

Graphical analysis of the orthopyroxene-pigeonite-augite-plagioclase equilibrium at liquidus temperatures and low pressure

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ABSTRACT

There are both natural and experimental observations of the coexistence of three pyroxenes—orthopyroxene (Opx) + pigeonite (Pig) + augite (Aug)—with plagioclase (Pl). Commonly, the assemblage occurs as an intermediate product in the following fractionation trend of a mafic magma: $\text{Opx} + \text{Aug} + \text{Pl} \rightarrow \text{Opx} + \text{Aug} + \text{Pig} + \text{Pl} \rightarrow \text{Aug} + \text{Pig} + \text{Pl}$. To clarify the phase-equilibria constraints on the existence of this mineral assemblage, we have graphically analyzed the change in topology of an isobaric–isoplethic section, Ol-Aug-Pl-Qtz [with $fe = 25\text{--}50\%$, where $fe = \text{Fe}/(\text{Fe} + \text{Mg})$, and $\text{An} = 50\%$], arising from an increase in the fe -value of silicate liquid. The analysis shows that the stability field of the mineral assemblage $\text{Opx} + \text{Aug} + \text{Pig} + \text{Pl}$ is restricted, and can only crystallize in the interval between two invariant points— T_1^4 (Ol + Opx + Pig + Aug + Pl + L) and T_2^4 (Qtz + Opx + Pig + Aug + Pl + L)—that emerge successively during expansion of a liquidus volume of pigeonite within the isobaric–isoplethic section Ol-Aug-Pl-Qtz. At fe -values lower than T_1^4 , the 3-pyroxene assemblage cannot exist due to the absence of a contact surface between the primary volumes of plagioclase and pigeonite. At fe -values greater than T_2^4 , the assemblage is unstable due to separation of the primary volumes of orthopyroxene and augite.