

## X-RAY STUDIES OF SYNTHETIC COFFINITE, THORITE AND URANOTHORITES\*

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### ABSTRACT

X-ray data are presented for synthetic coffinite, thorite, and several uranothorites. The cell constants obtained for coffinite are  $a = 6.981 \pm 0.004$  kX,  $c = 6.250 \pm 0.005$  kX; for thorite  $a = 7.128 \pm 0.004$  kX,  $c = 6.314 \pm 0.003$  kX. Intermediate constants determined for several uranothorites indicate a continuous solid solution between  $\text{USiO}_4$  and  $\text{ThSiO}_4$ . Coffinite and thorite are isostructural with zircon; the space group is  $D_{4h}^{10}-I_4/a\text{ md}$ . The oxygen positions for coffinite are  $u = 0.180 \pm 0.010$ ,  $v = 0.347 \pm 0.010$  and for thorite  $u = 0.166 \pm 0.010$ ,  $v = 0.347 \pm 0.010$ . No changes were observed either in line intensities or in cell constants when  $(\text{OH})$  was removed from the hydrothermal preparations.

The mineral coffinite, described as a uranous silicate with hydroxyl substitution, was identified on the basis of the similarity of its x-ray powder pattern to that of zircon ( $\text{ZrSiO}_4$ ) or thorite ( $\text{ThSiO}_4$ ) (Stieff, Stern and Sherwood, 1956). Pabst (1951) obtained single crystal patterns of a New Zealand detrital uranothorite (11.5 wt%  $\text{UO}_2$ ) and determined the space group to be the same as given for zircon. A similar assignment was made by Bonatti and Gallitelli (1951) on detrital thorite crystals from Nettuno Rome.

Although coffinite is isostructural with zircon and thorite, the naturally occurring mineral is reported to exist as the hydroxyl substituted form of  $\text{USiO}_4$  (Stieff, Stern and Sherwood, 1956). This consideration is based on the low silicon content which appears in the analyses of the mineral. However, our observations on a synthetic product (Hoekstra and Fuchs, 1956) indicate that neither cell dimensions nor line intensities appear to be modified when water is removed, suggesting therefore that the synthesized product may be  $\text{USiO}_4$  without hydroxyl substitution. Infrared examination of heated samples made in this laboratory clearly indicate that water has been expelled (Hoekstra and Fuchs, to be published).

In addition to coffinite, we have synthesized thorite and some uranothorites. The thorite can be prepared by a hydrothermal process or from the melt (Fuchs, 1958). The latter process yields water-free samples, and again it does not seem possible to distinguish between hydroxyl-containing and hydroxyl-free samples on the basis of x-ray powder patterns.

Powder data for the synthesized coffinite and thorite are presented in Table I. Cell dimensions and x-ray densities for these samples are given

\* Based on work performed under the auspices of the U. S. Atomic Energy Commission.

TABLE I. X-RAY DATA FOR COFFINITE AND THORITE. CuK $\alpha_{1,2}$  = 1.5386 kX

Coffinite					Thorite					
<i>hkl</i>	$\sin^2 \theta$ obs.	$\sin^2 \theta$ calc.	$I_s$	$I_e$	<i>hkl</i>	$\sin^2 \theta$ obs.	$\sin^2 \theta$ calc.	$I_s$	$I_e$	
101	0.02758	0.02759	197	214	101	0.02669	0.02666	174	208	
200	.04917	.04912	203	227	200	.04711	.04712	210	249	
211	.07636	.07653	96	105	211	.07348	.07357	99	111	
112	.08540	.08538	196	184	112	.08287	.08299	162	188	
220	.09784	.09792	53	58	220	.09376	.09376	63	68	
202	.1097	.1098	9	5	202	.1065	.1064	6	4	
301	.1252	.1252	62	54	301	.1202	.1202	63	57	
103	.1489	.1490	60	38	103	.1456	.1456	42	40	
321	.1737	.1739	69	58	321	.1672	.1669	67	63	
312	.1827	.1829	144	141	312	.1765	.1764	135	140	
400	.1964	.1953	1965	66	67	400	.1870	.1870	40	34
213		.1977				213	.1924	.1922	33	39
411	.2227	.2226	32	25	411	.2138	.2134	25	24	
004	.2437	.2429	2444	53	53	420	.2335	.2335	39	35
420		.2440				004	.2378	.2378	18	19
303		.2464				303	.2383	.2383		
402	—	.2559	0	<1	402	—	.2462	0	<1	
332	.2802	.2802	47	37	332	.2695	.2694	38	39	
204	.2913	.2916	48	30	204	.2847	.2843	36	41	
323	.2948	.2951	10	12	323		.2849			
422	—	.3045	0	<1	422	—	.2926	0	0	
501	0.3200	0.3199	34	29	501	0.3066	0.3064	27	30	
431		431	431	431	431		431			
224	.3405	.3402	47	45	224	.3315	.3309	36	43	
413	.3433	.3437	413	413	413		.3314			
314	—	.3646	0	<1	314	—	.3541	0	<1	
521	.3684	.3685	15	10	521	.3534	.3529	10	11	
512	.3777	.3774	36	38	512	.3627	.3624	38	39	
440	.3906	.3898	3906	18	14	440	.3732	.3728	8	8
105		.3914				105	.3829	.3828	5	5
404	.4390	.4375	4395	56	52	600	.4193	.4193	12	11
600		.4383				404	.4239	.4239		
215		.4400				433	.4245	.4248	26	29
503		.4409				503	.4248	.4245		
433		.4409				215	.4292	.4293	14	12
611	.4649	.4655	13	9	611	.4466	.4457	11	9	
532	.4753	.4745	37	30	532	.4557	.4553	27	31	
424	.4870	.4860	4877	60	59	620	.4657	.4658	17	16
620		.4868				424	.4703	.4703	33	42
305		.4886				523	.4705	.4713		
523		.4894				602	—	.4785	0	0
602	—	.4988	0	<1	541	.4928	.4922	7	7	

TABLE 1. (continued)

Coffinite					Thorite				
<i>hkl</i>	$\sin^2 \theta$ obs.	$\sin^2 \theta$ calc.	$I_{\text{B}}$	$I_{\text{c}}$	<i>hkl</i>	$\sin^2 \theta$ obs.	$\sin^2 \theta$ calc.	$I_{\text{B}}$	$I_{\text{c}}$
541	.5145	.5141	9	7	325	.5227	.5224	7	7
325	.5380	.5372	10	8	622	—	.5245	0	<1
622	—	.5474	0	<1	631	.5395	.5388	9	10
514	—	.5589	0	<1	514	—	.5398	0	0
631	.5628	.5626	14	10	116	.5573	.5576	10	14
116	.5697	.5701	20	14	613	.5649	.5644	9	10
415	.5866	.5858	17	18	415	.5682	.5690	7	9
613	.5866	.5862	206	701	206	—	.5809	0	<1
206	.5865	.5863	701	.5863	.5854	—	3	3	3
701	—	.5944	0	<1	640	.6051	.6055	12	13
444	.6122	.6111	6	4	444	—	.6094	—	—
640	.6340	.6317	29	34	543	.6099	.6099	15	21
543	.6324	.6330	721	.6323	.6103	—	5	6	6
534	—	.6560	0	<1	534	—	.6329	0	<1
721	.6601	.6596	14	9	552	—	.6411	.6408	20
316	.6672	—	552	—	712	—	.6424	.6439	22
552	.6684	.6685	56	56	552	—	—	—	—
712	.6685	—	712	—	712	—	—	—	—
604	.6802	—	316 $\alpha_1$	—	604	.6501	.6501	14	14
505	.6829	—	604	—	604	.6559	.6559	—	—
435	.6827	.6828	28	27	633	.6566	.6564	13	10
633	.6829	—	633	—	505	.6568	.6564	—	—
642	—	.6928	0	<1	642	—	.6614	10	7
624	.7287	—	435	—	642	—	.6641	0	<1
525	.7304	.7314	7307	32	624 $\alpha_1$	.7023	.7024	28	27
703	.7320	—	703	—	703	.7034	.7034	—	—
406	—	.7400	0	<1	624 $\alpha_2$	.7059	.7055	—	—
107	.7564	.7548	7558	10	525 $\alpha_1$	.7090	.7080	6	6
651	.7567	.7558	10	10	406	—	.7206	0	<1
336	.7643	—	651 $\alpha_1$	—	651 $\alpha_1$	.7247	.7243	8	7
732	.7660	.7650	33	38	107	—	.7257	—	—
800	.7656	—	732 $\alpha_1$	—	732 $\alpha_1$	.7336	.7272	15	17
723	.7779	.7792	14	15	732 $\alpha_2$	.7375	.7374	—	—
426	.7803	.7806	336	—	336	—	.7431	—	—
217	—	.7885	0	<1	$\alpha_1$	.7444	.7437	10	11
811	.8025	—	800	—	800	.7443	—	—	—
741	.8036	.8033	21	18	723 $\alpha_1$	.7493	.7498	8	5
741	.8036	—	426	—	426	—	.7671	0	0
217	.8053	—	811	—	811	.7715	.7708	8	9
811	.8075	.8068	741	—	741	—	.7746	.7746	—
741	.8075	—	811	—	811	—	—	—	—
820	.8265	.8275	17	21	217 $\alpha_1$	.7847	.7846	6	6
615	.8281	.8285	820 $\alpha_1$	—	820 $\alpha_1$	.7915	.7908	8	8

TABLE 1. (continued)

Coffinite					Thorite				
<i>hkl</i>	$\sin^2 \theta$ obs.	$\sin^2 \theta$ calc.	I <sub>o</sub>	I <sub>e</sub>	<i>hkl</i>	$\sin^2 \theta$ obs.	$\sin^2 \theta$ calc.	I <sub>o</sub>	I <sub>e</sub>
802	—	.8385	0	<1					
714	—	.8502	0	<1					
307	.8517	.8518	8	5	820 $\alpha_2$	.7951	.7948		
516 $\alpha_1$	.8610	.8501			615 $\alpha_1$	.8021	.8010	5	5
					802	—	.8046	0	
					307		.8319	0	
516 $\alpha_2$	.8646	.8639			516		.8361		<1
					660	.8363	.8368		4
							.8374		
					714		.8189		
644		.8744			516		.8393		
660					660	.8421	.8404		
545	.8769	.8751			644		.8415		
		.8771			653	.8421	.8424		
653		.8778					.8429		
822	—	.8871							
327		.9004			644		.8457		
831	.9018	.9014			653	.8467	.8463		
		.9024					.8468		
635		.9257			545 $\alpha_1$	.8467	.8475		
743	.9258	.9262			822		.8501		
813		.9264			545 $\alpha_2$	.8525	.8510		
		.9264			831 $\alpha_1$	.8641	.8538		
734	—	.9474							
417	.9480	.9490			831 $\alpha_2$	.8682	.8680		
536		.9586			327 $\alpha_1$	.8773	.8777		
752	.9585	.9593			327 $\alpha_2$	.8815	.8807		
840	.9729	.9723			743				
705	—	.9743			813	.8893	.8894		
606	—	.9831							
					743				
					813	.8939	.8935		
					635 $\alpha_1$	.8939	.8946		
					734	—	.9120		
					752 $\alpha_1$	.9202	.9199		
					752 $\alpha_2$		.9243		
					417 $\alpha_1$	.9247	.9242		
					536		.9292		
					840	.9292	.9298		
							.9304		
					536		.9328		
					840	.9345	.9339		
							.9350		
					705 $\alpha_1$	.9402	.9405		
					008 $\alpha_1$	.9485	.9489		
					606	—	.9535		
					901 $\alpha_1$	.9574	.9569		
					901 $\alpha_2$	.9622	.9616		
					804		.9815		
					833	.9818	.9820		
							.9824		
					804		.9860		
					833	.9870	.9866		
							.9871		

Observed intensities not resolved.

Observed intensities not resolved.  
Observed intensities not resolved.

TABLE II. CELL DIMENSIONS AND DENSITIES OF COFFINITE,  
THORITE AND URANOTHORITES

Material	<i>a</i>	<i>c</i>	<i>a/c</i>	$\rho$ gm./cm. <sup>3</sup> calc.
USiO <sub>4</sub> *	6.981 ± .004 kX	6.250 ± .005 kX	1.1170	7.15 ± 0.02
3 USiO <sub>4</sub> · ThSiO <sub>4</sub>	7.007 ± .005	6.275 ± .003	1.1167	7.04 ± .02
USiO <sub>4</sub> · ThSiO <sub>4</sub>	7.039 ± .003	6.294 ± .002	1.1184	6.91 ± .01
USiO <sub>4</sub> · 3 ThSiO <sub>4</sub>	7.071 ± .002	6.314 ± .003	1.1199	6.80 ± .01
ThSiO <sub>4</sub>	7.128 ± .004	6.314 ± .003	1.1289	6.67 ± 0.01

\* A value previously reported for USiO<sub>4</sub> (Hoekstra and Fuchs, 1956), varies slightly from the value given here, since the earlier value was based on an incomplete indexing of the powder pattern.

in Table II. We also include in this table *x*-ray data for the several uranothorite samples which were prepared by the hydrothermal process developed for coffinite. The only variation in technique involves the preparation of thorium and uranium tetrachloride solutions in the desired concentrations. Reference to Table II shows that the cell dimension changes are uniform throughout the composition range.

We have also attempted to determine the oxygen positions in coffinite and thorite from powder patterns. Although the oxygen scattering is very small, the assigned positions are necessary in order to obtain reasonable agreement between observed and calculated intensities.

The atomic positions are assumed to be those for zircon,  $D_{4h}^{19}$ — $I_4/a$  md and are as follows:

4 Th or 4 U in (*a*)

4 Si in (*b*)

16 O in (*h*) (International Tables)

The only variables are *u* and *v* for oxygen and these were determined by trial methods. The resulting values for the oxygen positions are given in Table III. The extent of agreement between observed and calculated intensities is shown in Table I. The observed intensities were deduced from microphotometer tracings and the calculated intensities were obtained from the formula

$$I \sim F^2 \rho \frac{1 + \cos^2 2\theta}{\sin^2 \theta \cos \theta}$$

TABLE III. OXYGEN POSITIONS

	<i>u</i>	<i>v</i>
USiO <sub>4</sub>	.180 ± 0.010	.347 ± 0.010
ThSiO <sub>4</sub>	.166 ± 0.010	.347 ± 0.010

TABLE IV. BOND LENGTHS

USiO <sub>4</sub>	U—4 O	2.32 Å ± .08 Å
	U—4 O	2.52 Å ± .09 Å
	Si—4 O	1.58 Å ± .09 Å
ThSiO <sub>4</sub>	Th—4 O	2.46 Å ± .08 Å
	Th—4 O	2.50 Å ± .09 Å
	Si—4 O	1.55 Å ± .09 Å
ZrSiO <sub>4</sub>	Zr—4 O	2.05 Å
	Zr—4 O	2.41 Å
	Si—4 O	1.62 Å

where  $\phi$  is the multiplicity factor and the other quantities have their usual significance.

The bond distances which result are shown in Table IV. The bond lengths in zircon (Wyckoff and Hendricks, 1927) are given for comparative purposes.

#### ACKNOWLEDGMENT

The writers wish to thank Dr. Stanley Siegel for helpful discussions during the course of the work.

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*Manuscript received July 15, 1957*