# A SODIAN STILBITE FROM ONIGAJŌ, MIÉ PREFECTURE, JAPAN, WITH SOME EXPERIMENTAL STUDIES CONCERNING THE CONVERSION OF STILBITE TO WAIRAKITE AT LOW WATER VAPOR PRESSURES

# KAZUO HARADA<sup>1</sup> AND KATSUTOSHI TOMITA, Institute of Geology and Mineralogy, Faculty of Science, Tokyo University of Education, Otsuka, Tokyo, Japan

#### ABSTRACT

A sodium-rich stilbite occurs in the drusy parts of a granite porphyry, Kii Peninsula, Japan. Chemical analysis gave SiO<sub>2</sub> 58.16, Al<sub>2</sub>O<sub>3</sub> 17.17, MgO 0.74, CaO 4.08, Na<sub>2</sub>O 3.60, K<sub>2</sub>O 0.85, H<sub>2</sub>O<sup>-</sup> 4.65 and H<sub>2</sub>O<sup>+</sup> 13.34. Sr, Ba, Rb and Cs were undetected. This corresponds to  $K_{0.52}$  Na<sub>3,30</sub> Ca<sub>2.07</sub> Mg<sub>0.57</sub> Al<sub>9.57</sub> Si<sub>26.55</sub> Or<sub>2</sub> · 28.87 H<sub>2</sub>O.

Unit-cell constants calculated from X-ray powder data are very close to those of normal stilbite; a=13.67, b=18.16, c=11.31 and  $\beta=129^{\circ}10'$ . The sodian stilbite is nearly colorless with brownish yellow tints, luster is vitreous to slightly pearly, and cleavage is {010} perfect. Hardness=3.5 and Sp. Gr.=2.18, 2V(-)=43.2,  $c \wedge X=9^{\circ}$ ,  $\alpha=1.482$ ,  $\beta=1.489$ ,  $\gamma=1.496$ ,  $\gamma-\alpha=0.014$ , all  $\pm 0.002$ .

Hydrothermal treatment of normal calcium stilbite at temperatures between 400°C and 200°C and at water vapor pressuers between 15 and 100 atm produced wairakite at higher pressures.

#### INTRODUCTION

Stilbite is a calcium zeolite usually containing a small amount of sodium in place of calcium, but none of them show Na>Ca in mole ratio. The writers found a sodium-rich stilbite in the drusy parts of a Miocene granite-porphyry. The present paper describes the mode of its occurrence and some mineralogical properties, together with the experimental data concerning the conversion of stilbite to wairakite at low water vapor pressures.

### OCCURRENCE

The granite-porphyry is a member of the Miocene Kumano Acidic Rocks (Kawano and Ueda, 1965) and widely distributed along southeastern Kii Peninsula, Japan (Sawatari, 1932; Harada, 1964; Aramaki and Hada, 1965; Aramaki, 1965).

The stilbite locality is about 0.5 mile north of Kumano city. The mineral is nearly colorless with a yellowish-brown tint, vitreous with a slightly pearly luster. It occurs as sheaf-like aggregates which contain small blades of about 5 mm in width, where phlogopitic biotite  $[(-) 2V=3^{\circ}\pm1^{\circ}, d_{001}=9.98$  Å, and  $\gamma=1.601\pm0.002]$  is associated with

<sup>1</sup> Present address: Chichibu Museum of Natural History, Nagatoro 1417, Nogamimachi, Chichibu-gun, Saitama Prefecture, Japan. minor laumontite and heulandite. The maximum size of aggregates is 3 cm across. Feldspars are lacking in the drusy parts of the granite-porphyry.

A zonal arrangement is apparent in the druses. Laumontite is abundant in the inner part (2-3 mm), heulandite in the middle part (2-3 mm) and stilbite in the outer part free from other zeolites (2-3 cm). According to Coombs (1961), laumontite is crystallized under the conditions of rather

	1	2	3	4	
$SiO_2$	56.16	68.90	55.39	55.53	
$\mathrm{TiO}_2$	none	0.73	none	none	
$Al_2O_3$	17.17	14.12	16.57	16.70	
$Fe_2O_3$	0.02	0.14	0.03	none	
FeO	0.11	3.31	none	none	
MnO	none	0.37	none	none	
MgO	0.74	1.39	0.29	0.15	
CaO	4.08	1.98	8.13	1,20	
$Na_2O$	3.60	3.32	1.10	7.00	
$K_{2}O$	0.85	3.31	0.05	0.03	
$H_2O(-)$	4.65 (105°C)	0.54 (105°C)	3.60 (105°C)	6.04 (110°)	
$H_2O(+)$	13.34 (105°C)	1.38	15.48 (105°C)	13.02 (110°)	
$P_2O_5$	0.01	0.09	none	0.01	
Total	100.74	99.58	100.64	99.68	

TABLE 1. CHEMICAL ANALYSES

1. Stilbite (sodium-rich) Onigajō, Mié Prefecture (Analyst, K. Harada, 1966).

2. Granite-porphyry, host rock of the sodium-rich stilbite (Analyst, K. Harada, 1966).

3. Stilbite (normal), Komuroyama, Shizuoka Prefecture (Analyst, K. Harada, 1966).

4. Komuroyama stilbite digested in 1N NaCl solution on water bath for 110 days (Analysts, K. Tomita and K. Harada, 1967).

Note: On specimens 1 and 3, Sr, Ba, Rb and Cs were determined by x-ray fluorescent analyses as being less than 0.1% in weight, if present (Analyst, A. Kato, 1966).

higher  $p_{\rm H_2O}$  and higher temperature than heulandite. In the light of the above results and mode of occurrence observed by the present writers, stalbite might have crystallized under lower  $p_{\rm H_2O}$ , lower temperature, or both, than heulandite did. In general, stilbites contain a larger proportion of H<sub>2</sub>O than heulandite.

### CHEMICAL COMPOSITION

Chemical analysis of material purified by hand picking was made by the oridinary wet method. The result is compared in Table 1 with normal stilbite in the drusy parts of andesitic tuff breccia from Komuroyama, Tagata-gun, Shizuoka Prefecture, Japan, and with the granite-porphyry host rock. The calculation of cation proportions assumes O=72 and  $H_2O=28$  as originally proposed by Sekanina and Wyart (1936) and as advocated by Černý (1965) to give the formulae:

$$\begin{split} & K_{0.52} Na_{3.30} Ca_{2.07} Mg_{0.52} Fe_{0.04}^{2+} Fe_{0.01}^{3+} Al_{9.97} Si_{26.55} \cdot O_{72} \cdot 28.87 \ H_2 O \ (Onigaj\bar{o}), \\ & K_{0.03} Na_{1.02} Ca_{4.17} Mg_{0.21} Fe_{0.01}^{3+} Al_{9.38} Si_{26.51} \cdot O_{72} \cdot 28.72 \ H_2 O \ (Komuroyama). \end{split}$$

Stilbite from Onigajō contains a remarkable amount of sodium showing Na>Ca in mole ratio and it also contains considerable amounts of

Specimens	Normal stilbite, Komuroyama.	Sodium-rich stilbite, Onigajō. $\alpha = 1.482 \pm 0.002$ $\beta = 1.489 \pm 0.002$ $\gamma = 1.496 \pm 0.002$ $\gamma - \alpha = 0.014$ $2V(-) = 43.2^{\circ}$ (mean) $XAc = 9^{\circ}$ (010) (perfect)	
Optical properties (Na light)	$\alpha = 1.492 \pm 0.002$ $\beta = 1.500 \pm 0.002$ $\gamma = 1.505 \pm 0.002$ $\gamma - \alpha = 0.013$ $2V (-) = 35.6^{\circ} (mean)$ $XAc = 2^{\circ} - 3^{\circ} (wavy)$		
Cleavage	(010) (perfect)		
Spec. gr. (mean)	2.19	2.18	
Hardness	4	3.5	

TABLE 2. Physical Properties of Sodium-Rich Stilbite with Corresponding Values for Normal Stilbite

potassium and magnesium. Such stilbite has not previously been reported.

## PHYSICAL PROPERTIES

The physical properties of the sodian stilbite are shown in Table 2 with corresponding values for normal stilbite from Komuroyama. These properties generally agree with those of normal stilbite, but closer observation of this mineral reveals that its refractive indices are lower, pearly luster is stronger, and it is more brittle and softer than normal stilbite.

### INFRARED ABSORPTION SPECTRA

Infrared absorption spectra were obtained using a DS-401-S gratingtype spectrophotometer with the Nujol past method. Figure 1 shows spectra of the sodian stilbite and stilbite from Komuroyama. Both spectra are quite similar and the bands of absorption around 3435 cm<sup>-1</sup>

1440

are probably due to adsorped water as indicated by comparison with montmorillonite data of Wolff (1965). The absorption bands around  $3600 \text{ cm}^{-1}$  may be attributed to structurally bound water. A clear asymmetrical absorption at 1640 cm<sup>-1</sup> due to adsorbed water is visible. The Si-O stretching region gives the band of adsorption at 1020 cm<sup>-1</sup>, which is in fair agreement with previous work (Milkey, 1960).

## DIFFERENTIAL THERMAL ANALYSES

DTA curves were taken by an automatic thermal analyser at a heating rate of 10°C/min. for the sodian stilbite and the normal stilbite from Komuroyama.

The endothermic peak at 230°C for the Komuroyama specimen was identical to those of previous workers (Koizumi, 1953; Černý, 1965), whereas the sodium-rich stilbite shows a characteristic endothermic





FIG. 1. Infrared absorption spectra of sodium-rich (upper) and normal (lower) stilbite. Note: Dotted line indicates the elimination of the peak due to Nujol.



FIG. 2. Differential thermal analysis curves of sodium-rich (upper) and normal (lower) stillbites.

doublet with maxima centered at 225°C and 285°C. Of the two, the lower temperature peak is correlated with that of the normal stilbite (Fig. 2).

## X-RAY DATA

On selected specimens of the sodium-rich and the normal stilbite, X-ray powder patterns were taken employing an X-ray diffractometer and CuK $\alpha$  radiation,  $\lambda = 1.5418$  Å, with a Ni-filter. The cell dimensions of the sodium-rich stilbite were determined by trial and error as a=13.62, b=18.16, c=11.31 and  $\beta=129^{\circ}10$ ; these values are very close to those of normal stilbite (Sekanina and Wyart, 1936; Mason and Greenberg, 1954). With these cell data, the sodium-rich stilbite can be indexed (Table 3), with a mean  $\Delta Q = \pm 3$ .

> Amounts of Exchangeable Cations and Cation-Exchange Capacities

Cation-exchange capacities and amounts of exchangeable cations were measured by Schollenberger and Simon's method (Schollenberger

$d_{\mathrm{obs}_*}$	$\begin{array}{c} I/I_0 \\ \times 100 \end{array}$	hkl	$d_{\rm obs}$	$\begin{array}{c} I/I_{\circ} \\ \times 100 \end{array}$	hkl
9.16	100	T10	2.489	2	<b>ī</b> 53
		020			062
6.85	1	201			331
6.39	0.5	021	2.350	3	263
5.31	4.5	200		1	444
5.24	2.5	130.130	2.277*	3	
4.67	11	$\overline{2}22$	2.229	2	371
4.56	2	220.220	2.125*	2.5	
		040	2.100	2	063
4.46*	2				510.510
4.29	8	312			282
4.07	45		2.090	2	335
4.02	7	041			370.370
		131	2.048	1.5	154
3.75	5	203			605
3.71	5		2.029	2.5	<b>4</b> 64
3.48	6	242			643
3.40	7	402			153
3.20	10	403	1.896*	2	
3.12	4	401	1.871*	1.5	
		333	1.821	7.5	641
3.02	32	152	1.807	2	316
		060	1.782	2	173
		423	1.728	1.5	620
3.000	11	151			620
2.974	3	421	1.636	1	045
2.876*	1.6				2103
2.825*	2		1.594	3	844
2.811	2	352	1.584	2.5	T16
2.744	12	023			826
		261	1.554	3	462.774
2.724	3	442			591
		204			175
2.599*	2		1.513	1	<b>4</b> 104.0120
2.566	5	334	1.470	1	865.264
2.532	1	350	1.448	1	801
		350	1.438	1.5	207
		420.420			776
2.510*	3		1.411	1	5113

TABLE 3. X-RAY POWDER DATA FOR SODIUM-RICH STILBITE

\* Unidentified lines.

### KAZUO HARADA AND KATSUTOSHI TOMITA

and Simon, 1945). The results are shown in Table 4 and the method is as follows:

Sodium-rich stilbite and normal stilbite were pulverized in an agate mortar and 0.2 gr of powder was mixed with 2 gr of quartz powder, and a complete mixture thus obtained was packed into a glass tube, which is 12 cm in length and 1.3 cm in inner diameter with a stem of 4 cm. length and 0.3 cm inner diameter. To prevent loss of the samples, absorbent cotton was placed at the bottom of the tube. Ammonium acetate solution (pH=7.0, 1N, 100 ml) was passed through the column for the duration between 8 and 10 hours. The amounts of exchangeable cations in the effluent were chemically determined. Then, the walls of the tube were washed three times with 80 percent ethyl alcohol (pH=7) until leaching of ammonium ions was not detected. In the next step, the column was washed with 10 percent KCl solution to leach adsorbed ammonium ions. The volume of the effluent was made to be 200 ml by adding distilled water, and the content of leached ammonium ions in 10 ml was determined by the distillation method.

TABLE 4.	CATION	Exchange	CAPACITIES	OF SODI	UM-RICH
	AND NO	RMAL STILE	BITE, IN med	1/100  g	

	1	2
$Ca^{2+}+Mg^{2+}$	98.07 (149.18)*	187.62 (304.32)*
$Na^++K^+$	124.07 (133.09)*	58.51 (36.53)*
Total	232.14 (282.27)*	246.13 (340.86)*
CEC	259.45	269.03

1. Sodium-rich stilbite from Onigajõ, Mié Prefecture.

2. Normal stilbite from Komuroyama, Shizuoka Prefecture. (Analysis by K. Tomita and T. Negishi, 1966).

\* Calculated on the basis of chemical analyses in Table 1.

The ideal values of exchangeable cations calculated on the basis of chemical analyses presented in Table 1 and the total amounts of exchangeable cations from the present examinations show considerable departures. These departures might be explained in that some cations might not have been leached from stilbite by the ammonium acetate solution. The CEC values agree well with experimental values of the total amounts of exchangeable cations.

## Artificial Replacement of Na<sup>+</sup> for Ca<sup>2+</sup> in Stilbite Structure at Room Pressures and at 100°C

The starting material was stilbite from Komuroyama. The stilbite was pulverized to fine powder, placed in a beaker (500 ml in capacity), and warmed on a water-bath with lN NaCl solution for 110 days. Evaporation was avoided by placing a glass plate on the top of the beaker.

1444

The lN NaCl solution was changed every two days. The obtained powder was washed with pure water 6 times in a centrifuge (2500 rpm), and examined with an X-ray diffractometer. No changes were apparent between starting material and the treated stilbite; the latter was analysed chemically (Table 1, Column 4). The chemical formula of the treated material is:  $Ca_{0.62}Na_{6.52}K_{0.02}Mg_{0.11}Al_{9.52}Si_{26.86}O_{72} \cdot 30.74H_2O$ 

The Na:Ca ratio of this formula is the reverse of the starting stillite from Komuroyama, but the total amount of Ca, Na, K and Mg in the treated stillite has diminished. This may be explained by substitution of some  $H^+$  for other cations during the experiments.

## Conversion Reactions of Stilbite into Wairakite at Low Water Vapor Pressures

### Experimental.

The starting material was stillite from Komuroyama. It was ground into fine powder in an agate mortar. Using closed autoclave (Morey type, 500 ml), water was put into the vessel and pressures were measured on heating. The powder was put into an open silver crucible placed on a stainless steel stand on the bottom of the vessel to avoid mixing of sample and water. The runs were made at heating rate 2°C/min. and were maintained under confining pressure at temperature for 2 to 94 hours. This type of autoclave is unquenchable, and it takes 3 hours at 100°C to reach room pressure and temperature. Temperature was controlled automatically, and the accuracy of measurements is believed to be within  $\pm$ 5°C. Pure-phase and converted-phase assemblages were obtained at water vapor pressures between 100 and 15 atms, between 200 and 400°C, from 2 to 94 hours, and were examined with an X-ray diffractometer immediately after cooling.

*Results*. Stilibite was easily converted into wairakite (Fig. 4) at water vapor pressures less than 100 atms. The 6.82 Å line which was discussed by Coombs (1955) and Ames (1966) for distinguishing wairakite and analcime was clearly exhibited and all diffractometer peaks agreed well with previous wairakite data, shown in Table 5 and Figure 3, together with synthesized wairakite from heulandite (at 300°C,  $p_{\rm H20} = 100$  atms, and 100 hours) and natural wairakite from Japan.

Wairakite is known to be easily synthesized from various starting materials under rather high pressure hydrothermal conditions, as reported by many workers (Ames and Sand, 1958; Barrer and Denny, 1961; Koizumi and Roy, 1960). From dehydrated heulandite Coombs *et al.*, (1959) obtained anorthite and wairakite at water vapor pressure and 370°C, at 2000–2600 bars 280–410°C assemblages containing wairakite, mordenite and analcime. From chabazite they obtained wairakite and phillipsite at temperatures 220°–310°C and water vapor pressure; and from stilbite they obtained epistilbite at 370°C, 5000 bars. The present writers obtained wairakite from stilbite at rather low water vapor pressures.



- FIG. 3. Selected X-ray diffraction patterns of natural and synthesized wairakites. A: Natural wairakite from Hanawa Mine, Akita Prefecture (wall rock of the Kurokō
- ore deposit of epithemal type). Specimen from Professor Y. Seki.
- B: wairakite after heulandite (300°C, 100 atms and 70 hours).

C: wairakite after stilbite (300°C, 100 atms and 70 hours).

Note: S indicates peaks of stilbite surviving conversion.

1446

In nature, wairakite occurs in active geothermal areas characterized by a steep thermal gradient (Steiner, 1953, 1955, 1958; Seki, 1966b), and also in some metamorphic terraines (Wise, 1959; Donnery, 1962; Whetten, 1965), or in some epithermal mineral deposits (Seki, 1966b). These



FIG. 4. Diagram showing the conversion reaction of stilbite into wairakite at low water vapor pressures. Hydrothermal conversion may have taken the following form:

Since no silica lines were observed, it was assumed to be in the noncrystalline state. Numbers represent the time in hours at confining water vapor pressures.

facts may agree well with our experimental studies, as well as with the result by Coombs *et al.*, (1959).

## ORIGIN OF NATURAL SODIUM-RICH STILBITE

According to Shibata *et al.*, (1959) and Shibata (1962), graniteporphyry and granite rocks from the Kii Peninsula show remarkable chemical compositions since  $Fe_2O_3+FeO$  surpasses CaO, and  $K_2O$ , Na<sub>2</sub>O and MgO occur in considerable amounts, as is well represented by the sodium-rich stilbite as well as the host rocks of the stilbite (granite-porphyry) (Table 1, Columns 1 and 2). Considerable amounts of Na<sub>2</sub>O,  $K_2O$  and MgO in sodium-rich stilbite might have come from the magma of the granite-porphyry in the final stage of the magmatic consolidation.

#### DISCUSSION

Our experimental studies suggest that almost all Ca<sup>2+</sup>, Mg<sup>2+</sup> and Na<sup>+</sup> in stilbite should be exchangeable, which is in contrast to the wairakiteanalcime substitutions. Such replacement is possible from the standpoint of the crystal structure of stilbite (Galli and Gottardi, 1966).

Since Na>Ca, a case can be made for considering the mineral a calcium-bearing representative of the sodium analogue of stilbite. If so, a new mineral name might be justified. However there are two ways in which Na (or Na,K) can replace Ca (or Ca,Mg) in a zeolite, either NaSi for CaAl or Na<sub>2</sub> for Ca.

Idealized formulae would be:

Ca4 Al<sub>8</sub> Si<sub>28</sub> O<sub>72</sub> 28H<sub>2</sub>O Na4 Al<sub>4</sub> Si<sub>32</sub> O<sub>72</sub> 28H<sub>2</sub>O Na8 Al<sub>8</sub> Si<sub>28</sub> O<sub>72</sub> 28H<sub>2</sub>O

or

In the present case the second type of replacement would appear to be dominant. Fifty percent replacement would give:

(Na, K)<sub>4</sub> (Ca, Mg)<sub>2</sub> Al<sub>8</sub> Si<sub>28</sub> O<sub>72</sub> 28H<sub>2</sub>O cf. Onigajō (Na, K)<sub>3.82</sub> (Ca, Mg)<sub>2.64</sub> Al<sub>9.57</sub> Si<sub>26.55</sub> O<sub>72</sub> 28.87H<sub>2</sub>O

On this basis, the Onigajō mineral is seen to have less than 50 percent of the Na<sub>2</sub> end member, and we believe is correctly termed sodian stilbite. Actually a further replacement, (Na, Al) (or Na, K, Ca, Al) for Si has clearly occurred to give Al>8, Si<28. This is a common phenomenon in zeolites.

#### ACKNOWLEDGEMENT

The writers wish to express their sincere thanks to Professor Dr. T. Sudo of the Geological and Mineralogical Institute, Tokyo University of Education, for his guidance and helpful comments through the study. Deep gratitude is also expressed to Drs. P. Černý of the Geological Institute, Czechoslovak Academy of Science, Praha, and R. C. Erd of the U. S. Geological Survey, Menlo Park, California, for the helpful comments concerning the chemical compositions of natural stilbite. Thanks are also due to Professor Dr. Y. Seki of the Department of Foundation Engineering, Saitama University, for his helpful comments concerning the stilbite reaction.

Special thanks are also due to Professors Drs. D. S. Coombs, University of Otago and W. G. Ernst, University of California who have been critically reviewed and materially improved the manuscript. Thanks are also due to Drs. K. Sakurai of the National Science Museum, and A. Kato of the Geological Institute, the University of Tokyo, for providing the specimens and kindly taking X-ray fluorescent analyses, as well as for their kind guidance.

The writers are greatly indebted to Professor Dr. H. Shibata and Professor Dr. Kozo Nagashima of Tokyo University of Education, for their guidance in chemical analyses, and

also to Professor Dr. K. Oinuma of the same University for supplying the data of infrared absorption spectra.

Stability relations of some calcium-zeolites are now being investigated in the laboratory of Professor Dr. W. G. Ernst and Dr. J. G. Liou, Department of Geology, University of California; their preliminary comments are appreciated.

#### References

- AMES, L. L. AND L. B. SAND (1958) Hydrothermal synthesis of wairakite and calciummordenite. Amer. Mineral., 43, 476-480.
- ———. (1966) Cation-exchange properties of wairakite and analcime. Amer. Mineral., 51, 903–909.
- ARAMAKI, S. AND S. HADA (1965) Geology of the central and southern parts of the acid igneous complex (Kumano Acidic Rocks) in southern Kii Peninsula. J. Geol. Soc. Japan, 71, 494-512.

---- (1965) Mode of emplacement of the acid igneous complex (Kumano Acidic Rocks) in the Southern Kii Peninsula. J. Geol. Soc. Japan, 71, 525–540.

- BARRER, R. M. AND P. J. DENNY (1961) Hydrothermal chemistry of the silicate, Part X. A partial study of the field CaO-Al<sub>2</sub>O<sub>3</sub>-SiO<sub>2</sub>-H<sub>2</sub>O. J. Chem. Soc. London, 202, 983-1001.
- ČERNÝ, P. (1965) Ionic substitutions in natural stilbite. Neues Jahrb. Mineral. Monatsh., 7, 198–208.

COOMBS, D. S. (1955) X-ray observations on wairakite and non cubic analcime. *Mineral.* Mag., 30, 699-708.

——, ELLIS, A. J., W. S. FYFE, AND A. M. TAYLOR (1959) The zeolite facies, with comments on the interpretation of hydrothermal synthesis. *Geochim. Cosmochim. Acta*, 17, 53-107.

—— (1961) Some recent work on the lower grades of metamorphism. Austr. J. Sci., 24, 203–215.

- DONNERY, T. J. (1962) Wairakite in West Indian spilitic rocks. Amer. Mineral., 47, 794– 802.
- GALLI, E. AND GOTTARDI, G. (1966) The crystal structure of stilbite. *Mineral. Petrogr. Acta*, 12, 1–10.

HARADA, K. (1964) Mineralogical notes on the alternations of liparite, Konoki, Mihamamachi, Mié Prefecture. J. Geol. Soc. Japan, 70, 296-299.

KAWANO, T. AND Y. UEDA (1964) K-A dating on the igneous rocks in Japan (1). J. Jap. Ass. Mineral. Petrol. Econ. Geol., 51, 127-148.

KOIZUMI, M. (1953) The differential thermal analysis courves and dehydration courves of zeolite. *Mineral. J.* (*Tokyo*)., 1, 36–47.

—— AND R. ROY (1960) Zeolite studies 1. Synthesis and stability of the calcium zeolites. J. Geol., 68, 41–53.

MASON, B. AND S. S. GREENBERG (1960) Zeolites and associated minerals from southern Brazil. Ark. Mineral. Geol., 1, 519–526.

MILKEY, R. G. (1960) Infrared spectra of some tectosilicate. Amer. Mineral., 45, 990-1007.

SAWATARI, M. (1932) Geology and petrology of the Kinomoto District, Province Kii. M.S. Thesis, Univ. Tokyo.

- SCHOLLENBERGER, C. J. AND R. N. SIMON (1945) Determination of exchange capacity and exchangeable bases in solil-ammonium acetate method. *Soil Sci.*, 59, 13–24.
- SEKANINA, J. AND J. WYART (1936) Sur la stilbite. Bull. Soc. Franc. Mineral. Crystallogr. 59, 377–383.
- SEKI, Y. (1966a) Wairakite in Japan (I). J. Jap. Ass. Mineral. Petrol. Econ. Geol., 55, 254-261.

— (1966b) Wairakite in Japan (II). J. Jap. Ass. Mineral. Petrol. Econ. Geol., 56, 30-39.

SHIBATA, H., S. OKADA AND Y. ÖKI (1958) Chemical composition of Japanese granitic rocks in regard to petrographic province (Part 6). Sci. Rep. Tokyo Univ. Educ. Sec. C., 50, 31-54.

-----, (1962) Chemical composition of Japanese granitic rocks in regard to petrographic province (Part 10). Sci. Rep. Tokyo Univ. Educ., Sec. C., 72, 33-47.

STEINER, A. (1953) Hydrothermal rock alteration at Wairakei, New Zealand. Econ. Geol., 48,1-3.

—— (1955) Wairakite, the calcium analogue of analcime, a new zeolite mineral. Mineral. Mag., 30, 691–698.

(1958) Occurrence of wairakite at the Geyser, California. Amer. Mineral., 43, 781. WHETTEN, J. T. (1965) Wairakite from low-grade metamorphic rocks on St. Croix, U. S.

 Virgin Islands. Amer. Mineral., 50, 752–755.
 WISE, W. S. (1959) Occurrence of wairakite in metamorphic rocks of the Pacific Northwest. Amer. Mineral., 44, 1099–1101.

WOLFF, R. G. (1965) Infrared absorption patterns (OH region) of several clay minerals. Amer. Mineral., 50, 240-244.

Manuscript received, August 5, 1966; accepted for publication, April 18, 1967.