	Ab initio study of structural, elastic and thermodynamic properties of Fe ₃ S at high pressure:					
	implications for planetary cores					
	Karen Valencia ^a , Aldemar De Moya ^{a,b} , Guillaume Morard ^{c,d} , Neil L. Allan ^e and Carlos Pinilla ^{a,e}					
	^a Departamento de Física y Geociencias, Universidad del Norte, Km 5 Via Puerto Colombia,					
	Barranquilla, Colombia					
	^b Departamento de Ciencias Naturales y Exactas, Universidad de la Costa, Calle 58 No. 55-66,					
	Barranquilla, Colombia					
С	Sorbonne Université, Institut de Minéralogie, de Physique des Matériaux, et de Cosmochimie					
((IMPMC), UMR CNRS 7590, IRD, Muséum National d'Histoire Naturelle, Paris, France.					
à	Now at: Université Grenoble Alpes, CNRS, ISTerre, F-38000 Grenoble					
^e School of Chemistry, University of Bristol, Cantock's Close. BS8 1TS. Bristol. United Kingdom						
	ABSTRACT					
	Using density functional theory electronic structure calculations, the equation of state,					
1	thermodynamic and elastic properties and sound wave velocities of Fe ₃ S at pressures up to 250 GPa					
	have been determined. Fe ₃ S is found to be ferromagnetic at ambient conditions but becomes non-					
	magnetic at pressures above 50 GPa. This magnetic transition changes the c/a ratio leading to a more					
İ	isotropic compressibility, and discontinuities in elastic constants and isotropic sound velocities.					
,	Thermal expansion, heat capacity and Grüneisen parameters are calculated at high pressures and					
(elevated temperatures using the quasiharmonic approximation. We estimate Fe-Fe and Fe-S force					
(constants, which vary with Fe environment, as well as the ⁵⁶ Fe/ ⁵⁴ Fe equilibrium reduced partition					
	function in Fe ₃ S and compare these results with recently reported experimental values. Finally, our					
	calculations under the conditions of the Earth's inner core allow us to estimate a S content of 2.7					

wt%S, assuming the only components of the inner core are Fe and Fe₃S, a linear variation of elastic
properties between end-members Fe and Fe₃S and that Fe₃S is kinetically stable. Possible
consequences for the core-mantle boundary of Mars are also discussed.

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	ARTICLE INFO
	Keywords:
	Fe ₃ S
	First-principles calculations
	High pressure
	Thermodynamic properties
	INTRODUCTION
	The Earth's core is composed of iron (Fe) alloyed with Ni (~5wt%) and lighter elements which
	account for a density deficit compared to pure Fe (~ 10 % for the liquid outer core and ~ 3 % for the
	solid inner core) (Poirier, 1994). The list of possible light element candidates in the outer core include
	S, Si, O, C, and H (Hirose et al. 2013; Litasov and Shatski, 2016). Sulfur is a promising candidate
1	because of its high abundance in iron meteorites (Chabot 2004), and its compatibility with the liquid
(outer-core density (Haijun et al., 2013; Morard et al., 2013) but it is also considered too volatile to be
	the major light element in the core, a constraint which imposes a limit of ~ 2 wt% on the maximum S
	content (Dreibus and Palme, 1996; McDonough, 2003). However, a recent geochemical study has
	revised this value to 6 wt%, potentially making S a major light element in the core (Mahan et al.,
	2017). To rationalise geophysical observation and to constrain the S concentration in the Earth's core,
	a detailed understanding from experiment and theory of the physical properties of compounds in the
	Fe-S binary system under extreme pressure is needed.
	Many experimental studies have been used to study the stability and physical properties of the solid
	and liquid phases in the Fe-FeS binary system under conditions of planetary interiors (Li et al., 2001;

57 Ono and Kikegawa, 2006; Chen et al., 2007; Morard et al., 2008; Kamada et al., 2012; Kamada et al.,

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58 2014). In the Fe-S phase diagram, the intermediate compound Fe_3S is stable from 21 GPa (Fei et al. 2000) up to approximately 250 GPa (Ozawa et al., 2013). Above this pressure, Fe₃S decomposes into 59 Fe₂S-B2 and Fe-HCP. Mössbauer spectroscopy (Lin et al. 2004) shows that a magnetic collapse 60 61 occurs in Fe₃S at 21 GPa. In addition, using synchrotron x-ray diffraction and laser-heated diamondanvil cells, the solubility of S in the eutectic melt has been measured at pressures from 34 GPa to 254 62

63 GPa (Mori et al., 2017).

64 It is now possible to use *ab initio* quantum mechanical methods, usually based on periodic density 65 functional theory (DFT), to study the behaviour of compounds under the high pressures and elevated 66 temperatures of the Earth's core, allowing detailed examination of the physical, mechanical and chemical properties of these materials and complementing experimental studies (Alfe et al., 2000; 67 Belonoshko et al., 2000; Alfè et al., 2002b; Belonoshko and Ahuja, 2003; Gavryushkin et al., 2016). 68 69 In this work we study the behaviour of crystalline Fe₃S structures under pressure using such methods, 70 considering the two most stable polymorphs: a face-centred cubic (AuCu₃-type) structure with $Pm\bar{3}m$ symmetry, with one formula unit per primitive unit cell (Sherman, 1997) and the 71 experimentally observed tetragonal (Fe₃P-type) structure with $I\overline{4}$ symmetry, with eight formula units 72 per primitive unit cell (Fei et al., 2000; Martin et al., 2004). We determine, as a function of pressure 73 74 and temperature, the structural and dynamical stability of these two structures shown in Figure 1. We 75 estimate properties such as thermal expansion coefficients, force constants and sound velocities as a 76 function of pressure and temperature. Such datasets can be used as a toolbox for modelling planetarycores dynamics and formation processes (Alfe et al., 2000; Aurnou, 2007; Alboussière et al., 2010; 77 Adams et al., 2015; Wicht and Sanchez, 2019). 78

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METHODS

Density functional theory (DFT) calculations were carried out using the Vienna *ab-initio* simulation package (VASP) (Kresse and Furthmüller, 1996), which expands the electronic wavefunction on a plane wave basis set and incorporates the projected augmented wave (PAW) method for pseudopotentials (Kresse and Joubert, 1999). The energy of electronic exchange and correlation was calculated using the generalized gradient approximation by the Perdew-Burke-Ernzerhof revised for solids (PBEsol) (Perdew et al., 1996; Perdew et al., 2008).

Martin et al. (2004) have previously performed first-principles calculations on the two polymorphs 87 88 of Fe₃S we consider here (Figure 1), but the authors took no explicit account of magnetic effects. We 89 have carried out calculations both spin-unrestricted (high spin, ferromagnetic) where atoms are allowed to have a net magnetic moment and also spin-restricted where the net magnetic moment is 90 constrained to be zero (low spin). Several studies on transition metal compounds have pointed out 91 92 that the GGA approximation often yields unsatisfactory results for highly correlated systems. Thus 93 in addition to GGA calculations, we have also tested the GGA+U method using the formulation of 94 Dudarev et al. (1998) with several U_{eff} parameters ranging from 0 to 4.0 eV in order to select the 95 value that reproduces best the static and magnetic properties of Fe₃S. However, most of these 96 calculations with a non-zero $U_{\rm eff}$ predict a total magnetic moment for this system at pressures over 97 100 GPa in disagreement with experiment. In particular, we have tested a value of $U_{\rm eff}$ of 1.7 eV, similar to the one suggested by Devey et al. (2009) for other FeS alloys and found using the method 98 99 of Cococcioni and de Gironcoli (2005). However, the calculated bulk modulus of the tetragonal phase 100 of Fe₃S with GGA+U is 20% larger than even the largest experimental value in Table 1. Thus, from 101 the available choices, we have selected the GGA method, which has been shown to yield good results 102 for other Fe alloys (Mookherjee, 2011; Wu et al., 2011; Gu et al., 2014).

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104	We have chosen PAW pseudopotentials for Fe with electron configuration [Ne] 3s ² 4s ¹ 3p ⁶ 3d ⁷ , while
105	for S we used a [Ne] 3s ² 3p ⁴ configuration. Calculations were performed using an energy cutoff of
106	600 eV (tested for convergence) and k-point grids were sampled ranging from $4 \times 4 \times 4$ to $11 \times 11 \times 11$
107	11 to test for total energy and force convergence with respect to grid size. We found the total energy
108	to be converged within 0.01 eV/formula unit for the tetragonal and AuCu ₃ -type phases of Fe ₃ S using
109	$4 \times 4 \times 4$ and $10 \times 10 \times 10$ mesh grids, both centred on the gamma point, respectively.

To understand the dynamical stability of these structures we have examined the vibrational properties. 110 111 Phonon frequencies and thermodynamic properties were calculated using the frozen-phonon approach as implemented in the PHONOPY code (Togo and Tanaka, 2015). All atoms not equivalent 112 by symmetry operations were displaced by 0.01 Å in a 2x1x2 supercell to obtain the forces in the 113 perturbed system and then used to construct the force constant matrix. Phonon frequencies are given 114 by the eigenvalues of the dynamical matrix. We found convergence for all calculated thermodynamic 115 116 properties using a 4x4x4 reciprocal q-space mesh. We calculate free energies and heat capacities as 117 a function of temperature and subsequently force constants and the equilibrium reduced partition function for ⁵⁶Fe/⁵⁴Fe in the solid phases of Fe₃S and Fe. These results are compared with those 118 119 recently obtained experimentally by Liu et al. (2017).

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RESULTS

122 Structural properties and equation of state

Fig. 2 shows the calculated energy vs. volume curves for the magnetic and non-magnetic solutions for the cubic and tetragonal phases. At all studied volumes the tetragonal phase has a lower internal energy than the cubic. At small volumes the difference in energy between the ferromagnetic and nonThis is the peer-reviewed, final accepted version for American Mineralogist, published by the Mineralogical Society of America. The published version is subject to change. Cite as Authors (Year) Title. American Mineralogist, in press.

DOI: https://doi.org/10.2138/am-2021-7268. http://www.minsocam.org/

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126	magnetic solutions is small, but at larger volumes above 38 Å ³ /formula unit the ferromagnetic
127	solution is appreciably lower than the non-magnetic. The results for the tetragonal ferromagnetic
128	phase in Figure 2 were fitted to a third order Birch-Murnaghan equation giving static values of V_0 =
129	339.0 Å ³ , $K_0 = 175$ GPa and $K'_0 = 5.1$. Note that V_0 here is not the volume per formula unit but the
130	volume per (tetragonal) unit cell which contains eight formula units. The calculated value of V_0 , the
131	volume at zero pressure is about 11% less than the room temperature experimental data summarised
132	in Table 1 (Fei et al., 2000; Seagle et al., 2006; Chen et al., 2007; Kamada et al., 2014). Comparison
133	with the values of Martin et al. (2004) is less straightforward since their calculations are non-
134	magnetic; our calculations suggest that V_0 for the non-magnetic state is about 6% smaller than that
135	for the magnetic, which is consistent with the conclusions of Vocadlo et al. (2002) and Ono and Mibe
136	(2010) for Fe ₃ C.

The calculated bulk modulus, K_0 , is 5 GPa higher than the largest experimental value (Fei et al., 2000) and is in reasonable agreement with all the experimental values as shown in Table 1. It it is worth noting a large variation in the experimental K_0 from 119 GPa to 170 GPa. The value of K'_0 , the first derivative of the bulk modulus, agrees with two of the larger values reported (Chen et al., 2007; Kamada et al., 2014). Overall agreement with experiment for V_0 and K_0 improves somewhat when zero point contributions and finite temperature corrections are included, as we show later.

The calculated magnetic moments of the three Fe in tetragonal Fe₃S at zero pressure are 2.12 μ_B , 1.40 μ_B and 1.82 μ_B for FeI, FeII and FeIII, respectively (see Fig. 3). All three magnetic moments decrease with increasing pressure, albeit at different rates and are predicted to go to zero at approximately 80 GPa. The variation with pressure for all three sites is very similar to those in calculations reported for the $I\bar{4}$ phase of Fe₃P (Gu et al., 2014). The calculated values are broadly consistent with values of 1.8 μ_B and 1.1 μ_B which we have estimated following Dubiel (2009) from the results of experimental synchrotron Mossbauer spectroscopy at zero pressure (Lin et al. (2004)). The experimental values are

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averages of only two magnetic hyperfine field sites derived from time spectra. This overestimation of the magnetic moment has been observed in previous first principles calculations for Fe₃C and Fe₃P, which adopt similar structures (Vocadlo et al., 2002; Gu et al., 2014). Temperature which affects the magnetic behaviour of materials due to atomic thermal vibrations that oppose the coupling forces between adjacent atomic dipoles, is not included in the calculations reported in Figure 3. Thus our calculated magnetic moments are expected to be overestimates.

156 The calculated average magnetic moment halves from $1.5\mu_B$ to $0.7\mu_B$ between 20-50 GPa. For 157 comparison, Shen et al. (2003) report a magnetic transition at 20-25 GPa using X-ray emission spectroscopy and Lin et al. (2004) a magnetic collapse at 21 GPa from time spectra of synchrotron 158 159 Mössbauer spectroscopy. The calculated pressure at which the magnetic moments are close to zero 160 is much higher than these experimentally reported pressures but, nevertheless, a broad range for the 161 calculated transition is consistent with different behaviour with pressure of the three Fe in different 162 environments in the unit cell, and with comparable calculations for Fe₃C (e.g., Ono and Mibe (2010), for Fe₃P (Gu et al. (2014)) and recent experiments on Fe₃P (Lai et al., 2020). The calculated transition 163

164 pressure in Fe₃P (Gu et al., 2014) is also considerably higher than the reported pressure.

165 The variation of the c/a ratio of the tetragonal phase (Fig. 4) with pressure from the spin unrestricted 166 calculations changes markedly around 50 GPa, consistent with the experimental values of (Kamada 167 et al., 2014). Our values for c/a at pressures above 50 GPa are around 2% lower than those reported 168 experimentally at room temperature by Seagle et al. (2006) and Kamada et al. (2014). There is considerable scatter in the experimental values which hinders comparison of experiment and theory, 169 170 but both suggest that after the magnetic transition above 75 GPa the structure compresses almost 171 isotropically. Previous experimental studies on similar compounds such as Fe₃P, have shown that in 172 the magnetic state the magnetic moments are aligned mainly in the *ab*-plane, with contributions along 173 the *c* direction close to zero (Liu et al., 1998; Liu et al., 2003). In the non-magnetic state the magnetic

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moments in the *ab*-plane are lost and the compressibility along the *a*-axis increases. Comparable behaviour has been seen in Fe₃P by Gu et al. (2016). It is tempting to associate the complex variation of c/a in Fig. 4 between 50 GPa and 75 GPa to differences in the magnitudes of the magnetic interactions of the three nonequivalent iron sites (as also seen in Fig. 3).

178 **Dynamical stability and thermal properties**

179 In order to take into account the dependence of volume and energy on temperature, we use the 180 quasiharmonic approximation (QHA), in which phonon frequencies are a function of volume but not 181 of temperature, to calculate the Helmholtz free energy (A):

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$$A(V,T) = U_0(V) + \frac{1}{2} \sum_{q,i} h \nu_i(q,V) + k_B T \sum_{q,i} \ln[1 - e^{-h\nu_i(q,V)/k_B T}]$$
(1)

183 The first, second and third terms are the static, zero point and vibrational contributions to the free energy respectively, the v_i are the frequencies of vibrational modes with wavevector **q**, h Planck's 184 constant and $k_{\rm B}$ Boltzmann's constant. The summation is performed over a 4x4x4 regular mesh in 185 reciprocal space using a 1x1x2 supercell which is sufficiently large to lead to convergence of the 186 187 vibrational frequencies to 0.01 cm⁻¹. Anharmonic effects become important at larger volumes and eventually the QHA fails. High pressures correspond to smaller volumes, so the quasiharmonic 188 189 approximation usually works well over a large range of temperatures at elevated pressures (Allan et 190 al., 1991; Allan et al., 1993).

In order to compare with experiment we have estimated the volumetric thermal expansion of tetragonal Fe₃S within the quasiharmonic approximation as a function of pressure, at temperatures for which there are experimental data (300-900 K). Calculation of the anisotropic thermal expansion requires determining the equilibrium structure at pressure P and temperature T by means of a full minimisation of the (non-equilibrium) Gibbs energy (A+PV) with respect to the lattice parameters aand c and the basis atom positions (Taylor et al., 1997; Taylor et al., 1998) at each pressure and

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197	temperature: this is too computationally expensive. Thus we have used what has been recently
198	referred to as the "statically-constrained" quasiharmonic approximation (Allan et al., 1996; Carrier et
199	al., 2007; Otero-de-la-Rosa et al., 2011). The calculated volumetric thermal expansion coefficient is
200	shown as a function of temperature for pressures of 50 GPa and above in Figure 5; all the curves in
201	this and subsequent figures are for the non-magnetic solution at 50 GPa; the spin-unrestricted results
202	are very similar. At a given pressure the variation of α with temperature is approximately linear at
203	higher temperatures. At a given temperature α decreases rapidly with pressure. At the very highest
204	volumes and temperatures (corresponding to pressures under 20 GPa at high temperatures above 1000
205	K) calculated expansion coefficients are unphysically high due probably to the breakdown of the
206	quasiharmonic approximation at large internuclear separations (Allan et al., 1991; Allan et al., 1993)
207	and so are not plotted. At very high pressures α only increases slowly with increasing temperature.
208	
209	The variation of volume with pressure for tetragonal Fe ₃ S (tetragonal phase) at different temperatures
210	is plotted in Figure 6. The inclusion of vibrational terms improves the agreement between theory and
211	experiment - the volume at 0 GPa (V_0) and 300 K extrapolated from our calculations at 50 GPa – 200
212	GPa is 3% larger than the static value and the isothermal bulk modulus decreases by 5%, with K_0
213	almost unchanged (Table 1). Nevertheless, the calculated volume is still underestimated by 8%; the
214	PBEsol functional often underestimates the equilibrium volume of metal and metal alloys.
215	
216	The dynamical stability of the tetragonal phase of Fe ₃ S was studied over a wide pressure range from
217	0 (magnetic) to 350 GPa (non-magnetic) by examination of the calculated phonon dispersion curves.

GPa; only above 50 GPa does the structure become dynamically stable. We have performed phonon

Dispersion curves for 0 and 50 GPa are shown in Figure 5. Imaginary frequencies are present at 0

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frequency calculations at pressures up to 350 GPa, values relevant to the inner core, and the $I\overline{4}$ phase is dynamically stable up to this pressure.

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Dynamic stability does not guarantee thermodynamic stability. Recently, it has been suggested that
Fe3S can decompose into pure Fe-HCP and a B2-type Fe₂S structure above 250 GPa (Tateno et al.,
2019; Thompson et al., 2020). The thermodynamic stability of Fe₃S is determined by the free energy
for the reaction:

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$$Fe_3S = Fe^{(hcp)} + Fe_2S^{(B2)}$$

Our calculated enthalpy in the static limit for this decomposition at 50 GPa is -0.5 eV. The enthalpy 228 229 of reaction decreases with pressure and at inner core conditions of 350 GPa is -1.25 eV. The negative 230 value shows that although Fe₃S is dynamically stable at 50 GPa, it is thermodynamically unstable 231 with respect to decomposition into Fe-HCP and Fe₂S-B2 even at pressures well below 250 GPa. Nevertheless, experiment shows Fe₃S is stable at least to 200 GPa (Kamada et al., 2012; Kamada et 232 233 al., 2014). Ozawa et al. (2013) only observed decomposition of Fe₃S into two phases at 271 GPa and 234 2530 Km. At the low pressure end, Fei et al. (2000) established the existence of Fe₃S at 21 GPa at 235 temperatures between 1220 K and 1673 K. Overall, observation of Fe₃S well beyond the calculated 236 range of thermodynamic stability suggests a substantial barrier to decomposition over a wider range 237 of high pressures at which Fe₃S is kinetically stable.

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The calculated temperature variation of the dimensionless thermal Grüneisen parameter, γ_{th} (= $\alpha K_T/C_V \rho$), widely used in approximate equations of state (Barron and White, 1999) is shown at different pressures in Figure 8. At 50 GPa this varies by only approximately 2% over 0-1500 K, and the variation with temperature at higher pressures even smaller. The logarithmic isothermal volume derivative of the Grüneisen parameter, $q (= \partial \ln \gamma_{\text{th}}/\partial \ln V)_T$, is only weakly dependent on the pressure, **Revision 3**

which is a common assumption made in the reduction of experimental data (Roberts and Ruppin,1971).

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Heat capacities at constant pressure and volume is plotted as a function of temperature in Figure 9. The constant pressure heat capacity is calculated using the relationship $C_P = C_V (1 + \alpha \gamma_{\text{th}} T)$, with C_V determined at the appropriate volume at the temperature and pressure of interest, determined from the calculated thermal expansion in figure 5. Our calculated value of C_V of 562 J mol⁻¹ K⁻¹ at 50 GPa and ambient temperature compares with a value of 620 J mol⁻¹ K⁻¹ estimated from inelastic neutron scattering measurements by Lin et al. (2004).

253 The increase in frequencies at high pressure is related to changes in the force constants of the Fe-Fe and Fe-S bonds, which increase as the bonds become stiffer. Fig. 10 shows the calculated force 254 constants taken from the dynamical matrix for the three types of Fe atoms in the $I\overline{4}$ structure as a 255 256 function of pressure. Our results are in good agreement with experimental data from NRIXS reported by Lin et al. (2004) and Liu et al. (2017), as shown in Fig. 10. A particular advantage of our 257 258 calculations is that, while experiment measures only an average force constant, we can assess the 259 variation of the force constants with local structural environment. The force constant for a FeI bond is slightly lower than those involving FeII and FeIII. FeII and FeIII are bonded to three sulfur atoms 260 261 whereas FeI is bonded only to two, so the effective charge on the FeI atom is lower. This difference in bonding environment leads to small differences in quantities such as the ⁵⁶Fe/⁵⁴Fe isotope 262 263 fractionation factor. The equilibrium reduced partition function (β -factor) that describes the isotopic 264 fractionation properties between a given condensed phase a and an ideal gas can be estimated within 265 the quasiharmonic approximation using the expression (Pinilla et al., 2015):

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$$\beta(a,Y) = \left[\prod_{i=1}^{3N_{at}} \prod_{\mathbf{q}} \frac{\nu_{\mathbf{q},i}^*}{\nu_{\mathbf{q},i}} \frac{e^{-h\nu_{\mathbf{q},i}^*/2k_BT}}{(1-e^{-h\nu_{\mathbf{q},i}^*/k_BT})} \frac{(1-e^{-h\nu_{\mathbf{q},i}^*/k_BT})}{e^{-h\nu_{\mathbf{q},i}^*/2k_BT}}\right]^{1/(NNq)}$$
(2)

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where the $v_{q,i}$ are the phonon frequencies with wave vector **q**. N_{at} is the number of atoms in the unitcell and *N* is the number of sites occupied by atom *Y* in the unit-cell. The $v_{q,i}^*$ correspond to analogous frequencies of the phase containing the heavy isotope. The product in equation 2 is performed on a grid containing $6x6x6 N_q$ **q**-vectors to ensure convergence with respect to grid size.

Using the expression in equation 2 we have calculated the β-factor for the isotope fractionation of ⁵⁶Fe/⁵⁴Fe obtaining an average value of δ Fe⁵⁶_{Fe₃S}=10³ln β(Fe₃S,Fe) of 9.58‰ at 50 GPa and 298 K which is in good agreement with the 10.2‰ reported by Dauphas et al. (2012). The different local environment of the three types of Fe atoms within the unit cell leads to a 1.3‰ difference in the βfactor between FeI and FeII and between FeI and FeIII which suggests that Fe₃S should be slightly richer in ⁵⁶Fe isotopes at the FeII and FeIII sites than at FeII.

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280 Elastic moduli and sound wave velocities

In Figure 11a we show the elastic constants at zero temperature calculated using the strain vs stress relationship for pressures up to 250 GPa where Fe₃S is predicted to decompose into Fe-HCP and FeS-B2. For the tetragonal structure of Fe₃S there are seven independent elastic constants: C_{11} , C_{33} , C_{44} , C_{66} , C_{12} , C_{13} and C_{16} (Mouhat and Coudert, 2014). These constants have been calculated in the static limit for the magnetic state at pressures below 50 GPa and for the non-magnetic thereafter. The elastic constants increase monotonically as a function of pressure (Fig. 11a). The increase with pressure is less marked in the non-magnetic phase (i.e., above 50 GPa) than in the magnetic (i.e., below 50 GPa)

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with the slight exception of C_{44} , C_{66} and C_{16} perhaps related to the resistance to shear by the aligned spins in the magnetic phase. Additionally, we have used these results to estimate the isotropic longitudinal (v_P) and shear-wave (v_s) velocities for pressures up to 250 GPa (Fig. 11b):

291
$$v_P = \sqrt{\frac{K + \frac{4}{3}G}{\rho}} \qquad v_S = \sqrt{\frac{G}{\rho}} \qquad (3)$$

where ρ is the density of the material and K and G the bulk modulus and shear modulus as shown in 292 Figure 11c. Our calculations include both regimes, magnetic and non magnetic. The response due to 293 294 the collapse in magnetic effects can be seen from the discontinuity in the velocities around 50 GPa. Our results are in excellent agreement with those reported by (Lin et al., 2004) for pressures up to 55 295 296 GPa. The computed P- and S- wave velocities for Fe₃S are 10% larger than those found 297 experimentally for pure Fe-HCP under the same pressure range. A linear regression of the curves in figure 11 at pressures above 100 GPa yields the equations $v_P/(m s^{-1}) = 1.1038\rho - 1883.5$ and 298 299 $v_S/(m \text{ s}^{-1}) = 0.65173\rho - 1940.5$, where the units of ρ are kg m⁻³.

300 Seismological observations indicate that the inner core is anisotropic with respect to the speed of 301 sound (Belonoshko et al., 2008; Lincot et al., 2015; Tkalcic, 2015), with the P-waves travelling faster 302 towards the pole than towards the equator. In order to address the pressure dependence of elastic 303 anisotropy in Fe₃S we calculated the directional dependence of the P and S wave velocities by solving the Christoffel equation using the method implemented by Jaeken and Cottenier (2016). Using the 304 305 calculated data for the full elastic tensor we see that Fe₃S exhibits a strong anisotropy in both 306 compressional (P) and shear (S) waves (Figure 12). In particular, there is a softness for P waves along 307 the [100] and [010] directions, and along the [110] for S waves. The effect of pressure on the 308 anisotropic behaviour is small with the hardening of P and S waves as a function of pressure (Table 309 2). The loss of magnetization has an effect on the xy plane leading to a softening of S wave vectors.

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IMPLICATIONS

Using these results for Fe₃S and assuming that this phase is still stable at 350 GPa, we assume a twocomponent ideal solid solution in order to estimate the amount of S present in the inner core at 350 GPa. Using a similar analysis to that proposed by Badro et al. (2007), the density and compressional sound velocity can be written as

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$$\rho = x\rho_1 + (1-x)\rho_2$$
 (4)

316 and

317
$$v = \frac{v_1 v_2}{(1 - x)v_1 + xv_2}$$
(5)

318 where x is the volume fraction of pure HCP iron, $\rho_1(v_1)$ and $\rho_2(v_2)$ the densities (velocities) of pure Fe-HCP and Fe₃S at inner core conditions respectively. Setting $\rho = \rho_{PREM} = 12.2 \times 10^3$ kg m⁻³ and 319 $v=v_{PREM}=10.6x10^3 \text{ m s}^{-1}$, using the equation of state of pure Fe from Badro et al. (2007) ($\rho_1=13.2x10^3$ 320 kg m⁻³ and $v_1 = (0.94-1466 \rho_1)$ m s⁻¹) and our values for v_2 as a function of density (Fig. 11b), we solve 321 equations 4 and 5 for x and ρ_2 . Thus we estimate the amount of S present in the inner core assuming 322 Fe and Fe₃S are the only components. We find x=0.56 vol% and ρ_2 =10.9x10³ kg m⁻³ leading to a 323 sulfur content in the inner core of ≈ 2.7 wt%. This is similar to the maximum silicon content expected 324 325 for the inner core (a maximum of 3 ±2wt.%Si) (Antonangeli et al., 2018), but is subject to the caution 326 about the possible thermodynamic decomposition of Fe₃S discussed above into Fe-HCP and Fe₂S-B2 327 at high pressures and temperatures. Taking into account the calculated partitioning of sulfur between 328 solid and liquid at ICB conditions (Alfè et al., 2002a), a value of 3.5 wt%S content in the liquid outer core is obtained. Such a sulfur content is not sufficient to account for the liquid outer core density 329 330 deficit (Huang et al., 2013; Morard et al., 2013), implying the requirement of another light element 331 more compatible with the liquid phase, such as oxygen (Alfè et al., 2002a).

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332	In recent years Fe isotopes have been proposed as a tool to study planetary differentiation and have
333	triggered a series of studies to investigate whether the formation of a metallic core could induce a
334	measurable isotope fractionation between chondrites and the silica part of the Earth (Poitrasson et al.,
335	2004; Weyer et al., 2005; Craddock and Dauphas, 2013). The results have been inconclusive so far
336	with some studies suggesting a non-resolvable isotope difference and others finding a clear trend
337	(Bourdon et al., 2018). Additionally, the effect of the concentration of light elements such as S, O,
338	Si, H in the Fe fractionation between the silicate part of the mantle and a metallic core at conditions
339	of core formation is not clear either (see for example (Poitrasson et al., 2009; Hin et al., 2012; Shahar
340	et al., 2015)). This effect imprinted by light elements could be a signature of the density deficit in
341	Earth's core that could lead to a way to identify the core's composition. In order to understand the
342	effect of a light element such as S in the isotope fractionation of iron we compare the average β -factor
343	of ⁵⁶ Fe/ ⁵⁴ Fe at conditions of the magma ocean (60 GPa, 3000K) calculated in this work for Fe ₃ S,
344	$\delta Fe^{56}_{Fe3S} = 10^3 ln(\beta) = 0.117\%$, with that for pure HCP iron at the same conditions $\delta Fe^{56}_{Fe} = 0.108\%$
345	(Pinilla et al., 2020). The difference $\delta Fe^{56}Fe^{-} \delta Fe^{56}Fe^{-} \delta Fe^{56}Fe^{-} \delta Fe^{56}Fe^{-} \delta Fe^{-} \delta Fe^{$
346	the findings of some authors (Shahar et al., 2015; Shahar et al., 2016) but is in better agreement with
347	Poitrasson et al. (2009); Hin et al. (2012); Liu et al. (2017) who conclude that the addition of light
348	elements have little effect on the iron isotone fractionation at the conditions of core formation up to

elements have little effect on the iron isotope fractionation at the conditions of core formation up to concentrations of 18wt% (Bourdon et al., 2018). The small dependence of the β -factor on concentration may be related to the Fe-S bond character and coordination number.

The elastic properties of Fe₃S calculated in the present study should be valuable in the near future for the study of the core-mantle boundary (CMB) in Mars. Following the results of the Insight mission (Banerdt et al., 2019), seismic wave travel times will be available and could potentially be used to reconstruct the internal structure of Mars. The pressure at the Mars CMB is estimated at 20-25 GPa (Sohl et al., 2005), pressures at which Fe₃S is experimentally stable. Mars does not possess a magnetic

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356 field, although, based on its moment of inertia and density, it is thought to have a partially liquid metallic core (Yoder et al., 2003). Recent estimates of sulfur present in the core of Mars (25-35 GPa, 357 2000-2200 K) range from 7 to 17wt% due to the large partition coefficient (log $Ds^{mel-sil} = 2.6$) 358 359 (Steentra and van Westrenen, 2018). Given such a large amount of S in the core, one of the first S 360 containing phases that should crystallize first is a magnetic tetragonal Fe₃S, the presence of which in 361 the core could be detected by the anisotropic seismic velocities we report in figures 11 and 12. If the 362 S content is over 15 wt% a possible top down crystallisation process could take place in the core of 363 Mars, with Fe₃S potentially present at the CMB (Breuer et al., 2015). Next seismological models of the Mars interior extracted from the Insight mission could benefit from our work, potentially looking 364 for a Fe₃S layer at the Mars CMB. 365

366

We have used ab-initio computer simulation methods to study the structural, mechanical and 367 thermodynamic properties of the $I\overline{4}$ phase of Fe₃S. We have shown that this phase is the most stable 368 369 at high pressure and provided an equation of state for this system. We have seen a collapse in magnetic order between 40-60 GPa. This magnetic transition affects the c/a ratio leading to a more isotropic 370 371 compressibility. Additionally, this change leads to a discontinuity in elastic constants and isotropic sound velocities. The different types of Fe atoms in the structure lead to different force constants, so 372 373 affecting quantities such as the equilibrium isotope fractionation factor. Finally, we have presented a 374 study of thermodynamic properties and provide information on thermal expansion, heat capacity and 375 Grüneisen parameters for this structure under conditions of interest for the cores of Earth and Mars. 376 We estimate that, provided Fe_3S is kinetically stable with respect to decomposition, the inner core 377 contains approximately 2.7 wt% S.

378

379	ACKNOWLEDGEMENTS
380 381 382 383	This work was performed using Granado-HPC from the Universidad del Norte and SCARF from the STFC of the UK. We acknowledge funding from COLCIENCIAS and ECOSNORD through research grants No. 2015-710-51568 (Contract No. 023-2016) and grant No. FP44842-143-2017 respectively.
384	REFERENCES CITED
385 386	Adams, M.M., Stone, D.R., Zimmerman, D.S., and Lathrop, D.P. (2015) Liquid sodium models of the Earth's core. Progress in Earth and Planetary Science, 2, 29.
387 388	Alboussière, T., Deguen, R., and Melzani, M. (2010) Melting-induced stratification above the Earth's inner core due to convective translation. Nature, 466, 744-747.
389 390	Alfe, D., Gillan, M., and Price, G. (2000) Constraints on the composition of the Earth's core from ab initio calculations. Nature, 405, 172-5.
391 392 393	Alfè, D., Gillan, M.J., and Price, G.D. (2002a) Composition and temperature of the earth's core constrained by combining ab initio calculations and seismic data. Earth and Planetary Science Letters, 195, 91- 98.
394 395	Alfè, D., Gillan, M.J., Vocadlo, L., Brodholt, J., Price, G.D., Trans, P., and Lond, R.S. (2002b) The ab initio simulation of the Earth ' s core Philosophical Transactions of the Royal Society A, 360, 1227-1244.
396 397 398	Allan, N.L., Barron, T.H.K., and Bruno, J.A.O. (1996) The zero static internal stress approximation in lattice dynamics and the calculation of isotope effects on molar volumes. Journal of Chemical Physics, 105, 8300-8303.
399 400	Allan, N.L., Braithwaite, M., Cooper, D.L., Mackrodt, W.C., and Wright, S.C. (1991) Ionic solids at high pressure and elevated temperatures: MgO (Periclase). Journal of Chemical Physics, 95, 6792.
401 402 403	Allan, N.L., Braithwaite, M., Cooper, D.L., Petch, B., and Mackrodt, W.C. (1993) Ionic halides and oxides at high pressure: Calculated hugoniots, isotherms and thermal pressures. Journal of Chemical Society, Faraday Transactions, 89, 4369–4374.
404 405 406 407	Antonangeli, D., Morard, G., Paolasini, L., Garbarino, G., Murphy, C.A., Edmund, E., Decremps, F., Fiquet, G., Bosak, A., Mezouar, M., and Fei, Y. (2018) Sound velocities and density measurements of solid hcp- Fe and hcp-Fe—Si (9 wt.%) alloy at high pressure: Constraints on the Si abundance in the Earth's inner core. Earth and Planetary Science Letters, 482, 446-453.
408 409	Aurnou, J.M. (2007) Planetary core dynamics and convective heat transfer scaling. Geophysical & Astrophysical Fluid Dynamics, 101, 327-345.

410 411	Badro, J., Fiquet, G., Guyot, F., Gregoryanz, E., Occelli, F., Antonangeli, D., and d'Astuto, M. (2007) Effect of light elements on the sound velocities in solid iron : Implications for the composition of Earth ' s				
412	core. Earth and Planetary Science Letters, 254, 233-238.				
413	Banerdt, W.B., Smrekar, S., Antonangeli, D., Asmar, S., Banfield, D., Beghein, C., Bowles, N., Bozdag, E., Chi,				
414	P., Christesen, U., John, C., Gareth, C., Ingrid, D., Véronique, D., Matthew, F., Bill, F., Raphael, G.,				
415	Jim, G., Domenico, G., Matt, G., John, G., Matthias, G., Jurek, G., Troy, H., Jessica, L. Catherine, J.,				
416	Günter K Taichi K Sharon K Scott K Brigitte K-F Mark I Philippe I Balph I Justin M				
417	Ludovic M Scott M Chloë M David M Antoine M Paul M Nils M L Seijchi N Claire N				
/10	Erancis N. Mark D. Tom D. Ana Catalina D. Jose A.P. M. Chris P. Nick S. Matt S. Avmeric S.				
410	Fidilits, N., Wark, P., Tom, P., And-Catalina, P., Jose, A.RWi, Chins, R., Nick, S., Watt, S., Ayment, S.,				
419	Tilman, S., Sabine, S., Nick, T., Jeroen, T., Nicholas, W., Renee, W., and Mark, W. (2019) Insight-the				
420	first three months on mars. 50th Lunar and planetary science conference 2019, Cont. 2132.				
421 422	Barron, T.H.K., and White, G.K. (1999) Heat capacity and thermal expansion at low temperatures. 356 p. Kluwer, New York.				
423	Belonoshko A B, and Abuja B. (2003) Stability of the body-centred-cubic phase of iron in the Earth's				
123	inner core Nature 424 1022-1024				
727	inner core. Nature, 424, 1032-1034.				
425	Belonoshko, A.B., Ahuja, R., and Johansson, B. (2000) Quasi – Ab Initio Molecular Dynamic Study of Fe				
426	Melting Physical Review Letters, 84, 3638				
120	Weiting. Thysical Neview Letters, 64, 5656.				
427	Belonoshko, A.B., Skorodumova, N.V., Rosengren, A., and Johansson, B. (2008) Elastic anisotropy of Earth's				
428	inner core. Science. 319. 797.				
429	Bourdon, B., Roskosz, M., and Hin, R.C. (2018) Isotope tracers of core formation. Earth-Science Reviews.				
430	181. 61-81.				
100					
431	Breuer, D., Rueckriemen, T., and Spohn, T. (2015) Iron snow, crystal floats, and inner-core growth: modes of				
432	core solidification and implications for dynamis and terrestrial planets and moons. Progress in Earth				
433	and Planetary Science 2, 39				
100					
434	Carrier, P., Wentzcovitch, R., and Tsuchiva, J. (2007) First-principles prediction of crystal structures at high				
435	temperatures using the quasiharmonic approximation. Physical Review B, 76, 64116				
155					
436	Chen, B., Gao, L., Funakoshi, Ki., and Li, J. (2007) Thermal expansion of iron-rich alloys and implications for				
437	the Earth's core. Proceedings of the National Academy of Sciences of the United States of America.				
438	104, 9162-9167				
100					
439	Cococcioni, M., and de Gironcoli, S. (2005) Linear response approach to the calculation of the effective				
440	interaction parameters in the LDA+U method. Physical Review R. 71, 035105				
110					
441	Craddock, P.R., and Dauphas, N. (2013) Iron isotopic composition of reference materials. geostandards and				
442	chondrites. Geostandards & Geoanalytical Research. 35, 101-123.				

- Dauphas, N., Roskosz, M., Alp, E.E., Golden, D.C., Sio, C.K., Tissot, F.L.H., Hu, M.Y., Zhao, J., Gao, L., and
 Morris, R.V. (2012) A general moment NRIXS approach to the determination of equilibrium Fe
 isotopic fractionation factors: Application to goethite and jarosite. Geochimica et Cosmochimica
 Acta, 94, 254-275.
- 447 Devey, A.J., Grau-Crespo, R., and de Leeuw, N.H. (2009) Electronic and magnetic properties of Fe₃S₄: GGA+U
 448 investigation. Physical Review B, 79, 195126.
- Dreibus, G., and Palme, H. (1996) Cosmochemical constraints on the sulfur content in the Earth's core.
 Geochimica et Cosmochimica Acta, 60, 1125-1130.
- 451 Dubiel, S.M. (2009) Relationship between the magnetic hyperfine field and the magnetic moment. Journal
 452 of Alloys and Compounds, 488, 18-22.
- Dudarev, S.L., Botton, G.A., Savrasov, S.Y., Humphreys, C.J., and Sutton, A.P. (1998) Electron-energy-loss
 spectra and the structural stability of nickel oxide: An LSDA+U study. Physical Review B, 57, 1505.
- Fei, Y., Li, J., Bertka, C.M., and Prewitt, C.T. (2000) Structure type and bulk modulus of Fe₃S, a new iron sulfur compound. American Mineralogist, 85, 1830-1833.
- Gavryushkin, P., Popov, Z.I., Litasov, K.D., Belonoshko, A.B., and Gavryushkin, A. (2016) Stability of B2-type
 FeS at Earth 's inner core pressures : Stability of B2-FeS at core pressures. Geophysical Research
 Letters, 43, 069374.
- Gu, T., Fei, Y., Wu, X., and Qin, S. (2014) High-pressure behavior of Fe₃P and the role of phosphorus in
 planetary cores. Earth and Planetary Science Letters, 390, 296-303.
- Gu, T., Fei, Y., Wu, X., and Quin, S. (2016) Phase stabilities and spin transitions of Fe₃(S_{1-x}P_x) at high pressure and its implications in meteorites. American Mineralogist, 101, 205-210.
- Haijun, H., Wu, S., Hu, X., Wang, Q., Wang, X., and Fei, Y. (2013) Shock compression of Fe-FeS Mixture up to
 204 GPa. Geophysical Research Letters, 40, 687-691.
- Hin, R.C., Schmidt, M.W., and Bourdon, B. (2012) Experimental evidence for the absence of iron isotope
 fractionation between metal and silicate liquids at 1 GPa and 1250–1300°C and its cosmochemical
 consequences. Geochimica et Cosmochimica Acta, 93, 164-181.
- Huang, H., Wu, S., Hu, X., Wang, Q., and Wang, X. (2013) Shock compression of Fe-FeS misture up tu 204
 GPa. Geophysical Research Letters, 40, 687-691.
- Jaeken, J., and Cottenier, S. (2016) Solving the Christoffel equation: Phase and group velocities. Computer
 Physics Communications, 207, 445-451.

- Kamada, S., Ohtani, E., Terasaki, H., Sakai, T., Miyahara, M., Ohishi, Y., and Hirao, N. (2012) Melting
 relationships in the Fe-Fe₃S system up to the outer core conditions. Earth and Planetary Science
 Letters, 359-360, 26-33.
- Kamada, S., Ohtani, E., Terasaki, H., Sakai, T., Takahashi, S., Hirao, N., and Ohishi, Y. (2014) Equation of state
 of Fe₃S at room temperature up to 2megabars. Physics of the Earth and Planetary Interiors, 228,
 106-113.
- 479 Kresse, G., and Furthmüller, J. (1996) Efficient iterative schemes for ab initio total-energy calculations using
 480 a plane-wave basis set. Physical Review B Condensed Matter and Materials Physics, 54, 11169481 11186.
- 482 Kresse, G., and Joubert, D. (1999) From ultrasoft pseudopotentials to the projector augmented-wave
 483 method. Physical Review B, 59, 1758.
- Lai, X., Zhu, F., Liu, Y., Bi, W., Zhao, J., Alp, E.E., Hu, M.Y., Zhang, D., Tkachev, S., Manghnani, M.H.,
 Prakapenka, V.B., and Chen, B. (2020) Elastic and magnetic properties of Fe3P up to core pressures:
 Phosphorus in the Earth's core. Earth and Planetary Science Letters, 531, 115974.
- Li, J.Y., Fei, Y., Mao, H.K., Hirose, K., and Shieh, S.R. (2001) Sulfur in the Earth's inner core. Earth and
 Planetary Science Letters, 193, 509-514.
- Lin, J.F., Fei, Y., Sturhahn, W., Zhao, J., Mao, H.K., and Hemley, R.J. (2004) Magnetic transition and sound
 velocities of Fe₃S at high pressure: Implications for Earth and planetary cores. Earth and Planetary
 Science Letters, 226, 33-40.
- Lincot, A., Merkel, S., and Cardin, P. (2015) Is inner core seismic anisotropy a marker for plastic flow of cubic
 iron? Geophysical Research Letters, 42, 1326-1333.
- Liu, H.-P., Andersson, Y., James, P., Satula, D., Kalska, B., Haggstrom, L., Eriksson, O., Broddefalk, A., and
 Nordblad, P. (2003) The antiferromagnetism of (Fe_{1-x}Mn_x)P, x>0.67 compounds. Journal of
 Magnetism and Magnetic Materials, 256, 117-128.
- Liu, H.-P., James, P., Broddefalk, A., Andersson, Y., Granberg, P., and Eriksson, O. (1998) Structural and
 magnetic properties of (Fe_{1-x}Co_x)P compounds: experimental and theory. Journal of Magnetism and
 Magnetic Materials, 189, 69-82.
- Liu, J., Dauphas, N., Roskosz, M., Hu, M.Y., Yang, H., Bi, W., Zhao, J., Alp, E.E., Hu, J.Y., and Lin, J.-f. (2017)
 Iron isotopic fractionation between silicate mantle and metallic core at high pressure. Nature
 Publishing Group, 8, 1-6.
- Mahan, B.M., Siebert, J., Pringle, E.A., and Moynier, F. (2017) Elemental partitioning and isotopic
 fractionation of Zn between metal and silicate and geochemical estimation of the S content of the
 Earth's core. Geochimica et Cosmochimica Acta, 196, 252-270.

506 507	Martin, P., Vocadlo, L., Alfe, D., and Price, G. (2004) An ab initio study of the relative stabilities and equations of state of Fe₃S polymorphs. Mineralogical Magazine, 65, 181-191.
508	McDonough, W.F. (2003) Compositional model for the Earth's core. Elsevier, New York.
509	Mookherjee, M. (2011) Elastic and anisotropy of Fe₃C at high pressure American Mineralogist, 96, 1530-
510	1536.
511	Morard, G., Andrault, D., Guignot, N., Sanloup, C., Mezouar, M., Petitgirard, S., and Fiquet, G. (2008) In situ
512	determination of Fe – Fe₃S phase diagram and liquid structural properties up to. Earth and
513	Planetary Science Letters, 272, 620-626.
514	Morard, G., Siebert, J., Andrault, D., Guignot, N., Garbarino, G., Guyot, F., and Antonangeli, D. (2013) The
515	Earth's core composition from high pressure density measurements of liquid iron alloys. Earth and
516	Planetary Science Letters, 373, 169-178.
517 518	Mori, Y., Ozawa, H., Hirose, K., Sinmyo, R., Shigehiko, T., Morard, G., and Ohishi, Y. (2017) Melting experiments on Fe–Fe₃S system to 254 GPa Earth and Planetary Science Letters, 464, 135-141.
519 520	Mouhat, F., and Coudert, FX. (2014) Necessary and sufficient elastic stability conditions in various crystal systems. Physical Review B, 90, 224104.
521	Ono, S., and Kikegawa, T. (2006) High-pressure study of FeS, between 20 and 120 GPa, using synchrotron X-
522	ray powder diffraction. American Mineralogist, 91, 1941-1944.
523 524	Ono, S., and Mibe, K. (2010) Magnetic transition of iron carbide at high pressures. Physics of the Earth and Planetary Interiors, 180, 1-6.
525	Otero-de-la-Rosa, A., Abassi-Perez, D., and Luana, V. (2011) First-principles prediction of crystal structures
526	at high temperatures using the quasiharmonic approximation. Computer Physics Communications,
527	182, 2232.
528	Ozawa, H., Hirose, K., Suzuki, T., Ohishi, Y., and Hirao, N. (2013) Decomposition of Fe₃S above 250 GPa.
529	Geophysical Research Letters, 40, 4845-4849.
530 531	Perdew, J.P., Burke, K., and Ernzerhof, M. (1996) Generalized gradient approximation made simple. Physical Review Letters, 77, 3865-3868.
532	Perdew, J.P., Ruzsinszky, A., Csonka, G.I., Vydrov, O.A., Scuseria, G.E., Constantin, X.Z., and Burke, K. (2008)
533	Restoring the density-gradient expansion for exchange in solids and surfaces. Physical Review
534	Letters, 100, 136406.
535 536 537	Pinilla, C., Blanchard, M., Balan, E., Natarajan, S.K., Vuilleiumier, R., and Mauri, F. (2015) Equilibrium magnesium isotope fractionation between aqueous Mg ²⁺ and carbonate minerals: Insights from path integral molecular dynamics. Geochimica et Cosmochimica Acta, 163, 126-139.

- Pinilla, C., De Moya, A., Blanchard, M., Morard, G., and Roskosz, M. (2020) First principles investigation of
 equilibrium iron isotope fractionation in FeS alloys at Earth's core formation conditions. In
 preparation.
- Poitrasson, F., Halliday, A.N., Lee, D.-C., Levasseur, S., and Teutsch, N. (2004) Iron isotope differences
 between Earth, Moon, Mars and Vesta as possible records of contrasted accretion mechanisms.
 Earth and Planetary Science Letters, 223, 253-266.
- Poitrasson, F., Roskosz, M., and Corgne, A. (2009) No iron isotope fractionation between molten alloys and
 silicate melt to 2000 °C and 7.7 GPa: Experimental evidence and implications for planetary
 differentiation and accretion. Earth and Planetary Science Letters, 278, 376-385.
- Roberts, R.W., and Ruppin, R. (1971) Volume Dependence of the Gruneisen Parameter of Alkali Halides.
 Physical Review B, 4, 2041-2046.
- 549Seagle, C.T., Campbell, A.J., Heinz, D.L., Shen, G., and Prakapenka, V.B. (2006) Thermal equation of state of550Fe₃S and implications for sulfur in Earth 's core. Journal of Geophysical Research, 111, B06209.
- Shahar, A., Hillgren, V.J., Horan, M.F., Mesa-Garcia, J., Kaufman, L.A., and Mock, T.D. (2015) Sulfur controlled iron isotope fractionation experiments of core formation in planetary bodies.
 Geochimica et Cosmochimica Acta, 150, 253-264.
- Shahar, A., Schauble, E.A., Caracas, R., Gleason, A.E., Reagan, M.M., Xiao, Y., Shu, J., and Mao, W. (2016)
 Pressure-dependent isotopic composition of iron alloys. Science, 352(6285), 580.
- Shen, G., Lin, J., Fei, Y., Mao, H., Hu, M., and Chow, P. (2003) Magnetic and Structural Transition in Fe₃S at
 High Pressures, American Geophysical Union, Fall Meeting 2003, abstract id.V31D-0961.
- Sherman, D.M. (1997) pressure and the composition of the Earth's core: constraints on S and Si vs.
 temperature. Earth and Planetary Science Letters, 153, 149-155
- 560
- 561Sohl, F., Schubert, G., and Spohn, T. (2005) Geophysical constraints on the composition and structure of the562Martian interior. Journal of Geophysical Research, 110, E12008.
- 563 Steentra, E.S., and van Westrenen, W. (2018) A synthesis of geochemical constraints on the inventory of
 564 light elements in the core of Mars. Icarus, 315, 69-78.
- Tateno, S., Ozawa, H., Hirose, K., Suzuki, T., I-Kawaguchi, S., and Hirao, N. (2019) Fe₂S: The most Fe-Rich Iron
 Sulfide at the Earth's inner Core Pressure. Geophysical Research Letters.
- 567Taylor, M.B., Barrera, G.D., Allan, N.L., and Barron, T.H.K. (1997) Free energy derivatives and structure568optimisation within quasiharmonic lattice dynamics. Physical Review B, 56, 14380-14390.

- Taylor, M.B., Barrera, G.D., Allan, N.L., Barron, T.H.K., and Mackrodt, W.C. (1998) SHELL a code for lattice
 dynamics and structure optimisation of ionic crystals. Computer Physics Communications, 109, 135 143.
- Thompson, S., Komabashi, T., Breton, H., Suehiro, S., Glazyrin, K., Pakhomova, A., and Ohishi, Y. (2020)
 Compresion experiments to 126 GPa and 2500 K and thermal equation of state of Fe3S:
 Implications for sulfur in the Earth's core. Earth and Planetary Science Letters, 534, 116080.
- Tkalcic, H. (2015) Complex inner core of the Earth: The last frontier of global seismology. Reviews of
 Geophysics, 53, 59-94.
- 577 Togo, A., and Tanaka, I. (2015) First principles phonon calculations in material sciences Scr. Mater., 108, 1-5.
- Vocadlo, L., Brodholt, J., Dobson, D.P., Kinight, K.S., Marshall, W.G., and Price, G.D. (2002) The effect of
 ferromagnetism on the equation of state of Fe₃C studied by first-principles calculations. Earth and
 Planetary Science Letters, 203, 567-575.
- Weyer, S., Anbar, A.D., Brey, G.P., Münker, C., Mezger, K., and Woodland, A.B. (2005) Iron isotope
 fractionation during planetary differentiation. Earth and Planetary Science Letters, 240, 251-264.
- Wicht, J., and Sanchez, S. (2019) Advances in geodynamo modelling. Geophysical & Astrophysical Fluid
 Dynamics, 113, 2-50.
- Wu, X., Mookherjee, M., Gu, T., and Qin, S. (2011) Elasticity and anisotropy of iron-nickel phosphides at
 high pressure. Geophysical Research Letters, 38, L20301.
- Yoder, C.F., Konopliv, A.S., Yuan, D.N., Standish, D.M., and Folkner, W.M. (2003) Fluid core size of mars
 from detection of the solar size. Science, 300, 299-303.
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600 Figure Captions

- **Figure 1.** Crystal structures of the most stable phases of Fe₃S: a) tetragonal Fe₃P-type structure $(I\overline{4})$
- and b) the cubic $AuCu_3$ -type structure.

Figure 2. Total energy (per formula unit) vs. volume for the tetragonal and cubic forms of Fe₃S

Figure 3. Calculated magnetic moment from spin-unrestricted calculations of the three nonequivalent iron atoms in the $I\overline{4}$ structure at different pressures. The black line is the calculated average.

Figure 4. Ratio of c/a as a function of pressure for tetragonal Fe₃S (ferromagnetic below 50 GPa, non-magnetic at larger pressures). The straight lines are guides for the eye.

Figure 5. Thermal expansion coefficient from spin-restricted calculations as a function of
 temperature for several pressures over the range 50-250 GPa.

Figure 6. Volume as a function of pressure for the Fe₃S tetrahedral phase at different isotherms. Full red line and green line correspond to calculations at 300 K and 900 K respectively (ferromagnetic below 50 GPa, non-magnetic at higher pressures). Brown diamonds, purple triangles, and green circles correspond to experimental values taken from Kamada et al. (2014), Chen et al. (2007) and Seagle et al. (2006) respectively.

Figure 7. Phonon dispersion curves and phonon density of states calculated for tetragonal Fe₃S at zero pressure (spin-unrestricted) and 50 GPa (spin-restricted).

Figure 8. Grüneisen parameter of tetragonal Fe₃S from spin-restricted calculations as a function of temperature at several pressures over the range 50-250 GPa.

Figure 9. Temperature dependence of the heat capacity C_p from spin-restricted calculations along the isobars of 50 GPa (green dashed line), 100 GPa (red dashed line), 150 GPa (blue dashed line) and

622 200 GPa (violet dashed line). Solid lines are corresponding values of C_v .

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623	Figure 10. Force constants of iron bonds in tetragonal Fe ₃ S as a function of pressure. "Fe average"
624	refers to the overall mean force constant. Below 50 GPa results are from spin-unrestricted calculations
625	and from spin-restricted at higher pressures.

Figure 11. Elastic properties for tetragonal Fe_3S as a function of pressure: a) calculated elastic constants. b) the isotropic P-wave and S-wave velocities, red-triangles and black-squares correspond to experimental values reported by Lin et al. 2014. c) Bulk and shear modulus estimated from the calculations. Below 50 GPa results are from spin-unrestricted calculations and from spin-restricted at higher pressures.

Figure 12: Variation of a) compressional and b) shear wave velocities as a function of propagation

632 direction and pressure, from spin-restricted calculations.

633 **Table 1.** Calculated thermal equation of state parameters for ferromagnetic tetragonal Fe₃S compared

634 with room temperature experimental values and the calculations of Martin et al. (2004). V_0 is the zero-

635 pressure volume per unit cell.

V_0 (Å ³)	K_0 (GPa)	K_0	Method	Reference
377.0	170	2.6	Exp	(Fei et al. 2000)
377.0 (fixed)	156	3.8	Exp	(Seagle et al., 2006)
377.0 (fixed)	134	5.1	Exp	(Chen et al. 2007)
377.0 (fixed)	125	5.1	Exp	(Kamada et al. 2014)
377.0 (fixed)	119	5.6	Exp	(Thompson et al. 2020)
338.3	250	4.6	DFT	(Martin et al. 2004)*
339.0	175	5.1	DFT	This study, static
346.0	166	5.2	DFT	This study, including
				thermal effects at 300 K**

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* The calculations of Martin et al. (2004) are non-magnetic (spin-restricted)

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- 638 ** Values extrapolated from our calculations at 50 GPa 200 GPa using a 3rd order Birch639 Murnaghan equation of state.
- 640
- 641
- 642 **Table 2.** Calculated sound velocity ranges (km s⁻¹) and anisotropic coefficients A (%) of Fe₃S

P(GPa)	V_P	V_{S1}	V_{S2}	A_P	A_S
50	7.95-	4.53-	3.43-	9.82	25.32
	8.81	4.60	4.53		
150	9.78-	5.34-	4.20-	7.24	25.48
	10.54	5.64	5.34		
250	10.81-	5.85-	4.63-	7.32	26.66
	11.67	6.31	5.85		

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Fig. 1



Fig. 2



Fig. 3



Fig.4













Fig. 10



Fig. 11a







Fig. 12a



Fig. 12b