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5 **Towards the wider application of ^{29}Si NMR spectroscopy**
6 **to paramagnetic transition metal silicate minerals and glasses, 2:**
7 **Fe(II), Co(II), and Ni(II) silicates**

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Abstract

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In studies of the structures of silicate minerals and glasses, ^{29}Si NMR spectroscopy has

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been applied almost exclusively to materials containing relatively low concentrations of ions

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with unpaired electrons spins, such as most transition metals and rare earths, because of

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sometimes severe broadening and shifting of resonances in such strongly paramagnetic systems.

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However, by adapting experimental methods to allow detection of very broad signals, and by

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examining a much expanded range of frequencies, we show here that accurate spectra can indeed

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be measured for a series of pure-phase transition metal silicates, including olivines (Fe_2SiO_4 ,

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Co_2SiO_4 , Ni_2SiO_4), Co-akermanite ($\text{CoCa}_2\text{Si}_2\text{O}_7$), and clinopyroxenes ($\text{CoCaSi}_2\text{O}_6$, $\text{NiCaSi}_2\text{O}_6$).

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For the latter two, we also present data for glasses of the same nominal compositions. For all of

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these phases, NMR peak broadening is large, in many cases to the point where magic-angle

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spinning (MAS) does not enhance resolution; in all cases the observed paramagnetic shifts fall

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far outside the known range for diamagnetic silicates. There are clearly large effects of local

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structure on shift, suggesting great potential sensitivity to variations in chemical, electronic and

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magnetic structure in both crystalline and amorphous phases. In particular, the spectra for the

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glasses are very different from those of crystals in both width and position. In most cases,

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measured spin-spin relaxation times are long enough to avoid major loss of signal during the

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NMR acquisition, but this may not always be the case in more magnetically dilute solid

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solutions, where small, broad paramagnetic resonances due to first cation neighbor interactions

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may in some systems be difficult to detect. To explore this issue, we present new data on a Ni-

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doped forsterite ($\text{Mg}_{1.9}\text{Ni}_{0.1}\text{SiO}_4$) and on natural San Carlos olivine ($\text{Mg}_{1.8}\text{Fe}_{0.2}\text{SiO}_4$), which

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improve upon the accuracy of our previous studies of these materials. It is clear that applications

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of NMR to paramagnetic silicates holds great promise not only for empirical studies of structure

55 of a much wider range of compositions of minerals and glasses, but for future testing and
56 application of advanced theoretical methods to more completely interpret such results.

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Introduction

59 ²⁹Si NMR spectroscopy has been widely applied to crystalline and glassy silicates for
60 decades, most commonly with resolution greatly enhanced by magic angle sample spinning
61 (MAS) methods (Engelhardt and Michel, 1987; Kirkpatrick, 1988; MacKenzie and Smith, 2002;
62 Stebbins and Xue, 2014). However, nearly all such studies have been of materials with low (less
63 than a few %) contents of cations with unpaired electron spins, such as most transition metal and
64 rare earth ions, because these can interact strongly with the observed nuclear spins, often leading
65 to severe broadening of spectra even to the point of apparent loss of signal (Grimmer et al., 1983;
66 Hartman et al., 2007; Sherriff and Hartman, 1985). This poses an unfortunate gap in our tools to
67 investigate questions of short-range structure and order/disorder, given the wide importance of
68 transition metal, rare earth, and actinide-rich silicates in the earth sciences and for technology.

69 Some initial steps have been made in closing this gap by the discovery and quantitation
70 of well-resolved, usually small, ²⁹Si (and ²⁷Al) NMR resonances that are displaced well away
71 from typical chemical shifts by relatively low concentrations (ca. 0.1 to 20%) of cations such as
72 Fe²⁺, Co²⁺, and Ni²⁺ in garnets, olivines, zircons and other minerals (Begaudeau et al., 2012;
73 Dajda et al., 2003; McCarty et al., 2015; Palke and Stebbins, 2011b; Palke et al., 2015; Stebbins
74 and Kelsey, 2009; Stebbins et al., 2017); analogous observations have been long been known for
75 ⁸⁹Y and ¹¹⁹Sn NMR spectra of rare-earth stannate pyrochlores (Grey et al., 1989; Grey et al.,
76 1990) and have been reported recently for ¹⁷O in MgO and CaO (McCarty and Stebbins, 2016b),
77 ²⁷Al in aluminate garnets (YAG) (George et al., 2013; McCarty and Stebbins, 2016a) and ³¹P in

78 REE-doped monazites (LaPO_4) and xenotimes (YPO_4) (Palke and Stebbins, 2011a; Palke et al.,
79 2013). However, a key part of understanding the complex physics behind such “paramagnetic
80 shifts” must be their measurement in pure-phase end member compounds (e.g. transition metal
81 silicates), where site occupancies, ordering, and short-range structure are well known, reducing
82 possible ambiguities in data interpretation. Very few such studies have been published, in part
83 because of significant challenges in even recording such spectra accurately (Saji et al., 1973).
84 We have begun to make progress in this area with an initial report on paramagnetic Cu(II)
85 silicates (Stebbins, 2017); a very recent study on low temperature magnetic ordering in
86 $(\text{Ba,Sr})\text{CuSi}_2\text{O}_6$ also reported high-resolution ^{29}Si spectra (Puphal et al., 2016). In these
87 compounds, the single unpaired electron spin on Cu^{2+} resulted in only moderate peak
88 broadening, which allowed the detection of ^{29}Si NMR peaks that ranged from about +1600 ppm
89 to -250 ppm, even in a single mineral with multiple Si sites; these are far outside the known
90 range for SiO_4 groups in diamagnetic silicates (about -60 to -120 ppm).

91 The feasibility, and eventual structural utility, of NMR for paramagnetic silicates is
92 strongly suggested, and inspired by, extensive recent work on NMR of ^{31}P , ^7Li , and other
93 nuclides in lithium/transition metal oxides and phosphates of key importance in advanced battery
94 technology, which has driven the development of both sophisticated NMR methods and
95 theoretical analysis (Carlier et al., 2003; Grey and Dupré, 2004; Middlemiss et al., 2013; Pecher
96 et al., 2017; Pigliapochi et al., 2017; Strobridge et al., 2014; Tucker et al., 2002; Wilcke et al.,
97 2007; Yoon et al., 2004; Zeng et al., 2007). Some of the now well-studied Mn, Fe, Co, Ni
98 phosphate materials (most notably LiMPO_4 olivines) have structures closely analogous to
99 important silicate minerals. Although acquisition of ^{29}Si spectra in analogous silicate minerals is
100 expected to be considerably more difficult because of the much lower natural abundance of this

101 nuclide (4.7% vs. 100% for ^{31}P), and by its lower resonant frequency at a given external
102 magnetic field (about half that of ^{31}P , reducing NMR signal strength further), the experimental
103 methodology, and to some extent theoretical interpretations, can be expected to be transferable.
104 Here we report new data on pure-phase Fe, Ni, and Co silicate olivines (M_2SiO_4) and on Ni and
105 Co clinopyroxenes having the diopside structure ($\text{M}\text{CaSi}_2\text{O}_6$), as well as glasses corresponding to
106 the latter two compositions. We also present improved data on two relatively dilute olivine solid
107 solutions studied previously, $\text{Mg}_{1.9}\text{Ni}_{0.1}\text{SiO}_4$ (5% Ni forsterite) and natural San Carlos olivine
108 ($\text{Mg}_{1.8}\text{Fe}_{0.2}\text{SiO}_4$) (McCarty et al., 2015), to allow better connections between data for pure-phase
109 and dilute paramagnetic minerals.

110

111 **Paramagnetic shifts in solid-state NMR spectra**

112 The physics of interactions between unpaired electron and nuclear spins is complex and
113 in many systems remains incompletely characterized. In-depth discussions of theoretical and
114 experimental considerations have been reported, often in the context of lithium/transition metal
115 phosphates and oxides (Grey and Dupré, 2004; Pecher et al., 2017; Pigliapochi et al., 2017); and
116 are summarized in the mineralogical context (Palke and Stebbins, 2011a; Stebbins et al., 2017;
117 Stebbins and Xue, 2014). In brief, two general types of interactions can cause shifts in resonant
118 frequencies (peak positions in spectra) away from those in diamagnetic materials. The first of
119 these, which probably is predominant in the systems studied here, is the Fermi contact shift,
120 sometimes referred to as “transferred hyperfine coupling”. This effect essentially involves the
121 through-bond transfer of unpaired electron spin density from a paramagnetic cation to the NMR-
122 observed nucleus. As such, the magnitude and even the sign of the shift depend strongly on the
123 extent and nature of orbital overlap, covalency, bond distances and angles. Such shifts can thus

124 be very sensitive to short-range structure and potentially informative when structural details (e.g.
125 order/disorder) are unknown, but can be correspondingly challenging to calculate from theory.
126 Impressive progress in this problem has been presented recently, particularly for Li battery
127 materials (Clément et al., 2012; Middlemiss et al., 2013; Pigliapochi et al., 2017), offering
128 promise for future results on silicates. Paramagnetic shifts for ^{31}P in such materials can be
129 positive or negative, and can be very large, up to 1000's of ppm and contrasting with a range for
130 this nuclide of less than 100 ppm in diamagnetic phosphates (Wilcke et al., 2007). A second kind
131 of interaction in such systems is called the pseudo-contact shift, and involves a through-space
132 magnetic dipolar coupling between a nuclear spin and unpaired electron spins in asymmetric
133 sites, and is sometimes described as “classical” in the sense of not depending on a detailed
134 quantum-mechanical description of electronic structure. This kind of coupling is generally less
135 than that for Fermi contact shifts, but can clearly be important, even predominant, in some
136 relatively ionic oxide materials (Grey et al., 1990).

137 NMR resonances with large paramagnetic shifts can be missed in conventional spectra
138 unless specific issues are addressed in experimental parameters, such as the transmitter central
139 frequency and bandwidth, instrumental deadtime, and data processing. More problematically,
140 such resonances are likely to be broadened by strong couplings to unpaired electron spins. In
141 many cases, this line broadening is severe enough so that MAS does not significantly narrow the
142 resonance, although this technology is advancing to allow faster and faster spinning, albeit on
143 smaller and smaller samples. Small, structurally interesting differences in the peak position
144 (chemical shift) can be obscured, and the total NMR signal will be spread out in frequency,
145 greatly reducing obtainable signal-to-noise ratios. In other cases broadening may be so severe
146 that the resonance becomes effectively unobservable. However, even in the latter scenario, the

147 effects of the paramagnetic cations can, in principle, be detected by calibrating total observed
148 NMR signal intensities relative to a diamagnetic standard. In this report, we encounter all of
149 these pragmatic issues, and give some examples of how they can be characterized and how they
150 affect analyses of results.

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152 **Samples and experimental methods**

153 **Samples**

154 **General procedures.** Dried, reagent-grade starting materials were used throughout.
155 These were thoroughly mixed by hand grinding in agate mortars; solid state reactions (excepting
156 that for fayalite) were carried out with powders pressed into pellets using a steel die and
157 hydraulic press. Except for the fayalite, mineral and glass syntheses were done in air in Pt
158 crucibles or tubes. Electron microprobe data were collected with a JEOL JXA-8230 instrument,
159 with Ni₂Si (Ni), Fe metal (Fe), Co metal or Co₂SiO₄ (Co), and diopside (Ca, Si, Mg) standards,
160 with a 3 μm beam diameter and 2 minute counting times.

161 **Olivine solid solutions.** The San Carlos olivine (Stanford University Research Mineral
162 Collection) sample comprised several gem-quality crystals crushed to a powder for MAS NMR
163 (McCarty et al., 2015). The solid-state synthesis and characterization of a 5% Ni-doped synthetic
164 forsterite (Mg_{1.9}Ni_{0.1}SiO₄) was also described previously (McCarty et al., 2015).

165 **Fayalite (Fe₂SiO₄).** Fayalite was synthesized by reaction of powdered Fe metal and
166 quartz at 1070 °C in a CO/CO₂ gas mixing furnace, with a total reaction time of about 100 h and
167 4 intermediate grinding steps. A small fraction (ca. 5%) of excess silica was added to ensure
168 complete reaction of the Fe metal. XRD data showed the presence of only fayalite and a minor

169 amount of excess cristobalite. The resulting powder was uniformly light greenish brown in color;
170 EPMA detected only fayalite and a silica phase.

171 **Co-olivine (Co₂SiO₄).** Co₂SiO₄ was synthesized from Co₃O₄ (known to decompose to
172 CoO above about 900 °C) and amorphous silica, heated at 1200 °C for a total of 57 h with one
173 intermediate grinding step. This relatively low temperature was chosen to avoid any melting
174 (O'Neill, 1987; Sugawara and Akaogi, 2003). A fine-grained, uniformly violet-pink sample
175 resulted. EPMA confirmed the sample purity and stoichiometry, detecting 1-2 % only of
176 unreacted silica.

177 **Ni-olivine (Ni₂SiO₄).** Ni₂SiO₄ was synthesized from NiO and amorphous silica, heated
178 for a total of 240 h at 1400 °C with two intermediate grindings (Sugawara and Akaogi, 2003). A
179 fine grained, uniformly pistachio-green product was the result. EPMA confirmed the
180 stoichiometry of the olivine, with small amounts of unreacted starting materials.

181 **NiCaSi₂O₆.** Crystalline NiCaSi₂O₆ (diopside structure, (Ghose et al., 1987)) was
182 synthesized from NiO, CaCO₃ and amorphous SiO₂, heated for a total of 96 h at 1300 °C with
183 one intermediate grinding, yielding a homogeneous medium-green product (Durand et al., 1996;
184 Navrotsky and Coons, 1976; Raudsepp et al., 1990). EPMA confirmed the stoichiometry of the
185 main phase but detected a few % unreacted NiO and an unidentified low-Ni phase. The same
186 starting mix was melted for about 1 h in a Pt tube in air for about 1 h. The resulting dark reddish-
187 brown glass contained a few spherules of bright green NiO crystals; fragments without visible
188 crystals were hand-picked for NMR. EPMA data showed that this glass and its Co equivalent
189 (below) were homogeneous, and were close to the corresponding crystals in composition with 1
190 to 2 mole % lower NiO or CoO, probably through loss to the Pt containers during melting.

191 **CoCaSi₂O₆**. The starting mixture of Co₃O₄, CaCO₃ and amorphous silica was first melted
192 to a homogeneous, very dark blue glass at 1300 °C. Part of this was crystallized at 1140 °C for 8
193 h. A second sample of crystalline CoCaSi₂O₆ (also diopside structure) was made by sintering the
194 same oxide mix at 1125 °C for a total of 192 h with two intermediate grinding steps (Navrotsky
195 and Coons, 1976). EPMA detected only stoichiometric CoCaSi₂O₆ in this product. Both
196 crystalline samples were pink in color, with a small percentage of dark blue grains, presumably
197 Co-akermanite. NMR data on the two crystalline samples were similar; those shown here are for
198 the sintered sample.

199 **Co-akermanite (CoCa₂Si₂O₇)**. Crystalline CoCa₂Si₂O₇ (akermanite structure, (Kusaka et
200 al., 2001)) was synthesized from the same reagents as CoCaSi₂O₆, ground, pelletized, and heated
201 for a total of 84 h at 1190 °C with two intermediate grindings. The product was a uniform, dark
202 blue color. EPMA detected only the desired phase and confirmed its composition.

203

204 **NMR data collection and analysis**

205 NMR methods used here are similar to those recently described for paramagnetic Cu(II)
206 silicates, which required exploration of large ranges in frequency of NMR signals, far beyond
207 those for diamagnetic silicates (Stebbins, 2017), as well as methods to allow accurate
208 observations of often very broad signals. Data were acquired with a Varian 400 MHz
209 spectrometer (9.4 Tesla magnet, 79.5 MHz for ²⁹Si), with 3.2 mm diameter MAS rotors, and
210 either non-spinning (“static”) samples or spinning frequencies of 10 to 22 kHz. Tetramethyl
211 silane (TMS) at 0 ppm was the frequency reference; a natural sample of pure diopside
212 (CaMgSi₂O₆) from Wakefield, Quebec was used as an intensity standard. All spectra shown here
213 (MAS and static) were collected with a standard spin-echo pulse sequence (90° - τ - 180°) with

214 echo delay time τ (chosen as an integral number of rotor periods for the MAS spectra) usually of
215 100 μ s but ranging from 60 to 900 μ s in experiments to estimate spin-spin relaxation times (T_2)
216 via the decay in signal with τ . The spin-echo method shifts the recorded time-domain signal
217 away from residual NMR probe “ringing” and allows more accurate measurement of very broad
218 lines, whose signals decay rapidly. For many of the samples studied here, standard 1-pulse
219 methods, with typical instrumental deadtimes of 10’s of μ s, miss most or even all of the NMR
220 signal because of this decay. However, the relatively long NMR pulses used here (typically 2.5
221 and 5 μ s) resulted in a relatively narrow excitation bandwidth (empirically estimated as roughly
222 150 kHz or 1800 ppm), requiring in most cases the acquisition of several spectra with different
223 settings of the transmitter frequency (re-tuning the NMR probe at each step), with offsets chosen
224 as 100 kHz. These were summed to obtain the final spectra. (Such “spin-echo mapping” methods
225 have often been applied to observe very NMR broad spectra for paramagnetic systems, for
226 example for ^{31}P NMR of phosphate sorbed onto, and ^{27}Al of alumina dissolved into, iron
227 oxyhydroxides (Kim et al., 2015; Kim et al., 2011), as well as pure-phase Fe(III) phosphates
228 (Kim et al., 2010)). In addition, the very large paramagnetic shifts of some of the materials
229 studied here (1000’s of ppm) required preliminary experiments, shifting the transmitter over a
230 wide range, to initially locate the resonances. Tests of varying times between pulse trains
231 indicated that all spectra were fully relaxed at the 0.1 s delay generally used; in contrast the pure
232 diopside intensity standard was run with a 1 h pulse delay to allow full relaxation and
233 quantitative signal intensity. For the spectra shown for the Fe, Ni, Co silicates, 200,000 to
234 500,000 acquisitions were collected for each transmitter setting, requiring several days for each
235 full data set. Spectral widths of 2 MHz were used; the time-domain data were processed from the
236 top of the echo, set at a fixed time for all comparable experiments. Gaussian apodization was

237 chosen to enhance signal-to-noise without significantly broadening spectra (<5% of linewidth).
238 Spin echo acquisition yielded visibly flat baselines without the need for back-predicting the FID
239 or large first-order phase corrections, procedures which are often required in single-pulse data
240 sets to compensate for instrumental deadtime and which can introduce distortions in line shapes.
241 Nonetheless, minor uncertainties in baselines of the very broad spectra contributed roughly a 10
242 % uncertainty in estimated peak areas.

243 In the MAS spectra, air friction caused temperature increases of a few 10's of °C; because
244 these can strongly affect positions of paramagnetically shifted peaks they were calibrated using
245 the ²⁰⁷Pb NMR peak shift in Pb(NO₃)₂ (Takahashi et al., 1999).

246 Areas of the composite, summed spectra were determined by integrating in non-
247 overlapping 100 kHz windows for the component spectrum centered at each transmitter
248 frequency; center frequencies reported here for the very broad static spectra are taken as the
249 centers of gravity of the summed line shapes, based on the centers of their overall integrals.
250 Reported areas are normalized per mg of SiO₂ in the sample in the probe, for comparison among
251 samples and with the diopside standard.

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Results

MAS spectra for pure-phase crystals

255 The crystalline, pure-phase Ni silicates showed the smallest amount of broadening in
256 both static and MAS ²⁹Si spectra, probably at least in part because Ni²⁺ has only 2 unpaired
257 electron spins per cation, compared with 3 for Co²⁺ and 4 for Fe²⁺ (assuming the normal high-
258 spin states of low-pressure silicates). For MAS data collected with samples spinning at
259 frequencies up to 22 kHz, full separation of spinning sidebands was observed for Ni₂SiO₄ (Fig.

260 1), and nearly full separation for NiCaSi₂O₆ (Fig. 2). Even in the most-resolved spectrum,
261 however, the widths of the individual central peak and sidebands were much greater than in
262 typical, ordered, diamagnetic silicates (ca. 100 ppm full width at half maximum, FWHM, for
263 NiCaSi₂O₆, 20 ppm for Ni₂SiO₄ vs. ca. 0.2 ppm for pure Mg₂SiO₄). Much less effective
264 narrowing by MAS, and only partial resolution, was seen for Co-silicates (Fig. 2), and very little
265 narrowing for pure-phase Fe₂SiO₄. For the Ni-silicates, MAS sideband (ssb) patterns roughly
266 mapped out the observed static (non-spinning, see below) line shapes (Fig. 1), but the sideband
267 manifold was typically somewhat narrower than that of the static peak, probably because of
268 excitation bandwidth limitations: the MAS spectra were generally collected with data from only
269 a single transmitter frequency. For all of the pure-phase transition metal silicates studied here,
270 peak positions are shifted far outside of the known range for SiO₄ groups in diamagnetic silicates
271 (about -60 to -120 ppm), either to higher or to lower frequencies.

272 As is often the case with spectra containing large number of sidebands (i.e. the spinning
273 frequency is much lower than the total static linewidth), the most intense peak does not
274 necessarily correspond to the isotropic chemical shift δ_{iso} , which can be somewhat difficult to
275 identify in such cases. Typically, spectra collected with different spinning rates can be compared
276 to find the position of the peak that is independent of the rate, which can then be identified as
277 δ_{iso} . However, paramagnetic shifts are often strongly temperature dependent, and sample heating
278 (up to a few 10's of °C) caused by air friction on the sample rotor can cause spin-rate dependent
279 changes in peak positions (McCarty et al., 2015; Palke and Stebbins, 2011a; Palke and Stebbins,
280 2011b; Stebbins, 2017): in general, the magnitudes of paramagnetic shifts become smaller at
281 higher temperatures as the Boltzmann distribution of low and high energy states for the unpaired
282 electrons approaches unity (Palke and Stebbins, 2011a; Wilcke et al., 2007). This effect can be

283 clearly seen in data for Ni₂SiO₄. Figure 3 shows expanded views of the central regions of spectra
284 collected at a wide range of spinning rates, with mean rotor temperatures. We thus identify the
285 peak near to –810 ppm in the near room temperature spectrum as the likely ‘central’ peak
286 representing δ_{iso} . An isotropic average shift near to –810 ppm is also most consistent with the
287 center of gravity of the static spectrum of -820 ± 20 ppm (see below). The increase in width and
288 the development of a complex peak shape in the higher spinning speed data are probably caused
289 by the growth with spinning rate of significant temperature gradients in the sample rotor, as was
290 reported recently for Cu(II) silicates (Stebbins, 2017).

291 The minor amount of unreacted silica known to be present in this sample (see
292 experimental section) was not detected by the NMR, presumably because of a much longer spin-
293 lattice relaxation time, as expected for a nearly pure cristobalite vs. a phase rich in paramagnetic
294 cations. This same finding applied to the fayalite sample (below).

295

296 **Static (non-MAS) spectra**

297 Because of the lack of full (or in some cases, of any) line narrowing by 20 to 22 kHz
298 MAS for most of the materials studied here, static (non-MAS, stationary samples) spectra were
299 collected. For all of the samples, however, static peak widths were broad enough so that spectra
300 collected at a fixed radiofrequency transmitter setting recorded only part of the line shape, due to
301 the relatively narrow excitation bandwidth of the echo pulses used to record the spectra (roughly
302 150 kHz or about 1800 ppm for ²⁹Si). This was a minor issue for the narrowest spectra (for
303 Ni₂SiO₄ and NiCaSi₂O₆), but of major importance for the other crystalline silicates and
304 especially for the glasses. As has been well-established in other ‘wideline’ spectra of
305 paramagnetic materials (Kim et al., 2015; Kim et al., 2011; Kim et al., 2010) (and other systems

306 with extremely broad NMR lines), this problem was overcome here by the relatively time-
307 consuming approach of systematically changing the spectrometer frequency (here, in 100 kHz \approx
308 1260 ppm increments) and recording a series of spectra that were summed to give a composite
309 spectrum that accurately represented the full static line shape. A typical example is shown for
310 Co_2SiO_4 in Fig. 4.

311 Static spectra for Fe_2SiO_4 , Co_2SiO_4 , and Ni_2SiO_4 are compared in Fig. 4; data for the
312 mean isotropic shifts and widths are in Table 1. Data for the Co-olivine can be compared to a
313 single previous report of an imprecisely measured room temperature shift of “ 0.03 ± 0.02 %”
314 relative to a glass sample tube (Saji et al., 1973). Adjusting to modern units and standard, this
315 would be roughly 200 ± 200 ppm relative to TMS, but the early study was primarily focused on
316 low-temperature magnetic ordering in the olivine and did not actually show any spectra. Our new
317 data show that shifts are in fact generally large relative to those in diamagnetic silicates, as
318 expected from our recent study of Cu(II) silicates, and, in particular, from extensive recent
319 studies of paramagnetic shifts for ^{31}P in analogous LiMPO_4 (M=Mn, Fe, Co, Ni) olivine phases
320 and related oxide and phosphate battery materials (Wilcke et al., 2007). Peak widths (Table 1)
321 are also large, and may be dominated by the large magnetic susceptibility anisotropy expected in
322 these minerals with relatively low symmetries, and/or large electron-nuclear dipolar couplings.
323 The asymmetric static line shapes, especially as seen for the clinopyroxene phases (below), are
324 somewhat analogous to ‘powder patterns’ observed for spin 1/2 nuclides in some diamagnetic
325 systems, which reflect anisotropy in the chemical shifts (CSA). The centers of gravity of the
326 peaks, which represent the average over all orientations of such anisotropic effects, range widely
327 from +3200 ppm for Fe_2SiO_4 to -820 ppm for Ni_2SiO_4 .

328 Static spectra for crystalline (diopside structure) and glassy $\text{NiCaSi}_2\text{O}_6$ are shown in Fig.
329 5. The anisotropic line shape is well-developed for the crystalline sample, and is seen in both the
330 static and MAS spectra (Fig. 2); its center of gravity (+3000 ppm) is far above that for Ni_2SiO_4
331 (−820 ppm). The spectrum for the glass is much broader and shifted to much lower frequency
332 than for the crystal, probably reflecting major short-range structural differences as well as
333 heterogeneity in medium-range distributions of paramagnetic cations in the glass (see below). A
334 small fraction of the total observed signal for the glass is narrowed significantly by MAS, again
335 suggesting heterogeneity, with some SiO_4 groups without nearby Ni^{2+} neighbors. The position of
336 this component (ca. −85 ppm) is within the range for diamagnetic silicates, but its width (ca. 100
337 ppm) remains abnormally large: MAS linewidths for diamagnetic amorphous silicates are
338 typically <25 ppm. This component could be inherent to the glass, or could be due to a minor
339 crystalline impurity or even to incipient liquid-liquid phase separation, undetected by EPMA
340 analysis and X-ray mapping. Determining its origin will be of interest in future studies.

341 Spectra for crystalline and glassy $\text{CoCaSi}_2\text{O}_6$, and crystalline $\text{CoCa}_2\text{Si}_2\text{O}_7$ are shown in
342 Fig. 6. For the isostructural clinopyroxenes, the shift for the Co phase, and its line width, are
343 considerably greater than those for the Ni phase, as was also seen for the olivines. Again, the
344 linewidth for the glass is much greater than for the crystal, and the average peak position is
345 shifted much less from the “normal”, diamagnetic range. A minor component in the glass, as for
346 the Ni analog, is again somewhat narrower in both the static and MAS (not shown) spectra. The
347 akermanite phase, $\text{CoCa}_2\text{Si}_2\text{O}_7$, which has Co in tetrahedral sites instead of the octahedral sites in
348 the clinopyroxene and olivine phases, also has a much lower shift than crystalline $\text{CoCa}_2\text{Si}_2\text{O}_7$.

349

350 **Spin-echo MAS and static spectra, solid solutions**

351 As reported previously (McCarty et al., 2015), the MAS spectrum (20 kHz spinning rate)
352 of San Carlos olivine (ca. $\text{Mg}_{0.9}\text{Fe}_{0.1}\text{SiO}_4$) at first appears to be dominated by a broad (75 ppm
353 FWHM) peak centered (“A” in Fig. 7) slightly above the normal range for diamagnetic silicates
354 (-50 ± 5 ppm); this width epitomizes the loss in resolution, or even (for early NMR methods) the
355 total loss in signal resulting from major contents of paramagnetic cations (Grimmer et al., 1983),
356 as the spectrum for pure Mg_2SiO_4 has a line width of about 0.2 ppm. Slightly faster spinning
357 enhances resolution further, revealing the presence of at least one additional paramagnetically
358 shifted peak (“B” in the figure), whose isotropic position appears to be at about -200 ppm but
359 which could actually be hidden under the shoulder of the main peak near to $+70$ ppm (with the $-$
360 200 ppm peak a sideband). Both of these possible positions are shifted farther from the
361 diamagnetic resonance (-61.8 ppm) than the tiny paramagnetic peaks reported for Mg_2SiO_4
362 (forsterite) with low (0.05 to 0.4 %) Fe_2SiO_4 component (McCarty et al., 2015), which ranged
363 from about -26 to -101 ppm (relative shifts of about $+39$ ppm to -39 ppm). However, it is
364 important to also note that when spin echo MAS spectra are collected, and combined from
365 several frequency offsets, it can now be seen that roughly 60 to 80% of the observable ^{29}Si signal
366 for this mineral remains in a very broad, unresolved peak (Fig. 7, 8). The latter presumably is
367 dominated by contributions from Si sites with one or more first-neighbor Fe^{2+} cation neighbors.

368 The composite, static spectrum of this olivine allows further comparison with the pure-
369 phase samples (Fig. 8). This line shape is complex, with a narrower and a broader component or
370 possibly a range of components with varying widths. The overall peak has intensity above
371 baseline from at least $+4000$ to -2000 ppm and a width (FWHM) of about 2500 ppm. A rough
372 estimate indicates that about 65 ± 5 % of the area of the composite static spectrum (at echo delay

373 of 100 μs) is in the broader component. The overall center of gravity of the composite static
374 spectrum is about 0 ± 100 ppm.

375 Spin-echo MAS spectra for Mg_2SiO_4 with a 5% Ni_2SiO_4 component ($\text{Mg}_{1.9}\text{Ni}_{0.1}\text{SiO}_4$)
376 displayed the same set of small, paramagnetically shifted peaks (at about -10 to -70 ppm) as
377 described in our recent study using the more conventional single-pulse excitation (McCarty et
378 al., 2015). As shown in Figure 9, however, one additional, relatively broad resonance was also
379 detected, centered at about $+135$ ppm, again well outside of the range previously reported (ca.
380 50 ppm FWHM, 6 ± 2 % of observed area for echo delays of 2 rotor periods). Spectra (both MAS
381 and static) collected with the transmitter offset up to 200 kHz above or below the main peak
382 center detected no additional resonances, either narrow or broad, at the obtainable signal to noise
383 ratio. The static spectrum (Fig. 8) is relatively narrow (FWHM=108 ppm) and featureless,
384 although slightly asymmetrical: a minor shoulder on the high frequency side may corresponds to
385 the $+135$ ppm shifted resonance. The overall spectrum is about twice as wide as the static
386 spectrum of pure forsterite (Stebbins et al., 2009), indicating only a relatively minor effect of the
387 paramagnetic cation on the overall line shape.

388

389 **Signal loss, spin-spin relaxation, and peak areas**

390 In simple “one pulse” NMR experiments often used to record high resolution (MAS)
391 spectra for solids, the time-domain signal (free-induction decay or FID) is recorded immediately
392 after a short instrumental dead time necessitated in part by a finite time required for the
393 excitation signal (observe pulse) to decay from the probe circuit (“probe ringing”). Typical dead
394 times of about 10 to 40 microseconds (μs) generally have no significant effect on narrow, high
395 resolution spectra, because the FID can persist for 10’s to 100’s of milliseconds (ms) or even

396 longer. However, for very broad spectra as described here, the FID may decay in 10's of μs ,
397 often making it nearly unobservable by "one pulse" methods. In such cases, the loss of signal can
398 be detected by intensity calibration of the observed residual signal with known standards. A spin-
399 echo pulse sequence, such as employed here, shifts the FID out in time from the end of the
400 observation pulses, in principle allowing its accurate observation without dead time or "probe
401 ringing" effects. However, spin-spin relaxation during the echo delay (typically 100 μs here) can
402 lead to decay of the signal, characterized by an exponential function that can be written as
403 $\ln(M/M_0) = -\tau/T_2$, where M is the observed signal intensity, M_0 is that at zero time (here taken as
404 the expected value from a standard calibration), τ is the echo delay time, and T_2 is the spin-spin
405 relaxation time, approximated here as a constant for each data set. Measurements of signal
406 strength (peak areas in the Fourier-transformed spectra) as a function of τ thus yield estimates of
407 T_2 , as given in Table 2; representative data are plotted in Figure 10. These data can then be used
408 to correct the observed signal intensity (obtained by integration and normalized to the mass of
409 SiO_2 in the sample) back to zero time, and thus obtain a more accurate comparison among
410 samples and between samples and standards.

411 The peak areas reported in Table 2 are the sums of the integrals of non-overlapping, 100
412 kHz-wide frequency windows centered at each offset of the transmitter frequency and spanning
413 the total range of observable signals for each material. The latter was defined as data for
414 transmitter settings that yielded signals greater than about 5 % of the total detected at all
415 frequencies. The T_2 given for the pure-phase crystalline samples are based on data for the
416 transmitter setting giving the most intense (or an average of the two most intense) signal. For
417 these, measurements on the lower intensity components at different center frequencies yielded
418 values that were similar but more uncertain.

419 In some reported cases, for example ^{29}Si spectra of the paramagnetic Cu(II) mineral
420 cuprorivaite ($\text{CaCuSi}_4\text{O}_{10}$) (Stebbins, 2017), a very short T_2 (ca. 35 μs for this example) can lead
421 to loss of much of the NMR signal during a typical spin-echo experiment. However, for the pure-
422 phase Ni, Co, Fe silicates here, T_2 values, although ranging widely from about 500 to 4000 μs ,
423 are all long enough so that signal loss during a 100 μs echo delay is less than 20% and most
424 commonly <10% (Table 2). Correcting the observed, integrated intensities to zero time, it can be
425 seen that the measured values are within 5 to 10 % of those expected from data collected for a
426 pure diopside standard with the same pulse sequence and instrumental settings (apart from a
427 delay between pulse trains of 1 hour for the standard instead of 0.1 s for the paramagnetic
428 samples). This is probably well within the errors associated with observing such wide and
429 inherently noisy spectra, and suggests that this approach can indeed produce relatively accurate
430 spectra even when overall linewidths are three or four orders of magnitude greater than those
431 typically measured in diamagnetic samples. The data for these samples thus rule out any large
432 spectral components that decay extremely rapidly in these samples (i.e. that are lost entirely in
433 less than the minimum echo delays of 50 to 100 μs explored here); in turn the lack of any large
434 “invisible” component of the NMR signal makes interpretation of data for materials with
435 unknown structures (e.g. glasses) somewhat more straightforward.

436 In two of the samples (Ni_2SiO_4 and Fe_2SiO_4) known to have minor amounts of unreacted
437 silica, NMR signals for this phase were not detected due to the long spin-lattice relaxation time
438 of pure silica (i.e. its NMR signal presumably ‘saturated’ with the short pulse delays used here).
439 Given that integrated peak areas were normalized to the total sample weights, this second phase
440 would introduce a corresponding error, i.e. the tabulated values would be proportionally
441 underestimated. This appears most obviously for Fe_2SiO_4 .

442 The two olivine solid solutions (San Carlos olivine and $\text{Mg}_{1.0}\text{Ni}_{0.1}\text{SiO}_4$) studied here by
443 spin-echo and variable transmitter offset methods were both previously investigated by one-pulse
444 NMR methods along with a series of more dilute Fe, Co, Ni olivines (McCarty et al., 2015). The
445 new data showed somewhat more complex T_2 and peak area behavior than for the pure-phase
446 samples. For the Ni-doped forsterite, in spite of high signal-to-noise resulting from long
447 acquisitions at a range of transmitter settings, the only observable components of the spectrum
448 were fully narrowed by MAS, with maximum widths of about 50 ppm. The total observed area
449 (with echo delays of 90 to 100 μs), out to ± 500 ppm from the center, is about 25 to 30 % below
450 that expected from the pure diopside standard (Table 2). The multiple, well-resolved
451 paramagnetically shifted peaks for this sample present an opportunity to measure T_2 values for
452 ^{29}Si in sites with varying degrees of interaction with the unpaired electron spins on the Ni^{2+}
453 cation(s). As shown in Figure 9, large changes in the echo time τ do indeed affect the relative
454 peak intensities: the central peak is reduced only slightly by increasing τ , the closer-in shifted
455 peaks more, and the newly-observed resonance at +135 ppm is reduced nearly to undetectability
456 at the longest τ explored (909 μs). This requires systematically shorter T_2 values for resonances
457 with larger shifts (Table 2, Figure 10). For the +135 ppm peak, extrapolation back to a zero echo
458 delay leads to a significant increase in its estimated contribution to the total observable signal,
459 i.e. from about 6% at $\tau = 91 \mu\text{s}$ to $10 \pm 2\%$ at $\tau = 0$. The static spectrum for this sample showed no
460 detectable features at higher or lower frequencies than seen in the MAS spectra (Fig. 8), but its
461 integrated intensity is significantly higher, only about 15 to 20% below the standard value.

462 In contrast to the Ni-doped forsterite, the MAS and static spectra for the San Carlos
463 olivine clearly show a major, broad component that is not narrowed by MAS. This broad feature
464 also has a relatively short T_2 , about 2 times shorter than that of the overall composite spectrum

465 (Table 2). This can be clearly seen by comparing its relative contribution to spectra collected
466 with short vs. long spin echo delays (Fig. 8) Correction back to a zero echo delay indicates that
467 most of the signal for this sample is detectable (Table 2).

468 The data for the $\text{CoCaSi}_2\text{O}_6$ glass also show faster spin-spin relaxation for the most
469 shifted (highest frequency) components, with T_2 estimated as low as about 500 μs , in contrast to
470 the average for the entire spectrum of about 1800 μs (Table 2). However, given the relatively
471 low obtainable signal to noise ratio, correction to zero echo delay for all components does not
472 yield a line shape very different from that shown in Figure 6.

473

474

Discussion

475 Paramagnetic shifts and short-range crystal structure

476 Analysis of paramagnetic shifts of NMR resonances in oxide and silicate materials
477 remains a challenging problem, as the physics of the underlying interactions between unpaired
478 electron and nuclear spins is complex and quantitative theoretical predictions remain difficult.
479 The first detailed studies of any such spectra in silicates with major, stoichiometric components
480 of magnetic transition metal cations (Cu(II) silicates) have been reported only recently and
481 probably raised more questions than answered, apart from clearly showing that collecting such
482 data is feasible and potentially that they can carry structurally important information (Stebbins,
483 2017). Another very recent study of low temperature magnetic transitions in $\text{Sr}_x\text{Ba}_{1-x}\text{CuSi}_2\text{O}_6$
484 solid solutions also presented room-temperature ^{29}Si MAS spectra with paramagnetic shifts of up
485 to 500 ppm (0.05%), although the frequency reference was not reported (Puphal et al., 2016). A
486 number of studies of ^{29}Si spectra of minerals with relatively dilute paramagnetic components (up
487 to a few to 20 %) do provide some initial guidance (McCarty et al., 2015; Palke and Stebbins,

488 2011b; Palke et al., 2015; Stebbins and Kelsey, 2009), as well, although the relatively narrow,
489 highly-resolved spectra obtainable in such systems are quite different from the “wideline” data
490 reported here.

491 Fortunately, there has been extensive work on ^{31}P (and ^7Li) NMR spectra of
492 paramagnetic transition metal phosphates and oxides, in large part because of their wide and
493 important utility in lithium ion batteries (Grey and Dupré, 2004; Pecher et al., 2017) and the
494 critical role played by solid solutions and short-range order/disorder in the performance of such
495 materials. Increasingly sophisticated experimental data have also driven advanced theoretical
496 work, which has led to considerable insight into relationships between short-range crystal
497 structure and spectra that have been important in understanding results for complex solid
498 solutions (Yoon et al., 2004; Zeng et al., 2007). Of particular interest here are a number of
499 studies of the LiMPO_4 phases ($\text{M} = \text{Mn}^{2+}, \text{Fe}^{2+}, \text{Co}^{2+}, \text{Ni}^{2+}$), which have the olivine structure, with
500 octahedrally coordinated Li^+ in the M1 site, octahedrally coordinated M^{2+} in M2, and P^{5+} in the
501 tetrahedral sites (Abrahams and Easson, 1993). For example, fast MAS at a relatively low
502 magnetic field of 1.5 T yielded well-resolved ^{31}P and ^7Li spectra, the former with paramagnetic
503 shifts of up to 1000’s of ppm (all positive) (Wilcke et al., 2007). The large magnitudes of these
504 shifts, together with theoretical calculations (Kim et al., 2010; Middlemiss et al., 2013;
505 Strobridge et al., 2014) have strongly suggested that they are dominated by through-bond
506 transfer of unpaired electron spin density from the paramagnetic cation to the observed NMR
507 nucleus (transferred hyperfine coupling or Fermi contact shift). The extent of this transfer, and
508 the resulting effect on the frequency (paramagnetic shift) of the NMR resonance, depends
509 strongly on the magnetic susceptibility and thus on the number of unpaired spins and their
510 interactions if any, but also is fundamentally controlled by the covalency and orbital overlap of

511 the bonds through which the spin density is transferred. Thus, paramagnetic shifts for ${}^7\text{Li}$ in these
512 phases are typically only a few percent of the ${}^{31}\text{P}$ shifts, due in large part to the higher covalency
513 of the P–O bond relative to the Li–O bond (Wilcke et al., 2007). Other effects of unpaired
514 electron spins, notably through-space dipolar coupling or the “pseudocontact” shift, appear to be
515 of lesser importance in these materials.

516 For the LiMPO_4 olivines, a strong, systematic increase in paramagnetic shift with the
517 number of unpaired electron spins (5 for Mn^{2+} , 4 for Fe^{2+} , 3 for Co^{2+} , 2 for Ni^{2+}) was observed,
518 which was linear when corrected to temperature-independent values using systematic data on
519 variation with temperature (Wilcke et al., 2007). Room temperature values from that study are
520 shown for the Fe, Co, and Ni phosphates in Figure 11. Our new results for the silicate olivines
521 also show a systematic, roughly linear trend on the same plot, but with a much steeper slope
522 presumably related to the greater number of paramagnetic cations per tetrahedral cation in the
523 silicate structure, as well as to other structural differences and variations in coupling for the two
524 different nuclides. Our two data points for the Co,Ni clinopyroxenes ($\text{M}\text{CaSi}_2\text{O}_6$) show a trend
525 parallel to that for the silicate olivines but displaced to much higher, more positive shift values.

526 As noted above, detailed theoretical understanding of such observed contact shifts is
527 challenging, but a great deal of progress with DFT-based methods has been reported recently for
528 the phosphate olivines and related oxide materials (Clément et al., 2012; Middlemiss et al., 2013;
529 Pigliapochi et al., 2017). Here, the basic approach has involved identifying all of the first cation
530 neighbor bond paths in the structure (e.g. M–O–P), then performing a high-level calculation of
531 the electron distribution for the path, in this case focusing on the redistribution of unpaired
532 electron spin density from the transition metal 3d orbitals to the 1s orbitals of the phosphorus
533 cation, which in turn interact with the nuclear spin. The overall paramagnetic shift observed in

534 the NMR spectrum can then be approximated as the sum over all the first neighbor bond paths.
535 Because of large variation in the electronic structure (e.g. orbital overlap) for different paths, the
536 calculated individual contributions can vary widely. Values depend strongly on structural details
537 and on the approximations made in the calculation, but tend to be most different for bond paths
538 with relatively narrow M–O–P angles (ca. 90°) common for edge-shared polyhedra vs. those for
539 more open M–O–P angles (ca. 130°) for corner-shared connections, for example from about –
540 400 ppm for the former and about +700 for the latter in LiCoPO₄ (Middlemiss et al., 2013).

541 Such calculations have not yet been made for silicates, but systematic trends in the data
542 and analogy to the phosphate olivines may provide important clues as to their structural origins
543 that could be the basis for beginning to think about results for unknown structures such as
544 glasses. Apart from the large effect of numbers of unpaired electrons noted above, the most
545 striking finding is the much larger shifts for the clinopyroxenes relative to those of the silicate
546 olivines for a given M cation, a difference of almost 4000 ppm for both the Ni and Co phases.
547 This is clearly not the result of simply the number of paramagnetic cation first neighbors or to
548 the number of first neighbor M–O–Si bond pathways, which are both much greater in the
549 olivines than in the clinopyroxenes (Table 1). However, theoretical studies of the phosphate
550 olivines, and their measured paramagnetic shifts, may provide a start in accounting for such
551 dramatic structural differences. Apart from stoichiometry and number of paramagnetic M
552 neighbors per Si, the most important differences in the short-range clinopyroxene vs. olivine
553 structures (Figure 12) may be that for the former, all of the M–O–Si connections (3 per Si)
554 involve octahedral-tetrahedral corner sharing, with bond angles typically about 120 to 150° and
555 relatively long M–O–Si distances of about 0.32 to 0.35 nm (Durand et al., 1996; Ghose et al.,
556 1987; Raudsepp et al., 1990). In contrast, for the silicate olivines there are 9 first neighbor M

557 cations for each Si. Six of these form corner-shared M–O–Si linkages with bond angles typically
558 in the range of 120 to 124° and M–Si distances of about 0.32 to 0.35 nm (Smyth, 1975; Tamada
559 et al., 1983); the other three cations are in edge-shared geometries and thus form 6 M–O–Si bond
560 paths (two each) with much narrower angles (ca. 90°) and shorter M–Si distances (about 0.27 to
561 0.29 nm).

562 As a zero-order starting point for thinking about the paramagnetic shifts in the silicates,
563 one could hypothesize that in the clinopyroxene structures each of the three corner-shared bond
564 paths provides a large, positive contribution to the total shifts, as suggested by the calculations
565 for the phosphates. This could be on the order of +1000 ppm per bond path, and larger for Co²⁺
566 than for Ni²⁺ given the greater number of unpaired spins (and possibly even greater for Fe²⁺,
567 although we lack data for FeCaSi₂O₆). It could then be assumed that the observed shifts for the
568 olivines result from the sum of contributions of these values for each of the 6 corner-shared M–
569 O–Si pathways, added to 6 large and *negative* contributions from the edge-shared linkages (due
570 to the three edge-shared octahedral M cations). This could imply a progressively greater effect of
571 the negative shift contributions from Fe²⁺ to Co²⁺ to Ni²⁺.

572 Observed shifts for the LiMPO₄ olivines may provide further support for the importance
573 of this crude but large distinction between corner-shared and edge-shared bond paths, given that
574 in these structures, the M1 site is occupied by Li⁺, which of course has no unpaired electron
575 spins. The result is that there are 4 instead of 6 corner-shared M–O–P linkages (4 cations) and
576 only 2 edge-shared linkages (1 cation). This could substantially reduce the negative contributions
577 to the total shift, leading to the higher overall observed values (all positive) for the phosphates
578 relative to the silicates.

579 However, at this point of our knowledge, we cannot be at all certain that this hypothesis
580 of large positive shifts for corner-shared linkages, partly balanced by large negative shifts for
581 edge-shared linkages, is responsible for the large observed differences between the data for
582 silicate olivines and clinopyroxenes. Other aspects of their electronic and/or magnetic structures
583 could play an important role. The olivine data, taken in isolation, could also suggest that each of
584 the nine first neighbor M cations contributes an *average* paramagnetic shift estimated as the
585 observed total simply divided by 9, yielding values of +356, +82, and -91 ppm/cation for
586 Fe₂SiO₄, Co₂SiO₄, and Ni₂SiO₄, respectively.

587 The coordination number of the paramagnetic cation itself is of course likely to also be
588 important for its contribution to NMR shifts, but this issue has rarely been explored either
589 empirically or theoretically. One result that bears on this question is the measured shift for
590 akermanite-structured CoCa₂Si₂O₇ (Kusaka et al., 2001), in which Co²⁺ is four-coordinated
591 instead of six-coordinated as in the olivines and clinopyroxenes. In the akermanite structure,
592 each Si has only two Co neighbors, with corner-shared Co–O–Si linkages. The observed shift of
593 +1380 is again relatively large and positive, but is lower than might be expected from the data
594 for octahedral M cation structures. This lower value comes in spite of the shorter, presumably
595 more covalent, Co–O bonds in the tetrahedra relative to octahedra, as well as shorter Co–O–Si
596 distances (ca. 0.31 nm) at similar angles (ca. 118°). A detailed analysis of the electronic structure
597 in this phase, in particular how crystal field effects and differences in d-orbital energy levels
598 change the charge distribution around transition metal cations in octahedra vs. tetrahedral sites,
599 and how this in turn affects the Si–O bonds, would be useful for further analysis of these data,
600 which are especially interesting for results on glasses discussed below. Another comparison with
601 our results here for transition metal cations in octahedral sites in olivine and clinopyroxene are

602 several studies of the effects of Fe^{2+} in dodecahedral sites in garnets on ^{29}Si paramagnetic shifts
603 in adjacent edge-shared sites (about +200 ppm) and corner-shared sites (about -20 ppm) (Palke
604 and Stebbins, 2011b; Palke et al., 2015).

605

606 **Crystals vs. glasses**

607 The reasonably good glass-forming ability of the $\text{NiCaSi}_2\text{O}_6$ and $\text{CoCaSi}_2\text{O}_6$
608 compositions provides a nearly unique opportunity to compare ^{29}Si spectra for transition metal-
609 rich glasses and nearly-isochemical crystals. As noted above and as illustrated by Figures 5 and
610 6, in both cases the glass spectra are much (at least 2 or 3 times) wider than those for the
611 corresponding crystals, and their average paramagnetic shifts are much (2000 to 3000 ppm)
612 lower. Also as noted above, the summed, integrated peak areas for the glasses and crystals are
613 equal within about 5%, showing that the large differences in peak shapes and positions cannot be
614 due to the selective loss of a major part of the NMR signal.

615 In a very general sense, the glasses are expected to have considerable structural disorder
616 and thus a possibly wide range of SiO_4 sites with varying numbers and varying types of bond
617 pathways to first-neighbor paramagnetic cations, as well as even more disorder with respect to
618 medium- and long-range structure. Some aspects of this disorder are relatively well known from
619 spectroscopic studies of diamagnetic glass compositions, for example $\text{MgCaSi}_2\text{O}_6$ (diopside). In
620 such compositions, instead of every SiO_4 group having exactly two bridging and two non-
621 bridging oxygens as in crystalline pyroxenes (“ Q^2 ” groups), a wide distribution of groups with 1,
622 2, 3, and probably 0 and 4, bridging oxygens (“ Q^n ”) (Stebbins, 1995) can be quantified by 2-
623 dimensional ^{29}Si NMR (Davis et al., 2011); one result is that the observed 1-dimensional MAS
624 NMR spectra for such glasses are typically at least 10 to 20 times broader (ca. 20 ppm) than

625 those of the corresponding crystals. Variations in bond distances and angles are also expected to
626 be large. Less is known about distributions of coordination environments for the lower-valent
627 network modifier cations, but ^{25}Mg NMR and X-ray methods have suggested that Mg^{2+} in
628 pyroxene-composition glasses probably has a mix of coordination numbers with a mean less than
629 the value of six seen in crystalline pyroxenes and olivines (George and Stebbins, 1998; Kroeker
630 and Stebbins, 2000; Trcera et al., 2009). Co^{2+} and especially Ni^{2+} have been extensively studied
631 by optical and X-ray spectroscopy and other methods in glasses including $\text{NiCaSi}_2\text{O}_6$ (Galoisy
632 and Calas, 1993; Hunault et al., 2014; Johnson et al., 1999). The short-range environments for
633 both can vary depending on glass composition, but again there are likely to be mixtures of
634 coordination numbers with means less than six: for example, the strong blue color commonly
635 caused by even $<1\%$ Co^{2+} added to silicate glasses is due to an unusually strong absorption
636 spectrum for the cation in tetrahedral coordination. M-Si distances derived from X-ray
637 spectroscopy are generally consistent with corner-shared connections. The +2 valence is
638 generally assumed to be predominant for Ni and Co in silicate glasses melted in air as done here,
639 and is consistent with optical and X-ray spectroscopic studies (Galoisy and Calas, 1993; Hunault
640 et al., 2014). However, minor concentrations of other valences, with different unpaired electron
641 spin populations, could further complicate the analysis of NMR spectra.

642 The data presented here on paramagnetic shifts in the transition metal silicate crystals of
643 known structure is (so far) too limited to uniquely address the observed large differences
644 between spectra for $\text{M}\text{CaSi}_2\text{O}_6$ glasses vs. crystals. We can say that Si environments very similar
645 to those in the clinopyroxene crystals (exclusively corner-shared octahedra, with large positive
646 shifts) are probably uncommon in the glasses. We have hypothesized that edge-sharing between
647 M^{2+} octahedra and SiO_4 tetrahedra could lead to large negative shifts, resulting in the much

648 smaller total shifts for the pure-phase olivines relative to the clinopyroxenes. However, it is
649 difficult to see why a glass with a much lower ratio of M^{2+} to Si, and presumably lower bulk
650 density, than the corresponding olivine, would have a preponderance of closely-packed edge-
651 shared geometries.

652 Our data for crystalline $\text{CoCa}_2\text{Si}_2\text{O}_7$ (akermanite structure), in which all Co^{2+} is four-
653 coordinated provides another clue: the observed paramagnetic shift is much lower than in
654 diopside-structured $\text{CoCaSi}_2\text{O}_6$, presumably because of as-yet poorly understood differences in
655 electronic structure and bonding. Given that low-coordinated Co^{2+} (and Ni^{2+}) are likely to be
656 abundant in such glasses (Galoisy and Calas, 1993; Hunault et al., 2014), their much lower
657 paramagnetic shifts may be at least in part related to this structural variable.

658

659 **Paramagnetic shifts in pure phases vs. dilute solid solutions**

660 Our previous studies of LaPO_4 (monazite) and YPO_4 (xenotime) doped with a few
661 percent of rare earth ions containing unpaired electron spins (REE, e.g. Nd^{3+} , Eu^{3+} , Ce^{3+})
662 described paramagnetically shifted resonances corresponding to single first neighbor dopant
663 cations in each of several distinct REE-O-P linkages (Palke and Stebbins, 2011a; Palke et al.,
664 2013). Single-neighbor shifts of up to about 250 ppm were reported. These can be summed over
665 each of the REE-O-P bond paths to predict what would be expected in the pure magnetic REE
666 phase. For example, the sum over the 7 first neighbor shifts (measured at 47 °C, relative to the
667 unshifted resonances at 0 ppm) in $\text{La}_{0.9}\text{Ce}_{0.1}\text{PO}_4$ is -59 ppm, and for $\text{La}_{0.9}\text{Nd}_{0.1}\text{PO}_4$ is -128 ppm
668 (Table 1, (Palke and Stebbins, 2011a)). These totals are roughly similar to measured values of -
669 87 for the former and -148 for the latter (correcting a typographic error in the original table). In
670 continued studies of REE phosphates, sums of shifts observed in dilute $\text{La}_{(1-x)}\text{Ce}_x\text{PO}_4$ solid

671 solutions predicted spectra for phases with x as high as 0.33, when additional peak broadening
672 was included (Palke et al., 2013). In these examples, additivity of shifts across a wide
673 concentration range is thus a reasonable approximation and, apparently, paramagnetic
674 broadening is small enough to allow first-neighbor shifts to be readily observed in the dilute
675 solid solutions. In more detail, differences in the summed individual shifts and the total for the
676 pure phases are probably to be expected from variations in local structure with composition.

677 For our new results on pure-phase Fe, Co, Ni silicates, it is apparent that at least some
678 first-neighbor paramagnetic shifts for corner-shared M-O-Si linkages are probably on the order
679 of at least 1000 ppm, for example to account for the observed total shifts of 3000 ppm and 4400
680 ppm in $\text{NiCaSi}_2\text{O}_6$ and $\text{CoCaSi}_2\text{O}_6$ respectively. For the silicate olivines, we have hypothesized
681 that these are partially balanced by large negative shifts (also on the order of -1000 ppm) for
682 edge-shared linkages to reach the much lower total observed shifts for Ni_2SiO_4 and Co_2SiO_4 ($-$
683 820 and $+740$ ppm respectively). Without theoretical calculations, these numbers remain as only
684 initial estimates. However, the data, as well as extensive theoretical work on analogous LiMPO_4
685 olivines (Middlemiss et al., 2013; Pigliapochi et al., 2017), do suggest that these effects should
686 be quite large.

687 In contrast, in our recent detailed study of resolvable paramagnetically shifted ^{29}Si NMR
688 resonances in Mg_2SiO_4 with up to 5% Fe, Co, or Ni substitution (McCarty et al., 2015), we
689 detected (with conventional “one pulse” NMR methods) only much smaller shifts, up to about 50
690 to 60 ppm from the central diamagnetic peak. There, we hypothesized that some of these could
691 be attributed to first cation neighbor interactions, smaller than for transition metal phosphate
692 olivines perhaps because of more ionic bonding in the silicates. The number and relative
693 intensities of such shifted resonances also indicated that at least some were due to magnetic

694 cations in more distant shells. With the spin-echo method, somewhat faster sample spinning, and
695 variable transmitter offsets used in our new study, we did locate an additional relatively broad
696 resonance at about +135 ppm in $\text{Mg}_{1.9}\text{Ni}_{0.1}\text{SiO}_4$, but this shift (200 ppm from the diamagnetic
697 peak) is still much less than those suggested for single M-O-Si linkages in the pure phase Fe, Co,
698 and Ni olivines. Another possible discrepancy of this type may be paramagnetic shifts for ^{31}P in
699 a few transition metal lithium phosphate solid solutions, which were reported to be much smaller
700 than in the pure end member phases, for unknown reasons (Wilcke et al., 2007).

701 For the silicate olivines, this apparent discrepancy between paramagnetic shifts in the
702 dilute solid solutions and the pure paramagnetic phases could be attributed to several causes or
703 their combination. First, there could be differences in local bonding/electronic structures around
704 the transition metal cations, which, given the great sensitivity of paramagnetic shifts to structure,
705 could lead to large differences in the transfer of unpaired electron spin density to the NMR
706 nuclides in similar sites in dilute vs. pure-phase systems. Given our own results for the REE
707 phosphates as well as other examples from the battery material literature, we consider this to be
708 an unlikely explanation for overall large differences. However, it could be responsible for some
709 differences in shift magnitudes and the number and areas of shifted resonances, if occupancies of
710 sites with varying electronic structures change: for example, in the dilute olivine solutions, Ni^{2+}
711 and to some degree Co^{2+} are ordered into M1 sites in preference to M2, but occupy both M1 and
712 M2 sites in Ni_2SiO_4 and Co_2SiO_4 (Ghose and Wan, 1974; McCarty et al., 2015; Rajamani et al.,
713 1975). Secondly, it is conceivable that in the pure-phase paramagnetic olivines, there is some
714 kind of short-range magnetic ordering that greatly enhances NMR shifts relative to those in the
715 dilute systems. The transition metal silicate olivines do undergo transitions to long-range
716 magnetic ordering at low temperatures (typically <70 K (Ballet et al., 1989; Newnham et al.,

717 1965)), but we are aware of no evidence for continued short-range effects near room
718 temperature. In a variable temperature ^{31}P NMR study of LiMPO_4 phases, large paramagnetic
719 shifts (1000's of ppm) persisted up to at least 300 °C, suggesting that magnetic ordering did not
720 be playing an important role in these materials (Wilcke et al., 2007). More generally, as
721 concentrations of cations with unpaired electron spins increase from dilute solid solutions to
722 pure-phase endmember, the probability of short-range magnetic interactions among them
723 increases dramatically, conceivably leading to important effects on couplings to NMR-observed
724 nuclei. However, any specific prediction of such effects remains a challenging theoretical
725 problem at this point. Further studies of magnetic properties, EPR, etc., could of course
726 contribute to resolving this hypothesis for the silicates. Thirdly, it is possible that the especially
727 large shifts that we have observed for the Co and Ni clinopyroxenes, which are the best evidence
728 for very large first-neighbor effects of >1000 ppm for each first neighbor M cation, are in some
729 way anomalous, with per-cation shifts in the olivines of smaller magnitude, closer to the average
730 values estimated above of +350 to -90 ppm/cation. Here, further studies of other end member
731 phases (e.g. $\text{FeCaSi}_2\text{O}_6$) and of intermediate solid solutions could prove informative.

732 Based on this limited new data set, it now seems possible that all (instead of some) of the
733 previously reported paramagnetic shifts in the dilute Fe, Co, Ni solid solutions in Mg_2SiO_4
734 (McCarty et al., 2015) are the result of cations beyond the first Si-O-M shell, with unpaired
735 electron spin density transferred through at least four bonds (e.g. Si-O-Mg-O-Ni) and/or through
736 space via the “pseudo-contact” mechanism. Our new data for $\text{Mg}_{1.9}\text{Ni}_{0.1}\text{SiO}_4$ (5% Ni forsterite)
737 confirms that, apart from the newly observed +135 ppm peak, hypothesized first-neighbor
738 paramagnetic resonances with shifts of many 100's or 1000's of ppm are not readily observable.
739 Our measurement of the total observable peak area for this sample, of about 10 to 20% below the

740 value expected from the standard, suggests also that there are important components of the
741 spectrum that are not detected by the methods used. However, a simple probability calculation
742 (McCarty et al., 2015; McCarty and Stebbins, 2016b) for this composition, assuming all Ni²⁺
743 cations are on the M1 site, predicts that 66% of the Si have no Ni²⁺ in the first shell. If all signals
744 from Si with one or more first shell Ni²⁺ are missed in the NMR spectra, an overall signal loss of
745 34% is thus expected. The observed loss is considerably less than this, although uncertainties are
746 fairly large. It could be that the newly-detected resonance at +135 ppm is indeed due to one
747 corner-shared first neighbor coupling that is relatively small compared to our estimates from the
748 end member transition metal silicates, or that some combination of paramagnetic shifts from
749 first-neighbor cations results in intensity within the observed manifold of peaks. Again, the
750 possibility that paramagnetic shifts deduced from the pure-phase samples do not accurately
751 predict the observations in dilute solid solutions could be clarified by further study of
752 intermediate compositions.

753 This apparent loss of some observable signal for Si with first-neighbor paramagnetic
754 cations in the Ni-doped forsterite seems surprising, given the readily observable spectra for the
755 pure-phase materials, with peak areas approximately as expected from standardization, and spin-
756 spin relaxation times T_2 long enough to avoid major signal losses during the spin echo delay
757 time. The relatively low intensity of such resonances in the dilute systems, and possible
758 broadening due to disorder, of course contributes to this problem, but our observations suggests
759 that additionally, signal decay and/or peak broadening caused by rapid relaxation may be
760 unusually severe in the magnetically more dilute silicate minerals. For the 5% Ni forsterite, the
761 systematically shorter T_2 values for the observed resonances that result from stronger Ni-Si
762 interactions (larger shifts) imply the possibility of even shorter values for hypothesized first-

763 neighbor interactions with even larger paramagnetic shifts. This could lead to signal loss during
764 the NMR acquisition, and/or more severe peak broadening, making direct detection of such
765 signals difficult, especially when their intensities are small in dilute solutions.

766 We do not yet have a definitive explanation for the unusually short T_2 values for the
767 highly-shifted resonances in the more dilute Ni silicates, but note that for the pure phases studied
768 here, T_2 is shortest in one of the minerals that has the *lowest* transition metal contents
769 (NiCaSi₂O₆). In our recent report on ²⁹Si NMR of Cu(II) silicates (Stebbins, 2017), we noted that
770 the only one of the minerals studied for which T_2 was short enough to lead to major signal loss
771 was that with the lowest Cu content, cuprorivaite (CaCuSi₄O₁₀). We speculated that this could be
772 the result of reduced magnetic Cu-Cu interactions, as this phase has isolated Cu²⁺ ions instead of
773 chains of Cu²⁺ in minerals such as diopside (Ca₂Mg₅Si₈O₂₂·2H₂O). This effect might in turn lead to
774 longer electron spin-lattice relaxation times, possibly enhancing the efficiency of coupling of
775 unpaired electron spins to nuclear spins and actually shortening the nuclear spin relaxation time.
776 We have no direct evidence to support this hypothesis for other transition metal silicates, but it
777 could be an important part of the difference in behavior between dilute and concentrated
778 systems.

779 The MAS data for the San Carlos olivine (Mg_{1.8}Fe_{0.2}SiO₄) provide another window onto
780 this complex problem. Relative to the Mg_{1.9}Ni_{0.1}SiO₄ olivine, this mineral has twice the
781 concentration of paramagnetic cations and these are probably distributed over both M1 and M2
782 sites instead of occupying primarily M1. Importantly, Fe²⁺ has 4 unpaired electron spins, Ni²⁺
783 only 2. Both concentration and spin numbers probably contribute to the much greater broadening
784 of the MAS-resolved resonances (ca. 2 ppm FWHM for the Ni solid solution, 75 ppm for the Fe
785 sample); conversely, however, the very broad, unresolved component for strongly-interacting

786 sites is detectable in the latter but not the former. A simple probability calculation for the
787 $\text{Mg}_{1.8}\text{Fe}_{0.2}\text{SiO}_4$ composition predicts 39% of the Si with no Fe^{2+} in first-neighbor positions,
788 which corresponds roughly with the estimate of $35\pm 5\%$ of the total area in the narrower
789 component of the composite static (or MAS) spectra. However, even the very broad observed
790 component seems to span a smaller range of shifts than suggested by our crude estimates for
791 individual bond path contributions described above, which are least well constrained for Fe^{2+}
792 relative to Ni^{2+} or Co^{2+} , given data for only a single pure-phase Fe silicate at this time. A
793 relatively rapid spin-spin relaxation (short T_2) is also seen for the broad component of the San
794 Carlos olivine spectra, although in this case this effect apparently does not lead to a large loss in
795 total observable signal.

796

797 **Implications**

798 Since the first widespread application of ^{29}Si NMR to silicate minerals and glasses in the
799 early- to mid-1980's, the method has become one of our most important experimental tools for
800 quantifying questions of short- to medium-range structure in materials with significant variations
801 from perfect ordering (Kirkpatrick 1988; MacKenzie and Smith, 2002; Stebbins and Xue, 2014).
802 Because order/disorder is a key issue in the crystalline solid solutions that are nearly ubiquitous
803 in natural and in technological systems, and is even more fundamental to the nature and
804 properties of glasses and liquids, NMR has had a major impact in our understanding of the
805 molecular-scale details that control many different processes. However, almost all of such
806 studies have been limited to materials with relatively low concentrations of ions with unpaired
807 electron spins, as early work suggested that minerals with more than a few percent of such
808 magnetic transition metal or rare earth cations would yield uninterpretable or even unobservable

809 spectra (Grimmer et al., 1983). Large ranges of composition of geologically and technologically
810 important materials have thus been excluded from study by NMR on this important nuclide, as
811 well as others of particular mineralogical interest such as ^{27}Al and ^{17}O . The most important
812 implication of our new study, along with our initial work on Cu(II) silicates (Stebbins, 2017), is
813 thus to demonstrate that ^{29}Si NMR experiments on silicate minerals and glasses with high
814 concentrations of paramagnetic transition metal cations can indeed provide informative spectra,
815 despite considerable challenges in data collection and interpretation. Although this type of
816 application of NMR is in its infancy, the potential expansion in the scope of the method offers a
817 wide range of future exploration and potential discovery of new insights into structure in
818 crystalline and amorphous solids.

819 In more technical detail, the results presented here have important implications for the
820 methods needed to obtain useful data, for the types of systems where future work may be most
821 feasible, and for what needs to be developed to make progress in a challenging new field. For the
822 olivines, clinopyroxenes, and glasses studied here, we have found that spin-spin relaxation times
823 (T_2) are generally long enough to allow quantitative acquisition of spectra, although peak
824 broadening can be severe and spectra are shifted far outside of the range generally observed for
825 diamagnetic silicates. Large effects of variation in short range structure and composition on
826 observable paramagnetic shifts, line widths, and relaxation processes suggest a sensitive new
827 approach for understanding more about bonding, electronic, and magnetic structure. Although
828 unique interpretations of these data are also difficult at this early stage of our knowledge, future
829 systematic experimental studies of individual transition metal substitutions over a wider range of
830 known structures and solid solutions are likely to be revealing of systematic trends in effects of

831 variables such as number of unpaired electron spins, cation coordination numbers, edge- vs.
832 corner-shared bonding, and structural and magnetic order/disorder.

833 Spectra for transition metal rich silicate glasses are extremely broad and time-consuming
834 to observe, but careful analysis of integrated peak intensities indicates that they can be accurately
835 measured. At least for the $\text{NiCaSi}_2\text{O}_6$ and $\text{CoCaSi}_2\text{O}_6$ compositions, the peak widths and
836 positions are very different from those of the corresponding crystals, requiring major differences
837 in structural order and average short range structure, probably due in part to differences in
838 transition metal cation coordination numbers. Again, future, more systematic studies are likely to
839 provide new insights into the many remaining unknowns of silicate glass and melt structure.

840 Our new results for pure-phase Fe, Co and Ni silicate minerals suggest that paramagnetic
841 shifts in ^{29}Si NMR spectra caused by single transition metal cation neighbors may be much
842 larger (100's to 1000's of ppm) than any of the resolved, shifted resonances detected in our
843 recent work on forsterite (Mg_2SiO_4) containing up to a few percent of these substituents
844 (McCarty et al., 2015). This implies that signals for such first-neighbor pairs may be especially
845 difficult to detect in dilute solid solution because of rapid relaxation and/or severe broadening of
846 such low-intensity features of the spectra. Independent evidence for this possibility comes from
847 new observations of especially rapid spin-spin relaxation (and thus signal decay during the NMR
848 experiment) for resonances with large shifts and/or broadening in a forsterite with a 5% Ni_2SiO_4
849 component and in natural San Carlos olivine (10% Fe_2SiO_4 component). This conclusion implies
850 that some re-interpretation of our previous data on the dilute solid solutions may provide
851 additional information on cation ordering in olivine. Furthermore, the large numbers of well-
852 resolved, observable resonances in the spectra of dilute silicate solid solutions (McCarty et al.,
853 2015) may present unique opportunities to better understand extended cation ordering, if

854 paramagnetic shifts are actually dominated by second neighbor (or even longer-range)
855 interactions.

856 Beyond the further development of empirical correlations between short-range structure
857 and NMR observables, future NMR studies of ^{29}Si (and other nuclides) in paramagnetic-rich
858 silicates are likely to be considerably enhanced by more advanced methodologies, including
859 more complex pulse sequences that more effectively excite very broad resonances (Pell and
860 Pintacuda, 2015) and MAS probes that can spin samples to much higher frequencies. Probably
861 most importantly, the coupling of experimental work with advanced computational approaches,
862 as has been so-well demonstrated for transition-metal rich Li phosphate battery materials (Pecher
863 et al., 2017; Pigliapochi et al., 2017) has the potential to provide unique and detailed connections
864 between spectra and short-range structure in a host of complex and incompletely known
865 materials.

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867

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- 1061
- 1062

1063 Table 1. Pure-phase crystalline silicates studied here: summary of first neighbor cation bond
 1064 paths and NMR results from static spectra for these and for glasses. Here, as throughout paper,
 1065 “M” denotes paramagnetic +2 transition metal cation.
 1066

composition (structure type)	no. of M first neighbors to Si	no. of unpaired electrons per M	no. of M- O-Si paths, corner- shared	no. of M- O-Si paths, edge- shared	²⁹ Si NMR shift, ppm (mean isotropic shift)	width, FWHM in ppm ^a
NiCaSi ₂ O ₆ (diopside)	3	2	3	0	3000	580
NiCaSi ₂ O ₆ (glass)					700	2000
CoCaSi ₂ O ₆ (diopside)	3	3	3	0	4400	1700
CoCaSi ₂ O ₆ (glass)					1200	3700
Ni ₂ SiO ₄ (olivine)	9	2	6	6 ^b	-820	1250
Co ₂ SiO ₄ (olivine)	9	3	6	6 ^b	740	2100
Fe ₂ SiO ₄ (olivine)	9	4	6	6 ^b	3200	2500
CoCa ₂ Si ₂ O ₇ ^c (akermanite)	2	3	2	0	1380	1400

1067 Notes:

1068 ^a Full width at half maximum (FWHM) is only a rough description of the width for asymmetric
 1069 peak shapes such as those observed here: see figures.

1070 ^b Note that each M cation in an edged-shared site contributes two M-O-Si bond paths to each
 1071 adjacent Si.

1072 ^c Four, instead of six-coordinated, transition metal cation as in olivine and pyroxene crystals

1073
 1074
 1075
 1076

1077 Table 2. Spin-spin relaxation time data and corrected ²⁹Si NMR peak areas
 1078

composition	phase or peak	T_2 , μ s	% signal loss in standard echo time ^b	corrected total peak area, per mg of SiO ₂ ^c
Ni ₂ SiO ₄	crystal	1700 ^a	6	20.3
Co ₂ SiO ₄	crystal	1400 ^a	7	20.2
Fe ₂ SiO ₄	crystal	1100 ^a	9	18.5 ^d
CoCa ₂ Si ₂ O ₇	crystal	750 ^a	13	20.8
NiCaSi ₂ O ₆	crystal	500 ^a	18	18.9
	glass	1650 ^e	6	18.2
CoCaSi ₂ O ₆	crystal	1600 ^a	6	18.7
	glass	1800 ^f	6	18.2
Mg _{1.8} Fe _{0.2} SiO ₄	olivine, ^g MAS	2100 ^a	5	17.8
	static	1600 ^f	6	18.5
	static, offset	650 ^h	15	-
Mg ₁ Ni _{0.1} SiO ₄	MAS, center pk	5600	2	15.0 (all) ⁱ
	MAS, -40 ppm pk	1500	6	
	MAS, +135 ppm pk	200	56	
	static	900	11	
MgCaSi ₂ O ₆	crystal ^j	-	≈ 0	19.0

1079
 1080 Notes:
 1081 ^a data are for the transmitter frequency centered near the most intense component of the line
 1082 shape
 1083 ^b 2 MAS rotor periods: 100 μ s for all but Mg₁Ni_{0.1}SiO₄ where 91 μ s was used to match 22 kHz
 1084 spin rate
 1085 ^c uncertainties in areas about 10%
 1086 ^d area is minimum value, not corrected for roughly 5-10% excess crystalline silica (see text)
 1087 ^e data for two largest components of spectrum
 1088 ^f data for entire composite spectrum
 1089 ^g from peridotite nodule, San Carlos, NM
 1090 ^h data for component centered 200 kHz (2500 ppm) above peak maximum
 1091 ⁱ corrected with mean value for all observed peaks, 6% total signal loss
 1092 ^j high-purity diopside standard, Wakefield, Quebec
 1093
 1094

1095

1096 **Figure captions**

1097

1098 **Figure 1.** ^{29}Si spin echo spectra for Ni_2SiO_4 , comparing results for static (non-spinning) sample
1099 vs. MAS at 20 kHz spinning rate. The static spectrum is the sum of two spectra collected with
1100 transmitter frequency shifted by 100 kHz (1260 ppm); the MAS data were collected at a single
1101 transmitter setting near to the highest peak.

1102

1103 **Figure 2.** ^{29}Si spin echo spectra for $\text{NiCaSi}_2\text{O}_6$ and $\text{CoCaSi}_2\text{O}_6$, collected with 20 kHz MAS and
1104 a single transmitter setting for each, centered near the peak maximum. Especially for the latter,
1105 the full peak width is not accurately recorded because of the limited excitation bandwidth.

1106

1107 **Figure 3.** Central regions of ^{29}Si spin echo MAS spectra for Ni_2SiO_4 , comparing results for
1108 sample spinning at different rates as labeled. The most likely location of the central peak,
1109 marking the isotropic chemical shift (δ_{iso}), is marked by the dashed line: its position shifts
1110 towards the diamagnetic value (ca. -61 ppm) with increasing temperature, the latter due to
1111 rotational air friction. All other peaks are spinning sidebands, separated by the spinning
1112 frequency. An example of a full spectrum is shown in Fig. 1.

1113

1114 **Figure 4.** ^{29}Si spin echo static (non-spinning) spectra for Fe_2SiO_4 (fayalite), Co_2SiO_4 , and
1115 Ni_2SiO_4 , showing summed spectra only for clarity. Centers of gravity, marking the isotropic
1116 shifts, are marked and labeled. Spectra collected at four transmitter frequencies successively
1117 offset by 100 kHz and summed are illustrated for Co_2SiO_4 . The known range of isotropic and

1118 anisotropic chemical shifts for SiO_4 groups in diamagnetic silicates is shown by the box with
1119 dashed lines.

1120

1121 **Figure 5.** ^{29}Si spin echo spectra for crystalline and glassy $\text{NiCaSi}_2\text{O}_6$. For the glass, both MAS
1122 and static spectra are shown, revealing a minor component with much lower paramagnetic shift
1123 and broadening than the bulk of the sample. Centers of gravity, marking the mean isotropic
1124 shifts, are marked and labeled. The arrow labeled “Ol” shows the center of the peak for Ni_2SiO_4
1125 olivine.

1126

1127 **Figure 6.** ^{29}Si spin echo static spectra for crystalline and glassy $\text{CoCaSi}_2\text{O}_6$, and crystalline
1128 $\text{CoCa}_2\text{Si}_2\text{O}_7$. Centers of gravity, marking the mean isotropic shifts, are marked and labeled. The
1129 arrow labeled “Ol” shows the center of the peak for Co_2SiO_4 olivine.

1130

1131 **Figure 7.** ^{29}Si spin echo spectra for San Carlos olivine (ca. $\text{Mg}_{1.8}\text{Fe}_{0.2}\text{SiO}_4$) collected with MAS
1132 at 20 and 22 kHz. “A” marks the most prominent, narrowed peak whose position is shifted only
1133 slightly from that of forsterite. (The position and width of the latter are marked by the vertical
1134 line labeled “Fo100”). “B” and “B?” mark possible positions of a second, partially resolved
1135 peak that has a significant paramagnetic shift: these are separated by close to the 22 kHz
1136 spinning rate and one is probably a sideband of the other. “*” marks obvious spinning sidebands.
1137 The central regions are shown in the larger spectra; the full, composite spectrum (22 kHz) is
1138 shown in the inset on a different scale.

1139

1140 **Figure 8.** ^{29}Si spin echo spectra for olivine solid solutions. Upper three are for San Carlos olivine
1141 (ca. $\text{Mg}_{1.8}\text{Fe}_{0.2}\text{SiO}_4$), comparing the static spectrum with the MAS spectrum (as in Fig. 7 but with
1142 an enlarged vertical scale factor). The upper curve (black) shows the data collected under
1143 standard conditions with spin echo delay τ of 100 μs , the lower overlaid curve (blue) shows data
1144 for τ of 700 μs . The lower pair shows the static spectrum for 5%Ni forsterite ($\text{Mg}_{1.9}\text{Ni}_{0.1}\text{SiO}_4$)
1145 with two different vertical scale factors to illustrate the absence of a detectable broad component.
1146

1147 **Figure 9.** ^{29}Si spin echo spectra for 5%Ni forsterite ($\text{Mg}_{1.9}\text{Ni}_{0.1}\text{SiO}_4$) collected with MAS at 22
1148 kHz. Inset (on a different frequency scale and offset to the right) shows the newly-observed peak
1149 at about +135 ppm, with vertical scale about 20x that of central region shown below. * marks a
1150 spinning sideband. In each plot, three spectra are overlaid, collected with echo delay τ of 91 μs
1151 (highest intensities) to 273 to 909 μs (lowest).
1152

1153 **Figure 10.** Plot of the natural logarithm of the observed NMR peak area ratioed to the area at the
1154 shortest echo delay tested (60 or 100 μs), vs. echo delay, for Ni-containing crystals, illustrating
1155 variation in T_2 (the inverse of the slope) as listed in Table 2.
1156

1157 **Figure 11.** Plot of isotropic paramagnetic shifts observed near to room temperature for transition
1158 metal silicates and lithium phosphates, vs. number of unpaired electron spins for each M^{2+}
1159 cation. Solid circles are published data for ^{31}P shifts in LiMPO_4 ($\text{M} = \text{Fe}^{2+}, \text{Co}^{2+}, \text{Ni}^{2+}$) olivines
1160 (Wilcke et al., 2007). Squares mark ^{29}Si shifts for M_2SiO_4 olivines; triangles show data for
1161 $\text{M}\text{CaSi}_2\text{O}_6$ clinopyroxenes.
1162

1163 **Figure 12.** Local structures around the SiO_4 tetrahedra in M_2SiO_4 olivine and $\text{M}\text{CaSi}_2\text{O}_6$
1164 clinopyroxene. Apart from the central Si and its 4 first neighbor O (red), only cation sites that
1165 may be occupied by Fe^{2+} , Co^{2+} or Ni^{2+} are shown: 4 M1 (green) and 5 M2 (brown) sites for
1166 olivine, 3 M1 (green) sites for clinopyroxene. M-O-Si bond paths for corner-shared octahedra are
1167 shown in gray, those for edge-shared octahedra are highlighted by dashed lines (color online).
1168
1169

Figure 1

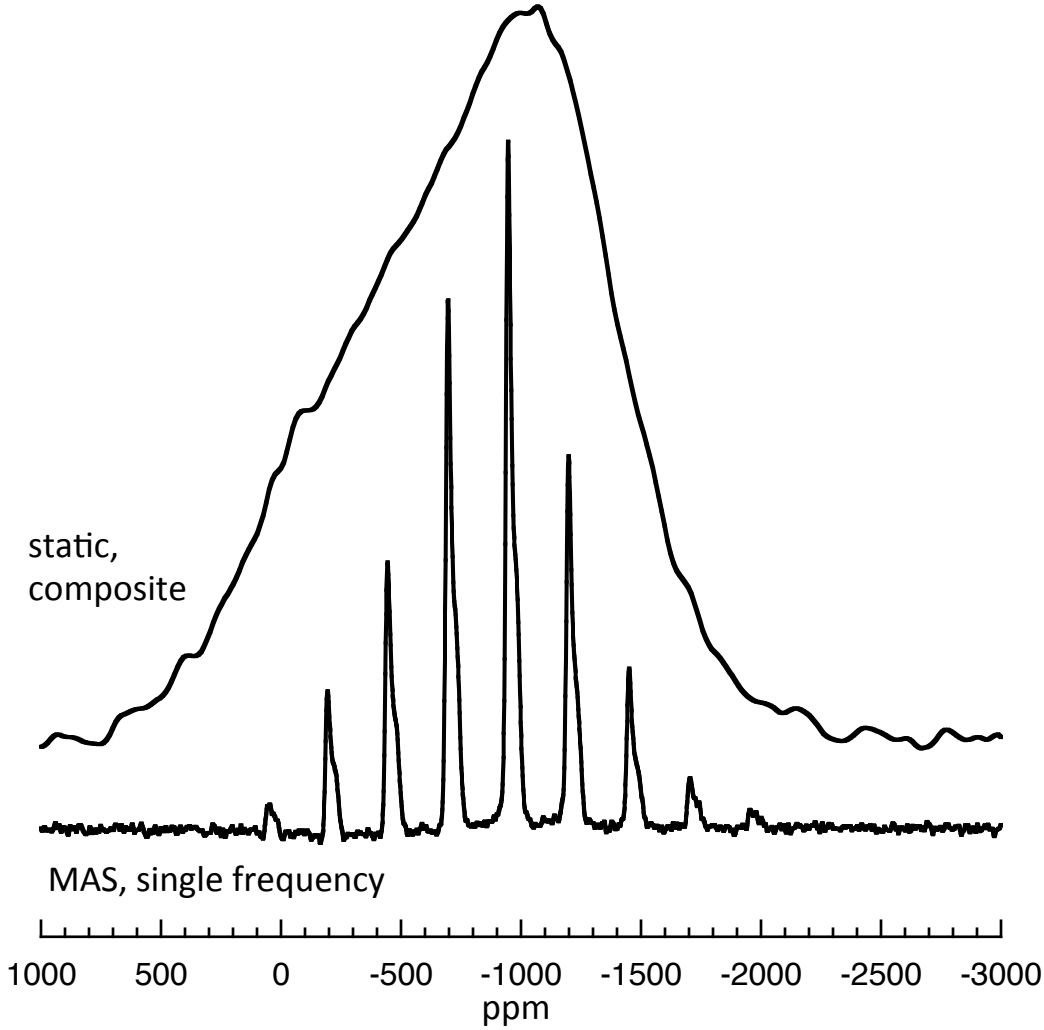


Figure 2

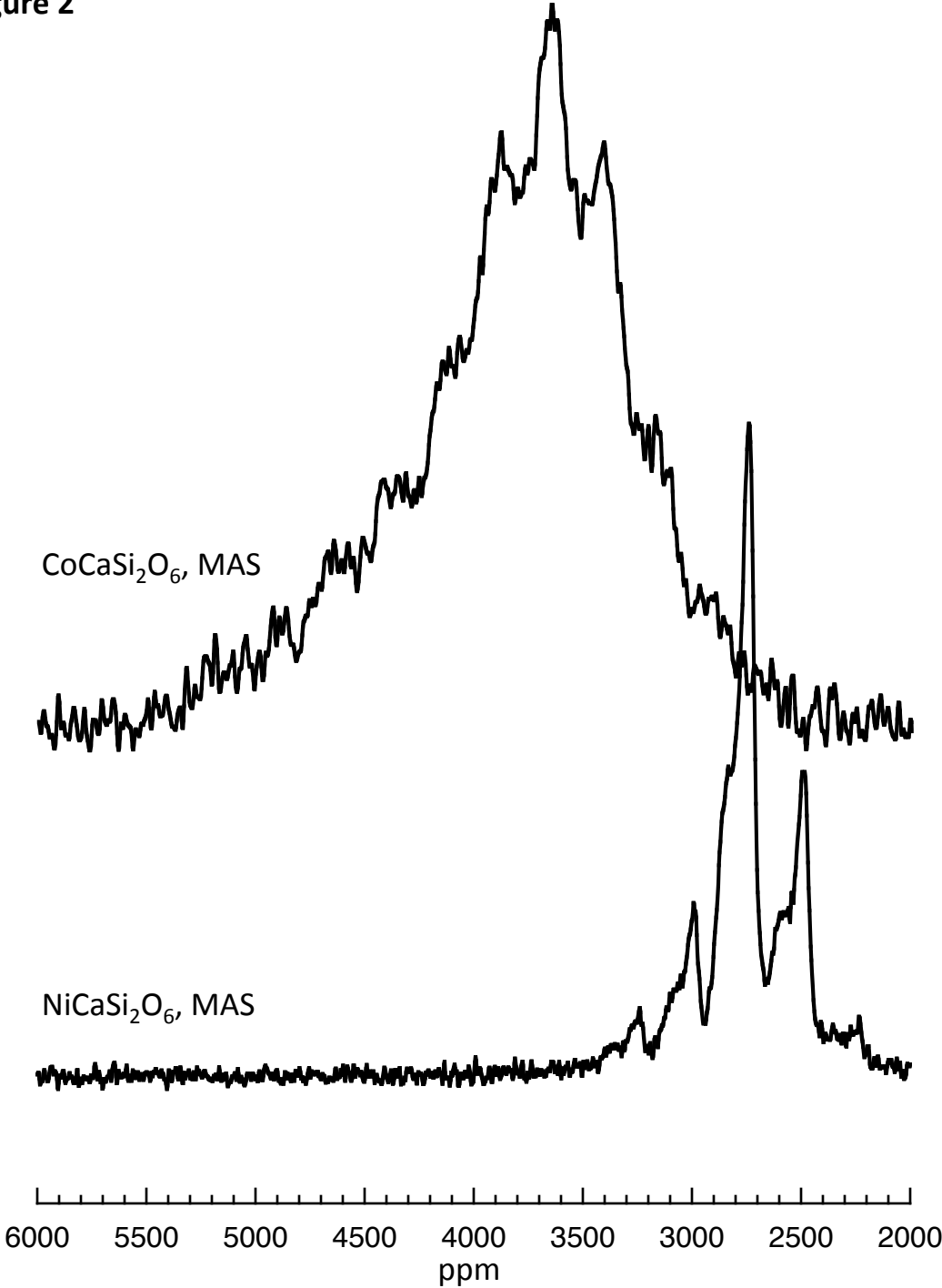


Figure 3

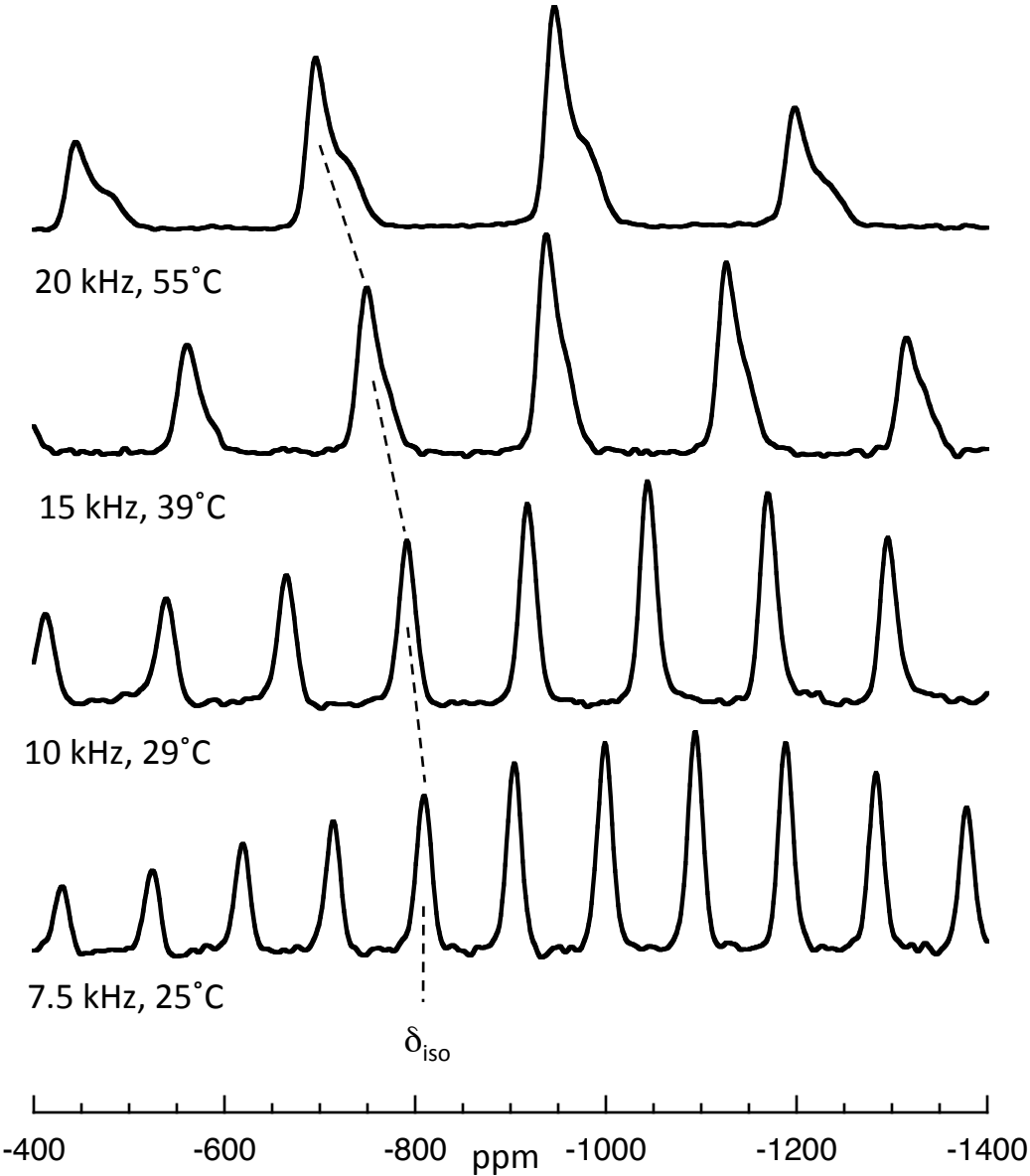


Figure 4

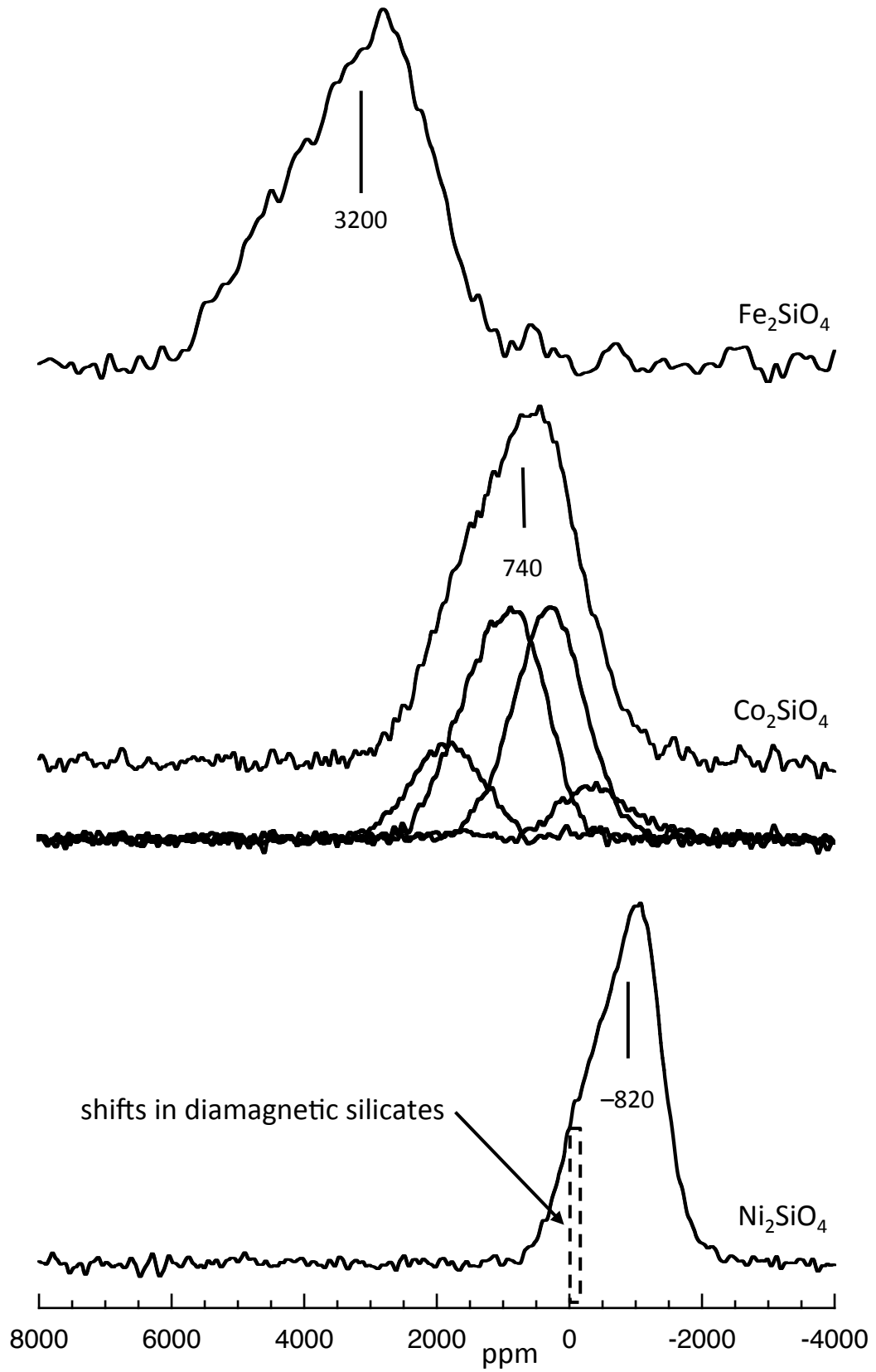


Figure 5

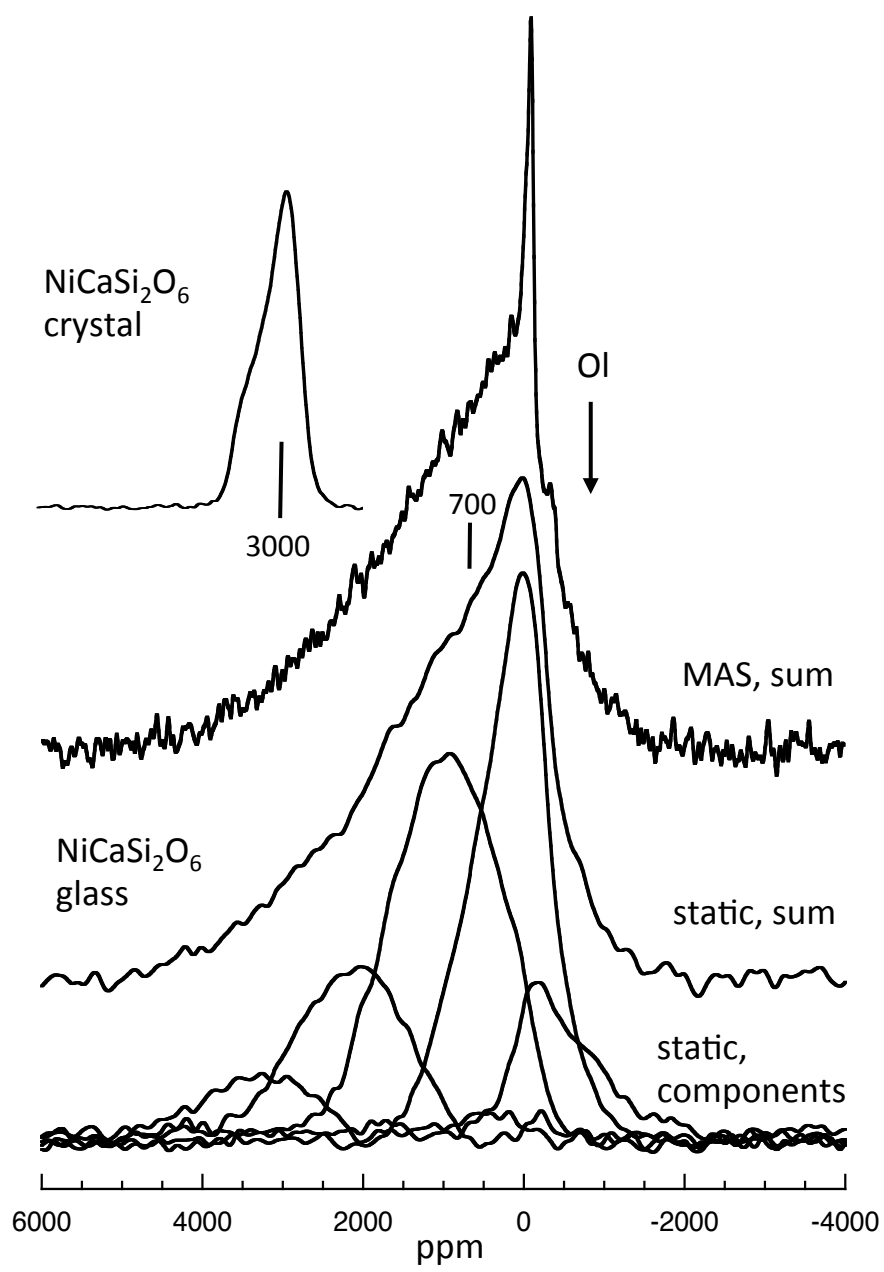


Figure 6

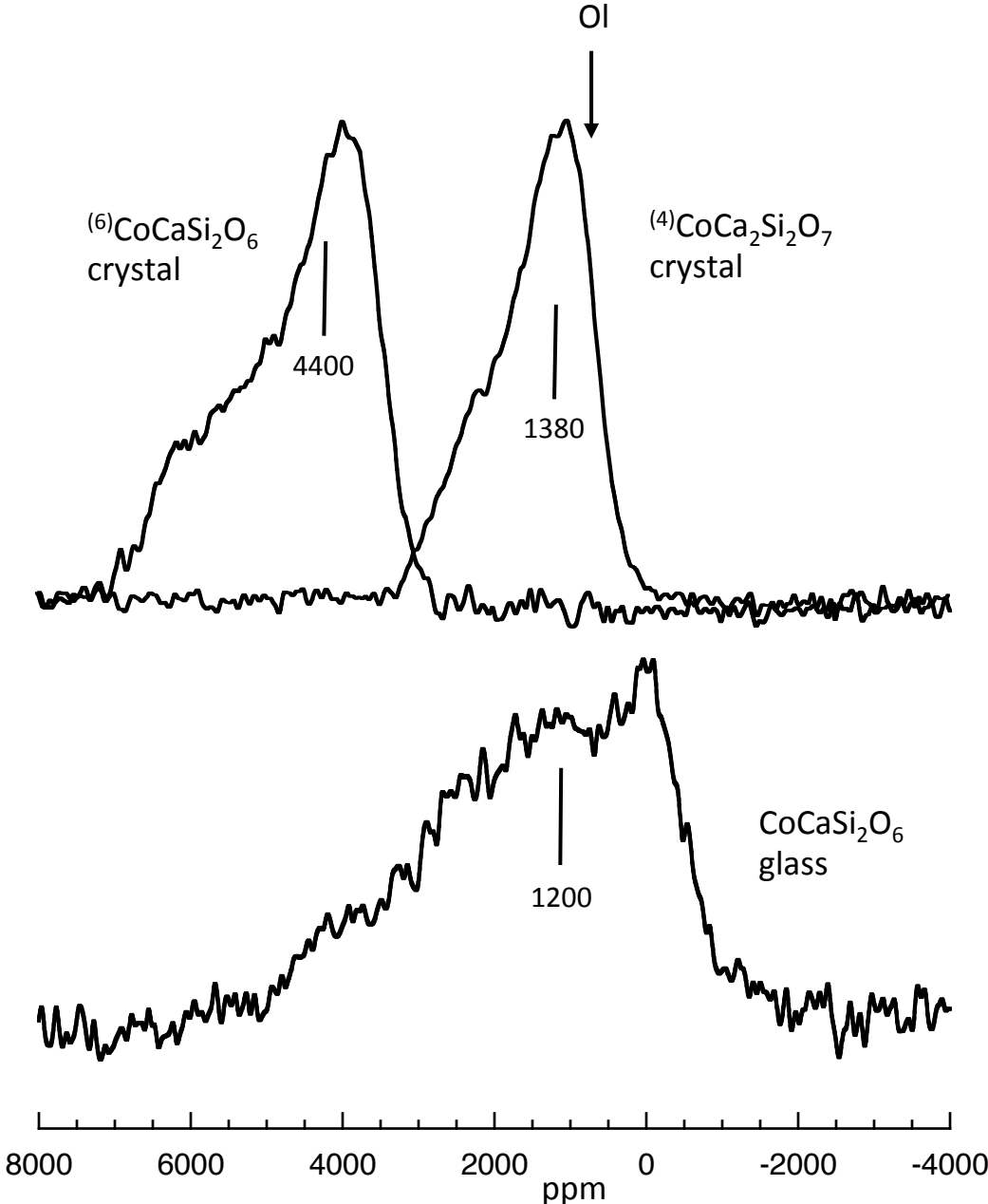


Figure 7

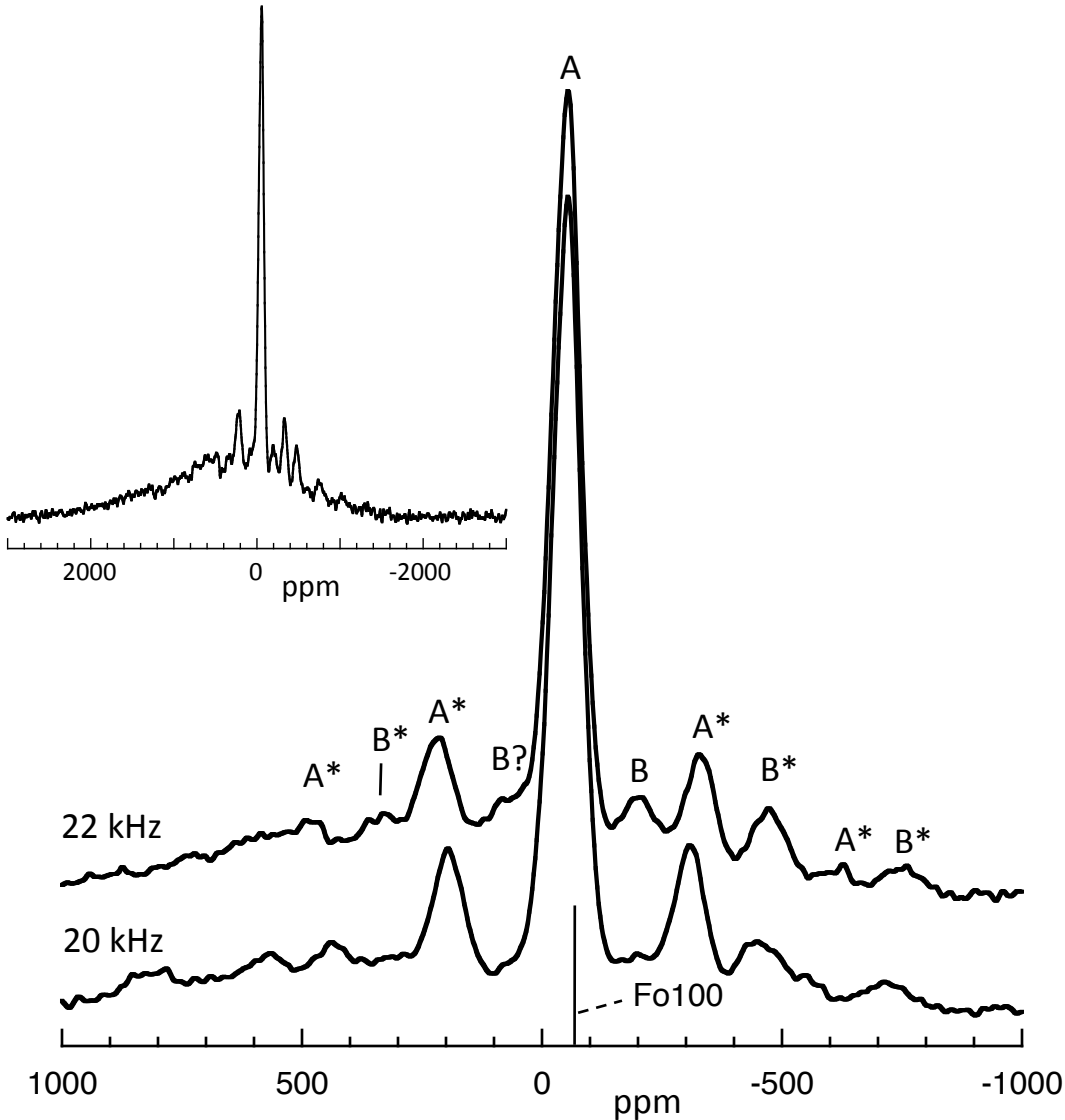


Figure 8

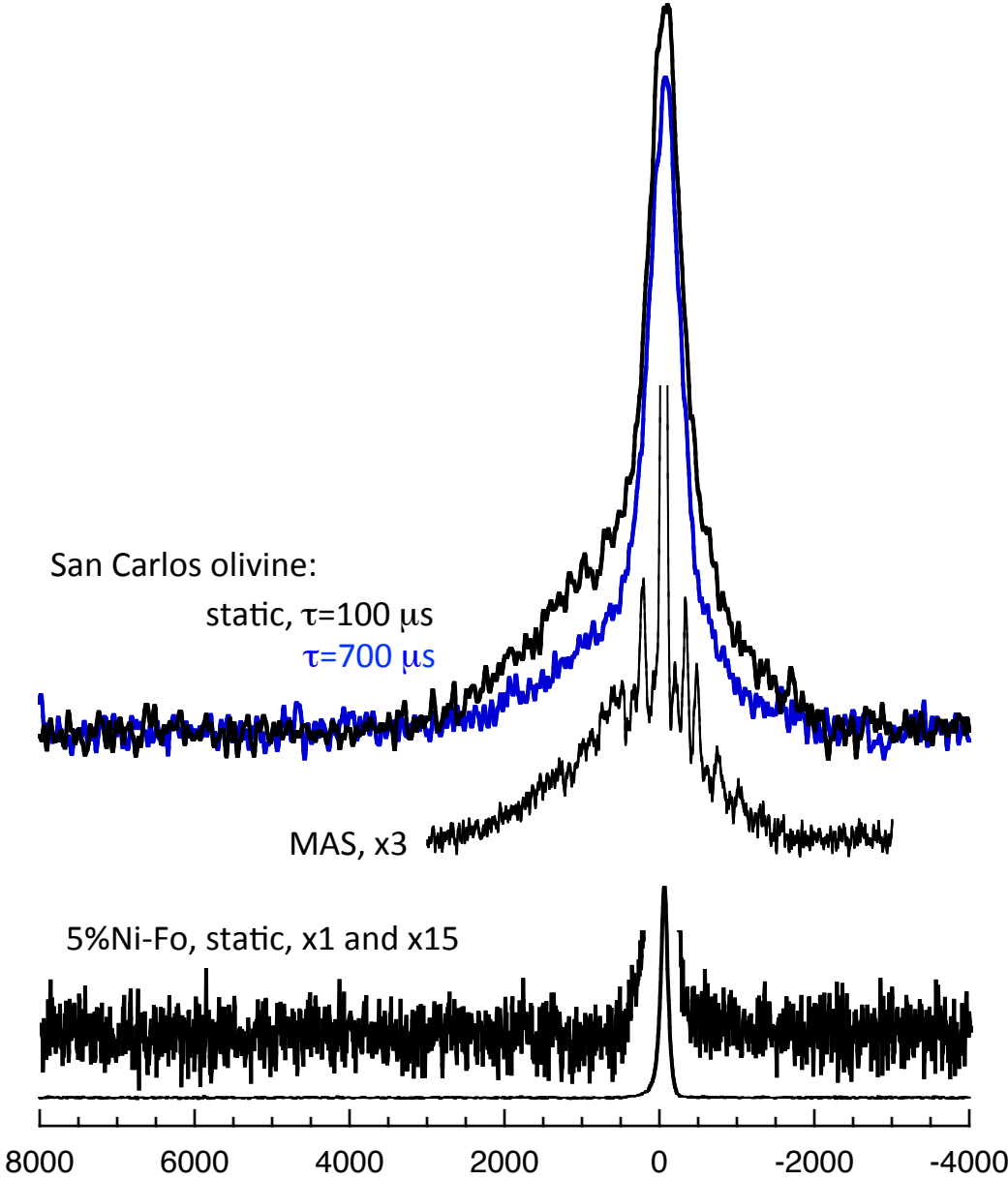


Figure 9

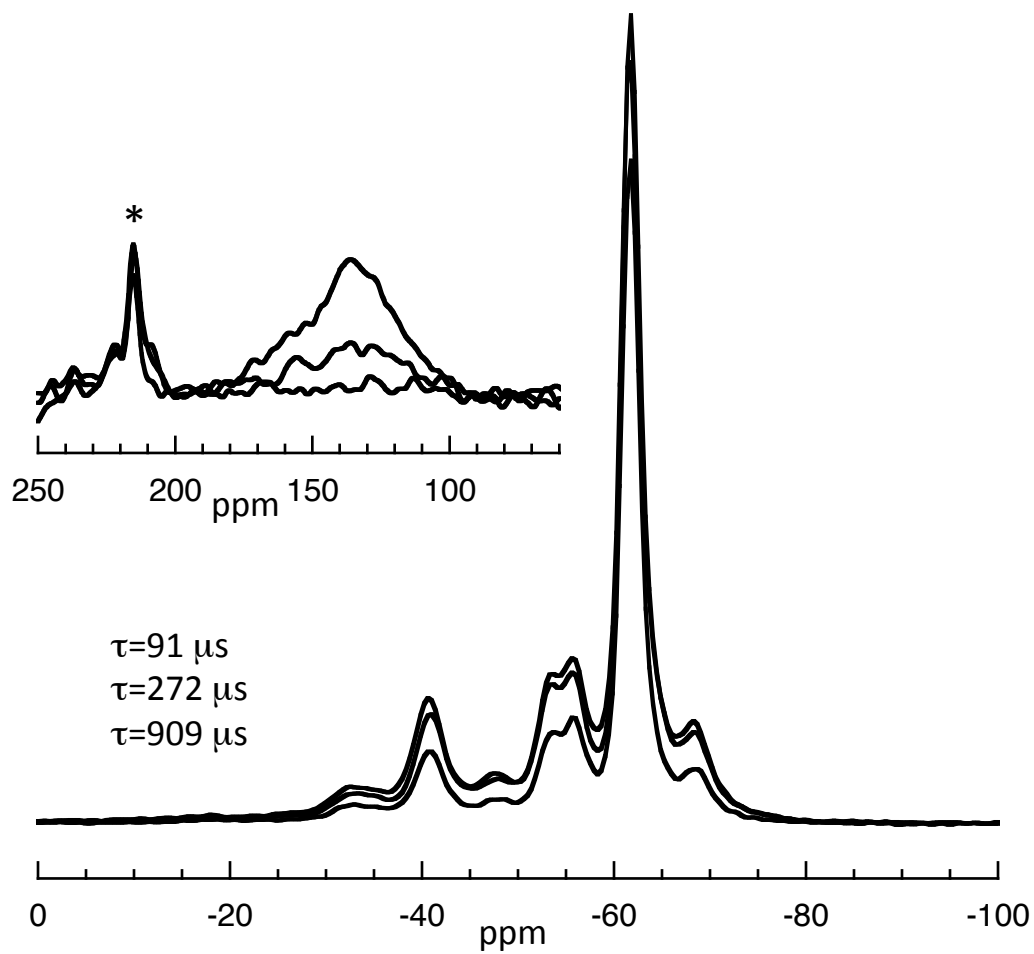


Figure 10

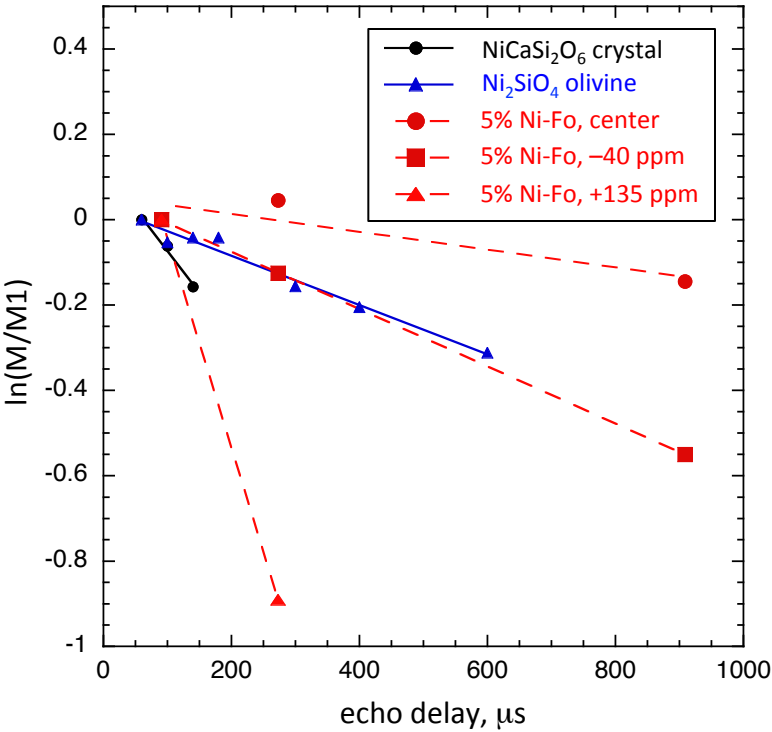


Figure 11

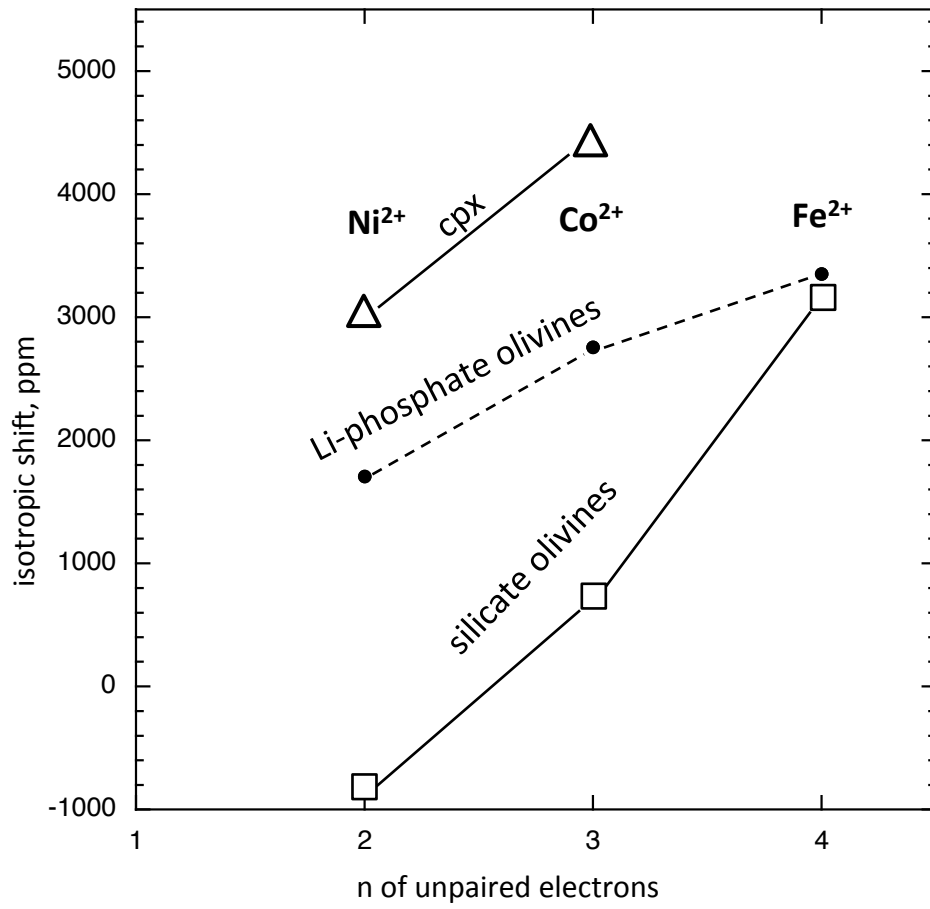


Figure 12

