Revision 2

2 The iron spin transition of deep nitrogen-bearing mineral Fe₃N_{1.2} at

- 3 high pressure
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10 Highlights:

- 1) In-situ X-ray emission spectroscopy measurements were performed on Fe₃N_{1.2} up to 45.8
- 12 GPa at room temperature.
- 13 2) The pressure-induced spin transition of iron in $Fe_3N_{1.2}$ starts at relatively low pressures and
- 14 completes at ~ 30.5 GPa.
- 15 3) The iron spin transition pressure is highly related to the nitrogen concentration of
- 16 hexagonal iron nitrides.
- 17 4) The identity and concentration of light elements, together with crystal structure, may take
- 18 charge of the spin transition pressure of iron-rich alloys.
- 19

20 Abstract

Nitrogen is an essential element for life, one of the most abundant volatiles in the 21 22 atmosphere, and an important component in the Earth's interior, where iron nitride is an 23 essential host of deep nitrogen. Here, we investigate the pressure-induced electronic spin-pairing transition of iron in siderazot (Fe₃ $N_{1,2}$) at pressures to 45.8 GPa at room 24 25 temperature, using diamond-anvil cell techniques coupled with synchrotron x-ray emission 26 spectroscopy. The integrated intensity of the satellite emission peak (K_{β}) decreases once upon compression and remains unchanged at pressures greater than 30.5 GPa. In other words, 27 28 the high-spin to low-spin transition of iron in $Fe_3N_{1,2}$ starts immediately at very low pressures 29 and completes at ~ 30.5 GPa. The iron spin transition completion pressures increase with the nitrogen concentration of hexagonal close-packed iron nitrides (i.e., Fe₃N_{1.2}, Fe₇N₃, and 30 31 Fe_2N). Moreover, the identity and concentration of light elements, together with their crystal 32 structure, could affect the iron spin transition pressure of binary iron-rich compounds such as Fe₃N, Fe₃C, Fe₃P, Fe₃S, Fe₇C₃, and Fe₇N₃. The spin transition of iron-rich alloys could alter 33 the bonding nature and the physical properties including the thermal and electrical 34 35 conductivity, influencing the thermal state and evolution of planetary interiors in turn. **Keywords:** spin transition, x-ray emission spectroscopy, high pressure, deep nitrogen

37 Introduction

38 Nitrogen is a key probe to understand the accretion and degassing of volatiles and the differentiation and evolution history of the Earth. It is also one of the most poorly studied 39 elements under the Earth's deep interior pressure and temperature (P-T) conditions. Surface 40 41 and deep nitrogen are largely connected via subduction and volcanic degassing [Halama et 42 al., 2012; Mao and Mao, 2020; Halama and Bebout, 2021]. Notably, the nitrogen concentration in the present-day bulk silicate Earth may be only tens of ppm, severely 43 44 depleted with respect to other volatiles like carbon and sulfur [Marty, 2012; Bergin et al., 2015; Yoshioka et al., 2018]. Apart from volatile degassing during accretion, the metallic core 45 has been considered a hidden reservoir to account for the missing nitrogen according to the 46 siderophile nature of nitrogen as revealed in metal-silicate partition experiments [Hirschmann, 47 2016; Li et al., 2016; Dalou et al., 2017; Dalou et al., 2019]. Up to 0.5 wt.% nitrogen, 48 together with other candidate light elements of Si, O, C, S, and H, has been suggested in the 49 50 Earth's core to interpret the core density deficit based on the global seismic models, e.g., the Preliminary Reference Earth Model (PREM) [Dziewonski and Anderson, 1981; Sugiura, 51 52 1998]. In addition, the deep mantle becomes relatively reduced with the $\sim 1 \text{ wt\%}$ metallic iron that is saturated at depths greater than 250 km [Rohrbach et al., 2007]. Based on the 53 54 presence of iron nitrides and carbonitrides in meteorites and deep diamond inclusions, nitrogen may be distributed heterogeneously and enriched in metal phases in the Earth's 55 interior [Li and Keppler, 2014; Litasov et al., 2017; Rubin and Ma, 2017; Huang et al., 56 2021]. 57

58	Iron nitrides have been widely suggested as the primary hosts of deep nitrogen in the
59	reduced Earth's interior; their thermostability and phase relations have attracted extensive
60	attention. Among these Fe-N intermediates, orthorhombic ζ -Fe ₂ N (space group: <i>Pbcn</i>), cubic
61	γ' -Fe ₄ N (space group: $Pm\bar{3}m$), and nonstoichiometric siderazot ϵ -Fe ₃ N _x (0.75 <x<1.5) are<="" td=""></x<1.5)>
62	stable at ambient conditions [Niellsen and Burhwald, 1981; Minobe et al., 2015; Litasov et al.,
63	2017; Yoshioka et al., 2018]. Under compression, ε -Fe ₃ N _x forms from Fe ₂ N and Fe ₄ N at 9-15
64	GPa and 1400-1600 K [Schwarz et al., 2009; Guo et al., 2013]. It has either a P3 ₁ 2 or P6 ₃ 22
65	space group with hexagonal close-packed iron atoms while nitrogen atoms occupy part of the
66	octahedral voids [Sifkovits et al., 1999; Guo et al., 2013; Bette et al., 2021]. The ε -Fe ₃ N _x
67	phase is the same as the ε -Fe ₇ N ₃ phase when x is equal to ~ 1.3. Fe ₇ N ₃ undergoes the
68	pressure-induced phase transition from ϵ to β crystal structures at pressures greater than 41
69	GPa and $\sim 1000~K$ while the $\beta\text{-}Fe_7N_3$ has been considered a candidate nitrogen-bearing phase
70	in the Earth's core [Minobe et al., 2015; Kusakabe et al., 2019]. Additionally, previous
71	studies reported that the ε -Fe ₃ N _x phase is stable at least up to 60 GPa at room temperature [Lv
72	et al., 2020; Huang et al., 2021]. To date, the effects of nitrogen concentration on the stability
73	and physical properties of nonstoichiometric ε -Fe ₃ N _x have not been investigated.
74	The iron $3d$ orbitals collapse under extremely high pressures of the deep Earth, leading to
75	the high-spin to low-spin transition of iron-bearing minerals. The configuration and
76	interaction of the outermost shell electrons of iron, which are dramatically altered by pressure,
77	are essentially significant for materials' physical and chemical properties [Liu et al., 2019;
78	Zhao et al., 2020]. The pressure-induced spin transition of iron greatly changes the chemical

79	bonding of relevant minerals and thus, affects thermodynamic properties, rheology, and the
80	thermal and electrical conductivity of iron-bearing minerals [Lin et al., 2013]. The iron spin
81	transition occurs widely in the deep mantle for both iron-bearing oxide-hydroxides and
82	metallic compounds [Lin et al., 2004a; Lin and Tsuchiya, 2008; Gu et al., 2014; Chen et al.,
83	2018; Su et al., 2021]. At ambient conditions, the 3d-orbit electrons of Fe in all iron nitrides
84	are in a high-spin ferromagnetic arrangement [Sifkovits et al., 1999]. The magnetic moment
85	of Fe-N compounds generally decreases with increasing nitrogen concentration, and it is
86	about 2.0 μ_B for ϵ -Fe ₃ N [Panda and Gajbhiye, 1997; Leineweber et al., 2001; Guo et al.,
87	2013]. The iron spin transition has been widely reported for the Fe-C, Fe-S, and Fe-P systems,
88	while it has only been investigated in the Fe-N system for $\gamma'\text{-}Fe_4N$ and $\epsilon\text{-}Fe_7N_3$ under high
89	pressure [Ishimatsu et al., 2003; Lv et al., 2020]. Lv et al. [2020] revealed that the iron spin
90	transition of ε -Fe ₇ N ₃ completes at 43 GPa and 300 K, but it may not apply to other ε -Fe ₃ N _x
91	compositions. Thus far, the effects of nitrogen on the iron spin transition pressure remains
92	poorly constrained with respect to other light elements because the crystal structure and
93	arrangement of iron 3d-orbital electrons could be affected by nitrogen concentration
94	[Widenmeyer et al., 2014].

In this work, the spin transition of iron in ε -Fe₃N_{1.2} is determined in diamond-anvil cells by synchrotron x-ray emission spectroscopy (XES) at high pressure. First, the magnetic spin state of iron in ε -Fe₃N_{1.2} is measured up to 45.8 GPa at 300 K. The high-spin to low-spin transition of iron in ε -Fe₃N_{1.2} starts at pressures below 9.8 GPa and completes at ~ 30.5 GPa. Then, we evaluate the effects of nitrogen concentration on iron nitrides' spin transition

pressure by comparing with previous studies on the spin transition of iron-rich alloys.
Furthermore, it is discussed how the identity and concentration of light elements regulate the
iron spin transition pressure of candidate iron-rich alloys.

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104 Methods

The ε -Fe₃N_x polycrystalline (Lot#: RA191101) sample was obtained from Antechnology 105 Co., Ltd, Beijing. Previous studies suggested the ideal space group for ε -Fe₃N_x is either P3₁2 106 or $P6_{3}22$ and their structural differences are subtle and challenging to be recognized by x-ray 107 108 diffraction measurements [Guo et al., 2013]. Moreover, the P3₁2 space group has been recommended for nonstoichiometric ε -Fe₃N_x while P6₃22 for stoichiometric ε -Fe₇N₃ 109 [Leineweber et al., 2001; Niewa et al., 2009; Lv et al., 2020]. In this work, the x-ray 110 111 diffraction pattern of ε -Fe₃N_x at ambient conditions confirms its hexagonal crystal structure in the space group of $P3_12$ and unit cell volume of 85.788(75) Å³ with unit cell parameters 112 a=b=4.740(1) Å and c=4.409(2) Å. On the basis of the empirical relationship between the 113 114 unit cell volume and nitrogen content of iron nitrides [Litasov et al., 2017], the 115 polycrystalline sample has a mildly higher nitrogen content than stoichiometric ε -Fe₃N with the chemical formula of $Fe_3N_{1,216(7)}$. For simplification, $Fe_3N_{1,2}$ is used to represent the 116 117 ε -Fe₃N_{1,216(7)} sample hereafter in this study.

The Fe₃N_{1.2} sample of ~ 60 μ m in diameter and ~ 8 μ m thick was loaded into a diamond anvil cell with a flat culet size of 300 μ m. The Be gasket was pre-indented to 20–25 μ m thick, and a hole of 200 um in diameter was drilled into the center of the indentation. Cubic boron 121 nitride (cBN)-epoxy powder was packed into this hole and then compacted to 20 GPa at room temperature. A hole of $\sim 100 \ \mu m$ in diameter was finally drilled into the center of the cBN 122 gasket insert to serve as the sample chamber using the laser drilling system at the Center for 123 High Pressure Science and Technology Advanced Research (HPSTAR), Beijing. Silicone oil 124 was used as a pressure-transmitting medium and two 5-8 µm-sized ruby spheres were placed 125 next to the sample as pressure calibration [Mao et al., 1986]. The pressure was measured and 126 averaged over multiple times before and after XES measurements at a given pressure, and its 127 128 uncertainty is mostly less than 3%. 129 High-pressure x-ray emission spectra of Fe₃N_{1.2} were collected between 2.9 and 45.8 GPa at room temperature at the HPCAT beamline 16-ID-D of the Advanced Photon Source, 130 Argonne National Laboratory. The incident x-ray beam of 11.3 keV with the full width at half 131 132 maximum (FWHM) of less than 1 eV was focused down to $\sim 20 \ \mu m$ vertically and horizontally at the sample position. The Fe emission signal was collected from 7020 to 7080 133 eV with an energy step of 0.3 eV using a Peltier-cooled silicon detector after being selected 134 135 by a silicon analyzer [Liu et al., 2019]. Helium gas was filled into the x-ray path to reduce the absorption by air. Each XES spectrum was collected for ~ 1 hour, and three XES spectra were 136 accumulated at each pressure. All spectra were normalized to the transmitted intensity, and 137 the K_{β} main peaks were aligned to 7058 eV [Liu et al., 2019]. The intensity difference of the 138 satellite peak K_{β} was integrated from 7030 to 7053 eV using the integrated relative difference 139 (IRD) method [Chen et al., 2014]. Compared with the integrated absolute difference (IAD) 140 141 method, the IRD method aligns the main peaks to the same energy position, which could

avoid the broadening effect of the $K_{\beta l,3}$ peak in modelling the spin momentum reduction with increasing pressure [*Mao et al.*, 2014; *Lin et al.*, 2016].

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145 **Results and Discussion**

The magnetic order and spin state of iron in Fe₃N_{1.2} were probed by the K_{β} fluorescence 146 peaks from the XES spectra up to 45.8 GPa (Figure 1). Each K_{β} emission spectrum at a given 147 pressure can be divided into a main $K_{\beta l,3}$ peak and a lower-energy satellite K_{β} peak due to the 148 149 final state interaction between the 3p core hole and the electrons of the partially filled 3dshell of iron [Gu et al., 2016; Liu et al., 2019]. All spectra were normalized to the transmitted 150 151 intensity and then shifted to set the $K_{\beta I,3}$ emission lines to 7058 eV. The intensity of the $K_{\beta'}$ peak is indicative of the spin state and total spin moment of iron in Fe₃N_{1.2}. As the pressure 152 increases, a gradual reduction of the K_{β}' peak corresponds to the pressure-induced high-spin 153 to low-spin transition of iron [Lin et al., 2005]. The integrated intensity of the K_{β} peak is 154 plotted between 7030 and 7053 eV and fitted to the Boltzmann function in Figure 2. Upon 155 compression, the intensity of the satellite K_{β}' peak decreases gradually up to 30.5 GPa and 156 remains unchanged at higher pressures. The steep slope in the relationship between the 157 high-spin fraction and pressure under relatively low pressures indicates that the spin 158 transition of iron in Fe₃N starts at pressure below 2.9-9.8 GPa. Moreover, the unchanged 159 intensity of the K_{β} peak at pressures over 30.5 GPa suggests the completion of the iron spin 160 161 transition of Fe₃N_{1.2} at ~ 30.5 GPa.

162	The magnetic state is highly related to the spin transition of iron in Fe ₃ N _{1.2} . At ambient
163	conditions, $Fe_3N_{1,2}$ is in a ferromagnetic (FM) state while all the 3 <i>d</i> orbitals of Fe are
164	occupied with four unpaired electrons [Leineweber et al., 2001; Lin et al., 2013]. The
165	high-spin to low-spin transition of $Fe_3N_{1.2}$ completes at ~ 30.5 GPa, likely corresponding to
166	the FM to nonmagnetic (NM) transition. We note that the FM to paramagnetic (PM)
167	transition cannot be detected with the XES measurements due to the relative orientations of
168	individual spins [Lv et al., 2020]. The compression behavior may provide an indirect
169	measurement for the FM-PM transition in the Fe-N system, which has been observed in other
170	Fe-light element compounds like Fe ₇ C ₃ [<i>Chen et al.</i> , 2014]. Among the Fe-N compounds, the
171	discontinuity of the unit cell volume of Fe ₄ N at ~ 24 GPa in Zhuang et al. [2021] may
172	correspond to the FM-PM transition, as suggested by Ishimatsu et al. [2003] using XMCD
173	measurements. However, this method might not be applicable for ϵ -Fe ₃ N _x because of its
174	relatively smooth compression behavior [Adler and Williams, 2005; Lv et al., 2020]. It should
175	be noted that the plateau in the relationship between the unit cell volume and pressure in
176	Adler and Williams [2005] is likely caused by the nonhydrostatic conditions. Previous ab
177	initio calculations predicted the spin transition of iron in ϵ -Fe ₃ N at ~ 125 GPa and 0 K, which
178	is inconsistent with our XES results [Popov et al., 2015]. Nevertheless, further work is
179	needed to decode the FM-PM transition in the ϵ -Fe ₃ N _x by using XMCD or Mössbauer
180	spectroscopy measurements.

181 The nitrogen concentration, together with the crystal structure, may control the 182 magnetism properties of iron nitrides such as ζ -Fe₂N, ε -Fe₃N, ε -Fe₇N₃, and γ' -Fe₄N. In the

183	lattices of these Fe-N compounds, nitrogen atoms occupy the octahedral sites with strong
184	covalent <i>p</i> - <i>d</i> bonds, which greatly influences the magnetic moments [Sifkovits et al., 1999;
185	Guo et al., 2013]. At ambient conditions, the γ' -Fe ₄ N is based on the face-centered cubic
186	structure of iron with 1/4 of the octahedral voids occupied by nitrogen, carrying a moment of
187	about 2.0 μ_B . Moreover, ζ -Fe ₂ N, ε -Fe ₇ N ₃ , and ε -Fe ₃ N are based on the (slightly distorted)
188	hexagonal close-packed lattice of iron with 1/2, 3/7, and 1/3 of the octahedral voids occupied
189	by nitrogen, respectively [Sifkovits et al., 1999]. The number of nearest neighboring nitrogen
190	atoms for each iron atom varies from 2 to 3, resulting in a magnetic moment reduction from
191	2.0 to 1.5 μ_B for these hexagonal iron nitrides [Panda and Gajbhiye, 1997]. In other words,
192	the nitrogen concentration could strongly influence the nature of the pressure-induced spin
193	transition of iron nitrides.

The completion pressures of the iron spin transition for all the iron nitrides are higher 194 than pure iron. As shown in Figure 2, the spin transition completion pressures for metallic 195 iron is ~ 13 GPa, which corresponds to the bcc-hcp phase transition at 300 K [Monza et al., 196 197 2011]. This suggests a positive effect of nitrogen on the spin transition completion pressure. The spin transition completion pressure of Fe₃N_{1,2} in this work is lower than that of Fe₇N₃ 198 reported by Lv et al. [2020], though the experimental settings are nearly identical. The 199 200 different completion pressures indicate that nitrogen concentration mostly dictates the iron spin transition of the hexagonal Fe-N compounds. In addition, this trend is further confirmed 201 because Fe₂N with higher nitrogen concentration shows a higher spin transition completion 202 203 pressure. These results suggest that more nitrogen in the Fe-N hexagonal lattice provides

204	more covalent bonds, which likely enhances the structural stability against compression
205	[Sifkovits et al., 1999]. On the other hand, γ' -Fe ₄ N shows a relatively high spin transition
206	completion pressure while its nitrogen content is much lower than ϵ -Fe ₃ N. This may be
207	attributed to the difference in the crystal structure and nature of Fe-N bonding. There are two
208	different Fe sites in γ' -Fe ₄ N: Fe(I) is similar to fcc-Fe, and Fe(II) has two nearest neighboring
209	nitrogen atoms, whereas all the Fe atoms are equally sited in hexagonal iron nitrides like
210	ε-Fe ₃ N [Sifkovits et al., 1999]. Therefore, nitrogen concentrations, together with crystal
211	structures, primarily control the iron spin transition of iron nitrides at high pressure. The
212	completion pressures of spin transition of iron in Fe ₃ C and Fe ₇ C ₃ also exhibit a similar trend
213	on how carbon concentration affects the pressure range of the iron spin transition in the Fe-C
214	system [Chen et al., 2018].

215

216 Implications

217 The iron spin transition has been widely observed in iron-rich Fe-light element alloys, such as the Fe-C, Fe-N, Fe-S, and Fe-P systems [Ishimatsu et al., 2003; Lin et al., 2004a; Lin 218 et al., 2004b; Chen et al., 2014; Gu et al., 2014; Gu et al., 2016; Chen et al., 2018; Lai et al., 219 2020; Lv et al., 2020]. The spin transitions of iron in these Fe-light element alloys are often 220 221 associated with a volume collapse. Consequently, the equation of state for the low-spin state differs from that for the high-spin state at given P-T conditions [Chen et al., 2014; Lai et al., 222 2020; Lv et al., 2020]. The spin transition pressures of various Fe-rich binary alloys are 223 224 compared in Figure 3. The spin transition completion pressures of iron in the two iron

carbides are greater than in iron nitrides, suggesting carbon has a stronger lattice stabilizing 225 effect than nitrogen. The large difference in the spin transition completion pressures between 226 Fe₃N_{1,2} and Fe₃C may result from different crystal structures. On the contrary, the spin 227 transition pressures are comparable between Fe₇N₃ and Fe₇C₃. We note that Fe₇N₃, Fe₇C₃, and 228 $Fe_3N_{1,2}$ have a hexagonal lattice, while Fe_3C has an orthorhombic structure [*Caracas*, 2016]. 229 Moreover, the spin transition completion pressures of iron sulfide Fe₃S and iron phosphide 230 Fe₃P are relatively lower than iron carbides and nitrides. Such pressure differences could be 231 ascribed to the tetragonal crystal structure of Fe₃S and Fe₃P because both sulfur and 232 233 phosphorus have larger atomic radii [Slater, 1964]. Similarly, the spin transition onset pressures are also different from each other. The iron spin transition of Fe₃N starts at 234 pressures lower than 5-10 GPa, whereas other iron nitrides and all iron carbides undergo the 235 236 electronic spin-pairing transition of iron at pressures greater than 10 GPa. Notably, the iron spin transition of all the iron-rich phosphides and sulfides starts at very low pressures under 237 compression. Therefore, the identity and concentration of light elements in the metallic cores 238 239 of terrestrial planets and asteroids may control the spin state of iron at high pressure.

The spin transition completion pressure may be positively correlated with high temperatures for iron alloyed with light elements. Previous studies found that the mixed spin state of iron widens with increasing temperature in ferropericlase, iron-bearing carbonate, and metallic iron [*Tsuchiya et al.*, 2006; *Lin et al.*, 2007; *Liu et al.*, 2014; *Ono*, 2015]. In addition, the onset pressure of the iron spin transition could be slightly increased with increasing temperature, while the completion pressure may be enhanced significantly with increasing

246	temperature. The HS-LS boundary slope varies among previous studies, and it is about 0.045
247	K/GPa for hcp-Fe [Ono, 2015]. Assuming that iron-rich alloys share a similar slope as pure
248	iron under high temperatures, all the Fe-C, Fe-N, Fe-S, and Fe-P systems might be in LS state
249	in the mantle and core of the present-day Earth [Katsura et al., 2010]. On the other hand, the
250	spin crossover of iron-rich alloys may occur under the P-T conditions of small terrestrial
251	planets, e.g., Mars and Mercury [Smith et al., 2012; Margot et al., 2018; Stähler et al., 2021].
252	Furthermore, the spin transition of iron in Fe ₃ N and other iron-rich alloys could have an
253	influence on the compressibility and conductivity of solid metallic cores of terrestrial planets,
254	likely controlling the nucleation pattern, thermal dynamo, and structure of planetary interiors.
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256

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267 References

- Adler, J. F., and Williams, Q. (2005), A high-pressure X-ray diffraction study of iron nitrides: Implications for
 Earth's core, *Journal of Geophysical Research*, *110*(B1), B01203, doi:10.1029/2004jb003103.
- 270 Bergin, E. A., Blake, G. A., Ciesla, F., Hirschmann, M. M., and Li, J. (2015), Tracing the ingredients for a 271 habitable earth from interstellar space through planet formation, *Proceedings of the National Academy of*
- 272 *Sciences*, *112*(29), 8965-8970, doi:10.1073/pnas.1500954112.
- Bette, S., Theye, T., Bernhardt, H.-J., Clark, W. P., and Niewa, R. (2021), Confirmation of siderazot, Fe₃N_{1.33},
 the only terrestrial nitride mineral, *Minerals*, *11*(3), 290, doi:10.3390/min11030290.
- 275 Caracas, R. (2016), Crystal structures of core materials, in Deep Earth: Physics and Chemistry of the Lower
- 276 Mantle and Core, edited by Terasaki, H. and Fischer, R. A., pp. 55-68, Wiley,
 277 doi:10.1002/9781118992487.ch5.
- Chen, B., Lai, X., Li, J., Liu, J., Zhao, J., Bi, W., et al. (2018), Experimental constraints on the sound velocities
 of cementite Fe₃C to core pressures, *Earth and Planetary Science Letters*, 494, 164-171,
 doi:10.1016/j.epsl.2018.05.002.
- Chen, B., Li, Z., Zhang, D., Liu, J., Hu, Y. M., Zhao, J., et al. (2014), Hidden carbon in Earth's inner core
 revealed by shear softening in dense Fe₇C₃, *Proceedings of the National Academy of Sciences*, 111(50),
 17755-17758, doi:10.1073/pnas.1411154111.
- Dalou, C., Furi, E., Deligny, C., Piani, L., Caumon, M.-C., Laumonier, M., et al. (2019), Redox control on
 nitrogen isotope fractionation during planetary core formation, *Proceedings of the National Academy of Sciences*, *116*(29), 14485-14494, doi:10.1073/pnas.1820719116.
- Dalou, C., Hirschmann, M. M., von der Handt, A., Mosenfelder, J., and Armstrong, L. S. (2017), Nitrogen and
 carbon fractionation during core-mantle differentiation at shallow depth, *Earth and Planetary Science Letters*,
 458, 141-151, doi:10.1016/j.epsl.2016.10.026.
- Dziewonski, A. M., and Anderson, D. L. (1981), Preliminary reference Earth model, *Physics of the Earth and Planetary Interiors*, 25(4), 297-356, doi:10.1016/0031-9201(81)90046-7.
- Gu, T., Fei, Y., Wu, X., and Qin, S. (2014), High-pressure behavior of Fe₃P and the role of phosphorus in
 planetary cores, *Earth and Planetary Science Letters*, *390*, 296-303, doi:10.1016/j.epsl.2014.01.019.
- Gu, T., Fei, Y., Wu, X., and Qin, S. (2016), Phase stabilities and spin transitions of $Fe_3(S_{1-x}P_x)$ at high pressure
- and its implications in meteorites, *American Mineralogist*, 101(1), 205-210, doi:10.2138/am-2016-5466.
- 296 Guo, K., Rau, D., von Appen, J., Prots, Y., Schnelle, W., Dronskowski, R., et al. (2013), High pressure
- high-temperature behavior and magnetic properties of Fe_4N : Experiment and theory, *High Pressure Research*,
- 298 *33*(3), 684-696, doi:10.1080/08957959.2013.809715.
- Halama, R., and Bebout, G. (2021), Earth's nitrogen and carbon cycles, *Space Science Reviews*, 217(3), 45,
 doi:10.1007/s11214-021-00826-7.
- 301 Halama, R., Bebout, G. E., John, T., and Scambelluri, M. (2012), Nitrogen recycling in subducted mantle rocks
- and implications for the global nitrogen cycle, *International Journal of Earth Sciences*, *103*(7), 2081-2099,
 doi:10.1007/s00531-012-0782-3.
- 304 Hirschmann, M. M. (2016), Constraints on the early delivery and fractionation of Earth's major volatiles from
- 305 C/H, C/N, and C/S ratios, American Mineralogist, 101(3), 540-553, doi:10.2138/am-2016-5452.

- Huang, S., Wu, X., Zhu, F., Lai, X., Li, J., Neill, O. K., et al. (2021), High-pressure phase stability and
 thermoelastic properties of iron carbonitrides and nitrogen in the deep Earth, *Journal of Geophysical Research: Solid Earth*, *126*(6), B021934, doi:10.1029/2021jb021934.
- 309 Ishimatsu, N., Maruyama, H., Kawamura, N., Suzuki, M., Ohishi, Y., Ito, M., et al. (2003), Pressure-induced
- 310 magnetic transition in Fe₄N probed by Fe *K*-edge XMCD measurement, *Journal of the Physical Society of* 311 *Japan*, 72(9), 2372-2376, doi:10.1143/jpsj.72.2372.
- Katsura, T., Yoneda, A., Yamazaki, D., Yoshino, T., and Ito, E. (2010), Adiabatic temperature profile in the mantle, *Physics of the Earth and Planetary Interiors*, *183*(1-2), 212-218, doi:10.1016/j.pepi.2010.07.001.
- 314 Kusakabe, M., Hirose, K., Sinmyo, R., Kuwayama, Y., Ohishi, Y., and Helffrich, G. (2019), Melting curve and
- 315 equation of state of β-Fe₇N₃: Nitrogen in the Core?, *Journal of Geophysical Research: Solid Earth*, *124*(4), 316 3448-3457, doi:10.1029/2018jb015823.
- Lai, X., Zhu, F., Liu, Y., Bi, W., Zhao, J., Alp, E. E., et al. (2020), Elastic and magnetic properties of Fe₃P up to
 core pressures: Phosphorus in the Earth's core, *Earth and Planetary Science Letters*, *531*(1), 115974,
 doi:10.1016/j.epsl.2019.115974.
- Leineweber, A., Jacobs, H., Hüning, F., Lueken, H., and Kockelmann, W. (2001), Nitrogen ordering and ferromagnetic properties of ϵ -Fe₃N_{1+x} (0.10 \leq x \leq 0.39) and ϵ -Fe₃(N_{0.80}C_{0.20})_{1.38}, *Journal of Alloys and Compounds*, *316*(1-2), 21-38, doi:10.1016/s0925-8388(00)01435-3.
- Li, Y., and Keppler, H. (2014), Nitrogen speciation in mantle and crustal fluids, *Geochimica et Cosmochimica Acta*, *129*, 13-32, doi:10.1016/j.gca.2013.12.031.
- Li, Y., Marty, B., Shcheka, S., Zimmermann, L., and Keppler, H. (2016), Nitrogen isotope fractionation during terrestrial core-mantle separation, *Geochemical Perspectives Letters*, 2, 138-147, doi:10.7185/geochemlet.1614.
- Lin, J.-F., Fei, Y., Sturhahn, W., Zhao, J., Mao, H.-k., and Hemley, R. J. (2004a), Magnetic transition and sound
 velocities of Fe₃S at high pressure: Implications for Earth and planetary cores, *Earth and Planetary Science Letters*, 226(1-2), 33-40, doi:10.1016/j.epsl.2004.07.018.
- Lin, J.-F., Speziale, S., Mao, Z., and Marquardt, H. (2013), Effects of the electronic spin transitions of iron in
 lower mantle minerals: Implications for deep mantle geophysics and geochemistry, *Reviews of Geophysics*,
 51(2), 244-275, doi:10.1002/rog.20010.
- Lin, J.-F., Struzhkin, V. V., Mao, H.-k., Hemley, R. J., Chow, P., Hu, M. Y., et al. (2004b), Magnetic transition in
 compressed Fe₃C from X-ray emission spectroscopy, *Physical Review B*, 70(21), 212405,
 doi:10.1103/PhysRevB.70.212405.
- Lin, J.-F., and Tsuchiya, T. (2008), Spin transition of iron in the Earth's lower mantle, *Physics of the Earth and Planetary Interiors*, *170*(3-4), 248-259, doi:10.1016/j.pepi.2008.01.005.
- Lin, J. F., Mao, Z., Yang, J., Liu, J., Xiao, Y., Chow, P., et al. (2016), High-spin Fe²⁺ and Fe³⁺ in single-crystal
 aluminous bridgmanite in the lower mantle, *Geophysical Research Letters*, 43(13), 6952-6959,
 doi:10.1002/2016gl069836.
- 342 Lin, J. F., Struzhkin, V. V., Jacobsen, S. D., Hu, M. Y., Chow, P., Kung, J., et al. (2005), Spin transition of iron in
- 343 magnesiowustite in the Earth's lower mantle, *Nature*, *436*(7049), 377-380, doi:10.1038/nature03825.
- Lin, J. F., Vanko, G., Jacobsen, S. D., Iota, V., Struzhkin, V. V., Prakapenka, V. B., et al. (2007), Spin transition
- 345 zone in Earth's lower mantle, *Science*, *317*(5845), 1740-1743, doi:10.1126/science.1144997.

- 346 Litasov, K. D., Shatskiy, A., Ponomarev, D. S., and Gavryushkin, P. N. (2017), Equations of state of iron nitrides
- 347 ε -Fe₃N_x and γ -Fe₄N_y to 30 GPa and 1200 K and implication for nitrogen in the Earth's core, *Journal of*
- 348 *Geophysical Research: Solid Earth*, *122*(5), 3574-3584, doi:10.1002/2017jb014059.
- Liu, J., Hu, Q., Bi, W., Yang, L., Xiao, Y., Chow, P., et al. (2019), Altered chemistry of oxygen and iron under deep Earth conditions, *Nature Communications*, *10*(1), 153, doi:10.1038/s41467-018-08071-3.
- 351 Liu, J., Lin, J. F., Mao, Z., and Prakapenka, V. B. (2014), Thermal equation of state and spin transition of
- magnesiosiderite at high pressure and temperature, *American Mineralogist*, 99(1), 84-93,
 doi:10.2138/am.2014.4553.
- Lv, M., Liu, J., Zhu, F., Li, J., Zhang, D., Xiao, Y., et al. (2020), Spin transitions and compressibility of ε-Fe₇N₃
 and γ'-Fe₄N: Implications for iron alloys in terrestrial planet cores, *Journal of Geophysical Research: Solid*
- 356 *Earth*, *125*(11), B020660, doi:10.1029/2020jb020660.
- Mao, H.-k., and Mao, W. L. (2020), Key problems of the four-dimensional Earth system, *Matter and Radiation at Extremes*, 5(3), 038102, doi:10.1063/1.5139023.
- Mao, H. K., Xu, J., and Bell, P. M. (1986), Calibration of the ruby pressure gauge to 800 kbar under
 quasi-hydrostatic conditions, *Journal of Geophysical Research: Solid Earth*, 91(B5), 4673-4676,
 doi:10.1029/JB091iB05p04673.
- Mao, Z., Lin, J. F., Yang, J., Wu, J., Watson, H. C., Xiao, Y., et al. (2014), Spin and valence states of iron in
 Al-bearing silicate glass at high pressures studied by synchrotron Mossbauer and X-ray emission
 spectroscopy, *American Mineralogist*, 99(2-3), 415-423, doi:10.2138/am.2014.4490.
- Margot, J.-L., Hauck II, S. A., Mazarico, E., Padovan, S., and Peale, S. J. (2018), Mercury's Internal Structure,
 in *Mercury: The View after MESSENGER*, edited by SC, S., LR, N. and BJ, A., pp. 85-113, Cambridge
 University Press, Cambridge
- Marty, B. (2012), The origins and concentrations of water, carbon, nitrogen and noble gases on Earth, *Earth and Planetary Science Letters*, *313-314*, 56-66, doi:10.1016/j.epsl.2011.10.040.
- 370 Minobe, S., Nakajima, Y., Hirose, K., and Ohishi, Y. (2015), Stability and compressibility of a new iron-nitride
- β -Fe₇N₃ to core pressures, *Geophysical Research Letters*, 42(13), 5206-5211, doi:10.1002/2015gl064496.
- Monza, A., Meffre, A., Baudelet, F., Rueff, J. P., d'Astuto, M., Munsch, P., et al. (2011), Iron under pressure:
 "Kohn tweezers" and remnant magnetism, *Physical Review Letters*, 106(24), 247201,
 doi:10.1103/PhysRevLett.106.247201.
- Niellsen, H. P., and Burhwald, V. F. (1981), Roaldite, a new nitride in iron meteorites, paper presented at
 Proceedings Lunar and Planet Science Conference.
- 377 Niewa, R., Rau, D., Wosylus, A., Meier, K., Wessel, M., Hanfland, M., et al. (2009), High-pressure 378 high-temperature phase transition of γ' -Fe4N, *Journal of Alloys and Compounds*, 480(1), 76-80, 379 doi:10.1016/j.jallcom.2008.09.178.
- Ono, S. (2015), Relationship between structural variation and spin transition of iron under high pressures and
 high temperatures, *Solid State Communications*, 203, 1-4, doi:10.1016/j.ssc.2014.11.010.
- Panda, R. N., and Gajbhiye, N. S. (1997), Magnetic properties of single domain ε-Fe₃N synthesized by
 borohydride reduction route, *Journal of Applied Physics*, *81*(1), 335-339, doi:10.1063/1.364115.
- 384 Popov, Z. I., Litasov, K. D., Gavryushkin, P. N., Ovchinnikov, S. G., and Fedorov, A. S. (2015), Theoretical
- 385 study of γ' -Fe₄N and ε -Fe_xN iron nitrides at pressures up to 500 GPa, *JETP Letters*, 101(6), 371-375,
- 386 doi:10.1134/s0021364015060090.

- 387 Rohrbach, A., Ballhaus, C., Golla-Schindler, U., Ulmer, P., Kamenetsky, V. S., and Kuzmin, D. V. (2007), Metal
- 388 saturation in the upper mantle, *Nature*, *449*(7161), 456-458, doi:10.1038/nature06183.
- Rubin, A. E., and Ma, C. (2017), Meteoritic minerals and their origins, *Geochemistry*, 77(3), 325-385,
 doi:10.1016/j.chemer.2017.01.005.
- 391 Schwarz, U., Wosylus, A., Wessel, M., Dronskowski, R., Hanfland, M., Rau, D., et al. (2009), High-pressure-
- 392 high-temperature behavior of ζ -Fe₂N and phase transition to ϵ -Fe₃N_{1.5}, European Journal of Inorganic
- 393 *Chemistry*, 2009(12), 1634-1639, doi:10.1002/ejic.200801222.
- 394 Sifkovits, M., Smolinski, H., Hellwig, S., and Weber, W. (1999), Interplay of chemical bonding and magnetism
- 395 in Fe₄N, Fe₃N and ζ -Fe₂N, *Journal of Magnetism and Magnetic Materials*, 204(3), 191-198, 396 doi:10.1016/s0304-8853(99)00296-6.
- Slater, J. C. (1964), Atomic radii in crystals, *The Journal of Chemical Physics*, 41(10), 3199-3204,
 doi:10.1063/1.1725697.
- Smith, D. E., Zuber, M. T., Phillips, R. J., Solomon, S. C., Hauck, S. A., Lemoine, F. G., et al. (2012), Gravity
 field and internal structure of Mercury from MESSENGER, *Science*, *336*(6078), 214-217.
- 401 Stähler, S. C., Khan, A., Banerdt, W. B., Lognonné, P., Giardini, D., Ceylan, S., et al. (2021), Seismic detection
 402 of the martian core, *Science*, *373*(6553), 443-448, doi:10.1126/science.abi7730.
- 403 Su, X., Zhao, C., Xu, L., Lv, C., Song, X., Ishii, T., et al. (2021), Spectroscopic evidence for the Fe^{3+} spin 404 transition in iron-bearing δ -AlOOH at high pressure, *American Mineralogist*, 106(11), 1709-1716, 405 doi:doi:10.2138/am-2021-7541.
- Sugiura, N. (1998), Ion probe measurements of carbon and nitrogen in iron meteorites, *Meteoritics & Planetary Science*, *33*(3), 393-409, doi:10.1111/j.1945-5100.1998.tb01645.x.
- Tsuchiya, T., Wentzcovitch, R. M., da Silva, C. R., and de Gironcoli, S. (2006), Spin transition in
 magnesiowustite in earth's lower mantle, *Physical Review Letters*, 96(19), 198501,
 doi:10.1103/PhysRevLett.96.198501.
- 411 Widenmeyer, M., Hansen, T. C., Meissner, E., and Niewa, R. (2014), Formation and decomposition of iron
- 412 nitrides observed by in situ powder neutron diffraction and thermal analysis, Zeitschrift fur Anorganische und
- 413 *Allgemeine Chemie*, *640*(7), 1265-1274, doi:10.1002/zaac.201300676.
- 414 Yoshioka, T., Wiedenbeck, M., Shcheka, S., and Keppler, H. (2018), Nitrogen solubility in the deep mantle and
- the origin of Earth's primordial nitrogen budget, *Earth and Planetary Science Letters*, 488, 134-143,
 doi:10.1016/j.epsl.2018.02.021.
- Zhao, C., Xu, L., Gui, W., and Liu, J. (2020), Phase stability and vibrational properties of iron-bearing
 carbonates at high pressure, *Minerals*, *10*(12), 1142, doi:10.3390/min10121142.
- Zhuang, Y., Su, X., Salke, N. P., Cui, Z., Hu, Q., Zhang, D., et al. (2021), The effect of nitrogen on the
 compressibility and conductivity of iron at high pressure, *Geoscience Frontiers*, 12(2), 983-989,
- 421 doi:10.1016/j.gsf.2020.04.012.
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423 Figure Captions

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Figure 1. High-pressure XES spectra of Fe₃N_{1.2} up to 45.8 GPa at 300 K. The XES spectra were normalized to unity in their integrated intensity. The insert is the full view of the Fe- K_{β} fluorescence spectra from 7020 to 7080 eV, with the main peak aligned to 7058 eV. The reduction of the satellite peak with increasing pressure indicates the high-spin to low-spin transition of Fe₃N_{1.2} under compression.



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Figure 2. The complete pressure of spin transition in iron nitrides as a function of nitrogen content. The upper panel shows the high-spin fraction of Fe in iron nitrides as a function of pressure. The red squares show the high-spin fraction of hexagonal $Fe_3N_{1.2}$ derived from the XES measurements in this study. The spectrum at 2.9 GPa is adopted as the high spin reference spectrum while the average of nearly-unchanged spectra at 35–50 GPa is used as

437	the low spin reference spectrum. The red curve is the Boltzmann function fitting to the XES
438	data of Fe ₃ N _{1.2} . The solid yellow and blue curves represent the high-spin fraction for
439	hexagonal Fe ₇ N ₃ and Fe ₄ N, respectively, derived from the XES experiments by Lv et al.
440	[2020]. The dashed olive curve is the iron spin transition of hexagonal Fe_2N based on the
441	crystal axis ratio as a function of pressure [Zhuang et al., 2021]. The volume jump of Fe ₂ N at
442	10-20 GPa in Zhuang et al. [2021] may correspond to an isostructural phase transition such
443	as the FM-PM transition rather than the HS-LS transition. The grey zone indicates the spin
444	transition pressure (magnetic collapse) of pure Fe, which is accompanied simultaneously by
445	iron's bcc to hcp transition at 300 K [Monza et al., 2011]. The lower panel shows the spin
446	transition complete pressure as a function of nitrogen content. Circle symbols represent
447	hcp-Fe-N phases, and the open square stands for fcc Fe ₄ N. The dashed lines are drawn to
448	guide the eye.





Figure 3. Spin transition pressure of iron-rich nitrides, carbides, phosphides, and sulfides. Each line shows the pressure range of the iron spin transition for a given phase. The left and right ends of the lines denote the onset and completion pressures of the HS-LS transition, respectively. The dotted line of $Fe_3N_{1.2}$ indicates the uncertainty of the onset pressure of the HS-LS transition. Literature data for Fe_3P and $Fe_3(P,S)$ are from XES experiments in *Gu et al.* [2016], Fe_3S from SMS experiments in *Lin et al.* [2004a], Fe_3C and Fe_7C_3 from XES experiments in *Chen et al.* [2018], and Fe_7N_3 from XES experiments in *Lv et al.* [2020].