1 1 2 **REVISION 1** 3 Word count: 8,986 4 Local strain heterogeneity associated with Al/Si ordering in anorthite, CaAl₂Si₂O₈, with 5 implications for thermodynamic mixing behaviour and trace element partitioning in plagioclase 6 7 feldspars 8 ALISON J. ATKINSON¹, TIZIANA BOFFA BALLARAN², AND MICHAEL A. CARPENTER¹ 9 10 ¹Dept. Earth Sciences, University of Cambridge, Downing Street, Cambridge CB2 3EQ, U.K. 11 ²Bayerisches Geoinstitut, Universität Bayreuth, D-95440 Bayreuth, Germany 12 13 14 ABSTRACT Hard Mode IR powder absorption spectroscopy has been used to characterise local strain relaxation 15 16 associated with Al/Si ordering in a suite of synthetic anorthite samples with structural states that vary 17 from a high degree of Al/Si order through a metastable incommensurate structure at intermediate states of order to long range order with $I\overline{1}$ symmetry. The dominant feature accompanying the changing 18 structural states is line broadening, which has been quantified by autocorrelation analysis and is 19 20 attributed to local heterogeneous strain variations on a length scale of at least 1-5 unit cells. The 21 autocorrelation results are consistent with contributions to the line broadening as being due to order parameters for both the $C\overline{1} \rightarrow I\overline{1}$ and $I\overline{1} \rightarrow P\overline{1}$ transitions which couple biquadratically, $\lambda Q_{od}^2 Q_{displ.}^2$ 22 Close correlation with enthalpy variations from previously published calorimetric data indicates that 23 24 the driving force for ordering can be understood in terms of elimination of strain fields arising from 25 accommodating more or less rigid AlO_4 and SiO_4 tetrahedra in the feldspar framework. The metastable Always consult and cite the final, published document. See http://www.minsocam.org or GeoscienceWorld

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incommensurate structure of anorthite is closely analogous to the stable incommensurate structure that develops at intermediate compositions in the plagioclase solid solution, confirming that the same strain relaxation mechanism dominates the properties and behaviour of all structural states across the solid solution. Elimination of strain heterogeneity by ordering on the basis of $I\overline{1}$ symmetry determines the form of non-ideal mixing shown by the solid solution at high temperatures, and changes in elastic properties may contribute to a break in slope of partitioning of trace elements between crystals and melt.

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Keywords: anorthite, Al/Si ordering, hard mode IR spectroscopy, local strain heterogeneity,
 enthalpy

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INTRODUCTION

38 A number of different strategies have been adopted to describe the thermodynamic properties of plagioclase feldspars in the context of efforts to produce quantitative descriptions of the stability fields 39 40 of mineral assemblages under the full range of possible geological conditions. Dubacq (2022) 41 summarised contrasting treatments of mixing properties for the $C\overline{1}$ and $I\overline{1}$ structures at high temperatures as those which smooth out the effects of the discrete $C\overline{1} \leftrightarrow I\overline{1}$ transition (Ghiorso 1984: 42 Benisek et al. 2004, 2010; Namur et al. 2012; Holland et al. 2022), those which treat states with 43 44 different symmetries as separate solid solutions (Holland and Powell 1992) and those which treat the mixing properties in terms of an Al/Si order/disorder process at anorthite-rich compositions (Carpenter 45 and Ferry 1984; Carpenter 1992, 1994). The differences revolve, in effect, around the need (or 46 otherwise) to include a physically correct description of the phase transition approximately half way 47 48 across the solid solution. This debate continues because mean field models of Al/Si ordering to give the 49 symmetry change explicitly, such as from Landau theory, do not readily yield equations that can 50 conveniently be incorporated into mixing models with a limited number of variable parameters. As a

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51 consequence, activity-composition descriptions based on conventional asymmetric mixing models are 52 used for applications in petrology, while models based on atomic interactions are used to understand 53 Al/Si ordering at a more fundamental level for albite, anorthite and the more complex incommensurate 54 structures that develop at intermediate compositions during slow cooling in nature.

55 A significant recent advance in the context of understanding the stability of plagioclase feldspars 56 from an atomic perspective has been provided by a Monte Carlo simulation by Dubacq (2022) which 57 focussed on interactions between nearest neighbour and next nearest neighbour tetrahedral cations. This author concluded that superposition of $I\overline{1}$ ordering on a $C\overline{1}$ solid solution, with large ordering 58 59 energies for anorthite, can provide reliable results in terms of energy and ordering across the solid 60 solution. However, he also drew attention to the fact that pairwise interaction models, such as used for 61 the Monte Carlo simulation, do not include the contribution of strain fields explicitly. If local strain 62 fields are less than the size of the supercell used in the model they are included implicitly but if they are 63 on a longer length scale they are not. This is a conundrum in all energetic modelling of crystal 64 structures and the present paper is intended to shed light on the influence of local strain effects from an 65 experimental point of view by analysing line broadening in IR absorption spectra.

66 The generalised result attributed to Eshelby (1957) that a strain field will decay through a medium in three dimensions according to $1/r^3$, where r is distance, implies that placing a specific limit on the 67 length scale of strain fields in a crystal is arbitrary. The expectation is at least that they extend further 68 69 than nearest or next nearest neighbour atoms, and there are experimental observations which confirm 70 this view for silicates. For example, diffraction contrast in a transmission electron microscope image of 71 exsolution of augite from pigeonite shows a strain field extending up to ~250 Å from the tip of an 72 augite precipitate (Carpenter 2002). Similarly, Figure 7.10 of Salje (1993) shows that the strain field 73 from a kink of individual domain wall in KAl₃SiO₈ can influence the topology of adjacent walls at 74 distances of ~ 1.000 Å. In a different context, the extent of the plateau of transition temperature as a 75 function of composition towards KAlSi₃O₈ for the displacive transition temperature near \sim 1325 K in

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albite, NaAlSi₃O₈, implies that the effective strain field around K⁺ has a diameter of 20-40 Å 76 (Carpenter et al. 1999). Length scales of 10's of Å overlap with estimates of the length scales of 77 phonons that are sampled in vibrational spectroscopy. The sensitivity of phonons to all the structural 78 variations in natural plagioclase crystals is seen in Raman spectra, for example (Bersani et al. 2018). 79 According to Atkinson et al. (1999) for feldspars, phonon length scales of 2-5 Å at 1500 cm⁻¹ would 80 become ~8-25 Å at 500 cm⁻¹ and ~60-150 Å at 50 cm⁻¹. The essential point is that phonon spectroscopy 81 provides, in principle, a probe with the correct length scale to detect and characterise structural 82 heterogeneity related to local strain relaxation in silicates. 83

84 The present study follows on from the Hard Mode IR powder absorption spectroscopy study of local 85 structural heterogeneity in natural and heat treated plagioclase feldspars by Atkinson et al. (1999). It 86 was designed to take advantage of a suite of synthetic anorthite samples prepared with a sequence of 87 metastable structural states starting from a high degree of Al/Si disorder, through incommensurate states with intermediate degrees of order to the equilibrium state with ordering based on $I\overline{1}$ symmetry. 88 A key advantage of this suite is that the composition is fixed (CaAl₂Si₂O₈) and, hence, that all changes 89 90 in spectral, physical and thermodynamic properties are due directly or indirectly to changes in Al/Si 91 order alone.

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EXPERIMENTAL METHODS

94 Synthesis and characterisation

The two series of synthetic anorthite samples used in the present study were selected from a larger suite of samples crystallised by heating glass of anorthite composition in air (Carpenter 1991a, 1991b, 1992). The first, referred to as a kinetic series, consisted of eight samples with crystallisation times ranging from 1 minute to 17 days at 1400 °C. The second, referred to as an equilibrium series, consisted of six samples equilibrated at temperatures between 850 °C and 1540 °C. A sample crystallised at 1100 °C for 15 minutes, ANC19, has been used as a reference state for maximum

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101 disorder of Al and Si. Times and temperatures for the original synthesis conditions are listed in Table 1, 102 together with details of structural states characterised by transmission electron microscopy, powder Xray diffraction and high temperature solution calorimetry (from Carpenter 1991a, 1991b, 1992). 103 The synthetic crystals all contained submicroscopic twinning parallel to (010) on the basis of albite 104 105 and Carlsbad twin laws (Carpenter 1991a, Xu et al. 1997). The presence of type-e and/or type-b 106 reflections in electron diffraction patterns defined whether the Al/Si order scheme was incommensurate or had evolved through to the commensurate $I\overline{I}$ state. Spacings and orientations of pairs of e-reflections 107 108 in samples from the kinetic series initially resembled the pattern seen in natural plagioclase samples with the e_2 structure, starting at composition $\sim An_{30}$ and evolving continuously through to the pattern for 109 110 crystals with composition $\sim An_{55}$. With further annealing, the spacing between e-reflections increased, but without further changes in orientation: ~25 Å for ANC19, ~30 Å for ANC60, extending to ~90 Å 111 for ANC73. Evolution of the microstructure was then of type-b antiphase domains with dimensions 112

increasing from ~200 Å in ANC73 Å to ~1,550 in ANC35. The size of type-b antiphase domains in crystals from the equilibrium series was in the range ~2,000-3,000 Å, but was not determined for ANC74/H2.

Diffuse type-c reflections arising from the $I\bar{1} \rightarrow P\bar{1}$ transition were present in diffraction patterns from all the samples, and type-d reflections were seen in diffraction patterns from the most ordered samples. The type-c and type-d reflections were sharp in diffraction patterns from ANC74/H2, which had been prepared initially by crystallisation from glass in air at 1400 °C and then annealed in the presence of water at 850 °C (Carpenter 1992). This treatment resulted in the highest degree of Al/Si order of all the samples listed in Table 1.

Macroscopic distortion of the structure due to Al/Si ordering was quantified in terms of the scalar strain, ε_s (Carpenter 1991b, 1992). Values were determined by first calculating six individual strain components, e_1 - e_6 , from changes in room temperature lattice parameters of each sample with respect to ANC19 as the reference state (e_1 - $e_6 = 0$) and then diagonalising the strain matrix to determine three Always consult and cite the final, published document. See http://www.minsocam.org or GeoscienceWorld

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principal strains $\varepsilon_1 - \varepsilon_3$ (Carpenter et al. 1990). The scalar strain was given by $\varepsilon_s = \sqrt{\sum_{1-3} \varepsilon_i^2}$ and increased with increasing degree of Al/Si order. ΔH_{soln} is the enthalpy of solution in lead borate flux at 700 °C (Carpenter 1991b, 1992). Samples with the lowest enthalpy (most ordered) have the largest values of ΔH_{soln} .

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131 Powder absorption infrared spectroscopy

Preparation of pellets used for powder absorption infrared spectroscopy measurements in the present study followed the procedure set out in detail in Atkinson et al. (1999). This involved a carefully repeated pattern of grinding the samples for 20 minutes in an agate ball mortar using a Spex microball mill, mixing in fixed proportions with CsI or polyethylene powder as the matrix phase and pressing under vacuum. CsI pellets weighed 300 mg and contained a ratio of sample to matrix of 1:450. Polyethylene pellets weighed 100 mg and contained a mixture of sample to matrix in the ratio 1:50.

138 Primary spectra were collected under vacuum at room temperature using a Bruker 113V FT-IR spectrometer and a Bruker 66V FT-IR spectrometer. Each spectrum was calculated by Fourier 139 transformation of 512 interferometer scans and recorded as absorbance, $a = -\log_{10}(I_{sample}/I_{reference})$, 140 where *I* is the single beam transmission intensity. The frequency ranges for spectra from the CsI pellets 141 were 150-700 cm⁻¹ and 400-2000 cm⁻¹, with an instrumental resolution of 2 cm⁻¹. The frequency range 142 of spectra obtained from the polyethylene pellets was 50-500 cm⁻¹, with an instrumental resolution of 4 143 cm⁻¹. Individual spectra were merged to produce the complete spectra shown in Figure 1 using 144 OPUS/IR (software of Bruker Analytische Messtechnik GmbH) by matching up sections in the region 145 $\sim 200-400 \text{ cm}^{-1}$ where they overlap. 146

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148 Quantitative analysis of primary spectra

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Rather than following the conventional approach of identifying individual absorption peaks and assigning their symmetries, the objective of the quantitative analysis presented here was to extract spectral parameters that relate specifically to the structural state and thermodynamic evolution of anorthite with respect to Al/Si order. This analysis was undertaken on the original, unmerged spectra, following the methodologies set out in Atkinson et al. (1999).

154 Three sections of the spectra were selected for peak fitting on the basis that they showed the 155 emergence of peaks as the degree of Al/Si order increased. The emerging peaks included a pair in the interval 300-320 cm⁻¹ (Fig. 2a, b) and another pair in the interval 460-500 cm⁻¹ (Fig. 2c, d). For each 156 pair, the area under a wider frequency interval was first normalised so as to minimise the influence of 157 158 any slight variations of the concentration of sample in the pellets or of background levels in the 159 spectrometers. This reduced scatter but did not change the trend shown by the same analysis of the original, un-normalised, spectra. Figure 3 illustrates six fits using Lorentzian profiles for individual 160 peaks in the frequency interval $\sim 260-340$ cm⁻¹, including shoulders of the peaks to either side of the 161 group of interest in order to obtain a realistic background. The frequencies, $\omega_{1,320}$, $\omega_{2,320}$, and line 162 widths, $\Gamma_{1,320}$, $\Gamma_{2,320}$, of the two peaks had uncertainties from the fits on the order of 0.2 cm⁻¹ (Salje et 163 al. 1989) for most of the spectra. The uncertainties were greater than this for spectra from the least 164 165 ordered samples. Areas of all the Lorentzian peaks within the group were summed to give a single overall intensity, I_{320} . This value had less scatter than the individual peak intensities as the relative sizes 166 of the peaks within the group were not well constrained by the fitting procedure. The same procedure 167 as set out above for the pair of peaks in the interval 300-320 cm⁻¹ was used to determine values of 168 frequencies, $\omega_{1,470}$, $\omega_{2,470}$, width, $\Gamma_{1,470}$, $\Gamma_{2,470}$, and combined intensity, I_{470} , of the two peaks between 169 460 and 500 cm^{-1} in Figure 2c.d. 170

The third focus of peak fitting was on an absorption peak with frequency close to 602 cm⁻¹, using
the same approach as for the peaks near 320 and 470 cm⁻¹. By comparison with spectra from natural
samples across the plagioclase solid solution (Atkinson et al. 1999) and the temperature dependence of
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spectra through the $I\overline{1} \leftrightarrow P\overline{1}$ transition in anorthite (Redfern and Salje 1992), this peak emerges most 174 175 clearly when the crystals have $P\overline{1}$ symmetry. It is not visible in the spectrum from ANC19 (0.25 hours 176 at 1100 °C), but developed with increasing intensity in the kinetic series (Fig. 1a) and is visible in all 177 spectra for the equilibrium series (Fig. 1b). It can also be seen in Figure 1 of Atkinson et al. (1999) 178 which includes spectra from natural samples with compositions in the range $\sim An_{71}-An_{100}$ and heat 179 treated natural samples in the range $\sim An_{78}$ -An₁₀₀. In the stack of spectra shown for Monte Somma anorthite as a function of temperature in Figure 2 of Redfern and Salje (1992), it emerges as an 180 increasingly distinct peak below ~700 K, ahead of the $I\overline{1} \leftrightarrow P\overline{1}$ transition at 515 ± 5 K. 181

Following the principles set out in Salje et al. (2000) and the approach of Atkinson et al. (1999) for analysing essentially the same groups of absorption peaks in spectra from natural plagioclase samples, the autocorrelation function was used to obtain an effective average linewidth for three different segments of the spectra: 270-450 cm⁻¹, 450-800 cm⁻¹, and 800-1500 cm⁻¹. These have been assigned labels Δcorr_{350} , Δcorr_{600} and $\Delta \text{corr}_{1000}$, respectively. Δcorr values in Atkinson (1999) for natural samples were for frequency ranges 260-530, 515-830, 820-1400 cm⁻¹.

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RESULTS

Primary spectra shown in Figure 1 closely resemble those reported previously from natural plagioclase samples (Salje et al. 1989; Redfern and Salje 1992; Atkinson et al. 1999). Low symmetry and large unit cell sizes result in multiple overlapping peaks with peak frequency variations, line widths, intensities and line broadening parameters that are sensitive to the degree of Al/Si order.

Figure 4 contains data for ω and Γ from fitting of the emerging peaks with frequencies near 320 and 470 cm⁻¹ in Figure 2 and near 600 cm⁻¹ in Figure 1. The kinetic and equilibrium series exhibit a welldefined trend of varying peak widths and almost no change in peak frequencies, with overlap for samples which contained type-b reflections in electron diffraction patterns ($I\overline{I}$ structure). There is no obvious break in slope at the point where type-e reflections gave way to type-b reflections. Data points Always consult and cite the final, published document. See http://www.minsocam.org or GeoscienceWorld

for ANC19, the sample with the lowest degree of Al/Si order, plot off the main trends but are relatively poorly constrained by the fitting procedure because the individual peaks were only weakly resolved in the primary spectra. Variations in peak widths correlate with changes in other parameters extracted from the spectra, such as peak intensities (Figure 5).

Each of the autocorrelation parameters $\Delta corr_{350}$, $\Delta corr_{600}$, and $\Delta corr_{1000}$, including those for ANC19, decreases systematically with increasing Al/Si order. $\Delta corr_{350}$, $\Delta corr_{600}$ correlate linearly with each other (Fig. 6a). Comparison of $\Delta corr_{350}$ and $\Delta corr_{600}$ with $\Delta corr_{1000}$ shows data points for the most ordered sample (which gave sharp type-c and type-d reflections) as deviating from a linear correlation (Fig. 6b). Figure 6c shows that decreasing $\Delta corr values correlate in a nonlinear manner with fitted peak$ widths, and that the nonlinearity increases with increasing order.

209 Line broadening parameters, Γ and Δ corr, provide measures of the local heterogeneity within each sample via the spread of frequencies of specific absorption modes on the phonon length scale. A 210 211 measure of the changes in Al/Si order averaged over a longer length scale, corresponding to at least the 212 coherence length of X-ray diffraction, and hence to the macroscopic order parameter(s), is provided by 213 the scalar strain, ε_s . $\Delta corr values display a nonlinear dependence on <math>\varepsilon_s$ (Fig. 7a), but Γ for the peak near 600 cm⁻¹ has a linear dependence (Fig. 7b). Data in Figure 7 taken from Atkinson et al. (1999) and 214 215 Atkinson (1999) for two natural samples of anorthite before and after being heat treated at 1300 °C or 216 1360 °C have been added and are consistent with the overall pattern of correlations.

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DISCUSSION

The clear outcome from analysing IR spectra from synthetic anorthite samples in this way is that the dominant influence of Al/Si disorder is seen as broadening of absorption peaks. In direct contrast, peak frequencies change very little through the full sequence of metastable states from those with the weakest development of incommensurate ordering through to equilibrium degrees of order based on the

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223 $I\overline{I}$ structure. ε_s represents the total long range elastic relaxation of the structure which occurs when 224 more or less rigid AlO_4 and SiO_4 tetrahedra arrange themselves in crystals of anorthite such that, in the 225 limit, there are no remaining Al-O-Al linkages. If ordering is incomplete, the net macroscopic strain, as 226 defined with respect to the disordered state, is less and local distortions where the arrangements of 227 adjacent tetrahedra do not conform to Al-avoidance occur instead. Line broadening parameters 228 obtained from the IR spectra provide a view of the resulting local strain heterogeneity depending on the 229 length scale of the phonons being sampled. As noted in the introduction above, these lengths scales are not well defined but are likely to be in the range $\sim 10-100$ Å. In this case, the observed $\Delta corr$ variations 230 231 are indicative of local structural heterogeneity on a scale of at least \sim 1-5 unit cells

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233 Hard modes and two order parameters

A conventional expectation of hard mode spectroscopy is that changes in peak widths, peak 234 235 frequencies and peak intensities will scale, in lowest order, with the square of the order parameter 236 (Salje 1992, 1994; Salje and Bismayer 1997). If this was simply Q_{od} for Al/Si ordering accompanying the change in macroscopic symmetry $C\overline{1} \rightarrow I\overline{1}$, each of the parameters, ω , Γ , Δ corr and Intensity would 237 238 be expected to vary linearly both with each other and with the macroscopic strain, since each would be expected to scale with Q_{od}^2 . While there is a general correlation among the line broadening parameters 239 240 and strain, the trends are not linear and a contributary factor to this must be that, in anorthite at room temperature, these parameters depend on both $Q_{\rm od}^2$ for the $C\bar{1} \rightarrow I\bar{1}$ transition and $Q_{\rm displ}^2$ for the $I\bar{1} \rightarrow$ 241 $P\overline{1}$ transition. 242

Samples with the highest degree of order display the most obvious deviations from linear relationships between spectral parameters. They also have changes in the intensity and sharpness of the type-c and type-d reflections arising from the displacive transition which indicate that they have the largest values of Q_{displ} . It is well understood that there is strong, biquadratic coupling between the two

order parameters, ie with the form $\lambda Q_{od}^2 Q_{displ}^2$, where λ is the coupling coefficient (Salje 1987; Redfern 247 248 and Salje 1987; Redfern et al. 1988; Redfern 1992). The linear relationship between Γ for the peak near 600 cm⁻¹ and the scalar strain seen in Figure 7b implies that both record the two order parameters in the 249 same proportion. Likewise, the linear relationship shown in Figure 6a implies that $\Delta corr_{350}$ and $\Delta corr_{600}$ 250 251 also have a dependence on the same proportions of the two order parameters as each other. Non-linear relationships seen in Figures 6b,c and Figure 7a imply that the dependence of $\Delta corr_{1000}$ and Γ for peaks 252 near 320 and 470 cm⁻¹ depend on a combination of Q_{od}^2 and Q_{displ}^2 which differs from those of Γ for the 253 600 cm⁻¹ peak, $\Delta corr_{350}$, $\Delta corr_{600}$ and ε_s . Some additional measure of the order parameters would be 254 255 needed to separate these dependences quantitatively, however.

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257 $I\overline{1} \rightarrow P\overline{1}$ transition

Due to biquadratic coupling, local variations in the degree of Al/Si order are necessarily 258 accompanied by local variations in displacements relating to the $P\overline{1}$ structure. Although the degree of 259 Al/Si order in the synthetic samples has not been determined directly, Q_{od} is known to vary between 260 0.92 for a natural metamorphic anorthite sample and 0.78 after equilibration at ~1535 °C (Angel et al. 261 262 1990; Carpenter et al. 1990). The observation that samples with relatively high degrees of Al/Si 263 disorder gave diffuse type-c and type-d reflections in electron diffraction patterns is consistent with the 264 original observation by Bruno et al. (1976) that anorthite crystals quenched from ~1530 °C did not have 265 type-c and type-d reflections in single crystal X-ray diffraction patterns.

Salje (1987) considered that the length scale of correlated displacements of atoms on the basis of $P\overline{1}$ symmetry would be reduced to less than ~200 Å, as an estimate for the typical length scale of X-ray diffraction, by local strain heterogeneity arising from the introduction of Al/Si disorder. This would also cause the $I\overline{1} \rightarrow P\overline{1}$ transition to become smeared rather than occurring at a discrete temperature. Spontaneous strain data extracted from lattice parameters measured as a function of temperature for a

sample equilibrated at 1302 ± 4 °C did not show evidence of such smearing (Redfern 1992), but the effect of disordering was seen increasingly in spontaneous strains from lattice parameters collected as a function of pressure through the related transition as a function of pressure in crystals equilibrated at temperatures between ~1350 ± 5 °C ($Q_{od} = 0.87$) and ~1533 ± 5 °C ($Q_{od} = 0.78$) (Angel 1992, 1994). Confirmation of the local heterogeneity responsible for this behaviour has been provided by the IR spectroscopy results reported here.

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278 Elastic strain heterogeneity and enthalpy

279 Figure 8 contains a compilation of Δ corr data for synthetic anorthites from the present study with 280 data from Atkinson et al. (1999) and Atkinson (1999) plotted as a function of composition. Dashed 281 lines are guides to the eye illustrating linear extrapolation to An_{100} of fits to data in the range $\sim An_{20}$ -282 An₆₀ for heat treated series samples which have $C\overline{1}$ symmetry. Each of the Δ corr parameters has the same pattern of a break in slope when $C\overline{1}$ members of the solid solution give way to $I\overline{1}$ members. Each 283 284 also hasvalues for synthetic anorthites which spread between natural, ordered anorthite and the extrapolated limit for $C\overline{1}$ structures. When expressed simply in terms of these parameters alone, 285 deviations from the trend for a solid solution with $C\overline{1}$ symmetry are clearly due to the Al/Si ordering. 286

287 Correlations between enthalpy and local strain heterogeneity from $\Delta corr$ for plagioclase feldspars 288 were discussed in detail by Atkinson et al. (1999) and are confirmed in Figure 9. Figure 9a shows 289 almost linear relationships between $\Delta corr$ and ΔH_{soln} for Al/Si ordering in anorthite, apart from the value of ΔH_{soln} for ANC19 which is considered to be lower than would be expected for the most 290 disordered crystalline state due to the presence of a small amount of residual glass (Carpenter 1991b). 291 Any non-linearity would be due predominantly to the additional contributions of Q_{displ}^2 to $\Delta corr$ but 292 might also include a slight dependence of the enthalpy of ordering on higher order terms than $Q_{\rm od}^2$ 293 294 alone (Carpenter 1992). Figure 9b is the equivalent for ΔH_{soln} as a function of composition to the $\Delta corr$ variations shown in Figure 8. Just as with $\Delta corr$, the spread of ΔH_{soln} values for synthetic anorthites is between values for ordered anorthite and the extrapolated end member of a solid solution with $C\overline{1}$ symmetry.

Spectral parameters showing line broadening in anorthite clearly correlate with enthalpy changes due to ordering, and the data in Figure 9b confirm the previous conclusion that deviations from ideality in the enthalpy of mixing for the solid solution of plagioclase in "high" structural states arises from Al/Si ordering in anorthite-rich compositions. The enthalpy of ordering in anorthite and anorthite-rich compositions is also closely associated with the elimination of local strain heterogeneity.

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304 Structural evolution from incommensurate to commensurate ordering

305 The initial Al/Si ordering scheme that develops in anorthite crystallised from chips of glass in air 306 displays characteristic features in diffraction patterns which mimic closely those of the incommensurate structures in natural samples with intermediate compositions in the plagioclase solid solution 307 308 (Carpenter 1991a). NMR spectroscopy on the same samples has shown that changes with annealing 309 time occur by apparently continuous reductions in the number of Al-O-Al linkages (Phillips et al. 1992). Elimination of Al-O-Al linkages might conventionally be understood in terms of the long 310 311 established principle of Al-avoidance or Loewenstein's rule whereby Al in silicates avoids occupying 312 tetrahedral sites adjacent to others also occupied by Al. However, as discussed at length by Dove et al. (1996), this view is simplistic because the Al/Si ordering transition temperature for some silicates is so 313 314 low that ordering does not occur even on a geological timescale.

At the length scale of phonons, between the coherence length of X-ray diffraction and the near neighbour scale of NMR, local ordering in the incommensurate structure of anorthite evolving into that of the $I\overline{1}$ structure is accompanied by elimination of local strain. On this basis, the driving force for ordering comes from the need to minimise the elastic strain energy in structures where a disordered configuration of more or less rigid AlO₄ and SiO₄ tetrahedra with different sizes cannot be Always consult and cite the final, published document. See http://www.minsocam.org or GeoscienceWorld accommodated in a given structure without introduction of local strain heterogeneity. The stable incommensurate arrangement of tetrahedra in intermediate plagioclase and the related metastable arrangement in incommensurate anorthite just provides one way in which AlO₄ and SiO₄ units can be arranged to reduce the local strain energy. At intermediate compositions it provides the lowest energy scheme, while ordering on the basis of the commensurate $I\overline{1}$ structure allows further elimination of the local strains in anorthite.

Advanced refinements of the structure of the e_1 form of the incommensurate structure in a sample with composition An_{51} from a metamorphic rock have revealed additional structural changes accompanying Al/Si ordering (Jin and Xu 2017). The modulation period of ~30 A for this sample is within the length scale over which local strain fields related to the arrangements of AlO₄ and SiO₄ tetrahedra would extend. Adjustments of Na and Ca positions in response to relaxations of the local heterogeneity provide additional stabilisation of the incommensurate structure and involve a small degree of chemical segregation.

Finally, a high degree of short range order in the $C\overline{1}$ solid solution provides further demonstration of the fact that the three dimensional feldspar framework structure is singularly unsuitable for accommodating disordered arrangements of AlO₄ and SiO₄ tetrahedra without introducing large local strain fields.

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338 Strain and elasticity in thermodynamic mixing models of silicate solid solutions

A defining role for local and macroscopic strain energies fits with understanding of the mixing properties of silicate solid solutions more widely. The two dominant properties controlling excess enthalpies of mixing on cation sites are the difference in size of the substituting ions and the elastic constants of the host crystal (Carpenter and Boffa Ballaran 2001; Tarantino et al. 2003). Particularly striking in this context is the jadeite (NaAlSi₂O₆) – augite (Ca(Mg,Fe)Si₂O₆) solid solution for which changes in Δ corr and ΔH_{soln} due to ordering of cations at intermediate compositions are exactly the

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reverse of changes due to mixing of the disordered structure (Boffa Ballaran et al. 1998; Carpenter and Boffa Ballaran 2001). As in the case of anorthite presented in the present paper, cation ordering involves the conversion of local strain fields into a single, homogeneous macroscopic strain. It is unlikely that pairwise atomic interaction models will be able to mimic this behaviour unless they include the longer range strain relaxations.

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351 Elastic moduli for trace element partitioning between plagioclase crystals and silicate melt

352 As discussed above for cation ordering and solid solution formation, the model of Blundy and Wood (1984) for partitioning of trace elements between crystals and melt in silicate systems attributes the 353 354 energies involved to the elastic strain energy associated with substitution of one cation by another with 355 a different size. It depends on an effective Young's modulus for the site on which the substitution 356 occurs and this, in turn, is related to the Young's modulus of the bulk crystal. Mutch et al. (2022) recently reviewed treatments of the partitioning of Mg^{2+} between plagioclase crystals and melt that had 357 included a smooth variation of properties with composition across the plagioclase solid solution. They 358 359 found that a better match with experimental data was obtained if an inflection point was included in the 360 composition dependence of the partition coefficient at temperatures and compositions corresponding to the locus of the $C\overline{1} \leftrightarrow I\overline{1}$ transition. They also noted, however, that the values of bulk and shear moduli 361 362 reported by Brown et al. (2016) do not show any equivalent reversal in trends at anorthite-rich compositions. Instead, they attributed the change in partitioning behaviour to a crossover from 363 preferential location of Mg²⁺ in M sites to preferential occupancy of tetrahedral sites across the $C\overline{1} \leftrightarrow$ 364 $I\overline{1}$ boundary. 365

No single dataset is yet available for the Young's modulus of plagioclase crystals with different degrees of Al/Si order corresponding to those that would develop at the pressure/temperature conditions in the earth where crystals would coexist with melt. Shear and bulk moduli given by Brown et al. (2016) were all obtained at room temperature using natural crystals from a mixture of geological Always consult and cite the final, published document. See http://www.minsocam.org or GeoscienceWorld

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370 environments. As such, they are unlikely to reveal any systematic dependence of the elastic moduli on the degree of Al/Si order. An additional consideration is that the $I\overline{1} \rightarrow P\overline{1}$ transition must necessarily 371 contribute to changes in elastic properties of anorthite-rich crystals measured at room temperature, as 372 opposed to the temperatures at which trace element partitioning occurs. The influence of coupling 373 374 between Q_{displ} and strain has not been determined but would give changes in moduli with the form 375 typical of a co-elastic transition, is not dissimilar from the pattern shown through the $\alpha \leftrightarrow \beta$ transition 376 in guartz (Carpenter et al. 1998; Carpenter and Salje 1998) and on the order of a few percent to low tens of percent. On the other hand, the compilation of data shown in Figure 10 for the bulk modulus 377 378 determined by single crystal X-ray diffraction measurements as a function of pressure is sufficient to 379 suggest there is indeed a significant change in elastic properties associated with the $C\overline{1} \leftrightarrow I\overline{1}$ transition. 380 Data for crystals with different compositions were taken from Figure 5 of Brown et al. (2016) and data for anorthite crystals with different degrees of Al/Si order were taken from Table 2 of Angel (1994). 381 The data for K_0 in Figure 10 show two overlapping trends: increasing values with increasing 382 383 anorthite content and decreasing values with increasing Al/Si disorder. Information about the disordering is provided by the observation of a displacive phase transition, $P\overline{1} \rightarrow I\overline{1}$, with increasing 384 385 pressure (Angel et al. 1988, 1989; Angel 1988, 1992, 1994). Although the high pressure structure of anorthite with $I\overline{1}$ symmetry has some subtle differences from the $I\overline{1}$ structure at high temperatures, K_0 386

387 for the former varies systematically according to the temperature at which the crystals had previously

been annealed. Anorthite crystals equilibrated at $1533 \pm 5 \,^{\circ}C \,(Q_{od} = 0.78)$ gave $K_o = 66.3 \pm 1.7$ GPa, crystals equilibrated at $1350 \pm 5 \,^{\circ}C \,(Q_{od} = 0.87)$ gave $K_o = 77.1 \pm 1.7$ GPa, and crystals equilibrated during slow cooling in nature ($Q_{od} = 0.92$) gave $K_o = 88.6 \pm 3.7$ GPa (Angel 1994). A similar dependence of K_o on Q_{od} for other plagioclase crystals with $I\bar{1}$ symmetry would decrease with reducing anorthite content towards the $C\bar{1}/I\bar{1}$ boundary.

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For reasons of stoichiometry, full ordering on the basis of $C\overline{1}$ symmetry is restricted to pure albite, 393 and members of the solid solution with compositions that fall on the $C\overline{1}$ side of the $C\overline{1}/I\overline{1}$ line do not 394 show the same extent of Al/Si order variation. The small break in slope at ~50 % An for K_0 of natural 395 samples in Figure 10 is not outside realistic experimental uncertainty limits but would be amplified 396 397 significantly if the anorthite-rich crystals had a degree of Al/Si disorder. The expectation, therefore, is 398 that for temperatures corresponding to those at which plagioclase crystals coexist with a basaltic liquid 399 there should be a maximum in K_0 half way across the solid solution. In the Blundy and Wood model, 400 such a break in slope of bulk modulus with composition would contribute to the pattern of partition coefficients for Mg²⁺ between plagioclase and melt reported by Mutch et al. (2022), though the 401 magnitude of the contribution remains undefined. 402

403 In the absence of a more comprehensive data set for crystals with wide ranges of composition and order, variations of spectral parameters such as $\Delta corr_{570}$ and $\Delta corr_{1000}$ can provide a prediction of the 404 form of variations of K_0 because of their sensitivity to Al/Si ordering. The relationship is explicit for 405 anorthite because observed values of K_0 for both the high pressure $I\overline{1}$ structure and the $P\overline{1}$ structure 406 vary linearly with Q_{od}^2 , due to a coupling term between spontaneous strain, e, and Q_{od} of the form 407 $\lambda e^2 Q_{od}^2$ (Carpenter 2006). On this basis, the data in Figure 8 imply that the dependence of K_0 on the 408 409 degree of Al/Si order should be greatest for pure albite and pure anorthite and smallest for crystals with compositions in the range ~20-50 % An. The sensitivity of K_0 to Q_{od} should increase with anorthite 410 411 content and there should be a non-linear variation of K_0 with composition for suites of crystals which 412 have crystallised at more or less the same temperature. The same arguments apply to the shear modulus but there are currently no data available to separate any influence of the $I\overline{1} \rightarrow P\overline{1}$ transition from the 413 414 room-temperature data of Brown et al. (2016). In order to correlate the real influence of changing bulk 415 and shear moduli on trace element partitioning, it will be necessary to determine their values for

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416 samples with degrees of Al/Si order equilibrated at liquidus temperatures as measured at temperatures 417 above the $I\overline{1} \rightarrow P\overline{1}$ transition.

418 A final consideration in this context is that, because there must be some degree of strain variation 419 through twin walls in a crystal, partition coefficients between bulk crystal and melt must be different 420 from those for partitioning between twin walls and melt. Changes in the concentration of cations have 421 been seen at transformation twin walls in anorthoclase, for example (Hayward et al. 1998). Albite-rich 422 crystals typically have a higher density of growth twins than An-rich crystals and this ought to 423 influence the overall pattern of partitioning with composition. It remains to be determined whether the 424 number of crystallographic sites on and immediately adjacent to twin walls, as a proportion of the total 425 number of sites in a crystal, would be sufficient for this effect to be detectable.

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CONCLUSIONS

Autocorrelation analysis of line broadening in IR powder absorption spectra from synthetic anorthite 428 429 samples with a wide range of structural states has shown that the predominant effect of Al/Si ordering 430 in the feldspar structure is to reduce local strain heterogeneity. Close correlation of line broadening 431 parameters with calorimetric data has then shown that the enthalpy of ordering can be understood in 432 terms of the accompanying reduction in elastic strain energy. Comparisons of line broadening parameters with calorimetric data for members of the plagioclase solid solution confirm the view that 433 the thermodynamic mixing properties of plagioclase feldspars depend essentially on these 434 435 order/disorder processes. Given that strain fields generate long ranging interactions within a structure, 436 it follows that mean field models, such as Bragg Williams or Landau, should provide physically 437 realistic descriptions of the overall thermodynamic behaviour. More generally, the Al-avoidance 438 principle or Loewenstein's rule is expected to apply when the geometry of a structure does not allow 439 relaxation of local strain effects arising from the difference in size of AlO₄ and SiO₄ tetrahedra but not when the geometry allow local strains to be eliminated by relaxation. The same constraints of strain 440

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and elasticity that apply to solid solution formation also apply to incorporation of trace elements and,hence, to partition coefficients.

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IMPLICATIONS

Key outcomes are firstly that the dominant effect of Al/Si disordering in anorthite is to give rise a 445 446 high degree of local strain heterogeneity, secondly that Al-avoidance can be understood completely in 447 terms of the need to minimise the strain associated with accommodating the different sizes of relatively 448 rigid AlO₄ and SiO₄ tetrahedra in a framework structure, and thirdly that the reduction of the local 449 heterogeneity with increasing order correlates with changes in macroscopic strain and enthalpy. While 450 activity models of the plagioclase solid solution can skate over these effects, any physically accurate 451 description of the structures will need to incorporate the effects of strain relaxation on a scale of at least 452 a few unit cells. This treatment completes the picture of strain and elasticity as being the predominant 453 factors in determining the properties of minerals which undergo displacive transitions, such as in 454 stishovite (Carpenter et al. 2000), the driving mechanism for cation ordering transitions, such as in 455 omphacite (Boffa Ballaran et al. 1998) and anorthite (this study), the mixing behaviour of solid 456 solutions such as feldspars, pyroxenes, amphiboles and garnet (Carpenter and Boffa Ballaran 2001), 457 and partitioning of trace elements such as in the Blundy-Wood model (Blundy and Wood 1994).

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FIGURE CAPTIONS

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FIGURE 1. (a) Merged infrared powder absorption spectra for the kinetic series of synthetic anorthites. The annealing time for each sample is given in hours. All samples were heated at 1400 °C except the lowest in the figure, which was heated at 1100 °C. (b) Merged spectra for the six equilibrium samples. Annealing temperatures are given in °C. The spectrum of the sample equilibrated at 1400 °C (with an annealing time of 1102 hours) has been also included in (a).

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FIGURE 2. Details from infrared powder absorption spectra showing the emergence of new peaks at around \sim 320 cm⁻¹ for (a) the kinetic series and (b) the equilibrium series samples and at \sim 470 cm⁻¹ for (c) kinetic and (d) equilibrium samples. Note the change in intensity with time or annealing temperature, reflecting changes in the degree of local order.

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FIGURE 3. Examples of peak fitting of peak group at 300-320 cm⁻¹ in spectra from synthetic anorthite samples with different degrees of Al/Si order. Equilibrium series: (a) 850 °C, (b) 1400 °C, (c) 1536 °C. Kinetic series: (d) 8.25 hrs, (e) 0.78 hrs, (f) 0.017 hrs. The fits are shown as solid lines and the original spectra as dashed lines. The number of peaks used in the fit was arbitrary and does not necessarily reflect the real number of modes present. The intensity of the peak group was obtained by the summation of the areas of the peaks used in the fit.

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FIGURE 4. Values of ω and Γ from peak fitting of selected peaks in primary spectra. Experimental uncertainties are shown as ± 0.2 cm⁻¹, but uncertainties in the data for sample ANC19 are larger than this. Data for samples with type-b reflections in the kinetic series generally overlap with data from the equilibrium series. Any sensitivity of phonon frequencies to Al/Si order is clearly substantially less than the sensitivity of line widths. (a) Peak with the higher frequency of the pair near 320 cm⁻¹ in Figure 2a,b. (b) The pair of peaks in Figure 2c,d. (c) Peak near 602 cm⁻¹ in Figure 1.

FIGURE 5. Variation of linewidths, Γ, from peak fitting with respect to the integrated intensities of peaks. Γ values labelled as 320 were taken directly from the fits to the peak at ~322 cm⁻¹. Γ values labelled as 470 are averages for the two peaks at ~470 and ~490 cm⁻¹. I_{320} and I_{470} are the sum of intensities for each of the two pairs of peaks. Black lines are guides to the eye showing linear dependences between Γ and intensity.

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FIGURE 6. Variations in parameters from autocorrelation analysis. $\Delta corr_{350}$, $\Delta corr_{600}$ and $\Delta corr_{1000}$ 616 refer to frequency intervals 270-450 cm⁻¹, 450-800 cm⁻¹, and 800-1500 cm⁻¹, respectively. (a) 617 618 Variations of $\Delta corr_{350}$ and $\Delta corr_{600}$ show essentially the same variation for the two frequency intervals. (b) $\Delta corr_{350}$ and $\Delta corr_{600}$ show linear correlation with $\Delta corr_{1000}$ apart from data points for the most 619 620 ordered sample (equilibrated at 850 °C). (c) Variations of Δ corr correlate with variations of peak width, Γ , from peak fitting but with increasing deviation from linearity with increasing Al/Si order. Γ 470 is 621 the average width of the two peaks near 470 cm⁻¹ in Figures 2c,d. Γ 600 is the width of the peak near 622 600 cm^{-1} in Figure 1. 623

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FIGURE 7. Comparison of spectral parameters, which relate to variations in structural state at a phonon length scale, with scalar strain, which represents structural state variations on an X-ray length scale. Data for two natural samples of anorthite (An₁₀₀, Val Pasmeda; An₉₈, Monte Somma) are from Atkinson et al. (1999) and Atkinson (1999). "nat" refers to the natural samples without heat treatment, "heat treated" refers to the natural samples after heat treatment at 1300 °C (An₁₀₀) and 1360 °C (An₉₈), as set out in Carpenter et al. (1985). (a) Δ corr values show a non-linear correlation with ε_{s} . (b) Γ for the absorption peak near 600 cm⁻¹ has a linear relationship with ε_{s} .

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FIGURE 8. Variations of Δ corr values as a function of composition across the plagioclase feldspar 633 634 solid solution. Data for natural and heated treated samples are from Atkinson et al. (1999) and Atkinson 635 (1999). Heat treatment for samples represented by open triangles involved annealing in air at 636 temperatures of ~1100-1400 °C to induce a degree of Al/Si disorder consistent with $C\overline{1}$ symmetry in 637 the composition range ~An0-An60 and $I\overline{1}$ ordering in the range ~An80-An100. The most An-rich samples had $P\overline{1}$ symmetry at room temperature. Dashed lines are guides to the eye from fitting to data 638 639 for heat treated samples in the range ~An20-An60, with extrapolation to An100. Acorr values for 640 synthetic anorthite from the present study spread essentially between the value of the extrapolated 641 straight line, $C\overline{1}$, and values for the most ordered natural anorthite samples.

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643 **FIGURE 9.** Correlation of \triangle corr data (coloured points = this study, black = from Atkinson et al. 1999) and Atkinson 1999) with ΔH soln data (Carpenter et al. 1985; Carpenter 1991b, 1992) for members of 644 645 the plagioclase feldspar solid solution. ΔH_{soln} values for ANC19 were lower than would be expected for a disordered crystalline sample due to the presence of some residual glass from the starting material 646 647 (Carpenter 1991b). Apart from these, there is a strong and almost linear correlation between $\Delta corr$ and $\Delta H_{\rm soln}$ in (a). The pattern of $\Delta H_{\rm soln}$ as a function of composition in (b) is almost exactly equivalent to 648 649 the pattern of Δ corr values in Figure 8, with a spread between values for ordered samples and the 650 extrapolated value for a solid solution with $C\overline{1}$ symmetry.

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FIGURE 10. Values of the room pressure bulk modulus, K_0 , of plagioclase feldspars determined by single-crystal high-pressure X-ray diffraction measurements at room temperature with the pressure dependence, K', set at 4. Values of Q_{od} refer to the degree of Al/Si order of anorthite samples equilibrated at different temperatures. Given that K_0 of pure anorthite decreases markedly with decreasing Al/Si order and that crystals with less than ~50% anorthite cannot change to the same

657 extent, the break in slope at \sim 50% An would become more marked with increasing equilibration 658 temperatures. 659 660 661 662 663 664 665 666 667 **TABLES** 668 669 **TABLE 1.** Synthesis and characterisation of two series of synthetic anorthite samples used in the present study. Data for structural states in terms of the presence of type-e or type-b 670 reflections in electron diffraction patterns, values of scalar strain, ε_s , and values of heat of 671 solution, ΔH_{soln} , are from Carpenter (1991b) (kinetic series) and Carpenter (1992) 672 (equilibrium series). As reported in Carpenter (1991b), sample ANC19 contained some 673 relict glass which accounted for its particularly low value of ΔH_{soln} . 674 Heat treatment e or b Scalar strain. Sample no. $\Delta H_{\rm soln}$ (kJ.mole⁻¹ $T(^{\circ}C)$ reflections *t* (h) $\mathcal{E}_{s}(\pm 1\sigma)$ $\pm 1\sigma$) Disordered reference state ANC19 1102 ± 3 0.25 0 e 37.32 ± 0.96 Kinetic series ANC60 1385 ± 5 0.017 0.0039(2) 55.23 ± 0.63 e ANC33 1392 ± 8 0.067 0.0047(2)e 57.15 ± 0.46 ANC65 1398 ± 2 0.25 e 0.0054(2) 58.53 ± 0.75 ANC73 1400 ± 3 0.78 e+b ANC38 1401 ± 1 1.5 0.0061(1)b 59.75 ± 0.67 ANC39 1400 8.25 b 0.0065(2) 62.01 ± 0.33 ANC₃₆ 1400 ± 2 48 0.0066(2) 63.56 ± 63 b ANC35 1402 ± 2 408 b 0.0069(2) 63.76 ± 1.09 Equilibrium series

ANC69 1400 ± 2 then 1536 ± 1 361, 66 b 0.0054(2) 59.96 ± 0.84

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	ANC56 ANC48 ANC62	1400 ± 1 then 1500 ± 1 1397 ± 1 1392 ± 2 then 1300 ± 1	356, 301 1102 380, 1342	b b b	0.0060(2) 0.0070(2) 0.0073(2)	61.97 ± 1.00 64.77 ± 0.88 66.27 ± 0.50	
	ANC71	1402 ± 2 then 1202 ± 2	331, 2141	b	0.0076(2)	68.03 ± 0.75	
	ANC74/H2	1393 ± 6 then 850 ± 7 , $P_{\rm H2O} = 750$ bars	0.25, 3166	b	0.0082(2)	-	
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689 Figure 2

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Figure 3 698

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Figure 5

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Figure 6





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713 Figure 8

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