1	HIGHLIGHTS AND BREAKTHROUGHS
2	
3	Seeking the most hydrous, primitive arc melts: The glass is half full
4	
5	Matthew Steele-MacInnis <sup>1,*</sup>
6	
7	<sup>1</sup> Dept. of Earth and Atmospheric Sciences, University of Alberta, Edmonton AB Canada
8	*steelema@ualberta.ca
9	
10	Keywords: Melt inclusions, glass, volatiles, quenching, H <sub>2</sub> O
11	
12	Experimental studies and petrologic constraints suggest that H <sub>2</sub> O contents of deep,
13	primitive melts in subduction settings may reach up to >15 wt% H <sub>2</sub> O (e.g., Krawczynski et al.,
14	2012). But curiously, mafic glasses preserved in melt inclusions - commonly the best available
15	tool to analyze $\rm H_2O$ contents of melts – seem to be limited to much lower values, mostly <6 wt%
16	(Plank et al., 2013). This apparent conundrum suggests that empirical results defy predictions,
17	and challenges our view of H <sub>2</sub> O in subduction-related magmatism. To address this issue,
18	Gravilenko et al. (2019) experimentally tested the quenching behavior of hydrous, mafic melts.
19	Their results demonstrate that quenching to glass becomes difficult at high H <sub>2</sub> O concentrations,
20	and that mafic melts exceeding ~9 wt% $H_2O$ are essentially unquenchable at realistic cooling
21	rates. This implies that glasses preserved in melt inclusions provide only a partial record of the
22	volatile contents of deep-seated melts, and are incapable of recording the deepest, most hydrous
23	melts. This work thus elegantly reconciles what previously appeared to be a stark contradiction
24	between prediction and observation, and adds a key piece to our evolving understanding of how
25	to analyze and interpret melt inclusions.
26	The H <sub>2</sub> O concentrations of melts exert a strong control on properties such as buoyancy
27	(Ochs and Lange, 1999), viscosity (Schulze et al., 1996), chemical diffusivity (Watson, 1994)
28	and explosivity (Sparks, 1978), as well as the ore-forming potential of arc magmas (Hedenquist
29	and Lowenstern, 1994). The $H_2O$ contents of arc magmas are also central to quantifying and
30	interpreting global geochemical cycling between Earth's surface and deep interior (Bodnar et al.,
31	2013). Moreover, H <sub>2</sub> O contents of melts are widely used to evaluate depths of magmatic

plumbing systems, based on the thermodynamic relationship between pressure and solubility of
volatiles (Audétat and Lowenstern, 2014, and references therein). However, the H<sub>2</sub>O contents of
pre-eruptive melts are also elusive parameters. Experimentally calibrated proxies have been
developed to estimate H<sub>2</sub>O contents of melts based on mineral equilibria (e.g., Krawczynski et
al., 2012), but commonly, the only available tool to directly quantify the H<sub>2</sub>O (and other volatile)
contents of pre-eruptive melts is by analysis of melt inclusions (Audétat and Lowenstern, 2014).

In recent years, a growing body of theoretical, experimental and analytical studies has 38 39 contributed new insights into the systematics of volatiles in melt inclusions, and how to best 40 analyze and interpret them. It is now widely recognized that bubbles within melt inclusions can 41 host a preponderance of the bulk CO<sub>2</sub> (Moore et al., 2015) and H<sub>2</sub>O (Esposito et al., 2016), and 42 that H<sub>2</sub>O concentrations can be rapidly modified by diffusive reequilibration (Gaetani et al., 43 2012). Careful attention to these phenomena has helped elucidate the record of pre-eruptive 44 volatiles and degassing. Yet even in light of these developments, still the growing body of 45 analytical data presents some enigmatic results.

46 One of the crucial and fundamental questions that has confounded our view of volatiles in 47 subduction-related melt inclusions arises from the growing recognition that H<sub>2</sub>O (as well as CO<sub>2</sub>) 48 contents of glasses preserved in melt inclusions seem to show an unexpectedly restricted range. 49 Specifically, mafic glasses in melt inclusions from arc settings seem to be limited to  $H_2O$ contents mostly less than ~6 wt%, and never exceeding ~9 wt% (Plank et al., 2013). In contrast, 50 experimental phase equilibria consistently predict much higher  $H_2O$  contents, up to >15 wt% 51 52 (Krawczynski et al., 2012). This apparent contradiction fundamentally challenges our view of 53 either the fidelity of melt inclusions, or how well our experiments reproduce nature, or both. 54 Although some  $H_2O$  is likely partitioned into bubbles (Esposito et al., 2016), such 55 partitioning is unlikely to have such a dramatic effect on the measured H<sub>2</sub>O concentration in the 56 glass (Steele-MacInnis et al., 2011). Diffusive reequilibration also likely plays a role in reducing water contents in melt inclusions (Gaetani et al., 2012). But neither process is not expected to 57 58 yield such a consistent threshold of H<sub>2</sub>O across the breadth of thousands of reported analyses, 59 which is moreover so far below experimental predictions. What then limits melt inclusion H<sub>2</sub>O 60 contents? Could it be that magmas related to subduction have only half the amount of water 61 implied by experimental studies?

62 On page XXXX of this issue, Gravilenko et al. (2019) test an alternative hypothesis, that 63 the upper limit of H<sub>2</sub>O contents of glasses preserved in melt inclusions reflects a quench control. 64 Specifically, Gravilenko et al. (2019) hypothesize that wetter melts are more difficult to quench, 65 and that the wettest melts simply cannot be quenched. This hypothesis is rooted in the well-66 known relationships between H<sub>2</sub>O concentration, viscosity and the glass transition (Mysen and 67 Richet, 2005): Wetter melts are less viscous, and less viscous melts are less easily quenched, 68 requiring either greater degrees of undercooling or faster cooling rates in order to be quenched as 69 glass. Gravilenko et al. (2019) test this hypothesis by conducting rapid-quench experiments on 70 mafic melts over a wide range of H<sub>2</sub>O contents. Importantly, the cooling rates achieved in their 71 experiments (20-90 K/s) are consistent with best estimates for cooling rates during eruption 72 (maximum ~22 K/s; Lloyd et al., 2013). The results are remarkable. Melts that contain modest 73  $H_2O$  concentrations up to ~6 wt% consistently quench to form optically clear glass. Melts 74 containing from ~6 to ~9 wt% H<sub>2</sub>O are somewhat difficult to quench, and consistently form 75 crystallites in addition to glass. Melts exceeding 9 wt% H<sub>2</sub>O do not quench to glass, and instead 76 form friable aggregates of crystallites, vapor bubbles, and material resembling devitrified glass. 77 And compellingly, the limiting values of H<sub>2</sub>O concentrations align perfectly with the empirical 78 results from melt inclusions.

79 The results by Gravilenko et al. (2019) strongly indicate that an apparent upper limit on 80 the H<sub>2</sub>O contents of mafic melts is a consequence of the inability of wetter melts to form glassy 81 inclusions. Recently, Maclennan (2017) used numerical modeling to investigate an apparent 82 upper limit of CO<sub>2</sub> contents of melt inclusions from low-H<sub>2</sub>O settings, and concluded that high 83 CO<sub>2</sub> concentrations (resulting from high trapping pressures) give rise to intense overpressure and 84 inevitable decrepitation during magma ascent. These two studies, Maclennan (2017) and 85 Gravilenko et al. (2019), indicate complementary phenomena that control and restrict the 86 observed ranges of both H<sub>2</sub>O and CO<sub>2</sub> in melt inclusions. Similarly, Esposito et al. (2016) argued 87 that H<sub>2</sub>O exsolved into vapor bubbles in melt inclusions rapidly reacts with the surrounding glass 88 causing devitrification, which would further obscure bulk volatile concentrations. It seems likely 89 that the quench control could work in tandem with exsolution of vapor, devitrification, 90 overpressure and decrepitation, as well as diffusive reequilibration, all conspiring to prevent 91 preservation of the most volatile-rich glasses in melt inclusions.

92 The major implication of the results by Gravilenko et al. (2019) is that wetter melts are unlikely to form glassy melt inclusions, which in turn implies that glassy inclusions are incapable 93 94 of preserving a complete record of the deepest, wettest melts. On the one hand, this is a sobering 95 message. But on the other hand, this work illuminates a fundamental control, and by 96 understanding this phenomenon we stand a better chance of obtaining robust interpretations of 97 water contents of melts. This conclusion underscores the need to analyze melt inclusions in the 98 context of robust petrographic and geochemical constraints on timing of trapping and post-99 entrapment processes (Esposito et al., 2014), and provides impetus for development and 100 application of complementary techniques that are not reliant on glassy inclusions. In any case, 101 previous empirical observations showing a restricted range of H<sub>2</sub>O contents can be viewed as a natural consequence of "quenchability." This result neatly reconciles the observations with 102 103 experimental predictions, without undermining the fidelity of glassy melt inclusions or predicted 104 H<sub>2</sub>O contents of arc magmas—instead, indicating that the elusive, H<sub>2</sub>O-rich primitive arc melts 105 may simply be missing from the record of glassy melt inclusions. Interestingly, this does not 106 necessarily preclude the possibility of more H<sub>2</sub>O-rich, non-glassy melt inclusions, which would 107 appear as being partially crystallized or devitrified. Such inclusions may be easily overlooked or 108 discarded, but future workers seeking wet, primitive melts may be wise to search for and 109 specifically target such inclusions. 110 **REFERENCE CITED** 111 112 Audétat, A., Lowenstern, J.B. (2014) Melt inclusions. Treatise on Geochemistry, 13, 143–173. 113 114 Bodnar, R.J., Azbej, T., Becker, S.P., Cannatelli, C., Fall, A., Severs, M.J. (2013) Whole Earth 115 geohydrologic cycle, from the clouds to the core: The distribution of water in the 116 dynamic Earth system. Geological Society of America Special Paper, 500, 431–461 Esposito, R., Lamadrid, H., Redi D., Steele-MacInnis, M., Bodnar, R.J., Manning, C.E., De Vivo 117 B., Cannatelli C., Lima A. (2016) Detection of liquid H<sub>2</sub>O in vapor bubbles in melt 118 inclusions: Implications for magmatic fluid composition and volatile budgets of magmas? 119 American Mineralogist, 101, 1691–1695. 120

- Esposito, R., Hunter, J., Schiffbauer, J., Shimizu, N., Bodnar, R.J. (2014) An assessment of the
  reliability of melt inclusions as recorders of the pre-eruptive volatile content of magmas.
  American Mineralogist, 99, 976–998.
- Gaetani, G.A., O'Leary, J.A., Shimizu, N., Bucholz, C.E., Newville, M. (2012) Rapid
  reequilibration of H<sub>2</sub>O and oxygen fugacity in olivine-hosted melt inclusions. Geology,
  40, 915–918.
- Gravilenko, M., Krawczynski, M., Ruprecht, P., Li, W., Catalano, J. (2019) The quench control
  of water estimates in convergent margin magmas. American Mineralogist, in press.
- Hedenquist, J.W., Lowenstern, J.B. (1994) The role of magmas in the formation of hydrothermal
  ore deposits. Nature, 370, 519–527.
- Lloyd, A.S., Plank, T., Ruprecht, P., Hauri, E.H., Rose, W. (2013) Volatile loss from melt
  inclusions in pyroclasts of differing sizes. Contributions to Mineralogy and Petrology,
  165, 129–153.
- Krawczynski, M.J., Grove, T.L., Behrens, H. (2012) Amphibole stability in primitive arc
  magmas: effects of temperature, H<sub>2</sub>O content, and oxygen fugacity. Contributions to
  Mineralogy and Petrology, 164, 317–319.
- Maclennan, J. (2017) Bubble formation and decrepitation control the CO<sub>2</sub> content of olivine hosted melt inclusions. Geochemistry, Geophysics, Geosystems, 18, 597–616.
- Plank, T., Kelley, K.A., Zimmer, M.M., Hauri, E.H., Wallace, P.J. (2013) Why do arc magmas
  contain ~4 wt% water on average? Earth and Planetary Science Letters, 364, 168-179.
- Moore, L.R., Gazel, E., Tuohy, R., Lloyd, A., Esposito, R., Steele-MacInnis, M., Hauri, E.H.,
  Wallace, P.J., Plank, T. & Bodnar, R.J. (2015) Bubbles matter: An assessment of the
  contribution of vapor bubbles to melt inclusion volatile budgets. American Mineralogist,
  100, 806–823.
- 145 Mysen, B.O., Richet, P. (2005) Silicate glasses and melts. Elsevier. 560 pp.
- Ochs, F.A., Lange, R.A. (1999) The density of hydrous magmatic liquids. Science, 283, 1314–
  1317.
- Schulze, F., Behrens, H., Holtz, F., Roux, J., Johannes, W. (1996) The influence of H<sub>2</sub>O on the
  viscosity of a haplogranitic melt. American Mineralogist, 81, 1155–1165.
- Sparks, R.S.J. (1978) The dynamics of bubble formation and growth in magmas: A review and
  analysis. Journal of Volcanology and Geothermal Research, 3, 1–37.

- Steele-MacInnis, M., Esposito, R. & Bodnar, R. J. (2011) Thermodynamic model for the effect
  of post-entrapment crystallization on the H<sub>2</sub>O-CO<sub>2</sub> systematics of vapor-saturated,
  silicate melt inclusions. Journal of Petrology, 52, 2461–2482.
- Watson, E.B. (1994) Diffusion in volatile-bearing magmas. Reviews in Mineralogy, 30, 371–
  411.