REVISION #1 1 The suppression of lone-pair-stereoactivity in $[Cu^{+}(As^{3+}O_{3})_{4}]$ clusters in 2 dixenite: a tribute to Paul B. Moore 3 4 FRANK C. HAWTHORNE^{1, \xi} AND JOHN M. HUGHES² 5 6 7 ¹Department of Geological Sciences, University of Manitoba, Winnipeg, Manitoba, R3T 2N2, Canada 8 ²Department of Geology, University of Vermont, Burlington, VT 05405, U.S.A. 9 10 11 ABSTRACT The crystal structure of dixenite, ideally Cu⁺Fe³⁺Mn²⁺₁₄(As⁵⁺O₄)(As³⁺O₃)₅(SiO₄)₂(OH)₆, 12 from Langban, Sweden, was refined to an R_1 -index of 1.58%, and the structure proposed by 13 Araki and Moore (1981) was confirmed and details elucidated. The structure, crystallizing in 14 space group R3 with a = 8.2204(3) and c = 37.485(3)Å, consists of layers of (Mn²⁺, 15 Fe^{3+})(O,OH)₆ octahedra linked by (As⁵⁺O₄) and (SiO₄) tetrahedra, (As³⁺O₃) trigonal pyramids 16 and $(Cu^{+}As^{3+})_{4}$ tetrahedra. There are five distinct layers in the repeat unit of the cell, four of 17 18 which are very similar to the layers in mcgovernite. An unusual aspect of one of the trimers of octahedra is that there is a triangular-prismatic hole through the center of the cluster. The 19 $(Cu^{+}As^{3+}_{4})$ tetrahedra are parts of larger clusters: $[Cu^{+}(As^{3+}O_{3})_{4}]$ in which four $(As^{3+}O_{3})$ groups 20 link to a central Cu⁺ which occupies the positions normally taken by the stereoactive lone-pairs 21 of electrons that generally characterize As³⁺ in triangular-pyramidal coordination by O. Thus the 22 stereoactive lone-pair behavior that is characteristic of $(As^{3+}O_3)$ trigonal pyramids is suppressed 23 by the coordination of Cu^+ by four As^{3+} ions. 24 25 Keywords: Dixenite, crystal structure, lone-pair electrons, Paul B. Moore

^{*ξ*} E-mail: Frank.Hawthorne@umanitoba.ca

27

INTRODUCTION

Paul Brian Moore (1940-2019) was one of the giants of Mineralogy, and his legacy lives 28 in the numerous papers he published in his distinguished career. His contributions to Mineralogy 29 were celebrated in a recent memorial (Hawthorne et al. 2019) that stated that he was the greatest 30 mineralogist of the 20th century. Professor Moore and his colleagues solved numerous complex 31 structures in an age before (usually) routine solution of crystal structures using Direct Methods 32 was possible, stretching the limits of equipment and computational resources of that time. As 33 noted in his Memorial, even in retirement Paul never stopped thinking about problems involving 34 35 crystal structures. In addition to being involved with new atomic arrangements, he urged colleagues to re-examine some of his earlier solutions that were done with the diffraction 36 equipment of an earlier day. One of those structures was that of dixenite, which was a frequent 37 source of conversations with his colleagues. At the time of his death, Professor Moore was 38 urging both of us to re-examine the dixenite structure, and we are pleased to offer the results of 39 that work here. 40

Dixenite, ideally $Cu^+Fe^{3+}Mn^{2+}{}_{14}(As^{5+}O_4)(As^{3+}O_3)_5(SiO_4)_2(OH)_6$, is of considerable 41 interest because it has a novel metallic cluster, ideally $(Cu^+As^{3+}_4)$, embedded in an oxide 42 environment (Araki and Moore 1981), an atomic arrangement that had not been observed 43 previously in any structure. At the same time, Moore and Araki (1979) found another metallic 44 cluster, $(Mn^+As^{3+}_{6})$, embedded in an oxide environment in magnussonite, ideally 45 $Mn^{2+}_{18}[Mn^+As^{3+}_{6}O_{18}]_{2}Cl_{2}$. As far as we are aware, these are the only examples of such clusters 46 found in minerals, and Paul B. Moore (personal communication) was keen to see confirmation of 47 48 these findings by more accurate modern instrumentation. Moreover, both these minerals occur 49 only in the (Fe-Mn)-oxide ore deposits at Långban, Sweden (Moore 1970; Holtstam and Langhof

50	1999; Lundström 1999; Bollmark 1999; Nysten et al. 1999), and we have a long-term interest in
51	the basic Mn-arsenate-silicate minerals from this locality (Hawthorne 2018; Hawthorne et al.
52	2013; Cooper and Hawthorne 1999, 2012).
53	
54	SAMPLE PROVENANCE
55	Both samples examined here are from Långban. One sample was obtained from the U.S.
56	National Museum of Natural History (C4440, provided to Paul B. Moore) and the other from the
57	late Mark Feingloss, a well-known mineral collector from Duke University. The structural results
58	are virtually identical and we present only one refinement here.
59	
60	CRYSTAL STRUCTURE
61	Data collection and refinement
62	A crystal fragment from NMNH C4440 was mounted on a Bruker Apex CCD
63	diffractometer equipped with graphite-monochromated Mo $K\alpha$ radiation. Refined cell
64	parameters and other crystal data are listed in the deposited CIF file. Redundant data were
65	collected for a sphere of reciprocal space, and were integrated and corrected for Lorentz and
66	polarization factors and absorption using the Bruker program SAINTPLUS.
67	The atomic arrangement was solved independently of that given by Araki and Moore
68	(1981), and hydrogen positions were located using difference-Fourier maps. The structure is in
69	close agreement with that of Araki and Moore (1981); we have retained their original atom
70	nomenclature in this paper. Refinement was done with anisotropic-displacement parameters for
71	all atoms except H. Table 1 lists the refined atom parameters, Table 2 lists selected interatomic
72	distances, and Table 3 gives the bond-valence table (vu; valence units) calculated with the

parameters of Gagné and Hawthorne (2015). A CIF has been deposited which contains further
 details of crystal data and structure refinement¹.

75 Site populations

The atomic arrangement is virtually identical with that given by Araki and Moore (1981), 76 although of much greater precision. There are three As sites with associated coordinations, bond 77 lengths and angles characteristic of As³⁺. Each As³⁺ is coordinated by three O²⁻ ions at distances 78 from 1.726-1.776 Å, in the range of ^[3]As³⁺-O distances shown by inorganic crystal-structures: 79 1.671-1.845 Å (Gagné and Hawthorne 2018) and close to the grand mean ^[3]As³⁺-O distance of 80 1.776 Å. There are three T sites that have site-scattering values between Si and As; each site is 81 coordinated by a tetrahedral arrangement of O atoms, and have T-O distances (Table 2) that are 82 intermediate between those expected for ^[4]Si-O and ^[4]As⁵⁺-O in oxide and oxysalt structures 83 (Gagné and Hawthorne 2018). There are six M sites that have site-scattering values characteristic 84 of Mn and Fe and that are octahedrally coordinated by O atoms. The observed interatomic 85 distances (Table 2) indicate that six of these sites are occupied by Mn^{2+} with minor Fe²⁺. The site 86 scattering at the M3 site indicates that this site is occupied by a transition metal. The <M3-O> 87 bond length is 2.057 Å, significantly longer than the grand $\langle Fe^{3+}-O \rangle$ distance of 2.015 Å 88 89 reported by Gagné and Hawthorne (2020) for oxide and oxysalt structures. Thus, in addition to Fe^{3+} , M3 must be occupied partly by Fe^{2+} and/or Mn^{2+} , in accord with the incident bond-valence 90 sum at M3 of 2.67 vu (Table 3). Polyhedra are labelled by the identity of the ion/atom at the 91 central site: thus the coordination octahedron of Mn^{2+} at the M1 site, $Mn^{2+}O_6$, is denoted as the 92

¹ Deposit item AM-20-xx1 for CIF file. Deposit items are available two ways: for paper copies, contact the Business Office of the Mineralogical Society of America (see inside front cover of recent issue) for price information. For an electronic copy, visit the MSA web site at http://www.minsocam.org, go to the *American Mineralogist* Contents, find the table of contents for the specific volume/issue wanted, and then click on the deposit link there.

M1 octahedron. The bond valences incident at the O anions (Table 3) show that OH1 and OH2 are hydroxyl groups and the remaining anions are O^{2^-} . The site-scattering values and mean bond lengths are in accord with the formula $Cu^+Fe^{3+}Mn^{2+}{}_{14}(As^{5+}O_4)(As^{3+}O_3){}_5(SiO_4){}_2(OH)_6$ assigned by Araki and Moore (1981).

97 Bond topology

98 The dixenite structure contains fifteen layers of approximately close-packed polyhedra
99 that comprise one translation along c. In space group *R*3 five symmetrically and bond-

topologically distinct layers are stacked orthogonal to the c-axis to give the ~37.5 Å repeat in the

101 **c** direction. The five distinct layers of cation-centered polyhedra are labelled m = 0-4 in Figure 1, 102 and these layers are shown in plan in Figure 2 where they are compared with similar layers in the

103 crystal structure of mcgovernite (Hawthorne 2018).

At m = 0 (Fig. 2*a*), the layer consists of M5 octahedra containing Mn²⁺ and two distinct 104 tetrahedra, T2As and T3As, that contain As⁵⁺ and Si with As⁵⁺ dominant at both T sites (Table 105 2). The M5 octahedron shows a large dispersion of bond lengths: 2.086-2.735 Å (Table 2) and 106 the question arises as to the coordination number of Mn^{2+} at the M5 site. Gagné and Hawthorne 107 (2020) list the range of observed ${}^{[6]}Mn^{2+}$ -O distances in oxide and oxysalt structures as 1.968-108 2.798 Å and the range of $<^{[6]}$ Mn²⁺-O> distances as 2.134-2.305 Å; the observed M5-O distances 109 fall within these ranges (Table 2) and hence we consider the M5 polyhedron as an octahedron. 110 The M5 octahedra share edges to form a very unusual trimer (Fig. 3a). Edge-sharing between 111 octahedra is extremely common in oxide and oxysalt structures, and the usual arrangement of a 112 trimer of edge-sharing octahedra is shown in Fig. 3b for the M6 octahedra in layer m = 4, in 113 which the shared edges (shown in red in Fig. 3) meet at the centre of the trimer. In the M5 trimer 114 (Fig. 3a), the shared edges form the vertical (sub-parallel to c) edges of a twisted triangular prism 115

116	that forms an empty channel through the center of the M5 trimer. The reason for this
117	arrangement is not apparent, but may be related to the different stoichiometries of the two
118	arrangements: M5: [M ₃ O ₁₂]; M6: [M ₃ O ₁₃]. The M5 trimers link into a sheet by sharing corners
119	with tetrahedra; the T2As tetrahedron points in the $+c$ direction and the T3As tetrahedron points
120	in the $-c$ direction (Fig. 2a). Fig. 2b shows the analogous layer in the structure of mcgovernite in
121	which the unit cell is shifted relative to that in dixenite by $1/3 2/3$ in the (001) plane. In
122	mcgovernite, there is only one crystallographically distinct tetrahedron in this layer and this
123	tetrahedron is occupied by Si. The Z1 tetrahedron in mcgovernite occupies a position similar to
124	the M5 octahedron in dixenite but does not link to other Z1 tetrahedra by sharing vertices;
125	however, the pattern of Z1 tetrahedra in the $m = 7$ layer in mcgovernite (Fig. 2b) resembles the
126	pattern of M5 polyhedra in the $m = 0$ layer of dixenite (Fig. 2a).
127	At $m = 1$ (Fig. 2c), the layer consists of trimers of edge-sharing M4 octahedra linked by
128	sharing corners with $As^{3+}O_3$ groups. In the topologically analogous layer in mcgovernite (m = 6;
129	Fig. 2d), the trimers of M5 octahedra are linked by SiO ₄ groups. In dixenite, the corners of the
130	unit cell are situated at holes in the layer (Fig. 2c). This hole has a very low occupancy by Cu
131	(M2A site; Table 2) and we have denoted the corresponding layer as $m = 1$ ' as it is disordered
132	with the $m = 1$ layer. In the $m = 1$ ' layer (Fig. 2e) in dixenite, there is an M2A octahedron at the
133	origin of the unit cell that shares edges with the octahedra of the M4 trimers to form an
134	interrupted sheet of octahedra with As ³⁺ O ₃ groups occupying the interstices. A similar layer
135	occurs in mcgovernite (Fig. 2f) except that the $As^{3+}O_3$ groups in dixenite are replaced with SiO_4
136	groups in mcgovernite.

The m = 2 layer in dixenite (Fig. 2g) consists of isolated $Cu^+As^{3+}_4$ groups that link to isolated M1 and M2 octahedra through two distinct $As^{3+}O_3$ groups: As2 and As3. There is no analogous layer in the structure of mcgovernite.

- In the m = 3 layer in dixenite (Fig. 2h), trimers of M7 octahedra share edges with a single M3 octahedron to form an interrupted sheet in which $As^{3+}O_3$ groups link to the surrounding M7 octahedra. The analogous m = 3 layer in mcgovernite (Fig. 2i) has the same connectivity except that there are $As^{5+}O_4$ tetrahedra in the interstices of the sheet and the M1 octahedron, although dominated by Mn^{2+} , is occupied by a considerable amount of Mg.
- 145 The m = 4 layer in dixenite (Fig. 2j) consists of edge-sharing trimers of M6 octahedra
- linked by sharing corners with SiO_4 tetrahedra; note that this layer differs from the m = 1 layer in
- 147 dixenite in that the trimers have different orientations (Figs. 2c, 2j). The analogous layer in

148 mcgovernite is m = 2A (Fig. 2k) in which the trimers are linked by $As^{3+}O_3$ groups.

149 The $(Cu^+As^{3+}_4)$ arrangement

There are two Cu sites, Cu1 and Cu2, that are 0.879(8) Å apart, and hence both cannot be locally occupied. The two sites are jointly coordinated by a trigonal bipyramid of As³⁺ ions (Fig. 4). The refined site-scattering at the two Cu sites, in accord with aggregate occupancy of the two Cu sites, is Cu⁺_{0.85} + $\Box_{0.15}$, an at any one Cu1-Cu2 pair, one Cu site is occupied by Cu⁺ and the locally associated Cu site is vacant. The neighboring As sites are fully occupied by As³⁺. Where both sites are locally not occupied, the lone pairs on the locally associated As³⁺ ions will point toward the unoccupied Cu1 and Cu2 sites.

The coordination of each of the Cu sites is illustrated in Figure 4 in which each Cu⁺ ion is [4]-coordinated by As^{3+} Each coordinating As^{3+} ion bonds to three O atoms at distances of ~1.76 Å and one Cu⁺ ion at distances of 2.34-2.50 Å. The oxygen atoms bonded to As^{3+} are arranged in

160	a trigonal prismatic arrangement; the range of As ³⁺ -O distances (1.73-1.78 Å) are well within the
161	range of 1.67-1.85 Å given by Gagné and Hawthorne (2018) for ^[3] As ³⁺ -O distances in all
162	inorganic oxide and oxysalt compounds, and the mean distance of 1.76 Å is close to their grand
163	mean ^[3] As ³⁺ -O distance of 1.776 Å. Thus the Cu ⁺ ions are each coordinated by four As ³⁺ -O ₃
164	groups in a tetrahedral arrangement. Such As ³⁺ -O ₃ groups are characterized by a stereoactive
165	lone-pair of electrons extending away from the As ³⁺ ion on the side opposing the three O atoms.
166	In the atom arrangements shown in Figure 4, it is apparent that the stereoactive character of the
167	lone electron pair is suppressed by the presence of Cu^+ as a fourth ligand to As^{3+} : $As^{3+}O_3Cu^+$.
168	The Cu1-As distances are 2.400 \times 3, 2.342 and 3.378 Å and the Cu2-As distances are 2.369 \times 3,
169	2.499 and 3.221 Å. It is apparent that Cu1 and Cu2 are each coordinated by four As^{3+} ions and
170	the longer distances > 3.2 Å are not bonded interactions. Inspection of Figure 4 shows that the
171	splitting of Cu into two separate sites, Cu1 and Cu2, is driven by the need for Cu ⁺ to shorten the
172	bond to As ³⁺ , either As1 or As2, and the 1:1 split suggests that this is not long-range ordered.
173	Cu ⁺ -As ³⁺ interactions are not common in crystal structures, but such bonds have been
174	reported in a few metallo-organic structures. Karagiannidis et al. (1991a) lists a single Cu ⁺ -As ³⁺
175	bond of 2.371(1) Å in [Cu(tclH) ₂ (AsPh ₃)Br] in a (Cu ⁺ As ³⁺ S ²⁻ ₂ Br ⁻) tetrahedron, and Karagiannidis
176	et al. (1991b) lists Cu^+ -As ³⁺ distances of 2.411 (2) and 2.372(2) Å in Cu(tclH)(AsPh ₃) ₂ Br in a
177	$(Cu^+As^{3+}_2S^2Br^-)$ tetrahedron. The distances are sufficiently similar to those reported for Cu ⁺ -
178	As ³⁺ in Table 2 to suggest that the atomic arrangements shown in Fig. 4 are chemically
179	reasonable.
180	Fitting the $[Cu^+(As^{3+}O_3)_4]$ clusters into the layered structure

181 The $[Cu^+(As^{3+}O_3)_4]$ clusters are quite complicated atom arrangements and it is surprising 182 to encounter them in a close-packed structure. The way in which they are incorporated is

illustrated in Fig. 5. There is a layer of trimers of Mn²⁺ octahedra (Fig. 5a) in which adjacent 183 trimers provide a triangle of O atoms that link to As^{3+} at the apical position of the $(Cu^{+}As^{3+})$ 184 tetrahedron centered at the Cu2 site. Figure 5b shows a view of the same arrangement in the 185 other direction where the linkage of the other $(As^{3+}O_3)$ groups to the underside of the M4 trimers 186 and to the underlying layer of M(7) trimers is apparent. Thus the two layers of M-trimers are 187 linked by the $[Cu^+(As^{3+}O_3)_4]$ cluster involving the Cu⁺ ion at the Cu2 site. As is apparent from 188 Figure 4, the corresponding $[Cu^+(As^{3+}O_3)_4]$ cluster involving the Cu1 site points in the opposite 189 direction. 190

Figure 6 compares the M(4)-As(1) layer in dixenite with the M(2)-As(3) layer in 191 mcgovernite. Both layers consist of trimers of octahedra linked by $(As^{3+}O^{3})$ groups but the 192 patterns of distribution of the $(As^{3+}O_3)$ groups are complementary. The layer in mcgovernite 193 (Fig. 6b) is disordered with the layer shown in Figure 6c; in that layer the $As^{3+}O_3$ groups are 194 replaced by a trimer of Z(2) octahedra in which the octahedra are only partly occupied. It is 195 striking that the other layers in mcgovernite with the same pattern of trimers and linking (SiO_4) 196 groups (layers m = 4 and m = 6; Hawthorne 2018) do not show this disorder. It is not clear at the 197 moment what is causing the presence of the unusual $[Cu^+(As^{3+}O_3)_4]$ cluster in dixenite and the 198 analogous disordered layers of trimers of partly occupied octahedra in mcgovernite. It is to be 199 hoped that a more comprehensive examination of all the basic manganese-silicate-arsenate-200 arsenite minerals (including two potentially new minerals, work in progress) will allow us to 201 understand why Cu⁺-As³⁺ bonds form in some of these structures. 202

203

IMPLICATIONS

The basic manganese-iron arsenate-arsenite-silicate minerals of the Långban-type
 deposits in Bergslagen, Sweden, form a family of very complicated layered structures, several of

206 which contain local exotic atomic arrangements embedded within their close-packed structures. Examples are the $[Cu^+(As^{3+}O_3)_4]$ cluster and the $[Mn^{2+}_{3}O_{12}]$ cluster reported here in dixenite and 207 the local replacement of an $As^{3+}O_3$ group by a [Mg,Mn²⁺₃O₁₃] cluster in mcgovernite and 208 carlfrancisite (Hawthorne 2018). In the same type of deposit, magnussonite (Moore and Araki 209 1979), although a defect-fluorite structure, contains a $[Mn^+As^{3+}_{6}]$ cluster. The presence of these 210 exotic clusters in densely packed Mn^{2+} octahedra is not understood at the moment, but may be 211 related to the relaxation of accumulated strain. There is the potential to incorporate exotic 212 clusters with unusual properties into dense oxide matrices, and hence develop materials with 213 desirable physical properties if the details of their incorporation can be understood. With only 214 two minerals that contain metallic clusters in an oxide matrix, little is known about such clusters 215 and compounds. In both minerals, the semimetal As is an essential component of the cluster, and 216 that class of elements may facilitate their formation in oxide matrices. Metal clusters on two-217 dimensional substrates are studied extensively in the science of non-volatile memory materials, 218 and the existence of such clusters in three-dimensional mineral structures may provide a template 219 for important industrial materials. The detailed re-examination of dixenite described herein, and 220 the discovery and description of increasingly complex mineral structures over the past decade, 221 illustrate that minerals provide a rich template for important synthetic materials. 222

223

ACKNOWLEDGMENTS

This work was supported by a Discovery grant from the Natural Sciences and Engineering Research Council of Canada and a Canada Foundation for Innovation Grant, both to FCH, and by NSF grants EAR-0003201 and MRI 1039436 to JMH. Both authors are also grateful for years of interaction with the late Paul Brian Moore and for his suggestion that we look at dixenite more closely. The authors are grateful for the editorial handling of Oliver

- 229 Tschauner and the detailed reviews of Anthony R. Kampf, Martin Kunz, and the American
- 230 *Mineralogist* Technical Editor.

231			
232			
233			
234			
235			
236			
237			
238			
239			
240			
241			
242			
243			
244			
245			

246

248	References Cited
249	Araki, T., and Moore, P.B. (1981) Dixenite, $Cu^{1+}Mn_{14}^{2+}Fe^{3+}(OH)_6(As^{3+}O_3)5(As^{5+}O_4)$: metallic
250	$[As_4^{3+}Cu^{1+}]$ clusters in an oxide matrix. American Mineralogist, 66, 1263-1273.
251	Bollmark, B. (1999) Some aspects of the origin of the deposit. In Långban: the Mines, their
252	Minerals, Geology and Explorers (D. Holtstam & J. Langhof, eds.). Raster Förlag,
253	Stockholm, Sweden (43-49).
254	Cooper, M.A. and Hawthorne, F.C. (1999) The effect of differences in coordination on ordering
255	of polyvalent cations in close-packed structures: the crystal structure of arakiite and
256	comparison with hematolite. The Canadian Mineralogist, 37, 1471–1482.
257	Cooper, M.A. and Hawthorne, F.C. (2012) The crystal structure of kraisslite,
258	$^{[4]}Zn_3(Mn,Mg)_{25}(Fe^{3+},Al)(As^{3+}O_3)_2[(Si,As^{5+})O_4]_{10}(OH)_{16}$, from the Sterling Hill mine,
259	Ogdensburg, Sussex County, New Jersey, USA. Mineralogical Magazine, 76, 2819-
260	2836.
261	Gagné, O. & Hawthorne, F.C. (2015) Comprehensive derivation of bond-valence parameters for
262	ion pairs involving oxygen. Acta Crystallographica, B71, 562-578.
263	Gagné, O. C. and Hawthorne, F.C. (2018) Bond-length distributions for ions bonded to oxygen:
264	Metalloids and post-transition metals. Acta Crystallographica, B74, 63-78.
265	Gagné, O.C. & Hawthorne, F.C. (2020) Bond-length distributions for ions bonded to oxygen:
266	Results for the transition metals and quantification of the factors underlying polyhedral
267	distortion via bond-length variation. IUCrJ, 7, 581-629.
268	Hawthorne, F.C. (2018): Long-range and short-range cation order in the crystal structures of
269	carlfrancisite and mcgovernite. Mineralogical Magazine, 82, 1101-1118.
270	Hawthorne, F.C., Abdu, Y.A., Ball, N.A. and Pinch, W.W. (2013) Carlfrancisite: Mn ²⁺ ₃
271	$(Mn^{2+},Mg,Fe^{3+},Al)_{42}[As^{3+}O_3]_2(As^{5+}O_4)_4[(Si,As^{5+})O_4]_6[(As^{5+},Si)O_4]_2(OH)_{42}$, a new

- arseno-silicate mineral from the Kombat mine, Otavi valley, Namibia. American
- 273 Mineralogist, 98, 1693–1696.
- Hawthorne, F.C., Kampf, A.R., and Hervig, R. (2019) Memorial of Paul Brian Moore, 1940-
- 275 2019. American Mineralogist, 104, 1062-1063.
- Holtstam, D. and Langhof, J. (editors) (1999) Långban, the Mines, their Minerals, History and
 Explorers. Raster Förlag, Stockholm, pp. 215.
- 278 Karagiannidis, P., Akrivos, P.D., Mentzafos, D. and Terzis, A. (1991a) New class of Cu(I)
- structure of [Cu(tclH)₂(AsPh₃)Br]. Inorganica Chimica Acta, 181, 263-267.
- 281 Karagiannidis, P., Akrivos, P., Aubry, A., and Skoulika, S. (1991b) Synthesis and study of
- 282 mixed ligand monomer Cu(I) compounds with Cu-As bonds. Crystal and molecular
- structure of bis(triphenylarsine)-(2-thioxohexamethyleneimine)copper(I) bromide.
- Inorganica Chimica Acta, 188, 79-83.
- Lundström, I. (1999) General geology of the Bergslagen ore region. In Långban: the Mines, their
- 286 Minerals, Geology and Explorers (D. Holtstam & J. Langhof, eds.). Raster Förlag,
- 287 Stockholm, Sweden (19-27).
- Moore, P.B. (1970) Mineralogy and chemistry of Långban type deposits in Bergslagen, Sweden.
 Mineralogical Record 1, 154-172.
- Moore, P. B. and Araki, T. (1979) Magnussonite, manganese arsenite, a fluorite derivative
 structure. American Mineralogist, 64, 390-401.
- 292 Nysten, P., Holtstam, D. & Jonsson, E. (1999) The Långban minerals. *In* Långban: the Mines,
- their Minerals, Geology and Explorers (D. Holtstam & J. Langhof, eds.). Raster Förlag,
 Stockholm, Sweden (89-183).

- **Table 1**. Atom coordinates and equivalent isotropic atom-displacement parameters (Å²)
- for dixenite. U(eq) is defined as one third of the trace of the orthogonalized U_{ij} tensor.

298

Atom	x/a	y/b	z/c	U(eq)	Occ.
As1	1/3	2/3	0.93013(3)	0.00961(13)	As _{1.00}
As2	2/3	1/3	0.74936(3)	0.00925(13)	As _{1.00}
As3	0.91124(5)	0.62586(6)	0.68425(3)	0.01006(9)	As _{1.00}
T1SI	1/3	2/3	0.81217(5)	0.0065(5)	Si _{0.946(7)} As _{0.054}
T2AS	0	0	0.85378(4)	0.0079(3)	Si _{0.458} As _{0.542(7)}
T3AS	2/3	1/3	0.88640(3)	0.0073(3)	Si _{0.275} As _{0.725(7)}
M1	0	0	0	0.01162(19)	Mn _{1.00}
M2	2/3	1/3	0.99435(5)	0.0177(4)	Mn _{0.937(5)}
M2A	2/3	1/3	0.9635(7)	0.030(6)	Cu _{0.063}
M3	0	0	0.74267(4)	0.0091(3)	Fe _{0.943(6)}
M4	0.95815(9)	0.73646(9)	0.93256(3)	0.01243(13)	Mn _{1.00}
M5	0.58214(9)	0.66184(10)	0.87042(3)	0.01397(13)	Mn _{1.00}
M6	0.89166(9)	0.60100(9)	0.80761(3)	0.01194(13)	Mn _{1.00}
M7	0.57710(9)	0.68485(9)	0.73874(3)	0.01028(12)	Mn _{1.00}
Cu1	2/3	1/3	0.68694(5)	0.0104(5)	Cu _{0.663(7)}
Cu2	1/3	2/3	0.99682(19)	0.0162(19)	Cu _{0.190(6)}
01	0	0	0.89791(15)	0.0155(10)	O _{1.00}
02	1/3	2/3	0.17418(14)	0.0153(10)	O _{1.00}
03	1/3	2/3	0.76863(13)	0.0091(8)	O _{1.00}
O4	0.9104(5)	0.5555(4)	0.97864(9)	0.0162(6)	O _{1.00}
05	0.5305(4)	0.8340(4)	0.90522(8)	0.0115(5)	O _{1.00}
06	0.7168(4)	0.5446(4)	0.90093(9)	0.0130(6)	O _{1.00}
07	0.8341(4)	0.7922(4)	0.83826(9)	0.0105(5)	O _{1.00}
08	0.4810(4)	0.8724(4)	0.82948(8)	0.0125(5)	O _{1.00}
09	0.6318(4)	0.4991(4)	0.77256(8)	0.0105(5)	O _{1.00}
010	0.8531(4)	0.7722(4)	0.70923(9)	0.0129(5)	O _{1.00}
011	0.5380(4)	0.8815(4)	0.70443(8)	0.0108(5)	O _{1.00}
OH1	0.7869(4)	0.8217(4)	0.96239(9)	0.0125(5)	O _{1.00}
H1	0.697(8)	0.719(8)	0.9679(15)	0.008(12)	$H_{1.00}$
OH2	0.7648(4)	0.9183(4)	0.77301(9)	0.0116(5)	O _{1.00}
H2	0.758(11)	0.852(11)	0.793(2)	0.04(2)	$H_{1.00}$

301 Table 2. Selected bond distances (Å) in dixenite. Bond distances in italics are to second

302

cation occupant in split site and are not included in mean

303						
304	As1 -	Distance	As2 -	Distance	As3 -	Distance
305	O5(×3)	1.7782(31)	O9(×3)	1.7567(29)	O4	1.7300(32)
306	Mean	1.778	Mean	1.757	O11	1.7634(28)
307					O10	1.7684(30)
308					Mean	1.754
309						
310	T1SI -	Distance	T2AS -	Distance	T3AS -	Distance
311	03	1.6320(53)	01	1.6544(56)	O6(×3)	1.6634(30)
312	O8(×3)	1.6438(32)	O7(×3)	1.6694(28)	O2	1.7076(54)
313	Mean	1.641	Mean	1.666	Mean	1.674
314						
315	M1 -	Distance	M2 -	Distance	M2A -	Distance
316	OH1(×3)	2.1533(33)	M2A	1.1580(250)	M2	1.1580(250)
317	O11(×3)	2.2318(30)	O4(×3)	2.0094(32)	O4(×3)	2.0036(78)
318	Mean	2.193	O10(×3)	2.4437(35)	Mean	2.004
319			Mean	2.227		
320						
321	M3 -	Distance	M4 -	Distance	M5 -	Distance
322	OH2(×3)) 2.0455(31)	O5	2.1537(30)	08	2.0859(30)
323	O10(×3)	2.0675(31)	OH1	2.1584(32)	O5	2.1177(30)
324	Mean	2.057	O6	2.1679(31)	O6	2.1282(31)
325			OH1	2.1701(31)	O7	2.1614(31)
326			O4	2.1832(32)	O5	2.4136(32)
327			O1	2.3985(30)	08	2.7354(32)
328			Mean	2.205	Mean	2.274
329						
330	M6 -	Distance	M7 -	Distance	Cul -	Distance
331	08	2.0777(31)	OH2	2.1803(30)	Cu2	0.8789(67)
332	07	2.1815(30)	09	2.1973(29)	As3(×3	3) 2.2361(4)
333	OH2	2.2454(32)	011	2.2112(29)	As2	2.3400(18)
334	09	2.2460(30)	03	2.2347(25)	Mean	2.262
335	09'	2.2805(30)	011	2.2762(30)		
336	02	2.3968(29)	010	2.2928(31)		
337	Mean	2.238	Mean	2.232		
338	a a	D: /				
339	Cu2 -	Distance				
340	Cul	0.8789(67)				
341	$As3(\times 3)$	2.3654(24)				
342	As1	2.5001(72)				
343	Mean	2.399				

	Table 3 Bond-valence (vu) table for discrite													
	<i>M</i> (1)	<i>M</i> (2)	<i>M</i> (3)	<i>M</i> (4)	M(5)	M(6)	M(7)	T(1)	T(2)	T(3)	As(1)	As(2)	<i>As</i> (3)	Σ
O(1)				$0.21^{x_3} \rightarrow$					1.17					1.80
O(2)						$0.21^{x3\rightarrow}$				1.08				1.71
O(3)							$0.31^{x3 \rightarrow}$	1.01						1.94
O(4)		$0.52^{x3\downarrow}$		0.34									1.11	1.97
O(5)				0.37	0.40						0.99 ^{x3↓}			1.96
					0.20									
O(6)				0.36	0.39					$1.22^{x3\downarrow}$				1.97
O(7)					0.36	0.35			$1.12^{x3\downarrow}$					1.83
O(8)					0.44	0.45		$0.98^{x3\downarrow}$						1.96
					0.09									
O(9)						0.30	0.33					$1.04^{x3\downarrow}$		1.94
						0.27								
O(10)		$0.18^{x3\downarrow}$	$0.43^{x3\downarrow}$				0.27						1.01	1.89
O(11)	$0.31^{x3\downarrow}$						0.32						1.03	1.94
							0.28							
OH(1)	$0.37^{x3\downarrow}$			0.37										1.10
				0.36										
OH(2)			$0.46^{x3\downarrow}$			0.30	0.35							1.11
Σ	2.04	2.10	2.67	2.01	1.88	1.88	1.86	3.95	4.53	4.74	2.97	3.12	3.15	
	2	3	3	2	2	2	2	4.05	4.54	4.72	3	3	3	

345







358

359

FIGURE 2. Plan views of layers in dixenite compared to similar layers in mcgovernite. Detailed
 comparisons given in text.

363





- 371
- 372
- **Figure 4.** Coordination of split Cu1 and Cu2 sites in dixenite.
- 374



376

377



379



Figure 6. Comparison of the M(4)-As(1) layer in dixenite with M(2)-As(3) layer in mcgovernite.