

Supplementary Figures and Tables

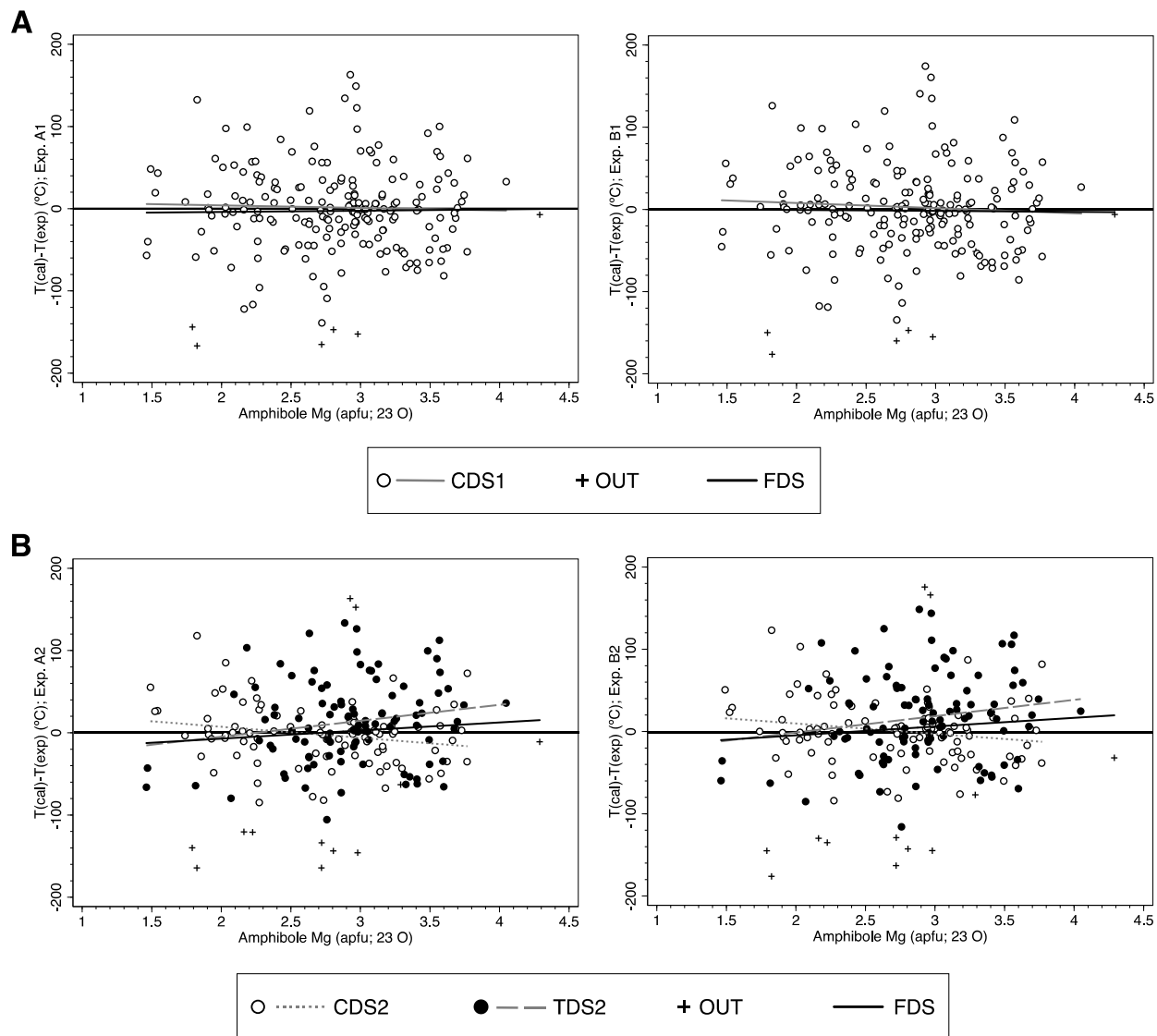


Fig. 1SB. Relationships of residuals versus amphibole Mg occupancy for the new amphibole-plagioclase NaSi–CaAl exchange thermometers. A) Expressions A1 and B1. B) Expressions A2 and B2. Abbreviations: CDS1, Calibration Data Set 1; CDS2, Calibration Data Set 2; TDS2, Test Data Set 2; FDS, Full Data Set; OUT, outliers.

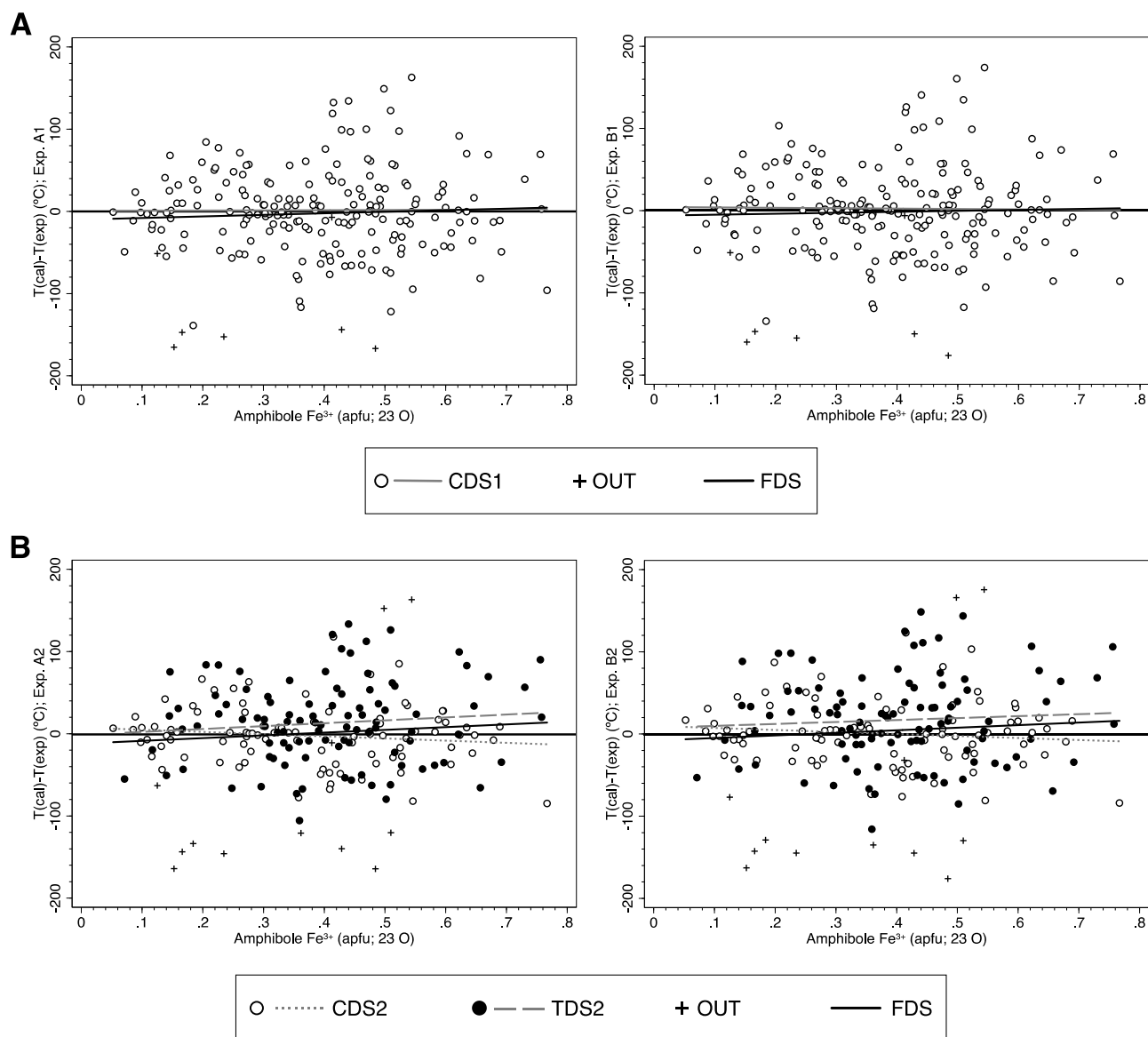


Fig. 2SB. Relationships of residuals versus amphibole Fe^{3+} occupancy for the new amphibole-plagioclase NaSi–CaAl exchange thermometers. A) Expressions A1 and B1. B) Expressions A2 and B2. Abbreviations: CDS1, Calibration Data Set 1; CDS2, Calibration Data Set 2; TDS2, Test Data Set 2; FDS, Full Data Set; OUT, outliers.

Table 1SB. Ranges of pressure, temperature and compositional variables of the full amphibole-plagioclase data set (203 observations)

| <i>FDS parameters</i> | <i>Amphibole composition (apfu, 23O)</i> | <i>Amphibole phase components</i> |
|--|--|-----------------------------------|
| $P(\text{kbar})=1-15$ | $Si=5.956-7.431$ | $p_1=0.034-0.214$ |
| $T(^{\circ}\text{C})=640-1000^{\circ}\text{C}$ | $Ti=0.004-0.349$ | $p_2=-0.277-0.311$ |
| $\ln K^{id}=-9.00- -1.99$ | $Al=0.668-3.392$ | $p_3=0.002-0.174$ |
| $RT\ln K^{id}=-76905- -17741 \text{ (J)}$ | $Fe =0.655-3.110$ | $p_4=0.026-0.383$ |
| $\ln K^{id}=-9.00- -1.99$ | $Mg=1.461-4.289$ | $p_5=0.023-0.708$ |
| $Den(A1)=-182- -122$ | $Ca=1.287-1.868$ | $p_1=0.034-0.214$ |
| $Den(B1)=-183- -124$ | $Na=0.229-0.785$ | $p_2=-0.277-0.311$ |
| $Den(A2)=-195- -132$ | $K<0.223$ | $p_3=0.002-0.174$ |
| $Den(B2)=-194- -132$ | $F<0.062$ | $p_4=0.026-0.383$ |
| <i>Plagioclase phase components</i> | $Cl<0.006$ | $p_5=0.023-0.708$ |
| $p_{ab}=0.108-0.820$ | $Al_{VI}=0.093-1.449$ | $p_6<0.223$ |
| $p_{an}=0.180-0.891$ | $Na_{M4}=0.067-0.427$ | $p_7=0.006-0.183$ |
| $p_{or}<0.109$ | $Na_A=0.023-0.708$ | $p_8=0.029-0.345$ |
| | $Fe_{2+}=0.285-2.676$ | $p_9=-0.215-0.321$ |
| | $Fe_{3+}=0.052-0.767$ | $p_{10}<0.053$ |
| | $Mg/(Mg+Fe_{2+})=0.353-0.928$ | $p_{11}<0.050$ |
| | $Fe_{3+}/Fe_T=0.039-0.699$ | |

Table 2SB. Test of the new amphibole-plagioclase NaSi–CaAl exchange thermometric expressions. Testing model1: $\hat{T}_i = a + b T_i$

| <i>Expression</i> | <i>A1</i> | <i>A1</i> | <i>B1</i> | <i>B1</i> | <i>A2</i> | <i>A2</i> | <i>A2</i> | <i>B2</i> | <i>B2</i> | <i>B2</i> |
|---------------------|--------------------------|--------------------------|-------------------------|-----------------------|--------------------------|--------------------------|--------------------------|--------------------------|--------------------------|--------------------------|
| <i>Data set2</i> | <i>CDS1</i> | <i>FDS</i> | <i>CDS1</i> | <i>FDS</i> | <i>CDS2</i> | <i>TDS2</i> | <i>FDS</i> | <i>CDS2</i> | <i>TDS2</i> | <i>FDS</i> |
| <i>Observations</i> | 196 | 203 | 196 | 203 | 92 | 99 | 203 | 92 | 99 | 203 |
| <i>R-squared</i> | 0.99 | 0.99 | 0.74 | 0.69 | 0.99 | 0.99 | 0.99 | 0.99 | 0.99 | 0.99 |
| <i>RMSE (°C)</i> | 50 | 55 | 49 | 54 | 35 | 51 | 55 | 36 | 52 | 56 |
| <i>AAD (°C)</i> | 37 | 40 | 37 | 40 | 26 | 42 | 39 | 26 | 42 | 40 |
| <i>MAD (°C)</i> | 27 | 29 | 28 | 30 | 18 | 35 | 29 | 21 | 35 | 28 |
| <i>Slope, b</i> | 1.000 ±0.004 t=250 | 0.995 ±0.005 t=199 | 0.85 ±0.04 t=21.3 | 0.84 ±0.04 t=21 | 0.996 ±0.004 t=249 | 1.014 ±0.006 t=169 | 1.000 ±0.005 t=200 | 0.998 ±0.005 t=200 | 1.013 ±0.006 t=169 | 0.999 ±0.005 t=200 |
| <i>Intercept, a</i> | 0 | 0 | 124 ±31 t=4.0 | 135 ±33 t= 4.1 | 0 | 0 | 0 | 0 | 0 | 0 |

Notes: 1, non-constant regression model used when $P(>/t/) > 0$ for a parameter in model $\hat{T}_i = a + b T_i$. 2: CDS1, Calibration Data Set 1, for expressions A1 and B1 (i.e., full data set excluding 7 outliers); CDS2, Calibration Data Set 2, for expressions A2 and B2; FDS, Full Data Set; TDS2, Test Data Set 2, for expression A2 and B2. Other abbreviations: AAD, Average Absolute Deviation; MAD, Median Absolute Deviation; RMSE, Root-Mean-Square Error.

Table 3SB. Test of the amphibole-plagioclase NaSi–CaAl exchange thermometer from Holland and Blundy (1994): expression B. Testing model₁: $\hat{T}_i = a + b T_i$

| <i>Data set</i> ₂ | <i>FDS</i> | <i>FDS</i> | <i>CDS2</i> | <i>CDS2</i> | <i>TDS2</i> | <i>TDS2</i> |
|---------------------------------------|-----------------------|-----------------------|-----------------------|-----------------------|------------------------|-----------------------|
| <i>Temperature range</i> ₃ | <i>FTR</i> | <i>CTR</i> | <i>FTR</i> | <i>CTR</i> | <i>FTR</i> | <i>CTR</i> |
| <i>Observations</i> | 201 | 158 | 91 | 70 | 98 | 76 |
| <i>R-squared</i> | 0.52 | 0.40 | 0.64 | 0.60 | 0.46 | 0.99 |
| <i>RMSE (°C)</i> | 60 | 61 | 52 | 51 | 67 | 72 |
| <i>AAD (°C)</i> | 56 | 53 | 48 | 42 | 60 | 58 |
| <i>MAD (°C)</i> | 42 | 41 | 37 | 35 | 49 | 49 |
| <i>Slope, b</i> | 0.660±0.045 t = 15 | 0.643±0.063 t = 10 | 0.665±0.053 t = 13 | 0.711±0.070 t = 10 | 0.674±0.075 t = 9.0 | 0.970±0.010 t = 97 |
| <i>Intercept, a</i> | 255±37 t = 6.9 | 268±50 t = 5.4 | 261±45 t = 5.8 | 226±56 t = 4.0 | 239±62 t = 3.9 | 0 |

Notes: 1, non-constant regression model used when $P(>/t/) > 0$ for a parameter in model $\hat{T}_i = a + b T_i$. 2: CDS2, Calibration Data Set 2; FDS, Full Data Set; TDS2, Test Data Set 2. 3: CTR, Calibration Temperature Range; FTR, Full Temperature Range.
Other abbreviations: AAD, Average Absolute Deviation; MAD, Median Absolute Deviation; RMSE, Root-Mean-Square Error.

Table 4SB. Test of the amphibole-only thermometer from Ridolfi and Renzulli (2012): expression 2. Testing model1: $\hat{T}_i = a + b T_i$

| <i>Data set2</i> | <i>FDS</i> | <i>FDS</i> | <i>CDS2</i> | <i>CDS2</i> | <i>TDS2</i> | <i>TDS2</i> |
|---------------------------|-----------------------|------------------------|------------------------|----------------------|------------------------|------------------------|
| <i>Temperature range3</i> | <i>FTR</i> | <i>CTR</i> | <i>FTR</i> | <i>CTR</i> | <i>FTR</i> | <i>CTR</i> |
| <i>Observations</i> | 166 | 123 | 84 | 60 | 75 | 57 |
| <i>R-squared</i> | 0.39 | 0.25 | 0.47 | 0.16 | 0.41 | 0.48 |
| <i>RMSE (°C)</i> | 51 | 44 | 51 | 42 | 43 | 33 |
| <i>AAD (°C)</i> | 65 | 45 | 79 | 59 | 51 | 33 |
| <i>MAD (°C)</i> | 49 | 36 | 67 | 54 | 34 | 25 |
| <i>Slope, b</i> | 0.436±0.043 t = 10 | 0.461±0.072 t = 6.4 | 0.489±0.057 t = 8.6 | 0.33±0.10 t = 3.3 | 0.407±0.057 t = 7.1 | 0.526±0.073 t = 7.2 |
| <i>Intercept, a</i> | 520±36 t = 14 | 499±63 t = 7.9 | 495±48 t = 10 | 641±89 t = 7.2 | 527±48 t = 11 | 421±65 t = 6.5 |

Notes: 1, non-constant regression model used when $P(>/t/) > 0$ for a parameter in model $\hat{T}_i = a + b T_i$. 2: CDS2, Calibration Data Set 2; FDS, Full Data Set; TDS2, Test Data Set 2. 3: CTR, Calibration Temperature Range; FTR, Full Temperature Range.
Other abbreviations: AAD, Average Absolute Deviation; MAD, Median Absolute Deviation; RMSE, Root-Mean-Square Error.

Table 5SB. Test of the amphibole-only thermometer from Putirka (2016): expression 5. Testing model: $\hat{T}_i = a + b T_i$

| <i>Data set</i> ² | <i>FDS</i> | <i>FDS</i> | <i>CDS2</i> | <i>CDS2</i> | <i>TDS2</i> | <i>TDS2</i> |
|---------------------------------------|------------------------|-----------------------|------------------------|------------------------|------------------------|-----------------------|
| <i>Temperature range</i> ³ | <i>FTR</i> | <i>CTR</i> | <i>FTR</i> | <i>CTR</i> | <i>FTR</i> | <i>CTR</i> |
| <i>Observations</i> | 152 | 132 | 65 | 52 | 76 | 70 |
| <i>R-squared</i> | 0.35 | 0.47 | 0.33 | 0.38 | 0.52 | 0.61 |
| <i>RMSE (°C)</i> | 52 | 48 | 46 | 41 | 48 | 43 |
| <i>AAD (°C)</i> | 60 | 45 | 83 | 62 | 41 | 34 |
| <i>MAD (°C)</i> | 41 | 35 | 59 | 50 | 30 | 26 |
| <i>Slope, b</i> | 0.412±0.046 t = 9.0 | 0.563±0.053 t = 11 | 0.295±0.053 t = 5.6 | 0.358±0.065 t = 5.5 | 0.605±0.067 t = 9.0 | 0.749±0.072 t = 10 |
| <i>Intercept, a</i> | 513±37 t = 14 | 382±44 t = 8.7 | 632±44 t = 14 | 578±55 t = 11 | 341±55 t = 6.2 | 218±59 t = 3.7 |

Notes: 1, non-constant regression model used when $P(>/t/) > 0$ for a parameter in model $\hat{T}_i = a + b T_i$. 2: CDS2, Calibration Data Set 2; FDS, Full Data Set; TDS2, Test Data Set 2. 3: CTR, Calibration Temperature Range; FTR, Full Temperature Range.
Other abbreviations: AAD, Average Absolute Deviation; MAD, Median Absolute Deviation; RMSE, Root-Mean-Square Error.

Table 6SB. Test of the amphibole-only thermometer from Putirka (2016): expression 6. Testing model: $\hat{T}_i = a + b T_i$

| <i>Data set</i> ² | <i>FDS</i> | <i>FDS</i> | <i>CDS2</i> | <i>CDS2</i> | <i>TDS2</i> | <i>TDS2</i> |
|---------------------------------------|------------------------|-----------------------|------------------------|------------------------|------------------------|-----------------------|
| <i>Temperature range</i> ³ | <i>FTR</i> | <i>CTR</i> | <i>FTR</i> | <i>CTR</i> | <i>FTR</i> | <i>CTR</i> |
| <i>Observations</i> | 152 | 132 | 65 | 52 | 76 | 70 |
| <i>R-squared</i> | 0.31 | 0.44 | 0.29 | 0.33 | 0.49 | 0.60 |
| <i>RMSE (°C)</i> | 53 | 48 | 46 | 41 | 47 | 42 |
| <i>AAD (°C)</i> | 61 | 46 | 87 | 65 | 42 | 33 |
| <i>MAD (°C)</i> | 41 | 33 | 63 | 53 | 31 | 29 |
| <i>Slope, b</i> | 0.383±0.046 t = 8.3 | 0.533±0.053 t = 10 | 0.273±0.053 t = 5.2 | 0.322±0.065 t = 5.0 | 0.559±0.066 t = 8.5 | 0.707±0.069 t = 10 |
| <i>Intercept, a</i> | 539±38 t = 14 | 409±44 t = 9.3 | 656±44 t = 15 | 613±55 t = 11 | 377±54 t = 7.0 | 250±57 t = 4.4 |

Notes: 1, non-constant regression model used when $P(>|t|) > 0$ for a parameter in model $\hat{T}_i = a + b T_i$. 2: CDS2, Calibration Data Set 2; FDS, Full Data Set; TDS2, Test Data Set 2. 3: CTR, Calibration Temperature Range; FTR, Full Temperature Range.

Other abbreviations: AAD, Average Absolute Deviation; MAD, Median Absolute Deviation; RMSE, Root-Mean-Square Error.

Table 7SB. Application of amphibole-plagioclase and amphibole-only thermometry to high-grade metamorphic rocks (see Appendix E for details)

| Metamorphic complex | Sample description | Geological setting |
|---|--|--|
| Berit meta-ophiolite, Anatolia (SE Turkey) | <i>Rock types:</i> amphibolites, Grt-Cpx amphibolites, Pl-rich non-coronitic granulites. <i>Compositions used in the calculations:</i> magnesiohornblende, pargasite, tschermakite ($Mg = 2.8\text{--}4.0$ apfu); oligoclase-bytownite ($p_{an} = 0.26\text{--}0.77$). <i>Temperatures calculated at 5–8 kbar</i> (amphibolite and Grt-Cpx amphibolite) and 11–15 kbar (Pl-rich non coronitic granulite). <i>Data source:</i> Awalt and Whitney (2018). | Late Mesozoic to Cenozoic ophiolite fragments that mark the suture of the Arabian and Eurasian plates appear in Anatolia. The Berit meta-ophiolite represents a rare example of metamorphosed ophiolitic rocks in the SE Anatolian orogenic belt. It contains Grt (\pm Cpx) amphibolites, granulite-facies metagabbros, which appear as lenses in amphibolites (formed after granulites) and metaperidotites with minor metapyroxenite layers. In the Doğanşir region, granulite-facies gabbroic rocks contain Grt+Cpx+Pl+Crn \pm Opx \pm Ky \pm Spr \pm Rt with coronitic and non-coronitic textures, and, in some samples, symplectites. <i>References:</i> Genç et al. (1993), Karaoğlu et al. (2013), Robertson et al. (2012), Awalt and Whitney (2018). |
| Amphibolites from the Fiskensæset complex (SW Greenland) | <i>Rock types:</i> amphibolites, high-Cr amphibolites. <i>Compositions used in the calculations:</i> ferropargasite, magnesiohornblende, edenite ($Mg = 1.7\text{--}2.6$ apfu); oligoclase-bytownite ($p_{an} = 0.29\text{--}0.79$). <i>Temperatures calculated at 7 kbar.</i> <i>Data source:</i> Weaver et al. (1982). | The Fiskensæset complex contains one of the best preserved Archean layered intrusions in the world. It consists of layered anorthosites, leucogabbros, gabbros, peridotites and chromitites that appear as dissected layers and lenses within amphibolites and quartzofeldspathic rocks. The complex was deformed and metamorphosed at amphibolite- to granulite- facies conditions at ca. 2.8 Ga. The amphibolites appear as up to ca. 1 km thick layers mostly limited by anorthosites, and locally by mica schists and marbles. <i>References:</i> Windley et al. (1973), Escher and Myers (1975), Rivalenti (1976), Page et al. (1980), Myers (1981), Weaver et al. (1981, 1982), Myers (1985), Ashwal et al. (1989), Riciputi et al. (1990), Peck and Valley (1996), Keulen et al. (2009), Polat et al. (2011), Polat and Longstaffe (2014), Polat (2015). |
| Kvalvåg granulites, Kristiansund Area, Western Gneiss Region (WGR, Norwegian Caledonides). | <i>Rock types:</i> Grt granulites. <i>Compositions used in the calculations:</i> pargasite ($Mg = 2.64$ apfu); andesine ($p_{an} = 0.31\text{--}0.33$). <i>Temperatures calculated at 13.1 kbar.</i> <i>Data source:</i> Krogh (1980). | The WGR presents the lowest exposed structural level in the southern Scandinavian Caledonides. It mostly consists of Proterozoic granitoid migmatitic and augen orthogneisses with, also common, ultramafic rocks, meta-anorthosites, mangerites, pelitic schists, quartzites, calc-silicates and marbles. Eclogites usually appear as lenticular pods or tabular layers within most of the other rock types. Based on the distribution of coesite or polycrystalline aggregates after coesite, two provinces, HP and UHP eclogites, were distinguished. In Kristiansund and other localities from the northern part of the WGR, eclogites tend to be overprinted by garnet granulite-facies assemblages, appearing trondhjemitic segregations in some eclogite bodies. <i>References:</i> Bryhni (1966), Krogh (1980, 1982), Smith (1984), Griffin et al. (1985), Mørk (1985), Jamtveit (1987), Krogh and Carswell (1995), Austrheim et al. (1997), Wain (1997), Austrheim (1998), Cuthbert et al. (2000), Austrheim et al. (2003). |
| Connaughton Terrane, Paleoproterozoic Rudall Complex, (Paterson Orogen, central Western Australia). | <i>Rock types:</i> Opx-Cpx amphibolites and Cpx-poor amphibolites. <i>Compositions used in the calculations:</i> pargasite, magnesiohornblende ($Mg = 2.27\text{--}3.36$ apfu); andesine-bytownite ($p_{an} = 0.36\text{--}0.82$). <i>Temperatures calculated at 7 kbar.</i> <i>Data source:</i> Smithies and Bagas (1997). | The Paterson Orogen is an arcuate, 2000 km long, belt that separates the West Australian Craton and North Australian Craton. It consists of Palaeoproterozoic to Neoproterozoic metasedimentary and meta-igneous rocks affected by the ca. 550 Myr Paterson (probably equivalent Petermann) orogeny. The Rudall Complex is exposed in the northwestern Paterson Orogen. It consists of deformed and metamorphosed Palaeoproterozoic and Mesoproterozoic sedimentary and igneous rocks. This complex has been subdivided into the Talbot, Connaughton and Tabletop terranes. The Connaughton terrane comprises a succession of mafic gneisses and schists, and paragneisses metamorphosed at the amphibolite–granulite-facies transition. It contains Opx-Cpx amphibolites, metamorphosed at granulite-facies conditions, and Cpx-poor amphibolites, generated by retrogression of the Opx-Cpx amphibolites. <i>References:</i> Williams (1990), Williams and Myers (1990), Perincek (1996), Smithies and Bagas (1997), Bagas (2004), Aitken et al. (2018). |

Table 8SB. Application of amphibole-plagioclase and amphibole-only thermometry to igneous rocks (see Appendix E for details)

| <i>Igneous association</i> | <i>Sample description</i> | <i>Geological setting</i> |
|--|--|---|
| Appinite suite from the Variscan Avila batholith (Central Iberian Zone) | <i>Rock types:</i> cortlandites. <i>Compositions used in the calculations:</i> pargasite ($Mg = 3.4\text{--}3.6$ apfu); andesine-labradorite ($p_{an} = 0.49\text{--}0.51$). <i>Temperatures calculated at</i> 3–4 kbar. <i>Data source:</i> Molina et al. (2009). | The Avila Batholith is exposed along the axis of the Central Iberian Zone that represents the most internal paleogeographic domain of the Variscan belt of Iberia. The batholith contains 325–295 Myr plutons of strongly peraluminous granodiorites to monzogranites, blocks of migmatites (age: 352–297 Myr), minor appinitic ultramafic-mafic to intermediate stocks dated at 319–300 Myr and late Permian lamprophyre dykes. The appinites appear as enclaves included in peraluminous granitoids or intruding migmatites. The largest bodies are zoned, with cortlandites and high-Mg gabbros at the core, and tonalites and granodiorites at the border that develop a brecciated mixed facies with mingling structures. <i>References:</i> Moreno-Ventas et al. (1995), Bea et al. (1999, 2003), Castro et al. (2003), Bea (2004), Bea et al. (2004), Montero et al. (2004a, b), Bea et al. (2006), Molina et al. (2009), Orejana et al. (2009), Scarrow et al. (2009), Bea (2010), Díaz-Alvarado et al. (2011), Villaseca et al. (2011), Molina et al. (2012), Rodríguez and Castro (2019). |
| Barcroft granodioritic pluton, northern White-Inyo Mountains (east-central California) | <i>Rock types:</i> gabbro-diorites, metadiorites, granodiorites. <i>Compositions used in the calculations:</i> magnesiohastingsite, magnesiohornblende ($Mg = 2.6\text{--}3.3$ apfu); andesine-bytownite ($p_{an} = 0.32\text{--}0.83$). <i>Temperatures calculated at</i> 3.8–6.6 kbar (stage 1), 1.5–6 kbar (stage 2) and 0.7–3 kbar (stage 3). <i>Data source:</i> Ernst (2002). | The White-Inyo Range lies between the Sierra Nevada fault block and the Basin and Range provinces. The 167–161 Myr Barcroft pluton is a tabular body that cuts the NNW structural trend of the White-Inyo Range and invades the Barcroft break, a NE-striking, SE-dipping high-angle reverse fault. It consists of Hbl- and Bio-rich (\pm relict Cpx)-bearing granodiorites, quartz gabbros and rare alaskites. Four stages of crystallization are distinguished: cumulate stage 1 at 3.8–6.6 kbar, intrusive stages 2 and 3 at 0.3–6 kbar and the deuteric stage 4 at 1–2 kbar. <i>References:</i> Ernst and Hall (1987), McKee and Conrad (1996), Ernst (2002), Ernst et al. (2003), Ernst and Rumble (2003), Michelsen (2003), Ferré et al. (2012), Ernst (2013), Chapman et al. (2015). |
| Bezmyannyi volcano (Kamchatka) | <i>Rock types:</i> andesites from the catastrophic 1956 eruption. <i>Compositions used in the calculations:</i> pargasite, tschermakite, magnesiohastingsite ($Mg = 2.7\text{--}2.8$ apfu); andesine-labradorite ($p_{an} = 0.37\text{--}0.66$). <i>Temperatures calculated at</i> 6–8 kbar. <i>Data source:</i> Almeev et al. (2002). | The Bezmyannyi volcano is located in the central part of the Klyuchevskaya volcanic group, southwest of Klyuchevskii volcano. It is the only active andesite volcano of this group; it experienced extremely powerful eruptions in 1956 and 1985. The andesites of Bezmyannyi volcano contain hornblende phenocrysts with opacite rims and inclusions of labradorite-andesite. <i>References:</i> Tolstykh et al. (1999), Almeev et al. (2002), Plechov et al. (2008), Ozerov et al. (2020). |
| Sheeted sills at Onion Valley (Mesozoic Sierra Nevada batholith) | <i>Rock types:</i> Ol hornblendites. <i>Compositions used in the calculations:</i> magnesiohastingsite, pargasite ($Mg = 3.2\text{--}3.3$ apfu); bytownite ($p_{an} = 0.86\text{--}0.90$). <i>Temperatures calculated at</i> 2–4 kbar. <i>Data source:</i> Sisson et al. (1996) | The Sierra Nevada batholith represents a very thick amalgamation of granitoid plutons dated at 248–80 Myr. It contains mostly granodiorites and granites, with minor mafic intrusions. Near Onion Valley, a mafic intrusive complex is exposed consisting of sheeted sills of Hbl gabbros to Hbl diorites, multiply-injected sills, mafic stocks and lens-shaped cumulate intrusions with layers of Hbl gabbros with minor Ol hornblendites, Pl hornblendites, and Hbl-Pl pegmatites. <i>References:</i> Bateman and Dodge (1970), Pickett and Saleeby (1993), Coleman et al. (1995), Sisson et al. (1996), Coleman and Glazner (1997), Ducea (2001), Lackey et al. (2008), Nadin and Saleeby (2008), Dumitru et al. (2010), Cecil et al. (2012), Gilbert et al. (2012). |
| Val Fredda Complex, Southern Adamello Massif (Central Alps, Italy). | <i>Rock types:</i> Hbl gabbros and diorite enclaves. <i>Compositions used in the calculations:</i> tschermakite, magnesiohornblende, actinolite ($Mg = 2.5\text{--}3.5$ apfu); andesine-labradorite ($p_{an} = 0.42\text{--}0.62$). <i>Temperatures calculated at</i> 2 kbar. <i>Data source:</i> Blundy and Holland (1990). | The Adamello batholith comprises several intrusive units dated at 42–29 Myr; it consists of tonalites, trondjemites and granodiorites, with minor mafic and ultramafic intrusions. The Val Fredda Complex contains granodiorites, tonalites and quartz diorites intruded by mafic sheets that range from Hbl gabbros to quartz diorites in the Mt. Cadino and Hbl-phyric, Px-bearing gabbros and subordinate hornblende blocks or bands in the Mt. Mattoni. <i>References:</i> Bianchi et al. (1970), Callegari and Dal Piaz (1973), Dupuy et al. (1982), Del Moro et al. (1983), Riklin (1983), Ulmer et al. (1983), Blundy and Holland (1990), Blundy and Sparks (1992), Callegari and Brack (2002), Mayer et al. (2003), Tiepolo et al. (2011). |

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