montmorillonitic materials in environments other than marine. Caution must therefore be exercised in the use of glauconite as an environmental indicator.

References

- BURST, J. F. (1958) Mineral heterogeneity in "glauconite" pellets. Am. Mineral. 43, 481– 497.
- CLOUD, P. E., JR. (1955) Physical limits of glauconite formation. Am. Assoc. Petrol. Geol. Bull. 39, 484–492.
- DYDCHENKO, M., G. AND A. Y. KHATUNTZEVA (1956) Cases of glauconite in a continental environment. Mem. Soc. Russe Min., Ser. 2, 85, 49 (Min. Abs. 13, 287).
- FOSTER, M. D. (1956) Correlation of dioctahedral potassium micas on the basis of their charge relations. U. S. Geol. Survey Bull. 1036-D, 57-67.
- KELLER, W. D. (1958) Glauconitic Mica in the Morrison Formation in Colorado. Clays and Clay Minerals, Fifth Natl. Conf. Clays Clay Minerals, 1956. Natl. Acad. Sci., Washington, D. C.
- MEIGS, C. C., H. P. BASSETT AND G. G. SLAUGHTER (1922) Report on Texas alkali lakes. Texas Bur, Econ. Geol. Bull, 2234.
- REEVES, C. C., JR. (1962) Pleistocene lake basins of West Texas. Geol. Soc. Am. Spec. Paper 72, 222-223.
- (1963) Subterranean natural brines produce sodium sulphate in West Texas. Ground Water. 1, 35-36.
- ------ AND W. T. PARRY (1965) Geology of West Texas pluvial lake carbonates. Am. Jour. Sci. 263, 606–615.
- ROBERSON, H. E. AND E. C. JONES (1965) Clay minerals intermediate between illite and montmorillonite. Am. Mineral, 50, 766-770.
- SABATIER, M. (1949) Recherches sur la glauconie. Bull. Soc. franc. Mineral. 72, 473. in, W. A. DEER, R. A. HOWIE AND J. ZUSSMAN (1962) Rock-Forming Minerals. Vol. 3. John Wiley and Sons, New York, p. 39.
- TAKAHASHI, J. (1939) Synopsis of glauconitization in Recent Marine Sediments. Am. Assoc. Petrol. Geol. 503–512.
- WEAVER, C. E. (1965) Potassium content of illite. Science. 147, 603-605.

THE AMERICAN MINERALOGIST, VOL. 51, JANUARY-FEBRUARY, 1966

DEHYDRATION OF DIASPORE AT WATER PRESSURES FROM 15 to 15,000 PSI

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The dehydroxylation of well-crystallized diaspore from Chester, Mass. has been investigated by differential thermal analysis at H_2O pressures ranging from 15 to 15,000 psi. Experimental techniques and the methods used to process data taken from the thermograms have been described in detail by Weber and Greer (1965).

(1) $80-100$ mesh grain size 524 14.7 594 520 14.7 596 522 14.7 596 520 14.7 600 533 41 601 554 78 610 554 115 623 548 169 620 548 282 619 553 465 513 1360 590 531 2360 590 544 5740 579 550 7000 519 10570 559 15600 519 14.7 581 523.5 14.7 583 524 14.7 583 525.5 14.7 583 524 14.7 581 525.5 14.7 583 524 14.7 583 525.5 14.7 583 524 14.7 585 531 572 594	Characteristic temperature ¹ ° C.	P_{H_20} psi	Peak temperature ¹ ° C.
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	(1) 80–100 mesh grain size		
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	524	14.7	594
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	520	14.7	596
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	522	14 7	596
$\begin{array}{cccccccccccccccccccccccccccccccccccc$			
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	520	14.7	600
$\begin{array}{c ccccccccccccccccccccccccccccccccccc$	533	41	601
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	554	78	610
$\begin{array}{cccccccccccccccccccccccccccccccccccc$			
$\begin{array}{c ccccccccccccccccccccccccccccccccccc$	554	115	623
$\begin{array}{c ccccccccccccccccccccccccccccccccccc$	548	169	620
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	548	282	619
$\begin{array}{cccccccccccccccccccccccccccccccccccc$			
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	553	465	
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	549	504	624
$\begin{array}{c ccccccccccccccccccccccccccccccccccc$	513	1360	590
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	531	01(0	500
$\begin{array}{c ccccccccccccccccccccccccccccccccccc$	531	2300	590
$\begin{array}{c ccccccccccccccccccccccccccccccccccc$	344	5740	579
$ \begin{array}{c ccccccccccccccccccccccccccccccccccc$	550	7000	
559 15600 $-$ (2) $-325 mesh grain size$ 519 14.7 581 523 14.7 582 527 14.7 583 525.5 14.7 583 525.5 14.7 585 524 14.7 591 531 51 572 542 84 600 538 112 602 545 325 598 555 345 595 546 394 593 546 425 607 551 626 599 547 1285 583 549 2420 576 561 8400 566	519	10570	_
$ \begin{array}{c ccccccccccccccccccccccccccccccccccc$	559	15600	
$ \begin{array}{c c c c c c c c c c c c c c c c c c c $	(2) _ 325 much and in side	20000	
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	(2) = 323 mesn grain size	4.4 27	501
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	519	14.7	501
527 14.7 583 525.5 14.7 581 523.5 14.7 585 524 14.7 591 531 51 572 542 84 600 538 112 602 545 182 590 545 325 598 555 345 595 546 394 593 546 425 607 551 626 599 547 1285 583 549 2420 576 561 8400 566	525	14.7	592
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	521	14.7	303
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	525 5	14 7	581
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	523.5	14.7	585
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	524	14.7	591
$\begin{array}{cccccccccccccccccccccccccccccccccccc$			
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	531	51	572
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	542	84	600
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	538	112	602
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	~ . ~	102	500
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	545	182	590
555 345 595 546 394 593 546 425 607 551 626 599 547 1285 583 549 2420 576 561 8400 566	545	325	598
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	555	345	595
546 425 607 551 626 599 547 1285 583 549 2420 576 561 8400 566	546	304	593
510 626 599 551 626 599 547 1285 583 549 2420 576 561 8400 566	546	425	607
547 1285 583 549 2420 576 561 8400 566	551	626	500
547128558354924205765618400566	001	020	077
549 2420 576 561 8400 566	547	1285	583
561 8400 566	549	2420	576
	561	8400	566

TABLE 1. DEHYDRATION DATA FOR DIASPORE

¹ "Characteristic temperature" is defined as the temperature at the point of intersection of the DTA base line and the tangent drawn at the point of maximum slope of the DTA curve. This measure of the reaction temperature is more reproducible than is peak temperature (temperature at which the DTA curve is most remote from the base line). The DTA apparatus has been calibrated with substances of known reaction temperature (fusion or inversion) so that the temperatures listed in column one closely represent the temperatures at which the dehydroxylation reaction began. Pressure-temperature data are listed in Table 1 and are compared graphically in Fig. 1 with the diaspore \Rightarrow corundum+water boundary determined by Kennedy (1959). The pressure-temperature (p-t) curve is essentially vertical at all water pressures attained, and grain size, at least in the range 80 to -325 mesh, appears to have little effect on the position



FIG. 1. Pressure-temperature curve for the dehydroxylation of diaspore (-325 mesh particle size and 80–100 mesh particle size) determined by differential thermal analysis. Ten points fall within the open circle at $P_{H_20}=1$ atm (four replicate determinations for 80–100 mesh particle size and six replicates of -325 mesh samples). The equilibrium p-t curve obtained by Kennedy (1959) is shown for comparison.

of the p-t curve. The dehydroxylation behaviour of diaspore differs from that of kaolinite, dickite, halloysite and serpentine studied under the same experimental conditions (Roy and Weber, 1964) in that the apparent dehydration temperature for the latter group of minerals increases with increasing $P_{\rm H_2O}$ until the region 500 to 1000 psi is reached, above which dehydration temperatures decrease irregularly with further increase in water pressure. Reversal in the slope of the p-t curve has been ascribed to metastable persistence of the hydrous minerals above the limit of stability in the relatively low water pressure range. At higher pressures, the "catalytic" or "solvent" effect of water facilitates dehydroxvlation of the metastable phase. Thus in the case of the kandites and serpentine, the temperature required for dehydration under the dynamic conditions of differential thermal analysis (DTA) above 1 atm $P_{H_{2}O}$ is always above the maximum equilibrium temperature at which these hydrates are stable, and increases with increasing water pressure between 1 atm and several hundred or more psi. At higher water pressures, the dehydration temperature approaches, with considerable scatter, the equilibrium p-t curve. The p-t curve for the dynamic dehvdration of diaspore is displaced from the equilibrium curve by approximately 150 to 200° C. and falls well within the stability field of corundum+water. The absence of any appreciable variation in dehydration temperature in the pressure range investigated demonstrates the sluggishness of this reaction even under water pressures as high as 15,000 psi.

The heat of dehydration, ΔH , determined from the area under the DTA curve for all analyses at $P_{\Pi_2O} = 1$ atm is 91.9 cal/gm for -325 mesh samples, and 91.0 cal/gm for samples whose grain size ranged from 80 to 100 mesh. The degree of dispersion of the ΔH measurements is indicated by a standard deviation of about 7 cal/gm.

Work reported here was sponsored by the Air Force Cambridge Research Laboratories, Office of Aerospace Research, under contract No. AF 19(628)-2773. I am indebted to Prof. Rustum Roy for discussion of the manuscript.

References

KENNEDY, G. C. (1959) Phase relations in the system Al₂O₃-H₂O at high temperatures and pressures. Am. Jour. 257, 563–573.

WEBER, J. N. AND R. T. GREER (1965) Dehydration of serpentine: heat of reaction and reaction kinetics at $P_{H_20} = 1 \text{ atm} \cdot Am \cdot Mineral.$ 50, 450-464.

ROY R. AND J. N. WEBER (1964) Stability-metastability relationships of hydrous minerals and their importance in designing facilities for the extraction of water from lunar rocks and minerals. Proc. Third Ann. Meet. Working Group Extraterrestrial Resources 117-130.