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2	EntraPT: an online platform for elastic geothermobarometry
3	Mattia Luca Mazzucchelli <sup>1</sup> , Ross John Angel <sup>2</sup> , Matteo Alvaro <sup>1</sup>
4	<sup>1</sup> Department of Earth and Environmental Sciences, University of Pavia, Via A. Ferrata, 1, Pavia,
5	27100, Italy
6	<sup>2</sup> IGG-CNR, Via G. Gradenigo 6, Padova, 35131, Italy
7	Corresponding author: Mattia L. Mazzucchelli (mattialuca.mazzucchelli@unipv.it)
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9	Abstract
10	EntraPT is a web-based application for elastic geobarometry freely accessible at the "Fiorenzo
11	Mazzi" experimental mineralogy lab website (www.mineralogylab.com/software). It provides an
12	easy-to-use tool to calculate the entrapment conditions of inclusions, with error propagation, from
13	the residual strain measured in mineral inclusions. EntraPT establishes a method and a workflow
14	to import and analyze the measured residual strains, correctly calculate the mean stress in the
15	inclusions, compute the entrapment isomekes with uncertainty estimation, and visualize all the
16	results in relevant graphs. It enables the user to avoid the many possible errors that can arise from
17	manual handling of the data, and from the numerous steps required in geobarometry calculations.
18	All of the data, parameters and settings are stored in a consistent format and can be exported as

20 researchers to store and/or share their data easily, making the checking and the comparison of data

project files and spreadsheets, and imported back to EntraPT for further analysis. This allows

21 and results reliable. EntraPT is an online tool that does not require any download and/or

installation, and it will be updated in the future with new functionalities made available fromadvances in the development of elastic geobarometry.

24 Keywords: EntraPT program, elastic geothermobarometry, host-inclusion systems, elasticity

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# Introduction

Mineral inclusions entrapped in their mineral hosts have provided fundamental information on 27 geological processes, first through simple phase identification (e.g. Chopin 1984). More recently, 28 advances in mineral physics data and the widesperead availability of analysis tools such as Raman 29 spectrometers has resulted in the rapid development of the field of piezobarometry for a wide 30 31 range of geological settings. Residual strains are developed in inclusions because of the contrast between their elastic properties and those of their host. If the residual strain is measured and 32 interpreted correctly, the conditions at entrapment  $(P_{trap}, T_{trap})$  can be determined by using the 33 elastic properties of the host and the inclusion. The basic concept has been known for a long time 34 (e.g. Rosenfeld and Chase 1961; Gillet et al., 1984; Izraeli et al., 1999; Zhang, 1998). Current 35 applications are based on the assumptions that both the host and inclusion minerals are elastically 36 isotropic and that the inclusion is an isolated sphere. Under such conditions the stress in the 37 inclusion will be a hydrostatic pressure. It can be determined by measuring the shift of the Raman 38 lines of the inclusion, or the unit-cell volume of the inclusion by diffraction. In absence of plastic 39 or viscous deformation (e.g. Zhong et al., 2018; Zhong et al., 2020), the determination of possible 40 entrapment conditions, the entrapment isomeke, is then a straight-forward calculation that is 41 implemented in a number of freeware programs such as EosFit7c (Angel et al., 2014a), EosFit-42

43 Pinc (Angel et al., 2017a), QuIB-calc (Ashley et al., 2014), or the algorithm published in Kohn
44 (2014).

However, no mineral crystal is elastically isotropic, and no inclusion is a perfect isolated sphere. 45 As a consequence the stress field inside the inclusion is not hydrostatic when the host is in the 46 laboratory (Eshelby, 1959; Murri et al., 2018; Mazzucchelli et al., 2018; Mazzucchelli et al., 2019). 47 The shifts of Raman lines of the inclusion are therefore different from those under hydrostatic 48 conditions (e.g. Anzolini et al., 2018; Thomas and Spear, 2018). Using hydrostatic calibrations of 49 Raman shifts (e.g. Enami et al., 2007; Ashley et al., 2016) can lead to significantly inaccurate 50 estimates of the residual "pressure" in the inclusion and, in turn, of the entrapment conditions, 51 leading to fundamental misinterpretations of geological processes. The shifts of Raman lines of 52 53 inclusions are correctly interpreted through the mode Grueneisen tensors of the inclusion crystal 54 (e.g. Grüneisen, 1926; Barron et al., 1980; Angel et al., 2019). These tensors allow the strains of 55 the inclusion crystal relative to a free crystal to be determined from the measured Raman shifts. Or, the strains can be determined by direct in-situ diffraction measurements of the inclusion (e.g. 56 Nestola et al., 2011; Alvaro et al., 2020). Once the strains are determined, the residual stress and 57 the residual pressure in the inclusion can be determined via the elastic tensor. Bonazzi et al. (2019) 58 showed that the entrapment conditions of synthetic quartz inclusions can be determined in this way 59 even in presence of deviatoric stress. 60

The determination of reliable geological conditions of inclusion entrapment from non-spherical anisotropic inclusions is thus far more complex, involving several intermediate steps with tensor calculations, than the isotropic approach embodied in existing software. Shape corrections are made with the relaxation tensor of the inclusion which, for non-ellipsoidal inclusions or anisotropic hosts, must be determined with numerical calculations such as finite element modeling

(Mazzucchelli et al., 2018; Mazzucchelli et al., 2019; Morganti et al., 2020; Alvaro et al., 2020). 66 Therefore, we have developed EntraPT, a web application which implements the methods 67 developed by Angel et al. (2019) and Murri et al. (2018) and applied by Bonazzi et al. (2019), to 68 correctly calculate the mean stress in inclusions from the measured strains, and then to calculate 69 entrapment isomekes with the isotropic model including the propagation of uncertainties. EntraPT 70 71 is freely accessible after registration. It can be used to process strain data from a single inclusion or large datasets of residual strains of multiple inclusions to obtain entrapment conditions and to 72 easily generate plots. In this paper we introduce the EntraPT program, document its methods and 73 illustrate its functionalities. A guided example of a step-by-step analysis of a dataset of guartz 74 inclusions entrapped in almandine (Bonazzi et al., 2019) is provided in the Supplementary 75 material. 76

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# **Program description**

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The successful determination of the P-T of metamorphism using elastic geothermobarometry 79 80 ideally requires the analysis of a large number of inclusions (e.g. Bonazzi et al., 2019). Because of the contrast in elastic properties between the two minerals, some host-inclusion systems are good 81 geothermometers (e.g. zircon in garnet) while other are good barometers (e.g. quartz in garnet), 82 83 and their coexistence in the same rock might provide a constraint on the metamorphic P-T path of the rock, because of the different slopes of their isomekes (e.g. Zhong et al., 2019). Furthermore, 84 to interpret correctly the conditions of metamorphism it is necessary to account for all of the 85 uncertainties from the measurement of the residual strain to the calculation of the entrapment 86 conditions. EntraPT is a MATLAB® based online application with a Graphical User Interface 87 (GUI) that can be used to (see Fig. 1): (1) import and visualize the residual strain and the associated 88

uncertainties from measurements in inclusions (Fig 1c-f); (2) correctly calculate the mean stress in inclusions from the measured strains (Fig. 1g); (3) calculate entrapment isomekes with the isotropic model of Angel, et al. (2017a) with uncertainty propagation; (4) plot the entrapment isomekes for multiple inclusions simultaneously (Fig. 1h); (5) export all the data and results as project files or spreadsheets to store and share them (Fig. 1i).

94 EntraPT is freely accessible upon registration as an online application at the website of the "Fiorenzo Mazzi" Experimental Mineralogy Lab (www.mineralogylab.com/software). EntraPT is 95 supported on most of the common browsers and operating systems (Windows®, macOS®, 96 Linux<sup>®</sup>) and no download and/or installation is required. The app was developed with 97 MATLAB® AppDesigner and deployed with the MATLAB® Compiler on Web App Server. For 98 all of the calculations based on the equations of state (EoS) of the minerals (e.g. the calculation of 99 the residual pressure and of the entrapment isomeke), EntraPT relies on Eosfit7 (Angel et al., 100 2014a) a stable and efficient Fortran code that has been validated over many years. Eosfit7n, a 101 fast version of the program without a console or GUI for direct user interaction, is embedded in 102 EntraPT. The instructions and EoS parameters for calculations are sent from EntraPT to Eosfit7n 103 as files. Once the calculations are terminated, the output generated by Eosfit7n is read back into 104 105 EntraPT. The use of EosFit7n as a separate executable ensures that in the future EntraPT will be able to access new EoS forms and other developments introduced into the CrysFML library 106 107 (Rodriguez-Carvajal and González-Platas, 2003) and EosFit program suite.

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The user interface of EntraPT consists of three main tabs: *Add New Analyses, Calculate Entrapment* and *View Data* (Fig. 2a, b, c) designed to guide the user through the required work path. The workflow, from the definition of the host-inclusion system to the calculation of

112	entrapment conditions and the data visualization, is illustrated in detail in the Supplementary
113	material, taking as an example the dataset of quartz inclusions entrapped in almandine reported by
114	Bonazzi et al. (2019). EntraPT works on analyses. Each analysis is a container of all of the
115	information relative to one measurement of residual strain and the calculations performed on it.
116	An <i>analysis</i> is defined at least by:
117	• A label
118	• A pair of mineral phases: the host and the inclusion phase
119	• The elastic properties of the host and the inclusion (EoS, stiffness tensors)
120	• The geometry of the host-inclusion system
121	• The relative orientation between the host and the inclusion
122	• The residual strain state defined by the components of residual strain vector, and its
123	covariance matrix
124	Other data that can be optionally associated with an <i>analysis</i> are:
125	• The labels of the sample, thin section, host, inclusion and of the point analysis
126	• A text description
127	• The results of the calculation of its entrapment conditions (i.e. <i>P</i> and <i>T</i> of the entrapment
128	isomeke, and all the intermediate results, such as residual stress, unrelaxed strain, their
129	covariance matrices, etc)
130	With this definition, two measurements of residual strain taken in the same inclusion crystal are
131	stored and treated as separate analyses in EntraPT. In fact, even if they belong to the same host-
132	inclusion system (same mineral phases, geometry and orientation) they can differ in the residual
133	strain. For the same reason, measurements from several inclusions in the same host crystal are also
134	considered as separate analyses. This allows for great flexibility in storing and processing the

measurements, since each of them can have different data and metadata and can be handled separately.

EntraPT makes use of the Voigt notation (Voigt, 1910) to represent tensors, and this notation will be assumed thoughout this paper. Therefore 2<sup>nd</sup> rank tensors are represented as vectors and 4<sup>th</sup> rank tensors as matrices. The mapping of the indices between the tensor notation and the vector (Voigt) notation is the following:

Tensor notation	11	22	33	23	13	12
Voigt notation	1	2	3	4	5	6

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### 143 **Properties of the host-inclusion system**

In EntraPT the user first adds one or more new analyses to the project from the Add New Analyses 144 tab (Fig. 2a), where the host and the inclusion minerals, the geometry of the system and their 145 relative orientation must be chosen. The calculation of the entrapment conditions requires the 146 147 elastic properties of the minerals. EntraPT is based on an internally-consistent database of elastic properties (Table 1). The database contains for each mineral phase the volume EoS and, for 148 inclusions, the 4<sup>th</sup> rank elastic tensor at room conditions. The elastic tensors are taken from 149 150 experimental determinations reported in the literature, and, if necessary, their components are rescaled to ensure that the Reuss bulk modulus equals that of the volume EoS (Table 1) at room 151 conditions. This ensures consistency throughout the entire calculation. During the rescaling, care 152 is taken that the degree of anisotropy (evaluated through the Universal anisotropic index, 153 Ranganathan and Ostoja-Starzewski, 2008) is not altered (see Mazzucchelli et al., 2019 for further 154 details). Only phases with published and validated elastic properties and EoS are included in the 155

program database. In the current version aluminosilicate garnet endmembers (e.g. pyrope, almandine and grossular) and diamond are available as hosts. The inclusion can be chosen among quartz, zircon, diamond and garnet. Additonal hosts and inclusion minerals will be added as further internally consistent elastic data and EoS become available.

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# 161 Residual strain

Once the properties of the system are set, the user must input the residual strain measured in the 162 inclusion for each analysis (Fig. 2g). As shown by Murri et al. (2019) and Angel et al. (2019) the 163 strain state can be obtained from the measured Raman shifts by applying the concept of the 164 phonon-mode Gruneisen tensor. This can be done easily for inclusions such as quartz and zircon 165 using stRAinMAN (Angel et al., 2019), a free program that can be downloaded at 166 http://www.rossangel.com/. stRAinMAN allows the user to load the measured Raman shifts and 167 to export the resulting strain components and the associated standard deviations and correlations 168 169 in a logfile. For a uniaxial inclusion without symmetry breaking, stRAinMAN provides an output with the sum of the strain components (in Voigt notation)  $\varepsilon_1 + \varepsilon_2$ , the component  $\varepsilon_3$ , their 170 estimated standard deviations (esd) and the covariance between  $\varepsilon_1 + \varepsilon_2$  and  $\varepsilon_3$ . If the strain is 171 measured with X-ray diffraction, measurements need to be carefully interpreted because irregular 172 and faceted inclusions may exhibit stress and strain gradients. In this case, diffraction 173 measurements only provide some average value of strain over the whole inclusion volume that 174 cannot be used directly in elastic geobarometry (e.g. Campomenosi et al., 2018; Mazzucchelli et 175 al., 2018; Murri et al., 2018, Alvaro et al., 2020). However, with XRD the value of each strain 176 177 component (together with their esd and covariances) can be determined independently.

The user can input the components of the residual strain and the associated uncertainties into 178 EntraPT, for each analysis. The esd's and the covariances associated with the residual strains are 179 needed to propagate the uncertainties into the residual pressure and the entrapment isomeke. Since 180 the application of elastic geobarometry usually requires the processing of several measurements 181 performed on several inclusions, the user has the option of importing a text file with a list of 182 183 analyses with their strains and the associated uncertainties. The input file can be selected with a file-browser window and it is uploaded to the server via a secure SSL connection. This procedure, 184 applied to the data of Bonazzi et al. (2019), is described in detail in the Supplementary material 185 and shown in Fig. S2. Template files are provided with the program on the website 186 mineralogylab.com that can be edited with a spreadsheet editor, to preserve the tab delimited 187 structure. Each row of this file stores the data of one *analysis*, i.e. of one measurement. All of the 188 analyses in one input file must belong to the same host-inclusion system (i.e. same host and 189 inclusion phases). The file format allows each analysis to be identified (sample, thin section, host, 190 inclusion and point analysis) and notes associated to the analysis to be recorded. The content of 191 the remaining columns depends on the symmetry of the residual strain, which corresponds to the 192 symmetry of the inclusion in absence of symmetry breaking (see Table S1 in the Supplementary 193 194 material and the \*.dat file included in the Deposit items for an example). The measured strain components and the statistical parameters (esd, covariances) can be easily copied and pasted from 195 196 the stRAinMAN logfile to this input file using any spreadsheet editor. During the import process, 197 EntraPT checks that the file is readable and that the data are consistent. Depending on the symmetry of the inclusion selected by the user in the app, a minimum number of independent strain 198 components is required. Moreover, the symmetry of the inclusion prescribes the equality between 199 200 some of the components of the strain. A check is also performed on the consistency of the provided

esd and covariances, since the resulting covariance matrix must be positive definite. Analyses that do not satisfy these requirements are discarded during the import process, and a detailed message describing the errors is shown to the user. EntraPT requires the full 6 x 6 covariance matrix  $V^{\varepsilon}$ of the residual strain values in order to propagate the uncertainties through the calculations of the residual pressure and the entrapment isomeke. For a general state of residual strain with six independent components in Voigt notation ( $\varepsilon_1, \varepsilon_2, \varepsilon_3, \varepsilon_4, \varepsilon_5, \varepsilon_6$ ), the fully-symmetric covariance matrix  $V^{\varepsilon}$  is defined as:

$$\boldsymbol{V}^{\varepsilon} = \begin{bmatrix} esd(\varepsilon_{1})^{2} & cov(\varepsilon_{1},\varepsilon_{2}) & cov(\varepsilon_{1},\varepsilon_{3}) & cov(\varepsilon_{1},\varepsilon_{4}) & cov(\varepsilon_{1},\varepsilon_{5}) & cov(\varepsilon_{1},\varepsilon_{6}) \\ & esd(\varepsilon_{2})^{2} & cov(\varepsilon_{2},\varepsilon_{3}) & cov(\varepsilon_{2},\varepsilon_{4}) & cov(\varepsilon_{2},\varepsilon_{5}) & cov(\varepsilon_{2},\varepsilon_{6}) \\ & & esd(\varepsilon_{3})^{2} & cov(\varepsilon_{3},\varepsilon_{4}) & cov(\varepsilon_{3},\varepsilon_{5}) & cov(\varepsilon_{3},\varepsilon_{6}) \\ & & & esd(\varepsilon_{4})^{2} & cov(\varepsilon_{4},\varepsilon_{5}) & cov(\varepsilon_{4},\varepsilon_{6}) \\ & & & & esd(\varepsilon_{5})^{2} & cov(\varepsilon_{5},\varepsilon_{6}) \\ & & & & & esd(\varepsilon_{6})^{2} \end{bmatrix}$$
(1)

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209 Where  $esd(\varepsilon_i)$  is the estimated standard deviation on the *i*-th component of the strain vector (in 210 Voigt notation), and  $cov(\varepsilon_i, \varepsilon_i)$  is the covariance between the strain components *i* and *j*.

The esds and the covariances of the residual strain of uniaxial and cubic inclusions are 211 automatically translated by EntraPT into the full 6 x 6 covariance matrix  $V^{\varepsilon}$  of the residual strain. 212 For uniaxial inclusions (i.e. trigonal, tetragonal and hexagonal crystal systems) without symmetry 213 breaking, two normal strain components are equal ( $\varepsilon_1 = \varepsilon_2$ ), while the third ( $\varepsilon_3$ ) is different. The 214 shear components are absent ( $\varepsilon_4 = \varepsilon_5 = \varepsilon_6 = 0$ ). When the residual strain of uniaxial inclusions 215 is determined from Raman spectroscopy through the stRAinMAN program, the strain sum  $\varepsilon_1 + \varepsilon_2$ 216 and  $\varepsilon_3$  are given as output together with the corresponding uncertainties  $esd(\varepsilon_1 + \varepsilon_2)$ ,  $esd(\varepsilon_3)$ . 217 The covariance between  $\varepsilon_1 + \varepsilon_2$  and  $\varepsilon_3$  is also computed and reported as  $cov(\varepsilon_1 + \varepsilon_2, \varepsilon_3)$ . Given 218

these parameters, and assuming that  $\varepsilon_1 = \varepsilon_2$  are completely correlated, the full covariance matrix

220 of the residual strain  $V^{\varepsilon}$  is assembled by EntraPT as:

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$$\mathbf{V}^{\varepsilon} = \begin{bmatrix} \left(\frac{esd(\varepsilon_{1} + \varepsilon_{2})}{2}\right)^{2} & \left(\frac{esd(\varepsilon_{1} + \varepsilon_{2})}{2}\right)^{2} & \frac{1}{2}cov(\varepsilon_{1} + \varepsilon_{2}, \varepsilon_{3}) & 0 & 0 & 0\\ & \left(\frac{esd(\varepsilon_{1} + \varepsilon_{2})}{2}\right)^{2} & \frac{1}{2}cov(\varepsilon_{1} + \varepsilon_{2}, \varepsilon_{3}) & 0 & 0 & 0\\ & & esd(\varepsilon_{3})^{2} & 0 & 0 & 0\\ & & & & 0 & 0\\ & & & & & 0 \end{bmatrix}$$
(2)

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For cubic crystals without symmetry breaking, the normal strain components are equal  $\varepsilon_1 = \varepsilon_2 = \varepsilon_3$ , the shear components are absent ( $\varepsilon_4 = \varepsilon_5 = \varepsilon_6 = 0$ ) and the residual volume strain is  $\varepsilon_V = \varepsilon_1 + \varepsilon_2 + \varepsilon_3$ . Given the uncertainty on the volume strain  $esd(\varepsilon_V)$ , the full covariance matrix of the residual strain  $V^{\varepsilon}$  is:

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Once the strains are imported or set in the program, the user can add the *analyses* to the current project.

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# 234 Calculation of the entrapment isomeke with uncertainties

- 235 The possible entrapment conditions, i.e. points on the entrapment isomeke (Rosenfeld and Chase,
- 1961; Angel et al., 2014b; Angel et al., 2015b), of each analysis can be calculated from the
- 237 Calculate Entrapment tab (Fig 2b). This calculation using isotropic elasticity and the full non-
- linear EoS of the host and inclusion requires the knowledge of the residual pressure of the inclusion

239 (Angel et al., 2017a). As shown by Bonazzi et al. (2019), when the residual strain of the inclusion

- is used there are two possible definitions of residual pressure  $(P_{inc})$  for anisotropic inclusions:
- 1) The residual stress can be calculated from the residual strain as:

242

$$\sigma_i = C_{ij} \varepsilon_j \tag{4}$$

243

where  $C_{ij}$  is the matrix representation in Voigt notation of the 4<sup>th</sup> rank elastic modulus tensor of the inclusion determined at room conditions. The pressure is then the negative of the mean normal stress:

247

 $P_{inc}^{strain} = -\frac{(\sigma_1 + \sigma_2 + \sigma_3)}{3}$ (5)

248

249

2) Alternatively, the residual volume strain can be found as the sum of the normal components ofthe strain:

$$\varepsilon_V = \varepsilon_1 + \varepsilon_2 + \varepsilon_3 \tag{6}$$

from which the residual pressure is obtained using the EoS of the inclusion as:

253

$$P_{inc}^{V} = f_{EoS}(\varepsilon_{V}) \tag{7}$$

254

By default, EntraPT uses model (1) for the calculation of residual pressure. However, the user can enable the *Expert mode* panel from the *Settings* menu and choose model (2) or both the models to explore a comparison of the results (the procedure is explained in details in the Example section of the Supplementary material). The covariance matrix of the residual strain is used to propagate the uncertainties on the calculated residual pressures and the isomekes. If model (1) is selected (eq. 4), the covariance matrix on the residual stress ( $V^{\sigma}$ ) in the inclusion is obtained from that on the residual strain ( $V^{\varepsilon}$ ) as:

$$\boldsymbol{V}^{\sigma} = \boldsymbol{C} \, \boldsymbol{V}^{\varepsilon} \, \boldsymbol{C}^{T} \tag{8}$$

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Where C is the matrix representation in Voigt notation of the 4<sup>th</sup> rank elastic modulus tensor of the inclusion, and  $C^{T}$  its transpose. Eq. (8) assumes that the uncertainties of the elastic components (usually less than 2% e.g. Lakshtanov et al., 2007) are negligible compared to the uncertainties of the measured strains which are typically larger than 5% (Bonazzi et al., 2019). Moreover, the covariances of the elastic components determined experimentally are often not reported in literature and cannot therefore be used for error propagation.

269 Once the residual stress and its covariance matrix are known, the standard deviation on the residual

270 pressure  $P_{inc}^{strain}$  (eq. 5) can be found as:

$$esd(P_{inc}^{strain}) = \frac{1}{3}\sqrt{V_{1,1}^{\sigma} + V_{2,2}^{\sigma} + V_{3,3}^{\sigma} + 2V_{1,2}^{\sigma} + 2V_{1,3}^{\sigma} + 2V_{2,3}^{\sigma}}$$
(9)

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On the other hand, if model (2) is selected, the uncertainty on the strain components is propagated

into the estimated standard deviation of the volume strain  $esd(\varepsilon_V)$  as:

$$esd(\varepsilon_{V}) = \sqrt{V_{1,1}^{\varepsilon} + V_{2,2}^{\varepsilon} + V_{3,3}^{\varepsilon} + 2V_{1,2}^{\varepsilon} + 2V_{1,3}^{\varepsilon} + 2V_{2,3}^{\varepsilon}}$$
(10)

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Given this uncertainty on the volume strain the standard deviation on the residual pressure  $P_{inc}^V$ (eq. 7) is found as:

$$esd(P_{inc}^{V}) = (1$$

$$(11)$$

$$(2) \left[ \left| f_{EoS}(\varepsilon_{V} + esd(\varepsilon_{V})) - f_{EoS}(\varepsilon_{V}) \right| + \left| f_{EoS}(\varepsilon_{V} - esd(\varepsilon_{V})) - f_{EoS}(\varepsilon_{V}) \right| \right]$$

277

The calculation of the possible entrapment conditions for the host-inclusion pair is performed 278 assuming isotropic elasticity with the model proposed by Angel et al. (2017a), using the non-linear 279 elasticity of the host and the inclusion. The calculation of the entrapment isomeke is performed 280 calling the specific routines in the Eosfit7n program. Since the calculation of the residual pressure 281 282 (eq. 4 - 5) requires the elastic tensors (C) of each inclusion, which are mostly only known for room P-T (1 bar, 25°C), EntraPT is necessarily restricted to calculations from strain measured when the 283 host is at room conditions. The user can set the range of temperatures for the calculation of the 284 entrapment isomeke. The uncertainties on the entrapment isomeke pressuures are estimated from 285

the standard deviation on the residual pressure  $esd(P_{inc})$ , by assuming an uncertainty on the residual pressure equivalent to one standard deviation as:

$$P_{inc}^{max} = P_{inc} + esd(P_{inc})$$

$$P_{inc}^{min} = P_{inc} - esd(P_{inc})$$
(12)

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The entrapment isomeke is computed for each value of residual pressure. Therefore, for each temperature step along the isomeke ( $T_{iso}$ ),  $P_{iso}$  is the mean value of the pressure on the isomeke, while  $P_{iso}^{max}$  and  $P_{iso}^{min}$  are the boundaries on the uncertainty, associated with one standard deviation on the residual pressure. For each  $T_{iso}$  on the isomeke the corresponding uncertainty on the  $P_{iso}$ is obtained as:

$$esd(P_{iso}) = \frac{1}{2} \left( |P_{iso}^{max} - P_{iso}| + |P_{iso}^{min} - P_{iso}| \right)$$
(13)

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Once the calculation is completed, the results of the calculation (residual stress, pressure, P-T points on the entrapment isomekes) of each *analysis* are stored in the current project. An example of the use of EntraPT to calculate the entrapment isomekes of multiple *analyses* is reported in the Supplementary material.

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### 300 Viewing and plotting data

In the *View Data* tab (Fig. 3) the user can view all of the data relative to each *analysis* and generate relevant plots. The residual strain of one or more *analyses* can be analyzed visually in a scatter plot from the *Plot Strain* page (Fig. 3d). Fig. 3f shows the  $\varepsilon_3$  vs  $\varepsilon_1$  plot with the strain data measured by Bonazzi et al. (2019) on several quartz inclusions in almandine (the procedures to

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generate such plot are described in the Supplementary material). For each *analysis*, the associated 305 confidence ellipse is shown, as obtained from the variance-covariance matrix  $V^{\varepsilon}$  determined from 306 307 the measurement of the residual strain. For a visual reference, the isochors and the lines of isotropic strain and hydrostatic stress state of a crystal of the same phase as the inclusion can be added to 308 the plot (Fig. 3h). The results of the calculation of the entrapment isomeke are shown in the *Results* 309 page (Fig. 3b). The user can also generate a *P*-*T* plot with the isomeke(s) of one *analysis* together 310 with the estimated uncertainty shown as a shaded area from the Plot Isomekes page (Fig. 3c and 311 Figs. S5 in the Supplementary material). The isomekes of multiple analyses can also be plotted at 312 the same time, even if they belong to different host-inclusion systems. This is particularly useful 313 for constraining the metamorphic conditions if several host-inclusion systems within the same rock 314 315 were measured. Since some host-inclusion systems are good barometers (e.g. quartz in garnet) while others are good thermometers (e.g. zircon in garnet), the intersection of their isomekes can 316 provide constraints on the *P*-*T* path of the rock (e.g. Zhong et al., 2019). 317

### 318 Export and import of data

319 Project files, that contain all of the data from all of the *analyses* (elastic properties, strain given as input, covariance matrices, calculation parameters, results, notes etc.) are stored in a database-like 320 format and can be downloaded to the user's computer. Project files can be imported back into 321 EntraPT, to view the data, generate the plots. Importing a project file puts EntraPT in the same 322 configuration as when the project file was created. Once a project is loaded, new *analyses* can be 323 added or existing *analyses* can be deleted. Moreover, multiple project files can be merged in 324 EntraPT to create larger databases. Such project files can be easily shared, making the checking 325 and the comparison of data and results reliable. The project file can be easily imported into 326 327 MATLAB® and custom scripts can be developed to further process the data and to generate

custom plots, taking advantage that all of the data are structured consistently in the project file. An 328 example of these scripts is included in the Deposit items, and a custom plot generated with it is 329 shown in Fig. S6 of the Supplemental material. During the export procedure from EntraPT, the 330 user can also choose to save the data to spreadsheets that can be read by any commonly used 331 spreadsheet application, such as Microsoft Excel® or LibreOffice®. To preserve the privacy of 332 the users and the complete control of their data, all of the data are deleted from the server when 333 the EntraPT session is terminated and cannot be recovered even by the server administrators. 334 Therefore, downloading the project to the users computer is the only way to have access to the 335 data after the EntraPT session is terminated. The upload and the download of the data to and from 336 the server is always performed over a secure SSL connection. 337

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### Implications

EntraPT is a web-based application freely accessible at www.mineralogylab.com/software that 339 provides an easy-to-use tool to calculate entrapment pressures with error propagation from the 340 residual strain measured in mineral inclusions. It obviates the many pitfalls that can arise from 341 342 manually handling large amounts of data in the multi-step calculations required for elastic geobarometry. EntraPT establishes a consistent workflow to import and visually analyze the 343 measured residual strains, correctly calculate the mean stress in the inclusions, compute the 344 entrapment isomekes with uncertainty estimations and plot the results. With EntraPT all of the 345 data are stored in a consistent format and can be exported as project files and spreadsheets. This 346 allows the data to be shared easily, making the checking and the comparison of data and results 347 reliable. EntraPT will be constantly updated, without requiring the user to download any program. 348 EntraPT is designed on a modular basis, which will allow its functionalities to be expanded in the 349

350	future, as autor	nated	corrections	for the	geometry	of the	system	and	the	use	of tł	ne ful	elastic
351	anisotropy of th	ne host	and the inc	lusion a	re develop	ed.							

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483	Captions
484	Figure 1. Workflow from Raman measurement to the determination of residual strain with
485	stRAinMAN (Angel et al., 2019), and the calculation of the entrapment conditions with EntraPT.
486	The Raman shifts for several modes are measured with Raman Spectroscopy (a). Knowing the
487	Gruneisen tensor of the inclusion and using stRAinMAN (b) the residual strain components of one
488	or several samples are determined and exported as a text file (c). The user imports to EntraPT these
489	strain components (d), sets the host and the inclusion minerals and the parameters of the system
490	(e). Then, with EntraPT the user can analyse the strain states of the inclusions (f) and calculate
491	their residual pressures from the individual components of the strain, using the 4 <sup>th</sup> rank stiffness
492	tensor, or from the volume strain, using the EoS of the inclusion (g). The entrapment isomekes can
493	be calculated and plotted (h), and all the data and results can be exported/imported as project files
494	(i). EntraPT can also be used to process residual strain states measured with X-ray diffraction,

following the steps (d-i) of the workflow. Inputs and outputs to and from the programs are
identified by grey and red arrows, respectively.

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Figure 2. EntraPT is structured with three main tabs: *Add New Analyses* (a), *Calculate Entrapment* (b) and View Data (c). The user has first to define one or more new analyses from the *Add New Analyses* tab. Using the panel on the right side the user can set the host and inclusion minerals (d), the geometry (e), the orientation (f) and the measured residual strains (g). This figure shows the Host&Inclusion page (d) where the host and the inclusion minerals are set, and a summary of their EoS is shown to the user.

Figure 3. Under the *View Data* tab the user can view all the data and results in the project, using 504 the Details (a), Results (b), Plot Isomekes (c) and Plot Strain (d) pages. Here is shown the visual 505 inspection of the measured residual strains that can be performed from the Plot Strain page. The 506 user can select one or more analyses from the Workspace (e). In this example, the strain 507 508 components determined at room temperature for quartz inclusions from experiments Alm-1 and Alm-2 by Bonazzi et al. (2019) are shown. By default, the strain data with their error bars and 509 510 confidence ellipses are shown in a  $\varepsilon_3$  vs  $\varepsilon_1$  plot (f), but other choices of axes can also be made. 511 The labels of the analyses, the confidence ellipses, the lines of isotropic strain and hydrostatic stress and the isochors can be shown or hidden (g, h). The range of the axes can be adjusted as 512 needed (i). 513

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**Table 1.** Equations of State (EoS) and independent components of the at room-conditions elastic (stiffness) tensor of the minerals currently available in the database of EntraPT. For internal consistency, the components of the stiffness tensor from literature were rescaled as described in the main text.

Mineral	Туре	EoS from:	$C_{ij}$ (GPa)	<i>C<sub>ij</sub></i> modified after:
Almandine	Host/inclusion	Milani et al. (2015)	$C_{11} = 300.2; \ C_{12} = 108.8; \ C_{44} = 93.3;$	Jiang et al. (2004)
Diamond	Host/inclusion	Angel, Alvaro, et al. (2015)	$C_{11} = 1078.4; \ C_{12} = 126.8; \ C_{44} = 575.7;$	Zouboulis et al. (1998)
Grossular	Host/inclusion	Milani et al. (2017)	$C_{11} = 316.7; \ C_{12} = 91.5; \ C_{44} = 102.2;$	Isaak et al. (1992)
Pyrope	Host/inclusion	Milani et al. (2015)	$C_{11} = 291.1; \ C_{12} = 100; \ C_{44} = 93;$	Sinogeikin and Bass (2002)
Zircon	Host/inclusion	Ehlers et al (submitted)	$C_{11} = 422.0; C_{12} = 70.0; C_{13} = 148.9; C_{33} = 488.0; C_{44} = 113.1; C_{66} = 48.3;$	Özkan et al. (1974)
Quartz	inclusion	Angel, Alvaro, et al. (2017)	$C_{11} = 86.1; C_{12} = 7.2; C_{13} = 11.7;$ $C_{14} = 17.7; C_{33} = 105.6; C_{44} = 59.2;$	Lakshtanov et al. (2007)

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Figure 2	(A) Fr	htroPT	

#### File Settings Help

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