This is a preprint, the final version is subject to change, of the American Mineralogist (MSA) Cite as Authors (Year) Title. American Mineralogist, in press. DOI: https://doi.org/10.2138/am-2020-7118

### **Revision 1**

1	Cassiterite crystallization experiments in alkali carbonate aqueous
2	solutions using a hydrothermal diamond-anvil cell
3	YONGCHAO LIU <sup>1</sup> , JIANKANG LI <sup>1,*</sup> , AND I-MING CHOU <sup>2</sup>
4	<sup>1</sup> MNR Key Laboratory of Metallogeny and Mineral Assessment, Institute of Mineral Resources,
5	Chinese Academy of Geological Sciences, Beijing 100037, China
6	<sup>2</sup> CAS Key Laboratory for Experimental Study under Deep-sea Extreme Conditions, Institute of
7	Deep-sea Science and Engineering, Chinese Academy of Sciences, Sanya, Hainan 572000, China
8	*Corresponding author: Li9968@126.com
9	
10	Abstract
11	Ore-forming fluids enriched in alkali carbonate are commonly observed in natural melt
12	and fluid inclusions associated with tin mineralization, particularly in granitic pegmatite.
13	However, the roles of alkali carbonates remain unclear. Hence, to investigate the roles of alkali
14	carbonate, herein, cassiterite (SnO <sub>2</sub> ) crystallization experiments in SnO <sub>2</sub> -Li <sub>2</sub> CO <sub>3</sub> -H <sub>2</sub> O and
15	SnO <sub>2</sub> -Na <sub>2</sub> CO <sub>3</sub> -H <sub>2</sub> O systems were conducted using a hydrothermal diamond-anvil cell. The
16	results showed that SnO <sub>2</sub> could dissolve into the alkali carbonate aqueous solution during
17	heating, and long prismatic cassiterite crystals grew during the subsequent cooling stage at
18	average rates of 0.61 $\times$ 10 <sup>-6</sup> –8.22 $\times$ 10 <sup>-6</sup> cm/s in length and 3.40–19.07 $\mu m^3/s$ in volume. The
19	mole fraction of cassiterite crystallized from the SnO <sub>2</sub> -Li <sub>2</sub> CO <sub>3</sub> -H <sub>2</sub> O system ranges from 0.03 to
20	0.41 mol%, which depends on the $Li_2CO_3$ content dissolved in the aqueous solution. In situ
21	Raman analysis of the alkali carbonate-rich aqueous solution in the sample chamber suggests
22	that the dissolution of SnO <sub>2</sub> can be attributed to the alkaline conditions produced by hydrolysis
23	of alkali carbonate, in which $Sn(OH)_6^{2-}$ may be a potential tin-transporting species. The

24 cassiterite crystallization conditions obtained in our SnO<sub>2</sub>-alkali carbonate-H<sub>2</sub>O systems 25 primarily fell within the 400-850 °C and 300-850 MPa temperature and pressure ranges, 26 respectively; further, cassiterite crystallization ended in rare metal pegmatite-forming conditions. 27 These crystallization features of cassiterite are similar to those formed in tin-mineralized granitic 28 pegmatites. It indicates that an alkali carbonate-rich aqueous solution or hydrous melt can work 29 as a favorable transport medium for tin and provides the necessary conditions for cassiterite 30 crystallization in granitic pegmatite, bearing the roles in decreasing the viscosity of hydrous 31 melts and enhancing the solubility of  $SnO_2$  in ore-forming melts or fluids. These roles of alkali 32 carbonate can also be extended for the mineralization of other rare metals (e.g., Li and Be) in granitic pegmatite. 33

Keywords: Cassiterite, alkali carbonate, pegmatite, crystallization, hydrothermal diamond-anvil
 cell

36

#### **37 INTRODUCTION**

38 Cassiterite (SnO<sub>2</sub>) is a predominant tin ore mineral (Kamilli et al. 2017); however, the 39 physicochemical parameters for its formation remain debatable (Audétat et al. 1998; Linnen 40 1998; Naumov et al. 2011; Korges et al. 2018). Thermodynamic calculations have indicated the presence of tin as hydroxide complexes in water (Jackson and Helgeson 1985), and it is well 41 42 known that the solubility of cassiterite is highly temperature-dependent (Štemprok 1990; Wilson 43 and Eugster 1990; Bhalla et al. 2005). The tin deposit formation process has been linked to the 44 mobility of tin as Sn(II)-Cl or F complexes in reduced acid fluids, while the precipitation of tin 45 as cassiterite requiring oxidation and acid liberation (Li 1989; Heinrich 1990; Taylor and Wall 46 1993; Linnen et al. 1995; Sherman et al. 2000; Müller and Seward 2001; Duc-Tin et al. 2007).

However, Schmidt (2018) demonstrated that tin can be transported in the form of Sn(IV)–Cl
complexes, such that HCl molality is a crucial parameter for cassiterite precipitation, and pH,
pressure, and temperature are less important parameters.

50 In addition, there is increasing evidence that the capability of CO<sub>2</sub>/carbonate-rich fluids 51 to transport some metals at high temperatures and pressures can be significant (Phillips and 52 Evans 2004; Jones et al. 2013; Kokh et al. 2017; Li et al. 2018). The mineral-forming fluids in 53 many tin deposits, especially in tin-mineralized granitic pegmatite, bear low salinity and high 54 alkali carbonate (e.g., Li<sub>2</sub>CO<sub>3</sub>, Na<sub>2</sub>CO<sub>3</sub>, Cs<sub>2</sub>CO<sub>3</sub>, K<sub>2</sub>CO<sub>3</sub>) concentrations, but only trace quantities 55 of the typical pegmatite-forming fluxes (e.g., F, Cl) (Fu et al. 1993; Fu and Kwak 1994; Rickers 56 et al. 2006; Thomas and Davidson 2016). For example, cassiterite hosts carbonate/bicarbonate-57 rich melt inclusions in the Ehrenfriedersdorf pegmatite, Germany (Thomas et al. 2006a). These 58 features suggest that alkali carbonates may play important roles in tin mineralization (Kosals 59 1976; Thomas et al. 2011). Furthermore, alkali carbonate-rich fluids are not rare in granite 60 pegmatites. Carbonate-rich melt and fluid inclusions are found in quartz, feldspar, spodumene, 61 and beryl of many pegmatites (e.g., Bikita in Zimbabwe, Tanco in Canada, Erzgebirge and 62 Zinnwald in Germany, Jiajika and Zhawulong in China) (Anderson et al. 2001; Thomas et al. 63 2006b, 2012; Li et al. 2017; Xiong et al. 2019). The carbonate content in fluid inclusions is 35.0% (g/g) NaHCO<sub>3</sub> in Orlovka pegmatite, Transbaikalia, Russia and 40.0% (g/g) KHCO<sub>3</sub> in 64 65 Nb-Ta-pegmatites from the northern part of Mozambique (Thomas et al. 2011). However, the 66 roles of alkali carbonates have rarely been investigated. Here, following the method of 67 crystallization experiment using a hydrothermal diamond-anvil cell (HDAC) (Li et al. 2013, 68 2018), cassiterite crystallization experiments in alkali carbonate-rich fluids were conducted to 69 further investigate the roles of alkali carbonates in tin mineralization.

70

# 71 EXPERIMENTAL METHODS

To explore the roles of alkali carbonate in cassiterite formation through in situ monitoring and measurement experiments, the starting materials prepared for the experiments were SnO<sub>2</sub> sintered lump (99.9% purity, Alfa Aesar), Li<sub>2</sub>CO<sub>3</sub> powder ( $\geq$ 99.0% purity, Alfa Aesar), and Na<sub>2</sub>CO<sub>3</sub> powder ( $\geq$ 99.5% purity, Alfa Aesar).

76 All experiments were conducted in an improved Bassett-type HDAC (HDAC-VT type, 77 Bassett et al. 1993, 1996; Chou 2003; Li et al. 2016). The sample chamber consisted of a hole 78 (0.5-mm diameter) at the center of a rhenium gasket (3.0-mm diameter, 0.25-mm thickness) that 79 was sealed by compressing two parallel diamond-anvil culets. The sample chamber was heated 80 via external resistive heating using two tungsten carbide seats, each of which was wrapped with a 81 Mo heating resistance wire. The temperature in the sample chamber was measured using two K-82 type thermocouples, whose junctions were attached to the diamonds near the sample chamber. 83 The thermocouples were calibrated via repeated measurements of the melting points of NaNO<sub>3</sub> 84 (306.8 °C) and NaCl (800.5 °C). The accuracy of the reported temperature was  $\pm 1$  °C. An 85 Olympus BX51TRF microscope was used to directly observe the sample, and images of the 86 entire sample chamber were continuously recorded during the experiments via a digital camera.

87 Considering the thermal expansion of a rhenium gasket during heating, the pretreatment 88 of a new rhenium gasket was conducted to ensure that the sample chamber was approximately 89 isochoric during the crystallization experiments (Chou 1996). During the pretreatment, the 90 sample chamber, which was loaded with pure water and a vapor bubble, was heated to 850 °C 91 and then cooled to ambient temperature at 20 °C/min in several repeated cycles until the vapor

4

92 bubble-disappearing temperature, measured for the vapor bubble generated during the cooling 93 process, was less than 1 °C below that measured during the previous heating process. 94 For the SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O experiments, a sintered SnO<sub>2</sub> lump, Li<sub>2</sub>CO<sub>3</sub> powder, and 95 distilled water were loaded into the HDAC sample chamber at room temperature (24 °C) (Fig. 96 1a). The sample chamber was first heated at a rate of 5 °C/min until the Li<sub>2</sub>CO<sub>3</sub> and SnO<sub>2</sub> were 97 completely dissolved into the aqueous solution, the temperature was then maintained for 30 min, 98 and the sample chamber was subsequently cooled at a rate of 1 °C/min. In situ cassiterite 99 crystallization in the sample chamber was observed and recorded, including the crystal habit, 100 initial and final crystallization temperatures, and growth rates. The vapor bubble-disappearing 101 temperature, measured for the vapor bubble generated during cooling process in the sample 102 chamber, was used to approximate the bulk density of the aqueous solution based on the equation 103 of the state of  $H_2O$  (Wagner and Pruß 2002) due to the low solubility of  $Li_2CO_3$  in water 104 (Anderson et al. 2001). The pressure in the sample chamber at any temperature during the 105 experiment could subsequently be obtained along the isochore of this bulk density, assuming a 106 constant volume of the sample chamber during the experiments (Shen et al. 1992; Bassett et al. 107 1993; Chou 2003; Schmidt and Chou 2012; Li et al. 2016). According to the data presented by 108 Schmidt (2014), the true pressure in the alkali carbonate aqueous solution in the sample chamber 109 is higher than that calculated in the way described above. Consequently, the pressure values 110 obtained in this study can be considered as minima at certain temperatures. In addition, 111 comparative experiments were performed using the same initial samples under different pressure 112 conditions in the SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O system. A proper amount of water was drained from the 113 chamber at room temperature without unloading the sample by loosening the driver screws upon 114 the completion of one crystallization experiment. This decreased the bulk water density in the 115 sample chamber, such that the pressure within the sample chamber was lowered at any specified 116 temperature during the reheating process along the new isochore. The aforementioned 117 experimental procedures were then repeated to investigate the effect of pressure on cassiterite 118 crystallization.

119 For the SnO<sub>2</sub>-Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O experiments, a sintered SnO<sub>2</sub> lump, Na<sub>2</sub>CO<sub>3</sub> powder, and 120 distilled water were loaded into the HDAC sample chamber at room temperature, and the same 121 experimental procedures outlined for the SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O experiments were employed. 122 Because there are no equations of state for Na<sub>2</sub>CO<sub>3</sub>–H<sub>2</sub>O, the pressure within the HDAC sample 123 chamber cannot be obtained during the SnO<sub>2</sub>-Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O experiments. Moreover, to determine the dissolution mechanism of SnO<sub>2</sub> in the alkali carbonate aqueous solution, the 124 125 experiments in the SnO<sub>2</sub>-H<sub>2</sub>O and SnO<sub>2</sub>-NaOH-H<sub>2</sub>O systems were conducted using similar 126 experimental procedures as those described for the SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O system.

127 To calculate the initial mass of SnO<sub>2</sub> and estimate the growth rates, the length and width 128 of the cassiterite crystal were measured in the crystallization experiments through observation 129 under an Olympus microscope, with an error of  $\pm 1 \mu m$  via repeated measurements. 130 Subsequently, the volume of cassiterite was calculated, assuming that it had a square-column 131 shape with equal width and thickness. The crystal mass of cassiterite was approximated based on the density (6.95 g/cm<sup>3</sup>) of cassiterite. Similarly, the initial mass of Li<sub>2</sub>CO<sub>3</sub> was approximated 132 133 based on the product of the density  $(2.11 \text{ g/cm}^3)$  and the volume of compacted Li<sub>2</sub>CO<sub>3</sub> powder 134 (approximated as spheres), which was calculated by measuring the diameters under the 135 microscope. The mass of H<sub>2</sub>O loaded into the HDAC sample chamber was obtained, which was 136 based on the H<sub>2</sub>O density calculated on the homogenization temperature of the aqueous solution in the chamber and the chamber volume that subtracted the volume of SnO<sub>2</sub> and Li<sub>2</sub>CO<sub>3</sub>. 137

138 To determine the species of dissolved  $SnO_2$  in the alkali carbonate aqueous solution, in 139 situ Raman analysis was performed in the SnO<sub>2</sub>-Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O, SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O experiments 140 and in the comparative experiments of Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O, SnO<sub>2</sub>-NaOH-H<sub>2</sub>O systems. Each Raman 141 spectrum was acquired using a JY Horiba LabLam HR Raman system with a long working-142 distance Olympus 50× SLMP objective (numeral aperture 0.35). Analyses were performed using 143 a 532-nm laser for excitation (~40 mW at the sample), a grating of 600 lines/mm, and a confocal 144 pinhole diameter of 100 µm. All Raman spectra were recorded with two accumulations of 60 s 145 each.

146

#### 147 EXPERIMENTAL RESULTS

### 148 SnO<sub>2</sub> dissolution and crystallization

We conducted 21 experiments in the SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O system and 11 experiments in the SnO<sub>2</sub>-Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O system (Tables 1 and 2). Five groups of comparative experiments in the SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O system were conducted (Nos. 1, 7, 8, 9, and 13 in Table 1), with the experiments in each group conducted using different bulk water densities for the same samples. The amounts of each component in these experiments are presented in Table 1. The unidirectional and volumetric growth rates of cassiterite were estimated for the experiments in which only one main crystal nucleated and grew (Tables 1 and 2).

In the  $SnO_2-Li_2CO_3-H_2O$  experiments,  $Li_2CO_3$  melted first, forming melt droplets in the sample chamber during the heating process (Fig. 1b). Subsequently,  $SnO_2$  dissolved in the aqueous solution (Figs. 1c and 1d). During the following cooling process, cassiterite nucleated and grew into long prismatic crystals in the 654–856 °C and 660–864 MPa temperature and pressure ranges, respectively (Fig. 1e); crystal growth stopped at 560–768 °C and 246–764 MPa

161 (Figs. 1e–1h). The Li<sub>2</sub>CO<sub>3</sub> melt droplets appeared and increased in size during the cassiterite 162 crystallization, until rapid zabuyelite crystallization at 515-614 °C in tabular shape and 163 occasionally spherulite (Figs. 1f-1i). In the SnO<sub>2</sub>-Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O experiments, SnO<sub>2</sub> dissolved in 164 the  $Na_2CO_3$  aqueous solution during heating. Long prismatic cassiterite crystals began to grow at 165 476–753 °C during the cooling process, and crystal growth ended at 390–619 °C (Fig. 2). Raman 166 spectroscopic analysis revealed that the long prismatic crystals (Figs. 1h and 2c) were cassiterite 167 (Fig. 3).

168 By plotting the crystal volume versus time, it was determined that the growth progress of 169 cassiterite in the SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O and SnO<sub>2</sub>-Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O systems is divided into three 170 segments: initial stage of unobstructed growth, steady growth, and slow growth in the late stage 171 (Fig. 4). In the SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O and SnO<sub>2</sub>-Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O systems, cassiterite initially grew quickly with average rates of 111.23 and 80.09  $\mu$ m<sup>3</sup>/s, respectively (Interval A in Fig. 4). 172 173 Subsequently, the plotted volumes versus time followed linear trends in the two systems (Interval 174 B in Fig. 4), which indicated constant volumetric growth rates of 15.77 and 17.31  $\mu$ m<sup>3</sup>/s, 175 respectively. Finally, the volume growth rates slowed down by one order of magnitude to 3.01 176 and 1.88  $\mu$ m<sup>3</sup>/s, respectively (Interval C in Fig. 4). Notably, the growth of larger crystals was 177 frequently accompanied by the dissolution of smaller crystals during the cassiterite crystallization process, indicating the occurrence of Ostwald ripening (Fig. 5) (Thomas and 178 179 Davidson 2016).

180 The SnO<sub>2</sub>-H<sub>2</sub>O experiments showed that the solubility of SnO<sub>2</sub> was very limited in 181 water; thus, cassiterite crystal was not formed during the experiment. In contrast, SnO<sub>2</sub> exhibited 182 significant solubility in the SnO<sub>2</sub>-NaOH-H<sub>2</sub>O system when heated to 300-600 °C. However, 183 cassiterite did not crystallize during the subsequent cooling process.

### 184 In situ Raman analysis

185 In the SnO<sub>2</sub>–Na<sub>2</sub>CO<sub>3</sub>–H<sub>2</sub>O experiments (Fig. 6), the Raman spectrum of the aqueous solution exhibited an intense  $CO_3^{2-}$  peak at 1065 cm<sup>-1</sup> and a weak  $HCO_3^{-}$  peak at 1017 cm<sup>-1</sup> at 186 24 °C (Hurai et al. 2015), and the  $CO_3^{2-}$  peak shifted to lower wavenumbers as the temperature 187 increased, similar to the Raman spectroscopy features for a Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O system (Rudolph et al. 188 2008; Schmidt 2014). A peak at ~971  $\text{cm}^{-1}$  emerged in the Raman spectrum of the aqueous 189 solution at 700 °C, shifting to lower wavenumbers with increasing temperature; a weak peak at 190 332 cm<sup>-1</sup> appeared at 850 °C. Raman peaks at 332, 919, 971, 1017 (HCO<sub>3</sub><sup>-</sup>) and 1065 cm<sup>-1</sup> 191  $(CO_3^{2^-})$  were detected in the aqueous solution after cooling to 24 °C as well as in the Na<sub>2</sub>CO<sub>3</sub>-192 H<sub>2</sub>O experiment (Fig. 6). The Raman peaks at 332, 919, and 971 cm<sup>-1</sup> can be attributed to the 193 194 tetrahedral ReO<sub>4</sub><sup>-</sup> that formed during the reaction of alkali carbonate aqueous solution with the 195 rhenium gasket at high temperature (Eysel and Kanellakopulos 1993; Chellappa et al. 2009; 196 Foustoukos and Mysen 2015). The addition of a certain amount of rhenium did not appear to 197 interfere with the crystallization kinetics of cassiterite in the alkali carbonate aqueous solutions. In the SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O experiments (Fig. 7), the Raman peaks of ReO<sub>4</sub><sup>-</sup> at 332, 919, 198

and 971 cm<sup>-1</sup> were clearly detected in the aqueous solution at 850 °C, in addition to the  $CO_3^{2-}$ peak at ~1065 cm<sup>-1</sup> while a significant dissolution of Li<sub>2</sub>CO<sub>3</sub> into water occurred at 600 °C. In addition to a weak peak at ~971 cm<sup>-1</sup>,  $CO_3^{2-}$  peaks at ~1072 cm<sup>-1</sup> (v<sub>1</sub>, the symmetric stretching vibration) and ~702 cm<sup>-1</sup> (v<sub>4</sub>, the in-plane symmetric vibration) were observed in the Raman spectra of the Li<sub>2</sub>CO<sub>3</sub> melt droplet (Bates et al. 1972; Carper et al. 2012; Foustoukos and Mysen 2015).

In the  $SnO_2$ -NaOH-H<sub>2</sub>O experiments (Fig. 8), a Raman peak at ~552 cm<sup>-1</sup> was detected at 300 °C after  $SnO_2$  dissolution into the aqueous solution. During further heating, this peak slightly shifted to lower wavenumbers, and the intensity significantly decreased. At 800 °C, the ReO<sub>4</sub><sup>-</sup> Raman peaks at 332, 919, and 971cm<sup>-1</sup> were detected. All of the aforementioned peaks remained in the Raman spectrum after cooling to 24 °C (Fig. 8).

210

### 211 **DISCUSSION**

### 212 Tin speciation dissolved in alkali carbonate-rich aqueous solution

SnO<sub>2</sub> is soluble in the SnO<sub>2</sub>–Li<sub>2</sub>CO<sub>3</sub>–H<sub>2</sub>O and SnO<sub>2</sub>–Na<sub>2</sub>CO<sub>3</sub>–H<sub>2</sub>O systems, and cassiterite crystallization occurred during cooling (Figs. 1 and 2). This result differed from the low solubility in water, as indicated in the experiments for the SnO<sub>2</sub>–H<sub>2</sub>O system. These results indicated that the presence of alkali carbonate can enhance the solubility of SnO<sub>2</sub> in an aqueous solution.

The hydrolysis of alkali carbonate in aqueous solutions can produce an alkaline condition, as indicated by the presence of  $HCO_3^-$  in the  $SnO_2-Na_2CO_3-H_2O$  system (Fig. 6) (Rudolph et al. 2008):

221 
$$CO_3^{2-} + H_2O = HCO_3^{-} + OH^{-}(1)$$

222 Under high pressure and temperature, the dissolved Sn(IV), as an amphoteric metal element,

should be present in an alkaline aqueous solution as  $Sn(OH)_6^{2-}$  (Jackson and Helgeson 1985;

- 224 Lothenbach 1999; Rai et al. 2011):
- 225  $\operatorname{SnO}_2 + 2\operatorname{OH}^- + 2\operatorname{H}_2\operatorname{O} = \operatorname{Sn}(\operatorname{OH})_6^{2^-}(2).$

However, in the SnO<sub>2</sub>-Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O and SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O experiments, the Raman signal

- from dissolved Sn(IV) was not detected, and only Raman peaks of  $CO_3^{2-}$  (1065 cm<sup>-1</sup>),  $HCO_3^{-1}$
- 228  $(1017 \text{ cm}^{-1})$ , and ReO<sub>4</sub><sup>-</sup> (332, 919, and 971 cm<sup>-1</sup>) were present (Figs. 6 and 7). On the other
- hand, in the SnO<sub>2</sub>–NaOH–H<sub>2</sub>O experiments, when SnO<sub>2</sub> was dissolved in the alkaline aqueous

solution by the reaction of Eq. 2, a Raman peak at 552 cm<sup>-1</sup> was detected, which was assigned to 230  $v_1$ [Sn(IV)–OH] of Sn(OH)<sub>6</sub><sup>2-</sup>, as reported by Taylor and Coddington (1992). The intensity of the 231 peak at 552 cm<sup>-1</sup> significantly decreased with an increase in temperature (Fig. 8), which can be 232 caused by the Sn(IV)-OH bonds in  $Sn(OH)_6^{2-}$  becoming less polarizable at high temperatures 233 (Schmidt 2014). The dissolution temperature of SnO<sub>2</sub> in the SnO<sub>2</sub>-NaOH-H<sub>2</sub>O system was much 234 235 lower than that in the SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O and SnO<sub>2</sub>-Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O systems, which indicated the higher solubility of SnO<sub>2</sub> in the SnO<sub>2</sub>–NaOH–H<sub>2</sub>O system. Thus, the weak polarization at high 236 237 temperature and lower solubility of SnO<sub>2</sub> may not allow to detect the Raman spectrum of  $Sn(OH)_6^{2-}$  at ~552 cm<sup>-1</sup> in the SnO<sub>2</sub>-Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O and SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O experiments, 238 wherein  $SnO_2$  dissolved into the alkaline aqueous solution as  $Sn(OH)_6^{2-}$  with low concentrations 239 240 at temperatures higher than 700 °C (Figs. 6 and 7). It could also be the reason that the Raman 241 peak of HCO<sub>3</sub><sup>-</sup> was not detected at high temperature in SnO<sub>2</sub>-Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O and SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-242 H<sub>2</sub>O experiments, as suggested from the Raman data by Schmidt (2014). Therefore, we inferred that Sn(IV) dissolved in alkali carbonate aqueous solutions as  $Sn(OH)_6^{2-}$  under high pressure and 243 temperature, rather than as a complex with carbonate such as Sn(IV) hydroxycarbonate 244 245 (Kuril'chikova and Barsukov 1971; Kosals 1976). According to Eqs. 1 and 2, the dissolution of  $SnO_2$  in an alkali carbonate aqueous solution can be expressed as the following reaction: 246

247 
$$\operatorname{SnO}_2 + 2\operatorname{CO}_3^{2-} + 4\operatorname{H}_2O = \operatorname{SnOH}_6^{2-} + 2\operatorname{HCO}_3^{-}(3)$$

248 The roles of alkali carbonate in the solubility of SnO<sub>2</sub> can be attributed to the alkaline nature of carbonate aqueous solutions. The back reaction of Eq. 3 to cassiterite +  $CO_3^{2-}$  + H<sub>2</sub>O occurred 249 250 upon cooling. In contrast, the dissolution reaction of SnO<sub>2</sub> in the strong alkaline SnO<sub>2</sub>-NaOH-251 H<sub>2</sub>O system (Eq. 2) was not reversible and did not result in cassiterite crystallization during the 252 cooling process. This indicates that strong alkaline conditions are unfavorable for the 253 crystallization of cassiterite, while the alkali carbonate aqueous solution provides suitable 254 alkaline conditions for cassiterite crystallization.

255 The moles fraction of cassiterite crystallized from the SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O system ranges 256 from 0.03 to 0.41 mol%, which mirrors the solubility of  $SnO_2$  in aqueous solution (Table 1). 257 Under certain ranges of temperatures and pressures (e.g., at approximately 700-800 °C and 700-258 800 MPa), the mole fraction of cassiterite crystallized from the aqueous solution is linearly 259 related to the mole fraction of Li<sub>2</sub>CO<sub>3</sub> dissolved in aqueous solution (Fig. 9). This suggests that the solubility of SnO<sub>2</sub> depends on the content of Li<sub>2</sub>CO<sub>3</sub> in the aqueous solution, further 260 261 indicating the effect of alkalinity, produced by the hydrolysis of alkali carbonate, on the 262 solubility of SnO<sub>2</sub> in the hydrothermal fluids.

263

#### 264 **Cassiterite crystallization conditions and behaviors**

265 The cassiterite crystallization occurred at 560–856 °C and approximately 246–864 MPa 266 in the SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O system and primarily ended at rare metal pegmatite-forming P-T 267 conditions (Table 1 and Fig. 10a). This temperature condition is higher than that (390–753 °C) in 268 the SnO<sub>2</sub>-Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O system (Table 2). It can be because of the low solubility of Li<sub>2</sub>CO<sub>3</sub> that 269 led to the relatively low solubility of  $SnO_2$  in the aqueous solution. Therefore, if more soluble 270 alkali carbonates (e.g., Na<sub>2</sub>CO<sub>3</sub>) dissolve in the aqueous solution, the crystallization P-T 271 conditions of cassiterite will be lower and much closer to the rare metal pegmatite-forming P-T conditions (Fig. 10a). This finding is consistent with the enrichment of  $Li^+$ ,  $Na^+$ ,  $K^+$ ,  $CO_2$ ,  $CO_3^{2-}$ , 272 273 and  $HCO_3^{-}$  in the melt and fluid inclusions of many tin-mineralized pegmatites (Rickers et al. 274 2006; Thomas et al. 2006a, 2012; Thomas and Davidson 2016). In some cases, the 275 concentrations of carbonates and bicarbonates can achieve as high as 20-40 wt% (Thomas et al.

2006b). Therefore, the P–T conditions of cassiterite crystallization in alkali carbonate-rich
environment can reflect the cassiterite crystallization conditions in the rare metal granitic
pegmatite.

In each of the five groups of comparison experiments in the  $SnO_2-Li_2CO_3-H_2O$  system, both the initiation and ending temperatures of cassiterite crystallization increased with pressure decreasing (Table 1 and Fig. 10b). This negative P–T trend of cassiterite crystallization can reflect the effect of pressure on  $SnO_2$  solubility, which was in contrast to the data presented by Wilson and Eugster (1990) and Schmidt (2018) that showed that the effect of pressure on  $SnO_2$ solubility in H<sub>2</sub>O+HCl is insignificant.

Growth rates in the order of 1 cm per 1000-10,000 years have been reported for K-285 feldspar crystals growing in granite (Pitcher 2012). Experimental studies on the kinetics of 286 287 crystallization of hydrous granitic melts reported that the crystal growth rates were in the order of  $10^{-6}$ - $10^{-8}$  cm/s (Fenn 1977; Swanson and Fenn 1992; Baker and Freda 2001; Li et al. 2013; 288 289 Sirbescu et al. 2017; Maneta and Anderson 2018). These growth rates are equivalent to or lower than the average length growth rates  $(0.61 \times 10^{-6} - 8.22 \times 10^{-6} \text{ cm/s})$  of cassiterite in alkali 290 291 carbonate aqueous solutions (Tables 1 and 2). The cassiterite volume versus time shows 292 unobstructed and steady linear trends (Intervals A and B in Fig. 4). In this study, the average volume growth rates of cassiterite range from 3.40 to 19.07  $\mu$ m<sup>3</sup>/s (Tables 1 and 2), which 293 confirms to the values  $(10^{-5}-10^2 \,\mu\text{m}^3/\text{s})$  measured for the undercooled pegmatitic melt (Sirbescu 294 et al. 2017). The high growth rates indicate the low viscosity of alkali carbonate-rich aqueous 295 296 solutions that enhances the diffusion of components and, therefore, speeds up the growth of the 297 crystals (Thomas and Davidson 2016). Obvious Ostwald ripening that occurred in some experiments, also indicates the high diffusion of  $Sn(OH)_6^{2-}$  in the alkali carbonate aqueous 298

solution because Ostwald ripening requires enough compositions to transfer to the cassiterite crystallization front during the dissolution of finer cassiterite crystals (Thomas and Davidson 2016). As a result, the long prismatic cassiterite crystals formed in the sample chamber (Figs. 1h and 2c) because the (110) prismatic surface of cassiterite with the lowest surface energy preferentially grows while high diffusion supplies nutrition to the crystal surface (Smith et al. 1995; Batzill and Diebold 2005).

305

### **306** Roles of carbonate in tin mineralization in granitic pegmatite conditions

307 The HDAC sample chamber loaded with alkali carbonate-rich aqueous solution is similar 308 to the alkali carbonate-rich inclusions that represent the hydrous melt or fluid from which tin-309 mineralized pegmatite crystallized (Thomas et al. 2011; Thomas and Davidson 2013). As 310 discussed above, the formation conditions of cassiterite can fall within the rare metal pegmatite-311 forming P–T conditions if the alkali carbonates in aqueous solution are as rich as they are in 312 many pegmatite-forming melts or fluids. Moreover, in the alkali carbonate-rich tin-mineralized 313 pegmatites, cassiterite occurs in the form of long prismatic crystals that are similar to those 314 grown in the present experiments, such as the cassiterite from the peralkaline zones of the 315 Ehrenfriedersdorf pegmatites in 5-cm-long prismatic shapes (Thomas et al. 2006a). The long 316 prismatic shape of cassiterite indicates the high diffusion rate of Sn in pegmatitic melt or fluid 317 (Smith et al. 1995; Batzill and Diebold 2005), resulting in high growth rates of cassiterite in 318 pegmatite-forming conditions (Chakoumakos and Lumpkin 1990; Morgan and London 1999; 319 Webber et al. 1999; Sirbescu et al. 2009; Li et al. 2013). It is also consistent with the growth rate 320 of cassiterite in the carbonate-rich aqueous solution as mentioned above.

#### This is a preprint, the final version is subject to change, of the American Mineralogist (MSA) Cite as Authors (Year) Title. American Mineralogist, in press. DOI: https://doi.org/10.2138/am-2020-7118

321 The above features indicate that carbonates can play important roles in the formation of 322 tin mineralization in high-fractionated pegmatite enriched in lithium or sodium. The carbonaterich aqueous solution has a high Sn(IV) transport capability in the form of  $Sn(OH)_6^{2-}$  to the 323 324 crystallization front of cassiterite, which helps to form long prismatic cassiterite macrocrystals 325 during the pegmatite-forming process. This observation agrees with the experiments of Schmidt 326 (2018), which showed that tin can be transported in Sn(IV). Moreover, alkali carbonates can 327 lower the crystallization temperature and solidi, increase the crystallization rate (Sirbescu and 328 Nabelek 2003) and decrease the viscosity of a hydrous melt by as much as 2-3 orders of 329 magnitude (Ejima et al. 1987; Thomas et al. 2012; Jones et al. 2013; Thomas and Davidson 2016). These properties are essential for the formation of pegmatite texture (London 2009). 330 331 Therefore, carbonate-rich silicate melts or highly concentrated fluids can serve as media for tin 332 mineralization in granitic pegmatite (Thomas et al. 2011; Li and Chou 2017; Maneta and 333 Anderson 2018).

334

#### 335 IMPLICATIONS

336 The roles of suitable alkaline condition produced by alkali carbonates in tin 337 mineralization can be extended to LCT-type pegmatite because many LCT pegmatites (e.g., 338 Amis, Borborema, Bornholm, Elba, Froland, Königshain, Orlovka, Muiane, and Tanco) have 339 plenty of alkali carbonate-rich inclusions (Thomas et al. 2011, 2012; Li and Chou 2016). Alkali 340 carbonates occur as daughter crystals in crystal-rich inclusions, mainly including nahcolite 341  $[NaHCO_3]$ , zabuyelite  $[Li_2CO_3]$ , potash  $[K_2CO_3]$ , and dawsonite  $[NaAl(CO_3)(OH)_2]$  in addition 342 to the volatile phases (CO<sub>2</sub>) and carbonates/bicarbonates dissolved in the aqueous solution 343 (Thomas et al. 2012). These crystal-rich inclusions contain ~7 mol% Li<sub>2</sub>CO<sub>3</sub> in the Tanco Li-Ta

#### This is a preprint, the final version is subject to change, of the American Mineralogist (MSA) Cite as Authors (Year) Title. American Mineralogist, in press. DOI: https://doi.org/10.2138/am-2020-7118

344 pegmatite (Anderson 2001), 35.0 wt% NaHCO<sub>3</sub> in the Orlovka pegmatite in Russia (Thomas et 345 al. 2011), and 10 vol%  $Li_2CO_3$  in the Jiajika lithium pegmatite in China (Li and Chou 2016). In 346 the crystal-rich inclusions hosted in spodumene in the Jiajika pegmatite, China (Li and Chou 347 2016), alkali carbonate can react with lithium aluminosilicate into soluble silicate or aqueous 348 solution (Thomas et al. 2011; Li and Chou 2017). Similarly, alkaline conditions enhance the 349 solubility of beryllium (Thomas et al. 2009) and tungsten (Li et al. 2018) in the form of ions (e.g.,  $Be^+$  and  $W^{6+}$ ) in a pegmatite melt or a fluid. The alkali carbonate-rich aqueous solution can 350 351 also dissolve more SiO<sub>2</sub> and Al<sub>2</sub>O<sub>3</sub> than pure water under pegmatite-forming P-T conditions (Ye 352 et al. 2008; Schmidt 2014), which are essential components for granitic pegmatite formation 353 (London 2008).

354 Moreover, an alkali carbonate-rich hydrous melt or fluid bears low crystallization 355 temperature and solidi and low viscosity, which helps to form the boundary layer liquid in the 356 "constitutional zone refining" pegmatite-forming model (London 2018). Carbonates promote the 357 occurrence of immiscibility in alkaline-rich granitic magma (Veksler 2004; Linnen et al. 2014) in 358 the pegmatite-forming model of immiscibility (Thomas and Davidson 2016). Consequently, CO<sub>2</sub> 359 or carbonate could be types of fluxes to promote the formation of rare metal pegmatites, and 360 alkali carbonate-rich aqueous solutions or hydrous melts are potential media for the formation of 361 rare metal mineralization.

362

### 363 ACKNOWLEDGMENTS

This study was supported by the National Natural Science Foundation of China (41872096), the Chinese National Non-profit Institute Research Grant of CAGS-IMR

- 366 (JYYWF201814), the National Key R&D Program of China (2016YFC0600208), and the Key
- 367 Frontier Science Program (QYZDY-SSW-DQC008) of Chinese Academy of Sciences.
- 368

#### **369 REFERENCES CITED**

- 370 Anderson, A.J., Clark, A.H., and Gray, S. (2001) The occurrence and origin of zabuyelite
- (Li<sub>2</sub>CO<sub>3</sub>) in spodumene-hosted fluid inclusions: Implications for the internal evolution of
   rare-element granitic pegmatites. Canadian Mineralogist, 39, 1513–1527.
- 373 Audétat, A., Gunther, D., and Heinrich, C.A. (1998) Formation of a magmatic-hydrothermal ore
- deposit: Insights with LA-ICP-MS analysis of fluid inclusions. Science, 279, 2091–2094.
- 375 Baker, D.R., and Freda, C. (2001) Eutectic crystallization in the undercooled orthoclase-quartz-
- H<sub>2</sub>O system: Experiments and simulations. European Journal of Mineralogy, 13, 453–466.
- 377 Bassett, W.A., Shen, A.H., Bucknum, M., and Chou, I-M. (1993) A new diamond anvil cell for
- 378 hydrothermal studies to 2.5 GPa and from -190 to 1200 °C. Review of Scientific
  379 Instruments, 64, 2340–2345.
- 380 Bassett, W.A., Wu, T.C., Chou, I-M., Haselton, H.T. Jr., Frantz, Z., Mysen, B.O., Huang, W.L.,
- 381 Sharma, S.K., and Schiferl, D. (1996) The hydrothermal diamond anvil cell (HDAC) and its
- 382 applications. In M.D. Dyar, C. McCammon, and M.W. Schaefer, Eds., Mineral
- Specptroscopy: A Tribute to Roger G. Burns, p. 261–272. The Geochemical Society, Special
  Publication No. 5.
- 385 Bates, J.B., Brooker, M.H., Quist, A.S., and Boyd, G.E. (1972) Raman spectra of molten alkali
- 386 metal carbonates. The Journal of Physical Chemistry, 76, 1565–1571.
- Batzill, M., and Diebold, U. (2005) The surface and materials science of tin oxide. Progress in
  Surface Science, 79, 47–154.

- 389 Bhalla, P., Holtz, F., Linnen, R.L., and Behrens, H. (2005) Solubility of cassiterite in evolved
- 390 melts: Effect of T,  $fO_2$ , and additional volatiles. Lithos, 80, 387–400.
- 391 Carper, W.R., Wahlbeck, P.G., and Griffiths, T.R. (2012) DFT models of molecular species in
- 392 carbonate molten salts. Journal of Physical Chemistry B, 116, 5559–5567.
- 393 Chakoumakos, B.C., and Lumpkin, G.R. (1990) Pressure-temperature constraints on the
- 394 crystallization of the Harding pegmatite, Taos country, New Mexico. Canadian Mineralogist,
- **395 28**, **287–298**.
- Chellappa, R.S., Somayazulu, M., and Hemley, R.J. (2009) Rhenium reactivity in H<sub>2</sub>O–O<sub>2</sub>
   supercritical mixtures at high pressures. High Pressure Research, 29, 792–799.
- 398 Chou, I-M. (1996) Precautions on the use of the hydrothermal diamond-anvil cell for the
- acquisition of volumetric and phase relation data of geologic fluids. PACROFI VI, Programand Abstracts, p. 31.
- 401 Chou, I-M. (2003) Hydrothermal diamond-anvil cell: application to studies of geologic fluids.
  402 Acta Petrologica Sinica, 19, 213–220.
- 403 Duc-Tin, Q., Audétat, A., and Keppler, H. (2007) Solubility of tin in (Cl, F)-bearing aqueous
- 404 fluids at 700 °C, 140 MPa: A LA-ICP-MS study on synthetic fluid inclusions. Geochimica et
- 405 Cosmochimica Acta, 71, 3323–3335.
- Ejima, T., Sato, Y., Yamamuro, T., Tamai, K., Hasebe, M., Bohn, M.S., and Janz, G.J. (1987)
  Viscosity of the eutectic Li<sub>2</sub>CO<sub>3</sub>–Na<sub>2</sub>CO<sub>3</sub>–K<sub>2</sub>CO<sub>3</sub> melt. Journal of Chemical & Engineering
  Data, 32, 180–182.
- Eysel. H.H., and Kanellakopulos, B. (1993) Raman spectra, absolute Raman intensities and
   electro-optical parameters of pertechnetate, perrhenate and periodate ions in aqueous
- 411 solution. 24, 119–122.

- 412 Fenn, P.M. (1977) The nucleation and growth of alkali feldspars from hydrous melts. Canadian
  413 Mineralogist, 15, 135–161.
- 414 Foustoukos, D.I., and Mysen, O.B. (2015) The structure of water-saturated carbonate melts.
- 415 American mineralogist, 100, 35–46.
- 416 Fu, M., and Kwak, T.A.P. (1994) Geology, geochemistry and fluid inclusions of the Gejiu tin-
- 417 polymetallic field, People's Republic of China. International Geology Review, 36, 272–292.
- 418 Fu, M., Kwak, T.A.P., and Mernagh, T.P. (1993) Fluid inclusions studies of zoning in the
- 419 Dachang tin-polymetallic ore field, the People's Republic of China. Economic Geology, 88,
  420 283–300.
- 421 Heinrich, C.A. (1990) The chemistry of hydrothermal tin (-tungsten) ore deposition. Economic
  422 Geology, 85, 457–481.
- 423 Hurai, V., Huraiová, M., Slobodník, M., and Thomas, R. (2015) Geofluids: Developments in
- 424 Microthermometry, Spectroscopy, Thermodynamics and Stable Isotopes, 504 p. Elsevier.
- 425 Jackson, K.J., and Helgeson, H.C. (1985) Chemical and thermodynamic constraints on the
- 426 hydrothermal transport and deposition of tin: I. Calculation of the solubility of cassiterite at
- 427 high pressures and temperatures. Geochimica et Cosmochimica Acta, 49, 1–22.
- Jones, A.P., Genge, M., and Carmody, L. (2013) Carbonate melts and Carbonatites. Reviews in
  Mineralogy and Geochemistry, 75, 289–322.
- 430 Kamilli, R.J., Kimball, B.E., and Carlin, J.F. Jr. (2017) Chapter S. Tin. In K.J. Schulz, J.H. Jr.
- 431 DeYoung, R.R. II Seal, and D.C. Bradley, Eds., Critical Mineral Resources of The United
- 432 States—Economic and Environmental Geology and Prospects for Future Supply, p. S1–S53.
- 433 U.S. Geological Survey Professional Paper 1802, Reston, VA.

- 434 Kokh, M.A., Akinfive, N.N., Pokrovski, G.S., Salvi, S., and Guillaume, D. (2017) The role of
- 435 carbon dioxide in the transport and fractionation of metals by geological fluids. Geochimica

436 et Cosmochimica Acta, 197, 433–466.

- 437 Korges, M., Weis, P., Lüders, V., and Laurent, O. (2018) Depressurization and boiling of a single
- 438 magmatic fluid as a mechanism for tin-tungsten deposit formation. Geology 46, 75–78.
- 439 Kosals, J.A. (1976) Main Features of Geochemistry of Rare Metals in Granitic Melts and
- 440 Solutions (Fluids), 232 p. Nauka, Novosibirsk (in Russian).
- 441 Kuril'chikova, G.Y., and Barsukov, V.L. (1971) Effects of CO<sub>2</sub> and of sodium and potassium
- 442 bicarbonates and carbonates on the formation of Sn (IV) complexes in solution.
- 443 Geochemical International, 8, 395–404.
- Li, T.J. (1989) Experimental studies of the solubility of cassiterite and the extraction of tin from
  granitic melts. Chinese Journal of Geochemistry, 8, 84–96.
- Li, J.K., and Chou, I-M. (2016) An occurrence of metastable cristobalite in spodumene-hosted
  crystal-rich inclusions from Jiajika pegmatite deposit, China. Journal of Geochemical
  Exploration, 2016, 29–36.
- Li, J.K., and Chou, I-M. (2017) Homogenization experiments of crystal-rich inclusions in
  spodumene from Jiajika lithium deposit, china, under elevated external pressures in a
  hydrothermal diamond-anvil cell. Geofluids, 2017, 1–12.
- 452 Li, J.K., Chou, I-M., Yuan, S.D., and Burruss, R.C. (2013) Observations on the crystallization of
- 453 spodumene from aqueous solutions in a hydrothermal diamond-anvil cell. Geofluids, 13,
  454 467–474.
- Li, J.K., Bassett, W.A., Chou, I-M., Ding, X., Li, S.H., and Wang, X.Y. (2016) An improved
- 456 hydrothermal diamond anvil cell. Review of Scientific Instruments, 87, 1513–1527.

- 457 Li, J.K, Liu, Y.C., Zhao, Z., and Chou, I-M. (2018) Roles of carbonate/CO<sub>2</sub> in the formation of
- 458 quartz-vein wolframite deposits: Insight from the crystallization experiments of huebnerite
- 459 in alkali-carbonate aqueous solutions in a hydrothermal diamond-anvil cell. Ore Geology
- 460 Reviews. 95, 40–48.
- 461 Linnen, R.L. (1998) Depth of emplacement, fluid provenance and metallogeny in granitic
- 462 terrains: A comparison of western Thailand with other Sn-W belts. Mineralium Deposita, 33,
- 463 461–476.
- Linnen, R.L., Pichavant, M., Holtz, F., and Burgess, S. (1995) The effect of fO<sub>2</sub> on the solubility,
- 465 diffusion, and speciation of tin in haplogranitic melt at 850 °C and 2 kbar. Geochimica et
- 466 Cosmochimica Acta, 59, 1579–1588.
- 467 Linnen, R.L., Samson, I.M., Williams-Jones, A.E., and Chakhmouradian, A.R. (2014)
- 468 Geochemistry of the rare-earth element, Nb, Ta, Hf, and Zr Deposits. In K.K. Turekian, and
- 469 H.D. Holland, Eds., Treatise on Geochemistry (Second Edition), p. 543–568. Elsevier,
  470 Oxford.
- 471 London, D. (2008) Pegmatites, 368 p. Canadian Mineralogist Special Publication, 10.
- 472 London, D. (2009) The origin of primary textures in granitic pegmatites. Canadian Mineralogist,
  473 47, 697–724.
- 474 London, D. (2018) Ore-forming processes within granitic pegmatites. Ore Geology Reviews,
  475 101, 349–383.
- 476 Lothenbach, B., Ochs, M., Wanner, H., and Yui, M. (1999) Thermodynamic date for the
- 477 speciation and solubility of Pd, Pb, Sn, Sb, Nb, and Bi in aqueous solution. 340 p. JNC
- 478 Report TN8400 99-011, Japan.

- 479 Maneta, V., and Anderson, A.J. (2018) Monitoring the crystallization of water-saturated granitic
- 480 melts in real time using the hydrothermal diamond anvil cell. Contributions to Mineralogy481 and Petrology, 173, 83.
- 482 Morgan, G.B., and London, D. (1999) Crystallization of the Little Three layered pegmatite-aplite
- dike, Ramona District, California. Contributions to Mineralogy and Petrology, 136, 310–
  330.
- 485 Müller, B., and Seward, T.M. (2001) Spectrophotometric determination of the stability of tin (II)
- 486 chloride complexes in aqueous solution up to 300 °C. Geochimica et Cosmochimica Acta,
- 487 65, 4187–4199.
- 488 Naumov, V.B., Dorofeev, V.A., and Mironova, O.F. (2011) Physicochemical parameters of the
- formation of hydrothermal deposits: A fluid inclusion study. I. Tin and tungsten deposits.
  Geochemistry International, 49, 1063–1082.
- 491 Phillips, G.N., and Evans, K.A. (2004) Roles of CO<sub>2</sub> in the formation of gold deposits. Nature
  492 429, 860–863.
- 493 Pitcher, W.S. (2012) The Nature and Origin of Granite, 2nd ed., 387 p. Springer Science &
  494 Business Media, Dordrecht, Netherlands.
- 495 Rai, D., Yui, M., Todd Schaef, H., and Kitamura, A. (2011) Thermodynamic model for SnO<sub>2</sub>(cr)
- 496 and  $SnO_2(am)$  solubility in the aqueous  $Na^+-H^+-OH^--Cl^--H_2O$  system. Journal of Solution
- 497 Chemistry, 40, 1155–1172.
- 498 Rickers, K., Thomas, R., and Heinrich, W. (2006) The behavior of trace elements during the
- 499 chemical evolution of the H<sub>2</sub>O-, B-, and F-rich granite–pegmatite–hydrothermal system at
- 500 Ehrenfriedersdorf, Germany: A SXRF study of melt and fluid inclusions. Mineralium
- 501 Deposita, 41, 229–245.

502	Rudolph, W.W., Irmer, G., and Königsberger, E. (2008) Speciation studies in aqueous HCO <sub>3</sub> <sup>-</sup> -									
503	$\mathrm{CO_3}^{2-}$ solutions. A combined Raman spectroscopic and thermodynamic study. Dalton									
504	Transactions, 2008, 900–908.									
505	Schmidt, C. (2014) Raman spectroscopic determination of carbon speciation and quartz									
506	solubility in H <sub>2</sub> O+Na <sub>2</sub> CO <sub>3</sub> and H <sub>2</sub> O+NaHCO <sub>3</sub> fluids to 600 °C and 1.53 GPa. Geochimica et									
507	Cosmochimica Acta, 145, 281–296.									
508	Schmidt, C. (2018) Formation of hydrothermal tin deposits: Raman spectroscopic evidence for									
509	an important role of aqueous Sn (IV) species. Geochimica et Cosmochimica Acta, 220, 499-									
510	511.									
511	Schmidt, C., and Chou, I-M. (2012) The hydrothermal diamond anvil cell (HDAC) for Raman									
512	spectroscopic studies of geological fluids at high pressures and temperatures. In J. Dubessy,									
513	M.C. Caumon, and F. Rull, Eds., Raman Spectroscopy Applied to Earth Science and									
514	Cultural Heritage, 12, 247–276. EMU Notes in mineralogy, Aberystwyth, UK.									
515	Shaposhnikov, V.V., and Aranovich, L.Y. (2015) Experimental study of model granite melting in									
516	the presence of alkali carbonate solutions at 400 MPa. Geochemistry International, 53, 838-									
517	844.									
518	Shen, A.D., Bassett, W.A., and Chou, I-M. (1992) Hydrothermal studies in a diamond anvil cell:									
519	pressure determination using the equation of state. In Y. Syono, and M.H. Manghnani, Eds.,									
520	High-pressure Research: Application to Earth and Planetary Sciences, p. 61-68. American									
521	Geophysical Union, Washington.									
522	Sherman, D.M., Ragnarsdottir, K.V., Oelkers, E.H., and Collins, C.R. (2000) Speciation of tin									
523	$(Sn^{2+} and Sn^{4+})$ in aqueous Cl solutions from 25 °C to 350 °C: An in situ EXAFS study.									

Chemical Geology, 167, 169–176. 524

#### This is a preprint, the final version is subject to change, of the American Mineralogist (MSA) Cite as Authors (Year) Title. American Mineralogist, in press. DOI: https://doi.org/10.2138/am-2020-7118

- 525 Sirbescu, M.C., and Nabelek, P.I. (2003) Crustal melts below 400 °C. Geology, 31, 685–688.
- 526 Sirbescu, M.C., Wilke, M., and Veksler, I.V. (2009) Understanding pegmatite texture: Kinetics of
- 527 crystallization in the haplogranite-Li-B-H<sub>2</sub>O system. American Geophysical Union, 90,
- 528 V43B–2233.
- 529 Sirbescu, M.C., Schmidt, C., Veksler, I.V., Whittington, A.G., and Wilke, M. (2017)
- 530 Experimental crystallization of undercooled felsic liquids: Generation of pegmatitic texture.
- 531 Journal of Petrology, 58, 539–568.
- 532 Smith, A., Laurent, J.M., Smith, D.S., Bonnet, J.P., and Clemente, R.R. (1995) Structural and
- 533 electrical studies on highly conducting spray deposited fluorine and antimony doped SnO<sub>2</sub>
- thin films from  $SnCl_2$  precursor. Thin Solid Films, 266, 20–30.
- 535 Štemprok, M. (1990) Solubility of tin, tungsten and molybdenum oxides in felsic magmas.
  536 Mineralium Deposita, 25, 205–212.
- 537 Swanson, S.E., and Fenn, P.M. (1992) The effect of F and Cl on the kinetics of albite 538 crystallization: A model for granitic pegmatites? Canadian Mineralogist, 30, 549–559.
- 539 Taylor, J.R., and Coddington, J.M. (1992) The constitution of aqueous tin(IV) chloride and
- 540 bromine solutions and solvent extracts studied by <sup>119</sup>Sn NMR and vibrational spectroscopy.
- 541 Polyhedron, 11, 1531–1544.
- Taylor, J.R., and Wall, V.J. (1993) Cassiterite solubility, tin speciation, and transport in a
  magmatic aqueous phase. Economic Geology, 88, 437–460.
- Thomas, R., and Davidson, P. (2012) Water in granite and pegmatite-forming melts. Ore Geology
  Reviews, 46, 32–46.
- 546 Thomas, R., and Davidson, P. (2013) The missing link between granites and granitic pegmatites.
- 547 Journal of Geosciences, 58, 183–200.

#### This is a preprint, the final version is subject to change, of the American Mineralogist (MSA) Cite as Authors (Year) Title. American Mineralogist, in press. DOI: https://doi.org/10.2138/am-2020-7118

548	Thomas, R., and Davidson, P. (2016) Revisiting complete miscibility between silicate melts and
549	hydrous fluids, and the extreme enrichment of some elements in the supercritical state-
550	Consequences for the formation of pegmatites and ore deposits. Ore Geology Reviews, 72,
551	1088–1101.
552	Thomas, R., Webster, J.D., Rhede, D., Seifert, W., Rickers, K., Förster, H.J., Heinrich, W., and
553	Davidson, P. (2006a) The transition from peraluminous to peralkaline granitic melts:

554 Evidence from melt inclusions and accessory minerals. Lithos, 91, 137–149.

- Thomas, R., Webster, J.D., and Davidson, P. (2006b) Understanding pegmatite formation: The
  melt and fluid inclusion approach. Mineralogical Association Canada Short Course, 36,
- 557 189–210.
- Thomas, R., Davidson, P., and Badanina, E. (2009) A melt and fluid inclusion assemblage in
  beryl from pegmatite in the Orlovka amazonite granite, East Transbaikalia, Russia:
  implications for pegmatite-forming melt systems. Mineralogy and Petrology, 96, 129–140.
- 561 Thomas, R., Davidson, P., and Schmidt, C. (2011) Extreme alkali bicarbonate- and carbonate-
- rich fluid inclusions in granite pegmatite from the Precambrian Rønne granite, Bornholm

563 Island, Denmark. Contributions to Mineralogy and Petrology, 161, 315–329.

- 564 Thomas, R., Davidson, P., and Beurlen, H. (2012) The competing models for the origin and
- 565 internal evolution of granitic pegmatites in the light of melt and fluid inclusion research.
- 566 Mineralogy and Petrology, 106, 55–73.
- 567 Veksler, I.V. (2004) Liquid immiscibility and its role at the magmatic-hydrothermal transition: a
- summary of experimental studies. Chemical Geology, 210, 7–31.

- 569 Wagner, W., and Pru $\beta$ , A. (2002) The IAPWS formulation 1995 for the thermodynamic
- properties of ordinary water substance for general and scientific use. Journal of Physical and
  Chemical Reference Data, 31, 387–535.
- 572 Webber, K.L., Simmons, W.B., Falster, A.U., and Foord, E.E. (1999) Cooling rates and
- 573 crystallization dynamics of shallow level pegmatite-aplite dikes, San Diego County,
- 574 California. American Mineralogist, 84, 708–717.
- Wilson, G.A., and Eugster, H.P. (1990) Cassiterite solubility and tin speciation in supercritical
  chloride solutions. Geochemistry Society Special Publication, 2, 179–195.
- 577 Xiong, X., Li, J.K., Wang, D.H., Li, S.P., and Lin, H. (2019) Fluid characteristics and evolution
- of the Zhawulong granitic pegmatite lithium deposit in the Ganzi-Songpan region,
  Southwestern China. Acta Geologica Sinica (English Edition), 93, 943–954.
- Ye, Y.P., Zeng, X.Q., Qian, W.L., and Wang, M.W. (2008) Synthesis of pure zeolites from
  supersaturated silicon and aluminum alkali extracts from fused coal fly ash. Fuel, 87, 1880–
- 582 1886.
- 583

## 584 **Figure captions**

585

586 Figure 1. Images from experiment No. 7-1 (Table 1) taken at various temperatures. The pressures 587 marked on the lower right corners represent minimum values in the sample chamber at 588 corresponding temperatures, because the true pressure is higher than that calculated based on the 589 equation of state of  $H_2O$  (Schmidt 2014). (a) The sample chamber was loaded with  $SnO_{2}$ , 590  $Li_2CO_3$ , distilled water, and vapor bubbles at room temperature (24 °C). (b)  $Li_2CO_3$  melted in the 591 water with increasing temperature, forming melt droplets at 639 °C. (c) SnO<sub>2</sub> was dissolving in 592 aqueous solution at 724  $^{\circ}$ C. (d) SnO<sub>2</sub> has almost completely dissolved in the aqueous solution at 593 766 °C. (e) Cassiterite nucleated and grew on the wall of the sample chamber during 1 °C/min 594 cooling, at 741 °C. (f) Li<sub>2</sub>CO<sub>3</sub> exsolved from aqueous solution and formed melt droplets during 595 the cassiterite crystallization process at 727 °C. (g) The Li<sub>2</sub>CO<sub>3</sub> melt droplets became larger 596 during further cooling. (h) Cassiterite growth ended at 660 °C. (i) Zabuyelite crystallized at 525 597 °C.

598

Figure 2. Images from experiment No. 11 (Table 2) taken at various temperatures. (a) Cassiterite nucleation and growth during 1 °C/min cooling, at 753 °C. (b) Cassiterite growth as long prismatic crystals. (c) Cessation of cassiterite growth at 619 °C.

602

Figure 3. Raman spectra of cassiterite crystallized in the systems of SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O (a),
SnO<sub>2</sub>-Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O (b), and from the RRUFF database (c, RRUFF ID: R040017).

605

606 Figure 4. Representative volumetric growth rate diagram for cassiterite in the SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O 607 and SnO<sub>2</sub>-Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O systems. Note, that the time was set to 0 s when the first measurement 608 was performed once the crystal appeared. The solid lines represent regression fit curves. For both 609  $SnO_2-Li_2CO_3-H_2O$  and  $SnO_2-Na_2CO_3-H_2O$  systems, the volumetric growth rates of cassiterite 610 changed in three intervals. Interval A: unobstructed growth with the average rates of 111.23 and 611  $80.09 \text{ }\mu\text{m}^3/\text{s}$ , respectively. Interval B: steady growth with the rates of 15.77 and 17.31  $\mu\text{m}^3/\text{s}$ . 612 respectively, which was obtained from the slopes of the linear regression fit curves. Interval C: slow growth with the average rates of 3.01 and 1.88  $\mu$ m<sup>3</sup>/s, respectively. 613

614

Figure 5. Ostwald ripening phenomenon observed in experiment No. 3 (Table 1). (a) Formation
of different-sized cassiterite crystals in the sample chamber during 1 °C/min cooling, at 796 °C.
(b) Gradual growth of the larger crystals and dissolution of the smaller crystals at 792 °C. (c)
Complete dissolution of the smaller crystals and continued growth of the larger crystals at 775
°C.

620

Figure 6. Raman spectra of the aqueous solution measured at different temperatures in the  $SnO_{2}$ -Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O and Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O experiments. The spectra colored in red and blue were obtained in the  $SnO_{2}$ -Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O system during the heating and cooling processes, respectively. The spectrum colored in black was obtained in the Na<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O system at 24 °C after cooling from 850 °C. The Raman peaks in the spectra are assigned to ReO<sub>4</sub><sup>--</sup> (332, 919, and 971 cm<sup>-1</sup>), HCO<sub>3</sub><sup>--</sup> (1017 cm<sup>-1</sup>), and CO<sub>3</sub><sup>2--</sup> (1065 cm<sup>-1</sup>) (Eysel and Kanellakopulos 1993; Hurai et al. 2015). For clarity, the spectra were shifted along the intensity axis.

628

629 Figure 7. Raman spectra of the aqueous solution and melt measured at different temperatures in 630 the  $SnO_2$ -Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O experiment. The spectra colored in red and blue were obtained from the 631 aqueous solution during the heating and cooling processes, respectively. The spectra colored in 632 purple were obtained from the melt droplet coexisting with the aqueous solution during cooling process. The Raman peaks in the spectra are assigned to  $\text{ReO}_4^-$  (332, 919, and 971 cm<sup>-1</sup>),  $\text{CO}_3^{2-1}$ 633 in aqueous solution (1065 cm<sup>-1</sup>), and  $CO_3^{2-}$  in melt (~702 and ~1072 cm<sup>-1</sup>) (Bates et al. 1972; 634 635 Eysel and Kanellakopulos 1993; Carper et al. 2012; Hurai et al. 2015). For clarity, the spectra 636 were shifted along the intensity axis.

637

Figure 8. Raman spectra of the aqueous solution in the SnO<sub>2</sub>-NaOH-H<sub>2</sub>O experiment measured 638 639 during heating from 24 °C to 800 °C (in red color) and at 24 °C after cooling from 800 °C (in blue color). The Raman peaks in the spectra are assigned to ReO<sub>4</sub><sup>-</sup> (332, 919, and 971 cm<sup>-1</sup>) and 640  $Sn(OH)_6^{2-}$  (552 cm<sup>-1</sup>) (Taylor and Coddington 1992; Eysel and Kanellakopulos 1993). For 641 642 clarity, the spectra were shifted along the intensity axis.

643

644 Figure 9. Mole fraction diagram of Li<sub>2</sub>CO<sub>3</sub> and crystallized cassiterite in the SnO<sub>2</sub>-Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O 645 experiments. The experimental data under similar temperature and pressure conditions 646 (approximately 700–800 °C and 700–800 MPa) were selected from Table 1, where the numbers 647 represent the experimental numbers. The solid line represents the linear regression fit curve.

648

649 Figure 10. (a) Cassiterite crystallization conditions in the  $SnO_2$ -Li<sub>2</sub>CO<sub>3</sub>-H<sub>2</sub>O system. The dashed 650 lines show the pressure-temperature (P-T) paths of each experiment listed in Table 1. The gray 651 area represents the formation conditions of rare element pegmatite proposed by London (2008).

652	(b) The P–T diagram of the five groups of comparative experiments in the $SnO_2-Li_2CO_3-H_2O$
653	system performed using the same initial samples under different pressure conditions were
654	marked with the experimental group numbers. The negative trends of the initiation and ending
655	P-T conditions of cassiterite crystallization in the same group of experiments were marked with
656	dotted lines, and overall P-T trends were indicated with thick gray arrows. In both (a) and (b),
657	the red circles and blue diamonds represent the crystallization-beginning and ending P-T
658	conditions, respectively.

659

<sup>a</sup> No.	${}^{b}m(SnO_{2})$ (10 <sup>-6</sup> g)	$^{c}m(\text{Li}_{2}\text{CO}_{3})$ (10 <sup>-6</sup> g)	$^{d}m(H_{2}O)$ (10 <sup>-6</sup> g)	$e_x(SnO_2)$ (mol%)	$f_x(\text{Li}_2\text{CO}_3)$ (mol%)	${}^{g}T_{h}$ (°C)	$^{h}T_{B}$ (°C)	$^{i}T_{\rm S}$ (°C)	<sup>j</sup> P <sub>B</sub> (MPa)	$^{k}P_{S}$ (MPa)	<sup>I</sup> N	$(10^{-6} \text{ cm/s})$	$^{n}v_{aveV}$ ( $\mu m^{3}/s$ )
1-1	0.66	2.32	30.59	0.25	1.81	225	824	631	864	602	1	0.62	8.20
1-2	0.66	2.32	26.46	0.29	2.08	281	833	750	642	552	1	8.22	19.07
1-3	0.66	2.32	24.91	0.31	2.21	290	856	768	628	537	3	°/	/
2	0.07	1.44	28.08	0.03	/	292	661	560	417	307	5	/	/
3	0.42	2.43	27.89	0.18	2.07	276	802	710	628	525	1	2.81	10.95
4	0.16	2.31	30.74	0.06	/	245	712	680	640	599	20	/	/
5	0.15	2.29	22.73	0.08	2.39	296	770	665	517	408	5	/	/
6	1.09	4.46	30.54	0.41	3.42	231	790	758	795	753	6	/	/
7-1	0.49	2.48	22.34	0.25	2.62	220	741	660	773	660	3	/	/
7-2	0.49	2.48	21.24	0.27	/	250	825	704	761	612	5	/	/
8-1	0.43	2.90	29.85	0.17	2.30	212	734	678	794	715	3	/	/
8-2	0.43	2.90	27.94	0.18	2.46	252	771	683	687	578	7	/	/
9-1	0.29	2.51	34.27	0.10	1.75	166	654	605	844	764	13	/	/
9-2	0.29	2.51	28.87	0.12	2.07	237	781	645	759	580	5	/	/
10	0.45	4.00	28.84	0.18	3.26	198	703	627	803	689	10	/	/
11	0.16	1.50	27.24	0.07	/	235	702	590	664	512	11	/	/
12	0.36	2.38	23.31	0.18	2.42	268	797	710	655	554	11	/	/
13-1	0.24	2.60	27.28	0.10	/	247	675	615	585	507	3	/	/
13-2	0.24	2.60	23.79	0.12	/	292	757	690	518	448	5	/	/
13-3	0.24	2.60	21.76	0.13	/	309	797	720	493	420	5	/	/
14	0.09	2.15	23.70	0.04	/	337	701	626	306	246	27	/	/

 $\label{eq:constant} 660 \qquad \text{Table 1. Experimental results of cassiterite crystallization in the $SnO_2-Li_2CO_3-H_2O$ system.}$ 

661 Note:

- <sup>662</sup> <sup>a</sup> No. Experimental number: a total of 21 experiments including five groups of comparative experiments. Each experimental group was labeled with the same
- prefix number (i.e., 1, 7, 8, 9, and 13). The experiments in each group were performed using the same initial samples but different bulk H<sub>2</sub>O density in the HDAC
- 664 sample chamber, to investigate the effect of pressure on cassiterite crystallization.
- $665 \quad {}^{b}m(\text{SnO}_2)$ . Initial mass of SnO<sub>2</sub> approximately calculated through the volume of cassiterite crystallized in the sample chamber and cassiterite density (6.95 g/cm<sup>3</sup>).
- $^{\circ}m(\text{Li}_2\text{CO}_3)$ . Initial mass of  $\text{Li}_2\text{CO}_3$  approximately calculated through the volume of compacted  $\text{Li}_2\text{CO}_3$  powder loaded in the sample chamber and the  $\text{Li}_2\text{CO}_3$ density (2.11 g/cm<sup>3</sup>).
- $^{d}m(H_2O)$ . Initial mass of  $H_2O$  approximated through the calculated  $H_2O$  volume and bulk density determined based on the measured vapor-disappearing
- 670 temperature and the equation of the state of H<sub>2</sub>O (Wagner and Pru $\beta$  2002).
- 671 ex(SnO<sub>2</sub>). Mole fraction of cassiterite crystallized from the aqueous solution calculated through the initial mass of SnO<sub>2</sub>, Li<sub>2</sub>CO<sub>3</sub>, and H<sub>2</sub>O in the sample
- 672 chamber.
- 673  $f_x(Li_2CO_3)$ . Mole fraction of  $Li_2CO_3$  in the aqueous solution at  $T_B$  calculated through the initial mass of  $SnO_2$ ,  $Li_2CO_3$ , and  $H_2O$  in the sample chamber. Note that
- 674 this calculation method only works for experiments in which no  $Li_2CO_3$  melt droplets appeared in the sample chamber at  $T_B$ , ensuring the  $x(Li_2CO_3)$  represented
- 675 the mole fraction of  $Li_2CO_3$  dissolved in aqueous solution at  $T_B$ .
- 676  ${}^{g}T_{h}$ . Temperature at which the vapor bubble disappeared in HDAC sample chamber for the vapor bubble generated during cooling process.
- 677 <sup>h</sup> $T_{\rm B}$ . Temperature at which cassiterite crystals began to grow.
- 678 <sup>i</sup> $T_{\rm S}$ . Temperature at which cassiterite crystals stopped growing.
- $^{5}P_{B}$ . Approximate pressure at  $T_{B}$  based on the equation of state of H<sub>2</sub>O (Wagner and Pru $\beta$  2002) due to the low solubility of Li<sub>2</sub>CO<sub>3</sub> in water (Anderson et al.
- 680 2001).
- 681 <sup>k</sup> $P_{\rm S}$ . Approximate pressure at  $T_{\rm S}$ .
- 682 <sup>I</sup>N. Number of cassiterite crystals formed in the HDAC sample chamber.
- 683 <sup>m</sup> $v_{aveL}$ . Average length growth rate of cassiterite for the experiments in which only one main crystal nucleated and grew during cooling.
- 684  $v_{aveV}$ . Average volume growth rate of cassiterite for the experiments in which only one main crystal nucleated and grew during cooling.
- 685 °/. Not determined.

<sup>a</sup> No.	${}^{\mathrm{b}}T_{\mathrm{h}}({}^{\mathrm{o}}\mathrm{C})$	$^{c}T_{\rm B}(^{\circ}{\rm C})$	$^{\mathrm{d}}T_{\mathrm{S}}(^{\mathrm{o}}\mathrm{C})$	еN	${}^{\rm f}v_{\rm aveL} (10^{-6} {\rm cm/s})$	$^{g}v_{aveV} (\mu m^{3}/s)$
1	354	645	522	1	0.92	3.40
2	220	664	467	27	h/	/
3	279	551	390	1	0.73	5.91
4	285	589	397	9	/	/
5	315	584	405	1	0.61	5.43
6	309	476	403	29	/	/
7	269	580	436	15	/	/
8	271	618	459	16	/	/
9	258	659	422	3	/	/
10	245	730	523	4	/	/
11	288	753	619	1	4.92	18.72

 $\label{eq:constant} 686 \qquad \text{Table 2. Experimental results of cassiterite crystallization in the $SnO_2-Na_2CO_3-H_2O$ system.}$ 

687 Note: <sup>a</sup>No., <sup>b</sup> $T_{h}$ , <sup>c</sup> $T_{B}$ , <sup>d</sup> $T_{S}$ , <sup>e</sup>N, <sup>f</sup> $v_{aveL}$ , <sup>g</sup> $v_{aveV}$ , and <sup>h</sup>/ bear the same representations as those described in Table 1.

688

Fig. 1













Fig. 6





Intensity (a.u.)







Temperature (°C)

