

Phase relations on the K_2CO_3 - CaCO_3 - MgCO_3 join at 6 GPa and 900–1400 °C: Implications for incipient melting in carbonated mantle domains

ANTON SHATSKIY^{1,2,*}, KONSTANTIN D. LITASOV^{1,2}, YURI N. PALYANOV^{1,2}, AND EIJI OHTANI^{1,3}

¹V.S. Sobolev Institute of Geology and Mineralogy, Russian Academy of Science, Siberian Branch, Novosibirsk 630090, Russia

²Novosibirsk State University, Novosibirsk 630090, Russia

³Department of Earth and Planetary Material Science, Tohoku University, Sendai 980-8578, Japan

ABSTRACT

To constrain the ternary K_2CO_3 - CaCO_3 - MgCO_3 T - X diagram at 6 GPa and to expand upon the known K-Mg, K-Ca, and Ca-Mg binary systems we have carried out multi-anvil experiments along the K_2CO_3 - $\text{Ca}_{0.5}\text{Mg}_{0.5}\text{CO}_3$ join. The diagram has primary phase fields for K_2CO_3 , $\text{K}_2\text{Mg}(\text{CO}_3)_2$, $\text{K}_2\text{Ca}_{0.1-0.5}\text{Mg}_{0.9-0.5}(\text{CO}_3)_2$, $\text{K}_4\text{CaMg}(\text{CO}_3)_4$, Ca-magnesite, and dolomite. The system has two liquidus minima near 1000 °C. At one minimum, a liquid with the composition of 36 K_2CO_3 :64($\text{Ca}_{0.65}\text{Mg}_{0.35}$) CO_3 is in equilibrium with three phases: Ca-magnesite, $\text{K}_2\text{Ca}_{0.1-0.5}\text{Mg}_{0.9-0.5}(\text{CO}_3)_2$, and $\text{K}_6\text{Ca}_2(\text{CO}_3)_5$. The other minimum, a liquid with the composition of 62 K_2CO_3 :38 $\text{Ca}_{0.72}\text{Mg}_{0.28}\text{CO}_3$ is in equilibrium with K_2CO_3 , $\text{K}_4\text{CaMg}(\text{CO}_3)_4$, and $\text{K}_6\text{Ca}_2(\text{CO}_3)_5$. At 900 °C, the ternary diagram contains two- and three-phase regions with Ca-magnesite, aragonite, $\text{K}_2\text{Ca}_3(\text{CO}_3)_4$, $\text{K}_2\text{Ca}(\text{CO}_3)_2$, $\text{K}_6\text{Ca}_2(\text{CO}_3)_5$, K_2CO_3 , $\text{K}_2\text{Ca}_{0.1-0.5}\text{Mg}_{0.9-0.5}(\text{CO}_3)_2$ solid solution, $\text{K}_2\text{Mg}_{0.9}\text{Ca}_{0.1}(\text{CO}_3)_2$, and $\text{K}_4\text{CaMg}(\text{CO}_3)_4$. We also expect an existence of primary phase fields for $\text{K}_6\text{Ca}_2(\text{CO}_3)_5$, $\text{K}_2\text{Ca}_3(\text{CO}_3)_4$ and aragonite.

We suggest that extraction of K from silicate to carbonate components should decrease the minimum melting temperature of dry carbonated mantle rocks up to 1000 °C at 6 GPa and yield ultrapotassic Ca-rich dolomite melt containing more than 10 mol% K_2CO_3 . As temperature increases above 1200 °C the melt evolves toward an alkali-poor, dolomitic liquid if the bulk molar CaO/MgO ratio > 1 , or toward K-Mg-rich carbonatite if bulk $\text{CaO}/\text{MgO} < 1$. The majority of compositions of carbonatite inclusions in diamonds from around the world fall within the magnesite primary field between the 1300 and 1400 °C isotherms. These melts could be formed by partial melting of magnesite-bearing peridotite or eclogite with bulk $\text{Ca/Mg} < 1$ at temperatures ≤ 1400 °C. A few compositions revealed in the Ebelyakh and Udachnaya diamonds (Yakutia) fall within the dolomite primary field close to the 1200 °C isotherm. These melts could be formed by partial melting of dolomite-bearing rocks, such as carbonated pelite or eclogite with bulk $\text{Ca/Mg} > 1$ at temperatures ≤ 1200 °C.

Keywords: Deep earth, high pressure, high temperature, melts, phase relations