Revision 2

The glass transition and the non-Arrhenian viscosity of carbonate melts.

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Abstract
We report the first calorimetric observation of the glass transition for a carbonate melt. A carbonate glass (55K$_2$CO$_3$-45MgCO$_3$ (molar)) was quenched from 780 °C at 0.1 GPa. The activation energy of structural relaxation close to the glass transition was derived through a series of thermal treatments comprising excursions across the glass transition at different heating rates. Viscosities just above the glass transition temperature were obtained by applying a shift factor to the calorimetric results. These viscosity measurements (in the range of $10^9$ Pa*s) at supercooled temperatures (ca. 230 °C), dramatically extend the temperature range of data for carbonates which were previously restricted to superliquidus viscosities well below 1 Pa*s. Combining our calorimetrically derived results with published alkaline-earth carbonate melt viscosities at high temperatures yields a highly non-Arrhenian viscosity-temperature relationship and confirms that carbonate liquids are “fragile”. Based on simulations, fragile behavior is also exhibited by Na$_2$CO$_3$ melt. In both cases the fragility presumably relates to the formation of temperature-dependent low dimensional structures and Vogel-Fulcher-Tammann (VFT) curves adequately describe the viscosity-temperature relationships of carbonate melts below 1000 °C.

Key words: Viscosity, glass transition, fragility, carbonate melt, carbonatite, shift factor

1. Introduction
Melt viscosity has a fundamental control on many Earth processes, from how melts are transported through the deep mantle and crust, to its influence on effusive or explosive eruption styles at the Earth surface (McKenzie 1985; Dingwell 1996). Although physical properties (e.g., viscosity, density, surface tension) are becoming well-established for multicomponent silicate systems of importance in earth sciences (Lange and Carmichael 1987; Knoche et al., 1995;
Bagdassarov et al., 2000; Giordano et al., 2008), carbonate melts are less well-studied in the geosciences, due to relatively rare volcanic expression at the Earth’s surface (Keller, 1989; Dawson, 1966). Carbonate volcanism is nevertheless a part of the Earth System and may be more common on other planets. Carbonatite volcanism has been invoked for so-called “canali” - channels on Venus which reach up to 6800 km in length, implying a very low viscosity flow (Kargel et al., 1994). Highly mobile carbonatite melts migrating within the Earth likely play a major role as metasomatizing agents (Blundy and Dalton, 2000; Hammouda and Laporte, 2000) and in the subduction-related recycling of carbon back to Earth’s surface as part of the global carbon cycle (Dasgupta and Hirschmann, 2010; Thomson et al., 2016). Their ability to scavenge high concentrations of REE and rare metals make them increasingly important as multi-commodity exploration targets for the green energy revolution (Simandl and Paradis, 2018). The extreme physical and chemical properties (e.g. density, viscosity, conductivity, reactivity) which set carbonatites apart from most silicate melts, must have a strong influence on the interaction and migration dynamics of buoyant carbonatite melts throughout the lithosphere after separating from their source region (Hunter and McKenzie, 1989; Minarik and Watson, 1995; Brooker, 1998; Hammouda and Laporte, 2000; O’Leary et al., 2015). The relationship between low viscosity and electrical properties of carbonate-rich melts has also been linked to regions of high conductivity in the upper and lower mantle (Sifré et al., 2014; Gaillard et al., 2018).

These physical properties also have commercial applications particularly in the design of molten carbonate fuel cells which generally offer a prospect of greener electricity, but may now also be designed with a capacity to efficiently capture CO\textsubscript{2} derived from the liquid, gaseous or solid hydrocarbon (i.e., coal) fuel used for generating electricity (Cassir et al., 2012). Such molten carbonates generally provide excellent performance at low cost in thermal energy storage or heat transfer fluid systems (Wu et al., 2011).
Recent investigations of carbonate melt at the molecular scale have revealed a temperature- and pressure-dependence of structure that likely affects physical properties (Wilding et al. 2016; Wilson et al. 2018; Wilding et al. 2019a, 2019b). Thus, a more complete understanding of the relationship between carbonate melt structure and physical properties is required. Here, we provide the first ever calorimetric determination of the glass transition of a carbonate glass, and generate a direct measurement of the viscosity of a carbonate liquid just above the glass transition interval. In this manner we are able to constrain the non-Arrhenian temperature-dependence and thereby the fragility of molten carbonates, providing a parameterization valuable for the extrapolation and interpolation in temperature of the transport properties of carbonates.

1.1 Viscosity Data – the carbonate challenge

Natural carbonatites are mainly composed of alkaline-earth (Ca, Mg and Fe) alkali (Na and K) carbonates. Rheological measurements on the most geologically relevant end-member carbonates have proven difficult. This is especially true for alkaline earth carbonates as they do not melt congruently at atmospheric pressure (1 atm). For example, CaCO$_3$ decarbonates to CaO + CO$_2$ below 0.004 GPa and 1240 °C and then melts incongruently (to CaCO$_3$ + CaO or CaCO$_3$ + CO$_2$ rich melts) up to 0.3 GPa and 1375 °C. Only above this pressure and temperature is congruent melting observed (Smyth and Adams 1923; Wyllie and Huang 1976). For MgCO$_3$ and FeCO$_3$ the conditions are even more extreme (see Wyllie and Huang 1976; Weidner 1972). Viscosity measurements are less complicated for the alkali carbonates, as they melt congruently at atmospheric pressure; although they can be subject to dissociative CO$_2$ loss whose intensity decreases in the order Li >> Na > K (Janz and Lorenz 1961).

In Figure 1a, the ambient pressure, single alkali carbonate measurements imply a clear compositional effect of the different alkali cations, but large scatter also exists even within a
single study. As is typically observed for very restricted ranges of temperature and viscosity, all
data (except (L,N,K)C ) can be described within error by an Arrhenian temperature dependence
(a linear fit on a log_{10} \eta vs. 1/T plot). As later noted by Janz et al. (1989) their initial ‘NIST
standard’ viscosity data were incorrect (and therefore are not included in Fig. 1a). For Na_2CO_3
they refer to the data of Ejima et al. (1984) which has a less steep slope, almost identical to the
Sato et al. (1999) data in Figure 1a. Janz et al. (1989) suggest that the error and differences in
general, arise from comparisons of various measurement methods, experimental assemblies, and
working equations, as well as matters of sample pre-treatments and impurities. Also included in
Figure 1a are green stars, representing new 50Na_2CO_3-50K_2CO_3 data from this study (see
Results), measured using the same equipment as described by Di Genova et al. (2016). This data
set covers a larger temperature range and so has a better-defined slope (green dotted line). The
higher pressure Na_2CO_3 data of Stagno et al. (2018) in Figure 1a perhaps illustrates the challenge
of performing accurate high-pressure measurements and it remains unclear whether they define
any systematic effect of pressure. Also included in Figure 1a are the data from Kim et al. (2015)
for a 22Li_2CO_3-33Na_2CO_3-45K_2CO_3 composition (purple stars). This eutectic in the Li_2CO_3-
bearing ternary system allows an even greater temperature range in measured viscosities. With
such a large temperature range considered, some curvature clearly becomes evident, indicating a
non-Arrhenian temperature dependence and raising questions regarding the validity of a linear fit
to the other data.

Figure 1b contains a much more diverse range of mainly mixed carbonate compositions.
Many involve alkaline-earth cations and therefore require pressure to ensure congruent melting. It
is difficult to untangle the effect of composition from that of pressure, but the Dobson et al.
(1996) study appear to include a range of data points on single lines for two compositional
datasets covering a range of pressures from 2.5 to 5.5 GPa. These two lines generally bracket the
different compositions of Kono et al. (2014) which cover an even larger pressure range from 0.9 to 5.3 GPa. In terms of composition, the green dotted line is representative of the alkali carbonates at 1 atm, when transferred to Figure 1b and could indicate either a lower viscosity for the alkali carbonates or some pressure effect with respect to the other high-temperature data. The lower viscosity of high-pressure measurements for Na$_2$CO$_3$ (yellow hexagons) compared to other compositions, but above the 1 atm data (green dotted line), appear to suggest both effects may be operating. Note, that the K$_2$CO$_3$ high pressure (4.0 GPa) data point of Dobson et al. (1996) was subsequently discredited (Liu et al. 2007, and therefore are not included in Fig. 1b), again illustrating the challenges in measuring physical properties using high pressure experiments (Kono et al., 2014; Stagno et al. 2018). All the other high temperature data in Figure 1b indicate the effect of pressure is less than half a log$_{10}$ unit between 0.9 and 5.5 GPa. The measured data appears at odds with the simulations of Vuilleumier et al. (2014), Du et al. (2018) and Desmaele et al. (2019a) which suggest a clear and more pronounced pressure effect. The seminal CaCO$_3$ data of Vuilleumier et al. (2014) is included in Figure 1b as an example, although these simulations may have over-estimated the pressure effect to some extent and the role of forces such as van der Waals interactions require further refinement for these ab initio calculations (Vuilleumier Pers. Comm.)

1.2 The Glass Transition and Viscosity of ‘Fragile’ Melts

As an aid to resolving some of the issues described above, an extension of the temperature range for viscosity determinations of carbonate liquids is a prime experimental goal. Most commercial glasses involve liquids with ‘covalent’ network-forming components such as silicate, aluminate or borate, and when cooled their increase in viscosity hinders nucleation and results in a glass with a characteristic, calorimetric glass transition reflecting this kinetic arrest (Moynihan 2019).
Glass formation is however not restricted to network-forming liquids (Tangeman et al. 2001; Dingwell et al., 2004: Kohara et al. 2011). Glasses can even be formed from organic liquids and metallic systems. Of possible relevance to this study are glasses formed in simple ionic molten salts such as sulphate (Förland and Weyl 1950; MacFarlane 1984) and nitrate systems (Van Uitert and Grodkiewicz 1971). For these ionic compositions, the temperature-dependence of viscosity, a vital parameter in glass-formation, rarely exhibits Arrhenian behavior. This departure from Arrhenius law behavior is at the heart of the concept of liquid “fragility”, a term introduced by Angell (1985) to describe and compare the viscosity-temperature relations of a variety of glass-forming liquids. The quantification of fragility has been provided by the model-independent, so-called melt fragility index $m$ (Böhmer et al. 1993). This index describes the rate of viscosity change of a melt with temperature at $T_g$, defined as the temperature where the viscosity is equal to $10^{12}$ Pa*s. The exact mechanism is thought to be related to medium range structural heterogeneity with faster and slower relaxing regions developing as liquids are supercooled. There are formal links to the changing in structure of supercooled liquids via the Adam-Gibbs model of structural relaxation (Gibbs and Adam 1995, Angell et al. 2000b, 2000a). Viscosity measurements made close to the glass transition can be used to evaluate liquid fragility (Gottsmann et al. 2002).

The linear Arrhenius law for viscosity ($\eta$) is expressed as:

$$\log_{10} \eta = A + \frac{E_a}{T}$$

and has two parameters to be fitted, $E_a$ the activation energy is considered as the energy barrier that must overcome for examples as bonds are broken allowing molecules to ‘flow’ from one position to another, and $A$ is a numerical constant.
Thus, although viscosity-temperature data such as those in Figure 1 are commonly approximated over limited temperature ranges to an Arrhenian fit, it has long been known that for many materials the value of $E_a$ changes with temperature yielding a non-Arrhenian curvature on a reciprocal absolute (Arrhenian) plot (Angell 1985; Böhmer et al. 1993). This is particularly evident as data are added from lower (supercooled) temperatures. One method of fitting such curvature is using the Vogel- Fulcher-Tammann equation (Vogel 1921; Fulcher 1925; Rault 2000) equation of the form:

$$\log_{10} \eta = A + \frac{B}{T(K) - C} \quad (2)$$

with an extra fitting parameter $C$ known at the ‘Vogel temperature’, which is non-zero for a non-Arrhenian viscosity-temperature relationship. This has the potential to reconcile a steep reciprocal temperature-dependence of low temperature viscosity data with a lower reciprocal temperature dependence of high temperature viscosity data.

1.3 Carbonate Glasses

At the accessible quench rates of most experimental equipment (100s-1000s K/s), carbonate melts rarely survive undercooling sufficient to produce glasses. As a result, direct investigation of their physico-chemical properties in the supercooled regime is currently limited to just the few known systems where glasses can be readily formed, MgCO$_3$-K$_2$CO$_3$ (Eitel and Skaliks 1929; Datta et al. 1964; Ragone 1966; Genge et al. 1995) and La(OH)$_3$-Ca(OH)$_2$-CaCO$_3$-CaF$_2$-BaSO$_4$ (Jones and Wyllie 1983; Genge et al. 1995). Eitel and Skaliks (1929) were apparently the first to report a carbonate glass, quenched from ~50:50 mol% MgCO$_3$-K$_2$CO$_3$ melts at 0.12 GPa. They also observed rapid devitrification of the glasses on an hour-timescale when held at a temperature between 205-300 °C. This is a common response to crossing the glass transition within poorly
glass-forming systems such as nitrates, sulphates and carbonates and even basaltic silicates (Böhmer et al. 1993; Angell et al. 2000a; Wilding et al. 2000; Nichols et al. 2009). This implies the presence of Tg in this region of temperature and that the glass transition temperature for K₂Mg(CO₃)₂ might be located near 200 °C. To the best of our knowledge this estimate has not been confirmed or used in the intervening 90 years, although Datta et al. (1964) did return to the system and mapped the glass-forming region.

Here, we report the first quantitative determination of a glass transition for a carbonate liquid. Specifically, we demonstrate the nature of the glass transition and viscosity of a mixed K₂CO₃-MgCO₃ carbonate melt by performing scanning calorimetry and viscometry on samples of a supercooled carbonate phase. Using this low-temperature viscosity data combined with published higher temperature (high pressure) data as well as both classical and ab initio molecular dynamics simulation estimates, we determine whether a non-Arrhenian Vogel- Fulcher-Tammann (VFT) law fit is possible and quantify the extent of fragility.

2. Materials and Methods

2.1 Glass synthesis: Carbonate glass were prepared using a starting mixture of 55 mol% K₂CO₃ and 45 mol% MgCO₃ (subsequently referred to as 55K₂CO₃-45MgCO₃) using a rapid quench, cold seal pressure vessel at University of Bristol. This composition is directly above the eutectic (at ~460 °C) on the binary join (Ragone 1966) and glasses can be formed easily at the quench rates achieved (~200 K/s) at a pressure of 0.1 GPa. Reagent grade K₂CO₃ (>99.9%) was dried at 400 °C, the MgCO₃ was in the form of a natural, optically clear and inclusion free magnesite crystal (Brumado, Brazil) and transmission FTIR was used on thin cleavage fragments to confirm that this material was virtually water-free. The starting materials were ground together and loaded in 3.8mm diameter, 20mm long gold capsules which were then welded shut and loaded into a...
Tuttle-type cold-seal pressure vessel with a rapid quench rod extension (Ihinger 1991). The experiments were run at 780 °C, 0.1 GPa for ~10-15 hrs. and quenched (>200 K/s). The resulting glass was removed from the gold capsule mostly as a single solid slug, representing the central part of the quenched melt, with an outer section that tends to spall off. This was stored in a desiccator. A small section was immediately set in dental resin and polished using dry SiC paper.

Attenuated Total Reflectance infrared spectroscopy (ATR) collected on this and similar samples indicate that some water may be present (see Wilding et al., 2019b). This was also confirmed by $^1$H NMR on similar samples (Wilding et al., 2019b). Wilding et al. (2019b) suggest this water is most likely introduced during loading or because the powders were improperly dried (in the case of the $^{13}$C enriched NMR sample). A comparison with the IR spectra of Genge et al. (1995) suggests our samples contain significantly less water although it is very difficult to quantify.

There is no evidence to suggest that the water content of the glass changes over the course of at least 24 hours as demonstrated by a time series of spectra published in Wilding et al (2019b). However, over the course of several weeks, samples can deteriorate even when stored in a desiccator, and change in appearance from a transparent, pristine glass to an opaque white powder. This particular sample was transferred in one piece from Bristol to Munich within a few days of production and broken to extract fresh interior sample for the measurements. The degree of hydration is unlikely to change over the course of the subsequent calorimetry experiments.

In addition, a 50Na$_2$CO$_3$-50K$_2$CO$_3$ (molar) carbonate mix was produced from Merck high purity carbonates ($\geq$99.9%; $\geq$K$_2$CO$_3$ $\geq$99.5%) to compare with the previously published end members in Figure 1a, but over a larger temperature range than pervious datasets to better define the slope on a 1/T plot.
2.2 Thermogravimetry, calorimetry and viscometry: Combined thermogravimetric analysis (TGA) and differential scanning calorimetry (DSC) was performed using a Netzsch STA 449 C/3/G Jupiter® simultaneous viscometry analyzer at LMU. A single, transparent chip of the carbonate glass, of about 10 mg, was heated up to 920 °C at a constant rate of 5 K/min in a Pt-crucible under a purging, high purity Argon atmosphere (30 ml/min; see Figure 2). The temperature calibration was based on the melting points of Indium, Zinc, Ba-carbonate and Gold. The measurements were corrected for thermal drift. Differential Scanning Calorimetry measurements at greater sensitivity were made using a Netzsch DSC 404 C/3/F Pegasus® calorimeter. This also involved a chip of carbonate glass (40 mg), placed in a platinum crucible and heated with constant rates of 5, 10 and 15 K/min up to a temperature of 250 °C under a high purity Argon atmosphere (30 ml/min; see Figure 3). The sample remained transparent until the end of the heating cycles. The melting points of Indium, Zinc, Barium carbonate and Gold were also used to calibrate temperature for this instrument.

In addition, one shear viscosity measurement has been performed at 230 °C using the micro-penetration method under a purging, high purity Argon atmosphere (30 ml/min). For this measurement, the sample was held for 15 min at this dwell temperature prior to the initiation of indentation in order to establish thermal equilibration. A hemispherical Iridium-indenter was used with a force of 1 N into the sample and the penetration depth was monitored with time. The corresponding viscosity and error was calculated (Hess et al. 1995) and is tabulated in Table 1. Optical inspection after the measurement revealed a surficial tarnishing of the sample below the indenter.

The viscosity value was immediately taken after the mechanical relaxation of the measurement setup (some seconds) and is assumed to be uncorrupted by crystallization effects.
High temperature $50\text{Na}_2\text{CO}_3\text{-}50\text{K}_2\text{CO}_3$ viscosity measurements were made using an adapted commercially-available concentric cylinder rheometer (see Di Genova et al. 2016) equipped with a graphite furnace, and featuring an air-bearing-supported synchronous motor and a specially designed Pt-Au concentric cylinder crucible and spindle assembly. With this adaptation, high accuracy viscosity measurements of highly fluid melts can be achieved at high temperatures, up to 1273 K and at extremely low torques. This allows accurate viscosity measurements as low as $10^{-3.5}$ Pa·s (and up to $10^{3.5}$ Pa·s at shear rates up to $10^2 \text{ s}^{-1}$). The apparatus was calibrated with distilled water, silicone oils, and the DGG-1 standard glass (see Di Genova et al. 2016). Superliquidus $50\text{Na}_2\text{CO}_3\text{-}50\text{K}_2\text{CO}_3$ melt was measured from 1023 to 1198 K at a shear rate of 20 s$^{-1}$ (see Table 2).

2.3 First Principle (Ab Initio) Molecular Dynamics: The ab initio MD simulations for the $55\text{K}_2\text{CO}_3\text{-}45\text{MgCO}_3$ melt composition were performed using the Vienna Ab Initio Software Package (VASP) (Kresse and Furthmüller 1996a, 1996b). The electronic interactions described by projector-augmented wave (PAW) pseudopotentials (Blöchl 1994; Joubert 1999) were computed at the Brillouin zone Γ-point only with an energy cut-off value of 550 eV and the Perdew-Burke-Ernzerhof (PBE) formulation of the generalized gradient approximation (GGA) exchange correlation functional (Perdew et al. 1996). Molecular dynamics trajectories were computed in the canonical (NVT) ensemble with periodic boundary conditions and a Nosé thermostat (Nosé 1984). The system was simulated using a randomly generated cubic cell of $N=222$ atoms (44 K, 18 Mg, 40 C, 120 O) with initial closest approaches defined from the lowest crystallographic interatomic bond lengths. The starting configuration was super-heated to 10,000 K for 2 ps, cooled isochorically to 2500 K over 4 ps, and equilibrated at final temperatures $T = 2500, 2100, 1800,$
1500, and 1100 K for 30 ps. The final 25 ps of the equilibrated trajectories were taken for computing melt structure and properties. To mitigate for the under-binding of chemical bonds inherent to the GGA pseudopotential, we constrained the melt simulation cell volume, V, at each T to experimentally derived values at 1 atm, as calculated following Liu and Lange (2003). Our simulations suggest that Mg is dominantly in 5-fold coordination so a molar volume of 36.9 cm$^3$/mol was used which is halfway between the values suggested for 4- and 6-coordinated Mg by Hurt and Lange (2019). The calculated density trend is only slightly lower than the ambient Dobson et al. (1996) measurements for 50K$_2$CO$_3$-50MgCO$_3$ melt (2.1945 vs 2.2621 at 723 K and 2.1439 vs 2.2522 at 837 K) but at 2.3318 we nearly match the value of 2.325 measured for the glass (which we take as the melt density at a Tg of ~500 K). However, it should be noted that we have used using the thermal expansivities of CaCO$_3$. Although this appears to be applicable to all alkaline earth carbonates (Hurt and Lange 2019), it is extrapolated outside the calibrated range for the high temperature of the simulations and assumes a temperature-independent thermal expansivity, which is not an accurate assumption for liquids (e.g., Knoche et al. 1992; Dingwell et al. 1993). For the simulations used in this study (including data from Wilson et al. 2018; Wilding et al. 2019a) the derived diffusion data are converted to a viscosity using the Stokes-Einstein relationship either for the alkali atoms or the average of alkali and carbon atoms.

3. Results

3.1 DSC and DTA Glass Transition Determination

The results of differential thermal and thermogravimetric analysis (DTA-TGA) and differential scanning calorimetry (DSC) are shown in Figure 2 and 3 and tabulated in Table 1 for the 55K$_2$CO$_3$-45MgCO$_3$ glass. The DTA analysis (Figure 2) shows clearly the glass transition onset at ~220 °C (493 K). There is a small peak at 120 °C which we interpret as the loss of surface
water from these samples. Above 260 °C (533 K) the DTA curves show a series of exothermic peaks which results from crystallization, followed by phase transitions (above 400 °C) and finally at higher temperatures, endothermic peaks that represent solid state reaction and partial decarbonation resulting in a mixture of carbonates and oxides. The onset of melting occurs at 900 °C (1173 K). On cooling, crystallization occurs at the same temperature, and thus liquidus and solidus temperature are consistent. This liquidus temperature for the 55K$_2$CO$_3$-45MgCO$_3$ composition is much higher than the 500°C expected from the 0.1 GPa phase diagram of Ragone (1966). The reported eutectic in the binary system (57K$_2$CO$_3$-43MgCO$_3$ at 0.1 GPa) is at 460°C although the liquidus rises very steeply with composition. This may suggest a change in composition and the loss of CO$_2$ at ambient pressure consistent with the observation of Eitel and Skaliks (1929) that high pressure is required to ensure melting of K$_2$CO$_3$ and MgCO$_3$ constituents without decomposition but we note that any such effect is post-Tg determination.

The results of the high sensitivity DSC measurements are shown in Figure 3. The as-quenched glass was heated through the glass transition into the supercooled liquid regime at 250 °C (523 K) where the liquid is fully relaxed. The glass sample was then subjected to a series of thermal cycles comprising excursions across the glass transition at different heating rates, with the heating rate matched to the prior cooling rate (i.e., for example, a heating rate of 5 K/min is matched with a previous cooling rate of 5 K/min) (Easteal et al. 1977). With this approach of matching cooling and heating rates, the so-called onset of the calorimetric glass transition corresponds to the fictive temperature (Moynihan 1993; Yue et al. 2004; Moynihan 2019) determined by the enthalpy-matching method (Moynihan et al. 1976; Yue 2008). As the sample is heated at rates of 5, 10, and 15 K/min, the temperature of the glass transition (taken as the peak position corresponding to the heat flow peak overshoot, hereafter referred to as $T_{peak}$) shifts systematically with matched heating/cooling rate ($|q|$) from 229 to 237 ± 1 °C (502 – 510 K) for 5
to 15 K/min. The combined thermogravimetry (TG) and calorimetry (DSC) measurements demonstrate the thermal stability of the glass sample, the glass transition is clearly identified at 229 ± 1 °C (for a cooling rate of 15 K/min) consistent with the original study of Eitel and Skaliks (1929) based on rapid devitrification of the glass between 200-300 °C.

3.2 Shift factor

The shift factor $K$ (here $K_{\text{peak}}$) (Scherer 1984; Stevenson et al. 1995; Gottsmann et al. 2002; Yue et al. 2004; Al-Mukadam et al. 2020) can be derived from the DSC measurements (Figure 3) by matching the characteristic glass transition temperature $T_{\text{peak}}$ for a given cooling/heating ($q$, K/s) cycle to the temperature of the viscosity measurement:

$$
K_{\text{peak}} = \log_{10} \eta(T_{\text{peak}}) + \log_{10}|q| \quad (3)
$$

The value (7.95) obtained at cooling/heating rates of 15K/min (where the temperatures overlap with the accuracy of the measurements) is then used for the other cooling/heating cycles:

$$
\log_{10} \eta(T_{\text{peak}}) = K_{\text{peak}} - \log_{10}|q| \quad (4)
$$

to calculate the viscosity for the 55K$_2$CO$_3$-45MgCO$_3$ liquid (Table 1) ranging from $\log_{10} 9.03 \pm 0.09$ Pa.s at 229 °C to $\log_{10} 8.55$ Pa.s \(\pm 0.09\) at 237 °C (error in temperature was converted to error in viscosity). We note that this shift factor is lower than that observed for silicate melts but is generally consistent with a high compressibility as expected for carbonate melts.

3.3 A viscosity-temperature relationship for carbonate melts.
For the 55K$_2$CO$_3$-45MgCO$_3$ sample, the three viscosities derived using DSC measurements and the shift factor together with the one data point from the micro-penetration measurement are plotted in Figure 4. These low temperature data can be fitted to a single Arrhenian trend but the slope greatly exceeds that of an Arrhenian fit solely though the high temperature 50K$_2$CO$_3$-50MgCO$_3$ data of Dobson et al. (1996) or Sifre et al. (2015), or the trends of simulation data plotted in Figure 4. It is clear that both Arrhenian subsets of data exhibit slopes that are Arrhenian only as artefacts of the very restricted temperature ranges of the individual datasets. Taken together the data describe a strongly non-Arrhenian temperature-dependence of viscosity. An unweighted Vogel-Fulcher-Tammann (VFT) fit through both data sets is shown in Figure 4. The resultant fit using the Dobson et al. (1996) data demonstrates the ‘fragility’ of this system. At high temperatures, this fit also agrees well with the 50K$_2$CO$_3$-50MgCO$_3$ data of Sifré et al. (2015) who measured the electrical conductivity and demonstrated how this high temperature data can be converted to a viscosity. It must be borne in mind however, that both Dobson et al. (1996) and Sifré et al. (2015) data are at high pressure (3-6 GPa). As the magnitude of any pressure effect remains controversial and it is not possible to conduct high temperature viscosity measurements at 1 atm to directly complement the DSC data, we have further extended (for further comparison) the 55K$_2$CO$_3$-45MgCO$_3$ dataset using ab initio simulation data in Table 2. These 1 atm results are included in Figure 4 where they lie close to the Dobson et al. (1996) high-pressure data. In fact, our simulations indicate that from 1 atm to 3 GPa there is less than half log$_{10}$ unit increase in viscosity. Also included in Figure 4 is an additional simulation data point for ‘ambient’ pressure 55K$_2$CO$_3$-45MgCO$_3$ melt derived from the classical simulations of Wilding et al. (2019a, 2019b), which are consistent with high energy XRD data used to elucidate the 55K$_2$CO$_3$-45MgCO$_3$ glass structure.
When evaluating the reliability of the Dobson et al. (1996) $50K_2CO_3-50MgCO_3$ viscosities, Kono et al. (2015) implied that they should lie 0.5 log$_{10}$ unit higher to match the $50K_2CO_3-50CaCO_3$ Dobson et al. (1996) data as no compositional effect was expected. Desmaele et al. (2019a) also conclude the compositions should have similar viscosity, but also indicate that the effect of pressure would reduce this value by a log$_{10}$ unit at 1 atm. Given all these unresolved discrepancies, we have chosen to fit to the Dobson et al. (1996) $50K_2CO_3-50MgCO_3$ data accepting they are possibly too high and that a reduction in pressure brings them down and closer to our 1 atm simulated data for $55K_2CO_3-45MgCO_3$. Regardless of which high temperature dataset is used, it is clear the $55K_2CO_3-45MgCO_3$ melt is highly fragile. The VFT fitting also indicates a value between -3 and -4 log$_{10}$ (in Pa*s) at infinite temperature consistent with other theoretical and statistical estimates (Angell et al. 1989). A calculated VFT for Na$_2CO_3$ is also illustrated for comparison in Figure 5. This represents the viscosity obtained from the diffusion data obtained from the classical MD simulations of Wilson et al. (2018), which are tabulated in Table 2. The ‘q’ value, the degree of charge separation across the carbonate anion (explained in Appendix 1) is 2.28, the simulated diffraction pattern at this value gives the best fit to the data obtained from high energy XRD measurements on molten Na$_2CO_3$. The diffusion (and viscosity) data obtained at this value of charge separation shows fragile (non-Arrhenian) behavior, correlated with the temperature-dependent abundance of CO$_3^{2-}$ rings and other complexes. Using this value of charge separation gives the best VFT fit to the experimental Na$_2CO_3$ viscosity data of Di Genova et al. (2016) in Figure 5 and illustrates the robustness of this simulation methodology.

Figure 5 also contains our new viscosity measurements for $50Na_2CO_3-50K_2CO_3$, listed in Table 3. This composition is consistent with the Di Genova et al. (2016) dataset plotting between viscosity data curves for Na$_2CO_3$ and K$_2CO_3$ when measured under an argon atmosphere (see
Figure 1a). In the context of comparing to the Na$_2$CO$_3$ VFT curve, this data has the advantage of a larger temperature range that the Na$_2$CO$_3$ measurements. The 50Na$_2$CO$_3$-50K$_2$CO$_3$ binary viscosity dataset can be fit with a linear trend within error, but a slight positive curvature would also be permitted. The curved trend for the 22Li$_2$CO$_3$-33Na$_2$CO$_3$-45K$_2$CO$_3$ eutectic composition of Kim et al. (2015) as noted in Figure 1, is reproduced in Figure 5 and is also consistent with our Na$_2$CO$_3$ VFT curve.

The other synthetic composition data shown in Figure 5 are simulations performed by Desmaele et al. (2019b) that reveal the effect of adding calcium to the sodium-potassium system. Starting from the 50Na$_2$CO$_3$-50K$_2$CO$_3$ composition of this study, there is a clear increase in viscosity with the addition of CaCO$_3$. Although Desmaele et al. (2019b) provide only two data points, they are consistent with a decrease in fragility with CaCO$_3$ addition.

4. Discussion

4.1 Carbonate melt structure

The structure of levitated alkali carbonate liquids has been studied directly using high energy X-ray diffraction (Wilding et al. 2016; Wilson et al. 2018). These diffraction data have been combined with classical molecular dynamic simulations and used to evaluate the changes in the liquid structure with temperature and accordingly the carbonate liquid fragility. Central to the simulation methodology is the flexibility of the carbonate anion geometry. Spectroscopic studies of the K$_2$CO$_3$-MgCO$_3$ glass (Sharma and Simons 1980; Genge et al. 1995; Wilding et al. 2019a, Wilding et al. 2019b) have suggested the presence of two structurally distinct populations of carbonate anion, one of which is more distorted and it is this distortion of the carbonate that has been used as a basis for the molecular dynamics simulation of the alkali carbonate liquids. In the
simulation of Na$_2$CO$_3$, flexibility is imposed by employing springs between the O-O and C-O pairs in the anion, the stiffness of these springs is constrained by comparing the liquid diffraction patterns at high and low values of scattering vector with the simulated liquid structure, the inter and intra-molecular contributions respectively. Once constrained, the simulated liquid structure can be explored as a function of temperature. The flexibility of the molecular carbonate anion allows the central carbon atom to be drawn out of the triangular plane and the anion becomes polarized. One consequence of this is that there is the development of a secondary length scale and development of carbonate chains and other carbonate complexes, the extent of which are strongly temperature-dependent. In a more recent study (Wilson et al. 2018), the fluctuation of charge across the molecular anion was introduced into the simulation to further explore the dynamics of sodium carbonate liquids and demonstrates a correlation between the development of carbonate complexes and liquid fragility, with the connectivity of these emergent structures dependent on the mean charge separation (see Appendix 1).

This modelling approach, with flexibility of the molecular anion, has also been applied to the K$_2$CO$_3$-MgCO$_3$ system where the simulations are used to identify the pressure-dependent changes in structure of the same 55K$_2$CO$_3$-45MgCO$_3$ glass studied here. At ambient pressure there is no evidence for formation of carbonate chains or other complex structures in the 55K$_2$CO$_3$-45MgCO$_3$ carbonate liquid, however as pressure is increased there is the development of a carbonate network associated with an increase in mean coordination number of the carbon with development of a CO$_3^{3-}$ configuration (Wilding et al. 2019a). The response to pressure in these liquids again reflects the flexibility of the carbonate and the complex interaction between the oxygen atoms in the CO$_3^{2-}$ anion and their strong electrostatic interaction with potassium cations. The ambient pressure glass does not show development of this network but $^{13}$C NMR confirms a distorted carbonate anion (Wilding et al. 2019b), whilst infrared spectroscopy confirms the
presence of two structurally distinct populations of carbonate anions identified in earlier studies.

In fact, the simulation of the liquids suggests different degrees of flexibility of the carbonate anion rather than two distinct populations with the stronger interactions with Mg$^{2+}$ and K$^+$ cations associated with the more distorted carbonate. None the less, the simulation of the K$_2$CO$_3$-MgCO$_3$ liquid shows distorted carbonate with magnesium and potassium cations occupying irregular channels with both types of cations adopting a network-forming role by bridging isolated carbonate anions (Wilding et al. 2019b). It has been suggested that glass formation in sulphate and nitrate systems also requires the presence of two different cations with different field strengths and different degrees of polarizability (Förland and Weyl 1950; van Uitert and Grodkiewicz 1971; MacFarlane 1984; Wilding et al., 2017).

The ab initio molecular dynamics simulations that form part of this study can be used to evaluate the changes in the ambient pressure liquid structure as a function of temperature. As expected, there is no formation of a Na$_2$CO$_3$ style carbonate network in 55K$_2$CO$_3$-45MgCO$_3$, however the main changes in structure occur in the local environment surrounding the K$^+$ cations. There is an increase in the mean K-O coordination number which increases as temperature is decreased, but no change in the average C-O coordination at ambient pressure. The distortion in the local environments for potassium is shown in the changes in the partial contributions to the pair distribution functions for K-O and K-K (Appendix 2). The coordination environments for both potassium and magnesium differ significantly from those in the equivalent crystalline phase, illustrated in a snapshot from the ab initio simulation in Figure 6.

As noted above, the K$_2$CO$_3$-MgCO$_3$ composition is not a naturally-occurring carbonatite liquid. However, this and related studies show that combining experimentally-derived structures and structure-related properties with classic and ab initio simulation provide insight into the viscosity-temperature relations of mixed carbonate liquids and that the same modelling approach can be
extended to naturally occurring systems. In contrast to the alkali carbonates, there is no evidence for formation of a carbonate network in the K$_2$CO$_3$-MgCO$_3$ liquids. However, the liquid dynamics still reflect the underlying flexibility of the carbonate anion.

4.2 The fragility of carbonate melts.

The carbonate melt viscosities are compared in reciprocal absolute temperature (Arrhenian) space, with other geologically relevant liquids in Figure 7 using the data in Table 4. This comparison confirms that under similar conditions, the carbonate liquids have very low viscosity and are potentially very mobile, at least at the ambient pressures measured in this study.

The VFT fit is a powerful tool in enabling the extrapolation of the temperature-viscosity relationship to higher temperatures and thereby constraining the pre-exponential term in the temperature-dependent viscosity equation (c.f. Russell et al. 2003). As previously noted, the pre-exponential term or viscosity limit at infinite temperature lies just above $10^{-4}$ Pa·s, a value consistent with literature data for other simple molecular liquid classes (oxides, halides, silicates, etc.). In Figure 7a, (see Table 4) the viscosities of 55K$_2$CO$_3$-45MgCO$_3$ and Na$_2$CO$_3$ liquids are compared with the viscosity of several silicates as well as 40CaNO$_3$-60K$_2$NO$_3$ (CKN), generally considered to be an archetypal fragile liquid. In Figure 7b these are compared in an Angell plot (Böhner et al. 1993; Angell and Moynihan 2000) with the temperature normalized to Tg, and the liquids show progressive departure from the Arrhenius behavior of “strong liquids” such as the network-forming liquid SiO$_2$, seen as a series of curves of increasing fragility. The different degrees of fragility reflect differences in the temperature-dependence of liquid structure, through their relative contributions to the configurational entropy (Richet 1984). The temperature-dependent structural elements might be locally-favored structures, dynamic heterogeneities or density fluctuations within the supercooled liquid state.
5. Implications

It will be apparent that the fragility of the carbonate liquid and its high temperature viscosity is very similar to that of a peridotite melt (Dingwell et al. 2004). Liquid peridotite is one of the most fragile silicate compositions ever measured and is extremely difficult to quench to a glass. Peridotite is an example of a ‘fully depolymerized’ silicate melt, where the silicate network structure is broken into isolated units surrounded by metal cations (cf. Kohara et al. 2011). Peridotite and carbonate melts lie at higher fragility than all the other petrologically and volcanologically relevant silicate melts, including basalts (Giordano and Dingwell 2003), extremely peralkaline phonolites and pantellerites (Whittington et al. 2001; Di Genova et al. 2013) and water-rich calc-alkaline rhyolites (Hess and Dingwell 1996). Peridotite melt may have formed the magma ocean in early Earth history, when carbonate content could also have been relatively high. In fact, at high enough pressures there is a continuum from carbonate melt to ‘depolymerized’ high-CO$_2$ silicate melt compositions (such as melilitites) generated during high temperature carbonated mantle melting (e.g. Brey and Green 1976; Gudfinnsson and Presnall 2005). Initially, SiO$_2$ that is added to dilute the carbonate melt will remain as isolated ‘SiO$_4$’ units with little effect on the ionic structure and therefore viscosity. But as SiO$_2$ (and Al$_2$O$_3$) reaches 15-30 wt% these transitional melt compositions (Brooker et al. 2011) form new structures with two sub-networks one consisting of regions with polymerised silicate structures and the other ionic carbonate. These may be precursors to silicate-carbonate immiscibility (Brooker et al 2001, Morizet et al. 2017). The presence of two sub-networks at the molecular level could lead to a complex viscosity temperature dependence related to this medium range structural heterogeneity, with the faster and slower relaxing regions that are often cited as a cause of fragility. However,
the viscosity determinations of Morizet et al. (2017) and those inferred by the conductivity measurements of Sifré et al. (2014) suggest that varying the amount of CO$_2$ in transitional or basaltic silicate melts and thus the proportion of the two sub-networks, has little effect on the viscosity which remains similar to the CO$_2$-free silicate composition. This is consistent with our observations here that all very high temperature interactions between extremely fragile very low silica silicate melts (e.g. peridotite, melilitite, kimberlite) and carbonatites will be interactions between two liquids of similar low viscosities.

The absolute values of the glass transition temperatures, together with the restricted range of glass-forming ability in carbonatite, means that the likelihood of encountering glassy behavior in natural magmatic systems is vanishingly small. The more rapid increase in viscosity at lower temperatures is obviously more important in silicate systems, particularly for silica-rich eruptions where volatiles are being exsolved. Gas bubbles are unable to escape, instead building up internal pressure that results in explosive behaviour and fragmentation of a super cooled melt (i.e. glass) as the viscosity approached some critical value (usually considered around $10^7$ Pa·s; e.g., Namiki and Manga 2008). The stoichiometry of the carbonate components within carbonatite melts suggest that the main volatile component (CO$_2$) is not necessarily exsolved, at least in alkali-rich systems. Even if other volatiles are exsolved, the general low viscosities at eruptive temperatures (see Figure 7a) will allow bubbles to escape through the melt more efficiently than in a silicate melt. The exception to this could be the apparently explosive alkaline-earth carbonatite volcanism that produced rounded lapilli melt droplets, preserved at the Kaiserstuhl volcanic complex (Keller 1989). These have an almost pure CaCO$_3$ composition and the shape and internal structure suggest these are airborne quenched melt. Pure CaCO$_3$ should dissociate at pressures below 0.004 GPa, so how this composition could erupt molten droplets, is still a mystery. This has prompted the usual idea that minor amounts of alkalis or fluorine allow the melt to exist at low pressure and
these are subsequently leached away (Gittins and Jago 1991; Brooker and Kjarssgard, 2011). However, one could speculate that a rapid increase in viscosity and cooling within some ‘CO$_2$ confining pressure’ in the vent, could combine to preserve a CaCO$_3$ melt as droplets. Zimanowski et al. (1986) have demonstrated droplet formation for 50Na$_2$CO$_3$-50K$_2$CO$_3$ and Oldoinyo Lengai carbonatites by interaction of the melt with injected water. As Zimanowski et al. (1986) point out these processes are perhaps also important for modelling explosive situations involving industrial molten salt cells where a full understanding of the physical processes becomes important as regards hazard mitigation.

Zimanowski et al. (1997) have indeed directly observed the fracturing of a high temperature carbonate liquid upon interaction with external water. They effectively induced the glass transition in the carbonate liquid and recorded the results by high-speed video. These observations are entirely consistent with the observation here that a glass transition can be encountered in molten carbonates. With knowledge of the viscosity-temperature curve observed here, together with the effective temperature of the melt-water interaction in the experiments of Zimanowski et al. (1997) one should be able in principle to estimate the volume strain rates at the point of brittle failure of the carbonate liquid. Alternatively, if the thermal stresses are well-estimated then one should be able to predict the effective temperature of the brittle failure-inducing melt – water interaction. That would have been impossible to achieve accurately with a degree of certainty by simply using extrapolations of superliquidus viscosity-temperature relations.

A non-linear rate change in viscosity is also an important parameter in understanding the nature of carbonatite lava flows, particularly the length and terminal velocity at Oldoinyo Lengai. The same applies to the calculations of Treiman and Schedl (1983) who calculated the carbonatite melt properties in magma chambers and their effect on turbulence and the rapid setting velocities.
of crystals and growth rates on chamber walls, with implications also for the settling rate in lava flows (Norton and Pinkerton 1997).

The most widespread occurrence of carbonate melts is most likely deep in the Earth’s mantle. Dalton and Wood (1993) demonstrated that carbonatites generated by melting a depleted carbonated mantle source are almost alkali-free, ranging from 25MgCO$_3$-75CaCO$_3$ to 8MgCO$_3$-92CaCO$_3$ (with some minor FeCO$_3$), but sodium can reach 15Na$_2$CO$_3$ for a fertile mantle source and subsequently be increased or decreased due to metasomatic wall-rock reactions. The range of compositions derived from subducted altered basalt (eclogite) in Thomson et al. (2016) is even more diverse, ranging to higher contents of iron- and alcalis (14-24FeCO$_3$, 9-15MgCO$_3$, 43-66CaCO$_3$, 2-29Na$_2$CO$_3$, 1-3K$_2$CO$_3$). At these conditions, the viscosity can be reasonably well approximated by a linear temperature dependence trend. The possible compositional effect of Na$_2$CO$_3$ inferred from Figure 1b and Figure 5 suggest this is an important component when considering the viscosity of mantle carbonatites, although these mantle melts are all produced at high temperatures (> 1200 °C) and even 1 atm sodic melts would have viscosities within 0.5 log$_{10}$ units of the other compositions in Figure 1b. However, the hydrous ‘fertile pyrolite’ composition of Wallace and Green (1988) produced an alkali-rich carbonatite melt (~ 5K$_2$CO$_3$-33Na$_2$CO$_3$-62CaCO$_3$) at temperatures between 930 and 1080 °C at 2.1 GPa. Wallace and Green (1988) suggest an even more alkali-rich carbonate melt may exist at even lower temperatures perhaps forming before the water is released by hydrous silicate minerals. These low temperatures are realistic for the shallow lithosphere beneath old continental crust and this is getting into the temperature range where a linear fit based on the high temperature data will become inaccurate and the high fragility of the Na$_2$CO$_3$ component may become important (Fig. 5). This would have implications for modelling the transport properties of these melts to their surface expression or as metasomatizing agents.
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Tables

Table 1: $K_2CO_3$-$MgCO_3$ melt viscosity data from differential scanning calorimetry (DSC), micro-penetration viscometry (MP) and falling sphere viscometry data (FS) used for an unweighted VFT fitting.

<table>
<thead>
<tr>
<th>Method</th>
<th>Temperature °C</th>
<th>Quench rate log q (K/s)</th>
<th>Viscosity log (Pa*s)</th>
<th>Error</th>
<th>Pressure</th>
</tr>
</thead>
<tbody>
<tr>
<td>DSC$^a$</td>
<td>237</td>
<td>-0.60</td>
<td>8.55</td>
<td>0.09</td>
<td>1 atm</td>
</tr>
<tr>
<td>DSC$^a$</td>
<td>234</td>
<td>-0.78</td>
<td>8.73</td>
<td>0.09</td>
<td>1 atm</td>
</tr>
<tr>
<td>VISC-MP</td>
<td>230</td>
<td></td>
<td>8.95</td>
<td>0.06</td>
<td>1 atm</td>
</tr>
<tr>
<td>DSC$^a$</td>
<td>229</td>
<td>-1.08</td>
<td>9.03</td>
<td>0.09</td>
<td>1 atm</td>
</tr>
<tr>
<td>VISC-FS$^b$</td>
<td>800</td>
<td></td>
<td>-1.44</td>
<td>0.43</td>
<td>3 GPa</td>
</tr>
<tr>
<td>VISC-FS$^b$</td>
<td>900</td>
<td></td>
<td>-1.66</td>
<td>0.43</td>
<td>3 GPa</td>
</tr>
<tr>
<td>VISC-FS$^b$</td>
<td>1200</td>
<td></td>
<td>-2.22</td>
<td>0.43</td>
<td>5 GPa</td>
</tr>
</tbody>
</table>

$^a$ using a shift factor of 7.95; error in T converted in error in viscosity

$^b$ from Dobson et al. 1996
Table 2: Diffusion coefficients obtained from simulations and used to derive viscosities

<table>
<thead>
<tr>
<th>Temperature (K)</th>
<th>Potassium</th>
<th>Carbon</th>
<th>Average</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>Log D (cm$^2$/s)</td>
<td>Log $\eta$ (Pa.s)</td>
<td>Error</td>
</tr>
<tr>
<td>1500</td>
<td>-4.875</td>
<td>-2.789</td>
<td>-0.244</td>
</tr>
<tr>
<td>1800</td>
<td>-3.998</td>
<td>-2.499</td>
<td>-0.200</td>
</tr>
<tr>
<td>2100</td>
<td>-3.857</td>
<td>-3.023</td>
<td>-0.193</td>
</tr>
<tr>
<td>2500</td>
<td>-3.538</td>
<td>-3.267</td>
<td>-0.177</td>
</tr>
<tr>
<td>1850</td>
<td>-4</td>
<td>-2.935</td>
<td>-0.200</td>
</tr>
</tbody>
</table>

Values derived from; a) ab initio calculations, b) the classical simulations presented in Wilding et al. (2019a) where only the K diffusivity is provided, and c = Wilson et al. (2018) for a favored ‘q’ value of 2.28 (see appendix).
Table 3: High temperature viscometry data for synthetic alkali carbonate melts.

<table>
<thead>
<tr>
<th>°C</th>
<th>10000/T (K⁻¹)</th>
<th>η (Pa*s)</th>
<th>log η (Pa*s)ᵃ</th>
<th>Shear rate (s⁻¹)</th>
</tr>
</thead>
<tbody>
<tr>
<td>750</td>
<td>9.775</td>
<td>0.00598</td>
<td>-2.22</td>
<td>20</td>
</tr>
<tr>
<td>775</td>
<td>9.542</td>
<td>0.00554</td>
<td>-2.26</td>
<td>20</td>
</tr>
<tr>
<td>800</td>
<td>9.320</td>
<td>0.00515</td>
<td>-2.29</td>
<td>20</td>
</tr>
<tr>
<td>825</td>
<td>9.107</td>
<td>0.00481</td>
<td>-2.32</td>
<td>20</td>
</tr>
<tr>
<td>850</td>
<td>8.905</td>
<td>0.00453</td>
<td>-2.34</td>
<td>20</td>
</tr>
<tr>
<td>875</td>
<td>8.711</td>
<td>0.00429</td>
<td>-2.37</td>
<td>20</td>
</tr>
<tr>
<td>900</td>
<td>8.525</td>
<td>0.00409</td>
<td>-2.39</td>
<td>20</td>
</tr>
<tr>
<td>925</td>
<td>8.347</td>
<td>0.00385</td>
<td>-2.41</td>
<td>20</td>
</tr>
</tbody>
</table>

ᵃ Error is ± 0.01 log (Pa*s)
Table 4: Fit parameters describing the temperature dependence of the viscosity (VFT), the glass transition temperature $T_{12}$, the activation energy $E_a(T_{12})$ and the fragility index (m) for carbonates in this study and selected silicate liquids of geological relevance.

<table>
<thead>
<tr>
<th>Composition</th>
<th>A (log $\eta$ (Pa s))</th>
<th>B (K$^{-1}$)</th>
<th>C (K)</th>
<th>$T_{12}$ (K)</th>
<th>$E_a(T_{12})$ (kJ/mol)</th>
<th>Fragility Index (m)</th>
</tr>
</thead>
<tbody>
<tr>
<td>$^a$Haplo-rhyolite</td>
<td>-6.219</td>
<td>15950</td>
<td>232.2</td>
<td>1108</td>
<td>489</td>
<td>23</td>
</tr>
<tr>
<td>$^b$Phonolite</td>
<td>-4.550</td>
<td>10261</td>
<td>263.8</td>
<td>884</td>
<td>399</td>
<td>24</td>
</tr>
<tr>
<td>$^b$Trachyte</td>
<td>-4.550</td>
<td>10449</td>
<td>303.7</td>
<td>935</td>
<td>439</td>
<td>25</td>
</tr>
<tr>
<td>$^c$SiO$_2$</td>
<td>-4.167</td>
<td>15336</td>
<td>508.5</td>
<td>1457</td>
<td>693</td>
<td>25</td>
</tr>
<tr>
<td>$^b$Basalt</td>
<td>-4.550</td>
<td>6101</td>
<td>567.0</td>
<td>936</td>
<td>753</td>
<td>42</td>
</tr>
<tr>
<td>55K$_2$CO$_3$-45MgCO$_3$</td>
<td>-4.010</td>
<td>1914</td>
<td>356.0</td>
<td>476</td>
<td>580</td>
<td>64</td>
</tr>
<tr>
<td>$^d$Peridotite</td>
<td>-4.310</td>
<td>3703</td>
<td>761.7</td>
<td>989</td>
<td>1345</td>
<td>71</td>
</tr>
<tr>
<td>Na$_2$CO$_3$</td>
<td>-3.470</td>
<td>832</td>
<td>329.0</td>
<td>383</td>
<td>807</td>
<td>110</td>
</tr>
<tr>
<td>$^e$CKN</td>
<td>-3.602</td>
<td>798</td>
<td>324.4</td>
<td>376</td>
<td>824</td>
<td>115</td>
</tr>
</tbody>
</table>

Glass transition temperature $T_{12}$ at log$_{10}$ 12 Pa.s, activation energy $E_a$ at $T_{12}$ derived from VFT parameter. Other data from $^a$Dorfman et al. (1996), $^b$Giordano et al. (2009), $^c$Hetherington, et al. (1964), Hofmaier and Urbain (1968), $^d$Dingwell et al. (2004), $^e$Tweer et al. (1971), Weiler et al. (1969).

Figure caption

Figure 1. Previously measured viscosity data for (a) alkali-carbonates at atmospheric pressure (unless labelled) and (b) high pressure experiments on alkali-earth and alkali carbonates (pressures indicated). The yellow hexagons for high pressure NC data are also reproduced from (a). Also included in (a) are new 50Na$_2$CO$_3$-50K$_2$CO$_3$ data from the results of this study. Note, in this figure and the text, molar amounts of carbonate are expressed in the form LC = Li$_2$CO$_3$, NC = Na$_2$CO$_3$, KC = K$_2$CO$_3$, MC = MgCO$_3$, CC = CaCO$_3$, FC = FeCO$_3$, except (Li,N,K)C = 22Li$_2$CO$_3$-33Na$_2$CO$_3$-45K$_2$CO$_3$. The error is less than symbol size for the 1 atm measurements, the errors for the high pressure are hard to access. Data from: $^a$Kim et al. (2015), $^b$Di Genova et al. (2016), $^c$Sato et al. (1999), TS This study, $^d$Stagno et al. (2018), $^e$Dobson et al. (1996), $^f$Kono et al. (2014). The green dotted line in (a) is an Arrhenian fit through the 50Na$_2$CO$_3$-50K$_2$CO$_3$ data and reproduced in (b) for comparison. Simulation data for CC are taken from Vuilleumier et al. (2014), gray lines, at pressure indicated.
Figure 2. DSC/DTA measurements of a single $55K_2CO_3-45MgCO_3$ glass chip heated to 750 °C. The onset of the glass transition is seen at about 225 °C (see inset). The relaxed glass is then devitrified shown by two exothermic crystallization peaks at ~300 and ~340 °C. On further heating the sample partly decomposes and forms a mixture of carbonates and oxides which cannot be quenched to form a glass. The stable mixture melts at 900 °C.

Figure 3. High sensitivity DSC measurement of a 40 mg sample of $55K_2CO_3-45MgCO_3$ glass with a series of excursions across the glass transition with heating rate matched to prior cooling rate at the rates 5-15 K/min.

Figure 4. Viscosity curve for $55K_2CO_3-45MgCO_3$ composition liquid obtained by VFT fit to the low temperature viscosity data from this study, and the $55K_2CO_3-45MgCO_3$ liquid viscosity data at 3-5 GPa from Dobson et al (1996). The fit is bracketed by the simulations of Wilding et al. (2019a, 2019b) and the VASP data (average diffusion coefficients) from this study, both at 1 atm. For comparison, the Sifré et al. (2015) conductivity data for a similar composition at 3 GPa, has been converted to a viscosity. Also shown are the simulations of Desmaele et al. (2019a).

Figure 5. The sodium carbonate VFT fit has been derived from the simulations of Wilson et al. (2018), but with selected parameters to best the fit the sodium carbonate data of Di Genova et al. (2016). Other published data for synthetic alkali carbonate mixes ranging from pure alkalis, to variable amounts of added calcium carbonate. Data from: a Desmaele et al. (2019b), b Kim et al. (2015), c Di Genova et al. (2016), d Sato et al. (1999).

Figure 6. A snapshot of the high temperature liquid configuration, carbon atoms are shown in black, oxygen atoms red, potassium atoms blue and magnesium atoms yellow, obtained directly from the ab initio MD (VASP) simulation trajectories for liquid $55K_2CO_3-45MgCO_3$ at 2500 K.

Figure 7. (a) Viscosity versus temperature in an Arrhenian plot. (b) Viscosity curves standardized to their glass transition ($T_{12}$) to give a fragility (Angell) plot of liquids of geological relevance, and also for the archetypical fragile liquids $40CaNO_3-60K_2NO_3$ (CKN). The VFT curves from this study are included and as expected the carbonates are significantly more fragile than the silicates.
Figure 2

DSC (mW/mg)

Transition to carbonate and oxide 1180 K

Endo

Exo

Glass transition

Crystallisation

Temperature (K)

400 600 800 1000 1200

450 500 550 600 650

Temperature (K)