

## Chemistry and genetic implications of tourmaline and Li-F-Cs micas from the Valdeflores area (Cáceres, Spain)

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### ABSTRACT

Pervasive metasomatism that involved the formation of tourmaline-rich rocks and influx of Li, F, and Cs into Ordovician psammo-pelitic metasediments occurred in the Valdeflores area (Cáceres, Spain). Numerous Li- and Sn-bearing, mineralized, greisen-type veins also can be observed here, in the vicinity of geochemically specialized granites. Tourmaline-rich rocks appear as: (1) massive, fine-grained, dark green to black rocks; and (2) fine-scale tourmaline-rich laminae, which alternate with quartz-rich layers parallel to the bedding.

Electron microprobe analyses indicate that the tourmaline lies mostly within the space defined by the exchange vectors:  $\text{FeMg}_{-1}$  (schorl),  $\square\text{AlNa}_{-1}\text{Mg}_{-1}$  (foitite),  $\text{AlOMg}_{-1}(\text{OH})_{-1}$  (olenite), and  $\text{CaMgNa}_{-1}\text{Al}_{-1}$  (uvite). The  $\text{Fe}/(\text{Fe}+\text{Mg})$  ratio ranges mainly from 0.87 to 0.54 and increases with Al in the Y-site. Analytical results and substitutional relations show an insignificant elbaite component. Mica in the tourmalinized rocks is very fine-grained (mostly  $<50\ \mu\text{m}$ ). White mica ranges from lithian muscovite-phengite to lepidolite/zinnwaldite, containing up to 8.40 wt% F, 6.0 wt%  $\text{Li}_2\text{O}$ , and 10.73 wt% FeO. Dark mica shows a variable color and has compositions characterized by relatively high contents of  $\text{Cs}_2\text{O}$  (1.14–2.78 wt%) and F (1.94–8.08 wt%), with a deficit in  $\text{K}_2\text{O}$  (5.75–9.04 wt%).  $\log(f_{\text{H}_2\text{O}}/f_{\text{HF}})$  of fluids in equilibrium with biotite in the tourmaline-rich rocks was 4.02–4.17 at  $T \approx 400\ ^\circ\text{C}$ .  $\log(f_{\text{H}_2\text{O}}/f_{\text{HF}})$  values of fluids in equilibrium with topaz ( $X_{\text{F}} \approx 0.8$ ) in country rock adjacent to contacts with veins, and in equilibrium with amblygonite-montebbrasite ( $X_{\text{amb}} = 0.2$ ) in the veins were about 4.30–4.60 and 6.4–6.7, respectively. These variations denote the existence of gradients in relative  $a_{\text{HF}}$  more than differences of temperature during metasomatism. The lack of tourmaline in the veins is interpreted to reflect the alkalinity and low Fe-Mg contents in the fluids, which precluded the formation of tourmaline. Consequently, most of the boron was expelled into metasediments where tourmaline was produced as a result.