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3	Anomalous elastic behavior of phase Egg, AlSiO ₃ (OH), at high pressures							
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11								
12	Abstract							
13	Phase Egg (AlSiO ₃ (OH)), is an aluminosilicate hydrous mineral that is thermodynamically stable							
14	in lithological compositions represented by Al ₂ O ₃ -SiO ₂ -H ₂ O (ASH) ternary i.e., a simplified							
15	ternary for the mineralogy of subducted sediments and continental crustal rocks. High- pressure							
16	and high-temperature experiments on lithological compositions resembling hydrated sedimentary							
17	layers in subducting slabs show that phase Egg is stable up to pressures of 20-30 GPa which							
18	translates to transition zone to lower mantle depths. Thus phase Egg is a potential candidate for							
19	transporting water into the Earth's mantle transition zone. In this study, we use first-principles							
20	simulations based on density functional theory to explore the pressure dependence of crystal							
21	structure and how it influences energetics and elasticity. Our results indicate that phase Egg							
22	exhibits anomalous behavior of the pressure dependence of the elasticity at mantle transition							
23	zone depths (~ 15 GPa). Such anomalous behavior in the elasticity is related to changes in the							

hydrogen bonding O-H...O configurations, which we delineate as a transition from a low-24 25 pressure to a high-pressure structure of phase Egg. Full elastic constant tensors indicate that phase Egg is very anisotropic resulting in a maximum anisotropy of compressional wave 26 velocity, $AV_P \sim 30$ % and of shear wave velocity, $AV_S \sim 17$ % at zero pressures. Our results also 27 28 indicates that the phase Egg has one of the fastest bulk sound velocities (V_P and V_S) compared to 29 other hydrous aluminous phases in the ASH ternary, which include topaz-OH, phase Pi, and δ -30 AlOOH. However, the bulk sound velocity of phase Egg is slower than that of stishovite. At depths corresponding to the base of mantle transition zone, phase Egg decomposes to a mixture 31 32 of δ -AlOOH and stishovite. The changes in compressional ΔV_P and shear ΔV_S velocity associated with the decomposition is ~ 0.42 % and -1.23 % respectively. 33

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Key-words: Phase Egg, equation of state, elasticity, anisotropy, symmetric hydrogen bonding,
subduction zone

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38 Introduction

Based on estimates of planetary accretion, geochemistry, and the influx and outgassing of water, 39 40 it is evident that our understanding of the size of the deep mantle reservoir of water is far from 41 complete. Water is absolutely essential for the sustenance of planetary activities, including 42 melting (Hirschmann, 2006) that eventually leads to large-scale geochemical differentiation. Water also helps in sustaining planetary geodynamics. In particular, water affects transport 43 44 properties including the rheology (Mei and Kohlstedt, 2000), viscosity (e.g., Ichikawa et al., 2015) and electrical conductivity (e.g., Wang et al., 2006). Without water the mantle convection 45 46 processes for a dry silicate rock will be extremely sluggish. Therefore, it is important to estimate

the size of such a water reservoir in the deep Earth. Knowing the extent of mantle hydration is 47 48 also likely to enhance understanding of the role of solid Earth in influencing the sea level over long time scales ($\sim 10^9$ years). However, constraining the water content of the deep Earth is not 49 50 straightforward. It is important to estimate how water is transported into the deep Earth and the 51 water-storage potential and stability of minerals and rocks at deep Earth conditions, which is an 52 upper bound of possible water content. Estimation of actual water storage requires geophysical 53 observations of elasticity and electrical conductivity of the deep Earth and the effect of water on 54 those physical properties.

Most of the mantle consists of nominally anhydrous minerals (NAMs) with trace 55 quantities of water, which are thermodynamically stable along a normal mantle geotherm i.e., do 56 57 not readily decompose at these temperatures. Hydrous mineral phases are distinct from the NAM phases and contain structurally bound hydroxyl groups i.e., OH⁻ groups occur in well-defined 58 59 crystallographic sites (Smyth, 2006). Although, water is transported into the deep Earth via 60 subduction of hydrated lithosphere containing hydrous mineral phases (Kawamoto, 2006), a 61 distinct limitation for hydrous phases as potential reservoir for water over geological time scale is that they are not thermodynamically stable in a normal mantle geotherm. Hence, water is 62 63 likely to be partitioned to NAMs whose water retention capacity is significantly lower. For 64 example, while hydrous phases can host several wt % of water (Kawamoto, 2006; Mookherjee et 65 al., 2015), at depths greater than 200 km, NAMs can host 70-700 wt ppm of H₂O as hydroxyl defects (Bolfan-Casanova, 2005).. However, based on high-pressure and high-temperature 66 experiments it is known that aluminous hydrous phases have significantly greater thermal 67 68 stability than the magnesio-silicate hydrous phases and are likely to be stable in normal mantle 69 geotherms. The aluminous hydrous phases can be represented by the Al_2O_3 -SiO₂-H₂O (ASH)

ternary i.e., a simplified ternary for the mineralogy of a subducted sedimentary layer (Peacock, 70 71 1990; Ono, 1999; Schmidt et al., 1998; Schreyer et al., 1995; Wunder et al., 1993a,b; Pamato et 72 al., 2015). Several hydrous mineral phases are stable in the ASH ternary system. These include 73 gibbsite (Al(OH)₃), diaspore (AlOOH), kaolinite (Al₂Si₂O₅(OH)₄), topaz-OH (Al₂SiO₄(OH)₂), 74 phase Pi (Al₃Si₂O₇(OH)₃), phase Egg (AlSiO₃(OH)), and dense high-pressure phases such as δ -AlOOH. In addition, recent experimental studies suggest that the dense hydrous magnesium 75 76 silicate minerals phase-D (MgSi₂O₄(OH)₂) could have an aluminum-rich end member with 77 stoichiometry of $Al_2SiO_4(OH)_2$ (Pamato et al., 2015). It is estimated that the net water contribution from oceanic crust is likely to be greater than from the sediments (Peacock, 1990). 78 79 However, owing to the greater thermal stability of the minerals stabilized in sedimentary layer, 80 they may be effective in transporting water into the deep Earth in warmer subduction zones 81 (Ono, 1998). Among these hydrous phases, phase Egg has the potential to carry water to mantle 82 transition zone depths and as a result it has been subject of numerous investigations that 83 elucidate its high-pressure and temperature stability and compressibility. Phase Egg was first 84 synthesized above 10 GPa (Eggleton et al., 1978) and since then the thermodynamic stability 85 have been investigated numerous times with the high-pressure limits ranging from 20 - 30 GPa 86 where phase Egg decomposes to δ -AlOOH and stishovite (Schmidt et al., 1998; Ono et al., 1999; 87 Sano et a., 2004; Fukuyama et al., 2017; Abe et al., 2018). Phase Egg has also been found as inclusions in natural diamonds from Junia, Brazil indicating that crustal materials are indeed 88 89 deeply subducted to the mantle transition zone and upper part of lower mantle (Wirth et al., 90 2007). Although phase Egg could potentially play an important role in transporting water into the 91 deep Earth, its elasticity remains unknown. The latter is a crucial physical property that will help 92 us to understand the mantle hydration caused by subduction of hydrated sediments or alumina 93 rich crustal rocks.

In this study, we use *first-principles* simulations to explore how pressure affects the crystal structure and proton environments, equation of state, full elastic constant tensor and elastic anisotropy of phase Egg at high pressures.

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98 Methods

99 Phase Egg, AlSiO₃(OH), is stable at pressures greater than 11 GPa and it has a monoclinic symmetry with space group $P2_{1/n}$ (Schmidt et al., 1998). Based on the powder X-ray 100 101 diffraction refinement of the crystal structure of phase Egg all Al and Si are located in the 102 octahedral coordination with oxygen with one unique Si and Al site each in the unit cell 103 (Schmidt et al., 1998). There are four crystallographically distinct positions for oxygen atoms-104 O(1), O(2), O(3), O(4). The O(1) and O(2) oxygen atoms are bonded to two silicon (Si) and one 105 aluminum (Al) atoms. The O(4) atom is bonded to the proton (H), two aluminum (Al) atoms and 106 a silicon (Si) atom. The structure consists of short chain segments connecting four edge-sharing 107 octahedral units in the *a*-axes direction with Al-Si-Si-Al ordering in the *a*-*c* plane. These 4member edge-sharing chains are linked in the *a*-axes direction by edge link between the AlO₆ 108 109 and the neighboring SiO_6 octahedral unit (Figure 1). These two-dimensional sheets are then 110 linked in the (011) direction through Al-Si corner linked octahedral unit. Hydrogen bonding 111 bridges corners of AlO₆ units, also linking the 2D structures of the edge-sharing octahedral units. Based on ¹H-²⁹Si NMR spectra of phase Egg, it has been suggested that there could be some 112 113 disorder of Si and Al between the octahedral sites (Xue et al., 2006).

114 We performed *first-principles* quantum mechanical calculations based on the density 115 functional theory (DFT) (Hohenberg and Kohn, 1964; Kohn and Sham, 1965). For the

116 investigation of the phase Egg, we employed a widely used approximation to the exchange-117 correlation functional: the generalized gradient approximation (GGA) and the highly accurate projector augmented wave method (PAW) as implemented in the Vienna ab initio simulation 118 119 package (VASP) (Kresse and Hafner, 1993; Kresse and Furthmüller, 1996a, b; Kresse and 120 Joubert, 1999). First, we performed a series of convergence tests by varying the energy cutoff 121 and k-points. We found that an energy cut-off $E_{cut} = 800$ eV and a k-point mesh of $6 \times 9 \times 6$ 122 Monkhorst-Pack grid (Monkhorst and Pack, 1976) with 100 irreducible k-points is sufficient for 123 describing the energetics of phase-Egg (Supplementary Figure SF1). We also used the van der 124 Waals corrections as implemented in VASP (Grimme, 2006). All computations were performed 125 on the primitive unit cell of phase Egg with the starting guess from X-ray diffraction crystal 126 structure (Schmidt et al., 1998). To test the assertion that the Al- and Si-octahedral sites may be 127 disordered (Xue et al., 2006), we have also explored the effect of octahedral cation disorder in 128 which we exchanged 25 % and 50% of the Al and Si octahedral sites.

129 The elastic tensor was derived using the finite strain approach. Starting from the 130 optimized cell at a given pressure, we have strained the lattice parameters by 1% and then relaxed the atomic positions and computed the corresponding stress tensor. Using Hooke's law, 131 132 the elastic constants were derived from the stress-strain relations, as outlined in previous studies 133 (Chheda et al., 2014). Finally, we computed the single crystal azimuthal anisotropy for compressional velocity (AV_P) i.e., $\left(\frac{(V_P^{max} - V_P^{min})}{(V_P^{max} + V_P^{min})} \times 200\right)$ and shear velocity (AV_S) i.e., 134 $\left(\frac{\left(V_{S}^{max}-V_{S}^{min}\right)}{\left(V_{S}^{max}+V_{S}^{min}\right)}\times 200\right)$ of phase-Egg using the formulation for maximum polarization anisotropy 135 (Mainprice, 1990). 136

138 Results

139 Crystal Structure

140 The converged structure has a monoclinic symmetry and space group (#14) $P2_1/c$. The 141 crystallographic space group P_{21}/n is an alternate setting for P_{21}/c . In spacegroup P_{21}/n the 142 translation is along a diagonal on the plane perpendicular to 2_1 -screw axis whereas the translation 143 is along the c-axis on the plane perpendicular to the 2_1 -screw axis. We note that the disordered phase has higher enthalpy than the ordered state by 0.5 eV, or a Gibbs free energy difference 144 145 between the ordered and the disordered states at 1800 K of ~0.2 eV/molecular unit (19 kJ/mol). 146 The enthalpy and Gibbs free energy difference remains largely insensitive to pressures 147 (Supplementary Figure SF2). Since the energetic cost of disordering is slightly greater than the 148 ordered state, we have used the ordered crystal structure for the determination of the full elastic 149 constant tensor.

150 At ambient conditions, the crystal structure of phase Egg has a single proton (H) with fractional coordinates of x_H =0.796, y_H = 0.553, z_H =0.422. The proton forms a hydroxyl O(4)-H 151 152 bond with the O(4) atom with fractional coordinate of $x_{04}=0.760$, $y_{04}=0.215$, $z_{04}=0.129$. The 153 proton also forms a hydrogen bond H....O with the O(3) atom with fractional coordinate of $x_{03}=0.880$, $y_{03}=0.201$, $z_{03}=0.512$ (Schmidt et al., 1998). Our results indicate that the O(4)-154 155 H...O(3) hydrogen bond is strengthened at high pressures, the O(4)-H bond is enhanced and H....O(3) length decreases. At around ~ 15 GPa the proton flips its position and is transferred to 156 157 the O(3) atom forming an O(3)-H group and an O(4)...H bond (Figure 2). This is slightly 158 different from more conventional symmetrization of hydrogen bonds at high pressures where the 159 proton is at the center of the O....O forming a O-H-O unit. At pressures lower than the proton 160 transfer ~ 15 GPa, we define the crystal structure as the phase Egg (*loP*), and at higher pressure,

we define the crystal structure as phase Egg (hiP), where "loP" and "hiP" refers to low- and 161 162 high-pressures respectively. The energetics and properties of the phase Egg (*hiP*) extrapolated to 163 pressures lower than 15 GPa differs significantly from that of phase Egg (loP). At pressures >15 164 GPa the structures are identical (Figure 2). Our pressure volume results and the evolution of 165 lattice parameter upon compression are in good agreement with previous experiments 166 (Vanpeteghem et al. 2003). High-pressure symmetrization of hydrogen bond has been reported in 167 many hydrous minerals including ice phases (Goncharov et al., 1996); δ-AlOOH (Tsuchiya et 168 al., 2002; Panero and Stixrude, 2004; Tsuchiya and Tsuchiya, 2009; Sano-Furukawa et al., 2010; 169 Mashino et al., 2016); δ -CrOOH analog to the δ -AlOOH phase (Jahn et al., 2012); ϵ -FeOOH 170 (Otte et al., 2009); phase D (Tsuchiya et al., 2005; Hushur et al., 2011); phase H (Tsuchiya and 171 Mookherjee, 2005; Panero and Caracas, 2017) and phase Pi (Peng et al., 2017). For symmetric 172 hydrogen bonds, often the O-H...O angle is $\sim 180^{\circ}$ or close to a linear arrangement. However, for 173 phase Egg we notice that the O(4)-H...O(3) angle to be around 160° at pressures where proton 174 transfer occurs (~ 15 GPa), leading to a new configuration of O(4)...H-O(3) (Figure 2). In a 175 recently examined hydrous aluminosilicate phase Pi, similar observation was made, few of the 176 O-H...O configurations remained non-linear even after compressions, however the O-H...O 177 configuration that exhibited symmetric hydrogen bonding at high pressure became linear (Peng 178 et al., 2017). In addition, a new monoclinic phase of AlOOH has been predicted which has non-179 linear and bent O-H-O linkages instead of symmetric hydrogen bonding at high pressures (Zhong 180 et al., 2016).

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182

184 *Equation of state*

The volume dependence of the total energy obtained using first principles simulation are adequately described by a third order Birch-Murnaghan equation of state (Birch, 1978), where total energy, E is expressed as

188
$$E = E_0 + \frac{9}{2} K_0 V_0 [f_V^2 + (K_0' - 4) f_V^3]$$
(1)

189 where, E_0 , K_0 , V_0 , and K'_0 represents the ground state energy, bulk modulus at zero-pressure, 190 unit-cell volume at zero pressure, and pressure derivative of zero-pressure bulk modulus 191 respectively. The Eulerian finite strain (f_V) and is defined as

192
$$f_V = \frac{1}{2} \left(\left(\frac{V_0}{V} \right)^{\frac{2}{3}} - 1 \right)$$
 (2)

We find that the phase Egg (loP) has lower energy than that of the phase Egg (hiP). The energies 193 194 of the two crystal structures approach each other gradually on expansion, becoming indistinguishable at a volume ~196 Å³ corresponding to a pressure ~ 15 GPa. The energy vs. 195 volume relationship of phase Egg (loP) and phase Egg (hiP) suggest that at lower pressures the 196 197 asymmetric arrangement of hydrogen i.e., O(4)-H...O(3) is favored, at pressures greater than 15 198 GPa, the proton flips its position to O(4)...H-O(3). The energy vs. volume relationship suggests 199 a second- order phase transition (Figure 3). Similar second order transitions characterized by 200 proton-order-disorder behavior and/or hydrogen bond symmetrization have been documented for 201 condensed phases including several H₂O-ice phases (Schweizer and Stillinger, 1980; Lee et al., 202 1993), dense hydrous magnesium silicate minerals phase- D (Tsuchiya et al., 2005), phase-H 203 (Tsuchiya and Mookherjee, 2015), and δ -AlOOH aluminum oxy-hydroxide (Tsuchiya et al., 2002; Panero and Stixrude, 2004; Tsuchiya and Tsuchiya, 2009). Such second order transitions 204 205 are often associated with discontinuity in second order derivatives of energy, such as bulk

modulus (Tsuchiya et al., 2002; Sano-Furukawa et al., 2009; Hushur et al., 2011). In contrast, the
pressure-volume relations for phases exhibiting second order transitions related to
symmetrization of hydrogen bonds are often continuous with slight or no difference in the zero
pressure volume between the low and high-pressure phases (for e.g., phase-D: Hushur et al.,
2011).

The bulk modulus at zero-pressure (K_0^{loP}) , unit-cell volume at zero pressure (V_0^{loP}) and 211 the pressure derivative of zero-pressure bulk modulus (K'_0^{loP}) for phase Egg (loP) i.e., at 212 pressures below the transfer of proton are, 164.4 (\pm 1.8) GPa, 210.21 (\pm 0.14) Å³, and 7.14 (\pm 0.24) 213 respectively. The bulk modulus at zero-pressure (K_0^{hiP}) , unit-cell volume at zero pressure (V_0^{hiP}) 214 and the pressure derivative of zero-pressure bulk modulus (K'_0^{hiP}) for phase Egg (hiP) at 215 pressures above the transfer of proton are, 222.8 (± 0.2) GPa, 207.74 (± 0.01) Å³, and 4.44 (± 0.02) 216 respectively. The zero pressure volume of the loP phase Egg is ~ 1.2 % larger than the hiP phase 217 Egg. In contrast, the zero pressure bulk modulus of *loP* phase Egg is ~35.5 % softer than the *hiP* 218 219 phase Egg. This behavior is similar to what has been observed in equation of state studies in phase-D where the zero pressure volume of the hydrogen off-centered (HOC) phase is $\sim 2.1\%$ 220 221 larger than the hydrogen centered (HC) phase (Hushur et al., 2011). In contrast the zero pressure 222 bulk modulus of the HC phase is $\sim 18\%$ stiffer than the HOC phase (Hushur et al., 2011).

The experimentally determined equation of state parameters for phase Egg, i.e., (K_0^{exp}) , (V_0^{exp}) and (K_0^{rexp}) are, 157.0 (±4.0) GPa, 212.99 (±0.01) Å³, and 6.5 (±0.4) respectively (Vanpeteghem et al., 2000). The unit-cell volume at zero pressure for phase Egg (*loP*) determined by *first-principles* simulation is 1.3 % smaller than the experimental results. Whereas, the bulk modulus at zero-pressure determined by *first-principles* simulation is 4.7 % greater than the experimental results. This is expected since the experiments are at 300 K and the 229 *first principles* simulations are at static conditions (0 K). The variation of the lattice parameters, 230 i.e., a, b, c, β with respect to the unit-cell volume for the low pressure and the high-pressure 231 phase Egg determined using *first-principles* simulations agree with the experimental results 232 (Figure 3, Supplementary Figure SF3). The linear moduli at zero-pressure for the lattice parameter- a, b, and $c^*=c \propto sin\beta$ i.e., K_a^{loP} , K_b^{loP} , and $K_{c^*}^{loP}$ for the phase Egg (loP) are 975.6 233 234 (±1.2) GPa, 344.6 (±0.6) GPa, and 531.1 (±0.8) GPa respectively. The linear moduli at zeropressure for the lattice parameter- a, b, and c^* i.e., K_a^{hiP} , K_b^{hiP} , and $K_{c^*}^{hiP}$ for the phase Egg (hiP) 235 are 815.6 (±2.8) GPa, 507.9 (±0.7) GPa, and 789.6 (±1.1) GPa (Supplementary Figure SF4). 236 The proton transfer stiffens with $K_b^{hiP} > K_b^{loP}$ and $K_{c*}^{hiP} > K_{c*}^{hiP}$ by ~ 32 %. This stiffening is likely 237 related to the proton transfer and the significant changes of compressibility along the b- and c^* -238 lattice directions are related to the orientation of the bent and non-linear O-H...O bonds. Upon 239 240 compression, of phase Egg exhibit anomalous elastic behavior at ~15 GPa, most likely related to 241 the observed proton transfer. Our results on equation of state and evolution of lattice parameters 242 with pressure are in good agreement with previous experimental results (Vanpeteghem et al. 243 2003). The experimental results showed a strong anisotropic behavior with larger compressibility 244 along the *b*-axis compared to the *a*- and *c*- axes directions. The measured change in the slope of 245 compressibility in *b*-direction with significant increase in stiffness above ~ 19 GPa, was 246 identified as a phase transition due to a hydrogen bond symmetrization (Sikka, 2007).

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248 Elasticity

Monoclinic phase Egg ($P2_1/c$) has 13 non-zero independent elastic constants. Upon compression almost all the individual components of the full elastic constant tensor becomes stiffer except the off-diagonal components c_{25} , c_{35} , and c_{46} (Figure 4). The pressure dependence of the elasticity data can be described with the finite strain formulation (Karki et al., 2001)

254
$$c_{ijkl} = (1 + 2f_V)^{\frac{1}{2}} [c_{ijkl0} + b_1 f_V + 2b_2 f_V^2] - P\Delta_{ijkl}$$
 (3)

255 where, f_V is the finite Eulerian strain as defined in eq (2)

256
$$b_1 = 3K_0(c'_{ijkl0} + \Delta_{ijkl}) - 7c_{ijkl0},$$
 (4)

257
$$b_2 = 9K_0^2 c_{ijkl0}'' + 3K_0' (b_1 + 7c_{ijkl0}) - 16b_1 - 49c_{ijkl0},$$
 (5)

258 and
$$\Delta_{ijkl} = -\delta_{ij}\delta_{kl} - \delta_{ik}\delta_{jl} - \delta_{il}\delta_{jk}$$
 (6)

where c'_{ijkl0} and c''_{ijkl0} are the first and second derivatives of c_{ijkl0} , with respect to pressure at ambient conditions. Δ_{ijkl} is equal to -3 for the principal constants (c_{iiii} in full tensor and c_{ii} in Voigt notation, with i=1, 2, 3), -1 for the off-diagonal elastic constants (c_{iijj} in full tensor and c_{ij} in Voigt notation, with $i=1, 2, 3, i\neq j$), -1 for the shear constants (c_{ijij} in full tensor notation with $i=1, 2, 3, i\neq j$ and c_{ij} in Voigt notation with i=4, 5, 6, i=j), and 0 otherwise. δ_{ij} is the Kronecker delta ($\delta_{ij}=1$ for i=j, and $\delta_{ij}=0$, for i \neq j).

265 At around ~15 GPa, i.e., pressure corresponding to mantle transition zone, many 266 components of the full elastic constant tensor exhibit anomalous behavior. At pressure greater 267 than 15 GPa, all components of the elastic constants become stiffer. This anomalous behavior in 268 elasticity is very likely related to the changes in the hydrogen bonding, in particular the proton 269 attached to O(4) atom flips to the O(3) atom. We analyzed the elasticity phase-Egg (*loP*) for 270 pressures below the proton transfer (\sim 15 GPa). And we also determined the elasticity of phase 271 Egg (*hiP*) for pressure > 15 GPa. We note that at low pressures, there is significant difference 272 between several components of the elastic constant tensor between the phase-Egg (loP) and the 273 metastable extension of elasticity of phase Egg (hiP) to pressure below 15 GPa. At pressures 274 greater than 15 GPa, elasticity of phase Egg (*loP*) and phase Egg (*hiP*) are identical (Figure 4). At zero-pressures conditions the principal elastic constants exhibit the relationship: $c_{11}^{loP} >$ 275 $c_{33}^{loP} > c_{22}^{loP}$. The same relationship holds for the zero-pressure metastable extension of high-276 pressure phase Egg (*hiP*), i.e., $c_{11}^{hiP} > c_{33}^{hiP} > c_{22}^{hiP}$. We also note that the $c_{11}^{hiP} \sim c_{11}^{loP}$, $c_{33}^{hiP} > c_{33}^{loP}$ 277 by 14 % and $c_{22}^{hiP} > c_{22}^{loP}$ by 30 %. This consistent with the fact that the orientation of the 278 279 hydrogen bond is mostly aligned along the *b*-axis but is tilted such that it has component along the *c*-axis of the crystal structure. The crystallographic *a*-axis direction remains largely 280 281 unaffected by the proton transfer. We also note that at pressures lower than the proton transfer pressures (~15 GPa), the pressure derivatives for the principal elastic constants $c_{ii}^{\prime loP} > c_{ii}^{\prime hiP}$, by 282 283 19 %, 29 %, and 58 % for i = 1, 2, and 3 respectively.

284 For the shear elastic components i.e., i = 4, 5, 6, i = j, at zero pressure conditions, the following relations holds $c_{44}^{loP} < c_{66}^{loP} < c_{55}^{loP}$. Similar relationship holds for the zero-pressure 285 metastable extension of phase Egg (*hiP*): $c_{44}^{hiP} < c_{66}^{hiP} < c_{55}^{hiP}$. However, unlike the principal 286 287 components, the shear components have different relationship between the low-pressure phase Egg (*loP*) and the metastable extension of high-pressure phase Egg (*hiP*), $c_{44}^{hiP} \sim c_{44}^{loP}$, $c_{66}^{hiP} < c_{66}^{loP}$ 288 by 1.8 % and $c_{55}^{hiP} < c_{55}^{loP}$ by 2.4 %. We also note that at P < 15 GPa, the pressure derivatives for 289 the shear elastic constants $c_{ij}^{\prime loP} > c_{ij}^{\prime hiP}$, by 33 %, 28 %, and 28 % for i = 4, 6, and 5, (i290 = i) respectively. 291

Our results show that the phase Egg exhibits significant elastic anisotropy (**Figure 5**). The P-wave azimuthal and S-wave anisotropy for phase Egg for the phase Egg (*loP*) are $AV_P \sim$ 30% and $AV_S \sim 17$ % at 0 GPa. The anisotropy reduces at high pressure and then increases closer to the proton transition pressures ~15 GPa. At pressures greater than 15 GPa, $AV_P \sim 10\%$ and $AV_S \sim 10\%$ and remains unchanged at higher pressures (**Figure 5**). 297 The isotropic bulk (K) and shear (G) moduli are determined using the relations

298
$$K_V = \left(\frac{1}{9}\right) [c_{11} + c_{22} + c_{33} + 2(c_{12} + c_{13} + c_{23})]$$
 (7)

299
$$K_R = [s_{11} + s_{22} + s_{33} + 2(s_{12} + s_{13} + s_{23})]^{-1}$$
 (8)

300
$$G_V = \left(\frac{1}{15}\right) [c_{11} + c_{22} + c_{33} - (c_{12} + c_{13} + c_{23}) + 3(c_{44} + c_{55} + c_{66})]$$
 (9)

301 and
$$G_R = 15[4(s_{11} + s_{22} + s_{33}) - (s_{12} + s_{13} + s_{23}) + 3(s_{44} + s_{55} + s_{66})]^{-1}$$
 (10)

302 where s_{ijkl} represents the elastic compliance tenor and is related c_{ijkl} by the relation 303 $s_{ijkl}c_{klmn} = \delta_{im}\delta_{jn}$, subscript *V* and *R* represents Voigt and Reuss bounds.

The Voigt limit of the bulk (K_V^{hiP}) and shear (G_V^{hiP}) moduli at zero-pressure for phase Egg (hiP)304 are 229 GPa and 159 GPa respectively. In contrast, the Voigt limit of the bulk (K_V^{loP}) and shear 305 (G_V^{loP}) moduli at zero-pressure for low- pressure phase Egg (loP) are 201 GPa and 151 GPa 306 307 respectively. Thus the Voigt limit for the phase Egg (hiP) is greater than phase Egg (loP) by 14 % and 6 % respectively. And the Reuss limit of the bulk (K_R^{hiP}) and shear (G_R^{hiP}) moduli for the 308 phase Egg (hiP) are 223 GPa and 157 GPa respectively. In contrast, the Reuss limit of the bulk 309 (K_V^{loP}) and shear (G_V^{loP}) moduli for the phase Egg (loP) are 165 GPa and 142 GPa respectively. 310 311 Thus the Reuss limit for the phase Egg (*hiP*) is greater than phase Egg (*loP*) by ~35 % and 10 % 312 respectively.

313

314 Discussions

The mineral phases in the Al₂O₃-SiO₂-H₂O (ASH) ternary show correlation between velocity and density. Stishovite and corundum are the densest phase and have fastest bulk sound velocity. Among the hydrous phases, phase Egg (~7.5 wt % H₂O) and δ -AlOOH (~15 wt % H₂O) are the densest and also have fast sound wave velocity (**Table 2, Figure 6**). Along the SiO₂-AlOOH join, we note that the zero pressure density follows the trend: $\rho_0^{pi} < \rho_0^{top-o} < \rho_0^{egg-loP} <$

 $\rho_0^{\delta-loP} < \rho_0^{st}$. Also, along the SiO₂-AlOOH join, the compressional velocity (V_P) extrapolated to 320 zero pressure, the following trend is observed $V_p^{pi} < V_P^{top-o} < V_P^{\delta-loP} < V_P^{egg-loP} < V_P^{st}$. 321 Similar trends for shear (V_S) velocity extrapolated to zero pressure i.e., $V_S^{pi} < V_S^{top-o} <$ 322 $V_S^{\delta-loP} < V_S^{egg-loP} < V_S^{st}$ is also observed (**Table 2**, **Figure 6**). The observed trend in the 323 324 density, and shear wave velocity could be explained in terms of the crystal structures. For 325 instance, the crystal structure of phase-Pi, that is stable at relatively low pressures between 2 and 326 7 GPa (Wunder et al., 1993a,b), consists of layers of distorted eight-membered rings formed by 327 AlO_6 units alternating with layers consisting of SiO_4 tetrahedral units (Peng et al., 2017). Crystal 328 structure of topaz-OH, that is stable up to 12 GPa (Ono, 1998), also consists of edge sharing 329 AlO₆ octahedral units forming crankshaft chains and SiO₄ tetrahedral units sharing corners. In 330 contrast, phase Egg is formed at higher pressures and has both aluminum and silicon in 331 octahedral coordination. The zero pressure density of phase Egg (*loP*) is slightly greater than the expected density based on an ideal solid solution of SiO₂ and δ -AlOOH (*loP*). The δ -AlOOH 332 333 (loP) refers to the hydrogen in off-center position (HOC), i.e., asymmetric hydrogen bond at 334 low- pressure condition (Tsuchiya and Tsuchiya, 2009). The compressional velocity (V_P) and 335 shear velocity (V_S) extrapolated to zero pressure for phase Egg (loP) is lower than the expected V_P and V_S based on an ideal solid solution of SiO₂ and δ -AlOOH (loP) by 6.3 and 6.5 % 336 337 respectively (Figure 6).

338

339 Implications

In a simplified ASH system, at the base of the mantle transition zone and along a warm
subduction zone geotherm phase Egg decomposes to a mixture of aluminum oxyhydroxide (δAlOOH) and stishovite (SiO₂) via the reaction-

343
$$AlSiO_3(OH) = \delta -AlOOH + SiO_2$$
 (11)

344 (phase Egg (*hiP*)) (aluminum oxyhydroxide) (stishovite)

345 This is likely to be associated with ΔV_P and ΔV_S of 0.42 % and -1.23 % respectively. Among all 346 the hydrous phases in the ASH ternary system, Phase Egg has one of the fastest compressional 347 (V_P) and shear (V_S) velocity (Figure 7). The velocity of phase Egg is slower than stishovite but 348 remains faster than δ -AlOOH phase at all the pressures within phase-Egg's thermodynamic 349 stability field. The δ -AlOOH phase with symmetric hydrogen bonding has sound wave velocities 350 faster than phase-Egg, however that is likely to occur at pressures corresponding to the lower mantle. In our analysis, we have used compressional (V_P) and shear (V_S) velocity of δ -AlOOH 351 352 (Tsuchiya and Tsuchita, 2009) and stishovite (Karki et al., 1997). The static sound wave velocity 353 results from the *first-principles* studies, are in very good agreement with experimental studies 354 including static Brillouin scattering experiments on stishovite (Jiang et al., 2009) and δ-AlOOH 355 (Mashino et al., 2016).

356 The decomposition of phase Egg at the base of the transition zone along a warmer subduction zone geotherm is ~ 1000 °C (Wirth et al., 2007). It is likely that the temperature will 357 358 affect the bulk sound velocity of all the hydrous phases including phase Egg. The effect of 359 temperature on the elasticity and seismic velocities of the hydrous phases in ASH ternary 360 remains mostly unknown, except for stishovite (Yang et al., 2014). The effect of temperatures on 361 compressional (V_P) and shear (V_S) velocity for natural topaz containing both hydrogen and fluorine have been recently constrained and are of the order of $\frac{dV_P}{dT} \sim -3.10 \times 10^{-4}$ and $\frac{dV_S}{dT} \sim -2.30$ 362 x 10^{-4} (Tennakoon et al., 2018). The effect of temperature on the elasticity of phase Egg (*loP*) 363 364 and its high-pressure polymorph, phase Egg (hiP) across the proton transfer needs to be 365 constrained in future work. Presence of minor amount of fluorine $(F)^{-}$ have been speculated in a

mantle minerals in transition zone (Grützner et al., 2018). Substitution of OH-1 by fluorine (F⁻)
occurs in natural minerals including topaz and recent study shows that such substitution does
indeed influence the bulk sound velocity (Tennakoon et al., 2018).

369 The sound wave velocity of phase Egg is significantly faster than the major mantle 370 phases such as wadsleyite (Zha et al., 1997), ringwoodite (Li, 2003), pyrope (Sinogeikin and 371 Bass, 2000), and majoritic garnet (Irifune et al., 2008) stable in the transition zone. Few wt % of 372 water tends to reduce the seismic velocity of wadsleyite (Mao et al., 2008) and ringwoodite 373 (Panero, 2010; Wang et al., 2006). Whereas, phase Egg could host almost 7.5 wt% of water and 374 vet have significantly faster velocity compared to hydrous wadslevite (Mao et al., 2008) and 375 ringwoodite (Wang et al., 2006). Although the presence of phase Egg is likely to be limited 376 mostly in the subducted sediments, it could store several wt % of water and be stable in a normal mantle geotherm and may not be readily detectable owing to its faster velocity compared to the 377 378 ambient mantle. Certainly, more research is warranted to have better constraints on the effect of 379 temperature, anisotropy, petrographic fabric or lattice preferred orientations on the seismic 380 velocity of phase Egg and other mineral phases in subducted sediments.

381

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391 References

- Abe, R., Shibazaki, Y., Ozawa, S., Ohira, I., Tobe, H., and Suzuki, A. (2018) In-situ X-ray
 diffraction studies of hydrous aluminosilicate at high pressure and temperature. Journal of
 Mineralogical and Petrological Sciences, 113, 106-111.
- Birch, F., (1978) Finite strain isotherm and velocities for single crystal and polycrystalline NaCl
 at high-pressures and 300 K. Journal of Geophysical Research, 83, 1257-1268.
- Bolfan-Casanova, N. (2005) Water in the Earth's mantle. Mineralogical Magazine, 69, 229-257.
- Chheda, T. Mookherjee, M., Mainprice, D., dos Santos, A. M., Molaison, J. J., Chantel, J.,
 Manthilake, G., and Bassett, W. A., (2014) Structure and elasticity of phlogopite under
 compression: Geophysical implications. Physics of the Earth and Planetary Interior,
 233,1-12.
- 402 Eggleton, R. A. Boland, J. N., and Ringwood, A. E., (1978) High pressure synthesis of a new
- 403 aluminum silicate: Al₅Si₅O₁₇(OH). Geochemichial Journal, 12, 191-194. Fukuyama, K.,
- 404 Ohtani, E., Shibazaki, Y., Kagi, H., and Suzuki, A. (2017) Stability field of phase Egg,
- 405 AlSiO₃OH at high pressure and high temperature: possible water reservoir in mantle
- 406 transition zone. Journal of Mineralogical and Petrological Sciences, 112, 31-35.
- 407 Goncharov, A.F., Struzhkin, V. V., Somayazulu, M. S., Hemley, R. J., and Mao, H. K. (1996)
- 408 Compression of ice to 210 gigapascals: Infrarerd evidence for a symmetric hydrogen409 bonded phase. Science, 273, 218-220.
- 410 Grimme, S. (2006) Semiempirical gga-type density functional constructed with a long-range
- 411 dispersion correction. Journal of Computational Chemistry, 27, 1787.

412	Grützner, T., Klemme, Rohrbach, A., Grevasoni, F., and Berndt, J. (2018) The effect of fluorine
413	on the stability of wadsleyite: Implication for the nature and depths of the transition zone
414	in the Earth's mantle. Earth and Planetary SciencesLetters, 482, 236-244.
415	Hirschmann, M.(2006) Water, melting, and the deep Earth H ₂ O cycle. Annual Reviews of Earth
416	and Planetary Sciences, 34, 629-653.
417	Hohenberg, P., Kohn, W. (1964) Inhomogeneous electron gas. Physical Review B, 136, B864-
418	871.
419	Hushur, A., Manghnani, M. H., Smyth, J. R., Williams, Q., Hellebrand, E., Lonappan, D., Ye,
420	Y., Dera, P., and Frost, D. J. (2011) Hydrogen bond symmetrization and equation of state
421	of phase D. Journal of Geophysical Research, 116, B06203.
422	Ichikawa, H., Kawai, K., Yamamoto, S., and Kameyama, M. (2015) Effect of water on
423	subduction of continental materials to the deep Earth, in "The Earth's Heterogenous
424	Mantle" edited by Khan, A. and Deschamps, F. 275-299.
425	Irifune, T., Higo, Y., Inoue, T., Kono, Y., Ohfuji, H., and Funakoshi, K. (2008) Sound velocities
426	of majorite garnet and composition of the mantle transition region. Nature, 451, 814-817.
427	Jahn, S., Wunder, B., Koch-Müller, M., Tarrieu, L., Pöhle, M., Watenphul, A., and Taran, M.N.
428	(2012) Pressure-induced hydrogen bond symmetrisation in guyanaite, β -CrOOH:
429	evidence from spectroscopy and ab initio simulations. European Journal of Mineralogy,
430	24, 839-850.
431	Jiang, F., Majzlan, J., Speziale, S., He, D., and Duffy, T.S.(2008) Single-crystal elasticity of
432	diaspore, AlOOH, to 12 GPa by Brillouin scattering. Physics of the Earth and Planetary
433	Interior, 170, 221-228.

- 434 Jiang, F., Gwanmesia, G. D.m Dyuzheva, T. I., and Duffy, T. S. (2009) Elasticity of stishovite
- and acoustic mode softening under high pressure by Brillouin scattering. Physics of theEarth and Planetary Interior, 172, 235-240.
- Karki, B.B., Stixrude, L., Crain, J. (1997) Ab initio elasticity of three high-pressure polymorphs
 of silica. Geophysical Research Letters, 24, 3269–3272.
- 439 Karki, B.B., Stixrude, L., Wentzcovitch, R.M. (2001) High-pressure elastic properties of major
- 440 materials of Earth's mantle from first principles. Reviews of Geophysics, 39, 507–534.
- 441 Katahara, K.W.(1996) Clay mineral elastic properties. SEG Annual Meeting Expanded abstracts,
- 442 1691-1694 http://dx.doi.org/10.1190/1.1826454.
- Kawamoto, T. (2006) Hydrous phases and water transport in the subducting slab. Reviews Of
 Mineralogy and Geochemistry, 62, 85-116.
- Kohn, W., and Sham, L. J.(1965) Self-consistent equations including exchange and correlation
 effects. Physical Review, 140, A1133-A1138.
- Kresse, G., and Furthmüller, J. (1996a) Efficiency of ab-initio total energy calculations for
 metals and semiconductors. Computational Material Sciences, 6, 15-50.
- Kresse, G., and Furthmüller, J. (1996b) Efficient iterative schemes for ab initio total-energy
 calculations using plane-wave basis set. Physical Review B 54, 11169-11186.
- Kresse, G., and Hafner, J.(1993) Ab initio Molecular Dynamics for Liquid-Metals. Physical
 Review B 47, 558-561.
- Kresse, G., Joubert, D. (1999) From ultrasoft pseudopotentials to the projector augmented-wave
 method. Physical Review B, 59, 1758-1775.
- Lee, C., Vanderbilt, D., Laasonen, K., Car, R., and Parrinello, M. (1993) Ab initio studies on the
 structural and dynamical properties of ice. Physical Review B, 47, 4863-4872.

- 457 Li, B. (2003) Compressional and shear wave velocities of ringwoodite γ -Mg₂SiO₄ to 12 GPa.
- 458 Americal Mineralogist, 88, 1312-1317.
- 459 Mainprice, D. (1990) An efficient FORTRAN program to calculate seismic anisotropy from the
 460 lattice preferred orientation of minerals. Computational Gesoscience, 16, 385-393.
- 461 Mao, Z., Jacobsen, S. D., Jiang, F., Smyth, J. R., Holl, C. M., and Duffy, T. S. (2008) Elasticity
- 462 of hydrous wadsleyite to 12 GPa: Implications for Earth's transition zone. Geophysical
 463 Research Letters, 35, L21305, doi:10.1029/2008GL035618.
- 464 Mashino, I., Murakami, M., and Ohtani, E. (2016) Sound velocities of δ-AlOOH up to core-
- 465 mantle boundary pressures with implications for the seismic anomalies in the deep
- 466 mantle. Journal of Geophysical Research- Solid Earth, 121, 595–609,
- 467 doi:10.1002/2015JB012477
- Mei, S., and Kohlstedt, D. L. (2000) Influence of water on plastic deformation of olivine
 aggregates 1. Diffusion creep regime. Journal of Geophysical Research, 105, 21,45721,469.
- 471 Monkhorst, H. J., and Pack, J. D. (1976) Special points for Brillouin-zone integrations. Physical
 472 Review B, 13, 5188-5192.
- 473 Mookherjee, M., Speziale, S., Marqardt, H., Jahn, S., Wunder, B., Koch-Müller, M., Liermann,
- 474 H.-P. (2015) Equation of state and elasticity of the 3.65 Å phase: implications for the X475 discontinuity. American Mineralogist, 100, 2199-2208.
- 476 Mookherjee, M., Tsuchiya, J., and Hariharan, A. (2016) Crystal structure, equation of state, and
- 477 elasticity of hydrous aluminosilicate phase, topaz-OH $(Al_2SiO_4(OH)_2)$ at high pressures.
- 478 Physics of the Earth and Planetary Interior, 251, 24-35.
- 479 Nye, J. F. (1985) Physical properties of crystals. Oxford University Press, Clarendon.

- 480 Ono, S. (1998) Stability limits of hydrous minerals in sediment and mid-ocean ridge basalt
 481 compositions: Implications for water transport in subduction zones. Journal of
 482 Geophysical Research, 103, 18253-18267.
- 483 Ono, S. (1999) High temperature stability of phase egg, AlSiO₃(OH). Contribution to
 484 Mineralogy and Petrology, 137, 83-89.
- Otte, K., Pentcheva, R., Schmal, W.W., and Rustad, J. R. (2009) Pressure-induced structural and
 electronic transitions in FeOOH from first principles. Physical Review B. 80, 205116.
- 487 Pamato, M.G., Myhill, R., Boffa Ballaran, T., Frost, D.J., Heidelbach, F. and Miyajima, N.
- 488 (2015) Lower-mantle water reservoir implied by the extreme stability of hydrous489 aluminosilicate. Nature Geosciences, 8, 75-79.
- Panero, W. R. and Stixrude, L. (2004) Hydrogen incorporation in stishovite at high pressure and
 symmetric hydrogen bonding in δ-AlOOH. Earth and Planetary Science Letters, 221,
 421-431.
- 493 Panero, W. R. (2010) First principles determination of the structure and elasticity of hydrous

494 ringwoodite. Journal of Geophysical Research, 115, B03203,

- doi:10.1029/2008JB006282.
- 496 Panero, W. R., and Caracas, R. (2017) Stability of phase H in the MgSiO₄H₂-AlOOH-SiO₂
 497 system. Earth and Planetary Science Letters., 463, 171-177.
- 498 Peacock, S. M. (1990) Fluid processes in subduction zones. Science, 248, 329-337.
- 499 Peng, Y., Mookherjee, M., Hermann, A., Bajgain, S., Liu, S., and Wunder, B. (2017) Elasticity
- 500 of phase-Pi (Al₃Si₂O₇(OH)₃)- A hydrous aluminosilicate phase. Physics of the Earth and
- 501 Planetary Interior, 269, 91-97, https://doi.org/10.1016/j.pepi.2017.05.016.

- 502 Perdew, J. P., Burke, K., and Erzerhof, M. (1996) Generalized gradient approximation made
 503 simple. Physical Review Letters, 77, 3865-3868.
- Sano, A., Ohtani, E., Kubo, T. and Funakoshi, K.-i. (2004) In situ X-ray observation of
 decomposition of hydrous aluminum silicate AlSiO₃OH and aluminum oxide hydroxide
 δ-AlOOH at high pressure and temperature. Journal of Physics and Chemistry of Solids,
- **507 65**, 1547-1554.
- 508 Sano-Furukawa, A., Kagi, H., Nagai, T., Nakano, S., Ushijima, D., Iizuka, R., Ohtani, E., and
- 509 Yagi, T. (2009) Change in compressibility of δ -AlOOH and δ -AlOOD at high pressure:
- A study of isotope effect and hydrogen-bond symmetrization. American Mineralogist, 94,
 1255-1261.
- Schmidt, M. W., Finger, L. W., Angel, R. J., and Robert, E. (1998) Synthesis, crystal structure
 and phase relations of AlSi₃OH, a high pressure hydrous phase. American Mineralogist,
 83, 881-888.
- 515 Schreyer, W. (1995) Ultradeep metamorphic rocks: The retrospective viewpoint. Journal of
 516 Geophysical Research, 100, 8353-8366.
- 517 Schweizer, K. S., and Stillinger, F. H. (1980) High pressure phase transitions and hydrogen-bond
 518 symmetry in ice polymorphs. Journal of Chemical Physics, 80, 1230-1240.
- 519 Sikka, S.K. (2007) On some bond correlations at high pressures. High Pressure Research, 27,
 520 313-319.
- 521 Sinogeikin, S.V., and Bass, J.D. (2000) Single-crystal elasticity of pyrope and MgO to 20 GPa
- by Brillouin scattering in the diamond cell. Physics of the Earth and Planetary Interior,120, 43-62.

- Smyth, J.(2006) Hydrogen in high-pressure silicate and oxide mineral structures. Reviews of
 Mineralogy and Geochemistry, 62, 85-116.
- 526 Tennakoon, S., Peng, Y., Mookherjee, M., Speiale, S., Manthilake, G., Besara, T., Andreu, L.,
- and Rivera, F. (2018) Single crystal elasticity of natural topaz at high-temperatures.
 Scientific Reports, 8, 1372, doi:10.1038/s41598-017-17856-3.
- Tsuchiya, J., Tsuchiya, T., Tsuneyuki, S., and Yamanaka, T. (2002) First principles calculation
 of high-pressure hydrous phase, δ-AlOOH. Geophysical Research Letters, 29, 1909.
- Tsuchiya, J., Tsuchiya, T., and Tsuneyuki, S. (2005) First-principles study of hydrogen bond
 symmetrization of phase D under high pressure. American Mineralogist, 90, 44-49.
- Tsuchiya, J., and Tsuchiya, T. (2009) Elastic properties of δ-AlOOH under pressure: First
 principles investigation. Physics of the Earth and Planetary Interior, 174, 122-127.
- 535 Tsuchiya, J., and Mookherjee, M. (2015) Crystal structure, equation of state, and elasticity of
- phase H (MgSiO₄H₂) at Earth's lower mantle pressures. Scientific Reports, 5, 15534,
 <u>https://doi.org/10.1038/srep15534</u>.
- Tulk, C.A., Gagnon, R.E., Kiefte, H., and Clouter, M.J. (1994) Elastic constants of ice III by
 Brillouin spectroscopy. Journal of Chemical Physics, 101, 2350.
- 540 Vanpeteghem, C. B. Ohtani, E. Kondo, T. Takemura, K. and Kikegawa, T. (2003)
 541 Compressibility of phase Egg AlSiO₃(OH): Equation of state and role of water at high
 542 pressure. American Mineralogist, 88, 1408-1411.
- Vaughan, M.T., and Weidner, D.J. (1978) The relationship of elasticity and crystal structure in
 andalusite and sillimanite. Physics and Chemistry of Mineralals, 3, 133–144.
- Wang, D. J., Mookherjee, M., Xu, Y., and Karato, S. (2006) The effect of water on the electrical
 conductivity of olivine. Nature, 443, 977-980.

- 547 Wang, J., Sinogeikin, S. V., Inoue, T., and Bass, J. D. (2006) Elastic properties of hydrous
- ringwoodite at high-pressure conditions. Geophysical Research Letters, 33, L14308,
 doi:10.1029/2006GL026441.
- Weidner, D.J., and Carleton, H.R. (1977) Elasticity of coesite. Journal of Geophysical Research,
 82, 1334-1346.
- 552 Winkler, B., Hytha, M., Warren, M.C., Milman, V., Gale, J.D., and Schreuer, J. (2001)
- 553 Calculation of the elastic constants of the Al₂SiO₅ polymorphs andalusite, silimanite and
 554 kyanite. Zeitschrift für Kristallographie, 216, 67–70.
- Wirth, R., Vollmer, C., Brenker, F., Matsyuk, S., and Kaminsky, F. (2007) Inclusion of
 nanocrystalline hydrous aluminum silicate "Phase Egg" in superdeep diamonds from
 Junia (Mato Grosso State, Brazil). Earth and Planetary Science Letters, 259, 384-399.
- Wunder, B., Medenbach, O., Krause, W., and Schreyer, W. (1993a) Synthesis, properties and
 stability of Al₃Si₂O₇(OH)₃ (phase Pi), a hydrous high-pressure phase in the system Al₂O₃SiO₂-H₂O (ASH). European Journal of Mineralogy, 5, 637-649.
- 561 Wunder, B., Rubie, D. C., Ross-II, C. R., Medenbach, O., Seifert, F., and Schreyer, W. (1993b)
- 562 Synthesis, stability, and properties of Al₂SiO₄(OH)₂: A fully hydrated analogue of topaz.
 563 American Mineralogist, 78, 285-297.
- Xue, X., Kanzaki, M., Fukui, H., Ito, E., & Hashimoto, T. (2006) Cation order and hydrogen
 bonding of high-pressure phases in the Al₂O₃-SiO₂-H₂O system: An NMR and Raman
 study. American Mineralogist, 91, 850-861.
- Yang, R., and Wu, Z. (2014) Elastic properties of stishovite and the CaCl₂-type silica at the
 mantle temperature and pressure: An ab inito investigation. Earth and Planetary Science
 Letters, 404, 14-21.

- 570 Zhong, X., Hermann, A., Wang, Y., and Ma, Y. (2016) Monoclinic high-pressure polymorph of
- 571 AlOOH predicted from first principles. Physical Review B, 94, 224110.

573 Figure Captions

Figure 1. Crystal structure of phase Egg: (*a*) projection down the *b*-axis, and (*b*) projection down
the *a*-axis. The octahedral units i.e., AlO₆ and SiO₆ are represented by grey and light blue
spheres respectively. The hydrogen atoms and the oxygen atoms are shown as silver and
red spheres respectively.

- Figure 2. (*a*) Plot of hydroxyl bond length: $d_{O(4)-H}$, hydrogen bond length: $d_{O(4)-H--O(3)}$, and separation of the oxygen atoms, i.e., O(4)---O(3) vs. unit-cell volume/pressure. (*b*) Hydrogen-hydrogen repulsion denoted by H---H separation distances. (*c*) The hydrogen bond angle, O(4)-H---O(3). The blue and red symbols represent the low-pressure (*loP*) and high-pressure (*hiP*) phase Egg.
- Figure 3. (a) Plot of total energy, (b) pressure, (c) lattice parameters, a (d) lattice parameter, b 583 (e) lattice parameter, c and (f) lattice parameter, β , for low-pressure (*loP*) (blue symbols) 584 585 and high-pressure (*hiP*) (red symbols) phase Egg vs. unit cell volume. The grey filled 586 symbols are from X-ray diffraction results (Vanepetghem et al., 2000). Unit-cell volume of (loP) (blue symbols) and high-pressure (hiP) (red symbols) phase Egg is also plotted 587 588 as a function of pressure and compared with experimental data (grey filled symbols) 589 (Supplementary Figure SF3). The volume/pressure dependence of the lattice parameters 590 is characteristics of a second order transition and has been documented in several hydrous 591 phases (Tsuchiya et al., 2002; 2005; Tsuchiya and Mookherjee, 2015).
- Figure 4. Full elastic constant tensor (C_{ij}) components, for phase Egg as a function of pressure (Table 1). There are 12 sub plots and an inset, representing the 13 independent elastic constants for the monoclinic symmetry. The low-pressure (loP) and high-pressure (hiP)phase Egg is represented by blue and red symbols respectively.

Figure 5. (a) Plot of elastic anisotropy AV_P and AV_S for phase Egg as a function of pressure. The low-pressure (*loP*) and high-pressure (*hiP*) phase Egg is represented by blue and red symbols respectively. The stereographic projection down the X₃ axes for AV_P , AV_S and V_S polarization planes at (**b**) ~0 GPa and (**c**) ~23 GPa.

600 **Figure 6.** (a) A ternary plot with Al_2O_3 -SiO₂-H₂O as the end member components, relevant for 601 the mineral phases stable in subducted sediments. Also shown is the blue shaded region 602 along the line joining the SiO₂ and AlOOH with a series of hydrous phases, δ -AlOOH, 603 topaz-OH, phase Egg, phase-pi and stishovite. Plot of $(b) V_P$, $(c) V_S$ as a function of 604 density for the mineral phases in the ASH ternary (Table 2). The vertical grey-brown 605 shaded line demonstrates that the density of low-pressure (*loP*) and high-pressure (*hiP*) 606 phase Egg are very similar but the velocities are distinct. This is related to the fact that second-order transition associated with the proton transfer affects the second derivative 607 608 of energy, i.e., elastic constants but the volume and density remains mostly unaffected by 609 the proton transfer. The blue and red dashed line represents linear regression fits for the compressional wave, V_P is given by $[2.77 (\pm 0.49) \times \rho + 0.07 (\pm 1.64)]$ and for the shear 610 wave, V_s is given by $[1.74 (\pm 0.27) \times \rho - 0.23 (\pm 0.91)]$. The light blue and red lines 611 612 represents 95 % confidence level for the primary and shear velocity- density regressions. 613 Plot of zero pressure (d) density (e) V_P , and V_S for hydrous mineral phase along the SiO₂ and AlOOH join. 614

Figure 7. (*a*) Plot of K_H and G_H i.e., Hill averaged (average of Voigt and Reuss limits) for lowpressure (*loP*) and high-pressure (*hiP*) phase Egg represented by blue and red symbols respectively. Also shown is an experimentally determined bulk modulus in grey symbol (Vanepetghem *et al.*, 2000). (*b*) Plot of pressure dependence of V_P and V_S for low-

619	pressure (loP) and high-pressure (hiP) phase Egg and other mineral phases including,
620	stishovite (Karki et al., 1997) topaz-OH (Mookherjee et al., 2016), phase Pi (Peng et al.,
621	2017), and δ -AlOOH (Tsuchiya and Tsuchiya, 2009) that are likely stable in subducted
622	sediments. (c) A pressure-temperature-depth phase diagram illustrating the stability of
623	phase Egg. The figure is modified from previous study (Wirth et al., 2007). Also shown
624	are the green shaded area representing the mantle transition zone. The blue shaded band
625	at \sim 15 GPa indicates the predicted pressure for the proton transfer, i.e., boundary
626	between phase Egg (loP) and phase Egg (loP) associated with the elastic anomaly as
627	observed in our study. The dashed circle at the intersection of warm subduction geotherm
628	with the high-pressure stability of phase Egg (hiP), where phase Egg will decompose to
629	d-AlOOH and stishovite is likely to be associated with ΔV_P and ΔV_S of 0.42 % and -1.23
630	% respectively.

|--|

V	Р	c_{11}^{loP}	C_{22}^{loP}	C_{33}^{loP}	c_{12}^{loP}	C_{13}^{loP}	c_{15}^{loP}	C_{23}^{loP}	c_{25}^{loP}	c_{35}^{loP}	C_{44}^{loP}	c_{55}^{loP}	c_{66}^{loP}	c_{46}^{loP}
220	-6.4	449.5	198.7	271.9	91.9	134.4	6.2	14.2	26.3	34.4	106.1	148.0	136.0	27.5
215	-3.4	473.9	243.1	345.0	92.2	135.5	6.6	56.9	18.5	25.2	130.3	161.6	147.8	22.7
210	0.2	504.7	280.4	401.0	98.6	141.6	7.5	87.9	13.5	19.8	150.3	174.0	159.7	18.6
205	4.5	538.9	314.3	448.8	109.3	152.1	9.1	113.0	10.1	16.7	168.1	186.2	171.7	15.6
200	9.8	575.4	342.7	492.7	121.9	165.4	8.8	133.6	7.3	14.9	185.5	197.3	184.3	13.1
198	12.2	590.4	334.8	508.7	125.7	171.3	9.1	135.6	7.4	14.6	192.7	201.7	189.7	12.5
196	14.7	605.0	361.7	499.1	138.1	183.7	9.8	125.9	12.7	19.2	200.7	205.4	195.6	12.3
195	16.1	608.7	467.8	555.3	143.5	188.6	4.4	192.0	-7.3	7.7	191.7	202.4	179.8	-9.7
190	23.7	653.1	519.8	600.2	169.5	212.4	5.7	221.7	-6.6	9.3	208.0	215.7	192.0	-7.5
185	32.8	701.1	573.8	648.1	198.7	238.6	7.2	252.3	-5.6	10.8	225.7	229.6	206.0	-4.6
180	43.7	752.0	627.4	699.5	228.6	267.6	8.3	284.6	-5.6	12.0	243.7	243.2	219.6	-2.9
finite strain fit														
	c_0^{loP}	503.4	278.9	398.5	98.1	141.3	7.5	86.5	13.7	20.1	149.5	173.5	159.2	18.8
	C'_0^{hiP}	8.02	9.60	14.78	1.69	1.76	3.17	8.15	1.68	1.59	7.27	5.22	5.11	2.01
	C''_0^{hiP}	-0.141	-0.775	-1.298	0.188	0.178	-0.008	-0.839	0.168	0.222	-0.400	-0.201	-0.137	0.102
V	Р	c_{11}^{hiP}	C_{22}^{hiP}	\mathcal{C}^{hiP}_{33}	c_{12}^{hiP}	\mathcal{C}_{13}^{hiP}	c_{15}^{hiP}	c_{23}^{hiP}	c_{25}^{hiP}	c^{hiP}_{35}	c_{44}^{hiP}	c_{55}^{hiP}	C_{66}^{hiP}	c_{46}^{hiP}
220	-11.3	431.9	281.6	380.8	65.2	110.3	2.4	81.1	-1.8	7.1	115.2	143.3	136.5	-11.6
210	-2.4	494.7	347.7	443.7	90.2	134.6	3.5	118.7	-6.2	6.8	144.8	165.5	153.1	-11.5
200	9.2	566.8	423.6	516.8	123.7	167.5	4.6	163.8	-8.9	8.2	175.0	189.2	171.1	-11.1
190	24.3	650.9	519.7	603.8	170.9	211.6	6.2	220.4	-9.1	11.3	206.9	215.0	193.0	-8.5
180	44.0	748.9	627.3	704.2	229.8	266.7	8.2	283.3	-8.3	14.5	243.2	242.3	220.1	-4.5
finite strain fit														
	c_0^{hiP}	509.8	362.8	456.0	98.9	141.9	3.7	127.8	-6.6	7.6	150.4	169.3	156.3	-11.5
	C'_0^{hiP}	6.48	6.82	6.16	2.97	2.80	3.10	3.88	2.69	3.06	4.85	3.78	3.66	3.04
	$C^{"hiP}_0$	-0.068	-0.050	-0.038	-0.002	-0.001	-0.0001	-0.021	0.015	0.005	-0.044	-0.010	-0.015	0.007

Table 2. Density, compressional (V_P) , and shear (V_S) wave velocity of mineral phases in Al₂O₃-SiO₂-H₂O ternary.

1 2

 V_P Vs H_2O ρ (g/cm^3) Abbreviation Stoichiometry (km/s)(wt %) Mineral (km/s)Reference Al_2O_3 Ohno et al. (1986) Corundum cor 3.95 10.94 6.41 0.00 6.09 Quartz SiO₂ 2.64 4.12 0.00 Ohno et al. (2006) qz 2.93 4.58 Weidner and Carleton (1977) Coesite SiO₂ 8.17 0.00 cs Jiang et al. (2009) Stishovite 4.30 7.28 st SiO₂ 11.91 0.00 Tulk et al. (1994) 2.01 100.00 Ice ice-iii H_2O 1.16 3.66 Al₂SiO₅ Vaughan and Weidner (1978) Andalusite and 3.15 9.76 5.65 0.00 Sillimanite Al₂SiO₅ 5.42 0.00 Vaughan and Weidner (1978) sil 3.24 9.65 Winkler et al. (2001) ky Al₂SiO₅ 9.68 5.87 0.00 Kyanite 3.76 Diaspore Alooh 3.38 Jiang et al. (2008) dia 9.42 5.83 15.00 δ-ΑΙΟΟΗ AlOOH 3.39 9.82 6.19 15.00 Tsuchiya and Tsuchiya (2009) δ-loP δ-AlOOH δ-hiP AlOOH 3.47 10.87 6.60 15.00 Tsuchiya and Tsuchiya (2009) Kaolinite 18.19 Katahara (1996) kl Al₂Si₂O₅(OH)₄ 2.52 6.23 3.55 Topaz-OH $Al_2SiO_4(OH)_2$ Mookherjee et al. (2016) top-m 3.43 9.75 5.81 10.00 Topaz-OH $Al_2SiO_4(OH)_2$ Mookherjee et al. (2016) top-o 3.39 9.51 5.51 10.00 phase-pi pi Al₃Si₂O₇(OH)₃ 3.21 8.86 5.28 9.00 Peng et al. (2017) phase-Egg AlSiO₃(OH) 3.79 7.50 this study egg-loP 9.99 6.22 7.50 Phase-Egg egg-hiP AlSiO₃(OH) 3.84 10.68 6.42 this study



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