1 Revision 1

2 Quantitative WDS Compositional Mapping

3 Using the Electron Microprobe

- 4 John J. Donovan¹, Julien M. Allaz², Anette von der Handt³, Gareth G. E. Seward⁴, Owen Neill⁵,
- 5 Karsten Goemann⁶, Julie Chouinard¹ and Paul Carpenter⁷
- 6 1. CAMCOR, University of Oregon, Eugene, Oregon, 97403, USA
- 7 2. Institute of Geochemistry and Petrology, ETH Zürich, 8092 Zürich, Switzerland
- 8 3. Department of Earth Sciences, University of Minnesota, Minneapolis, MN 55455, USA
- 9 4. Department of Earth Science, University of California Santa Barbara, Santa Barbara, CA
- 10 93101, USA
- 11 5. Department of Earth and Environmental Sciences, University of Michigan, Ann Arbor, MI
- 12 48013, USA
- 13 6. Central Science Laboratory, University of Tasmania, Hobart, Tasmania 7001, Australia
- 14 7. Department of Earth and Planetary Sciences, Washington University in St. Louis, One
- 15 Brookings Drive, St. Louis, Missouri, 63130 USA
- 16
- 17 Abstract word count: 345

18 Abstract

19 While much progress has been made in electron-probe microanalysis (EPMA) to improve the 20 accuracy of point analysis, the same level of attention has not always been applied to the 21 quantification of wavelength-dispersive spectrometry (WDS) X-ray intensity maps at the 22 individual pixel level. We demonstrate that the same level of rigor applied in traditional point 23 analysis can also be applied to the quantification of pixels in X-ray intensity maps, along with 24 additional acquisition and quantitative processing procedures to further improve accuracy, 25 precision and mapping throughput. Accordingly, X-ray map quantification should include pixel-26 level corrections for WDS detector deadtime, corrections for changes in beam current (beam 27 drift), changes in standard intensities (standard drift), high accuracy removal of background 28 intensities, quantitative matrix corrections, quantitative correction of spectral interferences, and, 29 if required, time dependent corrections (for beam and/or contamination sensitive materials). The 30 purpose of quantification at the pixel level is to eliminate misinterpretation of intensity artifacts. 31 inherent in raw X-ray intensity signals, which distort the apparent abundance of an element. 32 Major and minor element X-ray signals can contain significant artifacts due to absorption and 33 fluorescence effects. Trace element X-ray signals can contain significant artifacts where phases 34 with different average atomic number produce different X-ray continuum (bremsstrahlung) 35 intensities, or where a spectral interference, even an apparently minor one, can produce a false-36 positive intensity signal. The methods we propose for rigorous pixel quantification requires 37 calibration of X-ray intensities on the instrument using standard reference materials, as we 38 already do for point analysis, which are then used to quantify multiple X-ray maps, and thus the 39 relative time overhead associated with such pixel by pixel quantification is small. Moreover the 40 absolute time-overhead associated with this method is usually less than that required for 41 quantification using manual calibration curve methods, while resulting in significantly better

42	accuracy. Applications to geological, synthetic, or engineering materials are numerous as
43	quantitative maps not only show compositional 2D variation of fine-grained or finely zoned
44	structures, but also provide very accurate quantitative analysis, with precision approaching that
45	of a single point analysis, when multiple-pixel averaging in compositionally homogeneous
46	domains is utilized.

47 Keywords: EPMA, WDS, quantitative analysis, X-ray mapping, quantitative mapping.

48

49 Introduction

50 Ouantification of WDS X-ray intensities by electron-probe micro analysis (EPMA) has 51 progressed over the last few decades with the development of better instrument hardware for 52 stable operation, and improved software for physics-based background, matrix, and spectral 53 interference corrections (Armstrong, 1988; Donovan et al., 1993; Donovan and Tingle, 1996). 54 However, these advances have generally been limited to so-called "point analysis", by which we 55 mean quantitative characterization of elemental chemistry, typically at a nominal volume of a 56 cubic micrometer, when the beam and stage are stationary with respect to each other. This 57 quantitative point analysis now approaches and may be better than 2% relative accuracy for major and even minor elemental concentrations in many compositional matrices. Additionally, 58 59 these improvements in both hardware and software, combined with measurements at high beam 60 currents and extended integration times, have demonstrated sensitivity at ppm levels for 61 quantitative point analysis in selected materials (e.g., Donovan et al., 2011; Batanova et al., 62 2015).

63 In practice, quantitative EPMA point analysis can be performed in minutes for a suite of 64 elements and includes both background and matrix corrections. Conversely, X-ray maps obtain 65 spatial information by limiting the dwell-time per pixel to hundreds to thousands of milliseconds. 66 which, depending on the pixel resolution and sensitivity required, can take hours to days of map 67 acquisition time. These shorter pixel intensity integration times result in a reduction in precision 68 due to counting statistics, but in principle it does not reduce accuracy when the pixel intensity 69 values are corrected using the same methods that are already utilized for conventional point 70 analysis.

Some investigators have proposed using calibration curves for quantification of raw intensity Xray maps and such alternative quantification methods will be discussed below. However, except for specialized circumstances (e.g., trace carbon in steel [Eichen et al., 1972]; low voltage analysis [Moy at al., 2020]), such methods are problematic for optimal accuracy, especially in multi-phase materials often encountered by geologists and material scientists.

76 It should be noted that the evaluation of raw X-ray intensity maps may be no better than 77 comparing raw intensities from point measurements, if the map interpretation is dependent on 78 quantitative comparisons, as is usually the case. Analysts do not generally compare raw X-ray 79 count rates when interpreting their point analyses, as these count rates include the 80 bremsstrahlung ("background" or continuum X-ray signal) and are also subject to non-systematic 81 variations due to matrix effects; likewise, count rates obtained during X-ray mapping are subject 82 to these same effects, and therefore evaluation of raw intensity X-ray maps are subject to the 83 same pitfalls (i.e., the same X-ray intensity can represent widely different concentrations from 84 pixel to pixel depending on the composition of the specific pixel under examination). Science 85 depends on accurate numerical quantification and X-ray mapping should be no exception. This

paper describes a quantification protocol for the treatment of WDS element X-ray maps obtained
on the electron microprobe, in order to yield the most accurate and precise results, just as we
already do for our point analyses.

89

90 Acquisition of X-ray Maps for Quantification Using the

91 Electron Microprobe

92 X-ray element mapping can be performed either using a wavelength-dispersive spectrometer 93 (WDS) or a silicon-drift energy dispersive spectrometer (EDS) with pros and cons for each 94 detector. Modern silicon-drift EDS systems exhibit high throughput of X-ray intensities and can 95 be important for element mapping of major elements. The main advantage of EDS is to measure 96 simultaneously a wide range of X-ray energies during acquisition in a relatively short amount of 97 time, i.e. a few seconds per point, though artifacts in EDS spectra (see below) can become 98 problematic at typical high beam currents necessary for efficient mapping. Unfortunately, when 99 scanning over large areas and collecting X-rays as pixels, the counting statistics on a given pixel 100 at a given energy channel are comparatively poor, and therefore the peak-to-background (P/B) 101 ratios remain low. Moreover, EDS systems often exhibit high dead times at the elevated beam 102 current values typically used for element mapping of minor and trace elements (>100-200 nA). 103 At such high beam currents, these tradeoffs result in utilizing either a shorter pulse processing 104 time constant at the cost of poorer spectral resolution, which is not ideal for quantification, or a 105 longer pulse processing time constant by utilizing a reduced EDS-aperture size, which reduces 106 sensitivity and introduces spectral artifacts. Furthermore, the development of pileup peaks and 107 Si-escape peaks (Newbury, 1995) complicates the quantitative analysis and interference

108 correction of EDS spectra, although new EDS algorithms may offer better corrections for these109 artifacts (Newbury and Ritchie, 2019).

110 Measurement by WDS has the superior advantages of higher P/B and X-ray counting rates as the 111 X-ray wavelength of interest is selected by Bragg diffraction, and the counting electronics 112 process only these selected pulses and automatically rejects most high-order X-ray line 113 reflections, when differential pulse height analysis (PHA) parameters are properly adjusted. 114 These improved WDS P/B statistics and higher counting rates result in better analytical 115 sensitivity, and they allow WDS to be used not only for major, but also for minor and trace 116 element measurement. WDS is also the preferred choice for routine analysis and element 117 mapping of a wide range of phases, and excels in the measurement of elements at trace element 118 concentration and/or having low energy X-rays. (Goldstein, et al., 1992). In this paper we will 119 limit ourselves to a discussion of quantitative analysis of WDS X-ray maps acquired with EPMA 120 instruments.

121 Element mapping by WDS on the electron microprobe is usually performed by using a fixed 122 electron beam ("spot" mode) in combination with specimen stage scanning motion, an approach 123 referred to as "stage mapping". This fixed beam is used to avoid lateral defocusing of the 124 spectrometer when scanning sample areas larger than the WDS Bragg focus dimension. During 125 stage mapping, the image pixel dimension is defined by the stage motor step size, and the 126 electron beam diameter is usually set equal to or slightly less than this pixel dimension. On JEOL 127 microprobes the stage sample holder is then scanned in y-axis strips with incremental movement 128 of the x-axis starting on the top-right corner, and on Cameca microprobes the stage is scanned in 129 x-axis strips with incremental movement of the y-axis starting on the top-left corner. These 130 differences have implications for how the calculations for beam drift and standard intensity drift

corrections are applied to the pixel acquisition order. It should also be noted that when stage 131 132 maps are acquired over large sample areas or for long durations, both the flatness of the sample 133 and the stability of the instrument are of paramount importance for X-ray map quantification. 134 Any changes in the sample Z-axis focus or instrument stability can result in significant X-ray 135 measurement errors due to changes in Bragg defocus and column and/or spectrometer alignment. 136 For higher spatial resolution, a beam scanning method with sub-micron pixel size is preferred in 137 order to minimize stage reproducibility errors after each pixel strip and between multiple passes 138 over the same area. This beam mapping method can only be performed on small areas, where 139 Bragg defocusing effects are negligible ($\leq 20-40 \ \mu m$ lateral map dimensions, dependent on the 140 WDS Rowland circle and the specific Bragg crystal). In either case, each pixel records the total 141 X-ray count on each WDS channel, with each spectrometer tuned to an X-ray characteristic peak 142 position of an element of interest. For multiphase samples with numerous major and minor 143 elements to map, it is typically necessary to perform this acquisition using 1 to 3 or more passes 144 over the same sample area; one pass for each set of WDS elements to be measured. For instance, 145 on a 5-spectrometer microprobe it is necessary to do 2 passes of 5 elements each to acquire up to 146 10 elements for a typical silicate analysis (usually Si, Al, Fe, Mn, Mg, Ca, Na, and K, and as 147 required Ti, Cr, P, S, Ni or other trace or minor elements of interest). 148 For accurate map quantification, it is necessary to remove any intensities not pertaining to the 149 characteristic X-ray to be quantified. First and foremost is the bremsstrahlung, which is

150 commonly referred to as the X-ray continuum or X-ray background. The conventional

background correction method for point analysis is to drive the WDS spectrometer to two off-

152 peak positions, that is, to the sides of the characteristic X-ray peak position, in order to measure

153 the X-ray intensity at a spectrometer position where only continuum X-rays satisfy Bragg's law.

154 By extrapolating or interpolating such intensity values to the on-peak position, the continuum 155 intensity under the peak can subsequently be determined. Such off-peak background 156 measurements can also be acquired during quantitative X-ray mapping, but this requires the 157 acquisition of not only an on-peak map, but also the additional acquisition of at least one (e.g., 158 Cameca PeakSight software, https://www.cameca.com/), and ideally two background maps 159 (Probe Software Probe Image software, https://probesoftware.com/), particularly for emission 160 lines where the continuum spectrum is highly sloped or curved (e.g., at low spectrometer limits; 161 Jercinovic et al., 2012). If two background X-ray maps are acquired, an interpolation of the 162 intensities is made at each pixel using linear, polynomial, exponential or slope regression 163 methods in order to obtain the background intensity to subtract from the on-peak X-ray map 164 pixel intensity. 165 However, considerable acquisition time can be saved if no background X-ray maps are actually 166 acquired, but rather the background intensity is constrained by a standards-based background 167 calibration curve such as the mean atomic number (MAN) background correction method 168 (Kramers, 1923; Merlet and Bodinier, 1989; Donovan and Tingle, 1996). The use of the MAN 169 background corrections do not require the acquisition of off-peak X-ray maps at all, thus 170 resulting in not only significant savings in mapping acquisition time, but also improved 171 sensitivity due to the statistics of the MAN background correction method (Donovan et al., 172 2016). This approach is recommended and discussed further below. 173

174 Calibration Curve and Other Methods for X-ray Map

175 Quantification

176 Tinkham and Ghent (2005) have proposed X-ray map pixel quantification using Bence-Albee 177 (hyperbolic) calibration curve methods, combined with a single standard MAN background 178 correction, which they claim can produce quantitative X-ray maps, except in cases of high 179 absorption or fluorescence (Donovan, et al., 2019), or when mapping trace elements in multi-180 phase materials with different average atomic numbers (see below for more details). However, 181 since we already have more accurate analytical physics and background removal techniques 182 available from point analysis methods, it is reasonable to apply our full knowledge of electron-183 microprobe physics to the analytical situation at hand, whether that application is point analysis 184 or X-ray mapping.

185 Alternatively, some investigators (e.g., Lanari et al., 2014; Ortolano et al., 2018) have proposed 186 quantifying raw X-ray maps by means of calibration curve methods. These methods attempt to 187 relate the raw X-ray intensities in maps to concentrations determined from at least two or more 188 single-point analyses that are considered representative of the range of compositions of the 189 different phases within the map area. From this, one can interpolate concentrations for a given 190 arbitrary X-ray intensity. Although in principle this method appears to provide some usefulness 191 for quantification of X-ray maps of a single phase, the nature of electron and X-ray physics 192 informs us that when applied to situations with multiple phases, each with diverse matrix 193 physics, these assumptions of compositional linearity are unfounded (Barkman et al., 2013, 194 Donovan et al., 2019). Even worse are these assumptions, applied to trace minor and element 195 analysis, when multiple phases with different average atomic numbers are present, as each zero 196 concentration is represented by a different X-ray intensity. The calibration curve method is also 197 unable to accurately correct for spectral peak interferences in different phases.

198 In systems with multiple phases, the same X-ray intensity in different phases can represent quite 199 different elemental concentrations due to absorption and/or fluorescence effects. Similarly, the 200 X-ray intensity corresponding to a zero concentration of an element, i.e., the bremsstrahlung 201 intensity at the peak position of interest, depends wholly on X-ray continuum absorption and on 202 the average atomic number of the phase, which can vary significantly from one phase to another 203 and sometimes even within one a single phase exhibiting substantial compositional zonation. For 204 example. Si K α intensity data obtained in multiple silicate reference materials plotted against the 205 Si concentration do not lie on a linear line trend, and even an exponential calibration curve 206 cannot be properly fitted (Fig. 1). 207 Due to differences in the matrix correction physics for each of the phases plotted, there is 208 insufficient linearity to robustly correlate intensity to concentration when more than a single 209 phase is present or when a strong compositional zonation is observed and is not fully covered by 210 the single-point analysis obtained in the phase of interest. It is in fact, precisely for this reason 211 that such calibration curves are very limited in application for EPMA point analysis 212 quantification, with rare exceptions such as trace carbon in steel, where several reference 213 materials of similar range in composition are considered (Crisci et al., 2006). These non-214 linearities are even more obvious when examining the numerical results from a generic multi-215 standard calibration curve as seen in Table 1 for the analysis of several different standard 216 materials. Clearly, we should expect better accuracy than this for our X-ray map quantification 217 efforts.

218

A New Method for Rigorous Quantification of X-ray Maps

220 The method for quantitative X-ray mapping presented in this paper largely replicates the 221 quantitative methods already developed for point analysis. The main difference is that the full set 222 of correction algorithms are applied to intensities measured at each individual pixel in the X-ray 223 map, rather than just a single discrete point. This process, summarized in Figure 2 and detailed in 224 the following discussion, begins with the extraction of the array of raw counts obtained for all 225 elements on each pixel, one pixel at a time. The counts received are normalized to the counting 226 time and beam current to yield X-ray intensities in counts per second per nanoamp (cps/nA). 227 These resulting intensities are then corrected for dead time, beam drift, standard intensity drift, 228 background, and time dependent intensity (TDI) corrections when necessary, e.g. for beam 229 sensitive samples and for contamination issues such as trace carbon measurements. This whole 230 correction procedure is then run iteratively through a full matrix correction, either ZAF or $\phi(\rho z)$ 231 methods, as required by the analyst, along with any quantitative spectral interference corrections 232 specified by the user. Of course, the exact same normalization and correction procedures are also 233 applied to the standard intensities from point measurements, which will be utilized not only in 234 the construction of the intensity k-ratio prior to the matrix correction procedures, but also in the 235 calculation of the MAN (absorption corrected) background calibration curve for each analyzed 236 element, and in the peak interference correction (if necessary) for the primary standards.

237

238 Intensity Normalization and Dead Time Correction

The first step in the quantification of X-ray maps is to divide the total X-ray intensities in both
the standard point measurements and the unknown pixel measurements, by their respective
integration times and beam currents to obtain counts per second per nanoamp (cps/nA)
intensities (equation 1).

243
$$i_{cps/nA} = \frac{i_{counts}}{t*nA}$$
(Eq. 1)

244 Where :

245	i _{cps/nA}	Normalized intensity in counts per second per nanoamp
246	i _{counts}	Number of total raw X-ray intensity measured
247	t	Counting integration time in seconds
248	nA	Beam current in nanoamps (measured or interpolated)

The beam current values utilized for the standards are the measured beam currents from the standard intensity point measurements, while the beam current values for the map pixels are the interpolated beam currents measured prior and subsequently to the pixel intensity mapping measurements (see below for details).

Second, a dead time correction is applied using a single factorial expression (equation 2a).
However, when utilizing high beam currents to improve sensitivity, the count rate can exceed
50,000 counts per second, and a high-precision expression that extends the factorial to another
term is usually preferred (equation 2b).

257
$$i_{dt} = \frac{i_{cps}}{1 - i_{cps} * \tau}$$
(Eq. 2a)

258
$$i_{dt} = \frac{i_{cps}}{1 - \left(i_{cps} * \tau + \frac{\tau^2}{2}\right)}$$
 (Eq. 2b)

259 Where :

$$i_{dt}$$
 Dead time corrected intensity

261	τ	Dead time constant in seconds
262	i_{cps}	Intensity in cps
263		

264 Correction for Beam Drift and Standard Drift

265 Corrections for drift of the electron beam current are generally applied to the quantification of Xray maps due to the frequently extended acquisition times, even for modern instruments with 266 267 beam current regulation. This beam drift correction is usually based on the beam current 268 measurement before and after each map acquisition, and is normalized to the X-ray intensities 269 for each pixel based on the elapsed time between these beam current measurements. See 270 equation 3, which assumes the (slow) increment step motion is in the y-axis direction. Other 271 beam drift correction methods include measuring the beam current before and after each scan 272 line, though this procedure may not be necessary in modern instruments when the total map 273 acquisition time is less than 10 or 20 hours.

274
$$nA_{Interpolated} = nA_{Start} + (nA_{End} - nA_{Start}) * \frac{(P_x + (P_y - 1) * P_x)}{(N_x * N_y)}$$
 (Eq. 3)

275 Where :

276	$nA_{Interpolated}$	Interpolated mapping beam current
277	nA _{Start}	Starting mapping beam current
278	nA_{End}	Ending mapping beam current
279	P_{x}	Position of a pixel the x array dimension

280	P_y	Position of a pixel the y array dimension
281	N _x	Number of pixels in the x array dimension
282	N _y	Number of pixels in the y array dimension

Similarly, if standard intensities were measured before and after the X-ray map acquisition, a drift correction in the standard intensities can be applied (equation 4). This drift corrected standard intensity is calculated on an element by element basis for each pixel and is applied to the denominator of the k-ratio prior to the quantitative matrix correction.

287
$$I_{s} = I'_{s} + (I''_{s} - I'_{s}) \frac{(T_{P} - T'_{s})}{(T''_{s} - T'_{s})}$$
(Eq. 4)

288 Where :

289		Is	Interpolated (drift corrected) standard intensity
290		I'_{S}	Standard intensity from the preceding standardization
291		<i>I''s</i>	Standard intensity from the following standardization
292		T_P	Acquisition clock time of the X-ray map pixel
293		T'_{S}	Acquisition clock time of the preceding standardization
294		T''_{S}	Acquisition clock time of the following standardization
295	Both the beam drift a	nd stand	lard intensity drift corrections assume linear drift in between beam
296	current and standard i	ntensity	measurements, and of course are also dependent on the pixel
297	acquisition order whi	ch depe	nds on the map acquisition (slow) step increment direction, and the
298	instrument vendor as	previou	sly discussed.

299

300 Aggregate Intensities for Duplicate Elements

301 Sometimes we need to improve the geometric efficiency of our photon collection by measuring 302 the same element on multiple WDS spectrometers simultaneously (Donovan et al., 2011). 303 Different spectrometer/crystal combinations can be used as long as the same X-ray line emission 304 is utilized for each duplicate element. By aggregating intensities from several spectrometers, the 305 analytical sensitivity and thus the detection limit can be significantly improved. The intensities 306 for both the standard and the unknown are aggregated prior to the construction of the elemental 307 k-ratio (see section Trace Elements and Spectrometer Aggregation below for details). 308 309 Correction for X-ray Continuum 310 WDS background (i.e., X-ray continuum) measurement is central to accurate quantitative 311 microanalysis, especially for minor and trace elements. As mentioned previously, one or two 312 background maps can be acquired to subtract the interpolated background from the on-peak map 313 pixels, at the cost of at least doubling the total mapping time. Alternatively, the mean atomic 314 number (MAN) background method offers another background calibration method and provides 315 a number of benefits for microanalysis, especially with regards to acquisition time and precision, 316 particularly for X-ray mapping. Some investigators (Tinkham and Ghent, 2005), have utilized a 317 single MAN standard for X-ray map background corrections, but this is insufficiently accurate 318 for maps containing multiple phases with different average atomic numbers. However, by 319 measuring the X-ray intensity at the characteristic peak position for a given element on multiple 320 standards (that do not contain the element), and which includes the range of average atomic

number for the unknown phases to be mapped, we can create a robust relationship between these
continuum intensities and average atomic number of these standards. Such MAN standard plots,
usually acquired during the primary standard intensity calibration, yield an accurate WDS
background calibration, provided that a correction for continuum absorption is properly applied
to the measured intensities for each standard (Donovan et al., 2016).

326 The X-ray intensity data collected for these MAN plots are generally acquired at the same 327 conditions as any other standard acquisition, and therefore have a precision that is equal to that 328 for conventional spot analysis. For compositional mapping applications, the MAN background 329 correction also provides an accurate background calibration that is used for all pixels in the map 330 and at a precision that is superior to the counting statistics of conventional "off-peak" acquired 331 background maps. In most silicate and oxide materials, the absolute accuracy of the MAN 332 background correction is typically around 200 to 300 ppm, which is usually smaller than the 333 expected precision of most X-ray map pixel intensities. The greatest benefit of the MAN 334 background method is that all counting time in the mapping run is dedicated to measurement of 335 the characteristic X-ray peak, which significantly improves the precision and detection limit of 336 the element (Donovan et al, 2016). Since the MAN method is more time efficient and actually 337 provides better sensitivity, it is generally the preferred method for background correction of X-338 ray maps. Although MAN background calibrations might be slightly less accurate than using off-339 peak maps for background corrections, for example when phases contain unanalyzed (or 340 unspecified) elements such CO_2 in carbonates (depending on the details of the differences in 341 average atomic number), utilizing MAN background corrected X-ray maps does reduce the 342 necessity for repeated map acquisitions, thus also reducing concerns regarding stage 343 reproducibility and instrument stability. In addition, accuracy is further improved because the

MAN background method completely avoids the problem of spectral interferences on off-peak intensity measurements from other emission lines, since these off-peak intensities are no longer acquired when utilizing the MAN method. Finally, if necessary, a blank correction (Donovan et al., 2011) can be applied to the X-ray maps, if the map matrix is suitable, for even better accuracy.

349

350 Correction of Spectral Interferences

351 Correction of spectral interferences are applied to the pixel on-peak intensities during the matrix

352 correction iteration based on the concentrations of the interfering elements (Donovan et al.,

1993). This correction must be done iteratively with the matrix correction (Fig. 2; equation 5) as the interference correction can change the composition of the pixel significantly enough so that the matrix correction must be re-calculated based on the newly re-calculated concentrations of all elements; this is even more important when considering mutual interferences of two analyzed Xrays (e.g., Ti Kβ interferes with V K α , and V K β interferes with the trace levels of Cr K α in Al-Ti-V-Cr alloy).

$$359 \quad C_A^u = \frac{C_A^s}{[ZAF]_{\lambda_A}^s} \left[ZAF\right]_{\lambda_A}^u \frac{I^u(\lambda_A) - \frac{[ZAF]_{\lambda_A}^s}{C_B^{\bar{s}}} \frac{C_B^u}{[ZAF]_{\lambda_A}^u} I_B^{\bar{s}}(\lambda_A)}{I_A^s(\lambda_A)} \right]$$
(Eq. 5)

360 Where:

361 C_i^J Concentration of element i (A = measured element; B = interference 362 element) in matrix j

363	$[ZAF]^J_{\lambda_i}$	ZAF or $\phi(\rho z)$ correction term for matrix j at λ_i for element i
364	$\mathrm{I}_{i}^{j}(\lambda_{i})$	Measured X-ray intensity in matrix j at wavelength λ_i
365	u	Unknown sample which contains elements A and B
366	S	Primary standard which contains only element A
367	$\overline{\mathbf{S}}$	Interference standard which contains a known quantity of the interfering
368		element B, but none of the interfered with element A.

369

When necessary, this quantitative spectral interference correction routine is applied to each pixel of the acquired maps (and the standard point analysis intensities also), providing of course that the X-ray intensities of both the measured and the interfering elements are available.

373

374 Correction of Time Dependent Intensity Effects

Correction of time dependent intensity (TDI) effects, such as alkali-migration in glasses due to beam damage (cf. Nielsen and Sigurdsson, 1981; Morgan et al, 1996), is quite straight forward for point analysis, as it simply requires acquisition of "sub-interval" intensities during the onpeak measurement to monitor and correct for change in count rate over time. The same principle can be applied on beam mapping: by acquiring multiple frames of X-ray maps, we can obtain an evolution of the count rate for each pixel over time, and thus correct for possible changes in Xray intensity by extrapolating back to the intensity at time zero using equation 6.

₃₈₂
$$I_C = e^{Log(I_R) - m \cdot t \cdot 0.5}$$
 (Eq. 6)

383 Where:

384	I_R	X-ray intensity in counts per second per nA
385	m	Time Dependent Intensity (TDI) slope coefficient from linear fit of log
386		intensity versus time
387	t	Total elapsed integration time
388		
389	Linear, quadratic, o	r logarithmic TDI fitting models can be applied on an element-by-element
390	basis; quadratic or	logarithmic corrections are often necessary for extremely beam sensitive

391 materials (von der Handt, in preparation).

392

393 Correction of Compositional Matrix Effects

Once the k-ratios for each element are calculated from the normalized and background corrected unknown map pixel and standard point intensities, a matrix correction can be applied. Selection of the proper matrix correction routine is out of the scope of this paper and the analyst will be the judge of that choice. Typical analytical physics models can be used (e.g., ZAF or $\phi(\rho z)$) or even Monte Carlo methods (Donovan, 2019). In general, quantitative analysis using the electron microprobe is generally based on equation 7:

400 C = k * [ZAF] (Eq. 7)

401 Where C is the concentration result, [ZAF] is the total matrix correction, including the standard 402 k-factor correction for when the standard is not a pure element (e.g., Armstrong, 1988), and k is 403 the raw k-ratio intensity expressed as $(I)^{sample} / (I)^{standard}$, where I are the continuum corrected

404	peak X-ray intensities measured on samples and standards respectively. Modern matrix
405	correction algorithms correct for electron energy-loss and backscattering in the atomic number
406	factor Z, absorption of characteristic X-ray intensities in the absorption factor A (both which are
407	combined in $\phi(\rho z)$ matrix correction methods), and additional X-ray intensities resulting from
408	characteristic (and ideally continuum) X-ray fluorescence in the fluorescence factor F. This
409	matrix correction is made iteratively using a refined estimate of the pixel concentration. The final
410	analysis result is identified when no further improvement in the concentrations is obtained, just
411	as in the case of single point analysis. In fact we utilize the same source code for both point
412	analysis and map pixel analysis in all our results.
413	
414	Additional Corrections for Soft X-rays
414 415	<i>Additional Corrections for Soft X-rays</i> Not discussed in this paper are the effects of changes in peak shape on soft X-rays and the
414 415 416	<i>Additional Corrections for Soft X-rays</i> Not discussed in this paper are the effects of changes in peak shape on soft X-rays and the correction algorithm using area peak factors (Bastin and Heijligers, 1991). This approach is
414 415 416 417	Additional Corrections for Soft X-rays Not discussed in this paper are the effects of changes in peak shape on soft X-rays and the correction algorithm using area peak factors (Bastin and Heijligers, 1991). This approach is applied for the accurate quantification of low atomic number emitters. However, due to the
 414 415 416 417 418 	Additional Corrections for Soft X-rays Not discussed in this paper are the effects of changes in peak shape on soft X-rays and the correction algorithm using area peak factors (Bastin and Heijligers, 1991). This approach is applied for the accurate quantification of low atomic number emitters. However, due to the chemical bonding dependency of such corrections, they are of limited use in X-ray mapping
 414 415 416 417 418 419 	Additional Corrections for Soft X-rays Not discussed in this paper are the effects of changes in peak shape on soft X-rays and the correction algorithm using area peak factors (Bastin and Heijligers, 1991). This approach is applied for the accurate quantification of low atomic number emitters. However, due to the chemical bonding dependency of such corrections, they are of limited use in X-ray mapping when multiple phases are present. Nonetheless, it should be mentioned that the choice of
 414 415 416 417 418 419 420 	Additional Corrections for Soft X-rays Not discussed in this paper are the effects of changes in peak shape on soft X-rays and the correction algorithm using area peak factors (Bastin and Heijligers, 1991). This approach is applied for the accurate quantification of low atomic number emitters. However, due to the chemical bonding dependency of such corrections, they are of limited use in X-ray mapping when multiple phases are present. Nonetheless, it should be mentioned that the choice of different matrix correction models and mass absorption coefficients can have significant effects
 414 415 416 417 418 419 420 421 	Additional Corrections for Soft X-rays Not discussed in this paper are the effects of changes in peak shape on soft X-rays and the correction algorithm using area peak factors (Bastin and Heijligers, 1991). This approach is applied for the accurate quantification of low atomic number emitters. However, due to the chemical bonding dependency of such corrections, they are of limited use in X-ray mapping when multiple phases are present. Nonetheless, it should be mentioned that the choice of different matrix correction models and mass absorption coefficients can have significant effects on the accuracy of quantitative results and should be considered when performing quantitative
 414 415 416 417 418 419 420 421 422 	Additional Corrections for Soft X-rays Not discussed in this paper are the effects of changes in peak shape on soft X-rays and the correction algorithm using area peak factors (Bastin and Heijligers, 1991). This approach is applied for the accurate quantification of low atomic number emitters. However, due to the chemical bonding dependency of such corrections, they are of limited use in X-ray mapping when multiple phases are present. Nonetheless, it should be mentioned that the choice of different matrix correction models and mass absorption coefficients can have significant effects on the accuracy of quantitative results and should be considered when performing quantitative X-ray mapping with highly absorbed emission lines, such as low atomic number K-lines or other

425 Computation Time and Complementary Calculations

426 The processing of a typical set of 10 element X-ray maps of 512 x 512 pixels will require tens of 427 minutes to a couple hours of matrix physics calculation time depending on the speed of the 428 computer; most modern computers will process the data of smaller maps with fewer elements in 429 less than 10 minutes. Elemental concentrations, along with background-, net-, and k-ratio 430 intensities are generally calculated by default. An analytical total weight percent map is also 431 usually calculated to further help in judging the data quality. In addition, other information can 432 be mathematically extracted from the elemental concentration data, such as oxide weight-433 percent, atomic proportions, formula basis calculations, and logarithmic weight-% of 434 concentrations for trace element mapping. Statistical data such as detection limits and analytical 435 sensitivities can also be calculated for each pixel as well. Optionally, and if appropriate for the 436 compositions in question, several other map data can be extracted, such as a stoichiometric 437 oxygen map, an excess oxygen map (from assumed oxygen stoichiometry and/or measured 438 anions such as F and Cl), and a map of an element calculated by difference from 100%. To 439 facilitate the calculation of all specified output types it is more efficient to calculate them during 440 the map quantification, although they can be easily recalculated using the primary output 441 (elemental weight-%).

Because all pixels during the quantification of X-ray maps are treated to the same matrix
correction methods as we traditionally perform for point analyses, very large maps with many
elements can require considerable time for the calculation of concentrations. The calculationtime scales linearly with the number of pixels and exponentially with the number of elements
analyzed. The fully automated processing of these quantitative maps can be done off-line on any
computer without any action from the user besides the selection of the appropriate quantification

- 448 parameters (e.g., choice of primary and interference standards, definition of the MAN
- 449 background curves, calculated elements [by stoichiometry, by difference, etc.]). All maps
- 450 presented in this paper follow the procedures as presented above and were performed using the
- 451 Probe Software Probe for EPMA and CalcImage software packages
- 452 (https://probesoftware.com/), though other software codes could perform similar quantitative
- 453 corrections to raw X-ray map intensity data.
- 454

455 **Results and Discussions**

456 Primary standards required for quantitative mapping are usually acquired as point analyses at a 457 moderate beam current (typically 10-50 nA) with integration times of 10-30 s. However, stage or 458 beam scanned X-rays maps are usually acquired at higher currents, typically 20 to 500 nA, even 459 1 μ A in cases of trace element mapping, and using a pixel dwell time on the order of tens to 460 hundreds of milliseconds or more per pixel (which makes it essential that the actual beam 461 currents be recorded for both the standards and the X-ray map acquisitions in real-time). Since 462 both time and beam current should scale proportionally, doubling the count time is equivalent to 463 doubling the beam current (ignoring dead time effects). To compare data and evaluate the 464 precision of each measurement, the total electron dose (i.e., the product of beam current and 465 time) should be considered. Ignoring the effects of electron beam energy in beam power 466 calculations, we can simply say that single point analyses, for instance in standards, typically 467 yield a total electron dose in the range of 100 to 3000 nA s. Quantitative element maps runs 468 typically utilize higher currents (20-500 nA) to compensate for shorter count times (20-1000 ms), 469 which translate to a dose range of 0.4 to 500 nA \cdot s per pixel. In other words, even well-chosen

470 conditions for element maps can still represent an electron dose several orders of magnitude less 471 than for single-point quantitative analysis. This dose should be carefully considered with regards 472 to the beam size, and the applied dose should be normalized to the impacted surface (i.e., 473 expressed as $nA \cdot s/\mu m^2$). This calculation of electron dose is useful for evaluating the precision 474 and detection limit for conventional single-point versus map-pixel analysis and reveals that 475 mapping measurements have usually lower precision and therefore higher detection limits at 476 each individual pixel (Carpenter et al., 2017). It is therefore legitimate to ask how precise and 477 accurate quantitative element mapping truly is. Calculation of the electron dose can also be 478 useful to evaluate the maximum allowable dose for beam sensitive materials before beam 479 exposure damages the sample too much, resulting in inaccurate results. In such cases, large area 480 stage mapping is often done with a relatively large beam size that matches the pixel size to 481 further help dissipating the electron dose over a large area, thus minimizing the problem of beam 482 damage.

483

484 Evaluation of Accuracy

It is not immediately obvious that X-ray maps acquired using a dwell time per pixel of a few hundred milliseconds or even less, can be accurately quantified due to the apparently poor counting statistics per pixel. But accuracy depends primarily on matrix corrections, background and peak interference corrections, standard reference values, etc. All these parameters remain the same whether a single-spot analysis or quantitative element map pixel is considered, and therefore the map-derived element concentrations should be just as accurate as a single point measurement, with the caveat that the absolute accuracy of the pixel dwell times (for both

492 "continuous" and "step" mapping modes) at the millisecond level should be verified by both the 493 instrument vendor and user, for best quantitative mapping accuracy. In addition, when the 494 primary standards and X-ray maps are acquired at significantly different beam currents, the 495 linearity of the beam current picoammeter is critical for best accuracy, not to mention the 496 importance of the WDS detector dead time calibrations when high X-ray intensities from high 497 mapping currents are combined with standard intensities acquired with lower beam currents for 498 the primary standard calibrations, in the subsequent construction of the raw k-ratio prior to 499 quantification. 500 In any case, all quantitative measurements require an average and a variance regardless of 501 whether this is for point or pixel measurements. Similar to point analyses, a representative 502 sampling of the specimen composition requires the examination of an average and standard 503 deviation of each set of points or of pixels, especially in multi-phase materials. 504 To demonstrate the accuracy of the proposed quantitative mapping approach, a simple X-ray 505 map in the reference glass material NIST K-411 is quantified (Fig. 3). For this example, element 506 maps were acquired using a Cameca SX-100 electron probe microanalyzer at CAMCOR 507 (University of Oregon), using analyzing crystals LIF for Fe Ka, PET for Si Ka, LPET for Ca 508 Kα, TAP for Al Kα, and LTAP for Mg Kα. Analytical conditions were 15 keV, 30 nA using a 509 128 x 128 pixels stage scan with a pixel dwell time of 500 ms and a pixel size of $0.5 \,\mu m$. Results 510 of the quantification process are given in Figure 3. In this map quantification, oxygen was 511 calculated by stoichiometry with an additional 1.12 wt% oxygen from ferric iron (as determined 512 from colorimetry). All maps are corrected for background using the MAN background 513 correction. The calculation of a total concentrations map (g) provides for evaluation of the 514 analysis accuracy as all elements should total close to 100%. The calculated quantitative pixel

515	average results are subsequently compared to NIST published values (Table 2). Excellent
516	accuracy is achieved as the calculated weight-% of all four major elements (Si, Mg, Ca, and Fe)
517	are statistically equivalent to the certified NIST values, and Al accurately yields a zero-value
518	(nominally Al-free glass).

519

520 Evaluation of Precision

521 The shorter pixel integration times of X-ray mapping compared to point analysis generally leads 522 to poorer Poisson statistics. Therefore, the apparent precision per pixel is more limited, and it 523 should be evaluated carefully. Mapping precision and sensitivity are considerably improved by 524 pixel averaging, as it is commonly done for multiple point measurements in a homogeneous area 525 or by kernelling or binning such as it is already done for instance with EDS maps. Such pixel 526 averaging can provide enough precision for discriminating small differences in concentrations 527 and can ultimately increase the analytical sensitivity. 528 Table 3 compares point analyses (average of 4 points) of standard material NIST K-411, with 529 pixel averages from the quantitative X-ray map (Fig. 3), using pixel aggregates from 4 to 64 530 pixels. Comparing pixel averages with point averages reveals that accuracy and standard 531 deviation (pixel variance) are consistent and in close agreement with the average and variance of 532 4-point averaged measurements. Naturally, the error of the mean or standard error (pixel error), 533 improves significantly with the number of pixels being averaged (Table 3).

535 Bad Pixels

536 When considering pixel averaging in an unknown sample we must also consider the adverse 537 effects from so-called "bad pixels" that can degrade our quantitative maps with inaccurate data. 538 A bad pixel is often the consequence of breaking one of the essential requirements of WDS 539 measurement (flat and horizontal surface, homogeneous domain, proper working distance, etc.) 540 or a hardware issue (e.g., sudden beam current instability, noise on a WDS X-ray detector, etc.). 541 Excluding hardware problems, a typical bad pixel is commonly due to surface defects (holes, 542 cracks, poor polishing, etc.), phase boundaries (Barkman et al., 2013), or finely intergrown 543 materials or exsolutions. The choice of beam diameter and accelerating voltage along with the 544 density of the analyzed phase will control the size of the analyzed volume. When a larger 545 analytical volume is considered, a higher number of bad pixels is to be expected, especially when 546 the average grain-size or feature to be mapped is approaching the size of the analytical volume. 547 In other words, the larger the analytical volume and/or the smaller the phase dimensions, the 548 greater the number of pixels compromised by edge, crack and mixed-phase interaction volumes. 549 Large well-polished phenocrysts might give overall excellent results, but a fine-grained 550 heterogenous matrix may not. For this reason, the best approach would utilize a focused beam, 551 despite the fact that this would increase the current density and potential for beam damage. The 552 key point is to constrain the analytical volume and the pixel size to ultimately obtain a 553 representative number of pixels for the smallest feature to be mapped while minimizing the 554 number of bad pixels.

555 When considering pixel averaging, including "bad pixels" in the average can affect both the 556 precision and the accuracy. Therefore, the use of pixel extraction and filtering methods are 557 recommended and available in many image processing programs (e.g., Probe Software

558	CalcImage, NIH ImageJ, etc.). This approach allows the analyst to specify the exact pixel area to
559	be extracted from the map by providing for instance, a pixel filtering based on quantitative
560	compositional limits for selected analyzed elements. In this way one can avoid and/or filter out
561	these bad pixels from our compositional averages. An example is provided later in this paper.
562	
563	Quantitative Mapping Results
564	Geologists and materials engineers commonly require detailed element mapping of complexly

565 zoned phases. Whereas point analysis yields very precise and accurate data, it commonly fails at

revealing all the subtleties of compositional zonation (if present). Mapping is therefore

567 necessary, and as discussed above, both precision and accuracy can be obtained depending on the 568 analytical settings and the pixel averaging employed.

As a first example, a simple mapping of plagioclase with 5 major elements (Si, Al, Ca, Na, and 569 570 K) is considered. The sample is from the Adamello tonalite in Northern Italy (Fiedrich et al., 571 2017), and shows strong compositional zoning from an inherited anorthite-rich core to an albite-572 rich rim overgrowth showing a significant oscillatory zoning. After performing the quantitative 573 calculations for each pixel, the fully quantitative elemental or oxide weight-percent (wt%) 574 concentrations are available for output (Fig. 4a). In addition, these elemental concentrations can 575 be expressed in any number of various output types, such as atomic percent or formula basis 576 (Fig. 4b), and various end-member mineral normalizations. Additional normalizations based on 577 the previously calculated elemental or molar concentrations can also be performed. For example, 578 mineral end-member calculations where the Ca, Na, and K concentrations are normalized to 579 produce maps of anorthite, albite, and orthoclase end-member components for each pixel in the

original map (Fig. 4c), though multiple normalizations may be necessary when more than one
mineral phase is present.

582 If a single-phase is of interest to the researcher, pixel filtering and extraction based on chemical 583 composition and the analytical total can be applied. In Figure 4d, only the pixels corresponding 584 to a plagioclase composition are extracted; the conditions in this example are: 90% < Total < 110585 wt%, $40 < SiO_2 < 75$ wt%, $15 < Al_2O_3 < 40$ wt%, CaO < 25 wt%, $Na_2O < 15$ wt%, and $K_2O < 2$ 586 wt%. This compositional filtering also allows us to remove any of the previously discussed bad 587 pixel effects to improve accuracy. It is then possible to extract quantitative data using a pixel 588 averaging of multiple pixels of similar composition to obtain a more precise measurement, such 589 as discussed previously. It is also possible to work on other statistical data such as histograms of 590 composition to delineate for instance, the modal abundance of an anorthite-rich core versus an 591 albitized domain (Fig. 4e).

592 It is also possible to calculate the detection limits and analytical sensitivity just as it is done for 593 single point analyses (Fig. 5). At the analytical conditions considered here (50 nA, 50 ms per pixel), the per pixel calculated analytical sensitivity (~2-5% for major, >10-20% for minor 594 595 element) and detection limit (typically between 0.06 and 0.15 wt% in this example) remains of 596 course high, when compared to point analysis, but can be further improved through pixel 597 aggregation or with longer counting time, higher beam current, or more optimal spectrometer 598 selection (e.g., large area monochromator, P-10 vs. Xe counter). For cleanliness of the output, 599 the analytical sensitivity calculations can be skipped for any pixel yielding a concentration <1 600 wt%, as it usually yields very high and meaningless values (gray pixels).

602 Comparing Raw Intensity Maps to Matrix Effect Corrected Maps

603 Traditionally, raw intensity WDS and EDS X-ray maps were used for the documentation and 604 interpretation of chemical variation in solid materials. For several decades, EDS X-ray maps 605 have been quantitative (at least when standards are utilized), because it was relatively easy to do 606 so. While the importance of matrix corrections in WDS spot analyses is universally 607 acknowledged, lack of computing power and software development has posed a hindrance in 608 extending robust corrections to WDS X-ray maps. However, in situations that require significant 609 absorption correction for X-ray intensities, substantial discrepancies can be observed between 610 raw and quantified maps. Figure 6 is one example for such a case, showing a Fe-Ti oxide grain 611 with Fe-rich core and exsolutions of ilmenite (FeTiO₃) and titanite (CaTiSiO₅). 612 The raw X-ray map of oxygen (Fig. 6a) would suggest relatively constant oxygen contents across 613 the mineral grain, while the fully quantified map (Fig. 6b) shows pronounced differences in 614 oxygen concentrations. Maps of net and background intensities (Fig. 6c,d) calculated by MAN 615 show a slight increase in background intensities in the core but a negligible difference in net 616 counts due to the high peak-to-background ratios for oxygen across all phases. Therefore, neither 617 a background-corrected map by itself, nor a quantification using a calibration curve, would 618 accurately reveal the varying oxygen contents. 619 In this example, emitted O K α X-ray intensities in the Fe-Ti oxides are dominantly controlled by 620 the changes in mass absorption of O K α by Fe, Ti, or O itself. Mass absorption coefficients for O 621 in Ti and Fe, respectively, differ by almost an order of magnitude (Ti: 22422 vs. Fe: 3625; 622 FFAST database; Chantler, 2000; Chantler et al. 2005). Consequently, for a given oxygen wt%

623 content, areas high in Fe will show higher relative O K α intensities, whereas Ti-rich areas will

show lower X-ray intensities as they will be absorbed more. Only a full quantification, including a background and matrix correction (and interference correction), can adequately account for this change in X-ray emission rates. It is therefore essential that the analyst rigorously quantify each pixel using the methods described in this paper, so that they can be confident that differences in concentrations and zoning are not masked or introduced by variations in background intensities, absorption and fluorescence effects between grains or mineral domains.

630

631 Trace Element Quantification in X-ray Maps

632 Rigorous quantification is essential in order to facilitate accurate evaluation of X-ray intensity maps of elements with trace concentrations (>10 to 1000ppm). The crucial point for accurate 633 634 trace analysis is the correction for the X-ray continuum contribution to the measured X-ray 635 intensity at the Bragg conditions for the characteristic X-ray of interest (see Jercinovic et al., 636 2012; Allaz et al., 2019a), because continuum intensity varies as a function of the mean atomic 637 number of a given phase, and also from absorption effects associated with the composition of the 638 phase. An accurate EPMA trace element analysis, whether it is by single point analysis or by 639 pixel mapping, can only be accomplished with an accurate and precise background correction. 640 This means that the continuum must be accurately calculated or measured, for each pixel, with a 641 precision sufficient to answer the analytical questions at hand. 642 Bremsstrahlung correction can be problematic for multi-phase element mapping, where each 643 phase has a potentially different average atomic number. It can also be problematic for a single 644 phase that displays a significant compositional zonation inducing a significant change in average

645 Z-number (e.g., Z-bar ~10.7 to 11.9 between albite and anorthite). This issue gets even more

646 complicated when multiple substitutions are possible in a strongly heterogeneous phase, and 647 when both light and heavy elements are substituting (e.g., in zircon: $Zr \iff$ [Hf, U, Th], Si+Zr 648 \iff P+REE, etc.).

649 As previously described, when acquiring EPMA trace element point measurements, typically the 650 background X-ray intensity is sampled by de-tuning each WDS spectrometer to both the high 651 and low sides of the characteristic X-ray peak, implying at least a doubling of the acquisition 652 times. Fortunately, we can also apply the MAN background correction to X-ray maps, just as it is 653 applied to point analyses. Once a MAN background calibration curve is defined for each element 654 to be mapped, there is no need to measure a background map. Figure 7 compares the quantification of trace element maps of Hf La, Y La, P Ka, U Ma, and 655 656 Th M α in zircon using off-peak map acquisitions (employing two WDS spectrometer positions 657 with a linear interpolation of the off-peak map pixel intensities), and also using the MAN 658 background correction, with both methods utilizing the same on-peak map intensities. The ZrO_2 659 and SiO₂ contents were constrained by assuming a mineral formula by difference from 100%. 660 This is accomplished by including the calculated $ZrSiO_4$ (by difference from the measured 661 elements) into the matrix correction during the matrix iteration calculations so that an accurate 662 matrix and average atomic number (for the MAN interpolation) is obtained. In addition, it should 663 be noted that a correction for the interference of Th M β on U M α should normally be applied. 664 However, in this example, the Th content is barely above detection limit and therefore a spectral 665 interference correction is unnecessary in this particular case. 666 Somewhat unintuitively, the use of MAN background calibrations actually improves precision

because the background intensity calculation is no longer limited by continuum statistics, but is

668 instead based on our major element concentrations, which solely determine the average atomic

669 number of the matrix, and hence the precision (and accuracy) of the absorption-corrected MAN 670 background curve. Both background correction routines yield equally accurate results, yet the 671 MAN background corrected maps (Fig. 7a) yields more precise results with less noise (especially 672 visible on U and Th maps). These MAN background calibrations can result in accuracy errors for 673 trace elements on the order of a few hundred ppm in typical geological materials, though this can 674 be accurately compensated for by use of the "blank" correction (Donovan et al., 2011). However, 675 the pixel level measurement precision in our X-ray maps is limited by the pixel dwell time (and 676 beam current), and in most situations it is significantly worse than the accuracy obtained with the 677 MAN background correction.

Even with a cursory visual examination of the off-peak and MAN background corrected trace elements maps, the MAN-corrected quantitative element maps clearly provide improved precision/sensitivity compared to the off-peak background corrected maps, without the need of doubling or tripling the acquisition time in order to obtain background maps. Another advantage of the MAN correction is that is reduces the risk of acquiring a background map at a wavelength (off-peak) position that could be interfered with by another major or minor element present only in a few phases or in a specific zone of the area to be mapped.

The higher sensitivity seen in the MAN background corrected maps is due to the fact that when the on-peak intensities are corrected for background, the variances of the on and off-peak intensities are added in quadrature:

688
$$\sigma_{P-B} = \sqrt{\sigma_P^2 + \sigma_B^2}$$
(Eq. 8)

In practice, this means that as the measured concentration approaches zero, and therefore the
 variances of the on-peak and off-peak become similar, the precision of the background intensity

691 becomes as important as the precision of the on-peak measurement. In a classical two-point 692 background acquisition, the error on the two-background measurement is limited to only 693 continuum counting statistics, which due to the relatively low intensity, results in a relatively low 694 precision background correction. On the other hand, because the error on a MAN-calibrated 695 background correction is limited primarily by the measurement precision of the major elements 696 (e.g., the variance of the average atomic number of the material from the high concentration 697 elements), the variance of the interpolated background intensity is much smaller (see Donovan et 698 al., 2016 for a discussion of these statistical considerations). In fact, typically the resulting 699 precision of the MAN net intensity background is approximately 40% better than traditional off-700 peak corrections (by the square root of two), because the background variance term in Eq. 8 701 (σ_R^2) approaches zero, particularly when the major matrix elements are specified by difference or 702 fixed concentration). Therefore it is primarily only the precision of the on-peak measurement term (σ_P^2 in Eq. 8), which contributes to the net intensity precision, when the MAN background 703 704 method is utilized.

705

706 Correction of Spectral Interferences in X-ray Maps

Correction of spectral interferences is also particularly critical for accurate quantitation of trace elements where a trace element X-ray emission line is overlapped by a significant emission line from another element. This problem is illustrated by obtaining X-ray maps for S K α , Fe K α , Cu K α (Fig. 8a-c), and Mo L α (Fig. 8d-f) from a sample containing the phases pyrite (FeS₂) and chalcopyrite (CuFeS₂). These maps are followed by the extraction of a cross-section A-B across the pyrite-chalcopyrite interface (Fig. 8g, h). The Mo contents of both sulfides are expected to be

713	at the trace level, most likely below the detection limit of the map measurement sensitivity. Yet,
714	the raw counts (Fig. 8d) apparently reveal the presence of significant Mo. This apparent Mo
715	signal mostly comes from the strong interference of Mo L α by S K α on a PET monochromator,
716	and therefore a correction for spectral interference is essential for Mo accuracy. Without a peak
717	interference correction (Fig. 8e), up to 1.8 and 1.2 wt% Mo is erroneously measured in pyrite and
718	chalcopyrite, respectively. When a quantitative interference correction is applied (Donovan et al.,
719	1993), data essentially yield values at or below detection limit (~700 ppm at the mapping
720	conditions utilized), as seen in Figure 8f.
721	Interestingly, visualization of the spatial distribution of an element that has a large compositional
722	range, can be also problematic. For example, the Cu element map in Figure 8c shows levels of
723	Cu approaching zero in pyrite (phase on left), but around 33 wt% in chalcopyrite (phase on
724	right). One solution to this visualization problem is the use of logarithmic weight percent values
725	to scale the element map. Figure 9 shows the application of a logarithmic scale to the Cu wt%
726	values and subsequently elucidates an apparent increase in Cu up to $\sim 1 \text{ wt\%}$ in pixels adjacent to
727	the chalcopyrite boundary (Fig. 9b); an observation that otherwise might have been missed. In
728	this specific example, the apparent presence of Cu at the grain boundary is not related to
729	diffusion, but is rather an artifact from secondary fluorescence effect from the bremsstrahlung
730	(e.g., Borisova et al., 2018).

731

732 Quantification of Beam Sensitive Samples or TDI Scanning

733 The electron beam can cause many materials to suffer from ion diffusion, recrystallization,

amorphization, or even volatilization (e.g., loss of OH, H₂O, or CO₂ groups) chiefly due to the

heat (phonon excitation) generated by incident electron interaction with matter (Hughes et al., 735 736 2019). In such cases, where the beam current cannot be reduced or defocused, and a high thermal 737 conductivity coating (e.g., Al, Ag or Ir), is not appropriate, it is recommended to apply a time-738 dependent intensity (TDI) correction that can account for the ion migration and/or a change in 739 signal due to physical changes in the analyzed material, provided it follows a linear, exponential, 740 or polynomial trend over time. These so-called beam sensitive materials can be extremely 741 difficult to analyze using single-point analysis due to the limited analytical time available before the material is so badly damaged that even a TDI correction cannot account for the change (e.g., 742 743 Allaz et al. 2019b). 744 Among the classic examples in the literature are, for instance, phosphates (e.g., Goloff et al. 745 2012; Fialin and Chopin, 2006), alkali-rich phases (e.g., Gedeon et al., 2008), and carbonates 746 (e.g., Zhang et al., 2019). One extreme example concerns hydrated alkali-sulfate (natro-) jarosite 747 and (natro-) alunite (K,Na)(Al,Fe)₃(SO₄)₂(OH)₆ displaying fine compositional variation at the 748 scale of ~1-5 µm. Irremediable Na or K diffusion and a concomitant increase or decrease in Fe, 749 Al, and S happens within less than 10 seconds of analyzing a single point at less than 2 nA with a 750 slightly defocused beam (2-3 μ m). Beyond this, the beam damage induces a large change in the 751 major elements and a TDI correction must be applied. In this case, it is possible to run analyses 752 at 10 nA for 20 s using a focused electron beam, and to correct for the beam damage effects. 753 Beyond this point, the damage may become so severe that an accurate TDI correction cannot be applied. Notwithstanding a total electron dose of only 200 nA s, and with the use of the TDI 754 755 correction, the analytical precision is typically poor. Multiple point analyses in a homogeneous 756 domain are then required to obtain a good precision, which can be problematic when the material

757	to analyze is finely zoned at the micron-scale and in very small grains (typical grain-size in the
758	sample under investigation is $\sim 20 \ \mu m$; Potter-McIntyre and McCollom, 2018).

759 Another solution to obtain equally precise and sometimes better accuracy quantitative analyses 760 of such complexly zoned beam sensitive materials is to acquire element maps and apply a TDI 761 correction during the map quantification. To do this, a time series of beam (or stage) scanned 762 maps over an area are acquired with multiple (replicate) mapping frames. The TDI correction can then be applied on each individual pixel using these multiple frames to extrapolate the pixel 763 764 intensity at time t = 0 s (beginning of the mapping elapsed time for each individual pixel). Figure 765 10 is an example of such TDI-corrected element maps of Na-free jarosite (see Potter-McIntyre 766 and McCollom, 2018 for detail on sample MN05, notably their Fig. 10). Multiple test maps for K 767 K α , S K α , Fe K α , Al K α and As L α are run using either 50 or 100 nA beam currents, different 768 counting times (5, 10, 20, and 40 ms per pixel on one frame) and total number of frames (8 769 frames at 5 ms, 5 frames for all others), resulting in a variation of total beam dose of 2 to 20 770 nA·s. Due to the beam damage induced by each test, it is not possible to re-analyze the same 771 area. For the quantification of each map, the MAN background method is used, and the water-772 content is recalculated by stoichiometry assuming 6 hydrogen atoms for a total of 14 oxygens 773 and included in the iterated matrix correction routine (Fig. 2). Results obtained with and without 774 the TDI correction are then compared, along with a comparison of data obtained on the first and the last frame (labelled respectively "TDI", "Aggregated", "First", and "Last" in Figure 10). 775 776 At low beam dose (total of 2 nA s per pixel; Fig. 10a), both the TDI-corrected maps and the 777 aggregated maps yield identical results within their respective errors. It is clear from visual 778 inspection and counting statistics that the TDI results are slightly less precise due to the nature of 779 the TDI correction extrapolation. At intermediate to high beam doses (e.g., 4 and 10 nA·s; Fig.

780	10b,c), the results of the aggregated maps degrade: zones are getting enriched or depleted in
781	K_2O , and in parallel SO ₃ and Fe ₂ O ₃ either decrease or increase at high beam dose, and overall the
782	analytical totals increase. This can be clearly seen numerically in the averaged-pixel quantitative
783	data (Fig. 10d), especially when comparing the data from the first and the last maps. It is also
784	interesting to notice that there is no systematic diffusional loss of K, rather there are "sink and
785	hole" areas where K diffuses in or out. This effect is illustrated in Figure 11 with a map of the
786	TDI correction percentage for the measurement of K Ka, S Ka , and Fe Ka over five passes at
787	100 nA and 20 ms per pixel (see also Fig. 10c).

788

789 Trace Elements and Spectrometer Aggregation

790 Quantification of multiple phase X-ray maps can be easily done with great precision and 791 accuracy for major and minor element analyses, but the counting statistics usually prohibits 792 mapping for trace elements without excessively long acquisition times at high currents. If a high 793 current is not applicable (e.g., beam sensitive materials) or if the analyst aims at reducing the 794 total acquisition time, it is recommended to use multiple WDS detectors in parallel to analyze a 795 single element. Figure 12 presents an X-ray map of a multiphase area in a rhyolite sample 796 composed of glass, plagioclase, and K-feldspar. Previous single-point analyses and observations 797 in BSE images suggest a strong Ba-enrichment at the rim of K-feldspar. Maps were acquired in 798 two passes, with a first pass for the major elements (Si K α , Al K α , Ca K α , Na K α , and K K α), 799 and a second mapping pass with only Ba L α analysis on three PET-L spectrometers. Of the 800 major elements expected in this sample, only Fe was not mapped, as it is a trace concentration in 801 these specific feldspars, and only present at level around 2.5 wt% in the glass. All other elements

802	(Mg, Ti, Sr, etc.) are below 1000 ppm. The presence of Ti could potentially be a problem for Ba
803	analysis due to the mutual interference of Ti K α on Ba L α . However, single-point analyses
804	including a correction for this mutual Ti K $\alpha \ll$ Ba L α interference yield low Ti-content: <3000
805	ppm in glass and <500 ppm in feldspar minerals. At this level, a correction for Ti interference on
806	Ba is minimal and below the sensitivity of Ba L α element map (calculated Ba detection limit of
807	\sim 500 ppm). These duplicate element maps allows one to not only highlight and quantify the
808	high-Ba content in K-feldspar, but also to accurately and precisely analyze Ba even in the glass
809	(~1200 ppm). These element maps were essential for the data interpretation, as only the maps
810	clearly reveal a two-step growth. The primary feldspar, which is richer in albite (Na)-component,
811	suffered a partial resorption (e.g., irregular internal zonation in K and Ca, top-right corner of Fig.
812	12c), and the Ba-rich K-feldspar rim and the anorthite(Ca)-richer plagioclase rim both represent
813	a second magmatic event.

814

815 **Implications**

816 Rigorous quantification of X-ray maps is critically important for accurate interpretation of 817 element distributions, not only for minor and trace elements, but also elements whose emission 818 lines suffer significant absorption or fluorescence. Raw X-ray maps remain an essential tool to 819 reveal two-dimensional variation that cannot be fully ascertained by discrete point analysis. 820 However, because the same raw X-ray line intensity can represent quite different concentrations 821 in different phases, due most notably to significant differences in X-ray absorption effect, but 822 also from changes in background intensities (and potential issues stemming from spectral 823 interferences), without quantification at the pixel level it would be significantly more risky to

draw robust conclusions from merely qualitative X-ray maps. Moreover, this rigorous
quantification protocol ensures that all maps can be compared one-to-one, even if they were
acquired at different times, using different spectrometers, in different laboratories, or at different
analytical conditions, etc.

828 Applications of element map analysis by EPMA (WDS or EDS) are numerous, and the ability to 829 quantify them with accuracy and precision makes these applications even more attractive to 830 researchers. Firstly, geological materials are often the product of a complex mineral growth, 831 partial dissolution (resorption), and reprecipitation history, the interpretation of which is critical 832 to developing an understanding of large-scale geological processes. Up until now, element maps 833 were often only used for qualitative observations (e.g., identifying inherited core, retrogression, 834 diffusion profile, etc., which subsequently required quantification by careful single-point 835 analysis) or for approximate quantification (e.g., using a two-point calibration method in a single 836 compositionally zoned phase). However, with the level of accuracy and precision achieved with 837 the proposed map quantification methodology, the qualitative map data becomes quantitative 838 data, with an accuracy comparable to a single point analysis, and with an equally good precision 839 when pixel averaging is considered. Secondly, material scientists can now observe accurately 840 and with high sensitivity, diffusion profiles, identify and quantify exsolutions, or perform 841 homogeneity tests. Additionally, the "analytical totals" map can also uncover previously 842 unrecognized secondary fluorescence, i.e., pixels where the totals are consistently > 100%843 (Fournelle at al., 2005). By extension, a more accurate "local" composition can be obtained in 844 finely grained material (e.g., partially re-crystallized glass or quenched material, exsolutions 845 lamellae, etc.), without the need to acquire multiple discrete points on which the resulting 846 accuracy can be questionable, due to the inhomogeneous nature of the analyzed material.

847 Finally, the complementary use of the MAN background correction method allows one to apply 848 an accurate background correction, without the need to physically acquire additional background 849 intensity maps with a resulting doubling (at least) of the total acquisition time. The use of 850 spectral interference correction and TDI correction routines also ensure that no such artifacts are 851 present in the quantitative maps. As it follows the rigor and protocols of single point analysis, 852 this proposed mapping quantification method is most likely the best we can do, at present, in 853 terms of accuracy for WDS X-ray map quantification. The acquisition of quantitative X-ray 854 maps affords the analyst valuable quantitative data in less time, when compared to a series of 855 discrete point analysis in compositionally zoned materials or strongly heterogeneous materials, 856 or when the material is extremely beam sensitive.

857

858 Acknowledgements

859 The authors wish to thank all those who over the years who have contributed to the effort to 860 quantify WDS X-ray maps from EPMA instruments. A partial list would include Brian Gavnor, 861 Aoife McFadden, Michael Matthews, Gareth Hatton, Benjamin Wade, Jon Wade, Karen Wright, 862 Ben Buse, Stuart Kearns, Daniel Ruscitto, Malcolm Roberts, David Adams, David Wark, Philipp 863 Poeml, Philippe Pinard, John Fournelle, and Michael Jercinovic. The latter two are also thanked 864 for their most helpful and quite detailed reviews of the manuscript. We thank Alina Fiedrich for 865 providing the Adamello tonalite sample (Figs. 4, 5), Tom McCollom for the jarosite sample 866 (Figs. 10, 11), and Matthew Brueseke and Ben Ellis for the rhyolite sample (Fig. 12). All 867 electron microprobe works presented here were acquired on one of the following instruments: 868 Cameca SX-100 at the CAMCOR center (Oregon University, NSF EAR-0345908 and the

869 Murdoch Foundation), JEOL-8230 at the University of Colorado Boulder (NSF EAR-14)

- JEOL-8530F at the University of Minnesota Minneapolis (NSF EAR-1625422), and JEOL-8230
- at ETH Zürich. The authors are thankful to their respective National Science agencies and
- universities that made possible the purchase of these instruments.
- 873

874 **References**

- 875 Allaz, J.M., Williams, M., Jercinovic, M., Goemann, K., and Donovan, J. (2019a) Multipoint
- 876 Background Analysis: Gaining Precision and Accuracy in Microprobe Trace Element
- 877 Analysis. Microscopy and Microanalysis, 25(1), 30-46.
- Allaz, J.M., Popa, R.G., Reusser, E., and Martin, L. (2019b) Electron Microprobe Analysis of
- 879 Minor and Trace Elements in Beam Sensitive Materials: How Far Can We Go? Microscopy
- and Microanalysis, 25(82), 2312-2313.
- Armstrong, J.T. (1988) Quantitative Analysis of Silicate and Oxide Materials: Comparison of
- 882 Monte Carlo, ZAF, and Procedures. Microbeam Analysis, 239-246.
- Barkman, J., Carpenter, P., Zhao, J., and Donovan, J. (2013). Electron Microprobe Quantitative
 Mapping vs. Defocused Beam Analysis. Microscopy and Microanalysis, 19(S2), 848-849.
- 885 Bastin, G.F., and Heijligers, H.J.M (1991) Quantitative electron probe microanalysis of ultra-
- light elements (boron oxygen). In K.F.J. Heinrich and D.E. Newbury, Eds., Electron Probe
- 887 Quantitation, p. 145-161. Plenum Press, NY.

888	Batanova, V.G., Sobolev, A.V., and Kuzmin, D. V. (2015). Trace element analysis of olivine:
889	High precision analytical method for JEOL JXA-8230 electron probe microanalyser.
890	Chemical Geology, 419, 149-157.
891	Borisova, A.Y., Zagrtdenov, N.R., Toplis, M.J., Donovan, J.J., Llovet, X., Asimow, P.D., de
892	Parseval, P., and Gouy, G. (2018) Secondary fluorescence effects in microbeam analysis
893	and their impacts on geospeedometry and geothermometry. Chemical Geology, 490, 22-29.
894	Carpenter, P.K., Zeigler, R.A., Jolliff, B.L., Vicenzi, E.P., Davis, J.M., and Donovan, J.J. (2009)
895	Advances in electron-probe microanalysis and compositional mapping: Applications to
896	lunar samples. 40th Lunar and Planetary Science Conference, The Woodlands, TX, 2531.
897	Carpenter, P.K., and Hahn, T. (2017) New Developments in Compositional Stage Mapping by
898	EPMA and Micro-XRF. Microscopy and Microanalysis, 23(S1), 1068-1069.
899	Chantler, C.T. (2000) Detailed Tabulation of Atomic Form Factors, Photoelectric Absorption
900	and Scattering Cross Section, and Mass Attenuation Coefficients in the Vicinity of
901	Absorption Edges in the Soft X-Ray (Z=30-36, Z=60-89, E=0.1 keV-10 keV), Addressing
902	Convergence Issues of Earlier Work. Journal of Physical and Chemical Reference Data, 29,
903	597-1048.
904	Chantler, C.T., Olsen, K., Dragoset, R.A., Chang, J., Kishore, A.R., Kotochigova, S.A., and
905	Zucker, D.S. (2005) X-Ray Form Factor, Attenuation and Scattering Tables (version 2.1).
906	[Online] Available: http://physics.nist.gov/ffast [2020, July 29]. National Institute of

Standards and Technology, Gaithersburg, MD. 907

908	Crisci, F., Durand-Charre, A., and Jouanne, D. (2006) Practical Aspects of Carbon Content
909	Determination in Carburized Steels by EPMA. Microscopy and Microanalysis, 12(4), 331-
910	334.
911	Donovan, J.J., Snyder, D.A., and Rivers, M.L. (1993) An improved interference correction for
912	trace element analysis. Microbeam Analysis, 2, 23-28.
913	Donovan, J.J., and Tingle, T.N. (1996) An Improved Mean Atomic Number Correction for
914	Quantitative Microanalysis. Journal of Microscopy and Microanalysis, 2, 1-7.
915	Donovan, J.J., Lowers, H.A., and Rusk, B.G. (2011) Improved electron probe microanalysis of
916	trace elements in quartz. American Mineralogist, 96(2-3), 274-282.
917	Donovan, J.J., Singer, J.W., and Armstrong, J.T. (2016) A new EPMA method for fast trace
918	element analysis in simple matrices. American Mineralogist, 101(8), 1839-1853.
919	Donovan, J., Pinard, P., and Demers, H. (2019) High Speed Matrix Corrections for Quantitative
920	X-ray Microanalysis Based on Monte Carlo Simulated K-Ratio Intensities. Microscopy and
921	Microanalysis, 25(3), 735-742.
922	Eichen, E., Tabock, J., and Kinsman, K.R. (1972) An electron-microprobe technique for
923	detecting low carbon concentrations in iron. Metallography, 5(2), 151-162.
924	Fialin, M., and Chopin, C. (2006) Electron-beam (5-10 keV) damage in triplite-group
925	phosphates: Consequences for electron-microprobe analysis of fluorine. American
926	Mineralogist, 91(4), 503-510.

927	Fiedrich, A.M., Bachmann, O., Ulmer, P., Deering, C.D., Kunze, K., and Leuthold, J. (2017)
928	Mineralogical, geochemical, and textural indicators of crystal accumulation in the Adamello
929	Batholith (Northern Italy). American Mineralogist, 102(12), 2467-2483.
930	Fournelle, J. H., Kim, S., & Perepezko, J. H. (2005). Monte Carlo simulation of Nb Ka
931	secondary fluorescence in EPMA: comparison of PENELOPE simulations with
932	experimental results. Surface and Interface Analysis: An International Journal devoted to
933	the development and application of techniques for the analysis of surfaces, interfaces and
934	thin films, 37(11), 1012-1016.
935	Gedeon, O., Zemek, J., and Jurek, K. (2008) Changes in alkali-silicate glasses induced with
936	electron irradiation. Journal of Non-Crystalline Solids, 354(12-13), 1169-1171.
937	Goldstein, J.I., Newbury, D.E., Echlin, P., Joy, D.C., Firori, C., and Lifshin E. (1992) Scanning
938	Electron Microscopy and X-ray Microanalysis, 550 p., Springer-Verlag, New York (USA).
939	Goloff, B., Webster, J.D., and Harlov, D.E. (2012). Characterization of fluor-chlorapatites by
940	electron probe microanalysis with a focus on time-dependent intensity variation of
941	halogens. American Mineralogist, 97, 1103-1115.
942	Hughes, E.C., Buse, B., Kearns, S.L., Blundy, J.D., Kilgour, G., & Mader, H.M. (2019). Low
943	analytical totals in EPMA of hydrous silicate glass due to sub-surface charging: Obtaining
944	accurate volatiles by difference. Chemical Geology, 505, 48-56.
945	Jercinovic, M.J., Williams, M.L., Allaz, J., and Donovan, J.J. (2012). Trace analysis in EPMA.
946	IOP Conference Series: Materials Science and Engineering, 32, 1-22.

- Kramers, H.A. (1923) On the theory of X-ray absorption and of the continuous X-ray spectrum.
 Phil Mag, 46, p. 836
- Lanari, P., Vidal, O., De Andrade, V., Dubacq, B., Lewin, E., Grosch, E.G., and Schwartz, S.
- 950 (2014) XMapTools: A MATLAB©-based program for electron microprobe X-ray image
- processing and geothermobarometry. Computers and Geosciences, 62, 227-240.
- 952 Merlet, Claude, and Jean-Louis Bodinier. "Electron microprobe determination of minor and trace
- transition elements in silicate minerals: A method and its application to mineral zoning in
- the peridotite nodule PHN 1611." *Chemical Geology* 83.1-2 (1990): 55-69.
- Morgan, G.B. and London, D., 1996. Optimizing the electron microprobe analysis of hydrous
 alkali aluminosilicate glasses. *American Mineralogist*, *81*(9-10), pp.1176-1185.
- 957 Moy, Aurélien, John H. Fournelle, and Anette Von Der Handt. "Solving the iron quantification
- 958 problem in low-kV EPMA: An essential step toward improved analytical spatial resolution
- 959 in electron probe microanalysis—Olivines." American Mineralogist: Journal of Earth and
- 960 *Planetary Materials* 104.8 (2019): 1131-1142.
- 961 Newbury, D.E. (1995) Artifacts in Energy Dispersive X-Ray Spectrometry in Electron Beam
- 962 Instruments. Are Things Getting Any Better?. In D.B. Williams, J.I. Goldstein, D.E.
- 963 Newbury, Eds., X-Ray Spectrometry in Electron Beam Instruments, p. 167-201. Springer,
- Boston, MA.

965	Newbury, D., and Ritchie, N. (2019) Electron-Excited X-ray Microanalysis by Energy
966	Dispersive Spectrometry at 50: Analytical Accuracy, Precision, Trace Sensitivity, and
967	Quantitative Compositional Mapping. Microscopy and Microanalysis, 25(5), 1075-1105.
968	Nielsen, C.H., and Sigurdsson, H. (1981) Quantitative methods for electron microprobe analysis
969	of sodium in natural and synthetic glasses. American Mineralogist, 66, 547-552.
970	Ortolano, G., Visalli, R., Godard, G., and Cirrincione, R. (2018) Quantitative X-ray Map
971	Analyser (Q-XRMA): A new GIS-based statistical approach to Mineral Image Analysis.
972	Computers and Geosciences, 115, 56-65.
973	Potter-McIntyre, S.L., and McCollom, T.M. (2018) Jarosite and alunite in ancient terrestrial
974	sedimentary rocks: Reinterpreting Martian depositional and diagenetic environmental
975	conditions. Life, 8, 32.
976	Tinkham, D.K., and Ghent, E.D. (2005) XRMapAnal: A program for analysis of quantitative X-
977	ray maps. American Mineralogist 90, 737 - 744.
978	Warren, P.H. (1997) The Unequal Host-Phase Density Effect in Electron Probe Defocused Beam
979	Analysis: an Easily Correctable Problem. 28th Annual Lunar and Planetary Science
980	Conference, Houston, TX, 1497-1498.
981	Zhang, X., Yang, S., Zhao, H., Jiang, S., Zhang, R., and Xie, J. (2019) Effect of Beam Current
982	and Diameter on Electron Probe Microanalysis of Carbonate Minerals. Journal of Earth
983	Science, 30(4), 834-842.

985 **Table Captions**

Table 1. Differences in Si concentrations from standard reference published values, where the concentrations are derived from an exponential fit calibration-curve method to the raw X-ray intensities, for synthetic zircon ($ZrSiO_4$) and synthetic forsterite (Mg₂SiO₄).

989

990**Table 2.** Comparison of the average of all map pixels with published reference values for NBS991K-411 mineral composition glass from NIST. The standards were MgO synthetic for Mg K α ,992SiO₂ synthetic for Si K α , and NBS K-412 mineral composition glass for Fe K α , Al K α , Ca K α .993n.a. = not analyzed.

994

Table 3. Evaluation of accuracy and precision for NIST standard K-411 by averaging various
numbers of pixels from the quantitative X-ray maps compared to the average of 4 point analyses.
Values for pixel variance (standard deviation) are consistent, while the standard error (around the
average) decreases, with increasing pixel count as expected. n.a. = not analyzed.

999

1001 Figure Captions

1002	Figure 1. Example of an exponential calibration curve of Si concentrations as a function of X-
1003	ray intensities, for a number of standards each containing Si. Each green dot represents the
1004	average intensity for each standard. Errors on each standard measurement are smaller than the
1005	actual dot size (~0.5% relative for Si K α intensity).

1006

Figure 2. Flow chart of the proposed method for rigorously quantifying a series of raw intensity X-ray maps. These steps include normalization for pixel integration time, detector dead time, beam drift and standard drift (and optional off-peak map background corrections, TDI corrections and duplicate element aggregations), followed by iterative corrections for (MAN) background, matrix and spectral interferences on each pixel individually.

1012

1013 Figure 3. Quantified elemental X-ray maps for Si K α , Fe K α , Ca K α , Al K α and Mg K α for the 1014 NIST K-411 silicate glass reference material, with a total weight-% map and an oxygen map 1015 calculated by stoichiometry with an additional 1.12% O to account for the presence of ferric iron, 1016 shown along with quantitative histograms and measured average and 3 standard deviation bar 1017 and certified values for each element (in red). Conditions were 15 keV, 30 nA, and 500 ms per 1018 pixel. Field of view is 63.5 x 63.5 µm. Note the analytical total map which is important just as it 1019 is for point analysis. These simple maps demonstrate that mapped pixels can achieve accuracy 1020 similar to normal point analyses, even though single pixel precision can be problematic without 1021 pixel averaging as seen in the map color scale ranges. See table 3 for a numerical summary.

1023	Figure 4. (a) Quantitative maps in oxide concentrations (wt%) for a feldspar grain from the
1024	Adamello tonalite; K α lines were analyzed for all elements. Oxygen calculated by stoichiometry
1025	(not shown here). Conditions were 15 keV, 50 nA, and 50 ms per pixel. Field of view is 1.11 x
1026	1.644 mm at 3 μ m/pixel resolution. (b) Formula basis maps for a feldspar grain based on 8
1027	oxygen atoms (apfu = atoms per formula unit). (c) Recalculated mineral end-member maps
1028	albite (Na), anorthite (Ca), and K-feldspar (K). (d) Pixel filtering based on compositional limits
1029	and analytical totals. (e) Quantitative histograms of the previous pixel filtered quantitative maps.
1030	

Figure 5. Maps of the (a) relative analytical sensitivity (1σ) , and (b) detection limit at 99% confidence (3σ) for Si, Al, Ca, Na, and K calculated for each individual pixel for a feldspar grain from the Adamello tonalite.

1034

Figure 6. Element maps of a Fe-Ti mineral from the mafic-ultramafic Viravira Complex, Colombia. Conditions were 8 keV, 35 nA, and 100 ms per pixel. Oxygen K α measured on LDE1 monochromator, and PROZA91 matrix correction was applied with MAC values from FFAST database. Field of view is 414 x 274 µm. (**a**,**b**) The raw intensity X-ray map of O K α (**a**) in the Fe-Ti-oxide mixture shows significant differences when compared to the fully quantified X-ray map (**b**). (**c**,**d**) While the map of the calculated background intensities (**d**) correlates with Fe-rich core areas, the net intensities (**c**) are not affected due to the high peak-to-background ratios.

1042

Figure 7. Quantitative X-ray maps of trace elements in a zircon grain using off-peak maps for
the interpolated off-peak background map correction (a), and alternatively using MAN

background corrections (with the same on-peak map intensities) in (b). Conditions were 15 keV, 1046 100 nA, and 1800 milliseconds per pixel (on-peak and total off-peak). Field of view is 38.1 x 1047 38.1 μ m. Three separate map scans were acquired: the on-peak intensity maps, followed by the 1048 high off-peak intensity maps, followed by the low off-peak intensity maps. A comparison of the 1049 off-peak and MAN maps are seen in the cross sections for each element (c). Note the 1050 improvement in precision particularly for U and Th cross sections with the MAN corrected maps 1051 (red lines) compared to the off-peak corrected X-ray maps (blue lines).

1052

1053 Figure 8. (a-c) Quantitative element maps of Fe, S and Cu. Conditions were 15 keV, 100 nA, 1054 and 1000 ms per pixel. Field of view is 127 x 127 μ m. (d) Raw X-ray map of Mo L α expressed 1055 as counts per second, (e) quantified X-ray map without a spectral interference correction, and (f) 1056 with interference correction. (g,h) Cross section A-B across the pyrite-chalcopyrite interface 1057 with and without a spectral interference correction. (g) Major element values are within the range 1058 of expected values for pyrite and chalcopyrite. Totals are very close to 100% in pyrite, and 1059 slightly lower in chalcopyrite most likely due to the presence of minor elements not analyzed. (h) 1060 Comparison of the Mo content with (Mo) or without (Mo*) the peak interference correction. 1061 Numerical data on the right side are average of concentration measurements along the profile 1062 (see "Quant 1" and "Quant 2" aside the profile). Error (1σ) indicated in parentheses.

1063

Figure 9. Comparison of Cu element wt% map shown with (a) a linear or (b) a logarithmic scale. Using the logarithmic scale, the spurious low concentrations of Cu in pyrite due to secondary fluorescence across the phase boundary can be made visible, while avoiding

1067 oversaturation of pixels with higher Cu content in chalcopyrite. (c) Plot of Cu wt% in pyrite
1068 versus the distance to the boundary with chalcopyrite.

1069

1070 Figure 10. Quantified X-ray maps of jarosite in sample MN05 using a beam map at 15 keV with 1071 a focused beam. Field of view is 50 x 50 μ m at ~0.39 μ m/pixel resolution. Three separate 1072 examples are shown with increasing electron beam dose: (a) 50 nA, 5 ms dwell time, 8 frames (= 1073 2 nA·s), (b) 100 nA, 5 ms dwell time, 8 frames (= 4 nA·s), and (c) 100 nA, 40 ms dwell time, 5 1074 frames (= $20 \text{ nA} \cdot \text{s}$). For each test, the data was processed using either the TDI correction on each 1075 pixel, or by aggregating all intensities received on each frame (i.e., no correction for beam 1076 damage). All pixels yielding a low total below 80% were set to black in Figure 10. The table 1077 below lists the average of 100 pixels (area indicated on the SO_3 map) for the TDI corrected and 1078 aggregated data, respectively, along with the quantitative result considering only the first or the 1079 last acquired beam map frame. See text for discussion.

1080

1081 **Figure 11.** TDI correction (in relative percent) map for the analysis of K K α , S K α , and Fe K α 1082 at 100 nA, 40 ms per pixel, and with 5 passes (total beam dose is 20 nA·s).

1083

Figure 12. Quantitative element mapping in rhyolite, with focus on Ba-content in plagioclase, Kfeldspar, and glass. Conditions were 15 keV, 100 nA, and 40 ms per pixel, 2 passes. Field of view is 329.5 x 229.5 μ m at 0.5 μ m/pixel resolution. (**a-f**) Quantitative map of SiO₂, Al₂O₃, CaO, Na₂O, K₂O, and total wt%. (**g**) Quantitative BaO wt% map obtained by aggregating the intensity of three spectrometers equipped with L-type (large) PET monochromators. (**h**)

Quantitative cross-section extracted from the map, as shown in (g). Horizontal black lines over the BaO cross-section are precise data obtained from single-point quantitative analysis (longer counting time). (i) Quantitative single-point analysis shown for comparison; position of each analysis indicated in (g).

1093

Table 1

Si Ka, TAP, 20 keV, 30 nA	ZrSiO ₄	Mg ₂ SiO ₄
Published Concentration of Standard	15.32 wt%	19.96 wt%
Concentration from Calibration Curve (Fig. 1)	19.09 wt%	17.38 wt%
Relative error %	24.6%	-12.9%

Table 2	2
---------	---

	Si wt%	Fe wt%	Mg wt%	Al wt%	Ca wt%	O wt%	Calc. O	Total
Average	25.06	11.18	8.63	0.02	10.98	1.12	42.93	98.8 0
Published	25.38	11.21	8.85	n.a.	11.06	1.12	43.56	99.98
Std Dev	0.72	0.64	0.15	0.02	0.20	0.00	0.86	1.84
Rel %	2.89	5.74	1.75	99.23	1.85	0.00	2.01	1.86
Minimum	22.28	8.43	8.06	-0.05	10.16	0.00	39.56	91. 77
Maximum	27.84	14.17	9.25	0.13	11.82	0.00	46.38	106.08

Table 3

	Si wt%	Al wt%	Fe wt%	Mg wt%	Ca wt%	Integration time
NIST K-411 (Published)	25.38	n.a.	11.21	8.85	11.06	
NIST K-411 (Measured)	25.277	0.019	11.098	8.811	11.037	80 sec
4 Points Variance	0.037	0.004	0.082	0.027	0.025	
4 Points Error	0.019	0.002	0.041	0.014	0.013	
4 pixels Average	25.261	0.021	10.893	8.741	11.289	0.4 sec
4 pixels Variance	0.582	0.050	0.409	0.528	0.510	
4 pixels Error	0.291	0.025	0.205	0.264	0.255	
16 pixels Average	25.432	0.000	10.996	8.947	11.112	1.6 sec
16 pixels Variance	0.507	0.038	1.293	0.584	0.500	
16 pixels Error	0.127	0.010	0.323	0.137	0.125	
32 pixels Average	25.399	0.011	11.131	8.747	11.049	3.6 sec
32 pixels Variance	0.454	0.046	1.444	0.589	0.406	
32 pixels Error	0.076	0.008	0.241	0.098	0.068	
64 pixels Average	25.463	0.016	11.310	8.809	11.046	6.4 sec
64 pixels Variance	0.529	0.044	1.563	0.670	0.414	
64 pixels Error	0.066	0.001	0.195	0.084	0.052	
All pixels Average	25.353	0.016	11.057	8.748	10.994	1638 sec
All pixels Variance	0.494	0.048	1.427	0.593	0.460	
All pixels Error	0.004	0.001	0.011	0.005	0.004	





Figure 3





Figure 4 (continued)

d Oxide wt-%, after pixel filtering





Na₂O wt%





 K_2^{0} wt%

15

10



CaO wt%

40

-35

30

25

20

L₁₅

0.5

0.4

0.3

0.2

0.1

n



-25

F-20

<u>–</u>15

F10

-5

-0

F













0

-3



Quantitative maps (element wt%)



Raw image Mo K α in cps

d

Mo cps

Ó

200

400

6Ó0



No interference correction

20

40

60

ż



Mo wt%



0 10 20 30 ·

With interference corr.



Mo wt%

-0.2 0.0 0.2 0.4

Quant 1 Quant 2 Quant 1 Quant 2 60 2.5 h (g) Mo wt-%: No peak Quant 1 & 2 S 108 interference correction 2.0 50 1) FeS s 53.2(4) 104 Fe Fe 45.6(5) 1.5 Cu 0.07(8) **No kt-%** Total wt-% Mo 0.00(8) Total Mo* 1.84(8) 100 Mo wt-%: with peak 2) CuFeS, S 34.2(2) 0.5 interference correction 96 Fe 29.3(3) (tail of S K α peak on Mo L α peak) Cu Cu 33.1(3) ••• 10 Fig. 9 Mo 0.03(6) 0.0 92 Mo* 1.21(7) -0.5 0 100 120 **B** 100 120 **B** 20 40 60 80 20 40 60 80 Α Α Distance (µm) Distance (µm)





Any pixel with total below 80 wt% (e.g., other minerals, epoxy) is masked in black.

d) 100-pixel averages (10x10 pixels; see location on SO₃ map)

	50 nA, 5 ms, 8 frames			100 nA, 5 ms, 8 frames				100 nA, 40 ms, 5 frames				
	a1 TDI	a2 Agg.	FIRST	LAST	b1 TDI	b2 Agg.	FIRST	LAST	c1 TDI	c2 Agg.	FIRST	LAST
K,O	8.33(12)	8.24(8)	8.32	8.16	8.88 <i>(9)</i>	7.16 <i>(4)</i>	8.94	6.18	8.39 <i>(6)</i>	7.12(3)	8.44	6.49
SÔ	33.0(5)	32.6(2)	33.2	32.0	32.5(3)	33.2(1)	32.4	34.0	32.91(14)	34.16(10)	32.8	34.9
Fe,O,	46.0(4)	45.5(2)	45.5	45.6	45.4(3)	47.8(2)	45.1	49.7	44.00(14)	48.08(8)	44.7	51.4
Al ₂ O ₃	0.85(7)	0.79(5)	0.84	0.75	0.89(6)	0.87(4)	0.83	0.88	1.81(6)	1.92(5)	1.85	2.00
As ₂ O ₅	0.18(6)	0.06(1)	0.08	0.05	0.05(1)	0.06(1)	0.06	0.08	0.12(1)	0.13(1)	0.10	0.13
H ₂ O*	11.48	11.46	11.48	11.55	11.07	11.03	11.04	10.98	11.57	11.39	11.55	11.25
Total	99.30	98.10	98.88	97.61	98.70	100.13	98.23	101.91	98.28	102.43	98.88	105.86

* Calculated by stoichiometry: 6 H for 14 O.



Mostly positive TDI correction

Mostly negative TDI correction



Total	100.06	100.91	101.72	100.59	97.26
K ₂ O	7.98	8.71	4.12	1.72	6.50
Na₂O	4.80	4.31	6.22	6.67	2.17
BaO	1.28	2.75	0.83	0.32	0.12
SrO	0.09	0.08	0.11	0.09	0.01
CaO	1.42	0.89	3.71	6.12	0.78
MgO	0.01	0.00	0.01	0.01	0.05
MnO	0.00	-0.01	0.00	0.00	0.05
FeO	0.18	0.19	0.24	0.35	2.47
Al_2O_3	19.82	20.26	22.81	24.55	11.66
TiO ₂	0.04	0.04	0.03	0.02	0.29
2.2	•				

