Confirmation of the terrestrial occurrence of orthopyroxene with space group $P2_1ca$

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ABSTRACT

Based on 0kl electron diffraction patterns of two terrestrial orthopyroxene samples, all k-odd reflections violating a b glide are present. Such reflections are not caused by the overlap of diffractions from higher levels of orthopyroxene itself, or from the presence of C2/c clinopyroxene exsolution lamellae or G.P. zones, by diffuse streaks along a^* , or by the multiple diffraction effects. Observations of a hypersthene sample with a TEM equipped with a heating stage indicate that a phase transformation is produced in which all 0kl diffraction spots with k-odd systematically became unobservable when the sample was heated to approximately 900 °C. The existence of a terrestrial orthopyroxene with space group $P2_1ca$ is therefore confirmed.

INTRODUCTION

The space group of orthopyroxenes in the enstatiteferrosilite series has been regarded as being Pbca under normal pressure and temperature formation conditions from the time that the crystal structure of hypersthene was determined by Warren and Modell (1930). When Thompson (1970) investigated aspects of the amphibole structure, he indicated that orthorhombic pyroxenes, which obey the parity rule, should have space group $P2_1ca$. Papike et al. (1973), Matsumoto (1974), Pannhorst (1979), Law and Whittaker (1980), and Chisholm (1981, 1982) concluded that it is possible for orthopyroxene with space group $P2_1ca$ to exist. Smyth (1974) found k-odd reflections violating a b glide on 0kl precession photographs of bronzite from lunar troctolitic granulite. He interpreted the space group to be $P2_1ca$. Subsequently, Steele (1975), Harlow et al. (1979), Evensen et al. (1979), and Harlow (1980) reported space group $P2_1ca$ for one enstatite, four bronzites, and two hypersthenes. They occur in a lunar norite, a feldspar cumulate eucrite, the silicate inclusions of an iron meteorite, three achondrites, and one mesosiderite, respectively. Krstanović (1975) reported that there are $P2_1ca$ domains in enstatite occurring in harzburgite at Goleš Mountain, Yugoslavia. These domains are intimately intergrown with domains with space group Pbca that are richer in Mg. Harlow (1980) presented evidence that a terrestrial enstatite sample (AMNH G40403) had space group $P2_1ca$. This interpretation was subsequently reconsidered (Sasaki et al., 1984).

In a discussion of Smyth's (1974) report of $P2_1ca$ orthopyroxene, Nord et al. (1976) concluded that the appearance of diffractions violating the *b* glide was the result of multiple diffraction as is common in electron diffraction patterns. Sasaki and Matsumoto (1977) and Sasaki et al. (1981, 1984) subsequently proposed that the 0kl diffractions that appeared to violate the b glide were not caused by the lack of a b-glide plane. Recently, Smyth and Swope (1990) described further studies of lunar orthopyroxene 76535 and suggested a new explanation regarding the origin of the b-glide violations observed in that orthopyroxene.

Luo et al. (1987) ascertained that all of the k-odd diffraction spots are present on 0kl electron diffraction patterns of a hypersthene megacryst from Black Hill of Chengde, Hebei Province, China, thus confirming that pyroxene with space group $P2_1ca$ indeed exists. Luo et al. (1989) also derived the most probable structure topology, as shown in Figure 1. They indicated that it is possible to find domains related by twin plane (100). In such a structure, there are four kinds of tetrahedral chains, two M1 sites, and two M2 sites, all of which are symmetrically distinct. The ratio of O-rotated to S-rotated tetrahedral chains is 3:1, and the ++-- repeat sequence of skew of octahedral strips along the a axis is the same as in orthopyroxene with space group Pbca. Such a structure obeys the parity rule suggested by Thompson (1970) and has a as a polar axis. If there are two neighboring domains whose a axes are in reversed orientations, a twin with {100} as twin plane could be formed. The structures of such domains are fully continuous at the join. Therefore, it would be difficult to distinguish such a twin from a single crystal in morphology, optical properties, and diffraction patterns. It can also be seen from Figure 1 that the twin plane serves as a *b*-glide parallel to $\{100\}$, causing the intensity of 0kl diffractions with k-odd to be weakened or even to become almost extinct. In this case, it would be more difficult to distinguish orthopyroxene with space group $P2_1ca$ from one with space group Pbca. There are indications that such twinning exists in both the orthopyroxenes from Black Hill and Bamble, Norway (Luo et al., in preparation). We believe that this kind of



Fig. 1. Schematic I-beam model of the most probable structure and the structural relationship between the two domains in a twin on {100} of $P2_1ca$ orthopyroxene. The polar a axes are oppositely oriented in the twin domains. The skew of octahedral strips is symbolized with arrows with opposite directions instead of the conventional + and - to allow direct comparison with the skew of octahedral strips in the twin domains. The light and heavy arrows indicate octahedral strips that are symmetrically nonequivalent (after Luo et al., 1989).

twin is the cause of ambiguities regarding the space group of orthopyroxene. In this paper, the relations for the diffraction intensities violating the b glide in orthopyroxene are discussed.

SAMPLES AND EXPERIMENTS

Sample 1 is a hypersthene megacryst (average composition: $En_{56}Fs_{41}Wo_3$; a = 18.310, b = 8.900, c = 5.223 Å) from the Presinian noritic pegmatite from Black Hill, Chengde, Hebei Province, China. The sample contains exsolution lamellae of augite (approximately $En_{45}Fs_{30}Wo_{25}$; a = 9.76, b = 8.91, c = 5.22 Å, $\beta = 105.2^{\circ}$) and G.P. zones parallel to {100}. It always occurs as anhedral crystals over 2 cm in size, bronze in color, and with {100} parting. Sample 2 is an enstatite megacryst ($En_{96.6}Fs_{3.4}$; a = 18.224, b = 9.815, c = 5.179 Å) from Bamble (collection of Department of Earth Sciences, Nanjing University). It is green in color and over 10 cm in length, with {100} parting.

The samples were thinned by ion milling and coated with a C film. Selected area diffraction patterns were obtained using a JEM-200CX transmission electron microscope (TEM) operated at 200 kV, with camera lengths of 137 and 55 cm. Before the samples were examined by TEM, they were studied by polarized light microscopy, X-ray powder diffraction, and precession single crystal X-ray diffraction. Long exposure (120 h) [100] X-ray precession photographs revealed that 0kl diffraction spots with k-odd, such as 011, 031, 033, 051, 054, and 074,



Fig. 2. A [100] electron diffraction pattern of the Chengde hypersthene (sample 1). The 0kl diffraction spots with k = 2n + 1 systematically appear but are weaker than those with k = 2n.

are clearly observable, although they are weaker than those with k-even.

RESULTS AND DISCUSSION

Smyth (1974) first described the existence of pyroxene with space group P_{2_1ca} in a lunar sample. However, Nord et al. (1976), Sasaki and Matsumoto (1977), and Sasaki et al. (1981, 1984) questioned the existence of orthopyroxene with space group P_{2_1ca} for three reasons. They suggested that the presence of reflections incompatible with a *b* glide could be explained as (1) overlapping of diffractions from $C_{2/c}$ clinopyroxene exsolved on {100} of the host orthopyroxene, (2) diffuse streaks of orthopyroxene or clinopyroxene from levels separated from the 0kl net, or (3) multiple diffraction. Using our own experimental results, we discuss our interpretations of the orthopyroxenes as follows.

Relations for overlapping diffraction patterns of opx host and cpx lamellae

If the diffraction pattern of clinopyroxene {100} lamellae completely overlaps that of orthopyroxene with space group *Pbca*, cell dimensions of both phases must satisfy specific conditions because c^* of host and guest crystals are not parallel to each other. The reciprocal lattice translation of clinopyroxene that is approximately parallel to c^* of orthopyroxene is [T02]* (Sasaki et al., 1984). Cell dimensions of orthopyroxene and clinopyroxene can have values consistent with $S_{T02(cpx)}$ being almost parallel and equal to $S_{002(opx)}$ if the interfaces are coherent. Even so, the diffraction spots of clinopyroxene with space group C2/c that appear on the 0kl net of orthopyroxene, as pointed out by Sasaki et al. (1984), are limited to indices with k-odd, $l = \pm 2, \pm 6$, and so on, indexed on the



Fig. 3. A [100] electron diffraction pattern of Bamble enstatite (sample 2). The 0kl diffraction spots with k = 2n + 1 occur but are weaker than those with k = 2n.

cell of orthopyroxene. This is because the reciprocal net of clinopyroxene that is parallel to the 0kl net of the orthopyroxene is of hkl type with h + k = 2n and $l = \pm 2h$; therefore, when k is odd, l must be equal to 4n + 2. For instance, the allowed diffractions of clinopyroxene should be $\overline{2}04$, $\overline{4}08$, $\overline{1}12$, $\overline{3}16$, ..., and they would overlap with 004, 008, 012, 016, ... of orthopyroxene, respectively. Accordingly, it can be concluded that the 0kl spots with k-odd and $l \neq 4n + 2$ cannot be caused by the clinopyroxene lamellae. Selected area electron diffraction patterns of the 0kl net for samples 1 and 2 (Figs. 2 and 3) contain, in addition to k-even spots, k-odd spots (including those with $l \neq 4n + 2$) that are very sharp.

Intensity of 0kl reflections in relation to reflection shape

Differentiation of overlap from neighboring levels. Each reciprocal lattice point has a rodlike shape whose length increases with decreasing thickness of the sample. Nonzero level reflections can therefore overlap the zero level. However, our diffraction patterns show that the spots violating the b glide in the 0kl pattern of orthopyroxene are not caused by such overlap. Figure 4 is a [100] electron diffraction pattern of sample 1 that includes reflections from a higher-order Laue zone. The distance between levels is calculated to be 0.0546 Å⁻¹, which corresponds to $1/a^*$ of orthopyroxene (0.0549–0.0543 Å⁻¹). The higherorder Laue zone in Figure 4 must therefore correspond to reflections with indices 1kl. A gap is evident between the first and zero order Laue zones, also apparent in sample 2 (Fig. 5); this could not occur if 1kl reflections overlapped the zero level. Thus, none of the 0kl diffractions, including those with k-odd, result from overlap with higher-level reflections.

Streaking along a*. The [010] and [001] electron dif-



Fig. 4. A [100] electron diffraction pattern of Chengde hypersthene (sample 1) including reflections from the first level. It can be clearly seen that there is a gap between the two Laue zones.

fraction patterns of sample 1 (Figs. 6 and 7) show that, as observed by other authors, all diffraction spots are diffuse in the a* direction. However, the intensities of the diffuse streaks are not strong enough to form a sharp diffraction spot on the 0kl net. If the forbidden 0kl diffractions of a b glide were caused by diffuse streaks extending between the 1kl and $\overline{1}kl$ diffractions, there should have been diffraction spots between the first- and zero-order Laue zones in Figures 4 and 5, but there are no such reflections. Furthermore, h = 2n for hk0 reflections of *Pbca* and $P2_1ca$, and 1k0 and $\bar{1}k0$ diffractions are systematically extinct. Thus, 0k0 diffractions with k-odd would have to be caused by diffuse streaks between reflections 2k0 and $\overline{2}k0$. Such diffraction relations are highly unlikely because of the large separation between pairs of reflections. In fact, no streaking with intensity commensurate



Fig. 5. A [100] electron diffraction pattern of Bamble enstatite (sample 2) including reflections from the first-order Laue zone. There is a gap between reflections from the two Laue zones.

Fig. 6. A [010] electron diffraction pattern of Chengde hypersthene (sample 1).

with that of the 0k0 spots is observed in the [001] electron diffraction pattern (Fig. 7); formation of sharp 0k0 spots due to diffuse streaks between 2k0 and $\overline{2}k0$ diffractions is therefore not possible.

Multiple diffraction

Conditions for multiple diffraction. In order for multiple diffraction to occur, the indices of the secondary diffraction hkl and the two cooperating diffractions $h_1k_1l_1$ and $h_2k_2l_2$, which cause the diffraction hkl, should be $h = h_1 + h_2$, $k = k_1 + k_2$, $l = l_1 + l_2$. If the space group of orthopyroxene were *Pbca*, the nonextinct diffractions would have indices with values of k equal to 2n for 0kl diffractions. Obviously, the sum of even numbers cannot equal an odd number, so the k-odd diffractions on the 0kl pattern of the orthopyroxene cannot result from secondary diffractions originating from two 0kl diffractions of orthopyroxene with space group *Pbca*.

Multiple diffraction from higher levels. No matter whether the space group of orthopyroxene is Pbca or P2, ca, the following diffraction conditions will exist: h0l: l = 2n, no condition for h; 00l: l = 2n; hk0: h = 2n; h00: h = 2n; 0kl: no condition for l. Experimental observation shows that, as expected, all h0l diffractions with *l*-odd and hk0 diffractions with h-odd are extinct as shown by Figures 6 and 7. However, a few 00l diffractions with l-odd that appear on the [100] electron diffraction patterns (Figs. 2 and 3) are due to multiple diffraction for 0kl diffractions. The 00l diffractions with l-odd on the [010] pattern (Fig. 6) are extinct as a result of the absences of all h0l diffractions with l-odd. Similarly, because the hk0 diffractions with h-odd are extinct but h-odd indices are permitted for hol diffractions, the forbidden diffractions h00 with h-odd appear on the [010] pattern (Fig. 6) but not on the [001] pattern (Fig. 7). Therefore, at least in the [010] and [001] diffraction patterns, there are no

Fig. 7. A [001] electron diffraction pattern of Chengde hypersthene (sample 1). The hk0 diffraction spots with h + 2k = 4n + 2 occur but are weaker than those with h + 2k = 4n.

diffraction spots that are caused by multiple diffraction involving higher levels. Further experimentation (see below) shows that this also is true for the [100] pattern. In other words, the multiple diffractions all result from the diffractions for the zero level. We have already shown that the 0kl diffractions with k-even cannot produce any secondary 0kl diffractions with k-odd. Therefore, we conclude that the k-odd violations are not the result of multiple diffractions.

Distinguishing multiple diffraction. Multiple diffractions can be distinguished by tilting the sample and moving the source of multiple diffraction away from the Ewald sphere. For the [100] diffraction pattern, the forbidden diffractions of 00*l* become unobservable when the sample is tilted about c*. This indicates that these diffractions are caused by multiple diffraction. The 0k0 diffractions with either k-even or k-odd do not become unobservable when the sample is tilted about b*. To examine further whether the k-odd spots on 0kl diffraction patterns are caused by the multiple diffraction from higher levels, the (100) section of sample 1 was tilted about the reciprocal vector S_{032} , and the change in intensity for 032, an intense spot on the 0kl net, was observed. There was no perceivable change in its intensity. Seven selected-area diffraction patterns were obtained from $+2^{\circ}10'$ to $-3^{\circ}50'$ to confirm this observation. Therefore, the presence of 0kl diffractions with k-odd on the [100] pattern is not the result of multiple diffraction.

Extra extinction conditions

Chisholm (1981) has indicated, in his derivation of pyribole structure types, that in the case of projection down c, the point $\frac{1}{4}$, $\frac{1}{2}$ has the same environment as the origin 0,0 for pyroxene with space group $P2_1ca$, and this symmetrical relation gives rise to an additional diffrac-







Fig. 8. A [100] electron diffraction pattern of Chengde hypersthene (sample 1) when heated to more than 900 °C. The 0kl diffraction spots with k = 2n + 1 are systematically extinct, and the others are diffuse along **b***.



Fig. 9. A [001] electron diffraction pattern of Chengde hypersthene (sample 1) obtained from an area with abundant G.P. zones. A set of extra hk0 spots with h-odd is present.

tion condition 2h + k = 4n for hk0 diffractions. The additional diffraction condition 2h + k = 4n should instead be h + 2k = 4n, however. Nevertheless, in contrast to Chisholm's structure model of pyroxene with space group $P2_1ca$, the environment of the point $\frac{1}{4},\frac{1}{2}$ is different from, although similar to, that of the origin 0,0 in the model of $P2_1ca$ orthopyroxene (Fig. 1) proposed by Luo et al. (1989). Thus, it is expected that there would be no extra extinction condition for hk0 diffractions, but the intensities of the hk0 diffractions with h + 2k = 4n + 2 would be systematically weaker than those with h + 2k = 4n. This relation is evident in Figure 7.

Heating experiment

Heating of sample 1, carried out by means of a heating stage in the TEM (Xu et al., 1989), showed that there was no change in the [100] diffraction pattern until the sample was heated to approximately 900 °C. The 0kl spots (including 0k0 spots) with k-odd became systematically unobservable after the sample was held at 900 °C for 1 h. The remaining spots were streaked along **b*** (Fig. 8). This suggests that a phase transformation may occur at approximately 900 °C, and the observation of all 0kl diffractions at room temperature depends on the intrinsic symmetry of crystals and not on factors such as multiple diffraction.

Extra diffractions resulting from G.P. zones

Recently, in a further study of the origin of space-group violations in lunar orthopyroxene 76535, Smyth and Swope (1990) showed that forbidden diffractions for a b glide are not caused by diffractions from clinopyroxene lamellae, by diffuse streaks, or by multiple diffraction ef-

fects, based on automated four-circle X-ray diffractometer data. This corresponds with our above-described observations. They also suggested that the most reasonable explanation for the observed space-group violations is that both the a- and b-glide violations result from ordering of Ca into {100} G.P. zones.

In the electron diffraction patterns of sample 1, diffractions violating not only the b but also the a glide are present. Figure 9 is an [001] diffraction pattern that was obtained from an area that has abundant G.P. zones. Besides the streaks along a^* , extra hk0 spots with h-odd appear, but their intensities are much weaker than those with h-even. The dark-field image formed using the 300 diffraction shows that the extra diffractions violating the a glide result from G.P. zones. Since the G.P. zones are much smaller than the thickness of the selected area of sample and the diameter of the electron beam, the volume ratio of G.P. zones to host for any sample orientation is constant. However, if the intensities of 0k0 diffractions in different patterns are taken as a reference, it is found that the intensity ratio of hk0 (h-odd) to 0k0 in the [001] pattern (Fig. 9) is much smaller than that of 0kl(k-odd) to 0k0 in the [100] patterns (Fig. 2). In other words, the intensities of hk0 (h-odd) diffractions are obviously incommensurate with those of 0kl (k-odd) ones. These relations imply that the *b*-glide violations are caused by another factor. The dark-field image formed using the 010 diffraction (Fig. 10) indicates that the 010 diffraction intensity is due to the host orthopyroxene and not to G.P. zones or clinopyroxene lamellae. Therefore, the appearance of 0kl (including 0k0) diffractions with k-odd should mainly be attributed to the host orthopyroxene and not to the G.P. zones.



Fig. 10. Dark-field image of the Chengde hypersthene (sample 1) formed using the 010 diffraction. Only the host orthopy-roxene is in bright contrast. The horizontal dark strip is an augite lamella.

CONCLUSIONS

Based on the diffraction relations for the orthopyroxene samples, the following conclusions can be drawn.

1. For the two samples studied, the diffractions violating a b glide on [100] electron diffraction patterns are true violations. Their presence is not caused by overlap of diffractions from higher levels of orthopyroxene, from augite exsolution lamellae, or by diffuse streaks along a^* , and they are neither the result of multiple diffraction nor contributed by G.P. zones. The presence of diffractions violating a b glide is caused by the inherent symmetry of the crystal. Observations for sample 1 during heating tests support this view.

2. The conditions of diffraction for the two orthopyroxenes studied are: hkl: no condition; hk0: h = 2n; 0kl: no condition; h0l: l = 2n; h00: (h = 2n); 0k0: no condition; 00l: (l = 2n). The corresponding possible space groups are *Pmca* or $P2_1ca$. However, the true space group is most likely $P2_1ca$ (Luo et al., 1987) because there is no mirror plane in the crystal structure of any pyroxene. Therefore, we conclude that orthopyroxene with space group $P2_1ca$ can exist on the Moon, the Earth, and in meteorites.

ACKNOWLEDGMENTS

We are grateful to P.H. Ribbe for his encouragement, to S.W. Bailey, G.L. Nord, Jr., G.E. Harlow, J.R. Smyth, T.C. McCormick, and R.J. Swope for their critical reviews of the manuscript, particularly to G.L. Nord, Jr. for his helpful advice on sample orientation, to Fusheng Zhang and Jun Ma for help during TEM observations, and to Chengyi Lin for his helpful comments on the manuscript. This project was supported by the National Natural Sciences Foundation of China.

References cited

- Chisholm, J.E. (1981) Pyribole structure types. Mineralogical Magazine, 44, 205–216.
- (1982) Lowering of symmetry in pyribole structures. Mineralogical Magazine, 45, 25–34.
- Evensen, N.M., Hamilton, P.J., Harlow, G.E., Klimentidis, R., O'Nions, R.K., and Prinz, M. (1979) Silicate inclusions in Weekeroo Station: Planetary differentiates in an iron meteorite. Lunar and Planetary Science, 10, 367–377.
- Harlow, G.E. (1980) Low orthopyroxene: Achondritic abundance and planetary significance. Lunar and Planetary Science, 11, 396–397.
- Harlow, G.E., Nehru, C.E., Prinz, M., Taylor, G.J., and Keil, K. (1979) Pyroxenes in Serra de Magé: Cooling history in comparison with Moama and Moore County. Earth and Planetary Science Letters, 43, 173– 181.
- Krstanović, I. (1975) X-ray diffraction studies of orthoenstatite (abs.). Acta Crystallographica, A31, 73.
- Law, A.D., and Whittaker, E.J.W. (1980) Rotated and extended model structures in amphiboles and pyroxenes. Mineralogical Magazine, 43, 565-574.
- Luo, Gufeng, Xu, Huifang, Xue, Jiyue, Hu, Meisheng, Fan, Chenggao, and Zhang, Shuyuan (1987) The first discovery of terrestrial orthopyroxene with space group P2₁ca. Kexue Tongbao (Chinese Science Bulletin), 19, 1483–1486 (in Chinese).
- Luo, Gufeng, Lin, Chengyi, and Xu, Huifang (1989) The possible topologic structure types of orthopyroxene with space group $P2_1ca$. Acta Mineralogica Sinica, 9, 1–8 (in Chinese).
- Matsumoto, T. (1974) Possible structure type derived from *Pbca*-orthopyroxene. Mineralogical Journal, 7, 374–384.
- Nord, G.L., Jr., Ross, M., and Huebner, J.S. (1976) Lunar troctolite 76535: Mineralogical investigation. Lunar and Planetary Science, 7, 628-630.
- Pannhorst, W. (1979) Structural relations between pyroxenes. Neues Jahrbuch für Mineralogie Abhandlungen, 135, 1–17.
- Papike, J.J., Prewitt, C.T., Sueno, S., and Cameron, M. (1973) Pyroxenes: Comparison of real and ideal structural topologies. Zeitschrift f
 ür Kristallographie, 138, 254–273.
- Sasaki, S., and Matsumoto, T. (1977) Multiple diffraction in orthopyroxenes: Re-examinations of the space group. Proceedings of Japan Academy, 53(B), 84–89.
- Sasaki, S., Matsumoto, T., and Sawada, C. (1981) The influence of multiple diffraction on the space group determination of orthopyroxene, spodumene, low omphacite, and pigeonite. Physics and Chemistry of Minerals, 7, 260-267.
- Sasaki, S., Prewitt, C.T., and Harlow, G.E. (1984) Alternative interpretation of diffraction patterns attributed to low (P2₁ca) orthopyroxene. American Mineralogist, 69, 1082–1089.
- Smyth, J.R. (1974) Low orthopyroxene from a lunar deep crustal rock: A new pyroxene polymorph of space group P2₁ca. Geophysical Research Letters, 1, 27–29.
- Smyth, J.R., and Swope, R.J. (1990) The origin of space group violations in a lunar orthopyroxene. Physics and Chemistry of Minerals, 17, 438– 443.
- Steele, I.M. (1975) Mineralogy of lunar norite 78235: Second lunar occurrence of P2₁ca pyroxene from Apollo 17 soils. American Mineralogist, 60, 1086–1091.
- Thompson, J.B., Jr. (1970) Geometrical possibilities for amphibole structures: Model biopyriboles. American Mineralogist, 55, 292–293.
- Warren, B., and Modell, D.I. (1930) The structure of enstatite MgSiO₃. Zeitschrift für Kristallographie, 75, 1–14 (not seen; extracted from Rockforming minerals, vol. 2A: Single-chain silicates, 2nd edition, 1978).
- Xu, Huifang, Luo, Gufeng, and Hu, Meisheng (1989) Study on the phase transformation of orthopyroxene-P2₁ca. Acta Petrologica et Mineralogica, 8, 188–192 (in Chinese with English abstract).

MANUSCRIPT RECEIVED APRIL 11, 1990

MANUSCRIPT ACCEPTED AUGUST 18, 1991