| 1  | Revision 2  |
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| 2  | Phase transition boundary between fcc and hcp structures in Fe-Si alloy and   |
| 3  | its implications for terrestrial planetary cores  |
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### ABSTRACT

32The phase transition between a face-centered cubic (fcc) and hexagonal close-packed (hcp) structures in Fe-4wt% Si alloy was examined in an internally resistive heated 33 diamond anvil cell (DAC) under high-pressure (P) and -temperature (T) conditions to 343571 GPa and 2000 K by in-situ synchrotron X-ray diffraction. Complementary 36 laser-heated DAC experiments were performed in Fe-6.5wt% Si. The fcc-hcp phase transition boundaries in the Fe-Si alloys are located at higher temperatures than that in 37 pure Fe, indicating that the addition of Si expands the hcp stability field. The dP/dT 38slope of the boundary of the entrant fcc phase in Fe-4wt% Si is similar to that of pure Fe, 39 but the two-phases region is observed over a temperature range increasing with pressure, 40 going from 50 K at 15 GPa to 150 K at 40 GPa. The triple point, where the fcc, hcp, and 4142liquid phases coexist in Fe-4wt% Si, is placed at 90-105 GPa and 3300-3600 K with the melting curve same as in Fe is assumed. This supports the idea that the hcp phase is 43stable at Earth's inner core conditions. The stable structures of the inner cores of the 44other terrestrial planets are also discussed based on their P-T conditions relative to the 4546 triple point. In view of the reduced P-T conditions of the core of Mercury (well below the triple point), an Fe-Si alloy with a Si content up to 6.5 wt% would likely crystallize 47an inner core with an fcc structure. Both Venusian and Martian cores are believed to 48

| 49 | currently be totally molten. Upon secular cooling, Venus is expected to crystallize an      |
|----|---|
| 50 | inner core with an hcp structure, as the pressures are similar to those of the Earth's core |
| 51 | (far beyond the triple point). Martian inner core will take an hcp or fcc structure         |
| 52 | depending on the actual Si content and temperature.   |
| 53 |   |
| 54 | Key words   |
| 55 | Earth's core; high-pressure; diamond anvil cell; internal resistive heating; Fe-Si alloy    |
| 56 |   |
| 57 | INTRODUCTION  |
| 58 | Terrestrial core formation process has been discussed in relation to                        |
| 59 | metal-silicate equilibration during accretion stage (Li and Agee 1996; Wade and Wood        |
| 60 | 2005; Siebert et al. 2013) although some recent models considered disequilibrium            |
| 61 | processes at a later stage (Rubie et al. 2011). The metal-silicate equilibration inevitably |
| 62 | results in an impure iron rich metallic core (Wade and Wood; Siebert et al. 2013). The      |
| 63 | impurity includes nickel and several less dense elements which are also called light        |
| 64 | elements (Poirier 1994; Allègre et al. 1995).   |
| 65 | Birch (1952) pointed out that the density of pure iron might be greater than the            |
|    | Biten (1952) pointed out that the density of pure non inight be greater than the            |

| 67                         | with the presence of lighter element(s), and a recent internally consistent   |
|----------------------------|---|
| 68                         | thermodynamic model of pure iron estimated the core density deficit to be 7% for the  |
| 69                         | outer core and 4.5 % for the inner core (Komabayashi 2014). Other terrestrial planetary   |
| 70                         | cores also likely contain light elements considering metal-silicate partitioning during   |
| 71                         | their cores formation. In addition, the presence of a magnetic field found in some  |
| 72                         | terrestrial planets may indicate the presence of a light element-bearing partially molten   |
| 73                         | core (e.g., Sohl and Schubert 2007). Light elements would be expelled at the bottom of  |
| 74                         | the liquid outer core as it is less partitioned in the solid inner core and this would drive  |
| 75                         | convection in the outer core (Stevenson et al. 1983; Lister and Buffett 1995).  |
|                            |   |
| 76                         | Among the potential light elements, silicon is considered a plausible candidate   |
| 76<br>77                   | Among the potential light elements, silicon is considered a plausible candidate<br>for the terrestrial planetary cores for various reasons: i) silicon is the second most   |
|                            |   |
| 77                         | for the terrestrial planetary cores for various reasons: i) silicon is the second most  |
| 77<br>78                   | for the terrestrial planetary cores for various reasons: i) silicon is the second most abundant element in the mantle and series of high-pressure (P) and -temperature (T)  |
| 77<br>78<br>79             | for the terrestrial planetary cores for various reasons: i) silicon is the second most<br>abundant element in the mantle and series of high-pressure (P) and -temperature (T)<br>experiments demonstrated that silicon and oxygen could be dissolved from mantle  |
| 77<br>78<br>79<br>80       | for the terrestrial planetary cores for various reasons: i) silicon is the second most<br>abundant element in the mantle and series of high-pressure (P) and -temperature (T)<br>experiments demonstrated that silicon and oxygen could be dissolved from mantle<br>silicates to core melt (Takafuji et al. 2005; Ozawa et al. 2009), and then silicon is   |
| 77<br>78<br>79<br>80<br>81 | for the terrestrial planetary cores for various reasons: i) silicon is the second most<br>abundant element in the mantle and series of high-pressure (P) and –temperature (T)<br>experiments demonstrated that silicon and oxygen could be dissolved from mantle<br>silicates to core melt (Takafuji et al. 2005; Ozawa et al. 2009), and then silicon is<br>partitioned between solid and liquid during core crystallization; ii) silicon isotopic |

| 85  | formation models based on silicate-metal equilibration inevitably have silicon as a light   |
|-----|---|
| 86  | element in the core (Wade and Wood 2005; Rubie et al. 2011; Siebert et al. 2013).           |
| 87  | Phase relations and equations of state (EoS) of solid phases in the system                  |
| 88  | Fe-(Fe)Si have been extensively studied by both experiment and theory (Alfe et al.          |
| 89  | 2002; Dobson et al. 2002; Lin et al. 2002; Kuwayama and Hirose 2004; Lin et al. 2009;       |
| 90  | Tateno et al. 2015; Ozawa et al. 2016). An important phase relation is the transition       |
| 91  | between face-centered cubic (fcc) and hexagonal close-packed (hcp) structures, as this      |
| 92  | is central to address the solid inner core structure (Uchida et al. 2001; Asanuma et al.    |
| 93  | 2008; Komabayashi et al. 2009), and the P-T location of the phase boundary can be           |
| 94  | used to deduce thermodynamic properties (Wood 1993; Komabayashi 2014). Notably,             |
| 95  | the triple point P-T location where the hcp, fcc, and liquid phases coexist can be          |
| 96  | constrained from the fcc-hcp boundary and melting curve (Zhang et al. 2016). An             |
| 97  | experimental study in a laser-heated diamond anvil cell (DAC) reported that the             |
| 98  | transition temperature was greatly reduced when 3.4 wt%Si was added to Fe (Asanuma          |
| 99  | et al. 2008). In contrast, phase relations inferred by Fischer et al. (2013) suggested that |
| 100 | addition of silicon should increase the transition temperature. Tateno et al. (2015)        |
| 101 | experimentally demonstrated that the transition temperature was increased by the            |
| 102 | addition of 6.5 wt%Si to iron. As such the effect of Si on the transition temperature has   |

been a controversy and the P-T conditions of the actual boundaries in the system Fe-Si 103 104 are not unanimous. In this study, we present the investigation of the P-T locations of the fcc-hcp 105transition boundaries in Fe-Si alloys in an internally resistive heated DAC. The 106 107internally heated DAC heats the sample by its resistance, with an improved accuracy in 108 temperature with respect to conventional laser heated DAC (Komabayashi et al. 2009; 2012). Based on these experimental results, we will discuss the effect of Si on the Fe 109 properties under high P-T condition and address the stable structure of a solid Fe-Si 110 alloy at the conditions of the inner cores of the terrestrial planets of the solar system. 111 112113**EXPERIMENTAL PROCEDURE** 114We conducted high-P-T in-situ X-ray diffraction (XRD) experiments on Fe-Si 115samples at the beamline ID27, European Synchrotron Radiation Facility (ESRF). X-rays with a wavelength of 0.3738Å were focused to a  $3x3 \text{ }\mu\text{m}^2$  spot at sample position and 116 the diffracted X-rays were collected on a two dimensional detector (mar345 Image Plate 117

Detector). The collection time was 10 seconds for each measurement. Using the fit-2D
program (Hammersley 1996), the obtained data were converted to the conventional

120 one-dimensional XRD pattern.

| 121 | High pressure was generated in a DAC with a pair of diamond anvils with a                 |
|-----|---|
| 122 | culet size of 300 $\mu m$ or 150-450 $\mu m$ beveled depending on the pressure range. The |
| 123 | starting material was a 5-7 $\mu m$ thick Fe-Si alloy with 4 wt% Si (Rare Metallic. Co.,  |
| 124 | hereafter Fe-4Si), placed in the sample chamber and connected to platinum leads. The      |
| 125 | junction between the Fe sample and Pt leads was outside the sample chamber (see           |
| 126 | Komabayashi et al. (2009) for the sample geometry). $SiO_2$ glass layers served as a      |
| 127 | pressure transmitting medium and thermal insulator. High temperature was achieved         |
| 128 | with an internal resistive system (Komabayashi et al. 2009; 2012; Antonangeli et al.      |
| 129 | 2012). The sample was resistively heated by directly applying a DC voltage by an          |
| 130 | external power supply. The temperature was measured by a spectral radiometric system      |
| 131 | as conventional in laser heating experiments. Noteworthy, thanks to the improved time     |
| 132 | and spatial stability of the hotspot and the reduced thermal gradients, resulting         |
| 133 | uncertainties in temperature were about 50 K (Komabayashi et al., 2012)                   |
| 134 | Complementary laser heating experiments were conducted on an Fe-6.5Si                     |

Complementary laser heating experiments were conducted on an Fe-6.5Si sample (Rare Metallic. Co.) at ID27, ESRF (see Morard et al. (2011) for details of the laser heating experimental set up). The internal heating system was not applied because this alloy is so brittle that it was not possible to make it into thin foil. Irrespectively of the use of low numerical aperture and reflecting objectives which effectively minimize

| 139 | the chromatic aberration and improve reliability of temperature determination (Mezouar                                   |
|-----|--|
| 140 | et al. 2017), the laser heating experiments show larger temperature uncertainty due to                                   |
| 141 | large temperature gradient across the sample and laser fluctuations. The two microns                                     |
| 142 | diameter pinhole at the entrance of the spectrometer allows for a collection of signal                                   |
| 143 | only coming from the very central part of the hotspot and an optimal alignment of lasers                                 |
| 144 | and X-ray beam. Morard et al. (2011) discussed that the possible uncertainty in the                                      |
| 145 | temperature in this experimental setup might be 150 K at 3000 K. In the present study,                                   |
| 146 | we assumed a more conservative number, 10% of the generated temperature.   |
| 147 | In all the runs, the pressure at any given temperature was calculated with a   |
| 148 | thermal EoS for Fe-4Si or Fe-6.5Si with the hcp structure that was assessed based on                                     |
| 149 | pure iron (Dewaele et al. 2006) and Fe-9wt%Si (Tateno et al. 2015). The room   |
| 150 | temperature parameters for the Vinet EoS were obtained by averaging on the basis of                                      |
| 151 | mole fraction between the two compositions: for Fe-4Si, $V_0 = 22.56$ Å <sup>3</sup> , $K_0 = 166$ GPa,                  |
| 152 | K' = 5.4, and for Fe-6.5Si, $V_0 = 22.63 \text{ Å}^3$ , $K_0 = 167 \text{ GPa}$ , K' = 5.5, where $V_0$ , $K_0$ , K' are |
| 153 | the unit-cell volume, bulk modulus, and its pressure derivative at 300 K and 1 bar,                                      |
| 154 | respectively. We assumed the same thermal parameters as for pure iron (Dewaele et al.                                    |
| 155 | 2006; Tateno et al. 2015).   |

156

The thermal pressure effect on the sample pressure was checked against the

| 157 | pressure for the $SiO_2$ pressure medium. The EoS for stishovite by Wang et al (2012) was               |
|-----|---|
| 158 | used to calculate the pressure for $SiO_2$ . As the precise temperature for the pressure                |
| 159 | medium was unknown, we calculated the pressure for $\mathrm{SiO}_2$ at the sample temperature           |
| 160 | and 300 K. At 42.8 GPa and 1940 K for the sample, the pressure for $SiO_2$ were 44.9                    |
| 161 | GPa and 34.1 GPa at 1940 K and 300 K respectively. Since the crystallized portion of                    |
| 162 | the pressure medium should be at the same (or slightly lower) temperature of the sample,                |
| 163 | the above calculation independently supports our pressure determination for the iron                    |
| 164 | alloy. Also, after quench, the pressure for SiO <sub>2</sub> is $37.6 \pm 1.0$ GPa, which is consistent |
| 165 | with the sample pressure of $36.6 \pm 0.4$ GPa. Therefore, we can conclude that the                     |
| 166 | pressure estimation at high temperatures is reasonable.   |
| 167 | The use of the unit-cell volume of the hcp phase in the pressure calculation                            |

may introduce an uncertainty when the experimental condition is near the completion of the hcp-fcc reaction. Based on a binary temperature-composition (T-X) phase loop detailed below, the pressure could have been underestimated by less than 1.3 GPa at 40 GPa and 1870K. When the unit-cell volume for the hcp phase was not obtained, due to either grain growth or complete transition to the fcc phase, we assumed constant pressure upon further heating.

174

### RESULTS

175

176 **Fe-4Si** 

177 Six separate in-situ XRD experiments were carried out on the Fe-4Si sample in 178 the internally heated DAC. The results are illustrated in Fig. 1a and summarized in 179 Table 1.

In the first run, the sample was compressed to 16.4 GPa and the XRD pattern 180 181 shows coexistence of the bcc and hcp phases. Then the sample was heated to 1060 K 182and the fcc phase, with a minor amount of hcp phase, was observed. The hcp phase disappeared in the following XRD pattern at a similar temperature of 1050 K. 183 Temperature slightly increased with time to 1080 K at steady power from the DC power 184 supply without further changes in the pattern. Then, we tried to reverse the reaction. As 185186 the spectroradiometric method could not reliably measure temperatures below 1000 K, 187the temperature was estimated based on the linear power-temperature relationship established at 1060 K. During the cooling path, the reversal reaction started at 860 K. 188 200 K lower than the reaction in the forward heating cycle. Further decreasing 189 temperature to 770 K only slightly promoted the reaction, which suggests that 190 nucleation of the hcp phase is very sluggish and implies that the width of the reaction in 191192the backward cycle is much wider than the forward cycle. Accordingly, we constrained

the P-T conditions of the reaction in the forward cycle only and the results of thebackward cycle are not listed in Table 1 to avoid confusion.

In the following runs, we only employed heating cycles which started from the 195hcp phase towards the fcc stability field. In the runs 2, 3, and 4, we observed a transition 196 197 sequence from hcp to hcp+fcc and to fcc with increasing temperature. Figure 2 shows a 198 series of XRD patterns collected during the run 2 for increasing temperature at about 24 GPa. The temperature was first held at 1120 K for 4 minutes, observing only the hcp 199 200 phase. Then we increased the temperature to 1230 K and fcc peaks appeared. No further 201changes in the XRD pattern were recognized during the following 40 minutes, during which the temperature was kept constant. We further increased the temperature to 1260 202K, which instantaneously increased the intensity of the fcc peaks. During the following 20320430 minutes at constant temperature, the XRD patterns did not show significant changes. 205The transition was completed at 1330 K. In summary, the drastic changes in XRD were 206 observed only upon temperature increase. The transition from the hcp to fcc phase seems to be very fast, with minimal kinetic effects. 207 208In the runs 5 and 6, we observed no structural change to the highest temperatures and confirmed stability of the hcp phase to 71.0 GPa and 2020 K. 209

210

Overall, thanks to the performances of the internal heating system, we have

| 227 | DISCUSSION  |
|-----|---|
| 226 |   |
| 225 | results by Tateno et al. (2015) (Fig. 1b).  |
| 224 | laser heating experiment (i.e., $\pm 10\%$ ), the present experimental data are consistent with |
| 223 | fcc phase at 2340 K and 56.6 GPa. Considering the uncertainty in temperature in the             |
| 222 | coexistence with hcp peaks. In the second run, we observed a complete transition to the         |
| 221 | and then heated it by laser, reporting the appearance of fcc peaks at 1850 K in                 |
| 220 | pressures. In the first run, we compressed the sample to 24.0 GPa at room temperature           |
| 219 | 1b). Same as for the internal-heating runs, we increased the temperature under high             |
| 218 | Two separate laser-heating runs were conducted on the Fe-6.5Si sample (Fig.                     |
| 217 | Fe-6.5Si  |
| 216 |   |
| 215 | 50 K at 15 GPa to 150 K at 40 GPa (Fig. 1a).  |
| 214 | the temperature interval of the two-phase region expands with increasing pressure from          |
| 213 | the boundary of the entrant fcc phase in Fe-4wt% Si is similar to that of pure Fe, while        |
| 212 | particular, the width of the phase loop was accurately constrained. The dP/dT slope of          |
| 211 | been able to place tight constraints on the P-T location of the transition boundaries. In       |

# 228 Effect of Si on the fcc-hcp transition

| 229 | The present experiments confirmed the enlarged high temperature stability of             |
|-----|--|
| 230 | the hcp phase in Fe-4Si and Fe-6.5Si with respect to the case in pure Fe (Fig. 1), in    |
| 231 | agreement with Fischer et al. (2013) and Tateno et al. (2015) but in contrast to Asanuma |
| 232 | et al. (2008). Dissecting the XRD patterns in Asanuma et al. (2008), we noted that they  |
| 233 | assigned tiny shallow rises as peaks from the fcc phase, while the appearance of the fcc |
| 234 | phase is clearly marked by the presence of the (200) peak (Fig. 2) (Komabayashi et al.   |
| 235 | 2009; 2012). We conclude that the transition temperature between the hcp and fcc         |
| 236 | phases increases with Si content.  |
| 237 | Figure 3 shows a T-X diagram at 40 GPa based on the present data on Fe-4Si               |
| 238 | and Fe-6.5Si and existing experimental data on pure Fe and Fe-9Si (Komabayashi et al.    |
| 239 | 2009; Tateno et al. 2015). The fcc-hcp transition temperature increases with Si content. |
| 240 | The P-T conditions for a reaction $hcp = hcp + B2$ phase (Tateno et al. 2015) placed     |
| 241 | additional constraint on the phase diagram (Fig. 3). The maximum solubility of Si into   |
| 242 | the fcc phase should be about 7 wt%. A thermodynamic model will be made to fit the       |
| 243 | data in the near future.   |
| 244 |  |

245

# IMPLICATIONS

Figure 4 shows a phase diagram of iron alloys reporting the fcc-hcp boundaries

| 247 | in Fe, Fe-4Si, and Fe-6.5Si, together with the P-T ranges for the cores of Mars and        |
|-----|--|
| 248 | Mercury. Addition of Si to Fe expands the stability of the hcp phase as confirmed by the   |
| 249 | P-T locations of the fcc-hcp transitions observed in this study, which are consistent with |
| 250 | Tateno et al. (2015) based on a laser-heated DAC experiments (Fig. 1b). The triple point   |
| 251 | where the fcc, hcp, and liquid phases coexist in Fe-4wt% Si is placed at 90-105 GPa and    |
| 252 | 3300-3600 K (the melting curve is assumed to be the same as in Fe as Si inclusion at       |
| 253 | 4wt% level should not significantly affect the melting temperature (Morard et al. 2011)),  |
| 254 | supporting the idea that Earth's inner core at 330-364 GPa is made up with the hcp         |
| 255 | phase. Tateno et al. (2015) similarly proposed that the inner core would be made of a      |
| 256 | sole hcp alloy if the Si content is up to 7 wt.%.  |

More complex can be the cases for the other terrestrial planets of the solar 257258system, namely, Venus, Mercury, and Mars. Due to the lack of seismic data, 259information about their internal structures heavily depends on the average density. As such, the core density deficit cannot be pertinently discussed. Nevertheless geophysical 260modelling studies argued for the presence of light elements in the cores and kept this 261into consideration when simulating planetary core, and in particular when discussing 262their molten/solid state (Stevenson et al. 1983; Williams and Nimmo 2004; Rivoldini et 263264al. 2011; Dumberry and Rivoldini 2015; Knibbe and van Westrenen 2018). In the

| 265 | following, we will limit our discussion to the Fe-Si-S system. Tsujino et al. (2013)       |
|-----|--|
| 266 | summarized existing thermal models for terrestrial core-mantle boundaries in the           |
| 267 | system Fe-S (Stevenson et al. 1983; Sohl and Spohn 1997; Fei et al. 2000; Williams and     |
| 268 | Nimmo 2004) and assessed the adiabats across the cores on the basis of the Grüneisen       |
| 269 | parameter for pure fcc iron (Fig. 4). The known reduction of the crystallizing             |
| 270 | temperature of iron by addition of silicon (Kubaschewski 1993; Kuwayama and Hirose         |
| 271 | 2004) implies that the thermal profiles in Fig. 4, which are based on the liquidus of the  |
| 272 | system Fe-S, can be considered as the maximum estimates for the system Fe-Si-S.            |
| 273 | Venus' similar size to the Earth implies that its internal structure is                    |
| 274 | differentiated into crust, mantle, and core (Sohl and Schubert 2007). However,             |
| 275 | contrarily to the Earth, Venus does not have a global magnetic field .The pressure at the  |
| 276 | center of the planet was estimated to be 295 GPa which is slightly lower than of the       |
| 277 | Earth and crystallization of the liquid core might have not yet started (Stevenson et al.  |
| 278 | 1983). Since the core pressure is far greater than the triple point pressure in the system |
| 279 | Fe-4Si (Fig. 4), when upon secular cooling the inner core will start crystallizing,        |
| 280 | likewise the Earth, it should take an hcp phase.   |
| 281 | It is suggested that Mercury has a partially molten iron core as it shows a                |

dipole magnetic field from a spacecraft observation (Ness 1979). Further support of the

| 283 | presence of a liquid portion of the core comes from and the amplitude of its librations      |
|-----|--|
| 284 | (Margot et al. 2005). The suggested thermal structure of Mercurian core modeled on the       |
| 285 | basis of the system Fe-S argues for a temperature much higher than the fcc-hcp               |
| 286 | transition in Fe-6.5Si (Fig. 4). Recently proposed thermal models of Fe-Si cores showed      |
| 287 | a similar temperature range as in Fig. 4 (Knibbe and van Westrenen 2018). Hence              |
| 288 | Mercurian inner core is expected to take an fcc structure if the Si content is less than 6.5 |
| 289 | wt%. Only upon further cooling, the fcc-structured alloy in Mercurian inner core, will       |
| 290 | be transformed to the hcp phase.   |
| 291 | Mars currently does not have an active global magnetic field, although it has a              |
| 292 | metallic core in view of its average density and moment of inertia (Yoder et al. 2003).      |
| 293 | However, the presence of magnetized rock records found in southern highland area             |
| 294 | indicates that the planet should have had a magnetic field in the past. Stevenson et al.     |
| 295 | (1983) suggested that present Martian core should still be totally molten to account for     |
| 296 | the absence of the magnetic field and the rocks were magnetized by a field produced by       |
| 297 | a past, now extinct, thermal convection of the liquid core. The inferred thermal profiles    |
| 298 | of the core partially overlap with the fcc-hcp transition in Fe-4Si (Fig. 4). As such the    |
| 299 | first iron alloy crystal that will crystallize in the future will be either fcc or hcp phase |
| 300 | depending on the actual Si content and exact temperature. In the case of the fcc phase, it   |

301 will be transformed to the hcp phase as the core further cools down.

| 302 | In summary, the inner core structure of the terrestrial planets can be discussed          |
|-----|---|
| 303 | based on the phase relations in Fig. 4. The hcp core would show more anisotropic          |
| 304 | seismic properties (Steinle-Neumann et al., 2001) than the fcc core because of the        |
| 305 | anisotropic crystal structure. Also the hcp core would be denser than the fcc core as the |
| 306 | fcc-hcp transition in Fe-4Si shows about 0.8% density jump. Those changes are             |
| 307 | important in future attempts to construct a precise density model for the planetary       |
| 308 | interior.   |

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310

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# 317 References

- 318 Alfè, D., Gillan, M.J., and Price, G.D. (2002) Composition and temperature of the
- Earth's core constrained by combining ab initio calculations and seismic data.
- Earth and Planetary Science Letters 195, 91-98.
- Allègre, C.J., Poirier, J.P., Humler, E., and Hofmann, A.W. (1995) The chemical
- 322 composition of the Earth. Earth and Planetary Science Letters 134, 515-526.
- Anderson, J.D., Colombo, G., Esposito, P.B., Lau, E.L., and Trager, G.B. (1987) The

Mass, Gravity-Field, and Ephemeris of Mercury. Icarus 71, 337-349.

- 325 Antonangeli, D., Komabayashi, T., Occelli, F., Borissenko, E., Walters, A.C., Fiquet, G.,
- and Fei, Y.W. (2012) Simultaneous sound velocity and density measurements of
- hcp iron up to 93 GPa and 1100 K: An experimental test of the Birch's law at
- high temperature. Earth and Planetary Science Letters 331, 210-214.
- 329 Asanuma, H., Ohtani, E., Sakai, T., Terasaki, H., Kamada, S., Hirao, N., Sata, N., and
- 330 Ohishi, Y. (2008) Phase relations of Fe-Si alloy up to core conditions:
- Implications for the Earth inner core. Geophysical Research Letters 35, L12307,
- doi:10.1029/2008GL033863.
- Birch, F. (1952) Elasticity and constitution of the Earth's interior. Journal of
  Geophysical Research 57, 227-286.

- 335 Dewaele, A., Loubeyre, P., Occelli, F., Mezouar, M., Dorogokupets, P.I., and Torrent, M.
- 336 (2006) Quasihydrostatic equation of state of iron above2 Mbar. Physical Review
- 337 Letters 97, 215504.
- Dobson, D.P., Vocadlo, L., and Wood, I.G. (2002) A new high-pressure phase of FeSi.
  American Mineralogist 87, 784-787.
- 340 Dumbarry, M., Rivoldini, A. (2015) Mercury's inner core size and core-crystallization
- 341 regime. Icarus 248, 254–268.
- 342 Fei, Y.W., Li, J., Bertka, C.M., and Prewitt, C.T. (2000) Structure type and bulk
- modulus of Fe3S, a new iron-sulfur compound. American Mineralogist 85,
  1830-1833.
- Fischer, R.A., Campbell, A.J., Reaman, D.M., Miller, N.A., Heinz, D.L., Dera, P., and
- Prakapenka, V.B. (2013) Phase relations in the Fe-FeSi system at high pressures
- and temperatures. Earth and Planetary Science Letters 373, 54-64.
- 348 Fitoussi, C., Bourdon, B., Kleine, T., Oberli, F., and Reynolds, B.C. (2009) Si isotope
- systematics of meteorites and terrestrial peridotites: implications for Mg/Si
   fractionation in the solar nebula and for Si in the Earth's core. Earth and
- 351 Planetary Science Letters 287, 77-85.
- Georg, R.B., Halliday, A.N., Schauble, E.A., and Reynolds, B.C. (2007) Silicon in the

Earth's core. Nature 447, 1102-1106, doi:10.1038/nature05927. 353

- Hammersley, J. (1996) FIT2D V12.012 Reference Manual. Eur. Synchrotron Radiat. 354
- Facil., Grenoble, France. 355
- Hin, R.C., Fitoussi, C., Schmidt, M.W., and Bourdon, B. (2014) Experimental 356
- 357determination of the Si isotope fractionation factor between liquid metal and 358liquid silicate. Earth and Planetary Science Letters 387, 55-66.
- Knibbe, J.S., and van Westrenen, W. (2018) The thermal evolution of Mercury's Fe-Si 359

360 core. Earth and Planetary Science Letters 482, 147-159.

- Komabayashi, T. (2014) Thermodynamics of melting relations in the system Fe-FeO at 361high pressure: Implications for oxygen in the Earth's core. Journal of
- Geophysical Research 119, DOI: 10.1002/2014JB010980. 363
- 364 Komabayashi, T., Fei, Y., Meng, Y., and Prakapenka, V. (2009) In-situ X-ray diffraction
- 365measurements of the  $\gamma$ - $\varepsilon$  transition boundary of iron in an internally-heated diamond anvil cell. Earth and Planetary Science Letters 282, 252-257. 366
- Komabayashi, T., Hirose, K., and Ohishi, Y. (2012) In situ X-ray diffraction 367 measurements of the fcc-hcp phase transition boundary of an Fe-Ni alloy in an 368
- internally heated diamond anvil cell. Physics and Chemistry of Minerals 39, 369
- 329-338. 370

362

- 371 Kubaschewski, O. (1993) in: Okamoto, H. (Ed.), Phase diagram of binary iron alloys.
- 372 ASM International, Ohio, p. 380.
- Kuwayama, Y., and Hirose, K. (2004) Phase relations in the system Fe-FeSi at 21 GPa.
- American Mineralogist 89, 273-276.
- Li, J., and Agee, C.B. (1996) Geochemistry of mantle-core differentiation at high
  pressure. Nature 381, 686-689.
- Lin, J.F., Heinz, D.L., Campbell, A.J., Devine, J.M., and Shen, G.Y. (2002) Iron-silicon
- alloy in Earth's core? Science 295, 313-315.
- Lin, J.F., Scott, H.P., Fischer, R.A., Chang, Y.Y., Kantor, I., and Prakapenka, V.B. (2009)
- 380 Phase relations of Fe-Si alloy in Earth's core. Geophysical Research Letters 36,
- 381 L06306, doi:10.1029/2008GL036990.
- Lister, J.R., and Buffett, B.A. (1995) The Strength and Efficiency of Thermal and
- Compositional Convection in the Geodynamo. Physics of the Earth and Planetary Interiors 91, 17-30.
- Margot, J. L., Peale, S. J., Jurgens, R. F., Slade, M. A., and Holin, I. V. (2007) Large
- Longitude Libration of Mercury Reveals a Molten Core. Science 316, 710-714.
- 387 Mezouar, M., Giampaoli, R., Garbarino, G., Kantor, I., Dewaele, A., Weck, G., Boccato,
- 388 S., Svitlyk, V., Rosa, A.D., Torchio, R., Mathon, O., Hignette, O., and Bauchau,

- S. (2017) Methodology for in situ synchrotron X-ray studies in the laser-heated 389 diamond anvil cell. High Pressure Research 37, 170-180.
- Morard, G., Andrault, D., Guignot, N., Siebert, J., Garbarino, G., and Antonangeli, D. 391
- (2011) Melting of Fe-Ni-Si and Fe-Ni-S alloys at megabar pressures: 392
- 393 implications for the core-mantle boundary temperature. Physics and Chemistry
- 394of Minerals 38, 767-776.

390

- Ness, N.F. (1979) Magnetic-Fields of Mercury, Mars, and Moon. Annual Review of 395 Earth and Planetary Sciences 7, 249-288. 396
- Ozawa, H., Hirose, K., Mitome, M., Bando, Y., Sata, N., and Ohishi, Y. (2009) 397
- Experimental study of reaction between perovskite and molten iron to 146 GPa 398
- and implications for chemically distinct buoyant layer at the top of the core. 399
- 400 Physics and Chemistry of Minerals 36, 355-363.
- 401 Ozawa, H., Hirose, K., Yonemitsu, K., and Ohishi, Y. (2016) High-pressure melting
- experiments on Fe-Si alloys and implications for silicon as a light element in the 402
- core. Earth and Planetary Science Letters 456, 47-54. 403
- 404 Rivoldini, A., Van Hoolst, Verhoeven, O., Mocquet, A., Dehant V. (2011) Geodesy
- constraints on the interior structure and composition of Mars. Icarus 213, 451-405
- 472. 406

| 407 | Rubie, | D.C., | Frost, | D.J., | Mann, | U., | Asahara, | Y., | Nimmo, | F., | Tsuno, | К., | Kegler, | Р., |
|-----|--------|-------|--------|-------|-------|-----|----------|-----|--------|-----|--------|-----|---------|-----|
|     |        |       |        |       |       |     |          |     |        |     |        |     |         |     |

- 408 Holzheid, A., and Palme, H. (2011) Heterogeneous accretion, composition and
- 409 core-mantle differentiation of the Earth. Earth and Planetary Science Letters 301,
  410 31-42.
- 411 Shahar, A., Hillgren, V.J., Young, E.D., Fei, Y.W., Macris, C.A., and Deng, L.W. (2011)
- 412 High-temperature Si isotope fractionation between iron metal and silicate.
- 413 Geochimica et Cosmochimica Acta 75, 7688-7697.
- 414 Siebert, J., Badro, J., Antonangeli, D., and Ryerson, F.J. (2013) Terrestrial Accretion
- 415 Under Oxidizing Conditions. Science 339, 1194-1197.
- 416 Sohl, F., and Schubert, G. (2007) Interior structure, composition, and mineralogy of the
- 417 terrestrial planets, in: Schubert, G. (Ed.), Treatise on Geophysics, pp. 27-68.
- 418 Sohl, F., and Spohn, T. (1997) The interior structure of Mars: Implications from SNC
- 419 meteorites. Journal of Geophysical Research-Planets 102, 1613-1635.
- 420 Steinle-Neumann, G., Stixrude, L., Cohen, R.E., and Gülseren, O. (2001) Elasticity of
- iron at the tempeature of the Earth's inner core. Nature 413, 57-60.
- 422 Stevenson, D.J., Spohn, T., and Schubert, G. (1983) Magnetism and Thermal Evolution
- 423 of the Terrestrial Planets. Icarus 54, 466-489.
- 424 Takafuji, N., Hirose, K., Mitome, M., and Bando, Y. (2005) Solubilities of O and Si in

| 425 | liquid iron in equilibrium with (Mg,Fe)SiO <sub>3</sub> perovskite and the light elements in |
|-----|--|
| 426 | the core. Geophysical Research Letters 32, L06313,   |
| 427 | doi:10.1029/2005GL022773.  |
| 428 | Tateno, S., Kuwayama, Y., Hirose, K., and Ohishi, Y. (2015) The structure of Fe-Si alloy     |
| 429 | in Earth's inner core. Earth and Planetary Science Letters 418, 11-19.                       |
| 430 | Tsujino, N., Nishihara, Y., Nakajima, Y., Takahashi, E., Funakoshi, K., and Higo, Y.         |
| 431 | (2013) Equation of state of gamma-Fe: Reference density for planetary cores.                 |
| 432 | Earth and Planetary Science Letters 375, 244-253.  |
| 433 | Uchida, T., Wang, Y., Rivers, M.L., and Sutton, S.R. (2001) Stability field and thermal      |
| 434 | equation of state of $\varepsilon$ -iron determined by synchrotron X-ray diffraction in a    |
| 435 | multianvil apparatus. Journal of Geophysical Research 106, 21709-21810.                      |
| 436 | Wade, J., and Wood, B.J. (2005) Core formation and the oxidation state of the Earth.         |
| 437 | Earth and Planetary Science Letters 236, 78-95.  |
| 438 | Wang, F.L., Tange, Y., Irifune, T., and Funakoshi, K. (2012) P-V-T equation of state of      |
| 439 | stishovite up to mid-lower mantle conditions. Journal of Geophysical                         |
| 440 | Research-Solid Earth 117, B06209, doi:10.1029/2011JB009100.                                  |
| 441 | Williams, J.P., and Nimmo, F. (2004) Thermal evolution of the Martian core:                  |
| 442 | Implications for an early dynamo. Geology 32, 97-100.  |

| 443 | Wood, B.J. (1993) Carbon in the core. Earth and Planetary Science Letters 117,    |
|-----|---|
| 444 | 593-607.  |
| 445 | Yoder, C.F., Konopliv, A.S., Yuan, D.N., Standish, E.M., and Folkner, W.M. (2003) |
| 446 | Fluid core size of mars from detection of the solar tide. Science 300, 299-303.   |

- 447 Zhang, D.Z., Jackson, J.M., Zhao, J.Y., Sturhahn, W., Alp, E.E., Hu, M.Y., Toellner, T.S.,
- 448 Murphy, C.A., and Prakapenka, V.B. (2016) Temperature of Earth's core
- 449 constrained from melting of Fe and Fe0.9Ni0.1 at high pressures. Earth and
- 450 Planetary Science Letters 447, 72-83.

451

452 Figure captions

453

| 454 | Figure 1. Results of the experiments in (a) Fe-4Si and (b) Fe-6.5Si. The phases observed |
|-----|--|
| 455 | in XRD patterns are plotted: inversed triangle, hcp+bcc; square, hcp; normal triangle,   |
| 456 | fcc+hcp; circle, fcc. In (a), the boundaries between the fcc and hcp phases in pure iron |
| 457 | (Komabayashi et al. 2009) and Fe-3.4Si (Asanuma et al. 2008) are also plotted. The data  |
| 458 | with asterisk have larger temperature uncertainty. In (b), our data are shown together   |
| 459 | with experimental data by Tateno et al. (2015). The typical uncertainty (2.5 GPa and     |
| 460 | 200 K) is shown for a guide to the eye, see Table 1 for the uncertainty for each data    |
| 461 | point. The two datasets are fairly consistent considering the uncertainty in the laser   |
| 462 | heating experiments.   |
| 463 |  |
| 464 | Figure 2. Series of XRD patterns collected in the run 2 for increasing temperature. The  |
| 465 | presence of the fcc phase was unambiguously marked by the appearance of (200) peak.      |
| 466 |  |
|     |  |

Figure 3. Temperature-composition diagram for the fcc-hcp transition. The open circle denotes the transition temperature in pure Fe (Komabayashi et al. 2009); the star symbols are the fcc-hcp reactions constrained by the present experiments; the filled circle is the P-T condition for a reaction of hcp = hcp + B2 phase observed in Tateno et

- al. (2015). The stability fields of fcc+B2 and hcp+B2 were constrained by phase
  relations. In particular: (i) the boundary hcp = hcp+B2 should have a negative slope
  (Tateno et al. 2015) and (ii) the invariant boundary where the fcc, hcp, and B2 phases
  coexist should be placed at a temperature higher than the upper star at Fe-6.5Si.
- 476 Figure 4. The fcc-hcp boundaries in Fe-4Si and Fe-6.5Si (this study) with phase
  477 relations of pure iron (black lines) (Komabayashi et al. 2009; Komabayashi 2014).
  478 Thermal profiles of Mercurian and Martian cores in the system Fe-S are also shown
  479 (Tsujino et al. 2013).

| P, GPa    | $T, K^*$   | Phase  | V(hcp), Å <sup>3</sup>  | a (hcp), Å  | c (hcp), Å   | V(fcc), $Å^3$   | Remarks   |
|-----------|--|--|---|---|--|---|---|
|           |  |  |   |   |  |   |   |
|           |  |  |   |   |  |   |   |
| · · ·     |  | fcc+hcp(small  | ) 21.660(0)   | 2.4855(0)   | 4.04853(0)   | . ,   |   |
| · · ·     |  | fcc  |   |   |  | 43.615(19)  |   |
| 13.1(4)   | 1080   | fcc  |   |   |  | 43.683(19)  |   |
| 23.6(4)   | 300  | hcp  | 20.265(27)  | 2.4394(14)  | 3.9323(24)   |   |   |
| 22.6(4)   | 1120   | hcp  |   |   | 3.9866(3)  |   |   |
| 21.9(5)   | 1230   | hcp+fcc  | 20.955(3)   | 2.4596(2)   | 3.9998(3)  | 42.352(0)   |   |
| 21.2(4)   | 1260   | hcp+fcc  | 21.036(0)   | 2.4596(0)   | 4.0152(0)  | 42.343(38)  |   |
| 21.2      | 1260   | hcp+fcc  | -   | -   | -  | 42.330(5)   | one peak for he   |
| 21.2      | 1270   | hcp+fcc  | -   | -   | -  | 42.370(4)   | one peak for he   |
| 21.2      | 1330   | fcc  |   |   |  | 42.475(15)  |   |
| 21.2      | 1360   | fcc  |   |   |  | 42.520(12)  |   |
| 21.2      | 1430   | fcc  |   |   |  | 42.720(21)  |   |
| 21.2      | 1450   | fcc  |   |   |  | 42.854(69)  |   |
| 33.3(2)   | 300  | hcp  | 19.629(11)  | 2.4193(6)   | 3.8725(11)   |   |   |
| 32.7(5)   | 1030   | hcp  | 20.016(10)  | 2.4280(5)   | 3.9206(10)   |   |   |
| 30.6(5)   | 1190   | hcp  | 20.247(7)   | 2.4342(4)   | 3.9456(7)  |   |   |
| 30.2(5)   | 1240   | hcp  | 20.312(9)   | 2.4361(5)   | 3.9521(9)  |   |   |
| 29.3(1.0) | 1300**   | hcp+fcc  | 20.411(13)  | 2.4399(7)   | 3.9592(13)   | 41.463(0)   |   |
| 29.3(1.1) | 1370**   | hcp+fcc  | 20.453(18)  | 2.4417(9)   | 3.9611(18)   | 41.392(0)   |   |
| 29.0(6)   | 1440   | hcp+fcc  | 20.524(14)  | 2.4434(7)   | 3.9696(14)   | 41.285(0)   |   |
| 28.4(1)   | 1470   | hcp+fcc  | 20.559(5)   | 2.4465(3)   | 3.9663(3)  | 41.344(51)  |   |
| 28.6(8)   | 1500   | hcp+fcc  | 20.595(25)  | 2.4472(14)  | 3.9710(16)   | 41.480(37)  |   |
| 28.6      | 1580   | fcc  | ~ /   | . ,   | . ,  | 41.652(23)  |   |
| 28.6      | 1620   | fcc  |   |   |  | 41.650(19)  |   |
| 41.0(5)   | 300  | hcp  | 19.198(26)  | 2.3956(14)  | 3.8626(24)   |   |   |
| 43.7(6)   | 1130   | hcp  | 19.392(13)  | 2.4019(7)   | 3.8813(12)   |   |   |
| 43.6(6)   | 1190   | hcp  | 19.426(13)  | 2.4028(7)   | 3.8852(12)   |   |   |
| 43.3(6)   | 1250   | -  |   |   | 3.8899(9)  |   |   |
|           |  |  |   |   |  |   |   |
| . ,       |  | -  |   |   |  |   |   |
| 41.9(5)   | 1460   | -  | 19.659(7)   | 2.4103(4)   | 3.9074(7)  |   |   |
| 41.8(7)   | 1410   | -  | 19.637(14)  | 2.4106(8)   | 3.9021(13)   |   |   |
|           |  | -  |   |   |  |   |   |
|           |  | -  |   |   |  |   |   |
|           |  |  |   |   |  |   |   |
| · · ·     |  |  |   |   |  |   |   |
|           |  |  |   |   |  | 39.934(0)   |   |
|           |  |  |   |   |  |   |   |
| . ,       |  |  |   |   |  |   |   |
|           |  |  |   |   |  | • •   |   |
| . ,       |  | -  | 17.000(0)   | )(v)  | 5.5550(0)  |   |   |
|           |  |  |   |   |  | . ,   |   |
|           |  |  |   |   |  | . ,   |   |
|           |  |  |   |   |  |   |   |
| 12.0      | 2250   | 100  |   |   |  | ((-))   |   |
|           | $\begin{array}{r} \hline (internal r, 16.4(6) \\ 13.1(4) \\ 13.1(4) \\ 13.1(4) \\ 13.1(4) \\ 13.1(4) \\ 23.6(4) \\ 22.6(4) \\ 21.9(5) \\ 21.2(4) \\ 21.2 \\$ | i (internal resistive I16.4(6) $300$ 13.1(4)106013.1(4)105013.1(4)108023.6(4) $300$ 22.6(4)112021.9(5)123021.2(4)126021.2126021.2126021.2133021.2143021.2145033.3(2) $300$ 32.7(5)1030 $30.6(5)$ 1190 $30.2(5)$ 124029.3(1.0)1300**29.3(1.1)1370**29.0(6)144028.4(1)147028.6(8)150028.6158028.6162041.0(5)30043.7(6)113043.6(6)119043.3(6)125042.8(6)131042.3(5)138041.9(5)146041.5(7)156041.2(7)160041.5(8)168041.6(6)174042.8(5)194042.8205042.8209042.82160 | i (internal resistive heating)         16.4(6)       300       bcc+hcp         13.1(4)       1060       fcc         13.1(4)       1050       fcc         13.1(4)       1080       fcc         23.6(4)       300       hcp         22.6(4)       1120       hcp         21.9(5)       1230       hcp+fcc         21.2(4)       1260       hcp+fcc         21.2       1260       hcp+fcc         21.2       1260       hcp+fcc         21.2       1330       fcc         21.2       1360       fcc         21.2       1430       fcc         21.2       1430       fcc         21.2       1430       fcc         33.3(2)       300       hcp         30.6(5)       1190       hcp         30.2(5)       1240       hcp         29.3(1.0)       1300***       hcp+fcc         29.3(1.1)       1370**       hcp+fcc         29.3(6)       1440       hcp+fcc         28.6       1580       fcc         28.6       1500       hcp         43.6(6)       1190       hcp | (internal resistive heating)16.4(6)300bcc+hcp20.823(53)13.1(4)1060fcc+hcp(small)21.660(0)13.1(4)1050fcc23.6(4)300hcp20.265(27)22.6(4)1120hcp20.816(3)21.9(5)1230hcp+fcc20.955(3)21.2(4)1260hcp+fcc-21.21270hcp+fcc-21.21330fcc-21.21360fcc-21.21430fcc-21.21450fcc-33.3(2)300hcp19.629(11)32.7(5)1030hcp20.312(9)29.3(1.0)1300**hcp+fcc20.47(7)30.2(5)1240hcp20.312(9)29.3(1.1)1370**hcp+fcc20.524(14)28.4(1)1470hcp+fcc20.595(5)28.6(8)1500hcp+fcc20.595(5)28.61580fcc41.0(5)300hcp19.198(26)43.7(6)1130hcp19.392(13)43.6(6)1190hcp19.527(9)42.8(5)1380hcp19.592(7)41.9(5)1460hcp19.637(14)41.4(7)1450hcp19.637(14)41.4(7)1450hcp19.777(17)41.5(8)1680hcp19.736(13)41.5(7)1560hcp19.736(13)41.5(7)1560hcp19. | (internal resistive heating)16.4(6)300bcc+hcp20.823(53)2.4640(18)13.1(4)1060fcc+hcp(small)21.660(0)2.4855(0)13.1(4)1050fcc13.1(4)1080fcc23.6(4)300hcp20.265(27)2.4394(14)22.6(4)1120hcp20.816(3)2.4554(2)21.9(5)1230hcp+fcc20.955(3)2.4596(2)21.2(4)1260hcp+fcc21.21330fcc21.21330fcc21.21330fcc21.21360fcc21.21450fcc21.21430fcc21.21430fcc21.21450fcc21.21450fcc33.3(2)300hcp20.016(10)2.4280(5)30.6(5)1190hcp20.247(7)2.432(4)30.2(5)1240hcp20.312(9)2.4361(5)29.3(1.0)1300**hcp+fcc20.559(5)2.4465(3)28.6(8)1500hcp+fcc20.595(5)2.4465(3)28.6(8)1500hcp+fcc20.595(5)2.4465(3)28.6(8)1500hcp19.473(10)2.4028(7)43.3(6)1250hcp19.473(10)2.4028(7)43.3(6)1250hcp< | Internal resistive heating)16.4(6)300bcc+hcp20.823(53)2.4640(18)3.9604(83)13.1(4)1060fcc+hcp(small)21.660(0)2.4855(0)4.04853(0)13.1(4)1080fcc23.6(4)300hcp20.265(27)2.4394(14)3.9323(24)22.6(4)1120hcp20.816(3)2.4554(2)3.9866(3)21.9(5)1230hcp+fcc21.036(0)2.4556(0)4.0152(0)21.21260hcp+fcc21.21270hcp+fcc21.21330fcc21.21330fcc21.21430fcc21.21450fcc33.3(2)300hcp20.016(10)2.4280(5)3.9206(10)30.6(5)1190hcp20.247(7)2.4342(4)3.9456(7)30.2(5)1240hcp20.312(9)2.4361(5)3.9521(9)29.3(1.1)1370**hcp+fcc20.524(14)2.4443(7)3.9650(13)29.4(11)1470hcp+fcc20.524(14)2.4445(7)3.9663(3)29.4(6)1440hcp+fcc20.559(5)2.4445(3)3.9663(3)28.6(8)1500hcp+fcc20.595(25)2.4445(3)3.9663(3)28.6(8)1500hcp+fcc20.595(5)2.4465(3)3.8826(24)43.7(6)1130hcp19.3 | (internal resistive heating)16.4(6)300bcc+hcp20.823(53)2.4640(18)3.9604(83)13.1(4)1060fcc+hep(small)21.660(0)2.4855(0)4.04853(0)43.643(14)13.1(4)1080fcc43.615(19)13.1(4)1080fcc43.6683(19)23.6(4)300hcp20.265(27)2.4394(14)3.9323(24)22.6(4)1120hcp20.816(3)2.4554(2)3.9866(3)21.9(5)1230hcp+fcc20.955(3)2.4596(0)4.0152(0)42.343(38)21.21260hcp+fcc42.330(5)21.21260hcp+fcc42.330(5)21.21260hcp+fcc42.330(5)21.21360fcc42.720(21)2.121430fcc21.21430fcc42.720(21)2.121430fcc21.21430fcc42.720(21)2.122.1450fcc22.3(1.0)130**hcp20.016(10)2.4280(5)3.9206(10)30.6(5)1190hcp20.247(7)2.4342(4)3.9456(7)30.2(5)1240hcp20.312(9)2.4361(5)3.9521(9)29.3(1.0)130**hcp+fcc20.524(14)2.4437(15)3.9592(13)29.3(1.0)130**hcp+fcc20.5295(5)2.4465(3)3.9663(3)29.3(1.0)130**hcp+fcc20.5295(5)2.4465(3)3.9661(14)< |

# Table 1. Experimental conditions and results

| 46.0(3)                             | 300   | hcp  | 18.938(15) 2.3856(9) 3.8426(15)  |
|-------------------------------------|---|--|--|
| 50.0(6)                             | 1050  | hcp  | 19.020(12) 2.3882(7) 3.8507(11)  |
| 50.3(6)                             | 1120  | hcp  | 19.036(10) 2.3885(5) 3.8530(9)   |
| 50.6(5)                             | 1230  | hcp  | 19.068(7) 2.3890(4) 3.8579(6)  |
| 50.9(5)                             | 1350  | hcp  | 19.108(4) 2.3897(2) 3.8637(4)  |
| 52.0(5)                             | 1530  | hcp  | 19.124(5) 2.3901(3) 3.8657(5)  |
| 52.6(8)                             | 1620  | hcp  | 19.139(17) 2.3904(9) 3.8676(16)  |
| 52.8(1.3)                           | 1750  | hcp  | 19.186(44) 2.3914(24) 3.8739(41)   |
| 75.8(1.2)                           | 300   | hcp  | 17.705(43) 2.3355(25) 3.7481(42)   |
|                                     |   | -  | 18.135(48) 2.3563(22) 3.7716(72)   |
|                                     |   | -  | 18.133(42) 2.3553(19) 3.7745(62)   |
| · · ·                               |   | -  | 18.151(35) 2.3551(16) 3.7789(53)   |
|                                     |   | -  | 18.194(16) 2.3549(7) 3.7886(24)  |
|                                     |   | -  | 18.323(10) 2.3596(5) 3.8000(16)  |
|                                     |   | -  | 18.247(14) 2.3540(6) 3.8025(22)  |
| . ,                                 |   | -  | 18.301(21) 2.3552(8) 3.8096(36)  |
|                                     |   | -  | 18.347(24) 2.3570(9) 3.8134(42)  |
|                                     |   | -  | 18.387(25) 2.3585(9) 3.8169(43)  |
|                                     |   | -  | 18.415(21) 2.3595(8) 3.8193(36)  |
| . ,                                 |   | -  | 18.410(17) 2.3600(8) 3.8169(26)  |
| 71.0(7)                             | 2020  | hcp  | 18.444(9) 2.3616(4) 3.8187(14)   |
| 5<br>Si (laser hei                  | ating)  |  |  |
|                                     |   | hcp  | 20.321(14) 2.4382(8) 3.9471(14)  |
|                                     |   | -  | 20.539(10) 2.4438(5) 3.9713(9)   |
|                                     |   | -  | 20.544(10) 2.4440(5) 3.9717(9)   |
| . ,                                 |   | -  | 20.602(9) 2.4457(5) 3.9773(8)  |
| . ,                                 |   | -  | 20.612(9) 2.4459(5) 3.9784(9)  |
| . ,                                 |   | -  | 20.671(9) 2.4473(3) 3.9854(12)   |
| . ,                                 |   | -  | 20.683(36) 2.4465(15) 3.9900(51)   |
| . ,                                 |   | -  | 20.695(60) 2.4507(25) 3.9788(82)   |
|                                     |   | -  | 20.777(29) 2.4513(10) 3.9927(44)   |
| 29.5                                | 1850  | hcp+fcc  | - - $-$ 42.569(51) one peak for hcp  |
| 62,1(3,5)                           | 300   | hen  | 18.308(55) 2.3616(21) 3.7904(93)   |
| . ,                                 |   | -  | 18.620(32) 2.3702(10) 3.8271(57)   |
| 64.0(3.7)                           | 1600  | hcp  | 18.664(20) 2.3702(6) 3.8364(37)  |
| 0.1.0(0.7)                          |   | hcp  | 18.717(11) 2.3703(3) 3.8468(19)  |
|                                     | 1690  |  | 10,11,111,20,10,010,01,00,010,010,010,01   |
| 63.7(3.6)                           | 1690<br>1930  | -  |  |
| 63.7(3.6)<br>66.0(5.3)              | 1930  | hcp  | 18.711(77) 2.3733(8) 3.8360(157)   |
| 63.7(3.6)<br>66.0(5.3)<br>59.5(6.6) | 1930<br>2040  | hcp<br>hcp   | 18.711(77) 2.3733(8) 3.8360(157)<br>19.064(158) 2.3848(13) 3.8706(318)   |
| 63.7(3.6)<br>66.0(5.3)              | 1930  | hcp  | 18.711(77) 2.3733(8) 3.8360(157)   |
|                                     | $\begin{array}{c} 50.0(6)\\ 50.3(6)\\ 50.6(5)\\ 50.9(5)\\ 52.0(5)\\ 52.0(5)\\ 52.6(8)\\ 52.8(1.3)\\ \hline 75.8(1.2)\\ 70.1(1.6)\\ 70.7(1.5)\\ 70.8(1.3)\\ 70.5(8)\\ 68.7(7)\\ 71.0(8)\\ 71.0(1.0)\\ 70.8(1.0)\\ 70.6(1.1)\\ 70.6(9)\\ 71.2(9)\\ 71.0(7)\\ \hline Si (laser healer $ | 50.0(6)1050 $50.3(6)$ 1120 $50.6(5)$ 1230 $50.9(5)$ 1350 $52.0(5)$ 1530 $52.6(8)$ 1620 $52.8(1.3)$ 1750 $75.8(1.2)$ 300 $70.1(1.6)$ 1080 $70.7(1.5)$ 1140 $70.8(1.3)$ 1200 $70.5(8)$ 1300 $68.7(7)$ 1460 $71.0(8)$ 1510 $71.0(1.0)$ 1650 $70.8(1.0)$ 1760 $70.6(9)$ 1910 $71.2(9)$ 1960 $71.0(7)$ 2020 <b>55</b> (laser heating) $24.0(1.3)$ 300 $28.8(2.3)$ 1320 $30.5(2.5)$ 1520 $29.6(2.4)$ 1500 $28.2(2.3)$ 1360 $28.0(2.3)$ 1430 $29.5(2.8)$ 1610 $29.5(2.8)$ 1750 $29.5$ 1850 $62.1(3.5)$ 300 $64.0(3.9)$ 1480 | 50.0(6) $1050$ hcp $50.3(6)$ $1120$ hcp $50.6(5)$ $1230$ hcp $50.9(5)$ $1350$ hcp $52.0(5)$ $1530$ hcp $52.0(5)$ $1530$ hcp $52.6(8)$ $1620$ hcp $52.8(1.3)$ $1750$ hcp $75.8(1.2)$ $300$ hcp $70.1(1.6)$ $1080$ hcp $70.7(1.5)$ $1140$ hcp $70.8(1.3)$ $1200$ hcp $70.5(8)$ $1300$ hcp $71.0(8)$ $1510$ hcp $71.0(8)$ $1510$ hcp $70.6(1.1)$ $1840$ hcp $70.6(9)$ $1910$ hcp $71.0(7)$ $2020$ hcp $71.0(7)$ $2020$ hcp $70.5(2.5)$ $1520$ hcp $70.6(2.4)$ $1500$ hcp $70.5(2.5)$ $1520$ hcp $29.6(2.4)$ $1500$ hcp $29.5(2.8)$ $1610$ |

\* The uncertainty in temperature in the resistive and laser heating are typically 50 K and 10%, respectively.

\*\* Temperature uncertainty is as large as  $\pm 100$  K.

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figure 1

Always consult and cite the final, published document. See http://www.minsocam.org or GeoscienceWorld



figure 2



figure 3



figure 4