

Three-layer monoclinic lepidolite from Tørdal, Norway

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Abstract

A 3-layer monoclinic lepidolite from a pegmatite at Tørdal, Norway, previously described as 2-layer orthorhombic, has been identified as the $3M_2$ polytype. The space group is $C2$ with $a = 5.239(2)$, $b = 9.070(3)$, $c = 29.886(5)$ Å, and $\beta = 92.58(2)^\circ$. The X-ray powder pattern differs in detail from those of the conventional mica polytypes. Electron microprobe analysis gives SiO_2 48.78, Al_2O_3 23.87, FeO 1.39, MgO 0.02, MnO 4.61, TiO_2 0.07, K_2O 9.88, Na_2O 0.20, BaO 0.03, F 6.21 weight percent for the $3M_2$ flakes and SiO_2 53.41, Al_2O_3 21.43, FeO 0.06, MgO 0.02, MnO 1.82, TiO_2 0.06, K_2O 9.85, Na_2O 0.19, BaO 0.03, F 7.63 weight percent for associated $2M_1$ lepidolite flakes.

Heinrich *et al.* (1953) described a 3-layer monoclinic lepidolite from a pegmatite at Skuleboda, Sweden. Neumann *et al.* (1957) and Christie (1961) have described other lepidolite specimens from pegmatites at Varuträsk, Sweden, and Tørdal, Norway, that give X-ray powder patterns similar to that of the Skuleboda specimen. This note reports a more detailed study of the Tørdal material.

Single-crystal X-ray precession photographs of over 50 individual flakes of lepidolite from Tørdal showed three polytypic structures to be present— $1M$, $2M_1$, and $3M$. The $3M$ flakes represent the same material that was originally described as 2-layer orthorhombic by Christie (1961).

The $3M$ lepidolite has diffraction symmetry that could be described as $C2/m$, $C2$, or Cm . Single-crystal intensities were collected on an automated Syntex diffractometer and used for comparison with intensities calculated from the six possible 3-layer stacking sequences. Good agreement was found for only one model, which has ideal symmetry $C2$. Relative to a fixed initial set of axes, the intralayer shifts of $a/3$ in this model are directed in the sequence $-X_1$, $+X_2$, and $-X_3$ within the three successive layers. The resultant shift then is $a/3$ along $-X_2$ of the initial axes to give an ideal β angle of 93.4° .

The structure deduced above is identical to the ideal structure described as $3M_2$ by Ross *et al.* (1966) in their systematic study of possible mica polytypes. Comparison of Weissenberg photographs of the Tørdal specimen with those illustrated by Heinrich *et al.* (1953, Figs. 10–12) indicates that the latter specimens also have the $3M_2$ structure.

Table 1 lists the powder pattern of a pure sample of lepidolite- $3M_2$ from Tørdal. Indexing was achieved by direct comparison of the powder data with observed single-crystal intensities. The pattern differs in detail from those of the conventional mica structures and is characterized by the occurrence of several non-overlapping triplets of indices $\bar{1}1l$, $02l$, and $11l$. Least-squares refinement of the powder data gave cell dimensions $a = 5.239(2)$, $b = 9.070(3)$, $c = 29.886(5)$ Å, and $\beta = 92.58(2)^\circ$.

Table 2 lists electron microprobe analyses of flakes of $3M_2$ and $2M_1$ lepidolite from Tørdal. Although this method does not give Li or OH, the high F-contents and the individual oxide totals are characteristic of lepidolite. The $3M_2$ flake has less SiO_2 than the $2M_1$ flake but higher Al_2O_3 , MnO , and Fe (expressed as FeO) contents. The Li_2O contents must be approximately 4 to 5 weight percent for each flake,

Table 1. Indexed powder pattern of Tørdal lepidolite- $3M_2$

hkl	Int.	d(obs)	d(calc)	hkl	Int.	d(obs)	d(calc)
003	80	9.97 Å	9.952 Å	13.13	12	1.709	1.707
006	20	4.988	4.976	20.14	7	1.690	1.691
021	60	4.488	4.484	20.13			1.689
111			4.456	13.14	12	1.672	1.672
114	5	3.940	3.943	155	8 B	1.640	1.641
024	7	3.877	3.876	246			1.635
114	10	3.809	3.809	20.14	2	1.615	1.618
025	7	3.614	3.611	246			1.606
115	30	3.544	3.545	13.15	12	1.601	1.602
116	2	3.419	3.418	157			1.599
026	70 B	3.350	3.352	247	5	1.568	1.574
009			3.314	13.15			5
117	35	3.171	3.169	158	5	1.552	1.566
027	15	3.112	3.107	20.16			1.553
117	5	3.044	3.048	158	60	1.513	1.549
118	20	2.939	2.938	331			1.512
028	20	2.883	2.882	060	9	1.496	1.512
118	23	2.827	2.828	332			1.498
119	2	2.728	2.728	15.10	5	1.362	1.496
029	5	2.679	2.677	063			1.495
131	100	2.595	2.603	24.11	2	1.467	1.469
202			2.598	04.16			2
201			2.597	13.18	7	1.416	1.416
132			2.589	22.17			7
203	2.559	11.20	2	1.402	1.402		
202	2.558	04.17			2	1.388	
133	10 B	2.550	2.546	13.18	5	1.387	1.387
11.10	7	2.488	2.539	20.19			2
00.12			2.488	22.18	5	1.362	1.364
134	15	2.447	2.453	13.19			10
205			2.437	04.18	10	1.334	
204	15	2.425	2.435	13.19	2	1.322	1.323
135			2.418	20.20			2
135	5	2.364	2.378	22.18	25 B	1.309	1.314
206			2.360	13.20			25 B
205	5	2.349	2.358	261	25 B	1.309	1.310
136			2.339	260			25 B
136	2	2.293	2.296	400	25 B	1.309	1.308
207	2.277	262	25 B	1.307			
040	15	2.267	2.268	261	25 B	1.309	1.306
220			2.267	403			25 B
223	5	2.228	2.229	06.12	8 B	1.287	1.292
223	4	2.189	2.192	33.13			8 B
207			2.187	267	3	1.259	1.259
138	7	2.167	2.166	268			7
138	2	2.123	2.121	00.24	7	1.244	
045			2.120	424	13	B	1.228
209	2.101	356	13 B	1.229			
208	13	2.096	2.099	175	3 B	1.228	1.228
226			2.094	22.21			3 B
139	9	2.078	2.078	176	3 B	1.207	1.222
046			2.063	355			3 B
139	10	2.033	2.033	177	3 B	1.207	1.210
00.15	35	1.989	1.990	06.15			3 B
229	2	1.839	1.838				

Pattern taken with $\text{CuK}\alpha$ radiation in 114.59 mm diameter camera. Intensities estimated visually. d -values calculated on basis of $a = 5.239$, $b = 9.070$, $c = 29.886$ Å, and $\beta = 92.58^\circ$.

Table 2. Electron microprobe analyses of Tørdal lepidolites

	$3M_2$	$2M_1$
SiO_2	48.78 wt. %	53.41 wt. %
Al_2O_3	23.87	21.43
FeO	1.39	0.06
MgO	0.02	0.02
MnO	4.61	1.82
TiO_2	0.07	0.06
K_2O	9.88	9.85
Na_2O	0.20	0.19
BaO	0.03	0.03
F	6.21	7.63
Sums	95.06	94.50

Analyses run on ARL microprobe with an accelerating potential of 15 KV. The correction procedure devised by Bence and Albee (1968) has been followed, with incorporation of the alpha factors of Albee and Ray (1970).

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but lower for the $3M_2$ flake than for the $2M_1$ flake, as based on the observed F contents and oxide sums.

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