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2 Twinning in hydrous wadsleyite: symmetry relations, origin, and consequences

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14 Abstract:

Twins in hydrous wadsleyite were detected by polarized-light microscopy and characterized with transmission electron microscopy techniques including precession selected area electron diffraction and large-angle convergent beam diffraction. By inspecting diffracted intensities for high-order Laue zones, we found the symmetry of our hydrous wadsleyite samples to be reduced to monoclinic with respect to the orthorhombic symmetry of most anhydrous wadsleyite samples.

20 Twinned domains in hydrous wadsleyite share the (122) plane as composition plane and are related to each other by a two-fold rotation around a twin axis parallel to [212] or by reflection 21 22 on (122). The twin axis and twin plane in wadsleyite correspond to the <101> directions and the {101} planes of ringwoodite, respectively. The twin operations exchange the c^* and the [120]* 23 directions of wadsleyite, both of which correspond to the directions of the cubic a axes in 24 ringwoodite. Based on our analysis of symmetry relations and pseudo-symmetry in wadsleyite, 25 we conclude that the twins formed during crystal growth under quasi-hydrostatic conditions in 26 the presence of a hydrous fluid. Twinning in wadsleyite may affect the physical properties and 27 deformation behavior of wadsleyite as well as the phase transition to ringwoodite in the Earth's 28 mantle transition zone. 29

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31 Keywords

32 Twinning, hydrous wadsleyite, transmission electron microscopy, Earth's mantle transition zone

34 Introduction

As a high-pressure polymorph of the mineral olivine, wadsleyite $(\beta - (Mg, Fe)_2 SiO_4)$ is 35 expected to be a major mineral in peridotitic lithologies in the upper part of the transition zone of 36 Earth's mantle at depths between 410 and 520 km (e.g., Frost, 2008). Despite numerous studies 37 on the physical properties (e.g., Buchen et al., 2017; Zhou et al., 2021) and crystal-chemical 38 characteristics (e.g., Smyth et al., 2014; Kawazoe et al., 2016) of wadsleyite, little is known 39 about the morphology of wadslevite. Morphological properties, such as crystal shape, cleavage, 40 and twinning, however, can impact the physical response of polycrystalline aggregates, in 41 particular when interacting with the anisotropic properties of individual grains or domains. Here 42 we report on the discovery of twinning in hydrous wadsleyite and determine the twin law by 43 combining transmission electron microscopy with a theoretical analysis of symmetry relations. 44

Anhydrous wadsleyite most frequently crystallizes with orthorhombic symmetry (space-45 group type Imma, Moore and Smith, 1970; Horiuchi and Sawamoto, 1981). The crystal structure 46 of wadsleyite (wds) is closely related to the cubic spinel structure of ringwoodite (rwd), another 47 high-pressure polymorph of olivine. Both structures can be described by different stacking 48 sequences of structural modules along the [010]_{wds} or [110]_{rwd} directions (Price, 1983; Price et al., 49 1983; Putnis, 1992). In the presence of H₂O, wadsleyite incorporates hydrogen in the form of 50 hydroxyl groups (Inoue et al., 1995; Kohlstedt et al., 1996; McKay et al., 2019). The 51 incorporation of hydroxyl groups into wadsleyite and the related formation of cation vacancies 52 may reduce the crystal symmetry of wadsleyite to monoclinic (space-group type I2/m) with $\beta \neq 12/m$ 53 90° (Smyth et al., 1997). This change in symmetry can be described by the loss of the mirror 54 plane perpendicular to the *a* axis due to partial ordering of vacancies on octahedral sites (Smyth 55 et al., 1997; Kudoh and Inoue, 1999). Holl et al. (2008) reported broadening and possible peak 56

splitting of the 244 diffraction peak of a hydrous wadsleyite crystal by performing omega scans 57 along the a^* , b^* , and c^* directions on a four-circle diffractometer (Figure 7 in Holl et al., 2008). 58 They attributed the observed peak splitting to polysynthetic twinning, but did not specify any 59 twin relations. The existence of twins in wadsleyite, however, has not yet been demonstrated by 60 using electron microscopy techniques. Here we report, for the first time, twinning in synthetic 61 hydrous Fe-bearing wadslevite. We first detected the twins in thin sections of wadslevite samples 62 by the change in color across twin boundaries, which is enhanced by the strong pleochroism of 63 64 Fe-bearing wadsleyite when viewed under a polarized-light microscope (Fig. 1). To further characterize the twinning, we used transmission electron microscopy (TEM) assisted with 65 electron diffraction and scanning transmission electron beam techniques. Based on our 66 observations, we derive the twin law and discuss the origin of the twins in hydrous wadslevite. 67

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69 Experimental procedures

70 Sample synthesis at high pressure and temperature

Polycrystalline wadsleyite samples were synthesized from a powder made from San 71 Carlos olivine and saturated with distilled water. The wet powder was filled into a platinum 72 capsule, compressed to a pressure of 15 GPa, and held at a temperature of 1670 K for 3 h using a 73 74 Kawai-type multi-anvil apparatus at the Bayerisches Geoinstitut, Universität Bayreuth. Further details about the synthesis procedure have been described in Kawazoe et al. (2015). The obtained 75 76 crystals were several hundred micrometers in length, but less than 1 mm. The ratio Fe/(Mg+Fe) was determined by electron microprobe analysis to be 0.097(1). The water content was 77 78 determined by infrared spectroscopy (sample H4183 in Buchen, 2018) and corresponds to about

79 0.72 wt% H₂O (equivalent to 0.058 H₂O molecules per formula unit). The unit cell parameters of 80 a twinned wadsleyite crystal (Fig. 1b) were determined from 33 X-ray reflections, which were 81 centered on a four-circle diffractometer with a point detector and a Mo $K\alpha$ X-ray source operated 82 at 40 kV and 30 mA. We refined the unit cell parameters assuming first orthorhombic and then 83 monoclinic symmetry (**Table 1**).

84

Transmission electron microscopy

Two thin foils, one of which perpendicular to the twin boundary, were cut from a twinned 85 wadsleyite crystal with a dual-beam focused ion beam (FIB) milling machine (FEI Scios) (Fig. 86 2). The two extracted TEM thin foils were oriented normal to the $[2\overline{1}0]_{wds}$ (<100>_{rwd}) and 87 $[001]_{wds}$ (<001>_{rwd}) zone axes, respectively, and were inspected with scanning transmission 88 electron microscopes (FEI Titan G2 80-200 S/TEM and Philips CM20FEG) operating at 200 kV. 89 The foil perpendicular to the $[2\overline{1}0]_{wds}$ zone axis was prepared with the intention of identifying 90 91 the twin relation by using precession selected area electron diffraction (PsSAED) and large-angle convergent beam diffraction (LACBED) techniques. The foil perpendicular to the [001]_{wds} zone 92 axis was intended to investigate a potential symmetry reduction from the orthorhombic (point 93 group *mmm*) to the monoclinic structure (2/m) (Smyth et al., 1997; Holl et al., 2008) by using 94 95 high-order Laue zones (HOLZ) in the selected area electron diffraction (SAED) pattern. For comparison, a HOLZ pattern of the $[\overline{1}10]$ zone axis of olivine was recorded with the same 96 97 microscope. Scanning moiré fringe (SMF) imaging by scanning transmission electron microscopy (STEM) (Su and Zhu, 2010) has, for the first time, been applied to a high-pressure 98 99 mineral to visualize a change of moiré fringe pattern across the composition plane at a relatively 100 low electron dose due to lower magnification (a 160 k times magnification), about 70 times 101 lower than that of a conventional magnification (at 1300 k times) under high resolution STEM

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102	conditions (usually at 1000-5000 k times). The chemical homogeneity at the composition plane
103	was investigated in high-resolution TEM (HRTEM) and analytical STEM combined with a high-
104	angle annular dark field (HAADF) detector. The HRTEM image presented was processed using
105	a Wiener filter which was implemented in a Gatan DigitalMicrograph (Mitchell and Schaffer,
106	2005) to remove noise contrast (Kilaas, 1998).

- 107
- 108 **Results**

109 Investigation of the twin relation by transmission electron microscopy

110 The TEM observations assisted with SAED revealed that the orientations of twin domains can be brought into coincidence by a 180° rotation about an axis nearly perpendicular to 111 the $(122)_{wds}$ plane in wadsleyite or by reflection on the $(122)_{wds}$ plane, which corresponds to one 112 of the $\{101\}_{rwd}$ planes in the cubic system, i.e., a cubic axis system attached to structural units in 113 common with the cubic structure of ringwoodite (Figs. 3 and 8 and Supplemental Information 114 Figs. S1 and S2). Therefore, potential twin elements include a two-fold twin axis nearly 115 116 perpendicular to the (122)_{wds} plane and a mirror plane parallel to (122)_{wds}, or to one of the {122}_{wds} planes in orthorhombic symmetry and hereafter denoted as (122)_{wds} for simplicity. The 117 118 SAED patterns from twin-related domains display these symmetry relations (Supplemental Figs. S1 and S2). SAED patterns of the twin boundary recorded along the $[2\overline{1}0]_{wds}$ zone axis show 119 that 244 reflections (and higher orders) are split along the direction perpendicular to the twin 120 121 boundary (Fig. 3). For non-zero obliquity, reflections are typically split along a direction normal to the composition plane, i.e., the plane shared by two adjacent domains. The splitting of 122 reflections therefore indicates that the composition plane is parallel to one of the {244}_{wds} planes 123

127	wadsleyite crystal (Supplemental Fig. S3).
126	experiments, we observed a similar splitting of 244 reflections in omega scans of a twinned
125	244 reflections reported by Holl et al. (2008). During our single-crystal X-ray diffraction
124	and hence parallel to $(122)_{wds}$. This observation is consistent with the splitting and broadening of

Apart from the splitting of 244 diffraction spots in SAED patterns, we did not observe significant orientational misfits or remarkable discontinuities of the Bragg lines in the LACBED pattern recorded in the vicinity of the composition plane (**Fig. 4**). With the exception of a visible discontinuity along the composition plane, we did not detect visible structural defects such as dislocation arrays or chemical differences in the vicinity of or across the twin boundary in the conventional high-resolution lattice TEM image (**Fig. 5a**), by SMF-STEM (**Figs. 5b and 5c**), or by HAADF-STEM imaging (**Supplemental Fig. S4**).

135 Space-group identification and symmetry reduction

136 The electron diffraction pattern along the [001]_{wds} zone axis is crucial for identifying a potential reduction from orthorhombic to monoclinic symmetry (space-group type I2/m, b as 137 unique axis). In the diffraction pattern, the reflection conditions for hk0 reflections and the 138 existence of a mirror plane perpendicular to the *a* axis should be verified to distinguish between 139 monoclinic and orthorhombic symmetries. Reflections for $\mathbf{g} = hk0$: h, k = 2n + 1 are forbidden by 140 141 the orthorhombic space group Imma but not by the monoclinic space group I2/m reported for 142 hydrous wadsleyite by Smyth et al. (1997). Figure 6a displays SAED patterns of the [001]_{wds} zone axis without any detectable intensities of hk0 reflections with h, k = 2n + 1, while the 143 intensity of the 110 reflection was detected in a systematic diffraction condition along the [110]* 144 145 direction (Fig. 6b). In Fig. 6b, however, some dynamical diffraction effects, e.g., double

diffraction, might be visible. To test for mirror symmetries along the *a* and *b* axes, the intensity 146 distributions of high-order Laue zones for the [001]_{wds} zone axis were also examined to confirm 147 the absence of a mirror plane perpendicular to the *a* axis by using SAED patterns with smaller 148 149 camera lengths (Fig. 7 and Supplemental Fig. S5). In the HOLZ pattern of the [001]_{wds} zone axis using a reduced camera length of 340 mm, the intensity distributions along the a^* axis are 150 151 asymmetric, while those along the b axis are symmetric, which points to the absence of a mirror plane perpendicular to the *a* axis, consistent with the monoclinic space group I2/m (unique axis 152 153 b) but not with the orthorhombic space group Imma, i.e., I 2/m 2/m 2/a. In addition, the intensity distributions of the Bragg lines in a LACBED pattern of the [001]_{wds} zone axis were also 154 examined to test for the existence of a mirror plane perpendicular to the *a* axis (Supplemental 155 156 Fig. S6). In conclusion, our detailed analysis of diffracted intensities with TEM suggests our wadsleyite samples to be monoclinic, consistent with the significant deviation of the angle β 157 from 90° as detected by single-crystal X-ray diffraction (Table 1). 158

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160 Discussion

161 Symmetry analysis of twins in hydrous wadsleyite and derivation of the twin law

Our TEM observations suggest that individual twin domains in hydrous wadsleyite are related to each other by reflection on $(122)_{wds}$ or by a two-fold rotation about a twin axis nearly perpendicular to one of the $(122)_{wds}$ planes. Mallard's law requires the twin axis to be a lattice element in direct space. Since the direction $[212]_{wds}$, or an equivalent thereof, is the lattice row closest to be perpendicular to a $(122)_{wds}$ plane, this direction is a good candidate for the orientation of the twin axis. Indeed, exploring potential twin operations in the realm of the

reticular theory of twinning (Grimmer & Nespolo, 2006) by running the program 168 GEMINOGRAPHY (Nespolo & Ferraris, 2006) on the monoclinic unit cell given in Table 1 169 proposes the pair of twin elements (122)_{wds} and [212]_{wds}. According to this theory, twinning in 170 171 hydrous wadsleyite is by reticular pseudo-merohedry or, more specifically, by reticular pseudo-172 polyholohedry with the unique axis of the *I*-centered pseudo-monoclinic twin lattice (TL) oriented along the [212]_{wds} direction of wadsleyite. The twin lattice can be described by the 173 parameters $a_{\text{TL}} = 14.076$ Å, $b_{\text{TL}} = 23.137$ Å, $c_{\text{TL}} = 14.169$ Å, $\alpha_{\text{TL}} = 89.26^{\circ}$, $\beta_{\text{TL}} = 110.06^{\circ}$, $\gamma_{\text{TL}} = 100.06^{\circ}$, $\gamma_{\text{TL}} =$ 174 175 91.25° . The orientation of the axes spanning the twin lattice expressed relative to the crystal axes of wadsleyite is given by the matrix: 176

$$\mathbf{T} = \begin{bmatrix} 2 & 2 & 0 \\ 0 & 1 & -1 \\ -1 & 2 & 1 \end{bmatrix}$$

The second column of the matrix **T** is the direction of the twin axis in direct space expressed in coordinates of the crystal axis system of wadsleyite, the first and third columns are the shortest independent vectors parallel to the $(122)_{wds}$ twin plane, respectively. The associated obliquity, i.e., the orientational mismatch between twin domains, is 1.29° , and the twin index is 8. The theoretical analysis of twin operations is therefore fully consistent with our TEM observations, including the minor orientational mismatch as reflected in the splitting of 244 diffraction spots in SAED patterns.

To derive the twin law, we need to find the cosets of the point group of the twin lattice by decomposition with respect to the intersection group of the point groups of the individual domains in their relative orientations. As the twin operations do not map any of the mirror planes or twofold axes of the orthorhombic point group mmm of one twin domain onto their counterparts in another domain, the intersection group H of the individuals only contains the

identity operation and inversion, $H = \{1, \overline{1}\} = \overline{1}$. The point group of the twin lattice is a 189 190 supergroup of the intersection group and needs to contain the twin operations and reflect the 191 symmetry of the twin lattice. When we neglect the small obliquity and assume the twin lattice to be monoclinic, these conditions are met by the point group G of the twin lattice with $G = \{1, \overline{1}, 2, ..., T\}$ 192 *m*}, i.e., G = 2/m. The coset decomposition of G with respect to H then yields the cosets $\{1, \overline{1}\}$ 193 and $\{2, m\}$, the first of which is the intersection group H. The symmetry operations of reflection 194 on (122)_{wds} and two-fold rotation around [212]_{wds} as detected by our TEM work and confirmed 195 196 by symmetry analysis therefore constitute the twin law.

197

198 *Relation to ringwoodite*

Both a two-fold rotation around $[212]_{wds}$ and reflection on $(122)_{wds}$ map the c^* axis of one 199 twin domain onto the [120]* direction of another twin domain. Each of these directions 200 201 corresponds to one of the cubic a axes $(a_1, a_2, \text{ or } a_3)$ of ringwoodite. Similarly, both twin 202 operations map the $\{101\}_{wds}$ set of planes onto the $\{021\}_{wds}$ set of planes, both of which correspond to the $\{111\}_{rwd}$ planes in ringwoodite. In both minerals, close-packed layers of 203 204 oxygen anions are oriented parallel to the respective families of planes. The twin operations 205 therefore map the oxygen substructure onto itself. The relation between the crystal structures of wadsleyite and ringwoodite is illustrated in Figure 8. A vector w with coordinates w_1, w_2, w_3 206 207 defined with respect to the crystal axes of wadsleyite is transformed to a vector \mathbf{r} in the cubic axis system of ringwoodite by the matrix (see Supplemental Information) 208

$$\mathbf{M}\mathbf{w} = \frac{1}{2} \begin{bmatrix} 1 & 2 & 0 \\ -1 & 2 & 0 \\ 0 & 0 & 2 \end{bmatrix} \begin{pmatrix} w_1 \\ w_2 \\ w_3 \end{pmatrix} = \begin{pmatrix} r_1 \\ r_2 \\ r_3 \end{pmatrix} = \mathbf{r}$$

209 The twin lattice is then related to the crystal axes of ringwoodite by the matrix

$$\mathbf{MT} = \begin{bmatrix} 1 & 2 & -1 \\ -1 & 0 & -1 \\ -1 & 2 & 1 \end{bmatrix}$$

In terms of the axis system of ringwoodite, the twin lattice is spanned by two space diagonals through the unit cell and the face diagonal that is perpendicular to both of them. The two-fold twin axis is oriented along this face diagonal, i.e., along $[101]_{rwd}$. Note that while a two-fold rotation around $[101]_{rwd}$ is not a symmetry operation of the spinel structure of ringwoodite (space-group type $F\bar{4}3m$), reflection on the plane perpendicular to $[101]_{rwd}$, i.e., on $(101)_{rwd}$, is indeed a symmetry operation of ringwoodite. The symmetry characteristics of twins in hydrous wadsleyite might therefore indicate a structural relation to ringwoodite.

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218 Origin of twins in wadsleyite

The twins we detected in wadsleyite may have formed during crystal growth, in response 219 220 to deformation, or may have resulted from a phase transition with a change in point group symmetry. In the previous section, we showed that our wadsleyite samples are strictly 221 222 monoclinic, and a change from orthorhombic to monoclinic symmetry might well result in 223 twinning as suggested by Holl et al. (2008). In this case, at least one of the twin operations 224 should correspond to a symmetry operation of the orthorhombic space group *Imma* that is not 225 present in the monoclinic space group I2/m. We identified a two-fold twin axis parallel to 226 [212]_{wds} and a twin plane parallel to (122)_{wds} as the twin elements. None of the associated symmetry operations, however, is contained in either of the space groups. Alternatively, we may 227 consider twinning to result from a phase transition from ringwoodite $(F\overline{4}3m)$ to wadsleyite 228

(Imma or I2/m) given that both crystal structures share common structural modules. While a two-229 fold rotation around [212]_{wds} or [101]_{rwd} is not a symmetry operation of either of the relevant 230 231 space groups, reflection on $(101)_{rwd}$ is indeed a symmetry operation of ringwoodite. Based on 232 symmetry considerations alone, the twins in hydrous wadsleyite might have formed as a result of 233 the phase transition from ringwoodite to wadsleyite. The pressure and temperature conditions during sample synthesis, however, are inconsistent with the formation of ringwoodite as a 234 precursory phase to wadsleyite as they did not coincide with the stability field of ringwoodite. 235 236 Twinning as a result of a phase transition with a change in point group symmetry does therefore not seem to be a likely explanation for the here-detected twins in wadsleyite. 237

238 Twinning by deformation requires substantial deviatoric stresses in combination with a 239 structural mechanism that allows the crystal structure to fold in response to these stresses. The 240 stress state inside the sample capsule during a multi-anvil experiment can be expected to be close to quasi-hydrostatic, in particular at high temperatures and in the presence of a hydrous fluid or 241 melt as in our synthesis experiments. As a result, the synthesized wadsleyite crystals lack distinct 242 243 deformation features such as dislocation arrays, subgrains, or oriented fractures. As discussed 244 below, adjacent twin domains in hydrous wadsleyite mainly differ in the way the cations are ordered into the interstitial voids formed by the distorted cubic close-packing of oxygen anions, 245 which appears to be continuous across the twin interface. The rearrangement of the cation 246 247 substructures requires diffusive motion of cations rather than a directed displacement of the 248 crystal structure as a whole. The structural relationship between adjacent twin domains therefore 249 does not seem to be consistent with a folding of the crystal structure as expected for mechanical 250 twinning by deformation.

251 Twinning during crystal growth, or primary twinning (Bloss, 1971), does not require a specific structural mechanism or compliance with strict symmetry rules. Instead, growth twins 252 tend to exploit pseudo-symmetries of a crystal structure while conserving or extending 253 254 fundamental structural units across the interface between twin domains (Buerger, 1945). For wadsleyite, layers of close-packed oxygen anions parallel to $\{021\}_{wds}$ and $\{101\}_{wds}$ planes are 255 among the most fundamental structural units (Fig. 8). Both a 180° rotation about an axis parallel 256 to $[212]_{wds}$ and a reflection on a $(122)_{wds}$ plane map $\{021\}_{wds}$ planes onto $\{101\}_{wds}$ planes and 257 therefore preserve layers of close-packed oxygen anions while exploiting the pseudo-symmetry 258 of the oxygen substructure (Supplemental Fig. S2). As a result, layers of close-packed oxygen 259 anions may be continuous across the composition plane while the ordering of cations into 260 tetrahedral and octahedral voids on opposing sites of the interface is related by the twin operation. 261 The twin law that we derived for the here-reported twins in wadsleyite is therefore consistent 262 with growth twinning. 263

The twin relation can then be described by the following twin equation (e.g., Hirsch, 1977): $\{021\} + \{101\} = \{122\}$, in which one of the $\{122\}_{wds}$ planes should be one of the composition planes in orthorhombic wadsleyite (**Fig. 3**). The $\{122\}_{wds}$ planes in wadsleyite correspond to the $\{101\}_{rwd}$ planes in ringwoodite. As the mirror symmetry parallel to $\{101\}_{rwd}$ in ringwoodite is reproduced along the twin interface, the atomic structure in the vicinity of the interface between twin domains may locally resemble the spinel structure of ringwoodite.

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271 Relation to twins in spinel and to spinelloid phases

272	The twin relation in wadsleyite presented here may appear similar to that of the (111)
273	twin in MgAl ₂ O ₄ spinel (e.g., Fig. 5 in Daneu et al., 2007), i.e., the spinel twin law. According to
274	Daneu et al. (2007), (111) twins in spinel can be described by a 180° rotation of the oxygen
275	substructure about one of the $<111>$ axes which are perpendicular to the $\{111\}$ composition
276	planes. Despite these similarities, the twin axis and twin plane for twins in wadsleyite correspond
277	to the cubic $<101>_{rwd}$ directions and the $\{101\}_{rwd}$ planes, respectively, i.e., not the $<111>$
278	directions and {111} planes as in spinel. Instead, there might be some similarity with the iron
279	cross law in pyrite, FeS ₂ (Rečnik et al., 2016).

280 The twinning in wadsleyite could potentially be related to other spinelloid structures 281 (Horiuchi et al., 1982). However, we could not detect any other phases at the composition plane 282 nor a dense magnesium-iron silicate polymorph with a structure intermediate between olivine, 283 ringwoodite, and wadsleyite, as for instance poirierite (Tomioka et al., 2021), wadsleyite II (Smyth and Kawamoto, 1997), or other spinelloid phases (Horiuchi et al. 1982; Davies and 284 285 Akaogi, 1983). The crystal structures of two adjacent twin domains continue perfectly across the 286 interface, i.e., the composition plane, without any significant structural mismatches or chemical changes (Fig. 5). 287

288

289 Implications

The formation of twins might have consequences for the physical properties and deformation behavior of wadsleyite in the mantle transition zone. Shear deformation experiments on wadsleyite aggregates and the analysis of the resulting microstructures indicated a dislocation creep mechanism with slip mainly on the lattice planes (010)_{wds} and {101}_{wds} (Farla et al., 2015;

Kawazoe et al., 2013; Ohuchi et al., 2014; Miyajima and Kawazoe, 2015). None of these slip planes coincides with the composition planes $(122)_{wds}$ of the here-detected twins nor do the orientations of slip planes in adjacent domains coincide at the composition plane. The resulting discontinuity of slip planes at the interface between twin domains might hinder dislocations to move across the interface. In this way, the twins in wadsleyite might impede dislocation movement and reduce the strength of deformation textures related to the affected slip systems.

Twin walls in wadsleyite may only reduce deformation rates dominated by dislocation glide. More comprehensive models for the viscoelastic behavior of wadsleyite aggregates, however, suggest that dislocation glide might be an ineffective deformation mechanism at pressures and temperatures of the mantle transition zone (Castelnau et al., 2020; Ritterbex et al., 2016) with pure climb creep being more efficient than dislocation glide except for the coldest regions of the mantle transition zone (Ritterbex et al., 2020). For climb creep, twin interfaces might act as sources of cation disorder and vacancies and promote deformation.

Depending on the frequency of twinning, the modified mobility of dislocations might 307 have implications for the deformation of wadsleyite-bearing rocks in the mantle transition zone 308 309 and the associated seismic anisotropy (Kawazoe et al., 2013; Tommasi, 2004). In addition to 310 impairing the alignment of wadsleyite grains by deformation through dislocation glide, twins in 311 wadsleyite will reduce the elastic anisotropy as directions with P and S wave velocity maxima in 312 one twin domain will be mapped onto directions with intermediate wave velocities in the other twin domain (Buchen, 2018, Buchen et al., 2018). Although further research is needed to assess 313 314 the potential impact of twinning on the properties of wadsleyite-bearing rocks, we anticipate a 315 potential effect on the deformation behavior under the assumption that twinning is common in 316 natural wadsleyite.

In addition to affecting physical properties, twinning in wadslevite might have an impact 317 on the kinetics of the phase transition to ringwoodite. Our TEM observations indicate that the 318 atomic structure in the vicinity of twin boundaries resembles the structure of ringwoodite while 319 conserving the oxygen substructure across the composition plane. Because the c axes and $[120]^*$ 320 directions (corresponding to the *a* axis in ringwoodite) of adjacent domains meet at the 321 composition plane, the composition plane itself has the same structural unit as ringwoodite (e.g., 322 Horiuchi et al., 1982). We present direct evidence for local structural order along the 323 composition plane in the HRTEM image of Fig. 5a. Boundaries between twinned domains in 324 wadsleyite might therefore act as nucleation sites for lamellae of ringwoodite. Lamellar 325 intergrowth of spinelloid phases has been observed previously in the system Ni₂SiO₄-NiAl₂O₄ 326 327 (Davies and Akaogi, 1983). When wadslevite-bearing rocks are brought into the stability field of ringwoodite, for example, by downward flow in the mantle transition zone, ringwoodite lamellae 328 might preferentially nucleate along twin composition planes without need to overcome the 329 energy barrier associated with creating viable nuclei. Crystal growth of ringwoodite may then 330 proceed from the interface outwards by rearranging cations but retaining the oxygen substructure. 331 This replacement mechanism might be faster than other nucleation and growth processes. 332 However, more research is needed to understand the relation between wadsleyite twins and 333 ringwoodite nucleation and the potential impact on the kinetics of the wadsleyite-ringwoodite 334 335 phase transition.

336

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457 Figure captions

458	Fig. 1. Polarized-light microscopic images of twins in wadsleyite between crossed polarizers (a)
459	and in plane-polarized light (b). The thin section shown in (b) has been polished parallel to the
460	(120) plane of domain A. As a result, the darker domain B on the right is oriented parallel to
461	(001). Optical vibration directions (arrows) are also indicated for each domain.
462	Fig. 2. (a) Secondary electron image during the micro-sampling of a TEM lamella across a twin
463	interface detected with polarized light microscopy (Fig. 1b). (b, c) Optical microscopy images of
464	the TEM lamella.
465	Fig. 3. (a) Dark-field TEM image (with $g = 244$) of a twinned wadsleyite grain. The composition
466	plane runs along the diagonal direction of the FIB-sampling TEM foil. (b) The SAED pattern of
467	the $[2\overline{1}0]$ zone axis from an area including the composition plane and (c) the corresponding
468	PsSAED pattern. The precession angle is 3°. Note how the 244 diffraction spots (and higher
469	orders) are split perpendicular to the composition plane. The indexing of host domain (A) and twin
470	domain (B) in the diffraction pattern (b).
471	Fig. 4. (a) Bright-field TEM image and (b) LACBED pattern across a composition plane viewed
472	along the $[2\overline{1}0]$ zone axis of wadsleyite. The convergent angle is more than 100 mrad (~ 5°). (c)
473	The composition plane was moved off-center on the LACBED pattern for better visualization.
474	Fig. 5. (a) Filtered high-resolution TEM, (b) high-resolution bright-field STEM, and (c) bright-
475	field scanning moiré fringe (SMF) imaging of a composition plane viewed along the $[2\overline{1}0]$ zone
476	axis of wadsleyite. The directions of moiré fringes indicate a clear discontinuity across the
477	interface between two domains. Note that in (a) the foil thickness decreases towards the right-

hand side. The SMF image in (c) was taken under 66 times lower electron dose than that of thehigh-resolution bright-field STEM image in (b).

480 Fig. 6. SAED patterns of the [001] zone axis of wadsleyite. In (a) there is no detectable intensity

481 for *hk*0 reflections with h, k = 2n + 1, e.g. at the expected position of the $110_{\rm m}$ reflection (red

482 circle). In (b) the $110_{\rm m}$ reflection becomes visible in a systematic tilt condition along the [110]*

483 (and [120]*) direction.

Fig. 7. (a) A HOLZ pattern of the [001] zone axis of wadsleyite using a camera length of 340

485 mm in a Philips CM20FEG TEM. The observed intensity distributions along the a^* axis (red

ellipses) are incompatible with a mirror plane (*m*) perpendicular to the *a* axis. (b) Simulated

487 HOLZ pattern for wadsleyite with orthorhombic symmetry, using the EDA software (courtesy of

488 Prof. Kogure). (c) For comparison, a HOLZ pattern of the $[\overline{1}10]$ zone axis of olivine, space

489 group *Pbnm*, i.e., $P2_1/b 2_1/n 2_1/m$, recorded with the same TEM instrument. The mirror

490 symmetry of olivine perpendicular to the c axis is clearly reflected in the HOLZ pattern.

Fig. 8. Depiction of the crystal structures of wadsleyite (a,c) and ringwoodite (b,d). Red spheres 491 492 show oxygen anions; blue tetrahedra host silicon cations; green octahedra host magnesium and 493 iron cations. The axis systems and crystal structures are drawn to preserve the orientational 494 relationship between both crystal structures. In (a) and (b), orange planes indicate the orientation of close-packed layers of oxygen anions. In (c) and (d), the crystal structures of wadsleyite and 495 ringwoodite are projected along the $(\overline{1}22)_{wds}$ and $(011)_{rwd}$ planes, respectively, to highlight 496 structural similarities. The orientation of twin axes and twin planes with respect to the axis 497 systems of both crystal structures is indicated in red. Graphics were created with VESTA 498 499 (Momma and Izumi, 2011).

Space-	Imma	I2/m
group type		
a (Å)	5.6936(4)	5.6939(2)
<i>b</i> (Å)	11.5067(3)	11.5067(2)
<i>c</i> (Å)	8.2676(3)	8.2678(2)
$V(\text{\AA}^3)$	541.65(4)	541.68(3)
β (°)	90	90.030(3)
b/a	2.0210(2)	2.0209(1)
N^{a}	33	33

Table 1.Unit cell parameters of a
twinned wadsleyite crystal

Note: Standard deviations on the last digit are given in parentheses.

^a Number of reflections centered on a four-circle diffractometer.

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504 Fig. 1. Polarized-light microscopic images of twins in wadsleyite between crossed polarizers (a)

and in plane-polarized light (b). The thin section shown in (b) has been polished parallel to the

- 506 (120) plane of domain A. As a result, the darker domain B on the right is oriented parallel to
- 507 (001). Optical vibration directions (arrows) are also indicated for each domain.

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Fig. 2. (a) Secondary electron image during the micro-sampling of a TEM lamella across a twin
interface detected with polarized light microscopy (Fig. 1b). (b, c) Optical microscopy images of
the TEM lamella.











Fig. 3. (a) Dark-field TEM image (with g = 244) of a twinned wadsleyite grain. The composition 520 521 plane runs along the diagonal direction of the FIB-sampling TEM foil. (b) The SAED pattern of the $[2\overline{10}]$ zone axis from an area including the composition plane and (c) the corresponding 522 PsSAED pattern. The precession angle is 3°. Note how the 244 diffraction spots (and higher 523 524 orders) are split perpendicular to the composition plane. The indexing of host domain (A) and twin 525 domain (B) in the diffraction pattern (b).





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Fig. 4. (a) Bright-field TEM image and (b) LACBED pattern across a composition plane viewed
along the [210] zone axis of wadsleyite. The convergent angle is more than 100 mrad (~ 5°). (c)
The composition plane was moved off-center on the LACBED pattern for better visualization.









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Fig. 5. (a) Filtered high-resolution TEM, (b) high-resolution bright-field STEM, and (c) brightfield scanning moiré fringe (SMF) imaging of a composition plane viewed along the [210] zone
axis of wadsleyite. The directions of moiré fringes indicate a clear discontinuity across the
interface between two domains. Note that in (a) the foil thickness decreases towards the right-

- hand side. The SMF image in (c) was taken under 66 times lower electron dose than that of the
- 546 high-resolution bright-field STEM image in (b).



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Fig. 6. SAED patterns of the [001] zone axis of wadsleyite. In (a) there is no detectable intensity for *hk*0 reflections with h,k = 2n + 1, e.g. at the expected position of the $110_{\rm m}$ reflection (red circle). In (b) the $110_{\rm m}$ reflection becomes visible in a systematic tilt condition along the [110]* (and [120]*) direction.



Fig. 7a

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Fig. 7. (a) A HOLZ pattern of the [001] zone axis of wadsleyite using a camera length of 340 mm in a Philips CM20FEG TEM. The observed intensity distributions along the a^* axis (red ellipses) are incompatible with a mirror plane (*m*) perpendicular to the *a* axis. (b) Simulated HOLZ pattern for wadsleyite with orthorhombic symmetry, using the EDA software (courtesy of Prof. Kogure). (c) For comparison, a HOLZ pattern of the [110] zone axis of olivine, space group *Pbnm*, i.e., *P*2₁/*b* 2₁/*n* 2₁/*m*, recorded with the same TEM instrument. The mirror symmetry of olivine perpendicular to the **c** axis is clearly reflected in the HOLZ pattern.





Fig. 8. Depiction of the crystal structures of wadsleyite (a,c) and ringwoodite (b,d). Red spheres 571 show oxygen anions; blue tetrahedra host silicon cations; green octahedra host magnesium and 572 iron cations. The axis systems and crystal structures are drawn to preserve the orientational 573 relationship between both crystal structures. In (a) and (b), orange planes indicate the orientation 574 of close-packed layers of oxygen anions. In (c) and (d), the crystal structures of wadsleyite and 575

576 ringwoodite are projected along the $(\bar{1}22)_{wds}$ and $(011)_{rwd}$ planes, respectively, to highlight 577 structural similarities. The orientation of twin axes and twin planes with respect to the axis 578 systems of both crystal structures is indicated in red. Graphics were created with *VESTA* 579 (Momma and Izumi, 2011).

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