

1 **Revision 2**

2 **High-pressure electrical conductivity and elasticity of iron-bearing δ -AlOOH**

3
4 Xiaowan Su¹, Jin Liu^{2,3,*}, Yukai Zhuang^{2,4}, Chaojia Lv², Xuyong Pang⁵, Fuyang Liu², Xiaohui Yu⁶,
5 Qiang Sun^{1,*}

6
7 ¹*School of Earth and Space Sciences, Peking University, Beijing 100871, China*

8 ²*Center for High Pressure Science and Technology Advanced Research (HPSTAR), Beijing 100094, China*

9 ³*CAS Center for Excellence in Deep Earth Science, Guangzhou, 510640, China*

10 ⁴*Institute of Atomic and Molecular Physics, Sichuan University, Chengdu 610065, China*

11 ⁵*Institute of Mineral Resources, Chinese Academy of Geological Sciences, Beijing 100037, China*

12 ⁶*Beijing National Laboratory for Condensed Matter Physics, Institute of Physics, Chinese Academy of
13 Sciences, Beijing 100190, China*

14 *Correspondence to: J. Liu (jin.liu@hpstar.ac.cn) and Q. Sun (qiangsun@pku.edu.cn)

15
16
17 **Research Highlights:**

18 (1) Electrical conductivity and compressibility were determined for δ -(Al,Fe)OOH with 5 and 48
19 mol.% FeOOH at pressures up to 75 GPa.

20 (2) The conductivity of δ -(Al,Fe)OOH may be slightly affected by high iron content and spin
21 transition at high pressure.

22 (3) Subducted Fe-bearing δ -AlOOH may account for some high conductivity regions in the
23 lower mantle, e.g., the North Philippine Sea slab.

25 **Abstract**

26 The electrical conductivity and elasticity of deep hydrous phases are essential to constrain water
27 distribution, as well as decipher the origins of conductivity anomalies in the lower mantle. To
28 uncover the impact of iron-bearing δ -AlOOH on geophysical properties of the lower mantle, we
29 carried out synchrotron x-ray diffraction and electrical conductivity measurements on δ -
30 $(\text{Al}_{0.52}\text{Fe}_{0.48})\text{OOH}$ and $(\text{Al}_{0.95}\text{Fe}_{0.05})\text{OOH}$ in diamond-anvil cells at pressures up to 75 GPa at
31 room temperature. A sharp volume reduction of $\sim 6.5\%$ was observed in δ - $(\text{Al}_{0.52}\text{Fe}_{0.48})\text{OOH}$
32 across the spin transition at 40.8–43.3 GPa, where its electrical conductivity increases steadily
33 without abrupt changes. The electrical conductivity of δ - $(\text{Al}_{0.52}\text{Fe}_{0.48})\text{OOH}$ is greater than that of
34 pure δ -AlOOH at high pressure, suggesting that both small polaron and proton conduction
35 mechanisms dominate in iron-bearing δ -AlOOH. Furthermore, the high-pressure electrical
36 conductivity profiles are comparable between δ - $(\text{Al}_{0.95}\text{Fe}_{0.05})\text{OOH}$ and δ - $(\text{Al}_{0.52}\text{Fe}_{0.48})\text{OOH}$,
37 indicating that high iron content only marginally influence the conductivity of iron-bearing δ -
38 AlOOH. Notably, the electrical conductivity of iron-bearing δ -AlOOH along the North
39 Philippine geotherm is greater than the average 1D electrical conductivity profile in the mantle
40 (*Ohta et al.*, 2010a). This result suggests that δ - $(\text{Al,Fe})\text{OOH}$ is a promising candidate to account
41 for high conductivity in some subducting slabs.

42

43 **Keywords:** Hydrous minerals, Spin transition, High pressure, X-ray diffraction, Electrical
44 conductivity

45

46 1. Introduction

47 Knowledge of water distribution in the mantle is fundamental to understand our planet's
48 evolution and geodynamics (Mao and Mao, 2020; Ohtani, 2020). It is widely believed that water
49 is transported to the lower mantle via slab subduction in the forms of various hydrous phases
50 such as δ -AlOOH (Duan et al., 2018; Sano-Furukawa et al., 2008), phase D (Pamato et al.,
51 2015), phase H (Liu et al., 2019b), $(\text{FeH})_{1-x}\text{Ti}_x\text{O}_2$ (Nishihara et al., 2016), pyrite-type
52 $(\text{Mg,Fe})\text{O}_2\text{H}_x$ (Hu et al., 2021), and hexagonal $(\text{Mg,Fe})_2\text{O}_{3+\delta}\text{H}_x$ (Liu et al., 2021). The δ -AlOOH
53 phase, one of the most important hydrous minerals in subducting slabs, is stable at least up to
54 134 GPa and 2300 K, that is at the high pressure and temperature (P - T) conditions of the Earth's
55 lowermost mantle (Duan et al., 2018; Sano-Furukawa et al., 2008). Thus, it has been considered
56 as an important deep-water carrier throughout the lower mantle (Ohtani et al., 2001; Sano-
57 Furukawa et al., 2008; Tsuchiya and Tsuchiya, 2011). Additionally, δ -AlOOH forms a solid
58 solution with the isostructural hydrous phase ϵ -FeOOH in multicomponent systems (Kawazoe et
59 al., 2017; Ohira et al., 2019; Liu et al., 2019b; Buchen et al., 2021).

60 The incorporation of FeOOH greatly influences the physics and chemistry of δ -AlOOH.
61 Because of pressure effects on the electronic spin states of $3d$ transition metal iron, the δ -
62 $(\text{Al,Fe})\text{OOH}$ phase undergoes a high-spin to low-spin (HS-LS) transition of iron at ~ 40 GPa
63 (Ohira et al., 2019; Hsieh et al., 2020; Su et al., 2021a). Intriguingly, the unit cell volume,
64 compressibility, and thermal conductivity of δ - $(\text{Al,Fe})\text{OOH}$ exhibit abnormal behaviors
65 throughout the electronic spin-pairing transition. For instance, δ - $(\text{Al}_{0.908}\text{Fe}_{0.045})\text{OOH}_{1.14}$ and δ -
66 $(\text{Al}_{0.832}\text{Fe}_{0.117})\text{OOH}_{1.15}$ both display gradual volume collapses across the HS-LS transition (Ohira
67 et al., 2019). The thermal conductivity of δ - $(\text{Al}_{0.85}\text{Fe}_{0.15})\text{OOH}$ drastically decreases at 30–45 GPa
68 and approaches an exceptionally low value of $\sim 10 \text{ W m}^{-1} \text{ K}^{-1}$ in the LS state at 66 GPa (Hsieh et
69 al., 2020). Moreover, the incorporation of 5 mol.% FeOOH decreases the shear-wave velocity of
70 δ -AlOOH by $\sim 5\%$ at 20–135 GPa (Su et al., 2021b). However, the effect of iron on the electrical
71 conductivity of δ -AlOOH has not yet been reported under high pressures.

72 The electrical conductivity values of iron-free δ -AlOOH remain almost the same up to 50

73 GPa, which indicates proton conduction is the primary conduction mechanism (*Zhuang et al.*,
74 2021). ϵ -FeOOH exhibits much higher conductivity than δ -AlOOH up to 20 GPa at high
75 temperatures (*Wang and Yoshino*, 2021). We note that the electrical conductivity of mantle
76 minerals would be affected by the spin transition of iron at high pressure. In particular,
77 bridgmanite and ferropericlase are subject to conductivity reduction across iron spin transition
78 (*Lin et al.*, 2007; *Ohta et al.*, 2010b). It has been ascribed to the reduced unpaired electrons,
79 which are the charge transfer carriers in the small polaron conduction mechanism. Thus, the
80 incorporation of FeOOH and iron spin transition are expected to influence the electrical behavior
81 of δ -AlOOH with increasing pressure. Meanwhile, it still remains unclear whether the presence
82 of this hydrous phase could contribute to conductivity anomalies observed in the deep mantle,
83 which, in turn, can allow a better understanding of the Earth's deep-water cycle. Therefore, high-
84 pressure electrical conductivity measurements are highly demanded for iron-bearing δ -AlOOH.

85 In this work, δ -(Al_{0.95}Fe_{0.05})OOH and δ -(Al_{0.52}Fe_{0.48})OOH single-crystal samples were
86 synthesized at 26 GPa and 1473 K for 4 hours using a 1000-ton Kawai-type multi anvil
87 apparatus. Electrical conductivity measurements were conducted on these two samples up to 66.3
88 and 68.3 GPa at room temperature, respectively, in diamond-anvil cells (DAC). In addition, the
89 pressure-volume relation of δ -(Al_{0.52}Fe_{0.48})OOH was investigated at pressures to 75.4 GPa using
90 synchrotron x-ray diffraction (XRD). These results are used to decipher how iron content and
91 spin transition affect the electrical conductivity of iron-bearing δ -AlOOH at high pressure. Our
92 study gives insights into the electrical structure of the deep mantle, e.g., beneath northeast China
93 and the North Philippine Sea.

94

95 **2. Experimental Methods**

96 **2.1 Sample synthesis and characterization**

97 High P - T synthesis experiments were performed in a 1000-ton Kawai-type multi anvil
98 apparatus at the Institute of Physics, Chinese Academy of Sciences, Beijing. Details of the
99 synthesis procedure were similar to those described by *Kawazoe et al.* (2017) and *Ohira et al.*

100 (2019). Fe_2O_3 and $\text{Al}(\text{OH})_3$ powders were mixed as starting materials with chemical
101 compositions of $(\text{Al}_{1-x}\text{Fe}_x)\text{OOH}$ and excess water, where $x = 0.50$ and 0.05 , respectively. A
102 welded Au capsule was filled with powder mixtures. The capsule was placed in the MgO
103 cylinder, and the rhenium heater was enclosed by a ZrO_2 thermal insulator. A Cr_2O_3 -doped MgO
104 octahedron with a 10 mm edge length was used as the pressure medium. Tungsten carbide
105 second-stage anvils were used with a 4 mm edge length. The temperature was measured using a
106 $\text{W}_{97\%}\text{Re}_{3\%}$ - $\text{W}_{75\%}\text{Re}_{25\%}$ thermocouple (type D). The assembly was first compressed to 26 GPa and
107 then heated to 1473 K. This target temperature was maintained for 4 hours for grain growth. The
108 sample was rapidly quenched at 26 GPa by switching off the electrical power and then slowly
109 decompressed for 15 hours.

110 The chemical compositions of the recovered samples were measured using electron probe
111 microanalyzer (EPMA) at the Institute of Mineral Resources, Chinese Academy of Geological
112 Sciences, resulting in $(\text{Al}_{0.52}\text{Fe}_{0.48})\text{OOH}$. The recovered samples with the lower iron content were
113 examined using SEM-EDS (scanning electron microscope equipped with energy-dispersive
114 spectrometer) at Peking University, resulting in $(\text{Al}_{0.95}\text{Fe}_{0.05})\text{OOH}$. SEM and EPMA analyses
115 showed that the synthesized samples are chemically homogeneous within analytical
116 uncertainties. XRD patterns of the recovered samples were collected using Bruker APEX III D8
117 venture diffractometer at the Center for High Pressure Science and Technology Advanced
118 Research (HPSTAR). The lattice parameters of $(\text{Al}_{0.52}\text{Fe}_{0.48})\text{OOH}$ (space group: $P2_1nm$) are $a =$
119 $2.9078(3)$ Å, $b = 4.3274(5)$ Å, and $c = 4.8223(6)$ Å. In addition, $(\text{Al}_{0.95}\text{Fe}_{0.05})\text{OOH}$ (denoted as
120 “Delta95”) has an orthorhombic structure (space group: $P2_1nm$), yielding the unit cell parameters
121 of $a = 2.8465(2)$ Å, $b = 4.2448(2)$ Å, and $c = 4.7350(3)$ Å. These values are consistent with the
122 previous studies (*Su et al.*, 2021b; *Buchen et al.*, 2021; *Ohira et al.*, 2019; *Sano-Furukawa et al.*,
123 2009; *Suzuki et al.*, 2010). Figure 1 shows the unit-cell parameters of δ - $(\text{Al}_x\text{Fe}_{1-x})\text{OOH}$ samples
124 increasing with $\text{Fe}/(\text{Al}+\text{Fe})$ at ambient conditions.

125

126 2.2 Synchrotron x-ray diffraction experiments

127 δ -(Al_{0.52}Fe_{0.48})OOH (denoted as “Delta52”) samples were double-side polished to ~ 40 μ m
128 in diameter and 7–9 μ m thick. The prepared platelet was then loaded into a sample chamber
129 of 100 μ m in diameter and ~ 25 μ m thick, drilled into the center of the pre-indented tungsten
130 gasket in a DAC. Au and ruby were placed next to the sample to serve as pressure calibrants (*Fei*
131 *et al.*, 2007; *Mao et al.*, 1986). Neon was loaded into the sample chamber as the pressure-
132 transmitting medium using the high-pressure gas loading system at HPSTAR. *In situ* synchrotron
133 XRD experiments were conducted at the BL10XU beamline of the SPring-8 synchrotron
134 radiation facility (*Hirao et al.*, 2020). An x-ray beam with $\lambda = 0.4133$ Å was used in combination
135 with an image plate (IP) detector (Rigaku RAXIS-IV). The six diffraction peaks (110, 101, 200,
136 111, 210, 211) were observed in most of the diffraction patterns of the Delta52, whereas other
137 peaks were relatively weak. Therefore, these six diffraction peaks of the Delta52 were used to
138 calculate the lattice parameters via the Unit Cell and Dioptas programs (*Prescher and*
139 *Prakapenka*, 2015). Based on XRD patterns of the loaded sample at ~1.0 GPa, the single-crystal
140 Delta52 sample was crushed to a large-grained polycrystalline sample in closing the diamond
141 anvil cell.

142

143 **2.3 Electrical conductivity measurements**

144 The values of electrical conductivity of the Delta52 and Delta95 samples were determined
145 up to 66.3 and 68.3 GPa, respectively, at room temperature by using symmetrical DACs with a
146 culet size of 250 μ m. Delta52 and Delta95 samples were prepared into sample disks of 40 μ m in
147 diameter and ~ 8 μ m thick. Re gaskets were pre-indented to 35–40 μ m (c.a. 20 GPa) thick
148 between a pair of diamond anvils, and a hole of 180 μ m in diameter was drilled into the center of
149 the pre-indentation. A powder mixture of cubic boron nitride (cBN) and epoxy was packed into the
150 hole and further compressed up to 20–25 GPa. A 100 μ m diameter hole was then drilled into the
151 center of the cBN-epoxy gasket insert. The sample disk was further placed into the hole, and four
152 platinum electrodes were connected to the sample. The electrical resistance of the sample was
153 measured by using Solartron-1260 AC impedance spectroscopy with frequencies between 0.1 Hz

154 and 10 MHz using the four-terminal method (Zhuang *et al.*, 2021). The procedure of resistance
155 data analysis can be found in detail in *van der Pauw* (1958).

156

157 **3. Results and Discussion**

158 **3.1 Compressibility and spin transition of iron-bearing δ -AlOOH at high pressure**

159 XRD patterns of the Delta52 were collected up to 75.4 GPa (Fig. 2). The lattice parameters
160 (a , b , and c) and unit cell volume were determined for the Delta52 (Fig. 3 and Table 1). The
161 pressure-volume (P - V) compression data of the Delta52 were divided into three regions based on
162 our x-ray emission scattering (XES) and laser Raman spectroscopic results on analogous material
163 (*Su et al.*, 2021a): 0–10.4 GPa for the $P2_1nm$ structure (HS), 10.4–40.8 GPa for the $Pnmm$
164 structure (HS), and 43.3–75.4 GPa for the $Pnmm$ structure (LS). The third-order Birch-
165 Murnaghan (BM) EoS was fitted to the P - V data of the Delta52 using EosFit7Gui software
166 (*Gonzalez-Platas et al.*, 2016). The unit cell volume at ambient pressure conditions (V_0),
167 isothermal bulk modulus (K_{T0}), and pressure derivative of the bulk modulus (K_{T0}') were derived,
168 yielding: $V_0 = 61.6(1) \text{ \AA}^3$, $K_{T0} = 134(4) \text{ GPa}$, and K_{T0}' fixed at 4 for the $P2_1nm$ structure (HS); V_0
169 $= 60.4(1) \text{ \AA}^3$, $K_{T0} = 216(2) \text{ GPa}$, and $K_{T0}' = 3.98(9)$ for the $Pnmm$ structure (HS); and $V_0 = 56.5(1)$
170 \AA^3 , $K_{T0} = 234(2) \text{ GPa}$, and $K_{T0}' = 4.00(2)$ for the $Pnmm$ structure (LS) (Table 2).

171 The P - V profile of the Delta52 shows a sharp volume reduction of $\sim 6.5\%$ in a narrow
172 pressure range between 40.8 and 43.3 GPa at 300 K in Ne pressure medium (Fig. 3 and Table 3).
173 Such a volume collapse is related to the pressure-induced spin transition in the Delta52, whereas
174 a gradual volume collapse of $\sim 10.3\%$ was observed in ϵ -FeOOH at $\sim 45(2) \text{ GPa}$ (*Thompson et*
175 *al.*, 2017). The volume reductions are about 1.3% for δ -(Al_{0.95}Fe_{0.05})OOH and δ -
176 (Al_{0.908}Fe_{0.045})OOH_{1.14} and 2.8% for δ -(Al_{0.832}Fe_{0.117})OOH_{1.15} across the spin transition (*Ohira et*
177 *al.*, 2019; *Su et al.*, 2021b). In other words, the volume reduction of the δ -(Al, Fe)OOH increases
178 with increasing FeOOH content throughout the spin transition of iron (Table 3). On the other
179 hand, a thermal conductivity anomaly of δ -(Al_{0.97}Fe_{0.03})OOH was observed at 30–50 GPa and
180 room temperature using silicon oil as a pressure-transmitting medium by *Hsieh et al.* (2020),

181 which has also been related to the spin crossover. Similarly, δ -(Al_{0.88}Fe_{0.12})OOH and δ -
182 (Al_{0.85}Fe_{0.15})OOH with higher iron contents exhibit anomalous changes in thermal conductivity
183 approximately at 30–53 GPa (Hsieh *et al.*, 2020). Notably, the use of neon as a pressure-
184 transmitting medium in this study can greatly decrease the width of the spin crossover of iron in
185 δ -(Al,Fe)OOH, while the use of silicone oil generates a larger deviatoric stress that could
186 broaden the width of iron spin transition. Overall, the onset pressure of the spin transition
187 appears to increase gradually with the increasing iron content from 3 to 48 mol.%, approximately
188 from 30 to 41 GPa (Su *et al.*, 2021a).

189 Figure 4 presents the volume difference between the Delta52 and δ -AlOOH as a function of
190 pressure at room temperature (Sano-Furukawa *et al.*, 2009). It decreases dramatically at \sim 40
191 GPa and approaches 1 Å³ at 65 GPa. The unit cell volume of the Delta52 is close to that of δ -
192 AlOOH above 40 GPa, illustrating that the effective ionic radius of ferric iron Fe³⁺ in the LS
193 state is almost the same as that of Al³⁺. In addition, at approximately 10 GPa, the *a/c* and *b/c*
194 ratio values of the Delta52 reach their minima, whereas the *a/b* ratio values reach their maximum
195 (Fig. 5). Such trends were also observed for the axial ratios of δ -(Al_{0.832}Fe_{0.117})OOH_{1.15}, δ -
196 (Al_{0.908}Fe_{0.047})OOH_{1.14}, and δ -AlOOH at \sim 10 GPa (Ohira *et al.*, 2019; Su *et al.*, 2021b; Sano-
197 Furukawa *et al.*, 2009; Duan *et al.*, 2018). Similarly, the reversal of the axial ratios' pressure
198 dependences was observed for δ -(Al_{0.95}Fe_{0.05})OOH at around 14 GPa (Su *et al.*, 2021b). This
199 behavior has been related to the hydrogen bond order-disorder phase transition from *P2₁nm* to
200 *Pnnm* (Sano-Furukawa *et al.*, 2018; Su *et al.*, 2021b; Ohira *et al.*, 2019). In contrast, the pressure
201 dependences of ϵ -FeOOH axial ratios are different due to the absence of the hydrogen bond
202 order-disorder phase transition. The *K*₇₀ value of the Delta52 increases by about 52.7% at 10.4–
203 13.3 GPa and 7.6% at 40.8–43.3 GPa, respectively (Fig. 6a). These results show that both the
204 hydrogen bond order-disorder and HS-LS phase transitions influence the compressibility of iron-
205 bearing δ -AlOOH.

206 The bulk sound velocity of the Delta52 as a function of pressure is shown in Figure 6b. The
207 velocity of the Delta52 increases monotonically from 15 to 75 GPa. Compared to δ -AlOOH, the

208 substitution of 48 mol.% FeOOH decreases the velocity by $\sim 7.3\%$. Similarly, the bulk sound
209 velocity of the $\delta\text{-(Al}_{0.87}\text{Fe}_{0.13}\text{)OOH}$ is 4.0% lower than that of $\delta\text{-AlOOH}$ at 65 GPa (*Ohira et al.*,
210 2021). Additionally, the bulk sound velocity of the Delta95 is less than that of $\delta\text{-AlOOH}$ by $\sim 2.9\%$
211 at 20–135 GPa (*Su et al.*, 2021b). Moreover, $\epsilon\text{-FeOOH}$ decreases the velocity by $\sim 16.3\%$ with
212 respect to $\delta\text{-AlOOH}$. Thus, the magnitude of the velocity reduction in $\delta\text{-(Al,Fe)OOH}$ exhibits a
213 positive correlation with iron content. On the other hand, the bulk sound velocity of $\delta\text{-}$
214 (Al,Fe)OOH does not undergo a dramatic change across the spin crossover, although the spin
215 transition increases the bulk modulus (Fig. 6b). The reason for this is that it is counterbalanced
216 by the increased density of $\delta\text{-(Al,Fe)OOH}$ from the HS to LS states.

217

218 **3.2 Electrical conductivity and conduction mechanisms of $\delta\text{-(Al,Fe)OOH}$ at high pressure**

219 The electrical conductivity of the Delta52 and Delta95 is shown as a function of pressure in
220 Figure 7a and Table 4. The electrical conductivity profile of the Delta95 is present with a positive
221 spike of $\sim 10^{-1.50}$ S/m at around 6 GPa (Fig. 7a). This abrupt change may be related to the
222 hydrogen bond order-disorder transition, which would weaken the O–H bond and strengthen the
223 $\text{H}\cdots\text{O}$ bond (*Kuribayashi et al.*, 2014; *Cortona*, 2017; *Ohira et al.*, 2019; *Wang and Yoshino*,
224 2021). Moreover, the electrical conductivity of the Delta95 increases with increasing pressure at
225 10–66 GPa (Fig. 7a). We note that $\delta\text{-(Al}_{0.85}\text{Fe}_{0.15}\text{)OOH}$ (denoted as “Delta85”) shows two new
226 Raman bands at 6.0 GPa, in agreement with the transition from an ordered ($P2_1nm$) to a
227 disordered hydrogen bonding structure ($Pnmm$) for $\delta\text{-AlOOH}$ and the Delta95 (*Su et al.*, 2021a).
228 In this previous study, the Delta52 exhibits two new Raman bands between 8.5 and 15.8 GPa,
229 suggesting that the incorporation of 48 mol.% FeOOH would increase the order-disorder
230 transition pressure. Thus, like the Delta95, it is expected to observe a dramatic change in the
231 electrical conductivity of the Delta52 at 8.5–15.8 GPa. However, it is not present in the present
232 measurements due to limited data points that were collected below 20 GPa. The conductivity of
233 the Delta52 increases with pressure from $10^{-1.56}$ S/m at 23 GPa to $10^{-0.69}$ S/m at 68 GPa (Fig. 7a).
234 The electrical conductivity of iron-free $\delta\text{-AlOOH}$ is around $10^{-2.25}$ S/m and remains nearly

235 unchanged with increasing pressure up to ~ 50 GPa (*Zhuang et al.*, 2021). That is, the
236 incorporation of FeOOH significantly alters the electrical conductivity evolution of δ -AlOOH at
237 pressures greater than 20 GPa. Moreover, the electrical conductivity values of the Delta52 and
238 Delta95 monotonically increase without kinks throughout the spin crossover (Fig. 7a). This
239 suggests that the iron spin transition may have negligible effects on the electrical conductivity of
240 iron-bearing δ -AlOOH at high pressure. In general, iron spin transition greatly reduces unpaired
241 electrons and thus the charge carrier density in the small polaron conduction mechanism;
242 accordingly, mantle bridgmanite and ferropericlase undergo a reduction in electrical conductivity
243 from the HS to LS states (*Lin et al.*, 2007; *Ohta et al.*, 2010b). This reduction, however, does not
244 occur in δ -(Al,Fe)OOH, indicating that the spin transition-induced decrease in charge carrier
245 density may not alter the predominant electrical conductivity mechanism of δ -(Al,Fe)OOH at
246 high pressure.

247 Proton conduction is well known to be present in hydrous minerals owing to high proton
248 mobility (*Guo*, 2016). The number of free protons in hydrous phases and proton migration rate
249 through lattice sites are two significant factors influencing their conductive behavior. The O_1 -
250 $H\cdots O_2$ bonds of δ -AlOOH become symmetric above 20 GPa, hindering protons moving freely
251 (*Wang and Yoshino*, 2021). Therefore, the electrical conductivity values of δ -AlOOH keep
252 almost constant at 20–50 GPa (*Zhuang et al.*, 2021). Notably, the electrical conductivity values
253 of δ -(Al,Fe)OOH become greater than that of pure δ -AlOOH at pressures greater than 20 GPa. It
254 is mostly related to the increment of free protons in iron-bearing δ -AlOOH through the activation
255 reaction of $Fe^{3+} = Fe^{2+} + H^+$. The δ -(Al,Fe)OOH may thus contain a number of Fe^{2+} through this
256 substitution mechanism. Small polaron hopping conduction occurs through charge transfer
257 between neighboring ions of different valences like Fe^{2+} and Fe^{3+} or electron-electron hole pair
258 recombination. As *Wang and Yoshino* (2021) pointed out, the dynamical coexistence of Fe^{2+} and
259 Fe^{3+} in δ -(Al,Fe)OOH makes further contributions to the conductivity. Therefore, the primary
260 conduction mechanism of δ -(Al,Fe)OOH at 20–70 GPa is proton conduction coupled with small
261 polaron hopping conduction. It is the same as the conduction mechanisms of iron-bearing

262 antigorite, an important hydrous mineral in subduction zones (*Guo et al.*, 2011).

263 The electrical behavior of δ -AlOOH is greatly altered by the incorporation of FeOOH at
264 high pressure. At 20–70 GPa, the electrical conductivity values of iron-bearing δ -AlOOH phases
265 are greater than that of pure δ -AlOOH (Fig. 7a). Specifically, the electrical conductivity of δ -
266 AlOOH at 50 GPa is enhanced by about 25 times through the incorporation of 48 mol.%
267 FeOOH, corresponding to an increase of ~ 1.4 in decimal logarithm unit. Moreover, the linear
268 fitting between 20–70 GPa yields a pressure dependence of the decimal logarithm of electrical
269 conductivity 0.02–0.03 log[S/m] per gigapascal for the Delta52 and Delta95 at room temperature
270 (Fig. 7b). We note that the Delta52 and Delta95 have comparable electrical conductivity values at
271 pressures greater than 20 GPa. That is, the electrical conductivity behavior of iron-bearing δ -
272 AlOOH may be slightly affected by FeOOH content between 5 and 48 mol.%. Similarly, it has
273 been observed on the thermal conductivity of δ -(Al,Fe)OOH with 3–15 mol.% FeOOH above
274 40–50 GPa (*Hsieh et al.*, 2020).

275

276 **4. Implications**

277 Aluminous hydrous mineral δ -AlOOH may be stable up to the deep lower mantle
278 conditions based on the recent XRD experiments at high P - T conditions (*Duan et al.*, 2018).
279 Diaspore (α -AlOOH), a naturally occurring hydrous mineral in sediments, could transform into
280 its high-pressure phase (δ -AlOOH) when being transported to deep Earth through subducting
281 slabs (*Suzuki et al.*, 2000). We note that δ -AlOOH could also emerge out of the decomposition of
282 Al-rich phase D and phase Egg (AlSiO₃OH) at the uppermost lower mantle conditions (*Sano et*
283 *al.*, 2004; *Xu and Inoue*, 2019). In addition, δ -AlOOH likely coexists with bridgmanite and
284 remains stable down to the bottom of the lower mantle (*Ohira et al.*, 2014). More importantly, δ -
285 AlOOH is expected to accommodate some FeOOH in the deep mantle. *Wang and Yoshino* (2021)
286 suggested that δ -(Al,Fe)OOH could be entrained into the deep mantle through the North
287 Philippine Sea slabs. *Liu et al.* (2019b) experimentally synthesized a series of Al-rich hydrous
288 phases up to 15–17 vol% out of subducted oceanic crusts in the mantle, and found that

289 (Al,Fe)OOH components are predominant according to their large volume press experiments at
290 25–26 GPa and high temperatures. Therefore, the δ -(Al,Fe)OOH phase is likely stable in cold
291 subducting slabs when entering the lower mantle.

292 Furthermore, the electrical conductivity of the δ -(Al,Fe)OOH was extrapolated to high P - T
293 conditions, allowing us to assess how the δ -(Al,Fe)OOH phase contributes to the electrical
294 structure of the lower mantle (Figs 7b and 8). *Wang and Yoshino (2021)* reported the activation
295 enthalpy (ΔH) of δ -(Al,Fe)OOH with 0.82 wt% Fe₂O₃ is 0.38(1) eV at 500–1200 K and 20 GPa.
296 Assuming that the activation enthalpy of δ -(Al,Fe)OOH would not change significantly with
297 increasing pressure, the pressure dependence on electrical conductivity of δ -(Al,Fe)OOH could
298 not change with increasing temperature. Thus, the pressure dependence of the decimal logarithm
299 of electrical conductivity of δ -(Al,Fe)OOH could be extrapolated from at 300 K to higher
300 temperatures in this study. The electrical conductivity of the δ -(Al,Fe)OOH along the cold slab
301 geotherm (e.g., the North Philippine Sea) was calculated using Arrhenius formula (Fig. 8). The
302 extrapolated EC values of δ -(Al,Fe)OOH at depths of \sim 600 km could well match the
303 conductivity-depth profile of δ -(Al,Fe)OOH by *Wang and Yoshino (2021)*. Notably, the electrical
304 conductivity values of δ -(Al,Fe)OOH are greater than the average 1D electrical conductivity
305 profile in the range of 600–1500 km (*Ohta et al., 2010b*). It is conceivable that δ -(Al,Fe)OOH-
306 bearing domains in cold slabs may account for some high conductivity regions in the lower
307 mantle. Meanwhile, the incorporation of some FeOOH would significantly reduce the sound
308 velocity of δ -AlOOH, iron-bearing δ -AlOOH would have sound velocities much lower than
309 major lower-mantle minerals (*Su et al., 2021b; Fu et al., 2018*). Therefore, iron-bearing δ -
310 AlOOH is a promising candidate contributing to lower-mantle mantle seismic heterogeneities,
311 such as large low-shear-velocity provinces (LLSVPs) and ultralow velocity zones (ULVZs) (*Liu*
312 *et al., 2017; Mashino et al., 2016; Thompson et al., 2017*). More interestingly, δ -(Al,Fe)OOH
313 could account for the high conductivity layer and seismic anomalies at similar depths, including
314 Northeast China, East Africa, and the Philippine Sea (*Tarits et al., 2010; Tada et al., 2014*). In
315 turn, these results may serve as a means of locating deep-water reservoirs in the mantle.

316 Acknowledgments

317 This study is funded by the National Key Research and Development Program of China
318 (2019YFA0708502). Y. Zhuang is supported by the China Postdoctoral Science Foundation
319 (18NZ021-0213-216308). J. Liu acknowledges support from the National Natural Science
320 Foundation of China (Grant no. 42072052 and U1930401). We thank Bingmin Yan, Hu Tang,
321 Jiu Hua Chen, and Shuailing Ma for providing their technical support for the multi-anvil press
322 experiments, as well as Saori Kawaguchi and Naohisa Hirao for their assistance with synchrotron
323 XRD experiments. *In situ* XRD measurements were performed at the BL10XU beamline of
324 SPring-8 (Proposal 2019A1284 and 2019B1203). SPring-8 is supported by the National Science
325 Foundation Earth Sciences (Grant No. EAR1128799) and the Department of Energy-
326 GeoSciences (Grant No. DE-FG02-94ER14466). Some experiments are supported by the
327 Synergic Extreme Condition User Facility (SECUF).

328

329 References

- 330 Buchen, J., Sturhahn, W., Ishii, T., and Jackson, J.M. (2021), Vibrational anisotropy of δ -(Al,Fe)OOH
331 single crystals as probed by nuclear resonant inelastic X-ray scattering, *European Journal of*
332 *Mineralogy*, 33(4), 485-502.
- 333 Cortona, P. (2017), Hydrogen bond symmetrization and elastic constants under pressure of δ -AlOOH,
334 *Journal of Physics: Condensed Matter*, 29(32), 325505.
- 335 Duan, Y., Sun, N., Wang, S., Li, X., Guo, X., Ni, H., et al. (2018), Phase stability and thermal equation of
336 state of δ -AlOOH: Implication for water transportation to the deep lower mantle, *Earth Planetary*
337 *Science Letters*, 494, 92-98.
- 338 Fei, Y., Ricolleau, A., Frank, M., Mibe, K., Shen, G., Prakapenka, V. Toward an Internally Consistent
339 Pressure Scale, *Proceeding of the National Academy of Sciences*, 104, 9182-9186.
- 340 Fu, S., Yang, J., Zhang, Y., Okuchi, T., McCammon, C., Kim, H. I., et al. (2018), Abnormal Elasticity of
341 Fe-bearing Bridgmanite in the Earth's Lower Mantle, *Geophysical Research Letters*, 45(10), 4725-
342 4732.
- 343 Gleason, A. E., Quiroga, C. E., Suzuki, A., Pentcheva, R., Mao, W. L. (2013), Symmetrization driven spin
344 transition in ϵ -FeOOH at high pressure, *Earth and Planetary Science Letters*, 379, 49-55.
- 345 Gonzalez-Platas, J., Alvaro, M., Nestola, F., Angel, R. (2016), EosFit7-GUI: a new graphical user
346 interface for equation of state calculations, analyses and teaching, *Journal of Applied*
347 *Crystallography*, 49(4), 1377-1382.

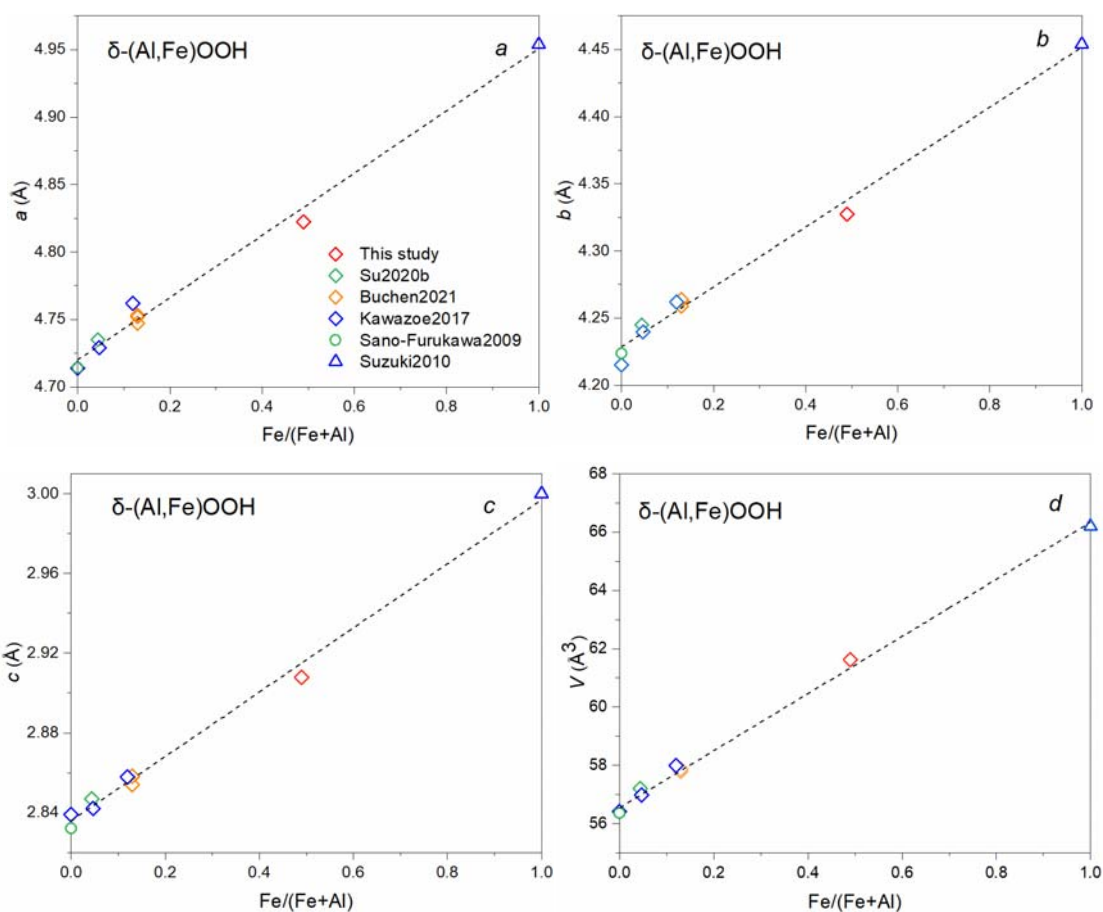
- 348 Guo, X. (2016), Experimental study of the electrical conductivity of hydrous minerals in the crust and the
349 mantle under high pressure and high temperature, *Science China Earth Sciences*, 59(4), 696-706.
- 350 Guo, X., Yoshino, T., Katayama, I. (2011), Electrical conductivity anisotropy of deformed talc rocks and
351 serpentinites at 3 GPa, *Physics of the Earth Planetary Interiors*, 188(1-2), 69-81.
- 352 Hirao, N., Kawaguchi, S. I., Hirose, K., Shimizu, K., Ohtani, E., Ohishi, Y. (2020), New developments in
353 high-pressure X-ray diffraction beamline for diamond anvil cell at SPring-8, *Matter and radiation at
354 extremes*, 5(1), 018403.
- 355 Hsieh, W. P., Ishii, T., Chao, K. H., Tsuchiya, J., Deschamps, F., Ohtani, E. (2020), Spin transition of iron
356 in δ -(Al, Fe)OOH induces thermal anomalies in Earth's lower mantle, *Geophysical Research Letters*,
357 47(4), 087036.
- 358 Hu, Q., Liu, J., Chen, J., Yan, B., Meng, Y., Prakapenka, V. B., et al. (2021), Mineralogy of the deep lower
359 mantle in the presence of H₂O, *National Science Review*, 8(4), nwaa098.
- 360 Kawazoe, T., Ohira, I., Ishii, T., Ballaran, T. B., McCammon, C., Suzuki, A., et al. (2017), Single crystal
361 synthesis of δ -(Al,Fe)OOH, *American Mineralogist*, 102(9), 1953-1956.
- 362 Kuribayashi, T., Sano-Furukawa, A., Nagase, T. (2014), Observation of pressure-induced phase transition
363 of δ -AlOOH by using single-crystal synchrotron X-ray diffraction method, *Physics Chemistry of
364 Minerals*, 41(4), 303-312.
- 365 Lin, J.-F., S. T. Weir, D. D. Jackson, W. J. Evans, Y. K. Vohra, W. Qiu, and C.-S. Yoo (2007), Electrical
366 conductivity of the lower-mantle ferropericlae across the electronic spin transition, *Geophysical
367 Research Letters*, 34(16), L16305.
- 368 Liu, J., Lin, J. F., Mao, Z., Prakapenka, V. B. (2014), Thermal equation of state and spin transition of
369 magnesiosiderite at high pressure and temperature, *American mineralogist*, 99(1), 84-93.
- 370 Liu, J., Wang, C., Lv, C., Su, X., Liu, Y., Tang, R., et al. (2021), Evidence for oxygenation of Fe-Mg
371 oxides at mid-mantle conditions and the rise of deep oxygen, *National Science Review*, 8(4),
372 nwaa096.
- 373 Liu, J., Hu, Q., Kim, D. Y., Wu, Z., Wang, W., Yuming, et al. (2017), Hydrogen-bearing iron peroxide and
374 the origin of ultralow-velocity zones, *Nature*, 551(7681), 494-497.
- 375 Liu, J., Hu, Q., Bi, W., Yang, L., Xiao, Y., Chow, P., et al. (2019a), Altered chemistry of oxygen and iron
376 under deep Earth conditions, *Nature Communications*, 10(1), 153-153.
- 377 Liu, X., Matsukage, K. N., Nishihara, Y., Suzuki, T., Takahashi, E. (2019b). Stability of the hydrous
378 phases of Al-rich phase D and Al-rich phase H in deep subducted oceanic crust. *American
379 Mineralogist*, 104(1), 64-72.
- 380 Mao, H.-k., Xu, J.-A., and Bell, P. (1986), Calibration of the ruby pressure gauge to 800 kbar under quasi-
381 hydrostatic conditions. *Journal of Geophysical Research*, 91, 4673-4676.
- 382 Mao, H.-k., and Mao, W. L. (2020), Key problems of the four-dimensional Earth system. *Matter and
383 Radiation at Extremes*, 5(3), 038102.
- 384 Mashino, I., Murakami, M., Ohtani, E. (2016), Sound velocities of δ -AlOOH up to core-mantle boundary
385 pressures with implications for the seismic anomalies in the deep mantle, *Journal of Geophysical
386 Research: Solid Earth*, 121(2), 595-609.
- 387 Nishihara, Y., Matsukage, K.N. (2016), Iron-titanium oxyhydroxides as water carriers in the Earth's deep

- 388 mantle, *American Mineralogist*, 101(4), 919-927.
- 389 Ohira, I., Ohtani, E., Sakai, T., Miyahara, M., Hirao, N., Ohishi, Y., and Nishijima, M. (2014), Stability of
390 a hydrous δ -phase, $\text{AlOOH-MgSiO}_2(\text{OH})_2$, and a mechanism for water transport into the base of
391 lower mantle, *Earth and Planetary Science Letters*, 401, 12-17.
- 392 Ohira, I., Jackson, J. M., Solomatova, N. V., Sturhahn, W., Finkelstein, G. J., Kamada, S., et al. (2019),
393 Compressional behavior and spin state of δ -(Al, Fe) OOH at high pressures, *American Mineralogist*,
394 104(9), 1273-1284.
- 395 Ohira, I., Jackson, J. M., Sturhahn, W., Finkelstein, G. J., Kawazoe, T., Toellner, T. S., et al. (2021), The
396 influence of δ -(Al,Fe)OOH on seismic heterogeneities in Earth's lower mantle, *Scientific Reports*, 11,
397 12036.
- 398 Ohta, K., Hirose, K., Ichiki, M., Shimizu, K., Sata, N., Ohishi, Y. (2010a), Electrical conductivities of
399 pyrolitic mantle and MORB materials up to the lowermost mantle conditions, *Earth and Planetary*
400 *Science Letters*, 289(3), 497-502.
- 401 Ohta, K., Hirose, K., Katsuya, S., Nagayoshi, S., Yasuo, O. (2010b), The electrical resistance
402 measurements of (Mg,Fe)SiO₃ perovskite at high pressures and implications for electronic spin
403 transition of iron, *Physics of the Earth Planetary Interiors*, 180(3-4), 154-158.
- 404 Ohtani, E. (2020), The role of water in Earth's mantle, *National Science Review*, 7(1), 224-232.
- 405 Ohtani, E., Litasov, K., Suzuki, A., Kondo, T. (2001), Stability field of new hydrous phase, δ -AlOOH,
406 with implications for water transport into the deep mantle, *Geophysical Research Letters*, 28(20),
407 3991-3993.
- 408 Pamato, M. G., Myhill, R., Boffa Ballaran, T., Frost, D. J., Heidelbach, F., Miyajima, N. (2015), Lower-
409 mantle water reservoir implied by the extreme stability of a hydrous aluminosilicate, *Nature*
410 *Geoscience*, 8(1), 75-79.
- 411 Prescher, C., Prakapenka, V. B. (2015), DIOPTAS: a program for reduction of two-dimensional X-ray
412 diffraction data and data exploration, *High Pressure Research*, 35(3), 223-230.
- 413 Sano-Furukawa, A., Komatsu, K., Vanpeteghem, C. B., Ohtani, E. (2008), Neutron diffraction study of δ -
414 AlOOD at high pressure and its implication for symmetrization of the hydrogen bond, *American*
415 *Mineralogist*, 93(10), 1558-1567.
- 416 Sano-Furukawa, A., Kagi, H., Nagai, T., Nakano, S., Fukura, S., Ushijima, D., et al. (2009), Change in
417 compressibility of δ -AlOOH and δ -AlOOD at high pressure: A study of isotope effect and hydrogen-
418 bond symmetrization, *American Mineralogist*, 94(8), 1255-1261.
- 419 Sano, A., Ohtani, E., Kubo, T., Funakoshi, K.-i. (2004), In situ X-ray observation of decomposition of
420 hydrous aluminum silicate AlSiO_3OH and aluminum oxide hydroxide δ -AlOOH at high pressure
421 and temperature, *Journal of Physics and Chemistry of Solids*, 65(8), 1547-1554.
- 422 Satta, N., Criniti, G., Kurnosov, A., Ballaran, T.B., Ishii, T., and Marquardt, H. (2021), High-Pressure
423 Elasticity of δ -(Al,Fe)OOH Single Crystals and Seismic Detectability of Hydrous MORB in the
424 Shallow Lower Mantle, *Geophysical Research Letters*, 48(23), e2021GL094185.
- 425 Spivak, A., Solopova, N., Cerantola, V., Bykova, E., Zakharchenko, E., Dubrovinsky, L., et al. (2014),
426 Raman study of MgCO_3 - FeCO_3 carbonate solid solution at high pressures up to 55 GPa, *Physics and*
427 *Chemistry of Minerals*, 41(8), 633-638.

- 428 Su, X., Zhao, C., Xu, L., Lv, C., Song, X., Ishii, T., et al. (2021a), Spectroscopic evidence for the Fe³⁺
429 spin transition in iron-bearing δ -AlOOH at high pressure, *American Mineralogist*, 106(11), 1709–
430 1716.
- 431 Su, X., Zhao, C., Lv, C., Zhuang, Y., Salke, N., Xu, L., et al. (2021b), The effect of iron on the sound
432 velocities of δ -AlOOH up to 135 GPa, *Geoscience Frontiers*, 12(2), 937-946.
- 433 Suzuki, A., Ohtani, E., Kamada, T. (2000), A new hydrous phase δ -AlOOH synthesized at 21 GPa and
434 1000 °C, *Physics and Chemistry of Minerals*, 27(10), 689-693.
- 435 Tada, N., Baba, K., Utada, H. (2014), Three-dimensional inversion of seafloor magnetotelluric data
436 collected in the Philippine Sea and the western margin of the northwest Pacific Ocean, *Geochemistry,*
437 *geophysics, geosystems*: 15(7), 2895-2917.
- 438 Tarits, P., Mandéa, M. (2010), The heterogeneous electrical conductivity structure of the lower mantle,
439 *Physics of the Earth Planetary Interiors*, 183, 115-125.
- 440 Thompson, E. C., Campbell, A. J., and Tsuchiya, J. (2017), Elasticity of ϵ -FeOOH: Seismic implications
441 for Earth's lower mantle, *Journal of Geophysical Research: Solid Earth*, 122, 5038-5047.
- 442 Thompson, E. C., Davis, A. H., Brauser, N. M., Liu, Z., Prakapenka, V. B., Campbell A. J. (2020), Phase
443 transitions in ϵ -FeOOH at high pressure and ambient temperature, *American Mineralogist*, 105(12),
444 1769-1777.
- 445 Tsuchiya, J., Tsuchiya, T. (2011), First-principles prediction of a high-pressure hydrous phase of AlOOH,
446 *Physical Review B*, 83(5), 054115.
- 447 Tsuchiya, J., Tsuchiya, T., Tsuneyuki, S., Yamanaka, T. (2002), First-principles calculation of a high-
448 pressure hydrous phase, δ -AlOOH, *Geophysical Research Letters*, 29(19), 15-1-15-4.
- 449 Xu, C., Inoue, T. (2019), Melting of Al-Rich Phase D up to the Uppermost Lower Mantle and
450 Transportation of H₂O to the Deep Earth, *Geochemistry, Geophysics, Geosystems*, 20(9), 4382-4389.
- 451 Wang, R., Yoshino, T. (2021), Electrical conductivity of diaspore, δ -AlOOH and ϵ -FeOOH, *American*
452 *mineralogist*, 106(5), 774-781.
- 453 Zhuang, Y, Gan, B., Cui, Z., Tang, R., Tao, R., Hou, M., et al. (2022), Mid-mantle water transportation
454 implied by the electrical and seismic properties of ϵ -FeOOH, *Science Bulletin*, 67(7), 748-754.
455

456 **Figure Captions**

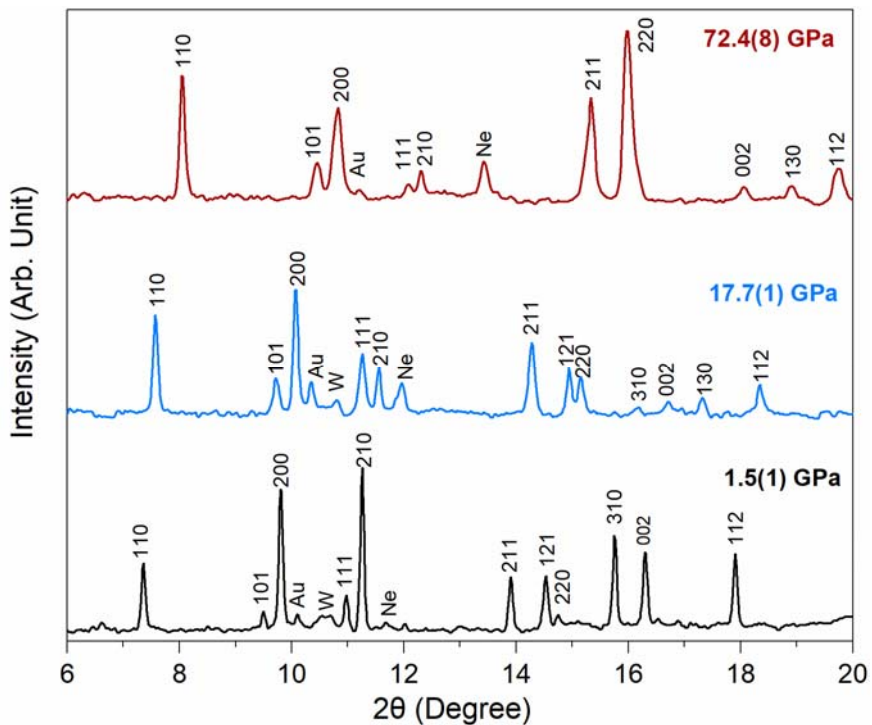
457



458

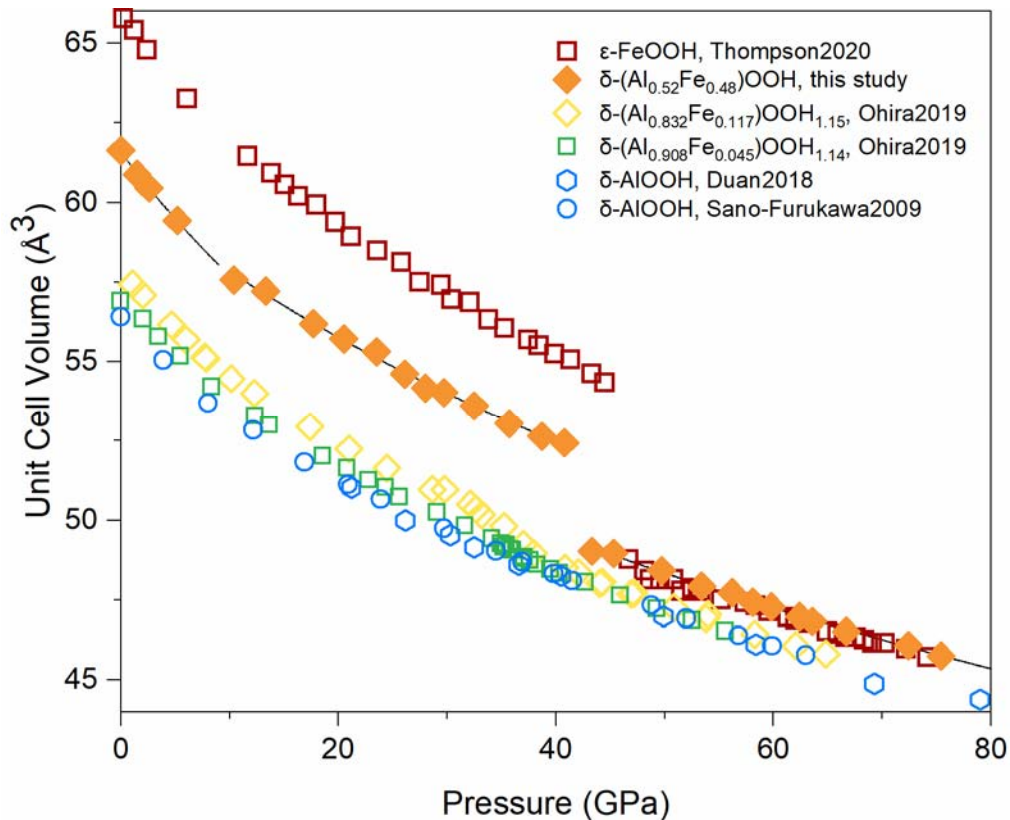
459

460 **Figure 1.** Unit-cell parameters of δ -(Al,Fe)OOH at ambient conditions. (a) a . (b) b . (c) c . (d) V .
461 Red diamonds: δ -(Al_{0.52}Fe_{0.48})OOH, this study; green diamonds: δ -(Al_{0.95}Fe_{0.05})OOH, *Su et al.*
462 (2021b); orange diamonds: δ -(Al_{0.87}Fe_{0.13})OOH, *Buchen et al.* (2021); blue diamonds: δ -AlOOH,
463 δ -(Al_{0.832}Fe_{0.117})OOH_{1.15}, and δ -(Al_{0.908}Fe_{0.047})OOH_{1.14}, *Ohira et al.* (2019); green circles: δ -
464 AlOOH, *Sano-Furukawa et al.* (2009); blue triangle: ϵ -FeOOH, *Suzuki et al.* (2010). Dashed
465 lines connect the end-member values.



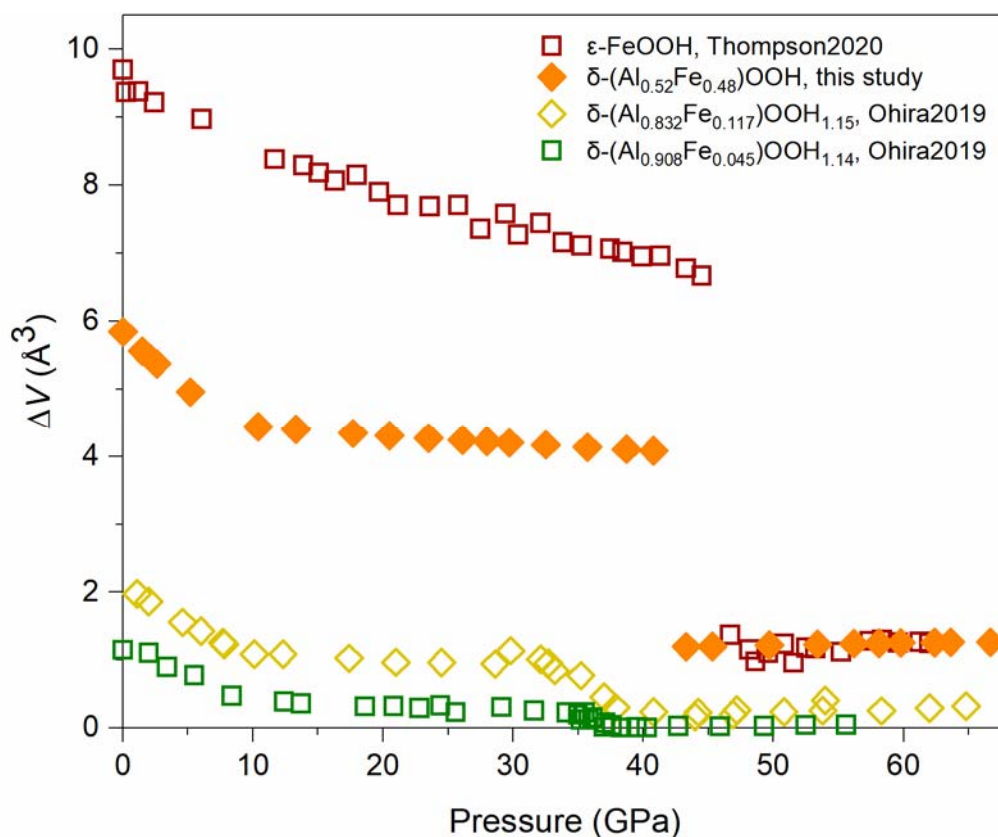
466

467 **Figure 2.** Representative X-ray diffraction patterns of δ -(Al_{0.52}Fe_{0.48})OOH at high pressure and
468 300 K. The equation of state of Au was used to determine pressure and its uncertainty (*Fei et al.*,
469 2007). Neon was used as the pressure medium. W indicates the diffraction lines of the tungsten
470 gasket. The wavelength of the monochromatic X-ray beam was 0.4133 Å (corresponding to an
471 X-ray energy of 30.0 keV).



472

473 **Figure 3.** Pressure-volume relations of δ -(Al,Fe)OOH at high pressure and 300 K. Solid orange
474 diamonds: experimental measurements of δ -(Al_{0.52}Fe_{0.48})OOH, this study; black curves: the 3rd
475 Birch-Murnaghan (BM) EoS fits to the experimental data at 0–10.4, 13.3–40.8, and 43.3–75.4
476 GPa, respectively, this study; yellow diamonds and green squares: δ -(Al_{0.832}Fe_{0.117})OOH_{1.15} and
477 δ -(Al_{0.908}Fe_{0.047})OOH_{1.14}, respectively, *Ohira et al.* (2019); red squares: ϵ -FeOOH, *Thompson et*
478 *al.* (2020); blue circles and hexagons: δ -AlOOH, *Sano-Furukawa et al.* (2009) and *Duan et al.*
479 (2018), respectively. Errors for the Delta52 are smaller than the symbols and are not shown for
480 clarity.

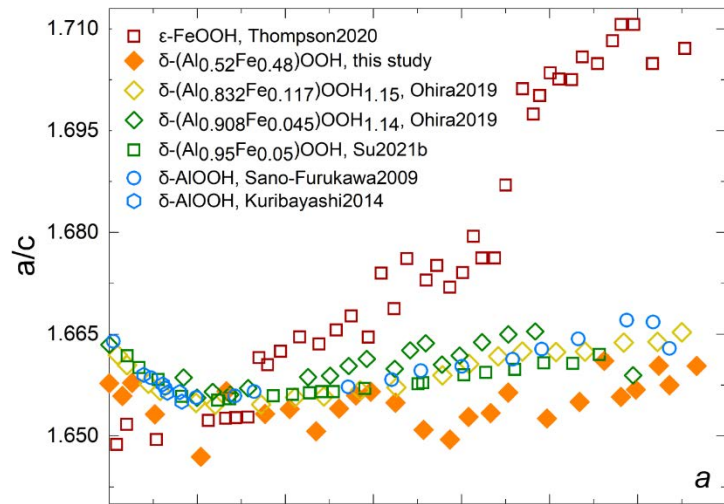


481

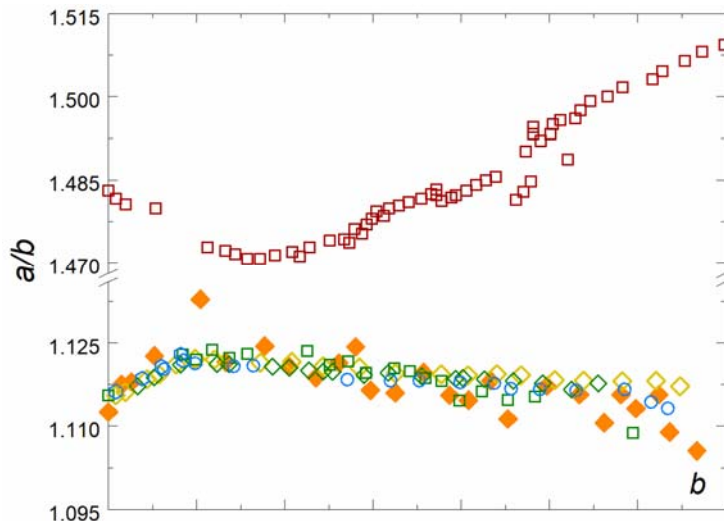
482 **Figure 4.** The unit cell volume difference (ΔV) between δ -(Al,Fe)OOH and pure δ -AlOOH at
483 high pressure. The equation of state for δ -AlOOH is from *Sano-Furukawa et al. (2009)*. A sharp
484 discontinuity was observed in δ -(Al_{0.52}Fe_{0.48})OOH (denoted as “Delta52”) across the spin
485 transition of iron. Yellow diamonds and green squares: δ -(Al_{0.832}Fe_{0.117})OOH_{1.15} and δ -
486 (Al_{0.908}Fe_{0.047})OOH_{1.14}, respectively, *Ohira et al. (2019)*; red squares: ϵ -FeOOH, *Thompson et al.*
487 (2020). Errors for the Delta52 are smaller than the symbol size and are not be shown for clarity.

488

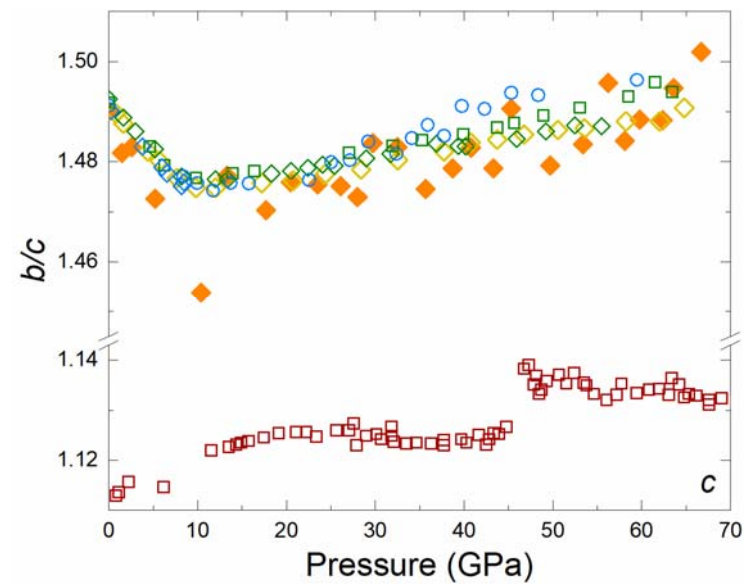
489



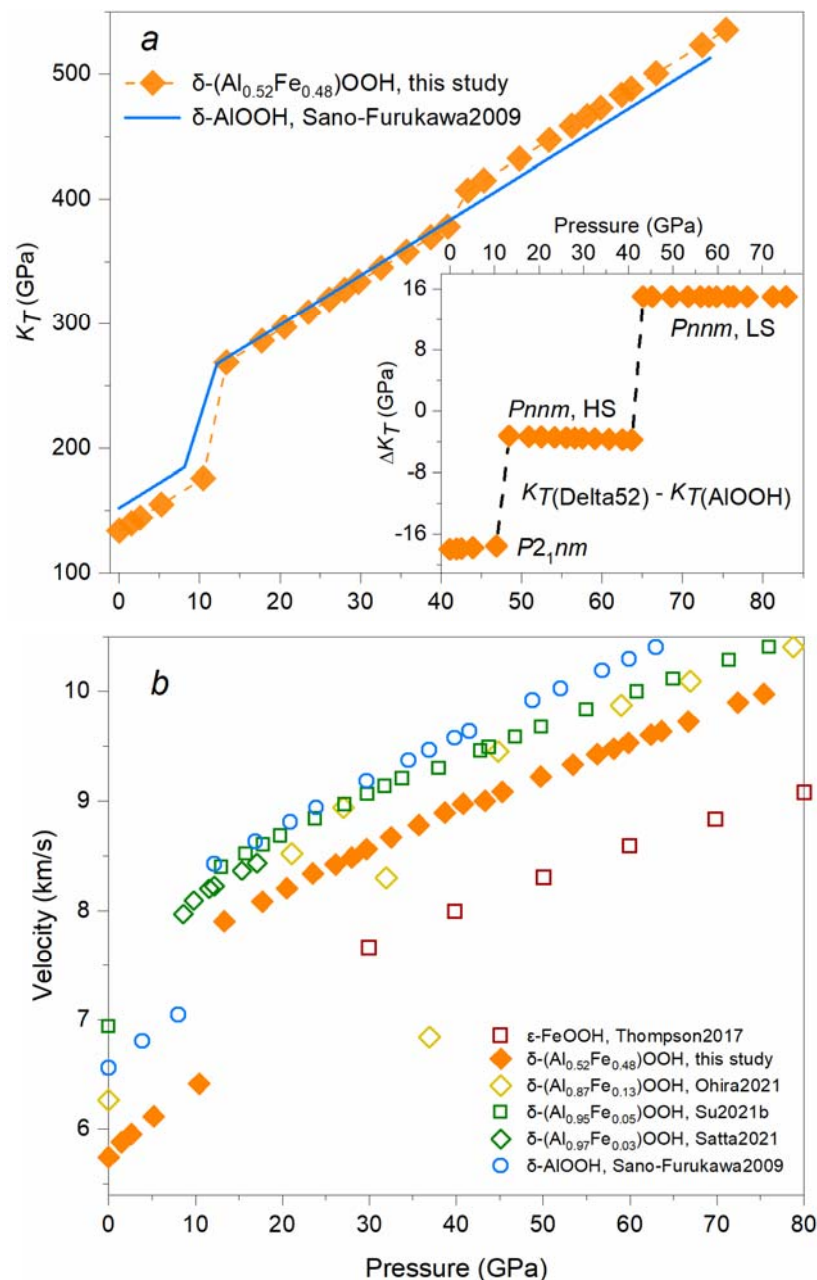
490



491



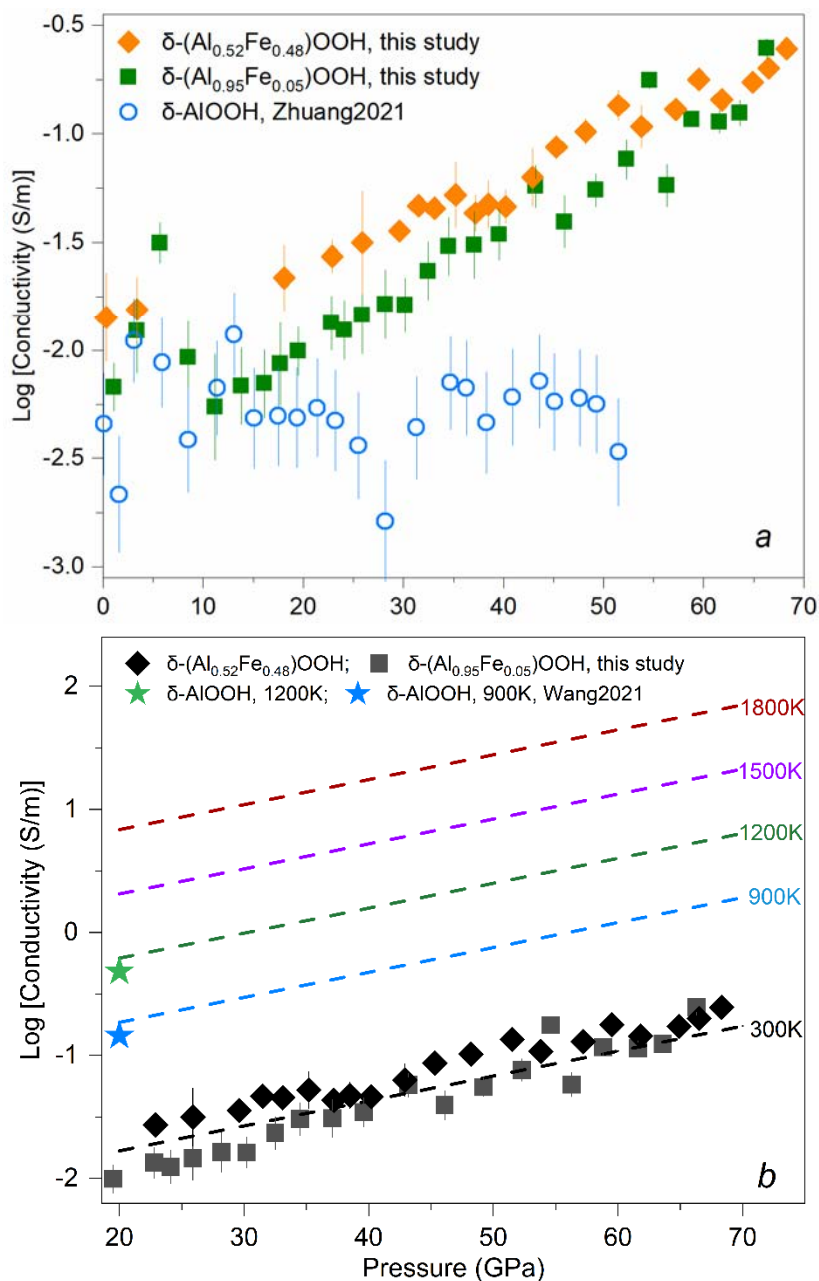
492 **Figure 5.** Axial ratios as a function of pressure for δ -(Al,Fe)OOH. The axial ratio values of δ -
493 (Al_{0.52}Fe_{0.48})OOH undergo discontinuity at approximately 10 GPa, in good agreement with the
494 two new Raman bands observed in this sample at 8.5–15.8 GPa by *Su et al.* (2021a). Yellow
495 diamonds and green squares: δ -(Al_{0.832}Fe_{0.117})OOH_{1.15} and δ -(Al_{0.908}Fe_{0.047})OOH_{1.14}, respectively,
496 *Ohira et al.* (2019); red squares: ϵ -FeOOH, *Thompson et al.* (2020); blue circles and hexagons:
497 δ -AlOOH, *Sano-Furukawa et al.* (2009) and *Kuribayashi et al.* (2014), respectively. Errors for
498 the Delta52 are smaller than the symbol size and are not be shown for clarity.
499



500

501

502 **Figure 6.** Isothermal bulk modulus (K_T) (a) and bulk sound velocities (b) of δ -(Al,Fe)OOH as a
 503 function of pressure at 300 K. Inset: the isothermal bulk modulus difference (ΔK_T) between δ -
 504 (Al_{0.52}Fe_{0.48})OOH and pure δ -AlOOH at high pressure. Orange solid diamonds: δ -
 505 (Al_{0.52}Fe_{0.48})OOH, this study; blue curve and open circles: δ -AlOOH, Sano-Furukawa *et al.*
 506 (2009); yellow diamonds: δ -(Al_{0.87}Fe_{0.13})OOH, Ohira *et al.* (2021); green squares: δ -
 507 (Al_{0.95}Fe_{0.05})OOH, Su *et al.* (2021b); green diamonds: δ -(Al_{0.97}Fe_{0.03})OOH, Satta *et al.* (2021).

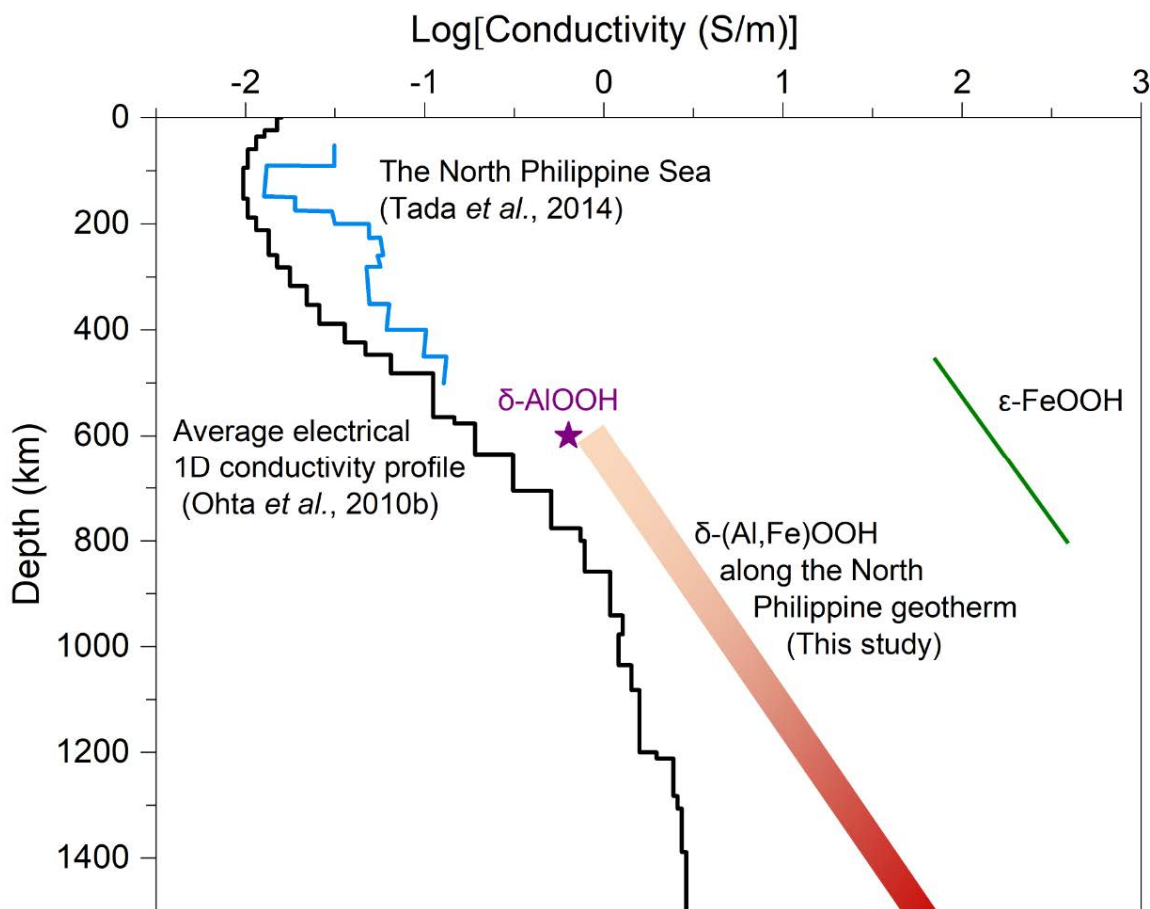


508

509

510 **Figure 7.** Electrical conductivity of the δ -(Al,Fe)OOH phase as a function of pressure at room
511 temperature (a) and high temperature (b). Solid squares and diamonds: δ -(Al_{0.95}Fe_{0.05})OOH and
512 δ -(Al_{0.52}Fe_{0.48})OOH, this study; open circles: δ -AlOOH, *Zhuang et al.* (2021); star symbols: δ -
513 AlOOH at 900 K (blue) and 1200 K (olive), *Wang and Yoshino* (2021); dashed curves: the EC
514 fits at 300 K (black), 900 K (blue), 1200 K (olive), 1500 K (purple), and 1800 K (red),
515 respectively. Vertical ticks represent the uncertainties for pure and iron-bearing δ -AlOOH.

516



517

518 **Figure 8.** The conductivity-depth profile of the δ -(Al,Fe)OOH phase along the North Philippine
519 geotherm. The orange-red envelope is a 95% confidence interval for the EC values of the δ -
520 (Al,Fe)OOH phase. Blue curve: the conductivity-depth profile of the North Philippine Sea, *Tada*
521 *et al.* (2014); black curve: average electrical 1D conductivity profile, *Ohta et al.* (2010b); purple
522 star symbol: the electrical conductivity of δ -AlOOH at 20 GPa, *Wang and Yoshino* (2021); olive
523 curve: the conductivity-depth profile of ϵ -FeOOH along the North Philippine geotherm, *Wang*
524 *and Yoshino* (2021).

525 **Table 1.** Lattice parameters of δ -(Al_{0.52}Fe_{0.48})OOH at high pressure and 300 K

<i>P</i> (GPa)	<i>a</i> (Å)	<i>b</i> (Å)	<i>c</i> (Å)	<i>V</i> (Å ³)
0.0(0)	4.8440(3)	4.3540(5)	2.9220(6)	61.63(1)
1.5(1)	4.8298(8)	4.3220(2)	2.9167(5)	60.88(3)
2.6(1)	4.8208(8)	4.3122(2)	2.9079(5)	60.45(3)
5.2(1)	4.7958(7)	4.2718(2)	2.9009(4)	59.43(3)
10.4(1)	4.7534(1)	4.1960(1)	2.8862(7)	57.57(7)
13.3(1)	4.7371(2)	4.2238(8)	2.8595(7)	57.22(7)
17.7(1)	4.7095(1)	4.1884(1)	2.8486(7)	56.19(7)
20.5(2)	4.6918(6)	4.1866(9)	2.8367(4)	55.72(3)
23.5(1)	4.6748(7)	4.1786(1)	2.8320(6)	55.32(3)
26.1(1)	4.6616(6)	4.1571(8)	2.8183(4)	54.62(3)
28.0(2)	4.6548(6)	4.1402(1)	2.8109(4)	54.17(3)
29.7(2)	4.6404(9)	4.1564(9)	2.8013(6)	54.03(7)
32.5(3)	4.6263(9)	4.1454(8)	2.7955(6)	53.61(6)
35.7(3)	4.6115(9)	4.1187(7)	2.7933(6)	53.05(6)
38.7(5)	4.5925(9)	4.1168(7)	2.7842(6)	52.64(6)
40.8(7)	4.5882(9)	4.1162(1)	2.7760(5)	52.43(6)
43.3(7)	4.4921(8)	4.0173(4)	2.7169(5)	49.03(6)
45.3(8)	4.4834(8)	4.0344(4)	2.7067(5)	48.96(6)
49.7(8)	4.4714(8)	4.0022(4)	2.7057(5)	48.42(6)
53.4(7)	4.4563(8)	3.9942(4)	2.6926(5)	47.93(6)
56.2(6)	4.4487(8)	4.0057(3)	2.6782(9)	47.73(6)
58.1(6)	4.4414(5)	3.9810(5)	2.6824(4)	47.43(2)
59.8(6)	4.4349(5)	3.9840(4)	2.6767(8)	47.29(2)
62.4(6)	4.4314(5)	3.9720(3)	2.6689(8)	46.98(2)
63.6(7)	4.4156(8)	3.9818(3)	2.6640(9)	46.84(6)
66.7(8)	4.4034(7)	3.9830(2)	2.6521(8)	46.52(6)
72.4(8)	4.3872(7)	3.9773(2)	2.6413(8)	46.09(5)
75.4(9)	4.3790(7)	3.9565(1)	2.6411(8)	45.76(5)

526

527 **Table 2.** Equation of state parameters of δ -(Al_{0.52}Fe_{0.48})OOH at 300 K.

Composition	V_0 (Å ³)	K_0 (GPa)	K_{T0}'	EoS	P range	References
δ -(Al _{0.52} Fe _{0.48})OOH ^h	61.6(1)	134 (4)	4 (fixed)	2nd BM	0.0–10.4	This study
δ -(Al _{0.52} Fe _{0.48})OOH ^l	60.4(1)	216 (2)	3.98(9)	3rd BM	13.3–40.8	This study
δ -(Al _{0.52} Fe _{0.48})OOH ^j	56.5(1)	234 (2)	4.00(2)	3rd BM	43.3–75.4	This study
δ -(Al _{0.832} Fe _{0.117})OOH _{1.15} ^h	57.85(2)	147(1)	4 (fixed)	2nd BM	1.1–10.1	Ohira et al. (2019)
δ -(Al _{0.832} Fe _{0.117})OOH _{1.15}	55.2(4)	241(14)	4 (fixed)	SC EoS	36.1–64.8	Ohira et al. (2019)
δ -(Al _{0.908} Fe _{0.045})OOH _{1.14} ^h	57.03(7)	152(7)	4 (fixed)	2nd BM	0.0–8.4	Ohira et al. (2019)
δ -(Al _{0.908} Fe _{0.045})OOH _{1.14}	55.4(3)	223(11)	4 (fixed)	SC EoS	34.9–55.6	Ohira et al. (2019)
δ -AlOOH ^a	56.408(9)	152(2)	4 (fixed)	2nd BM	0.0–10.0	Sano-Furukawa et al. (2009)
δ -AlOOH ^b	55.47(8)	219(3)	4 (fixed)	2nd BM	10.0–63.5	

528 ^a $P2_1nm$ at the high-spin state of δ -AlOOH; ^b $Pnmm$ at the high-spin state of δ -AlOOH; ^h $P2_1nm$ at the high-spin state
 529 of δ -(Fe,Al)OOH; ^l $Pnmm$ at the high-spin state of δ -(Fe,Al)OOH; ^j $Pnmm$ at the low-spin state of δ -(Fe,Al)OOH; SC
 530 EoS= Spin crossover EoS; BM = Birch-Murnaghan.

531 **Table 3.** Volume collapse of δ -(Al,Fe)OOH across the spin transition at room temperature.

Composition	PTM ^a	Pressure calibrant	Spin crossover (GPa)	Volume collapse	Reference
δ -(Al _{0.52} Fe _{0.48})OOH	Ne	Au	41–45	6.5%	This study
δ -(Al _{0.95} Fe _{0.05})OOH	Ne	Au	30–42	1.3%	Su et al. (2021b)
δ -(Al _{0.908} Fe _{0.045})OOH _{1.14}	Ne	Ruby	32–40	1.3%	Ohira et al. (2019)
δ -(Al _{0.832} Fe _{0.117})OOH _{1.15}	He	Ruby	32–40	2.8%	Ohira et al. (2019)
ϵ -FeOOH	Ne	Ruby	~45(2)	10.3%	Thompson et al. (2020)

532 ^a PTM: Pressure-transmitting medium.

533

534 **Table 4.** The electrical conductivity of iron-bearing δ -AlOOH at high pressures and 300 K.

δ -(Al _{0.95} Fe _{0.05})OOH		δ -(Al _{0.52} Fe _{0.48})OOH	
<i>P</i> (GPa)	Log[σ (S/m)]	<i>P</i> (GPa)	Log[σ (S/m)]
1.1(1)	-2.17(32)	0.3(1)	-1.84(27)
3.4(3)	-1.90(28)	3.4(4)	-1.80(26)
5.7(6)	-1.50(22)	18.1(19)	-1.66(24)
8.5(9)	-2.03(31)	22.9(25)	-1.56(23)
11.2(11)	-2.26(33)	25.9(28)	-1.49(22)
13.8(15)	-2.16(32)	29.6(32)	-1.44(21)
16.1(17)	-2.15(32)	31.5(34)	-1.33(19)
17.7(19)	-2.06(31)	33.1(36)	-1.34(21)
19.5(21)	-2.00(31)	35.2(38)	-1.28(19)
22.8(25)	-1.87(28)	37.2(41)	-1.36(21)
24.1(26)	-1.90(27)	38.5(42)	-1.32(19)
25.9(28)	-1.83(26)	40.2(44)	-1.33(21)
28.2(31)	-1.78(23)	42.9(47)	-1.19(18)
30.2(33)	-1.78(24)	45.3(49)	-1.06(16)
32.5(35)	-1.63(22)	48.2(53)	-0.98(14)
34.5(37)	-1.51(23)	51.5(56)	-0.86(13)
37.1(41)	-1.51(22)	53.8(59)	-0.96(14)
39.6(43)	-1.46(19)	57.2(62)	-0.88(13)
43.2(47)	-1.24(21)	59.5(65)	-0.74(11)
46.1(51)	-1.40(18)	61.8(67)	-0.84(12)
49.2(54)	-1.25(19)	64.9(71)	-0.76(11)
52.3(57)	-1.11(16)	66.5(73)	-0.69(9)
54.6(61)	-0.75(11)	68.3(75)	-0.60(9)
56.3(62)	-1.23(18)		
58.8(64)	-0.93(14)		
61.6(67)	-0.94(13)		
63.6(69)	-0.90(15)		
66.3(71)	-0.60(9)		

535 Conductivity error was estimated by the following equation:

536 $\Delta\sigma \approx \sigma \frac{-d^2}{16D^2 \ln 2}$, where *d* is the electrode contact length, and *D* is the diameter of the sample. The

537 uncertainty was estimated to be ~ 15%.