

1 **REVISION 1**

2
3 **SODIUM NANOPARTICLES IN ALKALI HALIDE MINERALS:**

4 **WHY IS VILLIAUMITE RED AND HALITE BLUE?**

5
6 Georges Calas^a, Laurence Galois^a and Alexis Geisler^a

7
8 ^a *Sorbonne Université, Muséum National d'Histoire Naturelle, CNRS, Institut de Minéralogie,*
9 *de Physique des Matériaux et de Cosmochimie, IMPMC, 75005 Paris, France*

10 **corresponding author : Georges Calas, Sorbonne Université, IMPMC, BC 115, 4 place*
11 *Jussieu 75005 Paris, France. georges.calas@sorbonne-universite.fr*

12
13
14 **ABSTRACT**

15 The presence of metal Na nanoparticles causes the bright, thermally unstable colors of
16 villiaumite, NaF, and halite, NaCl. These nanoparticles have been suspected since a long time to
17 be caused by external irradiation. Metal nanoparticles, often referred to as metal colloids, cause
18 surface plasmon resonance effects, characterized by a single Lorentzian-shaped absorption band.
19 The color of these minerals is due to metal Na nanoparticles of 2.5-3 nm. A key point is that the
20 resonance wavelength, which corresponds to the maximum of the absorption band, is inversely
21 related to the value of the refractive index of the embedding mineral. This causes the position of
22 the main absorption band to be offset downwards by 140 nm in halite relative to villiaumite. As a
23 consequence, the optical transmission window is shifted from the long to the short wavelength
24 domain, explaining the color of blue halite and red villiaumite, respectively. Similar refractive

25 index dependence may explain the purple color of fluorite, caused by metallic Ca nanoparticles.
26 Finally, the origin of the villiaumite irradiation may be the presence of Th-rich (about 8.8wt%
27 ThO₂) nano-inclusions, about 500 nm large, illustrating the specific geochemistry of peralkaline
28 rocks where villiaumite is found.

29
30 **Keywords:** villiaumite, halite, fluorite, peralkaline rocks, color, radiation damage, UV-visible
31 spectroscopy, nanoparticles

32

33 INTRODUCTION

34 One of the outstanding properties of alkali halides is their optical transparency from the
35 vacuum UV to the far infrared. Though they are intrinsically colorless, they sometimes show a
36 thermally unstable coloration, which, in the absence of chemical impurities, has long been
37 recognized to result from radiation damage (e.g., Prizibram, 1953; Stormer and Carmichael, 1970).
38 The simplest radiation-induced defect is the so-called F center, i.e. an electron trapped in an
39 anion vacancy. The F center creates a singly occupied electronic level in the band gap, which
40 colors the crystal. The aggregation of two or three F centers on nearest neighbor sites, gives rise
41 to binary M- and ternary R-centers, respectively. Eventually, color centers aggregate to form
42 metallic alkali nanoparticles, also referred to as colloids, caused by an irradiation at room
43 temperature or higher temperatures (see e.g., Schwartz et al., 2008). Optical properties of metal
44 nanoparticles have unique characteristics that give rise to brilliant colors. For instance, during
45 external irradiation of synthetic halite, the change from a yellow color due to *F*-centers to a bright
46 blue color indicates the formation of Na metal nanoparticles (Kreibig and Vollmer, 1995).

47 Natural halite, NaCl, presents various radiation-induced colors (Zelek et al., 2015), but blue
48 hues with a broad range of saturation are the most frequent (Supplemental Fig. 1). The blue color
49 has been assigned to colloidal Na metal formed by irradiation with ionizing radiation (Rossman,
50 2010). Radiation defects in halite provide information on the sedimentary history of salt deposits
51 (Sonnenfeld, 1995; Zelek et al., 2014 and 2015) or the stability of nuclear waste repositories
52 (Levy et al., 1983). Recently, they have been used to assess the exposure duration of the surface
53 of icy moons (Poston et al., 2017) and ordinary chondrites (Chan et al., 2018), subject to cosmic
54 radiation. Villiaumite, NaF, is an accessory phase formed during a late pegmatitic stage
55 associated with peralkaline nepheline syenites (Stormer and Carmichael, 1970; Marks and Markl,
56 2017). It is generally characterized by its intense carmine red color (Supplemental Fig. S2).

57 The origin of the red color of natural villiaumite has not yet been investigated, though it is
58 suspected to arise from radiation damage (Rossman, 2010). By contrast, blue halite received
59 much attention, with pioneering studies on samples from Strassfurt, Germany (Przibram, 1953;
60 Doyle, 1960; Howard and Kerr, 1960; Arun et al., 2017) and, more recently, Kłodawa, Poland
61 (Weselucha-Birczynska et al., 2012; Zelek et al., 2014 and 2015) and Morleben, Germany (Arun et
62 al., 2017) salt mines. This study shows that the red color of villiaumite is a result of a surface
63 plasmon resonance (Kreibig and Vollmer, 1995) due to Na metal nanoparticles. The optical
64 spectra of blue halite have a similar origin. The outstanding color change between blue halite and
65 red villiaumite is rationalized in terms of the dependence of the wavelength resonance of the
66 nanoparticles on the refractive index of the embedding mineral. The same formalism may be
67 extended to explain the origin of the purple color of fluorite, due to the presence of calcium
68 metallic nanoparticles. The origin of the villiaumite irradiation may be the presence of Th-rich
69 (about 8.8wt% ThO₂) nano-inclusions, about 500 nm large, illustrating the geochemistry of
70 peralkaline rocks where villiaumite is found.

71

72

MATERIALS AND METHODS

73 Villiaumite crystals come from the agpaitic suite of nepheline syenites of the type locality of

74 the Los Archipelago, Guinea (Lacroix, 1908; Moreau et al., 1998). A sedimentary navy blue

75 halite from Strassfurt (Germany) was used for comparison. Optical absorption spectra were

76 measured at room temperature and at 10K in the spectral range 200–3300 nm ($50,000\text{ cm}^{-1}$ - $3,030$

77 cm^{-1}), using a double-beam computerized Perkin-Elmer Lambda 1050 UV–Visible-NIR

78 spectrophotometer. The spectral resolution varies from 0.8 nm in the UV region to 2 nm in the

79 near IR - visible region. A He-cryostat under vacuum (around 3.4×10^{-7} mbar) was used to

80 record spectra at 10 K. The optical absorption spectra were obtained in transmission mode on

81 cleaved crystals (Supplemental Fig. S3). The spectra were normalized to sample thickness and

82 background corrected using a polynomial function. The data are presented and analyzed as a

83 function of wavelength, as the absorption is caused by surface plasmon resonance effects

84 interpreted in terms of the Mie theory (see e.g., Kreibig and Vollmer, 1995). Preliminary

85 scanning electron microscopy with field emission gun (SEM-FEG) analyses were obtained at 15

86 kV and a beam current of 200 nA with a Zeiss Ultra 55 instrument fitted with a high-resolution

87 Schottky FEG and a UHR Gemini® column. Semi-quantitative analyses were performed using a

88 Bruker Quantax XFlash 4010 energy-dispersive X-ray spectrometer.

89

90

RESULTS AND DISCUSSION

Optical absorption spectra of villiaumite

92 The samples investigated present an intense red color, which disappears in a couple of minutes

93 at 500 °C and after 2 hrs at 400°C. The room temperature optical absorption spectrum

94 (Supplemental Fig. S4) shows the presence of a background due to light scattering by inclusions

95 and fractures that result from the easy {100} cleavage of this cubic mineral. After subtracting this
96 background contribution, the most salient feature is an intense absorption band at 494 nm, which
97 extends over most of the visible spectrum because of its tail extending at long wavelengths. The
98 two other contributions are a shoulder near 415 nm and a small band at 328 nm (Fig. 1). This
99 spectrum is similar to the one of villiaumite from Mont Saint Hilaire, Canada (Rossman, 2010),
100 which is dominated by a main absorption band peaking at 510 nm.

101 The optical absorption spectra have been fitted using a minimum of components, chosen to
102 correspond to spectroscopic events. Gaussian and Lorentzian lineshapes of the spectral
103 components have been tested. Indeed, by contrast to the Gaussian shape of absorption bands
104 caused by color centers (see e.g. Jenkins et al., 2000; Hoya et al., 2017), surface plasmon
105 resonances exhibited by metallic nanoparticles give rise to Lorentzian-shaped absorption bands,
106 as predicted by Mie theory (Kreibig and Vollmer, 1995; Seinen et al., 1994; Ruiz-Fuertes et al.,
107 2019). A good fit of the main band is achieved by using a Lorentzian function peaking at 494 nm.
108 The full width at half-maximum (FWHM) of this band, 86 nm, is larger than that of the other
109 contributions due to color centers. This lineshape explains the presence of a long tail extending
110 towards long wavelengths, at the origin of the intense red hue of villiaumite. Such a Lorentzian
111 line-shape characterizes the optical spectra of free neutral Na clusters (Selby et al., 1991), Na
112 nanoparticles in NaCl (Seinen et al., 1994) and NaF (Seifert et al., 1994) or Ca nanoparticles in
113 CaF₂ (Ruiz- Fuertes et al., 2019; Ryskin et al., 2020).

114 The two minor contributions at 328 nm and 414 nm have a Gaussian lineshape. The former
115 corresponds to a *F*-center, widely investigated in synthetic NaF (Seifert et al., 1994; Tiwald et al.,
116 2015; Hoya et al., 2017). The latter may be assigned to a *R*-center (Amenu-Kpodo and Neubert,
117 1965; Bryukvina and Martynovich, 2012; Bryukvina et al., 2018). Recent ab-initio calculations

118 (Tiwald et al., 2015; Hoya et al., 2017) have shown that the absorption energy E_a of the F -center
119 (in eV) may be approximated by:

$$120 \quad E_a = 16.5 a^{-1.76} \quad (1)$$

121 where a is the anion-cation distance. Relation (1) gives a physical ground to the classical
122 empirical Mollwo-Ivey relation (Ivey, 1947). The predicted values, 345 and 415 nm, are in good
123 agreement with the experimental values, 328 and 414 nm. In the optical absorption spectrum of
124 villiaumite from Mont Saint Hilaire, Canada (Rossman, 2010), the main absorption band at 510
125 nm shows a slightly different lineshape and additional weak contributions around 400 nm. As in
126 natural blue halite (Zepek et al., 2015), the relative proportion of isolated defect centers and
127 metallic nanoparticles may vary among the samples.

128

129 **Evidence of a plasmon resonance in villiaumite**

130 As indicated above, a Lorentzian lineshape is consistent with a surface plasmon resonance.
131 The position of this resonance is similar to that in synthetic NaF (Chandra and Holcomb, 1969;
132 Bryukvina and Martynovich, 2012; Bryukvina et al., 2018). The resonance wavelength may be
133 predicted from the Mie theory by considering spherical metal particles (Doyle, 1958; Hunault et
134 al., 2017). In this approximation, the average size of the nanoparticles is derived from:

$$135 \quad R = V_f \lambda_p^2 / (2\pi c \Delta\lambda) \quad (2)$$

136 where R is the average radius of the metallic clusters, V_f is the Fermi velocity of the electrons in
137 the bulk metal (for Na, $V_f = 1.07 \times 10^6 \text{ m.s}^{-1}$), λ_p is the characteristic wavelength at which the
138 surface plasmon resonance (SPR) occurs, $\Delta\lambda$ is the full width at half-maximum and c is the speed
139 of light. This predicts an average diameter of Na nanoparticles slightly smaller than 3 nm. At 10
140 K, this absorption band only slightly shifts by about 10 nm towards higher wavenumbers, without
141 any narrowing. This is consistent with an assignment to a surface plasmon resonance, which does

142 not change in energy or amplitude with temperature, by contrast to the transitions related to color
143 centers (Kreibig and Vollmer, 1995).

144 Color centers in alkali halides have a limited thermal stability above room temperature relative
145 to Na colloids, (Schwartz et al., 2008). This explains why, in the geological samples, a plasmon
146 resonance is predominant over the electronic transitions expected from these color centers.
147 Similar mechanisms have been observed in glasses where alkali ions can act as electron traps
148 upon irradiation and form metal colloids (Boizot et al., 2000). Such processes are thermally
149 activated and demonstrate that alkali ions agglomerate to form bigger complexes after trapping
150 electrons at temperatures reaching a few hundreds °C.

151

152 **Comparison with halite**

153 The shape of the optical absorption spectrum of blue halite from Stassfurt is similar to some of
154 the blue halites above mentioned. It is almost identical to the first absorption spectra published on
155 blue halite (also from Stassfurt: Doyle, 1960) (Supplemental Fig. S5). The main absorption band
156 peaks at 640 nm, the same value as reported by previous authors (e.g., Doyle, 1960; Howard and
157 Kerr, 1960). Using several Gaussian-Lorentzian spectral components, assigned to Na colloids and
158 various color centers, gives a good fit (Zelek et al., 2014 and 2015). However, fitting the spectra
159 is non equivocal. Here, we use a minimum number of spectral components, as for villiaumite.
160 The main band, at 640 nm, is fitted with a single Lorentzian component (Fig. 2), because it arises
161 from a surface plasmon resonance of Na nanoparticles (e.g., Seinen et al., 1994). The FWHM,
162 105 nm, is similar to the values found in the Klodawa mine (Weselucha-Birczyńska et al., 2012)
163 and larger than in villiaumite (86 nm) . Minor additional Gaussian-shape contributions at 430,
164 531 and 744 nm may be assigned to *F*-, *R*- (3 neighboring *F*-centers) and *M* (2 neighboring *F*-
165 centers) color centers, respectively. The actual position of these bands is shifted relative to that

166 expected from relation (1), at 473, 550 and 705 nm for the *F*-, *R*- and *M*-centers-, respectively.
167 The same centers are found in the halite from Klodawa mine and a similar discrepancy with the
168 Mollwo-Ivey relation is observed (Weselucha-Birczyńska et al., 2012). Finally, a fourth minor
169 Gaussian contribution occurs at 385 nm as in other natural halites (Doyle, 1960; Weselucha-
170 Birczyńska et al., 2012), but its origin is unclear.

171

172 **Scaling the spectra of red villiaumite and blue halite**

173 The halite spectrum presents similarities with that of villiaumite, but occurs at larger
174 wavelengths. As a consequence, the transmission window does not occur in the same spectral
175 region, being located on the long or short wavelength side of the main band in red villiaumite and
176 blue halite, respectively (Figs. 1 and 2). This explains the difference in the color of these minerals,
177 despite both show a broad Lorentzian shape plasmon resonance with spectral properties
178 independent of temperature, which is assigned to the presence of Na nanoparticles. For spherical
179 particles that are smaller than the wavelength of light, the value of the resonance wavelength
180 depends on the refractive index of the surroundings (Kreibig and Vollmer, 1995). The
181 wavelength of maximum absorption λ_{\max} may be predicted within the Mie theory, following the
182 Doyle relation (Doyle, 1958; Davenas et al., 1973; Seifert et al., 1994):

$$183 \lambda_{\max} = \lambda_c (1 + 2n_0^2)^{1/2} \quad (3)$$

184 where n_0 is the refractive index of the host medium and λ_c is the critical wavelength for the onset
185 of ultra-violet transparency of sodium. The ratio between the λ_{\max} values in villiaumite and halite,
186 allows elimination of the λ_c term. The ratio calculated from the refractive index of these minerals
187 is 1.13 to compare to an experimental value of 1.29. This explains qualitatively the redshift of the
188 plasmon band with the increasing refractive index of the mineral. The underestimation of this

189 shift when using relation (3) may come from non-spherical shape effects or size distribution of
190 the Na colloids in geological samples with a complex history relative to laboratory samples...

191

192 **IMPLICATIONS**

193 The comparison of the optical absorption spectra of halite and villiaumite provides a nice
194 illustration of Mie theory, through the dramatic influence of the refractive index of halides on the
195 energy of the plasmon resonance of the embedded metal nanoparticles. Temperatures up to 300-
196 500°C (Zelek et al., 2015; Weerkamp et al., 1994), crystal dislocations and substituted impurities
197 favor the aggregation of color centers leading to the formation of Na colloids. This explains the
198 predominance of these colloids in natural halides (Seifert et al., 1994), as demonstrated by their
199 bright colors. The formation of Na colloids goes with that of free dihalogen molecules, as in the
200 villiaumite from Kola Peninsula, Russia (Celinski et al., 2016) and can only occur once the
201 crystal cooled down to avoid the annealing of these colloids. It is of interest that natural purple
202 fluorites also owe their color to the presence of Ca nanoparticles. Their optical absorption spectra
203 show an intense, broad absorption band near 560 nm that has been suggested to come from Ca
204 metal colloids (Bill and Calas, 1983; Rossman, 2010; Gaft et al., 2020; Ryskin et al., 2020). The
205 same band is found in additively colored synthetic CaF₂ (Angervaks et al., 2018). This
206 absorption band has a Lorentzian shape, which indicates a plasmon resonance origin (Ryskin et
207 al., 2020). Its position, near 560 nm, is intermediate between that in red villiaumite, 494 nm, and
208 blue halite, 640 nm. This absorption band allows light transmission in both the red and blue
209 regions of the spectrum, resulting in the characteristic purple color of irradiated fluorites. It may
210 be pointed out that the position of the colloid band ranges in the order villiaumite (494 nm)-
211 fluorite (560 nm)-halite (640 nm), i.e. the same ranking as for the refractive indices, 1.3253,
212 1.4338 and 1.5442, hence qualitatively following the prediction of Relation (3).

213 Natural halite exhibits a broad range of colors, navy-blue, blue, purple or colorless arising
214 from various proportions of color centers and Na colloids (Weselucha-Birczynska et al., 2012).
215 Zelek et al. (2014) have shown that this may result from the complex sedimentary geology that
216 governs in the mine the spatial distribution of halite and sylvinite, as ^{40}K is suspected to be the
217 main radiation source. In villiaumite, radiation damage is caused by a specific geological context.
218 Aegirite rocks are always enriched in Th and U (Sorensen, 1992). Though XRD only indicates
219 the presence of NaF, preliminary SEM-EDS investigations of our samples (Fig. 3) show the
220 presence of nanospheres, about 500 nm large, which show a preferential alignment that
221 apparently guides the {100} cleavage steps. These nanoinclusions contain about 8.8wt% ThO_2
222 (Supplemental Fig. S6) and may be at the origin of a permanent irradiation of the mineral, able to
223 create defects once the mineral cooled down to a few hundreds °C. This explains the presence of
224 isolated color centers, which otherwise have a limited thermal stability (Schwartz et al., 1994)
225 and the stability of the Na nanoparticles that are at the origin of the red color. The presence of
226 thorium in the villiaumite from the Khibiny and Lovozero alkaline massifs in the Kola Peninsula,
227 Russia (Chukanov et al., 2006) has been explained by its complexation by organic matter, which
228 was revealed by infrared spectroscopy. Preliminary infrared spectra on the villiaumite from the
229 Los Archipelago also reveal the presence of aliphatic hydrocarbon groups and carboxylates in our
230 samples (A. Geisler, unpublished report). The red color of villiaumite, prized in mineral
231 collections, appears to be also a useful "color indicator" of the conditions of emplacement of the
232 evolved stages of peralkaline systems.

233

234

ACKNOWLEDGMENTS

235

236

We are grateful to Jean-Claude Boulliard for providing the samples and their pictures from the mineral collection of Sorbonne-Université, Maxime Guillaumet for his help on the IMPMC

237 spectroscopy platform and Imene Esteve for the preliminary SEM analyses. We thank Charles
238 Geiger and George Rossman for fruitful suggestions that greatly improved the manuscript.

239

240

REFERENCES CITED

241 Amenu-Kpodo, K., and Neubert, T.J. (1965) Color centers in sodium fluoride. Journal of
242 Physics and Chemistry of Solids, 26,1615-1619.

243 Angervaks, A.E., Veniaminov, A.V., Stolyarchuk, M.V., Vasilev, V.E., Kudryavtseva, I.,
244 Fedorov, P.P., Ryskin, A.I. (2018) Optical study of calcium precipitates in additively colored
245 CaF₂ crystals. Journal of the Optical Society of America B, 35, 1288-1294.

246 Arun, T., Ram, S. S., Karthikeyan, B., Ranjith, P., Ray, D. K., Rout, B., Krishna, J. B. M.,
247 Sengupta, P., and Satyam, P.V. (2017) Ion beam radiation effects on natural halite crystals.
248 Nuclear Instruments and Methods B, 409, 216-220.

249 Bill, H., and Calas G. (1978) Color centers, associated rare-earth ions and the origin of
250 coloration in natural fluorites. Physics and Chemistry of Minerals, 3, 117-131.

251 Boizot, B., Petite, G., Ghaleb, D. Pellerin, N., Fayon, F., Reynard, B., and Calas, G. (2000)
252 Migration and segregation of sodium under β -irradiation in nuclear glasses. Nuclear Instruments
253 and Methods B, 166, 500–504.

254 Bryukvina, L., Ivanov, N., and Nebogin, S. (2018) Relationships between lithium and sodium
255 nanoparticles and color centers formation in LiF and NaF crystals with hydroxide and
256 magnesium ions impurities. Journal of Physics and Chemistry of Solids, 120,133-139.

257 Bryukvina, L.I., and Martynovich, E.F. (2012) Formation and properties of metallic
258 nanoparticles in lithium and sodium fluorides with radiation-induced color centers. Physics of the
259 Solid State, 54, 2374-2378.

- 260 Celinski, V.R., Ditter, M., Kraus F., Fujara F., and Schmedt auf der Günne, J. (2016) Trace
261 determination and pressure estimation of fluorine F₂ caused by irradiation damage in minerals
262 and synthetic fluorides. *Chemistry-A European Journal*, 22,18388–18393.
- 263 Chan Q. H. S., Zolensky M. E., Kebukawa Y., Fries M., Ito M., Steele A., Rahman Z., Nakato
264 A., Kilcoyne A. L. D., Suga H., Takahashi Y., Takeichi Y., and Mase K. (2018) Organic matter
265 in extraterrestrial water-bearing salt crystals. *Science Advances* 4, eaao3521.
- 266 Chandra, A., and Holcomb, D.F. (1969) Taxonomy of F-aggregate centers in NaF. *Journal of*
267 *Chemical Physics*, 51, 1509–1523.
- 268 Chukanov, N.V., Pekov, I.V., Sokolov, S.V., Nekrasov, A.N., Ermolaeva, V.N., and Naumova,
269 I.S. (2006) On the problem of the formation and geochemical role of bituminous matter in
270 pegmatites of the Khibiny and Lovozero alkaline massifs, Kola Peninsula, Russia. *Geochemistry*
271 *International*, 44, 715-728.
- 272 Davenas, J., Perez, A., Thevenard, P., and Dupuy, C.H.S. (1973) Correlation between
273 absorption bands and implanted alkali ions in LiF. *Physica Status Solidi A*, 19, 679–686
- 274 Davidson, A.T., Kozakiewicz, A.G., Comins, J.D., and Derry, T.E. (1996) Optical effects in
275 NaF crystals implanted with 100 keV ions. *Nuclear Instruments and Methods B*, 116, 216-219.
- 276 Doyle, W.T. (1958) Absorption of light by colloids in alkali halide crystals. *Physical Review*,
277 111, 1067-1072.
- 278 Doyle, W.T. (1960) Coagulation, optical absorption and photoconductivity of colloid centres
279 in alkali halides. *Proceedings of the Physical Society*, 75, 649–663.
- 280 Gaft, M., Waychunas, G.A., Rossman, G.R., Nagli, L., Panczer, G., Cheskis, D., and Raichlin,
281 Y. (2020) Red photoluminescence and purple color of naturally irradiated fluorite. *Physics and*
282 *Chemistry of Minerals*, 47, 46.

- 283 Howard, C.L.H., and Kerr, P.F. (1960) Blue halite. *Science*, 132,1886-1887.
- 284 Hoya, J., Laborde, J.I., Richard, D., and Renteria, M. (2017) Ab initio study of F-centers in
285 alkali halides. *Computational Materials Science*, 139, 1-7.
- 286 Hunault, M., Loisel, C., Bauchau, F., Lemasson, Q., Pacheco, C., Pichon, C., Moignard, B.,
287 Boulanger, K., Hérold, M., Calas, G., and Pallot-Frossardt, I. (2017) Nondestructive redox
288 quantification reveals glassmaking of rare French gothic stained glasses. *Analytical Chemistry*,
289 89, 6277-6284.
- 290 Ivey, H.F. (1947) Spectral location of the absorption due to color centers in alkali halide
291 crystals. *Physical Review*, 72, 341–343.
- 292 Kikuchi, A. and Ozawa K. (1973) Induced F aggregate centres in NaF crystals irradiated with
293 charged particles at low temperature. *Physica Status Solidi A*, 18, 567-578.
- 294 Kreibig, U., and Vollmer, M. (1995) Optical properties of metal clusters. Springer:
295 Berlin/Heidelberg.
- 296 Lacroix, A. (1908) Sur l'existence du fluorure de sodium cristallisé comme élément des
297 syénites néphéliniques des îles de Los. *Comptes Rendus Hebdomadaires des Séances de*
298 *l'Académie des Sciences*, 146, 213-216.
- 299 Levy, P. W., Loman, J. M., Swyler, K. J., and Klaffty, R. W. (1983) Radiation damage studies
300 on synthetic NaCl crystals and natural rock salt for radioactive waste disposal applications. In:
301 *Scientific Basis for Nuclear Waste Management*, Ed. P.L. Hofman, 136–167.
- 302 Marks, M.A.W., and Markl, G. (2017) A global review on agpaitic rocks. *Earth-Science*
303 *Reviews*, 173, 229–258.
- 304 Moreau, C., Ohnenstetter, D., Demaiffe, D., and Robineau, B. (1998) The Los Archipelago
305 nepheline syenite ring-structure: a magmatic marker of the evolution of the Central and
306 Equatorial Atlantic. *Canadian Mineralogist*, 34, 281–299.

- 307 Okada, T. (2006) Nano-scale particles of mixed alkali-metals in KCl crystals. Journal of the
308 Physical Society of Japan, 75, 094705.
- 309 Poston, M.J., Carlson, R.W., and Hand, K.P. (2017) Spectral behavior of irradiated sodium
310 chloride crystals under Europa-like conditions. Journal of Geophysical Research, Planets, 122,
311 2644–2654.
- 312 Przibram K. (1953) Verfärbung und Lumineszenz. Beiträge zur Mineralphysik. Springer-
313 Verlag, Vienna.
- 314 Rossman G.R. (2010) Mineral spectroscopy server. Caltech. [https://doi.org/10.7907/JYWR-](https://doi.org/10.7907/JYWR-QQ57)
315 [QQ57](https://doi.org/10.7907/JYWR-QQ57).
- 316 Ruiz-Fuertes, J., Ibáñez, J., Monteseguro, V., Alencar, I., and Cazorla, C. (2019). Reversible
317 tuning of Ca nanoparticles embedded in a superionic CaF₂ matrix. The Journal of Physical
318 Chemistry C, 123, 19945–19951.
- 319 Ryskin, A.I., Fedorov, P.P., Lushchik, A., Generalov, M.E., Angervaks, A.E., Stolyarchuk,
320 M.V., Vasil'chenko, E., and Kudryavtseva, I. (2020) Absorption spectrum of dark purple fluorite,
321 Kent deposit, Kazakhstan. Journal of Fluorine Chemistry, 240, 109654.
- 322 Schwartz, K., Volkov, A.E., Sorokin, M.V., Trautmann, V.C., Voss, K.O., Neumann, R., and
323 Lang, M. (2008) Effect of electronic energy loss and irradiation temperature on color-center
324 creation in LiF and NaCl crystals irradiated with swift heavy ions. Physical Review B, 78,
325 024120.
- 326 Seifert, N., Ye, H., Tolk, N., Husinsky, W., and Betz G. (1994) The influence of defects and
327 defect clusters on alkali atom desorption stimulated by low energy electron bombardment of
328 alkali halides. Nuclear Instruments and Methods B, 84, 77-88.
- 329 Selby K., Kresin, V., Masui, J., Vollmer, M., de Heer, W.A., Scheidemann, A., and Knight,
330 W.D. (1991) Photoabsorption spectra of sodium clusters. Physical Review B, 43,4565-4572.

- 331 Sonnenfeld, P. (1995) The color of rock salt -a review. *Sedimentary Geology*, 94, 267-276.
- 332 Sørensen, H. (1992) Agpaitic nepheline syenites—a potential source of rare elements. *Applied*
333 *Geochemistry*, 7, 417-427.
- 334 Stormer, J.C., and Carmichael, I. (1970) Villiaumite and the occurrence of fluoride minerals in
335 igneous rocks. *American Mineralogist*, 55, 126-134.
- 336 Tiwald, P., Karsai, F., Laskowski, R., Gräfe, S., Blaha, P., Burgdörfer, J., and Wirtz, L.
337 (2015) Ab initio perspective on the Mollwo-Ivey relation for F centers in alkali halides. *Physical*
338 *Review B*, 92, 144107.
- 339 Weerkamp, J.R. W., Groote, J.C., Seinen, J., and Den Hartog, H.W. (1994) Radiation damage
340 in NaCl. I. Optical-absorption experiments on heavily irradiated samples. *Physical Review B*, 50,
341 9781– 9786.
- 342 Weselucha-Birczynska, A., Zelek, S., and Stadnicka, K. (2012) Blue halite colour centre
343 aggregates studied by micro-Raman spectroscopy and X-ray diffraction. *Vibrational*
344 *Spectroscopy*, 60 124–128.
- 345 Zelek, S.M., Stadnicka, K.M., Tobola, T., and Nowak, L.N. (2014) Lattice deformation of blue
346 halite from Zechstein evaporate basin: Klodawa Salt Mine, Central Poland. *Mineralogy and*
347 *Petrology*, 108, 619-631.
- 348 Zelek, S.M., Weselucha-Birczyńska, A., Szklarzewicz, J., and Stadnicka, K.M. (2015)
349 Spectroscopic properties of halite from Klodawa salt mine, Central Poland. *Mineralogy and*
350 *Petrology*, 109, 45–51.

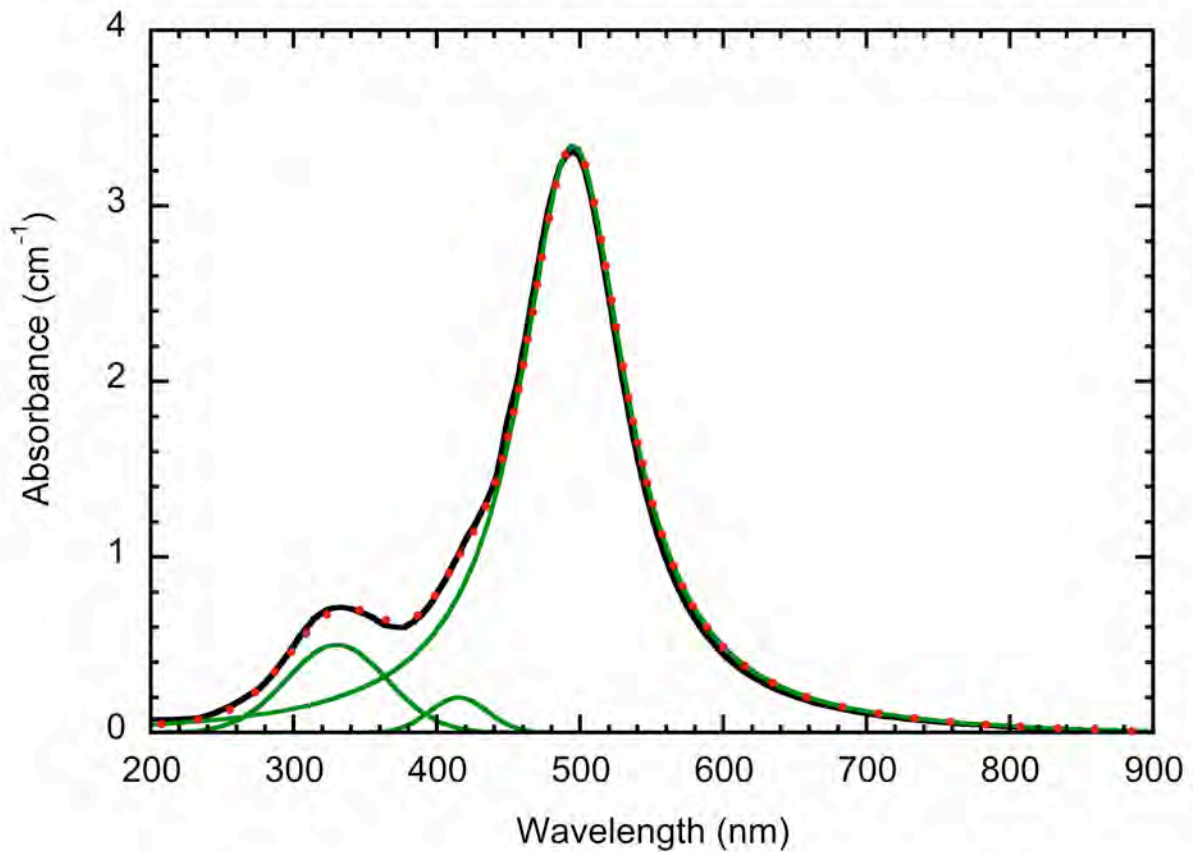


Figure 1. Room temperature optical absorption spectrum of red villiaumite from nepheline syenites of the Los Archipelago, Guinea. The optical spectrum has been background corrected for light scattering caused by crystal inhomogeneities. The main band, assigned to surface plasmon resonance for metallic Na, shifts by about 140 nm in halite relative to villiaumite. As a consequence, light is transmitted in the long wavelength side of the main absorption band. The fit uses a Lorentzian function for the surface plasmon resonance (main band) and Gaussian components for the minority color centers. The functions used for the fit are displayed in green. The resulting fit gives the red dots.

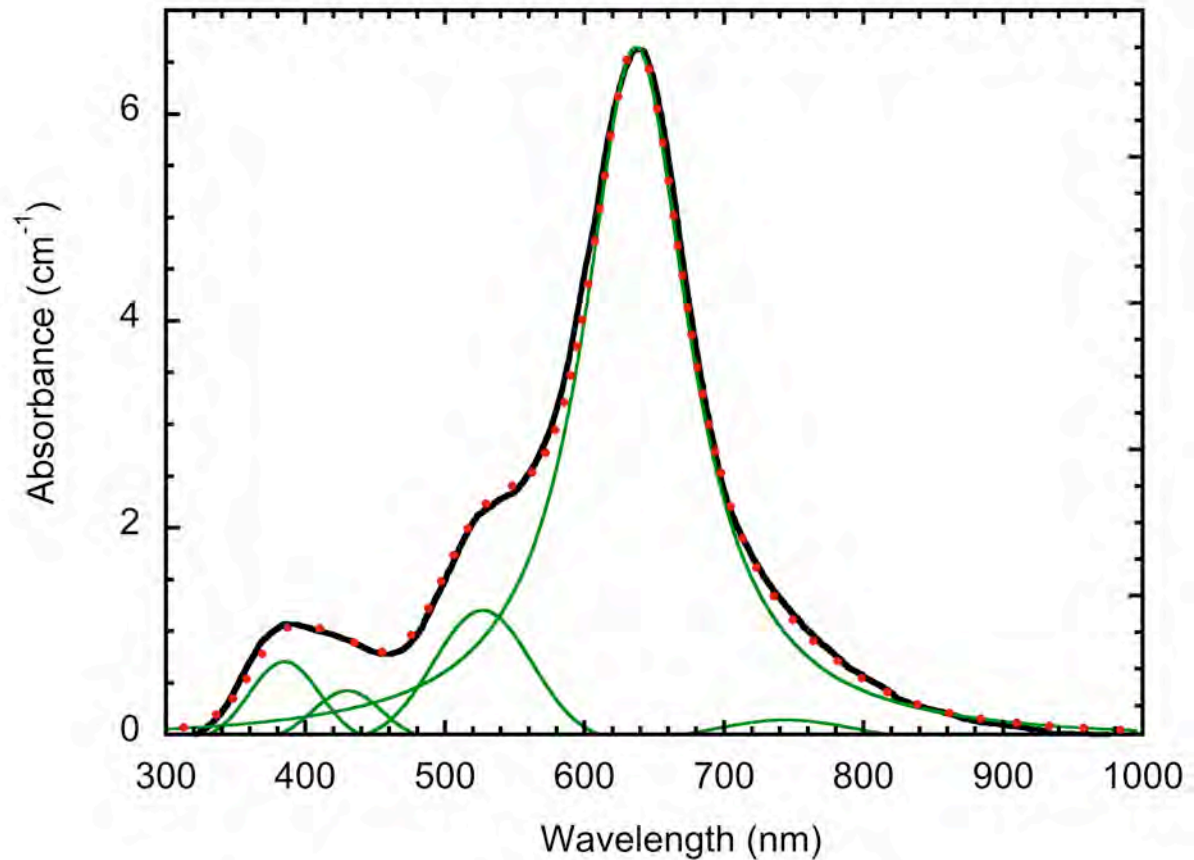


Figure 2. Room temperature optical absorption spectrum of blue halite from Stassfurt, Germany, background corrected for light scattering by crystal inhomogeneities. The main band shifts by about 140 nm relative to villiaumite. As a consequence, light is transmitted in the short wavelength side of the main absorption band, hence the spectacular color difference. The fit uses a Lorentzian function for the surface plasmon resonance (main band) and Gaussian components for the minority transient color centers. The functions used for the fit are displayed in green. The resulting fit gives the red dots.

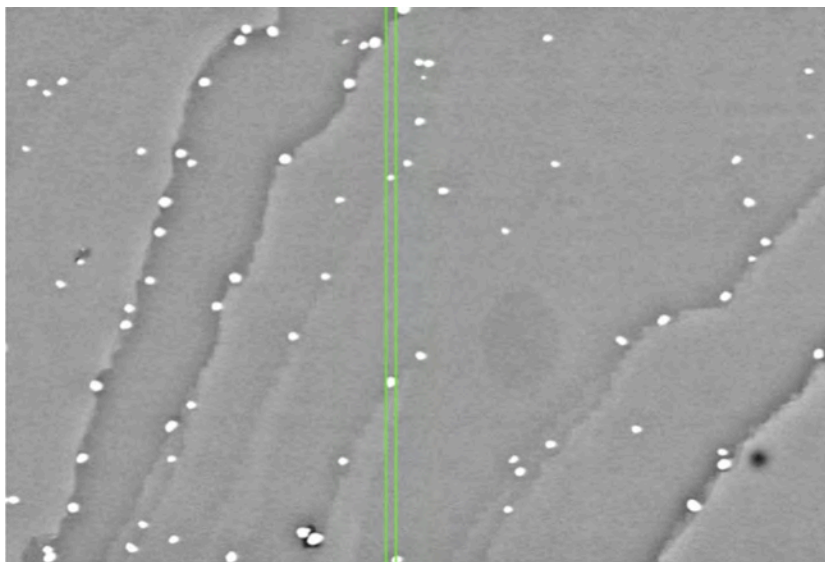


Figure 3. Scanning Electron Microscopy micrograph in backscattered electron mode of a villiaumite cleavage, showing the presence of thorium-rich nano-inclusions. The scale is given by the green bars, separated by 500 nm. EDS analysis of these nano-inclusions (Supplemental information, Fig. S6) shows that thorium is only accompanied by minor amounts of uranium. The apparent alignment of these inclusions inherits from the growth of the mineral, in which they guide the $\{100\}$ cleavage steps. These solid inclusions are thought to cause a permanent internal irradiation of the mineral. Once the villiaumite-bearing rock cools down to low temperatures, typically below 400°C, this irradiation will progressively cause the formation of radiation-induced color centers and metallic Na nanoparticles at the origin of the intense red coloration, characteristic of villiaumite.

REVISION 1

Supplemental Material



Figure S1. Blue halite (Sample # N00341_AJM4313) from Stassfurt, Germany (Mineral Collection, Sorbonne Université, Paris).



Figure S2. Red villiaumite (Sample # N13100_AJM5773) from the Los Islands Archipelago, Guinea (Mineral Collection, Sorbonne Université, Paris).



Figure S3. Cleavage section of euhedral crystals of villiaumite from the Los Islands Archipelago, Guinea.

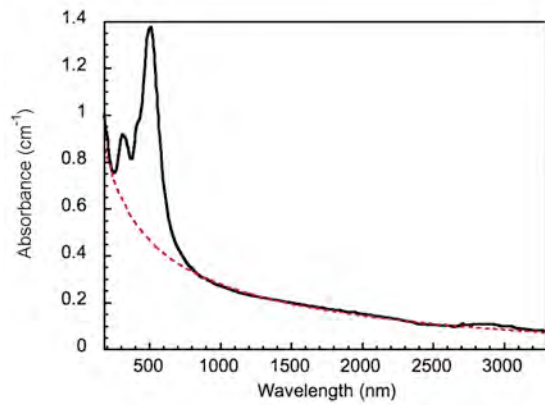


Figure S4. Room temperature optical absorption spectrum of villiaumite. The dotted red line corresponds to the background absorption due to a scattering of the incident light by optical heterogeneities (inclusions and fractures)

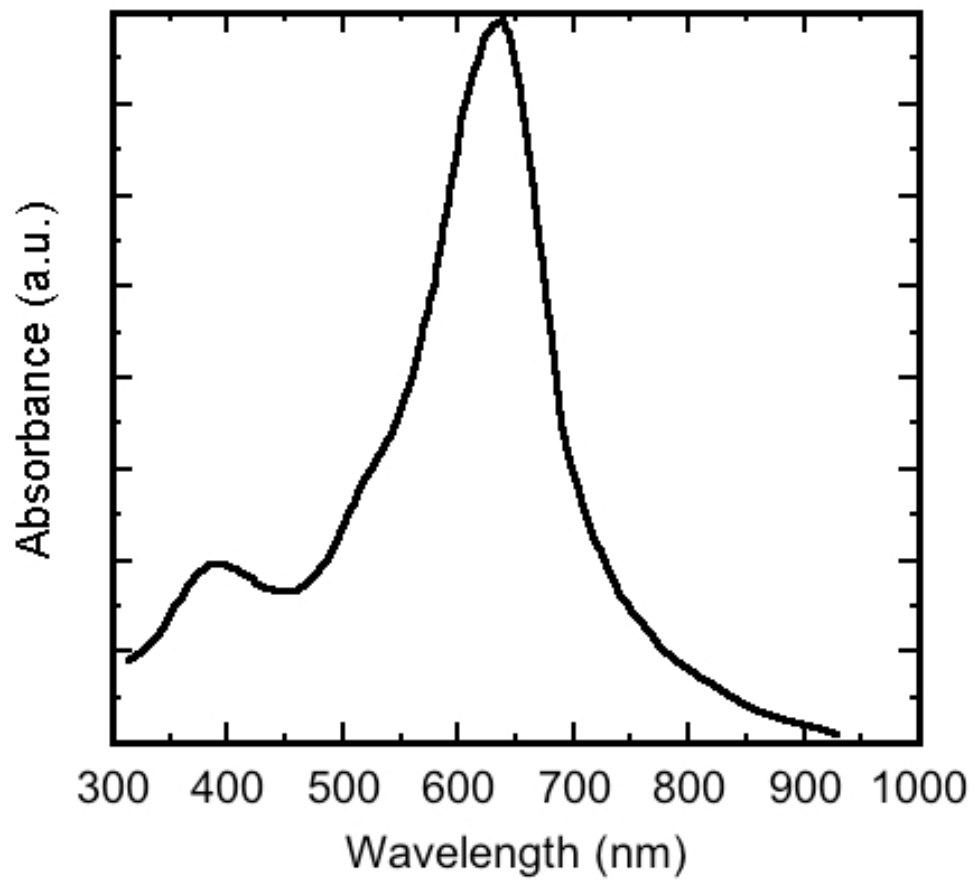


Figure S5. Scan of the historical absorption spectrum of the natural blue Stassfurt halite published by Doyle in 1960. The main absorption band peaks at 640 nm, a value consistent with that reported in this study.

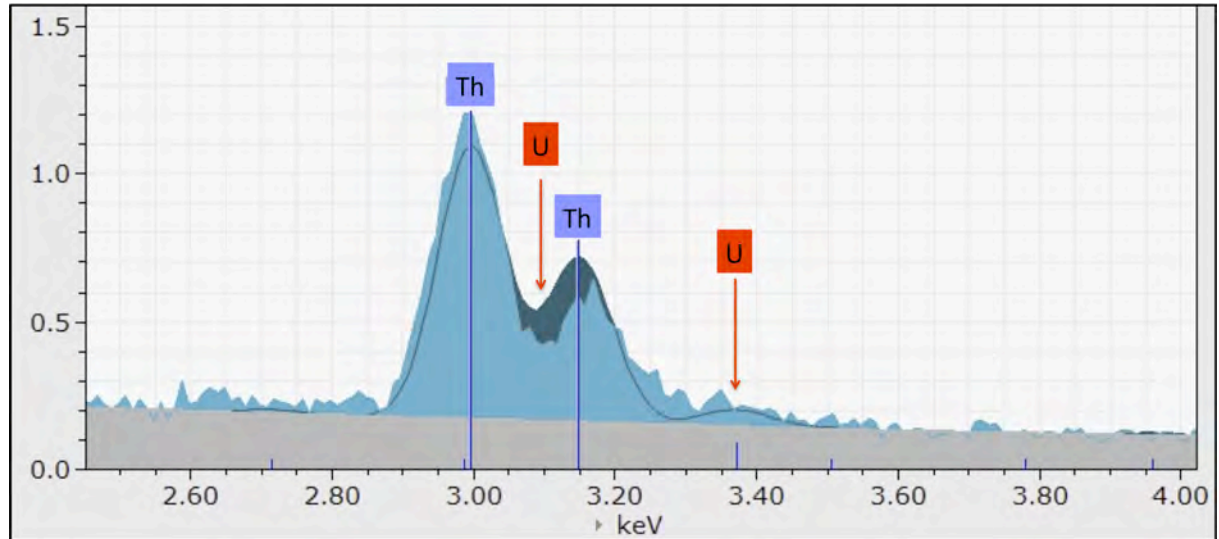


Figure S6. SEM/energy-dispersive X-ray spectroscopy of the nanoinclusions, about 500 nm large, encountered in the villiaumite from the Los Archipelago and depicted in Figure 2. The Th concentration has been calculated using the fit of the Th-M α_1 and Th-M β_1 signals, at 2.996 and 3.149 keV, respectively. The weak contribution at 3.35-3.40 keV may arise from the U-M β_1 signal, expected at 3.336 (the signal expected for U-M α_1 at 3.171 keV is hidden by the more intense Th-M β_1 emission).

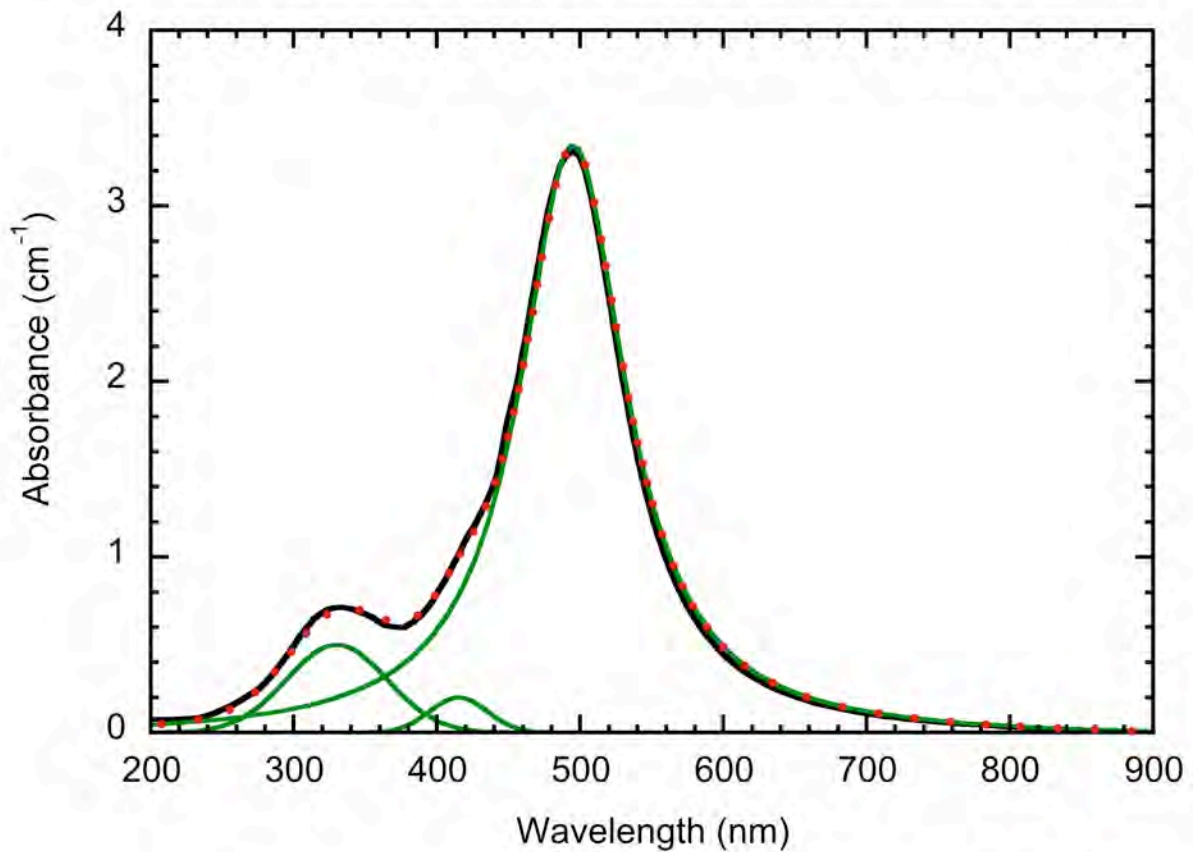


Figure 1. Room temperature optical absorption spectrum of red villiaumite from nepheline syenites of the Los Archipelago, Guinea. The optical spectrum has been background corrected for light scattering caused by crystal inhomogeneities. The main band, assigned to surface plasmon resonance for metallic Na, shifts by about 140 nm in halite relative to villiaumite. As a consequence, light is transmitted in the long wavelength side of the main absorption band. The fit uses a Lorentzian function for the surface plasmon resonance (main band) and Gaussian components for the minority color centers. The functions used for the fit are displayed in green. The resulting fit gives the red dots.

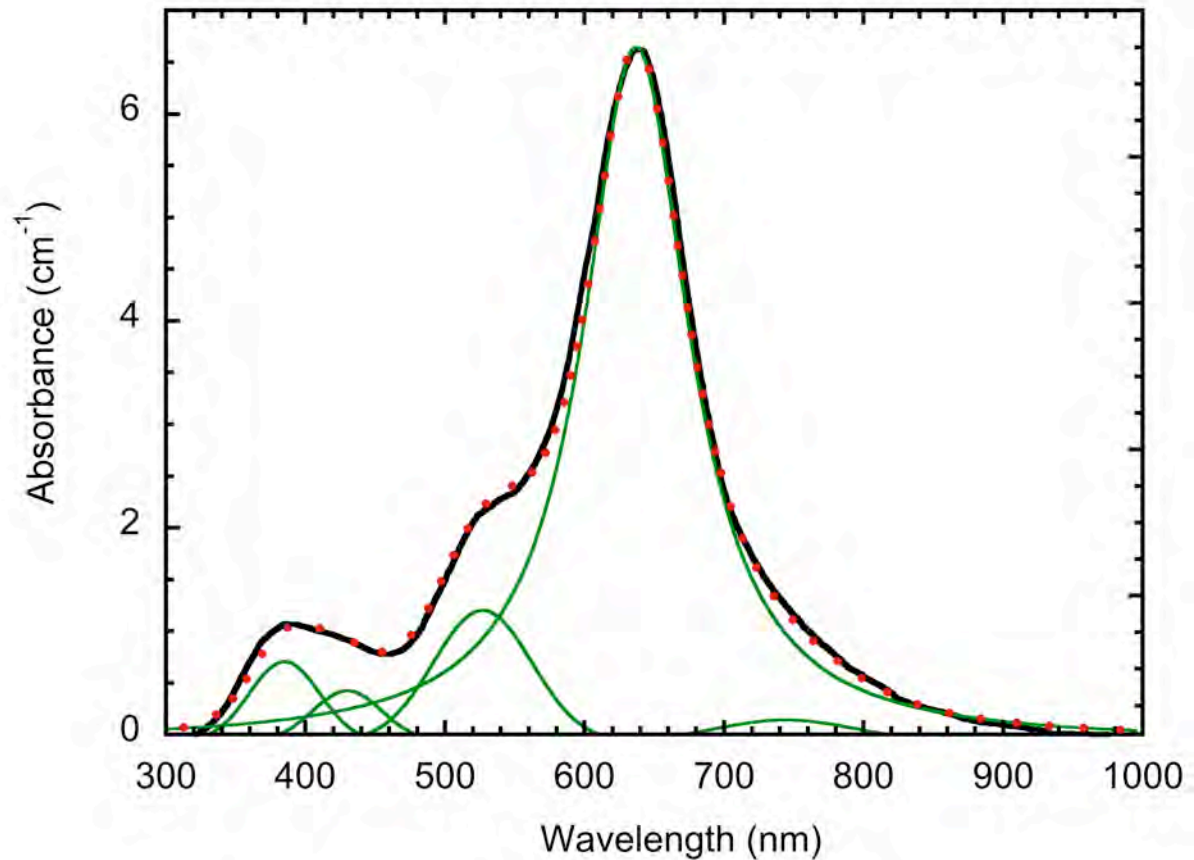


Figure 2. Room temperature optical absorption spectrum of blue halite from Stassfurt, Germany, background corrected for light scattering by crystal inhomogeneities. The main band shifts by about 140 nm relative to villiaumite. As a consequence, light is transmitted in the short wavelength side of the main absorption band, hence the spectacular color difference. The fit uses a Lorentzian function for the surface plasmon resonance (main band) and Gaussian components for the minority transient color centers. The functions used for the fit are displayed in green. The resulting fit gives the red dots.

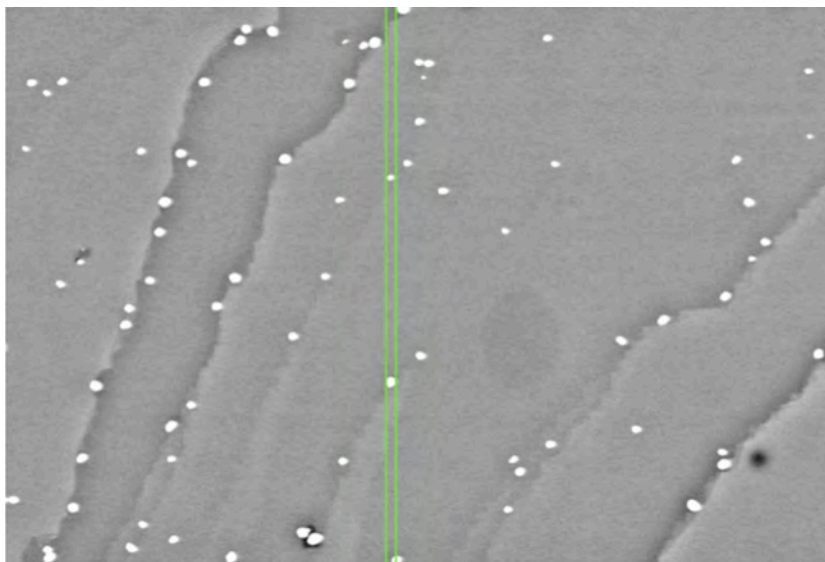


Figure 3. Scanning Electron Microscopy micrograph in backscattered electron mode of a villiaumite cleavage, showing the presence of thorium-rich nano-inclusions. The scale is given by the green bars, separated by 500 nm. EDS analysis of these nano-inclusions (Supplemental information, Fig. S6) shows that thorium is only accompanied by minor amounts of uranium. The apparent alignment of these inclusions inherits from the growth of the mineral, in which they guide the $\{100\}$ cleavage steps. These solid inclusions are thought to cause a permanent internal irradiation of the mineral. Once the villiaumite-bearing rock cools down to low temperatures, typically below 400°C, this irradiation will progressively cause the formation of radiation-induced color centers and metallic Na nanoparticles at the origin of the intense red coloration, characteristic of villiaumite.