated over 3 scans.

188

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Results

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Fe content in Au capsules

191 For exp. LMD670, the profile of Fe in the Au capsule appears as a flat line, indicating an
192 equilibrium between the sample and the capsule (Fig. 3a). The average Fe content in the Au
193 capsule is about 0.072 ± 0.004 wt%; while the FeO content in melt is 0.32 ± 0.05 wt%. Therefore
194 an “apparent” Fe partition coefficient between melt and the Au capsule, calculated as the ratio
195 between FeO in melt and Fe in Au, is ~ 4.4 . According to the calibration of Ratajeski and Sisson
196 (1999) (Eq. 3 and 5), the $f\text{O}_2$ for exp. LMD670 is calculated to be $\sim \text{FMQ}-3.5$, which is ~ 2 log
197 units higher than the upper limit estimated based on graphite stability.

198 For the MW-buffered exp. LMD690, the profile of Fe in the Au inner capsule shows a general
199 increasing trend from the sample-inner capsule boundary to the inner capsule-buffer boundary
200 (Fig. 3b), which can be attributed to the diffusion of Fe from the buffer inwards. However, on
201 closer inspection, the profile begins with a section (~ 40 μm) of decreasing Fe content, which
202 may represent the diffusion of Fe from the sample to the Au capsule. Analyses of Fe in Au close
203 to the sample returned an average of 0.083 ± 0.007 wt%. With an FeO content of 0.86 ± 0.07 wt%
204 in melt, the “apparent” Fe partition coefficient between melt and the Au capsule is ~ 10.4 ,
205 suggesting the $f\text{O}_2$ is ~ 0.8 log units higher than that for exp. LMD670. Such a variance in $f\text{O}_2$ is

206 far smaller than what would have been anticipated. This may be due to the fact that Fe
207 partitioning between the sample and the Au capsule had yet to reach equilibrium. If we take the
208 lowest point of the profile (Fig. 3b), 0.021 wt%, which is the apparent meeting point of the two
209 diffusion sections, the calculated difference in fO_2 is ~ 2 log units.

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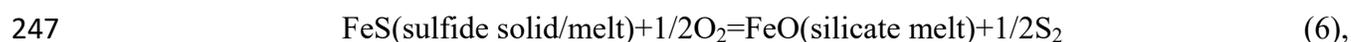
211 **The QIF-buffered experiment**

212 For the QIF-buffered exp. LMD776, the profile of Fe in the inner Au capsule also shows a
213 general increasing trend from the sample towards the buffer (Fig. 3d), reflecting the diffusion of
214 Fe from the buffer inwards. The wavy pattern observed in the Fe concentration profiles for both
215 inner and outer Au capsules may indicate a progressive approach toward equilibrium (Fig. 3d, e).
216 The capsule Fe content is generally higher than that observed for the MW-buffered exp.
217 LMD690, which may arise from either the lower fO_2 or the faster diffusion rate of metallic Fe
218 due to a higher concentration gradient or a combination of both. The Fe content in the region of
219 the Au capsule close to the sample has an average of 0.30 ± 0.04 wt%, while the FeO content in
220 melt is 1.36 ± 0.04 wt%, resulting in an apparent partition coefficient of ~ 4.5 , similar to that for
221 exp. LMD670. The fact that the beginning of the profile shows a minimal slope (Fig. 3d) may
222 suggest equilibrium has been attained between the sample and the Au capsule. If this is indeed
223 the case, the fO_2 imposed by the BN capsule sleeve would appear to be similar to that of the QIF
224 buffer.

225 The post-run buffer assembly has a rather interesting morphology, which is worth
226 documenting (Fig. S1). Note that we are observing only one longitudinal cross-section of the
227 capsule, and as such the possibility of heterogeneous distribution of buffer materials latitudinally
228 cannot be ruled out. Large olivine and ferrosilite crystals form a shell surrounding the inner

229 capsule. The formation of ferrosilite may be due to the reaction of Fayalite and SiO₂, or
230 alternatively to the reduction of olivine, which produces ferrosilite and metallic Fe. At one end of
231 the capsule, we observe the intergrowth of skeletal olivine and coesite between the olivine-
232 ferrosilite layer and the outer capsule. While the Ferrosilite-fayalite-Fe assemblage imposes
233 slightly lower f_{O_2} than QIF (Woodland and O'Neill 1997), it can be considered similar within
234 the accepted degree of experimental uncertainty. We found no presence of Fe metal in the outer
235 capsule, which may imply complete loss of Fe to the Au capsule, and consequently lower Fe
236 activity. Woodland and O'Neill (1997) observed lowered H₂O activity in the QIF assemblage at
237 1.7 GPa, 700-1000°C due to increased solubility of silicates in H₂O at high pressure and high
238 temperature. If the skeletal olivine and coesite represent quench crystal formation, i.e., high
239 solubility of coesite and olivine at 3 GPa, 1050°C, then H₂O activity will also be significantly
240 lowered. The final f_{H_2} in the outer capsule depends on the counterbalance of these two effects.
241 The leveling off of the Fe profile in the outer capsule of exp. LMD776 (Fig. 3e) suggests that the
242 buffer assemblage may have reached, or be close to reaching, equilibrium with the f_{H_2} imposed
243 by the assembly.

244 The FeO content in the melt of exp. LMD776 is far higher than would be predicted for such
245 low f_{O_2} conditions, as shown by the decreasing trend of FeO content in melt with decreasing f_{O_2}
246 for our experimental series (Li et al. 2021). As depicted by the FeS saturation reaction



248 at sulfide saturation, lower f_{O_2} leads to lower f_{S_2} , or lower FeO content in melt, or both. We have
249 observed decreasing trends for both f_{S_2} and FeO content in melt for our experimental series with
250 fixed bulk S content (Li et al. 2021). The higher FeO content in the melt of exp. LMD776 may
251 result from Fe contaminating the starting composition due to Fe diffusion from the buffer. The

252 melt S content falls within the trend observed in our previous study, that is, the lnS (ppm) in melt
253 has a negative linear correlation with $\ln X_{\text{FeO}}$ in melt with a slope close to -1 (Fig. 4). Such a
254 correlation for low FeO (<~5 wt%) melts is predicted by the thermodynamic framework for
255 modeling sulfur content in melt at sulfide saturation (SCSS) as a function of the standard state
256 Gibbs free energy of the FeS saturation reaction (Eq. 6), sulfide capacity (C_{S}^{2-}), the activity of
257 FeO in melt ($a_{\text{FeO}}^{\text{melt}}$) and the activity of FeS in sulfide melt ($a_{\text{FeS}}^{\text{sulfide}}$) (O'Neill and Mavrogenes
258 2002; O'Neill 2021); and in the case of hydrous rhyolitic melt with the relation (Li and Zhang
259 2022):

$$260 \quad \ln[\text{HS}]_{\text{SCSS}} = -\Delta G_{\text{FeS-FeO}}^{\circ}/RT + \ln C_{\text{HS}} - \ln a_{\text{FeO}}^{\text{melt}} + \ln a_{\text{FeS}}^{\text{sulfide}} \quad (7),$$

261 where C_{HS} represents the hydrosulfide capacity, defined for the dissolution of H_2S in hydrous
262 melt. Interested readers may find further details in the above references. The main point being
263 that such an observation (Fig. 4) provides indirect evidence that the sample and the near side of
264 the inner capsule may be at or close to an equilibrium state.

265 Further evidence supporting similar $f\text{O}_2$ conditions for exps. LMD670 and 776 comes from
266 the Raman spectra of their respective melts (Fig. 1). Firstly, their spectra show similar patterns in
267 terms of silicate peaks. As shown in our previous study, such patterns are notably different
268 between reduced (FMQ-3.8 and below) and oxidized (FMQ-2.2 and above) experiments (see Fig.
269 4 of Li et al. 2021). Secondly, they present H_2O and CH_4 peaks with similar intensity, suggesting
270 similar H_2O and CH_4 fugacity, and consequently similar H_2 and O_2 fugacity.

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272 **The Fe-(Mg,Fe)O sensor experiment**

273 Fe diffusion into the Au capsule is observed clearly as a grey band under stereomicroscope for
274 the Fe-(Mg,Fe)O sensor exp. LMD783 (Fig. 5a), demonstrating reduced redox conditions. The

275 achievement of equilibrium in the sample is demonstrated by the even distribution of Fe metal
276 and (Mg,Fe)O in the sample (Fig. 5b), and by homogeneous (Mg,Fe)O compositions. The band
277 of (Mg,Fe)O near the edge of the sample is interpreted to be a result of loss of Fe metal to the
278 capsule (Fig. 5a). This band of (Mg,Fe)O may have acted as a physical barrier preventing further
279 Fe loss from the sample, and consequently enabling the attainment of equilibrium.

280 The fO_2 of the experiment can be calculated based on the reaction $Fe + 1/2O_2 = FeO$ according
281 to the calibration of O'Neill et al. (2003), with the composition of (Mg,Fe)O reflecting deviation
282 from the IW buffer. The molar ratio of Mg/Fe in (Mg,Fe)O (~1) results in a calculated fO_2 of
283 IW-0.6, i.e., QIF+0.2, at 3 GPa, 1050°C. Considering analytical and experimental uncertainties,
284 this demonstrates that the assumption that BN imposed fO_2 conditions similar to that of the QIF
285 buffer is valid for our experimental assembly.

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287 **The diffusion coefficient of Fe in Au**

288 We are able to obtain the diffusion coefficient for Fe in Au by fitting the Fe concentration
289 profile in the Au capsule of exp. LMD783 (Fig. 5c) to the following equation from Zhang (2010):

$$290 \quad \ln C = \ln C_0 - \frac{x^2}{4Dt} \quad (8)$$

291 where x is the distance, t is time and D is the diffusion coefficient. This equation describes thin-
292 source diffusion into a semi-infinite medium. By plotting $\ln C$ vs x^2 , the coefficient of this linear
293 trend can be used to derive the diffusion coefficient, with a resulting value of 1×10^{-13} m²/s (Fig.
294 5d). Through fitting with the same equation, the Fe concentration profiles in both the inner and
295 outer capsules of exp. LMD690 return slightly lower diffusion coefficients, 3×10^{-14} m²/s and
296 5×10^{-14} m²/s, respectively (Fig. 6). These values compare favorably with published diffusion
297 coefficients at 1028°C: 1×10^{-12} m²/s at 0.1 MPa in He (Duhl et al. 1963), 1×10^{-13} m²/s at 5 GPa

298 in H₂, and 3×10^{-14} m²/s at 5 GPa without H₂ (Yamazaki et al. 2004). Considering that the
299 diffusion coefficient at 5 GPa in H₂ is three times higher than that in the absence of H₂, the lower
300 values obtained for exp. LMD690 may be due to its more oxidized conditions compared to those
301 for exp. LMD783. The possibility that the profile for the inner capsule of exp. LMD690 may
302 have been affected by reverse diffusion of Fe from the sample and that Fe diffusion in the
303 capsule of exp. LMD783 may have reached its outer boundary represent other possible factors
304 contributing to the variation in obtained diffusion coefficients.

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Discussion

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Determining fO_2 based on Fe content in Au capsules

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Calculating fO_2 based on the calibration from Ratajeski and Sisson (1999) has the advantage
of avoiding the uncertainties associated with calculating the FeO activity in hydrous rhyolitic
melt and the Fe activity in Au. However, the resultant fO_2 values are too high to be considered
reasonable. One possible cause is that the assumed fO_2 of FMQ-1 for the experiments of
Ratajeski and Sisson (1999) is overestimated. This possibility also agrees with the fact that
similar experiments from Patiño Douce and Beard (1994, 1995), performed without BN, have an
estimated fO_2 range of FMQ-2~FMQ. The calibration of Ratajeski and Sisson (1999) is similar
to the calibration for Fe partitioning between Pt metal and silicate melt obtained by Grove (1981).
However, the calibrations from Ratajeski and Sisson (1999) and Grove (1981) are not necessarily
equivalent considering the difference in γ_{FeO}^{melt} between rhyolitic melt and basaltic melt, and the
difference in γ_{Fe}^{alloy} between Fe-Au and Fe-Pt systems. Considering that Fe diffuses more readily
into Pt than Au at similar fO_2 conditions, the apparent equilibrium constant K' should have a
higher value for the Fe-Pt system than that for the Fe-Au system. Similar K' values obtained by

321 Ratajeski and Sisson (1999) for the Fe-Au system can therefore be attributed to overestimated
322 fO_2 (Eq. 4).

323 The fO_2 conditions for the equilibrium between FeO in melt and Fe in Au can be assessed in
324 comparison to the calibration for the IW buffer by O'Neill (1988). Apart from the melt and alloy
325 compositions, other parameters required include the equilibrium constant for the exchange
326 reaction between the FeO melt component and the FeO solid phase (wustite), and activity
327 coefficients γ_{FeO}^{melt} and γ_{Fe}^{metal} , as demonstrated by the equation below:

$$328 \quad \Delta IW = 2 \log \left(\frac{X_{FeO}^{melt}}{X_{Fe}^{alloy}} \right) + 2 \log \gamma_{FeO}^{melt} - 2 \log \gamma_{Fe}^{alloy} - 2 \log K(FeO^{melt} - FeO^s) \quad (9)$$

329 The estimated value is ~ 0.18 for $\log \gamma_{FeO}^{melt}$ based on the expression from O'Neill (2021), while the
330 $\log K$ term for the FeO melt component and wustite exchange reaction is calculated to be -0.37
331 for the 3 GPa, 1050°C conditions of exp. LMD670, based on the thermodynamic database of
332 Holland and Powell (2011). Although a positive deviation from ideal mixing has been observed
333 for Fe-Au alloys at ambient pressures (Seigle 1956), the activity of Fe in the Au capsule is
334 unknown for the very low Fe end of the Au-Fe solid solution. The low fO_2 for exp. LMD670
335 suggests that $\log \gamma_{Fe}^{alloy}$ has a positive value (~ 2).

336 When referring to redox sensors employing Fe alloys, Woodland and O'Neill (1997) state that
337 “the method loses accuracy as the activity of Fe decreases, leading to unacceptably large errors
338 when a_{Fe}^{alloy} is less than about 0.1”. This could also be the case for our experiments.

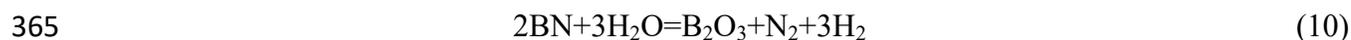
339

340 **An explanation for the QIF-like intrinsic fO_2 conditions for our BN experiment**

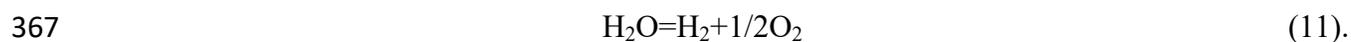
341 Both the comparability with the QIF-buffered experiment and direct fO_2 measurement using
342 the Fe-(Mg,Fe)O sensor, suggest that BN in our experimental assembly imposed fO_2 conditions
343 similar to that of the QIF buffer and thus far more reduced than the estimated FMQ-like

344 conditions assumed in previous studies. Such experimental studies used natural samples as their
345 starting materials, which would have imposed higher fO_2 when compared to the reduced starting
346 material used in our experimental series. However, as the difference in estimated fO_2 is so large,
347 we suspect that this likely results from the difference in the intrinsic fH_2 of the experimental
348 setup.

349 Truckenbrodt et al. (1997) investigated the redox conditions in piston-cylinder experiments
350 employing fired BN or unfired pyrophyllite as packing materials surrounding the capsule in
351 NaCl-graphite assemblies. They reported that C-O-H fluids from experiments using fired BN as
352 a packing material maintained the initial fH_2 of the starting organic compounds ($C_4H_4O_4$,
353 $C_9H_{10}O_2$, and $C_{14}H_{22}O$), while experiments with unfired pyrophyllite resulted in equivalent fH_2
354 for all three starting compositions, similar to cobalt-cobalt oxide (CoCoO)-buffered conditions,
355 and representative of the intrinsic fH_2 of the NaCl-graphite assembly. This interesting result
356 demonstrates that fired BN does not impose reduced fO_2 on the sample charge. This is likely due
357 to the fact that heat treatment of BN increases its crystallinity and purity, and consequently
358 reduces its ability to react with oxygen (Udayakumar et al. 2011). Moreover, their experimental
359 results demonstrate that the dehydration of pyrophyllite is key for transmitting the intrinsic fH_2 .
360 Matjuschkin et al. (2015) state that “As pointed out by Luth (1989), the availability of H_2 to
361 “transmit” any intrinsic fO_2 to the sample might depend on how well the larger parts of the
362 assembly are dried or on the dehydration of talc if used as an assembly component.” Under such
363 a consideration, we may derive the following reaction describing the oxidation of BN involving
364 H_2O :



366 through the addition of the BN oxidation reaction (Eq. 1) and the H_2O dissociation reaction



368 The $f\text{H}_2$ imposed by the BN sleeve is therefore a function of the H_2O activity, i.e., higher H_2O
369 activity, higher $f\text{H}_2$. Previously reported experiments in Li et al. (2021) were performed using a
370 cubic press. The experimental assembly used has a cylindrical configuration, similar to piston-
371 cylinder assemblies, but further encased in a pyrophyllite cube (Fig. 2). The more reduced
372 conditions observed in our BN experiment may be explained by a higher H_2O activity, which
373 may arise from dehydration of pyrophyllite in our experimental assembly.

374 The intrinsic H_2O activity of such assemblies remains unknown. At the commencement of
375 each run, gaskets form along the edges of the pyrophyllite cube during compression. If we can
376 consider the experimental assembly as a closed system after gasket formation, then H_2O activity
377 of the experimental assembly depends on the dehydration of pyrophyllite, and also the
378 consumption of H_2O due to BN oxidation. The BN oxidation reaction according to reaction (10)
379 produces both N_2 and H_2 . The H_2 fugacity can be calculated by assuming similar conditions to
380 those imposed by the QIF buffer (O'Neill 1987) plus pure H_2O and subsequently the H_2 activity
381 can be estimated to be ~ 0.6 . Based on the stoichiometry of reaction (10), and an assumption of
382 ideal mixing between gas species and similar fugacity coefficients, N_2 activity can be estimated
383 to be ~ 0.2 . The H_2O activity is also ~ 0.2 if the gas phase is dominated by H_2 , N_2 and H_2O . The
384 reduced N_2 activity implies that the $f\text{O}_2$ value would be ~ 1.4 log units lower than that of the QIF
385 buffer (Eq. 1), however, the similarly reduced H_2O activity means that $f\text{H}_2$ remains the same as
386 that of the QIF buffer (Eq. 10). There may be other gas species present, but not considered, for
387 example, CH_4 , from the reaction of H_2O with graphite. Although much-simplified, such an
388 inverse modeling approach suggests that BN in our experimental assembly can maintain $f\text{H}_2$

389 similar to that of the QIF buffer plus H₂O, especially in the limiting case when a gas phase is
390 present due to the dehydration of pyrophyllite.

391

392 **The application of BN as an fO_2 control in solid media assemblies**

393 The intrinsic fO_2 of our experimental assembly, using BN as a capsule sleeve and spacer,
394 appears significantly more reduced than that reported by previous studies using common piston-
395 cylinder assemblies containing BN. The primary difference lies in the usage of the pyrophyllite
396 cube in our experimental setup. This represents a simple way to impose a reduced redox state
397 relevant for investigating lunar conditions or reduced endmembers of the Martian environment,
398 although the resultant chemical system will be H₂O-bearing. According to reaction (10),
399 lowering H₂O activity will reduce or inhibit the BN oxidation reaction and its fO_2 buffering
400 ability. Piston-cylinder experiments normally employ NaCl, talc or BaCO₃ tube as the pressure
401 transmitting material encasing the graphite furnace, and MgO, Al₂O₃, pyrophyllite, BN or pyrex
402 as assembly components within the graphite furnace. An available source of H₂O is required in
403 order to impose the intrinsic fO_2 of the experimental assembly onto the sample charge. This
404 could be either moisture absorbed by the pressure transmitting material, or H₂O from
405 dehydration of talc and pyrophyllite.

406 The material immediately surrounding the capsule exerts the dominant control on the fO_2 of
407 the sample. It is evident from previous studies that BN can impose more reduced conditions than
408 MgO, which is useful when variation of fO_2 is a factor of research interest, although the exact
409 value of fO_2 needs to be determined independently. The starting material can also have a
410 buffering effect on fO_2 due to the exchange of H₂ between the sample and the assembly
411 (Holloway et al. 1992). Whether it is the intrinsic fO_2 of the assembly or that of the sample which

412 is desired, the redox potential between the sample and assembly can result in either gain or loss
413 of H₂O and Fe from the charge. Therefore, BN provides a suitable choice of pressure
414 transmitting material for reduced experimental redox conditions as it minimizes possible
415 modification of the starting composition.

416 A cautionary note on using BN as a pressure transmitting material: the experimental charge
417 may be contaminated with boron. As investigated and reported by Matjuschkin et al. (2015),
418 when BN powder was packed around a Pt capsule, 2000-3000 ppm boron was detected in glasses
419 produced from experiments run at 1200-1265°C for 24-26 hours. It appears that the magnitude of
420 boron contamination is similar to that of carbon infiltration (Matjuschkin et al. 2015). It is worth
421 keeping this issue in mind, especially for high temperature experiments.

422 An alternative approach to fO_2 control is to use the packing material surrounding the capsule
423 as a barrier between the capsule and assembly, in order to either maintain the redox state of the
424 sample or prolong the life of the oxygen buffer assemblage. Matjuschkin et al. (2015)
425 investigated the H₂ permeability of different pressure transmitting materials, and demonstrated
426 that pyrex functions better than MgO or Al₂O₃ at limiting the diffusion of H₂, therefore
427 maintaining the intended fH_2 within the capsule. The experimental results of Truckenbrodt et al.
428 (1997) suggest that fired BN also acts as an excellent H₂ barrier.

429

430

Implications

431 Controlling fO_2 in solid media assemblies is a challenging task. More often than not, the fO_2
432 of such experimental runs is not directly measured or controlled, but reflects the intrinsic fO_2 of
433 either the assembly or the sample, or somewhere in-between. The double-capsule configuration
434 employing a buffer assemblage is currently the most effective approach for achieving fO_2 control,

435 but requires a large capsule volume and is more difficult and time-consuming to prepare.
436 Therefore, imposing approximate fO_2 by selecting the appropriate capsule or pressure
437 transmitting material for the assembly is still the first choice for the majority of experimental
438 studies. Through our efforts to determine the fO_2 imposed by our chosen experimental assembly
439 using BN as a capsule sleeve, we are able to demonstrate the reducing ability of BN as an
440 assembly component, and furthermore shed light on the process by which BN imposes such
441 reduced fO_2 . We hereby present what we have learnt during the course of this investigation, in
442 the hope that the effect of BN on fO_2 control is both recognized and further exploited in future
443 experimental studies.

444

445

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447 microprobe analyses and Jesse Jones for helpful comments. Constructive reviews from Bruno
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569

570 **Figure captions:**

- 571 Fig. 1 A comparison of the Raman spectra of quenched melt from the BN experiment LMD670,
572 the MW-buffered exp. LMD690 and the QIF-buffered exp. LMD776.

573 Fig. 2 A schematic diagram of the experimental cubic press assembly used for our experiments.

574 Fig. 3 The Fe content profiles in both inner and outer Au capsules for the BN experiment

575 LMD670 (a), the MW-buffered exp. LMD690 (b, c), and the QIF-buffered exp. LMD776 (d,

576 e). Note that scales for the y axis vary between graphs for different experiments. The

577 variation in profile length reflects the variability of capsule width, both regionally within

578 each capsule and between capsules from individual experiments, caused by the compression

579 encountered during experimental runs. The two profiles for exp. LMD670 represent data

580 obtained from firstly a position closer to the center of the capsule (red line) and secondly a

581 distal position (blue line). The temperature gradient may explain the smoother profile for

582 the hotter position (red line).

583 Fig. 4 $\ln S$ (ppm) versus $\ln X_{\text{FeO}}$ in melt for 3 GPa, 1050°C experiments reported previously in Li

584 et al. (2021) and exp. LMD776 from this study. Note that the fitting of the trendline is based

585 on previous experiments.

586 Fig. 5 Stereoscopic photos of exp. LMD783 showing (a) the diffusion of Fe into the Au capsule

587 and (b) the intergrowth of Fe metal and (Mg,Fe)O. The Fe diffusion profile is plotted in (c).

588 The fitting of the Fe diffusion profile to equation (8) is shown in (d).

589 Fig. 6 The fitting of Fe diffusion profiles in both the inner (a) and outer (b) capsules from exp.

590 LMD690 to equation (8). Note that x represents the distance measured from the buffer for

591 both profiles. The profile from the inner capsule of exp. LMD690 is shown in blue in Fig.

592 3b with the first 40 μm section excluded for the purpose of fitting.

593

594 Supplementary figures:

595 Fig. S1 The positions for Fe in Au profiles shown as arrows for exp. LMD670 (a), LMD690 (b)

596 and LMD776 (c). Arrow colors correspond to the color of the profiles in Fig. 3. The black
597 rectangle in (c) outlines the magnified area shown in (d). Mineral abbreviations: Coe,
598 Coesite; Fs, ferrosilite; Ol, olivine.

599 Supplementary data tables:

600 Table S1 Raw data for Fe diffusion profiles.

601 Table S2 Fe content in Au capsules in contact with the sample charge.

602 Table S3 Phase compositions in exp. LMD776.

603 Table S4 Phase compositions in exp. LMD783.

604

605

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607

608 Table 1 Summary of experimental conditions and run products.

609

Exp. No.	Starting comp.	<i>P</i> (GPa)	<i>T</i> (°C)	Time (hours)	Buffer	Estimated <i>f</i> O ₂	Phase assemblage
^a LMD670	EPSM-S1	3	1050	96	BN sleeve	FMQ-7.5	melt ^b (95) Po ^c <i>Grt Coe Ap</i>
LMD690	EPSM-S1	3	1050	46	MW	FMQ-3.8	melt (95) Po Grt Coe Ap Gr
LMD776	EPSM-S1	3	1050	48	QIF	FMQ-7.5	melt (85) Po Grt Coe Ap
^a LMD783	Fe+Mg(OH) ₂	3	1050	24	BN sleeve	FMQ-5.2	Fe, (Mg,Fe)O

610

611 Mineral abbreviations: Ap, apatite; Coe, Coesite; Gr, graphite; Grt, garnet; Po, pyrrhotite.

612 Buffer abbreviations: QIF: quartz-iron-fayalite; MW, magnetite-wustite; FMQ: fayalite-magnetite-quartz.

613 ^aNo buffers were employed in these experiments.

614 ^bNumbers in brackets following “melt” are melt percentages in the charge, estimated by sight.

615 ^cMineral phases in italics appear at the colder end of the capsule only.

616

Figure 1

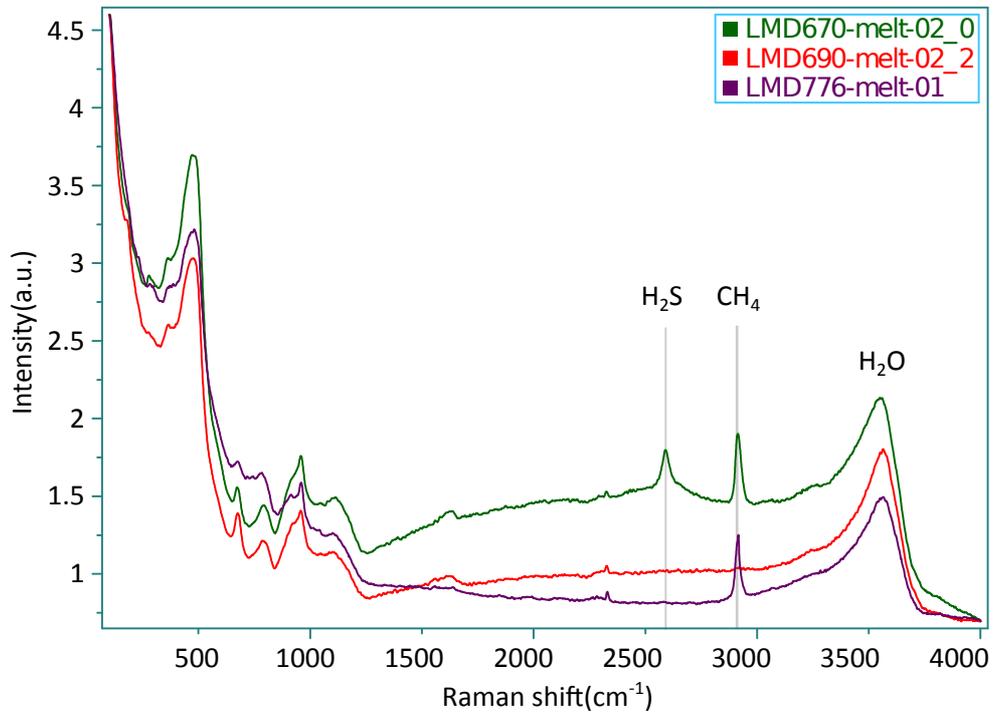


Fig.2

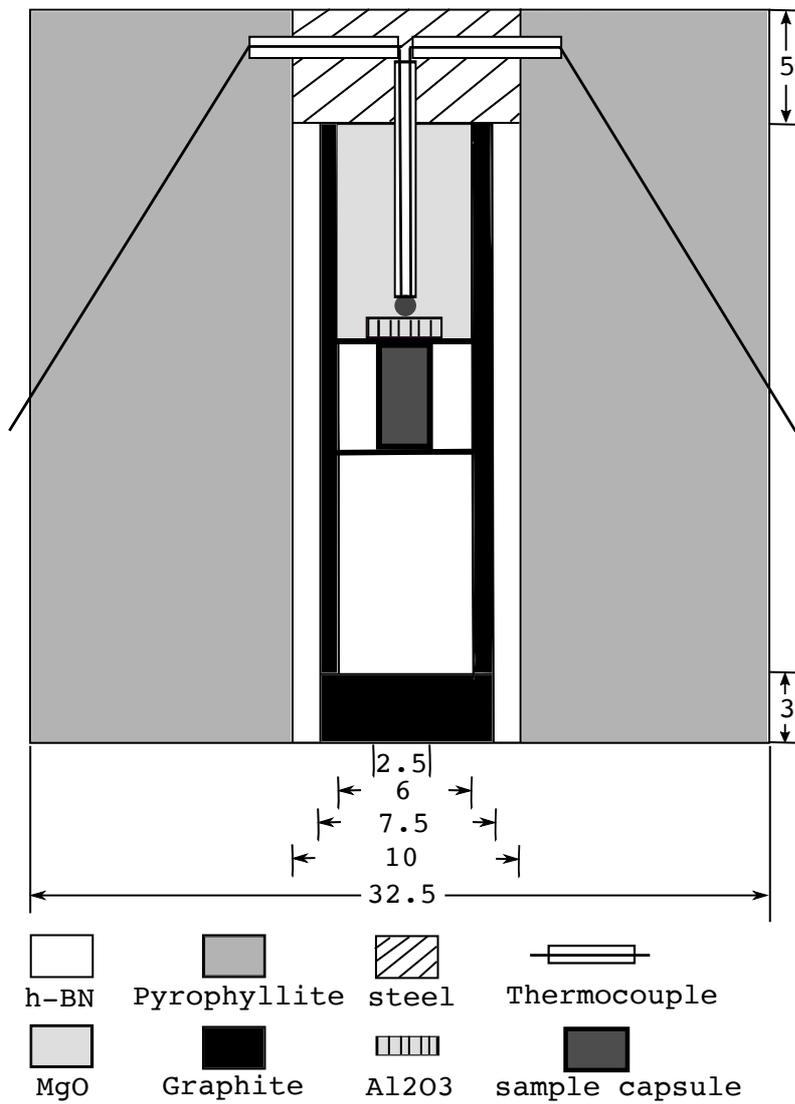


Fig. 3

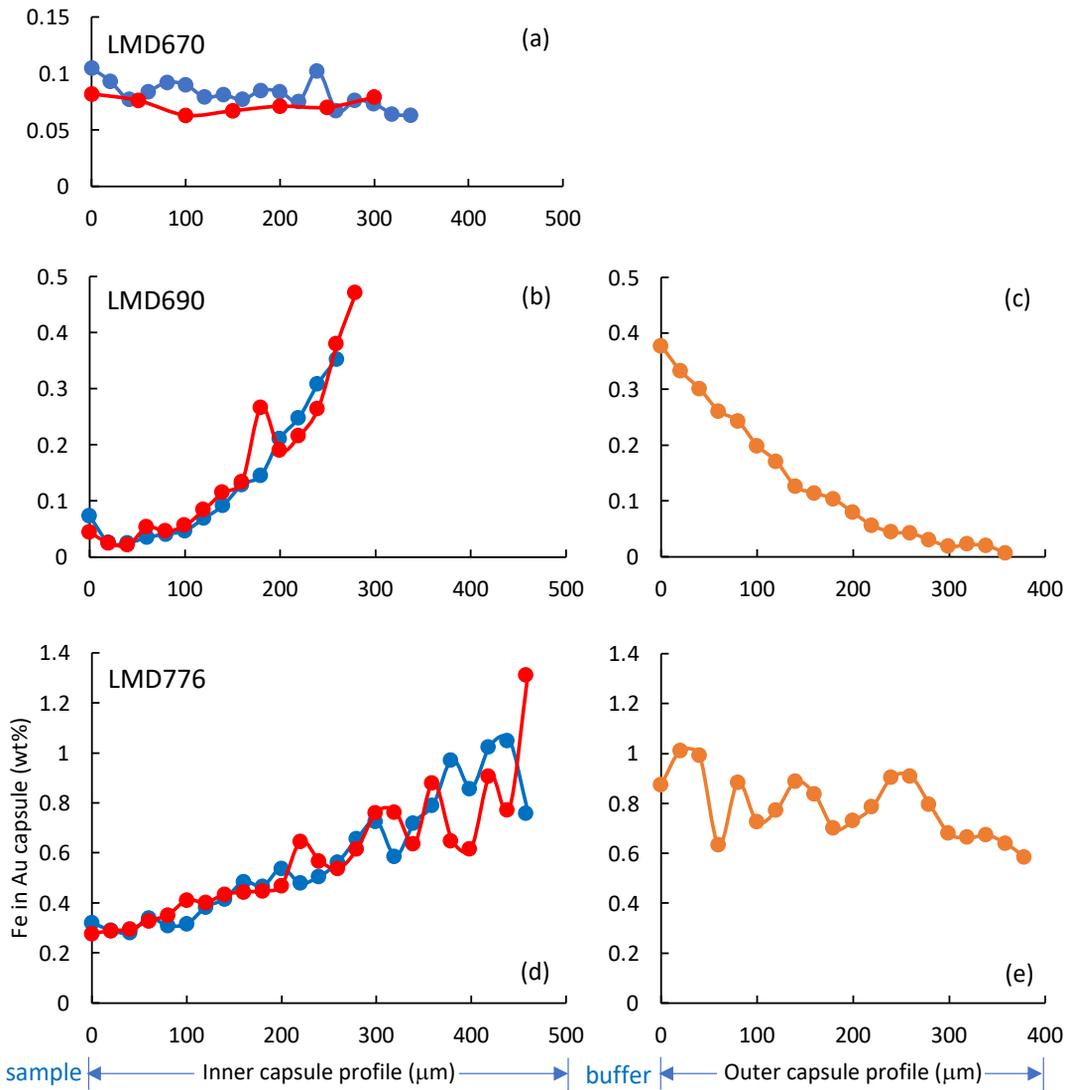


Fig. 4

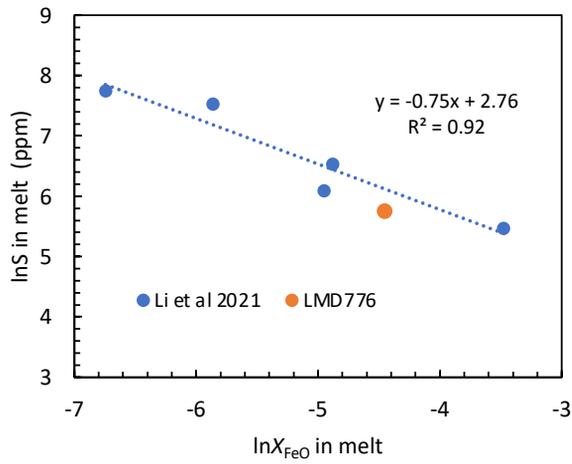


Fig. 5

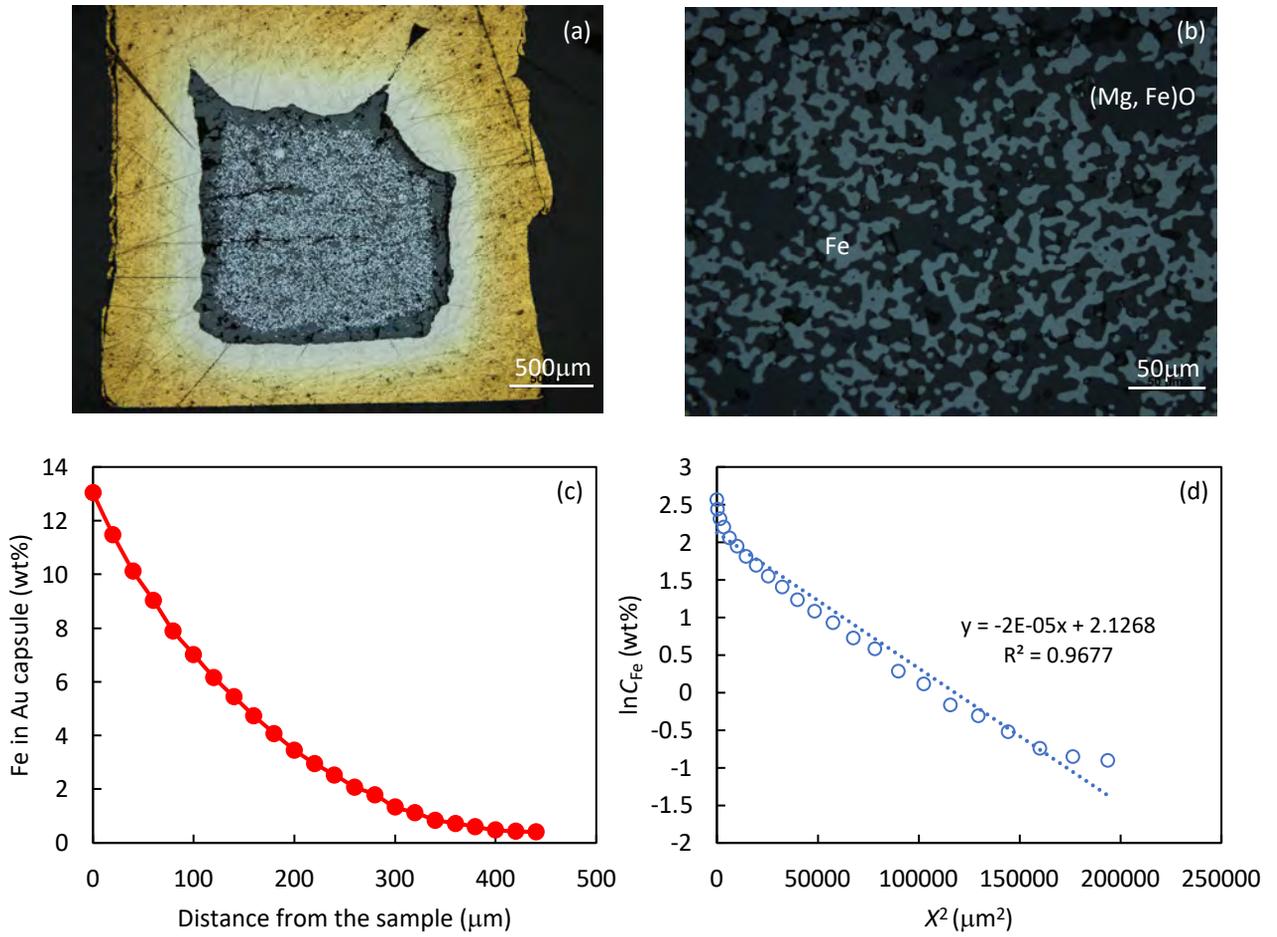


Fig. 6

