

1 REVISION 1

2 **Mechanism and kinetics of the pseudomorphic replacement of anhydrite by calcium**
3 **phosphate phases at hydrothermal conditions**

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11

12 **Abstract**

13 Mineral replacement reactions mediated by fluids are common in sedimentary basins, where
14 they influence geochemical cycles. Phosphorous (P) pollution of soils, sediments and water
15 bodies is currently a widespread problem. Some apatite accumulations in sediments may have
16 formed through the interaction of P-bearing aqueous solutions with mineral surfaces that result
17 in mineral replacement reactions. Here, we investigate the pseudomorphic replacement of
18 anhydrite single crystals by aggregates of β -tricalcium phosphate and hydroxyapatite upon
19 interaction with a P-bearing solution at temperatures between 120°C and 200°C. SEM
20 imaging is used to study the texture of the aggregates. Rietveld refinement of the X-ray
21 diffraction patterns and Raman spectra analysis of the reacted samples provide information on
22 the kinetics of the replacement. At all temperatures β -tricalcium phosphate forms alongside
23 hydroxyapatite at early stages of the replacement reaction. At $T \geq 180^\circ\text{C}$, hydroxyapatite/ β -

24 tricalcium phosphate ratio rapidly increases, and hydroxyapatite is the only phase in fully
25 replaced samples. At $T < 180^{\circ}\text{C}$ hydroxyapatite/ β -tricalcium phosphate ratio increases slowly
26 and fully replaced samples still contain significant amounts of β -tricalcium phosphate. The
27 progress of the replacement is facilitated by the formation of porosity. The evolution of the
28 hydroxyapatite/ β -tricalcium phosphate ratio and the crystal habit of both phases strongly
29 influence the arrangement of this porosity. The empirical activation energy E_a (kJ/mol) of the
30 replacement reaction is determined by the Avrami and the iso-conversion methods. Both
31 approaches yield an E_a of ~ 40 kJ/mol. Anhydrite dissolution appears as the rate-limiting
32 process and the overall kinetics of the replacement reaction is controlled by the rate diffusion
33 of dissolved species through the porosity network. The ripening of the metastable β -tricalcium
34 phosphate into hydroxyapatite affects the characteristics of the porosity network and further
35 modulates the kinetics of the replacement. These results may improve the understanding of
36 the mechanisms of P-sequestration by mineral surfaces through coupled dissolution–
37 precipitation reactions and shed light on the origin of apatite accumulations associated to
38 evaporitic sedimentary rocks.

39 **Keywords:** anhydrite, hydroxyapatite, β -tricalcium phosphate, mineral replacement,
40 pseudomorphism, kinetics, textures, coupled dissolution–precipitation.

41

42 **Introduction**

43 Mineral replacements are common processes in sedimentary, diagenetic, metamorphic, and
44 metasomatic environments, where they take place mediated by a fluid that facilitates the
45 dissolution of the primary mineral and the concomitant precipitation of the secondary one(s).
46 The coupling through the interface of the dissolution and precipitation rates allows for the
47 preservation of original shape and volume of the primary mineral (Putnis 2002, 2009; Putnis

48 and Putnis 2007; Pollok et al. 2011; Ruiz-Agudo et al. 2014). Moreover, when the dissolution
49 of the latter is the rate-limiting step of the interface coupled dissolution-precipitation reaction
50 (ICDP), original microscopic, and even nanoscopic, features also are accurately preserved
51 during the replacement (Xia et al. 2009; Ruiz-Agudo et al. 2014; Altree-Williams et al. 2015).
52 The progress of ICDP reactions requires a continuous communication between the interface,
53 where the reaction takes place, and the bulk fluid. This communication is ensured by the
54 formation of a network of porosity and/or fractures during the mineral replacement. This
55 network facilitates mass transport to and from the interfacial fluid (Zhao et al. 2009; Jonas et
56 al. 2013; Ruiz-Agudo et al. 2014; Putnis 2015).

57 In this work we describe the replacement of anhydrite by mixtures of two calcium phosphate
58 phases, β -tricalcium phosphate (β -Ca₃(PO₄)₂) (β -TCP) and hydroxyapatite (Ca₅(PO₄)₃(OH))
59 (Hap) through a ICDP mediated by an aqueous fluid containing phosphorous (P). Anhydrite
60 is a rock-forming mineral common in sedimentary basins, where it appears as a major
61 component of evaporites as well as forming nodules and cements in sandstones, limestones
62 and dolostones (Murray 1964; Rahimpour-Bonab et al. 2010; Olivarius et al. 2015).
63 Anhydrite also forms massive accumulations in modern submarine hydrothermal vents (Kuhn
64 et al. 2003). Apatite (Ca₅(PO₄)₃(OH,Cl,F)) is the main source of phosphorous, an element
65 that is scarce in the Earth's crust but has great economic interest in the fertilizer industry
66 (Filippelli 2002). Phosphorous is an essential macronutrient for biota as well as a pollutant
67 which in excess causes eutrophication of water bodies, contributes to the growth of toxigenic
68 algae, and boosts the development of potentially pathogenic microbes in the water column
69 and in soils (Mallin and Cahoon 2020). The structural characteristics of apatite make it a
70 relevant phase for immobilizing and storing of radioactive and other pollutant elements
71 (Rakovan and Reeder 1996; Ewing and Wang 2002; Mavropoulos et al. 2002).

72 There is abundant evidence that mineral replacement reactions play an important role in
73 controlling the fate of a variety of pollutants in soils and sedimentary basins (Pinto et al. 2009,
74 2012; Wang et al. 2012, 2015; Hövelmann and Putnis 2016; Callagon et al. 2017; Di Lorenzo
75 et al. 2019; Roza Llera et al. 2021). Anhydrite is highly reactive to dissolved ions. Thus, cycles
76 of anhydrite dissolution/precipitation influence the fate of trace elements (Sr, Y, REE) in
77 hydrothermal systems. Moreover, anhydrite surfaces effectively remove metal pollutants from
78 aqueous solutions through coprecipitation phenomena (Morales et al. 2014; Forjanés et al.
79 2020b, 2020a). Pollution of running waters and aquifers by phosphorous currently is
80 widespread due to the extensive use of ammonium phosphate as a fertilizer in intensive
81 agriculture (Smith 2003; Wei and Bailey 2021). Understanding the processes that result from
82 the interaction of phosphorous containing waters with common sedimentary rock-forming
83 minerals can help to mitigate the effects of phosphorous pollution in sedimentary basins.

84 Moreover, apatite is a main constituent of sedimentary phosphorites (Bentor 1980; Sheldon
85 1981; Hughes et al. 1989; Knudsen and Gunter 2002). Baturin (1989) reported on the
86 diagenetic origin of some apatite accumulations and connected their formation to changes in
87 the chemistry of pore waters due to organic matter decay and sulfate reduction. The
88 concentration of phosphorous in diagenetic fluids in pores can reach values up to 8-9 mg/L,
89 according to Sheldon (1981). Filippelli (2002) and Dzombak and Sheldon (2020) have
90 reported that P concentrations in young soils can reach values as high as 5340 mg/L,
91 highlighting the potential importance of P fluxes from soils to rivers in the current scenario of
92 climate change. In contact with the surface of sediment particles, these high concentrations
93 are sufficient to guarantee the precipitation of apatite, which opens the question of the role
94 that interface coupled dissolution-precipitation reactions (ICDP) may play in the formation of
95 diagenetic apatite. It has been experimentally demonstrated that the interaction of phosphate-
96 bearing fluids with other sedimentary rock-forming minerals like aragonite and calcite results

97 in their replacement by aggregates of apatite crystals (Kasioptas et al. 2010, 2011; Jonas et al.
98 2013; Reinares-Fisac et al. 2017).

99 The purpose of the present work is to study the interaction of a P-bearing aqueous solution
100 with anhydrite single crystals. The main goal is to determine the kinetics and mechanisms of
101 anhydrite replacement by calcium phosphate at temperatures between 120°C and 200°C. Two
102 calcium phosphate phases, β -TCP (β -Ca₃(PO₄)₂) and Hap (Ca₅(PO₄)₃(OH)), formed during
103 this replacement, which were identified by Raman spectroscopy and X-ray powder diffraction
104 (XRD). The amount of each phase as a function of temperature and reaction time was
105 calculated by conducting the Rietveld analysis of the X-ray diffraction (XRD) patterns. This
106 information was applied to determine the kinetics of the process. The textural relationships
107 and the crystal habit of the product phases were studied by scanning electron microscopy
108 (SEM). The influence of the transformation of β -TCP into Hap, which occurs concomitantly
109 to the progress of the replacement, as well as the specific morphological and textural features
110 of both calcium phosphates in the overall kinetics of the replacement reaction, was considered.

111

112 **Materials and methods**

113 **Hydrothermal experiments**

114 Hydrothermal interaction experiments were performed by reacting anhydrite single crystals
115 with a 2M (NH₄)₂HPO₄ aqueous solution at different reaction times (1 h to 5 days) and
116 temperatures (120, 150, 180 and 200°C) under autogenous pressure. The P concentration in
117 the aqueous solution is around 10 times higher than that reported in young soils (Filippelli
118 2002; Dzombak and Sheldon 2020). Anhydrite single crystals from Naica mine (Chihuahua,
119 México) were used in all experiments. X-ray fluorescence spectroscopic analysis confirmed
120 them as highly, with less than 0.4 wt% foreign elements and Sr as the major impurity. The

121 anhydrite crystals were cleaved along {100}, {010} and {001} using a sharp stainless-steel
122 knife edge to obtain similarly sized (about 3 x 3 x 3 mm) anhydrite subsamples. Prior to their
123 use in the experiments, these subsamples were cleaned in an ethanol bath to remove surface
124 impurities and then left to dry overnight in a desiccator at room temperature. Four anhydrite
125 subsamples (average weight $\sim 250 \pm 1$ mg) were used in each experiment, and the volume of
126 the solution (~ 4 mL) was adjusted to guarantee a solid-to-liquid ratio of 0.06 g/mL. Anhydrite
127 subsamples were placed together with the P-bearing aqueous solution in pre-heated individual
128 stainless-steel Teflon-lined autoclaves (2.5 \varnothing x 10 cm). Within the autoclave, all four
129 anhydrite subsamples laid with one of their surfaces in contact with the bottom of the reactor,
130 while their other five surfaces were directly exposed to the interaction with the P-bearing
131 solution. The crystallographic orientation of the laying surface was not taken into
132 consideration when setting the different experimental runs. The P-bearing aqueous solution
133 was prepared by dissolving reagent-grade $(\text{NH}_4)_2\text{HPO}_4$ (Acros Organics) in high purity
134 deionized water (MilliQ) (18 M Ω cm). The pH of the fluid was measured prior to and after
135 reaction by using a pH Meter basic 20-CRISON. Experiments were run triplicated to confirm
136 reproducibility. The aqueous solutions were modeled using the geochemical code PHREEQC
137 (Parkhurst and Appelo 1999) and llnl.dat database to calculate saturation indexes (SI) with
138 respect to relevant solid phases at the beginning of the experiments.

139 After reaction, the autoclaves were removed from the oven. To accelerate their cooling to
140 room temperature (25°C) a flow of compressed air was applied during 20 minutes. The
141 reaction time does not comprise this cooling period. Afterwards, the product samples were
142 removed from the solution, washed with distilled water, and dried overnight at 30°C in a
143 thermostatic chamber. Two of the four product samples from each experiment were powdered
144 in an agate mortar and used for powder X-ray diffraction (XRD) analyses. The other two

145 product samples were crosscut using a steel cutter and used for Raman spectroscopy and
146 scanning electron microscopy (SEM) imaging analyses.

147 **Analytical methods**

148 Diffraction patterns of powdered product samples were measured at room temperature with a
149 PANalytical Xpert Pro diffractometer, using a $\text{CuK}\alpha 1$ radiation. XRD data were collected in
150 the range $7\text{--}90^\circ 2\theta$ with a step size of 0.001° and a dwell time of 1 s per step. Samples were
151 hold in thin-walled glass capillaries to minimize preferential orientation effects. General peak
152 matching runs were conducted for the diffraction patterns of product samples. These runs
153 confirmed that in all diffraction patterns all peaks could be assigned to anhydrite (PDF 98-
154 900-1234) and/or Hap (PDF 98-006-8592) and β -TCP (PDF 00-009-0169). Furthermore,
155 anhydrite, Hap and β -TCP fractions in the reacted samples were determined by performing
156 Rietveld refinement analyses using the X'Pert HighScore Plus (Version 3.0) software package
157 from PANalytical. The aforementioned structural models were used as references.

158 Crosscut reacted samples were gold-coated and imaged using a JEOL-6610LV microscope
159 equipped with Energy-dispersive X-ray spectroscopy (EDX, INCA Energy 350). The contrast
160 resulting from the different composition of reacted and unreacted areas provided information
161 on the progresses of the anhydrite-(P)-bearing fluid interaction. Secondary electron images of
162 the external surfaces of the reacted solid samples were also obtained.

163 Raman spectra were collected from freshly crosscut sections of the reacted subsamples.
164 Raman spectroscopy was performed by means of a labRam HRU using JYVJobin Yvon
165 equipment with an excitation source laser CDPS532-DPSS at 24.3 mW and excitation times
166 of 10 s. This equipment was connected to a confocal microscope Olympus BXFM-ILHS with
167 a 100x objective.

168 **Kinetic analysis**

169 We assume that the isothermal kinetics of the anhydrite phosphation reaction can be described
170 by the Avrami equation (Eq.1):

$$171 \quad y_{CaP} = 1 - e^{-(kt)^n} \quad (1)$$

172 where y_{CaP} corresponds to the fraction of calcium phosphate formed ($y_{Anh} = 1 - y_{CaP}$), this is the
173 extent of the phosphation reaction, t (hours) represents the reaction time, k is the Avrami rate
174 constant (hour^{-1}) and n is the Avrami exponent. By linearizing the Avrami equation taking
175 twice natural logarithms the following expression is obtained:

$$176 \quad \ln(-\ln(1 - y_{CaP})) = n \ln k + n \ln t \quad (2)$$

177 The Avrami equation has successfully been used to describe the kinetics of a variety of
178 dissolution, crystallization, transformation, and mineral replacement processes (Lasaga 1998;
179 Xia et al. 2009; Kasiopas et al. 2010; Altree-Williams et al. 2017; Pedrosa et al. 2017). If a
180 reaction follows the Avrami equation, different isokinetic curves will have the same n value.
181 Moreover, their plot as $\ln(-\ln(1 - y_{CaP}))$ against $\ln t$ (Hancock-Sharp plot) will show
182 approximately equal gradients (Hancock and Sharp 1972). Changes in the value of n are
183 indicative of changes in the mechanisms of the reaction. If the activation energy of the reaction
184 is constant, the Hancock-Sharp plot will yield a straight line, whose slope corresponds to the
185 value of n and from whose intercept with the y -axis the rate constant k can be estimated.
186 Deviations from linearity of the Hancock-Sharp plot indicate that the reaction equation differs
187 from the Avrami one.

188 The rate constant k of most reactions shows a dependence of temperature that follows an
189 Arrhenius type equation:

$$190 \quad k = A \cdot e^{\frac{-Ea}{R \cdot T}} \quad (3)$$

191 where A is a frequency factor, R is the gas constant, T is the temperature of the experiment
192 (Kelvin) and E_a is the empirical activation energy. Taking logarithm, the equation 4 is
193 obtained:

$$194 \quad \ln k = \ln A - E_a/RT \quad (4)$$

195 If A and E_a are constant, the plot of $\ln k$ versus $1/T$ is a straight line whose slope is E_a/R and
196 its intercept with the y-axis is $\ln A$. Thus, the value of E_a can be estimated using the rate
197 constants k derived from the Hancock-Sharp plots of a set of isothermal runs.

198 Commonly the empirical activation energy sums up contributions of different events involved
199 in the reaction (dissolution of primary phases, mass transfer through interfaces, ion
200 dehydration, nucleation and growth of secondary phases, etc). Moreover, the method
201 described above fails to reveal changes in the activation energy that occur as the reaction
202 progresses. It is possible to derive information on the evolution of the empirical activation
203 energy E_a along the reaction by rewriting equation 1 and calculating the time required to
204 achieve the transformation of a given fraction of the primary phase, in this case, anhydrite,
205 into the secondary phase(s), in this case, Hap and β -TCP or a mixture of both phases,
206 according to equation 5:

$$207 \quad t_Y \propto k^{-1} \propto A^{-1} e^{\frac{E_a}{RT}} \quad (5)$$

208 where t_Y is the time for a given fraction to transform. This approach is described as the
209 isoconversional method or the time to a given fraction method (Putnis 1992).

210 Taking logarithms, equation 5 can be rearranged as:

$$211 \quad \ln t_Y = \text{const} - \ln A + E_a/RT \quad (6)$$

212 Empirical activation energies can be derived by plotting $\ln t_Y$ versus $1/T$. If there is no change
213 in the reaction mechanism as the reaction progresses, all plots will have similar gradients and,

214 consequently, E_a will have a constant value, regardless of the given fraction transformed. On
215 the contrary, changes in plot slopes are indicative of different mechanisms operating at
216 different stages of the reaction, each one characterised by a different E_a .

217

218 **Results**

219 **Reaction Pathway and Textural Features**

220 The hydrothermal interaction of anhydrite single crystals with a phosphate-bearing aqueous
221 solution (2M $(\text{NH}_4)_2\text{HPO}_4$) results in their partial to total transformation into Hap
222 ($\text{Ca}_5(\text{PO}_4)_3(\text{OH})$). Hydroxyapatite can be accompanied by varying amounts of β -TCP (β -
223 $\text{Ca}_3(\text{PO}_4)_2$) depending on the temperature and duration of the interaction. During the
224 transformation, the pH of the solutions progressively decreases from an initial value of 8.1 (1)
225 to 6.8 (1) at experiment termination time. Figure 1a depicts the XRD patterns from subsamples
226 reacted during 6 to 72 h at 120°C. All the diffraction patterns show sharp peaks that can be
227 assigned to anhydrite and/or the calcium phosphate phases Hap and β -TCP. Although the
228 XRD patterns of other calcium phosphates like whitlockite ($\text{Ca}_{18}(\text{Mg,Fe})_2\text{H}_2\text{PO}_4$)₁₄) show
229 similar features as the pattern of β -TCP (Gopal and Calvo 1972; Jonas et al. 2014), the
230 formation of any amount of the former phase is disregarded, since the pristine anhydrite is
231 highly pure and the aqueous phase contains no dissolved Fe and Mg. The intensity of anhydrite
232 peaks progressively decreases with the reaction time. Anhydrite peaks are absent in the pattern
233 corresponding to a reaction time of 72 h. Conversely, the intensity of peaks assigned to Hap
234 progressively increases with time. Peaks assigned to β -TCP show a more complex evolution:
235 their intensity initially grows to latter decrease. In all the experiments, regardless of the
236 reaction temperature, Hap rapidly becomes the main constituent of the transformed fraction.
237 For example, in experiments conducted at 120°C, the transformed fraction contains ~ 20 wt%

238 Hap after 12 h, 37 wt% Hap after 24 h and 73 wt% Hap after 48 h (Fig. 1b). In experiments
239 conducted at temperatures $\geq 200^{\circ}\text{C}$, Hap becomes the only solid phase present in the system
240 after 10 h since the beginning of the experiment. In addition, longer experiments were carried
241 out at temperatures below 180°C , once the anhydrite had been completely replaced, to observe
242 the evolution of the relative content of β -TCP and Hap in the samples. The results of these
243 experiments indicated that the amount of β -TCP continued to decrease with time. After 5 days
244 interaction at 120°C , reacted samples consist of 7.2% β -TCP and 92.8 anhydrite. The Raman
245 spectra collected on crosscut sections of reacted anhydrite subsamples indicate that all
246 transformed rims exclusively consist of calcium phosphate phases in good agreement with the
247 conclusions of XRD analyses. Figure 2 shows SEM images and Raman spectra taken on
248 crosscut sections of anhydrite reacted with the P-bearing fluid at 150°C during 12 h (Figs. 2a-
249 2d) and 24 h (Figs. 2e-2h). The spectrum in Figure 2a shows the main vibration bands of SO_4^{2-}
250 in anhydrite: a symmetric stretching (ν_1) band at 1019 cm^{-1} , three asymmetric stretching (ν_3)
251 bands at 1113 cm^{-1} , 1131 cm^{-1} and 1160 cm^{-1} , two symmetric bending (ν_2) bands at 416 cm^{-1}
252 and 502 cm^{-1} , and three asymmetric bending (ν_4) bands at 610 cm^{-1} , 627 cm^{-1} and 678 cm^{-1} .
253 Raman spectra collected from transformed regions of samples reacted during 12 h (Fig. 2c)
254 and 24 h (Figs. 2e and 2g) show bands that can be assigned to Hap: the symmetric stretching
255 (ν_1), the symmetric bending (ν_2) and the asymmetric bending (ν_4) vibration modes of PO_4^{3-}
256 are located at 961 cm^{-1} , 432 cm^{-1} and 588 cm^{-1} , respectively. The presence of a minor broad
257 band around 880 cm^{-1} (Figs. 2c and 2e) that can be attributed to the P-OH stretching is also
258 consistent with Hap (Penel et al. 1998). Moreover, a triplet at 1007 , 1045 and 1075 cm^{-1} ,
259 corresponds to the PO_4^{3-} asymmetric stretching (ν_3) as well as a band at 3571 cm^{-1} can be
260 assigned to the OH-stretching (Fig. 2). In addition, a broad band near $946\text{--}949\text{ cm}^{-1}$ assigned
261 to the symmetric stretching (ν_1) (Jonas et al. 2014; Pedrosa et al. 2016) indicates the presence
262 of β -TCP in the core of the sample reacted during for 24 h (Fig. 2e). The absence of bands

263 around 925 cm^{-1} (ν_1 mode) in all Raman spectra, confirms that whitlockite is not present in
264 the system at any stage, in good agreement with the results of XRD analyses that only identify
265 two calcium phosphate phases (β -TCP and Hap).

266 The main PO_4^{3-} vibrational mode (ν_1) in β -TCP, which is located at 971 cm^{-1} , is not apparent
267 in any of the Raman spectra, most likely due to overlapping with the main band in the spectra
268 of Hap (PO_4^{3-} ν_1 at 961 cm^{-1}). No bands that can be attributed to β -TCP are present in areas
269 close to the surface of the anhydrite sample reacted during 24 h (Fig. 2g). In summary, XRD
270 and Raman results corroborate that the transformed samples consist of mixtures of Hap and
271 β -TCP at early stages of the replacement reaction. At $T \geq 180^\circ\text{C}$, hydroxyapatite/ β -tricalcium
272 phosphate ratio rapidly increases, and hydroxyapatite is the only phase in fully replaced
273 samples. At $T < 180^\circ\text{C}$ hydroxyapatite/ β -tricalcium phosphate ratio increases slowly.
274 Furthermore, reacted samples still contain significant amounts of β -tricalcium phosphate even
275 after full replacement of anhydrite is reached.

276 The anhydrite by calcium phosphate replacement initiates at the surface of the anhydrite
277 subsamples and advances inwards, defining a sharp reaction front (Figs. 2b and 2d). This front
278 defines the interface between a shell-like transformed rim and the unreacted anhydrite core
279 (Fig. 2d). Figure 3a shows the contact between the transformed rim and the unreacted
280 anhydrite core. As can be seen, the rim mainly consists of needle-like crystals arranged in
281 fan-like bunches that appear oriented roughly perpendicular to the surface of the anhydrite
282 core, and a smaller amount of euhedral rhombohedron-shaped crystals (Fig. 3a). Based on
283 Raman spectroscopy analysis and the known crystal morphology of calcium phosphates, we
284 interpret the needle-like crystals as Hap (Zhu et al. 2008; Kasiopas et al. 2010; Yang et al.
285 2014; Li et al. 2016) and the euhedral rhombohedron-shaped ones as β -TCP (Roy and
286 Linnehan 1974). Closer inspection of the latter evidences the presence on their surfaces of

287 dissolution pits and nanometric Hap crystals that appear to grow in close spatial relation with
288 the dissolution pits (Fig. 3b).

289 Regardless the reaction temperature, the replacement reaction takes place with preservation
290 of both the volume and external shape of the anhydrite crystals and results in the formation of
291 calcium phosphate pseudomorphs. The pseudomorphs obtained in series of isothermal
292 experiments were crosscut and SEM imaged. Figure 4 depicts SEM microphotographs of
293 samples reacted during 12 hours at 120°C (a), 150°C (b), 180°C (c) and 200°C (d). The
294 transformed rim undergoes progressive thickening with increasing reaction time. The
295 anhydrite core concomitantly shrinks. It is important to note that, before full anhydrite
296 replacement is reached, the thickness of the transformed rim can significantly vary within each
297 crosscut section, as is apparent in Figures 4a-4c. Differences in rim thickness, measured from
298 opposite sides of a subsample along the same crystallographic direction, reflect the reduced
299 exposure of the subsample surface that lies in contact with the reactor bottom to the interaction
300 with the fluid. Differences in rim thickness measured along different crystallographic
301 directions reflect the different reactivity of anhydrite cleavage surfaces.

302 As can be seen, the reaction front advances faster with increasing temperature. Thus, after 12
303 h of reaction, in samples reacted at 120°C and 150°C and 180°C, the average transformed rim
304 thickness was 250, 550 and 1500 μm , respectively (Figs. 4a-4c). After 12 h reaction, no
305 unreacted core is observed in the crosscut sections of samples from experiments conducted at
306 200°C reacted (Fig. 4d). This is consistent with the results of both, X-ray diffraction and
307 Raman spectroscopy analyses, which also support the complete replacement of anhydrite by
308 calcium phosphate phases in samples reacted 12 hours at 200°C.

309 **Kinetics of the replacement reaction**

310 Several sets of hydrothermal interaction experiments were conducted to derive information
311 on the kinetics of the replacement of anhydrite single crystals by calcium phosphate crystal

312 aggregates. Table 1 summarises the fraction of anhydrite, Hap and β -TCP, as determined from
313 the Rietveld refinements of X-ray powder diffraction patterns of samples interacted with the
314 phosphate-bearing aqueous solution at temperatures between 120°C and 200°C during times
315 that varied between 2 and 72 hours. In Figure 5, the fraction of sample transformed into
316 calcium phosphate phases (y_{CaP}) is plotted against the reaction time (t) for each isothermal
317 experiment. Figure 6 shows the temperature dependence of the anhydrite-by- calcium
318 phosphate reaction, evidenced in the Hancock-Sharp plots derived by fitting the linearized
319 Avrami expression (equation 2) to the experimental data. All fitted lines are approximately
320 parallels. The Avrami parameters calculated for each set of isothermal data are also shown in
321 Figure 6. The activation energy can be obtained by plotting the slopes of the linear regressions
322 in Figure 6 against reciprocal temperature in an Arrhenius plot (Fig. 7). The rate constant (k)
323 clearly increases with temperature and the linear regression yields an empirical activation
324 energy, E_a of 40.2 ± 1.9 kJ/mol.

325 Alternatively, empirical activation energies can also be determined by using the
326 isoconversional method. E_a values were calculated for three different fractions of anhydrite
327 replacement by calcium phosphate, 0.40, 0.60, and 0.80 (Fig. 8). In all cases, the data are well
328 described by linear correlations and show closely parallel trends. The empirical E_a values
329 yielded from these fittings slightly decreases, from 40.4 ± 2.3 to 39.3 ± 2.3 kJ·mol⁻¹, as the
330 replaced fraction increases from 40 to 80%.

331

332 **Discussion**

333 **Replacement reaction mechanism**

334 Upon interaction with a P-bearing aqueous solution at temperatures between 120°C and
335 200°C, anhydrite single crystals are replaced by mixtures of β -TCP and Hap. Textural features

336 of the reacted samples are consistent with the replacement taking place through an interface
337 coupled dissolution-precipitation reaction (Putnis 2002, 2009; Putnis and Putnis 2007; Ruiz-
338 Agudo et al. 2014). Firstly, the reaction starts at the surface of anhydrite and progresses from
339 rim to core through the advancement of a reaction front that appears sharp, well defined, and
340 approximately parallel to the original anhydrite surface. Secondly, the external shape as well
341 as most surface features of anhydrite crystals are preserved throughout the reaction, which has
342 a pseudomorphic character. Thirdly, replaced samples consist of aggregates of crystals of the
343 product phases and contain large amounts of porosity. All these features are characteristic of
344 mineral replacement processes that occur mediated by the presence of a fluid phase. The
345 reaction takes place in the interfacial layer of fluid. When the process starts, the fluid is
346 undersaturated with respect to the primary phase. This drives its dissolution and determines
347 that the interfacial layer of fluid soon becomes supersaturated with respect to the product(s).
348 Anhydrite dissolution releases Ca^{2+} and SO_4^{2-} ions to the interfacial layer of fluid. To evaluate
349 the physicochemical conditions in the interfacial fluid at early stages of replacement process
350 we simulate the dissolution of a small amount of anhydrite in a small volume of fluid using
351 the geochemical code PHREEQC and the *lnll.dat* database (Parkhurst and Appelo 1999). We
352 consider the successive dissolution of layers of anhydrite whose thickness corresponds to one
353 unit cell ($a_0 = 6.993 \text{ \AA}$, $b_0 = 6.995 \text{ \AA}$ and $c_0 = 6.245 \text{ \AA}$; Hawthorne and Ferguson 1975). For a
354 $3 \times 3 \times 3 \text{ mm}$ sized crystal the dissolved anhydrite volume is $3.64 \times 10^{-8} \text{ cm}^3$ of CaSO_4 , which
355 corresponds to 7.94×10^{-10} moles of anhydrite. We consider two realistic thickness for the
356 fluid boundary layer, 100 nm and 1000 nm, which corresponds to a volume of interfacial
357 solution of $5.40 \times 10^{-6} \text{ cm}^3$ and $5.4 \times 10^{-5} \text{ cm}^3$, respectively (Ruiz-Agudo et al. 2015). The
358 state of supersaturation is defined by the value of the saturation index (SI), which is expressed
359 as:

360
$$\text{SI} = \log (\text{IAP}/\text{K}_{\text{sp}}) \quad (7),$$

361 where IAP is the ion activity product and K_{sp} is the solubility product of the phase considered.
362 $SI = 0$ indicates that the system is in equilibrium with a given phase, this is saturated, while
363 $SI < 0$ and $SI > 0$ indicate that the system is undersaturated and supersaturated, respectively.
364 At 150°C, the dissolution of one anhydrite monolayer in both 100 nm and 1000 nm boundary
365 layers results in the interfacial fluid layer becoming supersaturated with respect to Hap and β
366 -TCP. Thus, the 100 nm thick boundary layer becomes supersaturated with respect to both,
367 Hap and β -TCP ($SI_{Hap} = 23.86$, $SI_{\beta-TCP} = 3.42$) after the dissolution of one anhydrite
368 monolayer, while still remaining undersaturated with respect to this latter phase ($SI_{Anh} = -$
369 7.36). Similarly, the 1000 nm thick boundary layer reaches supersaturation with respect to
370 Hap and β -TCP ($SI_{Hap} = 19.11$, $SI_{\beta-TCP} = 0.56$) after the dissolution of one anhydrite
371 monolayer. The simulation of the dissolution of one anhydrite monolayer at other
372 experimental temperatures also yields $SI > 0$ for both, Hap and β -TCP. The SI value for all
373 the phases involved increases as successive anhydrite monolayers dissolve. Once the
374 interfacial fluid reaches the supersaturation threshold for Hap and β -TCP nucleation, both
375 calcium phosphate phases will precipitate at the reaction front. As a replaced rim forms, the
376 presence of pores within it guarantees a continuous communication between the bulk solution
377 and the reaction front, facilitating mass transport to and from the interface (Putnis et al. 2005;
378 Putnis and Putnis 2007; Pollok et al. 2011; Forjanés et al. 2020b). The progress of the reaction
379 involves the definition of a dissolution-crystallization feedback loop. Thus, anhydrite
380 dissolution promotes calcium phosphate precipitation in the interfacial fluid and *vice versa*.
381 The preservation of the external shape of anhydrite crystals requires that the rates of anhydrite
382 dissolution and calcium phosphate precipitation are coupled (Putnis 2002, 2009; Putnis and
383 Putnis 2007; Pollok et al. 2011; Ruiz-Agudo et al. 2014). In the temperature range of the
384 experiments, the solubility product of anhydrite varies between $10^{-4.36}$ at 120°C and $10^{-4.83}$ at
385 200°C (Freyer and Voigt 2003). The solubility product of Hap strongly decreases with

386 increasing temperature within this temperature range, being $10^{-58.33}$ at 120°C and $10^{-70.64}$ at
387 200°C (Kaufman and Kleinberg 1979). In contrast, the solubility product of β -TCP slightly
388 increases with temperature in the temperature range of the experiments from $10^{-26.63}$ at 120°C
389 to $10^{-25.42}$ at 200°C (Wang and Nancollas 2008). Considering the very large differences in
390 solubility between anhydrite and both calcium phosphates, it is likely that the latter phases
391 form under very high supersaturations and can be expected that this process occurs rapidly at
392 any experimental temperature. Therefore, anhydrite dissolution is most likely the rate-limiting
393 process. However, the progress of the reaction requires continuous chemical exchange with
394 the bulk solution and the rate of mass transport through the porous replaced rim may play a
395 significant role modulating the overall kinetics of the anhydrite by calcium phosphate
396 replacement reaction.

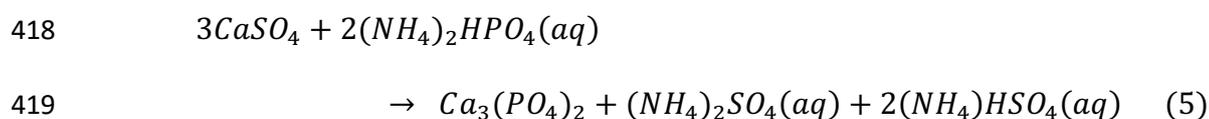
397 **Reaction pathway**

398 The mole ratio of β -TCP/Hap vary with temperature and reaction time. At 120°C, β -TCP/Hap
399 mole ratio first increases with reaction time and later slowly decreases. At all other
400 temperatures β -TCP/Hap mole ratio progressively decreases with increasing reaction time.
401 This decrease takes place much more rapidly with increasing temperature. In the temperature
402 range of the experiments, β -TCP is more soluble than Hap (lInl.dat) and, consequently, less
403 stable. Therefore, β -TCP forms as a thermodynamically metastable phase. Crystallization of
404 metastable phases often occurs under high supersaturations. The large solubility difference
405 between anhydrite, β -TCP and Hap determines that the layer of fluid at the interface will be
406 very highly supersaturated with respect to both, β -TCP and Hap. Under these conditions, both
407 calcium phosphate phases can nucleate. Indeed, the presence of Hap at very early stages of
408 the replacement process suggests a competition between the nucleation and growth of β -TCP
409 and Hap. Metastable β -TCP latter transforms into Hap according to an Ostwald ripening
410 process. This transformation most likely occurs through a dissolution-reprecipitation reaction

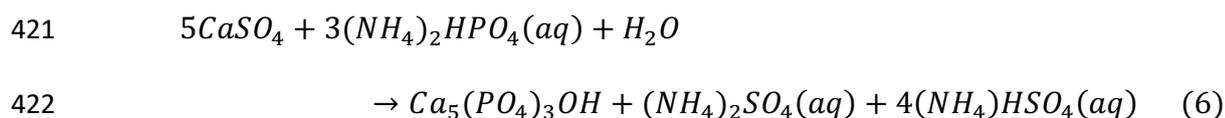
411 since dissolution pits are apparent in the surface of β -TCP and needle-like Hap crystals grow
412 spatially associated to those pits (Fig. 3b). The competition between the nucleation and growth
413 of β -TCP and Hap operates longer and plays a more important role at lower temperatures as
414 is evidenced by the initial increase in the β -TCP/Hap mole ratio as well as the large amount
415 of β -TCP in almost fully replaced samples at 120°C.

416 The replacement reaction can be described through the following equations:

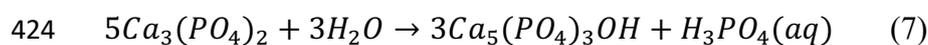
417



420



423



425

426 The precipitation of Hap, as a basic salt, results in a pH decrease in the fluid phase (equations
427 6 and 7). This is in good agreement with the change from an initial $pH_i = 8.1$ (1) to a pH of
428 6.8 (1) measured at the end of the experiments.

429 **Porosity generation**

430 Several factors explain the generation of porosity during pseudomorphic replacement
431 reactions. Porosity balances the loss of volume in reactions that involve a negative molar
432 volume change and guarantees the preservation of the original external shape of the primary
433 phase. In the case of the replacement of anhydrite ($V_{Anh} = 45.84 \text{ cm}^3/\text{mol}$) by mixtures of β -

434 TCP ($V_{\beta\text{-TCP}} = 33.14$) and Hap ($V_{\text{Hap}} = 31.79 \text{ cm}^3/\text{mol}$), all relevant molar volume changes are
435 negative: $-12.70 \text{ cm}^3/\text{mol}$ (27,7%) and -14.05 (30.65%), the changes associated to the
436 transformation of anhydrite into β -TCP and Hap, respectively, and -1.35 (4.07%) the change
437 associated to the transformation of β -TCP into Hap. These molar volume changes only very
438 slightly vary ($< 1\%$) depending on the experimental temperature (Evans 1979; Nakamura et
439 al. 1990). The porosity required to balance the molar volume change associated to the
440 replacement reaction will be around 30%, although the exact value will vary depending on the
441 β -TCP/Hap mole ratio in the calcium phosphate mixture. Since this ratio changes with reaction
442 time and temperature, small differences in porosity can be expected depending on the
443 experimental conditions.

444 A second source of porosity is the difference in solubility between the primary and product
445 phases (Putnis 2002, 2009; Pollok et al. 2011). β -TCP and Hap are less soluble than anhydrite.
446 This means that part of the dissolved anhydrite is lost to the solution when this phase is
447 replaced by β -TCP and Hap. The negative solubility change associated to the replacement will
448 also be balanced by generation of porosity that will add up to that that balances the negative
449 molar volume change. The volume of solubility change-related porosity is also influenced by
450 reaction time and temperature as both parameters determine the β -TCP/Hap in the reacted
451 sample. Differences in the porosity formed during the replacement may influence the mass
452 transport rate and be reflected by the overall kinetics of the replacement reaction (Jonas et al.
453 2013; Pedrosa et al. 2017).

454 **Textural features and evolution**

455 Mass transport from and to the interface is not only affected by the increase of porosity during
456 the replacement but also by the specific characteristics of this porosity regarding pore shape,
457 size and connectivity (Putnis et al. 2005; Jonas et al. 2013; Putnis 2015; Yuan et al. 2018;
458 Forjanés et al. 2020a). Yuan et al. (2018) studied the porosity generated during the

459 replacement of calcite by cerussite and identified three types of pores: Open pores that directly
460 connect the reaction front and the bulk solution, trapped pores, which are connected through
461 grain boundary diffusion between the crystals of the product phase and open pores to the bulk
462 solution, and isolated pores, located between differently oriented crystals of the product. All
463 types of pores are filled by fluid, but the chemistry of this fluid will vary due to differences in
464 the mass transport rate. Chemical exchange between the interfacial fluid and the bulk will be
465 more efficient if the porosity mainly consists of open pores, while the predominance of
466 isolated pores will strongly hinder mass transport and the progress of the reaction. Both habit
467 and textural relationships of the product crystals are key features that define the relative
468 amount of open, trapped and isolated pores in replaced samples. The existence/absence of
469 epitactic relationships between the product and the primary phase is another key factor. We
470 observe no evidence of epitatic growth of either β -TCP or Hap on anhydrite surfaces. β -TCP
471 crystals grow randomly oriented, while Hap grow as fan-like bunches of needle-like crystals
472 arranged approximately perpendicular to the anhydrite surface (Fig. 3). β -TCP crystals
473 accumulated in the central region of this bunches. Needle-like crystals within bunches that
474 seem to originate at the region occupied by β -TCP crystals are arranged with their long axis
475 approximately parallel between them. This arrangement is characteristic of competitive
476 growth between crystals that nucleate randomly oriented on a surface and have a preferential
477 growth direction. Those crystals oriented with the preferential growth direction perpendicular
478 to the surface can grow freely, while the growth of all differently oriented crystals rapidly
479 becomes prevented by the lack of space. It is noteworthy that the porosity generated during
480 replacement reactions has a transient nature, and its features evolve along time. In the case of
481 the replacement of anhydrite by calcium phosphate, the very different habit of β -TCP and Hap
482 crystals will determine differences in the relative abundance of different types of pores
483 depending on the mole β -TCP/Hap ratio in the replaced sample. As pores between elongated,

484 parallel crystals tend to be open, it can be expected that the number of open pores in replaced
485 samples will increase as more β -TCP transforms into Hap.

486 **Rate-limiting process**

487 Coupled dissolution-precipitation reactions are heterogeneous processes. Their kinetics is
488 either limited by the dissolution of the primary phase or by the precipitation of the product
489 phase(s). Several observations give support to the interpretation that the dissolution of
490 anhydrite is the limiting step of the anhydrite by calcium phosphate replacement reaction
491 Moreover, textural features of the reacted samples indicate that the reaction progresses from
492 rim to core. This has been interpreted by different authors as evidence of the precipitation of
493 the product(s) taking place at a faster rate than the dissolution of the primary phase. Altree-
494 Williams et al. (2015, 2017) pointed out that mineral replacement reactions which are rate-
495 limited by the precipitation of the product also result in the formation of pseudomorphs.
496 However, external features of these pseudomorphs are less accurately preserved as the
497 formation of overgrowths blurs them. Moreover, these pseudomorphs show hollow cores.
498 Neither the development of overgrowths nor the formation of a hollow core is observed during
499 the formation of calcium phosphate pseudomorphs after anhydrite.

500 Altree-Williams et al. (2019) calculated an average empirical activation energy (E_a) of $46.2 \pm$
501 7.6 kJ/mol for the carbonation of anhydrite. These authors interpreted the slightly lower value
502 of this activation energy compared to previously reported values of the activation energy for
503 anhydrite dissolution [50 kJ/mol, Bildstein et al. (2001); 61.0 ± 1.0 kJ/mol, Kontrec et al.
504 (2002)] as indicative of anhydrite dissolution being the rate-limiting event of anhydrite
505 carbonation kinetics. Moreover, they concluded that, since the reported values of the
506 activation energy for the diffusion of dissolved calcium are much smaller ($E_a = 12.6$ kJ/mol),
507 mass transport of species from the bulk solution to the reaction interface and backwards must
508 occur a faster rate compared to anhydrite dissolution. The average empirical activation energy

509 of the anhydrite by calcium phosphate replacement reaction determined in this work using the
510 isoconversional method is 39.9 ± 1.6 kJ/mol. The similarity of this value to the previously
511 reported activation energy for anhydrite carbonation supports that anhydrite dissolution also
512 controls the kinetics of anhydrite phosphation.

513 The experimental data fitted the Avrami model yields an $E_a = 40.2 \pm 1.9$ kJ/mol, which is
514 identical to that obtained using the isoconversional method regardless of the transformed
515 fraction considered. The calculated A is 9045.3 (hours)⁻¹ and the resulting integral law is Y_{CaP}
516 $= 9045.3e^{(-40/RT)t}$. Applying the law equation and assuming that the reaction mechanism is
517 the same as that controlling the process in the 120-200°C range, it can be estimated that at
518 20°C the time required to for the complete replacement of anhydrite single crystals by calcium
519 phosphate will be around 4 months. However, similar experiments performed under lower
520 temperature conditions need to be performed to confirm the soundness of this approach.

521 The shape of the rate curves in Figure 5 is approximately the same for those corresponding to
522 150°C, 180°C and 200°C and the curves can be described as isokinetic. Consistently, the
523 Avrami reaction exponents (n) corresponding to these three temperatures show almost
524 identical values (~ 1.35). The shape of the rate curve corresponding to 120°C is slightly
525 different, which correlates with a larger Avrami exponent (~ 1.66 at 120°C) (Fig. 5). The value
526 of Avrami exponents is characteristic of the rate-limiting kinetic mechanism that operates at
527 each temperature (Redfern 1987). All calculated Avrami exponents in this work stand between
528 those defined by Hancock and Sharp (1972) for first-order and higher-order processes.
529 However, it is important to keep into mind that Hancock and Sharp (1972) studied solid-state
530 transformations, while the replacement of anhydrite by calcium phosphate phase takes place
531 mediated by the presence of a fluid phase. Several authors have interpreted changes in the
532 value of the Avrami exponent with temperature as reflecting a balance between changes in the
533 solubilities of the primary and secondary phases combined with changes that affect mass

534 transport processes (Zhao et al. 2009; Aitree-Williams et al. 2019). The solubility of anhydrite
535 and apatite decrease with increasing temperature. However, this decrease follows a different
536 trend depending on the phase (Gregory et al. 1974; McDowell et al. 1977; Otálora and García-
537 Ruiz 2014; Inl database from PHREEQC code). Moreover, in the temperature range of the
538 experiments the solubility of β -TCP shows a positive dependence with temperature. These
539 differences in solubility evolution may be reflected by the Avrami exponent. Anhydrite
540 phosphation does not only involve the dissolution of anhydrite to form Hap. β -TCP
541 precipitation is metastable and, as the reaction progresses, β -TCP dissolves to form Hap. The
542 formation of β -TCP and its transformation into Hap may also weigh differently on the overall
543 kinetics of the replacement depending on the reaction temperature. Reported values of the
544 empirical activation energy of β -TCP dissolution are around 16.3 kJ/mol, which is much
545 smaller a value than that of the activation energy for the dissolution of anhydrite and very
546 close to the activation energy for the diffusion of dissolved calcium (Bohner et al. 1997).
547 Although its activation energy is small, the influence of β -TCP dissolution may sufficiently
548 influence the kinetics of Hap formation as to explain changes in the Avrami exponent.
549 Moreover, the β -TCP into Hap transformation differently affect textural features of the
550 reacted samples depending on the temperature. After complete replacement of anhydrite by
551 calcium phosphate, β -TCP content is highest in samples reacted at 120°C (~20%), followed
552 by those reacted at 150°C (~13%). Given the very different morphology of β -TCP crystals
553 compared to those of hydroxyapatite as well as the slightly larger molar volume of the former,
554 differences in porosity and pores arrangement may further explain small kinetics changes at
555 different reaction temperatures (Putnis et al. 2005; Jonas et al. 2013; Putnis 2015; Pedrosa et
556 al. 2017).

557

558 **Implications**

559 In this work we have analysed the kinetics of the replacement of anhydrite single crystals by
560 mixtures of the calcium phosphates phases, β -TCP and Hap, *via* a dissolution-precipitation
561 reaction that takes place under hydrothermal conditions (from 120 to 200°C). The fitting of
562 the experimental results to both, the Avrami equation and the isoconversional method yields
563 experimental activation energies, E_a , around 40 kJ/mol. This value is similar to that
564 determined by Aitree-Williams et al. (2019) for the carbonation of anhydrite, which is rate-
565 limited by anhydrite dissolution. However, Avrami exponents stand between those defined
566 by Hancock and Sharp (1972) for first-order and higher-order processes. We interpret that this
567 apparent contradiction reflects the complexity of coupled dissolution-precipitation reactions,
568 whose kinetics is modulated by a number of factors, including the dissolution rate of the
569 primary phase, the volume, connectivity and permeability of the porosity formed during the
570 replacement, the precipitation rate of the secondary phase(s) and the textural evolution of the
571 latter. In the case under consideration, we conclude that there exists a competition between
572 the nucleation and growth of β -TCP and Hap at early stages of the replacement, which last
573 longer the lower the reaction temperature is. Due to the very different morphological features
574 of β -TCP and Hap crystals, this competition strongly influences the texture of the replacement,
575 its pores arrangement and, thereby, mass transfer from and to the interfacial fluid, emphasizing
576 the important role of metastable precursor phases and Ostwald ripening processes in defining
577 the overall kinetics of coupled dissolution-precipitation reactions. The results of this study
578 indicated that the replacement of sedimentary rock-forming minerals during diagenesis is a
579 feasible mechanism for the formation of some calcium phosphate accumulations in
580 sedimentary basins. In this study anhydrite single crystals were used. Anhydrite sedimentary
581 rocks are polycrystalline materials that consist of grains bounded by differently oriented
582 crystal surfaces. The reactivity of these surfaces may differ from that of the anhydrite cleavage
583 surfaces. Moreover, the presence of differently oriented crystal surfaces may facilitate the

584 epitactic growth of the product phases on the parent phase substrate. It is well established that
585 the formation of epitactic overgrowths contributes to passivate the substrate and commonly
586 hinders the progress of replacement reactions (Rodríguez-Blanco et al. 2007; Forjanés et al.
587 2020a). On the other hand, the network of grain boundaries in rocks facilitates fluid infiltration
588 and provides an increased reactive surface compared to single crystal, which may influence
589 the reaction mechanism (Jonas et al. 2013, 2014). Future experimental work using anhydrite
590 evaporitic rock samples will help to elucidate this influence. The results of this study also shed
591 light on the mechanisms involved in the removal and immobilization of an important
592 pollutant, P, by mineral surfaces. Due to the main role of P in the life cycle, the scarcity of
593 this element in the Earth's crust and its ability to pollute soils, sediments and ground and
594 running waters, developing methods for the recovery of P through the precipitation of
595 phosphate phases on the surface of pre-existing minerals would have an outstanding
596 importance. The results of this study together to further experimental work using
597 polycrystalline samples and fluids with P concentrations closer to those found in natural
598 environments can contribute to reach this goal.

599

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607

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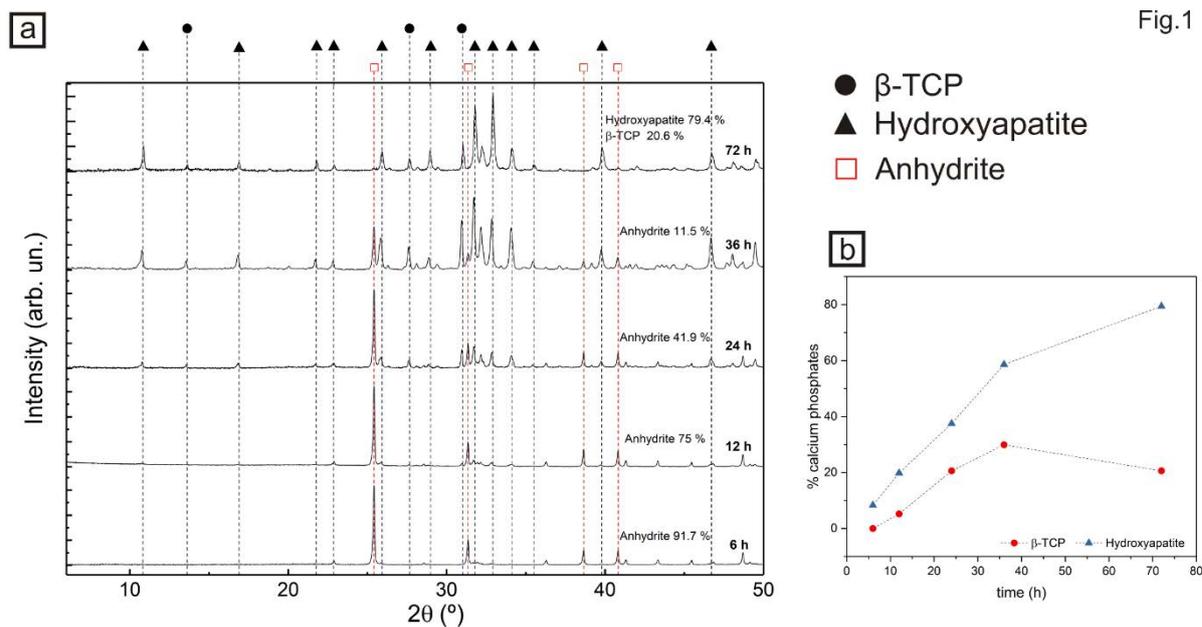
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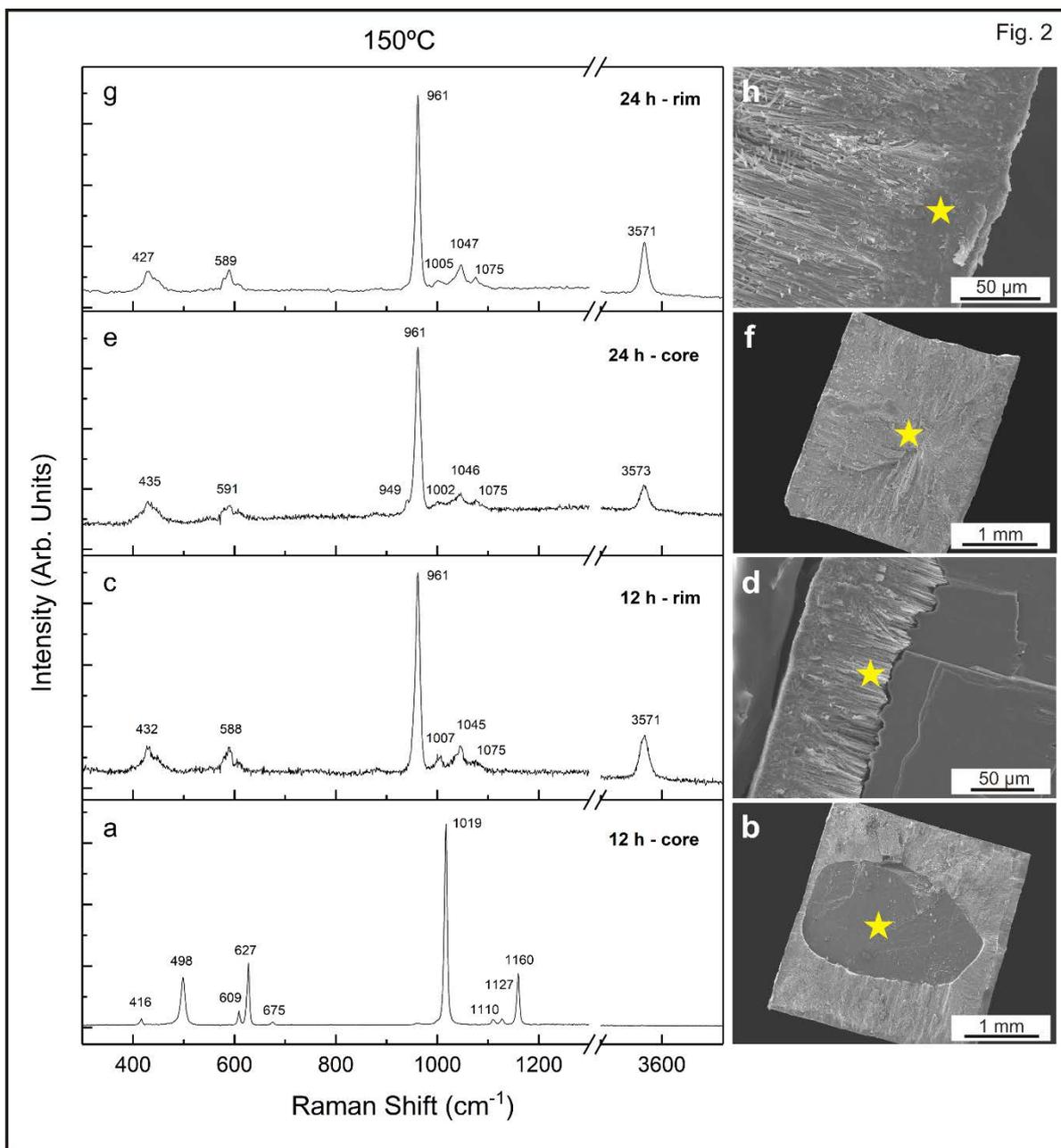
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806 Figure 1. (a) X-ray powder diffraction patterns showing the mineralogical evolution of the
807 replacement of anhydrite by calcium phosphates after hydrothermal interaction experiments
808 carried out at 120°C at different reaction times. The reflections of anhydrite, hydroxyapatite
809 and β -TCP were marked with squares, triangles and circles, respectively. (b) % calcium
810 phosphates, calculated by Rietveld refinement, against reaction time.

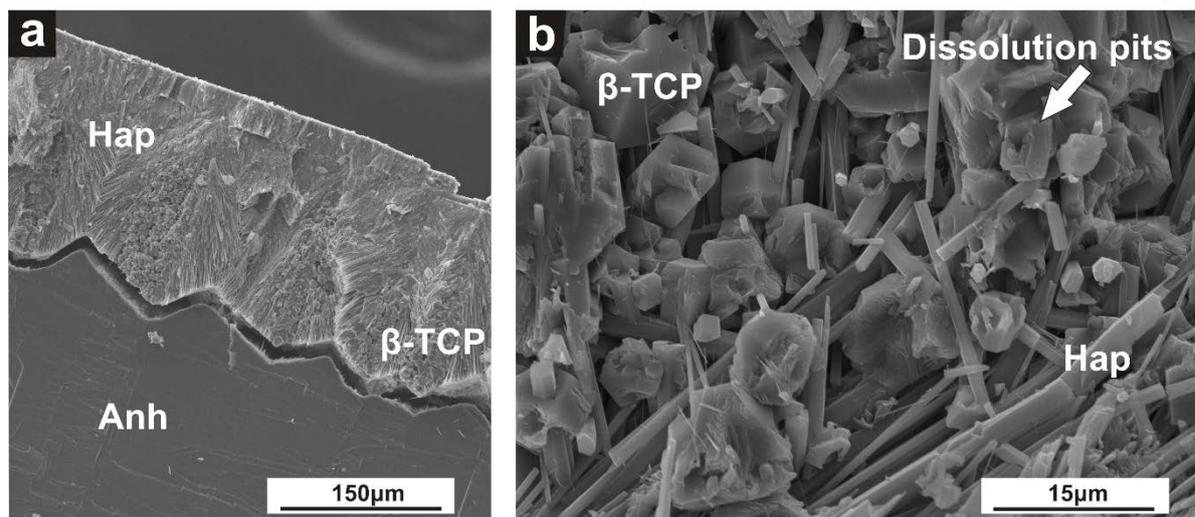
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812
813 Figure 2. Raman spectra and SEM images of cross-section of partially replaced anhydrite
814 samples for 12 h and 24 h at 150°C. The main vibrational bands of SO₄²⁻ identified (a) in
815 the core of anhydrite sample reacted (b) match well with anhydrite. The main vibrational
816 bands of PO₄³⁻ (c, e and g) identified in the rim (d and h) and core (f) of reacted samples
817 corresponds to hydroxyapatite. The broad band at 949 cm⁻¹ and the presence of the band at
818 1045 cm⁻¹ (d) are characteristic of β-TCP. SEM micrographs of the reacted samples show
819 the position where the spectra were measured.

820

Fig. 3



821

822

823 Figure 3. SEM micrographs of anhydrite (Anh) reacted with 2M (NH₄)₂HPO₄ aqueous

824 solution at different temperatures and reaction times. (a) After 3 h at 150°C a cross cut section

825 image shows the unreacted anhydrite core and the two calcium phosphates that can be

826 distinguished, β-tricalcium phosphate (β -TCP) and hydroxyapatite (Hap). Hydroxyapatite is

827 clearly more abundant and is growing in elongated hexagonal needle-like crystals that are

828 arranged perpendicular to the unreacted anhydrite core and β-TCP shows an euhedral habit.

829 (b) At higher temperatures and shorter reaction times (200°C and 2 h), β -TCP appears with

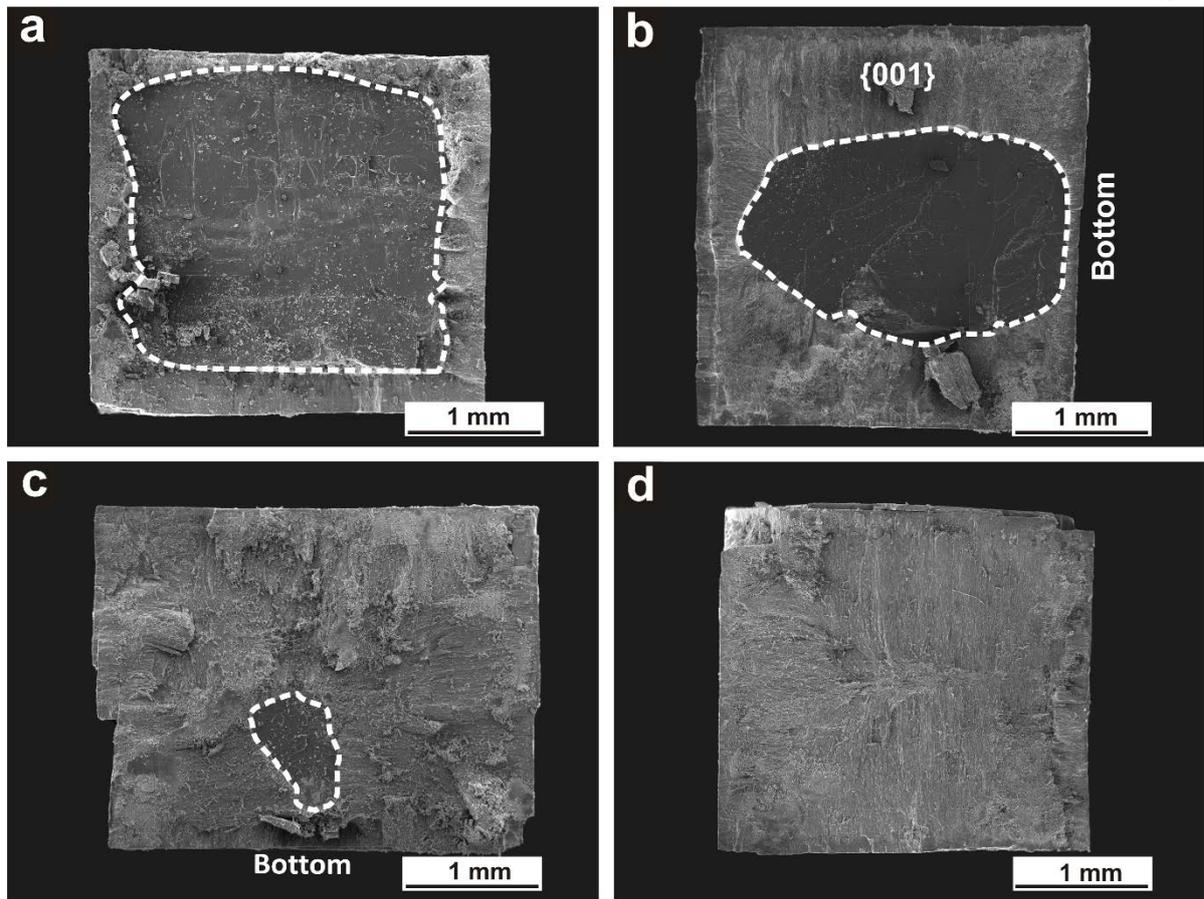
830 dissolution features and at the same time the development of smaller and thinner

831 hydroxyapatite crystals from β -TCP faces.

831

832

Fig. 4

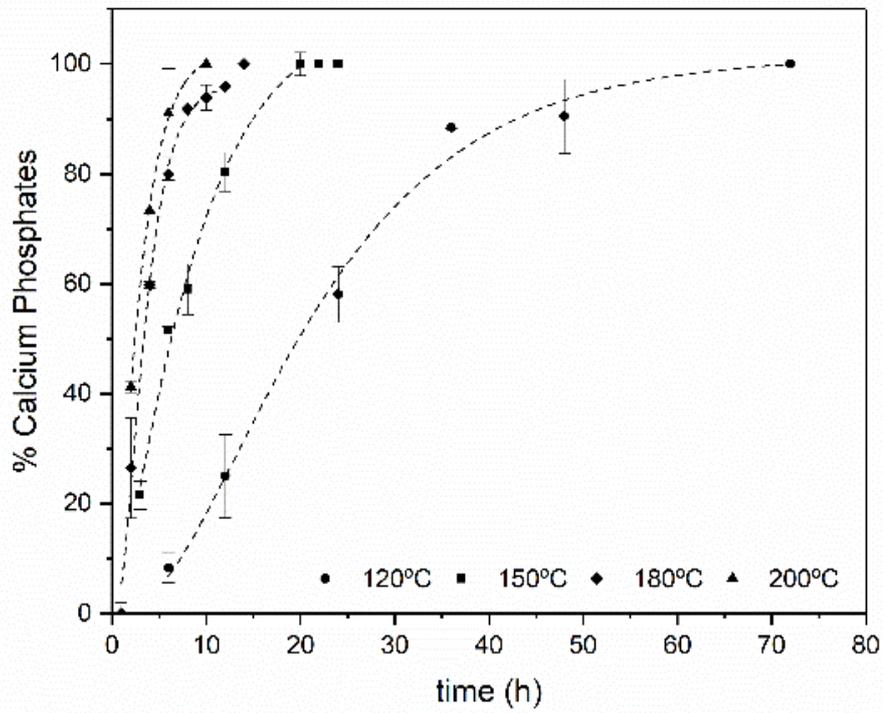


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Figure 4. SEM micrographs of cross-sections surfaces showing the anhydrite replacement by calcium phosphates after 12 hours of reaction at temperatures of 120 (a), 150 (b), 180 (c) and 200 (d).

838

Fig. 5



839

840 Figure 5. Fraction of calcium phosphates replacing anhydrite crystals against the reaction time
841 for each isothermal experiments. The lines are fitted to the Chapman equation.

842

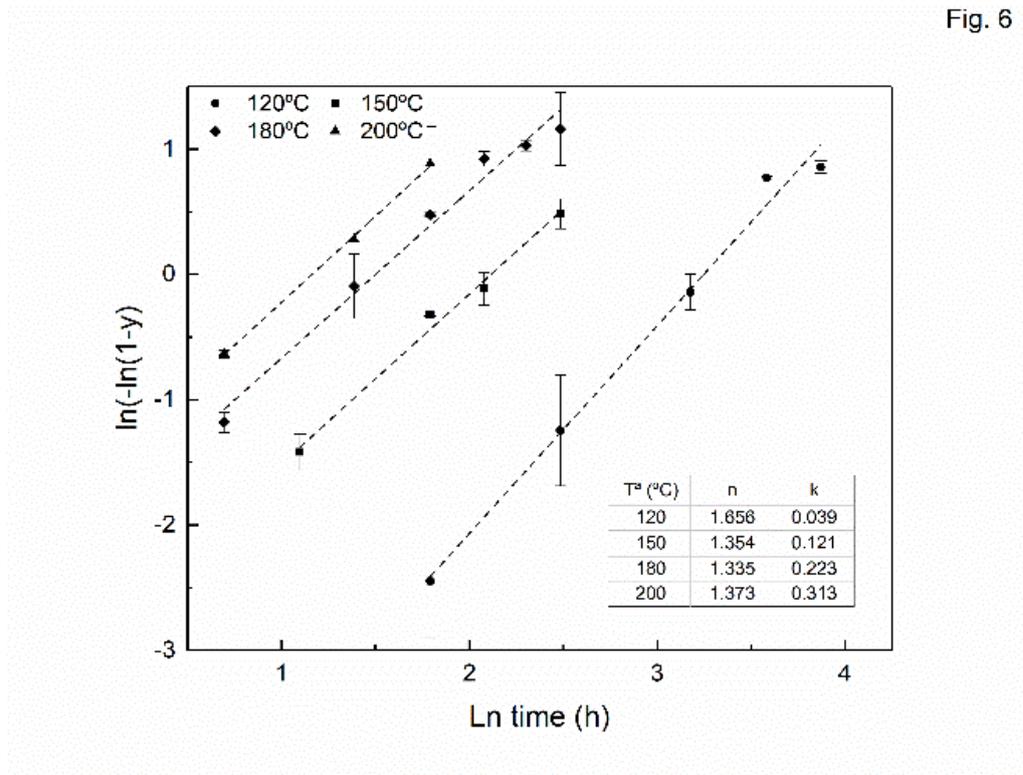
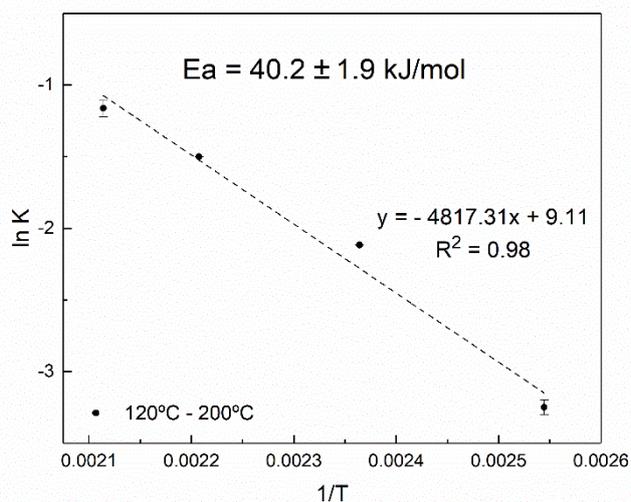


Fig. 6

843

844 Figure 6. Plot of the calcium phosphates transformed fraction (y) versus time. From the slope
845 and the intercept, Avrami rate law parameters (n and k) are determined.

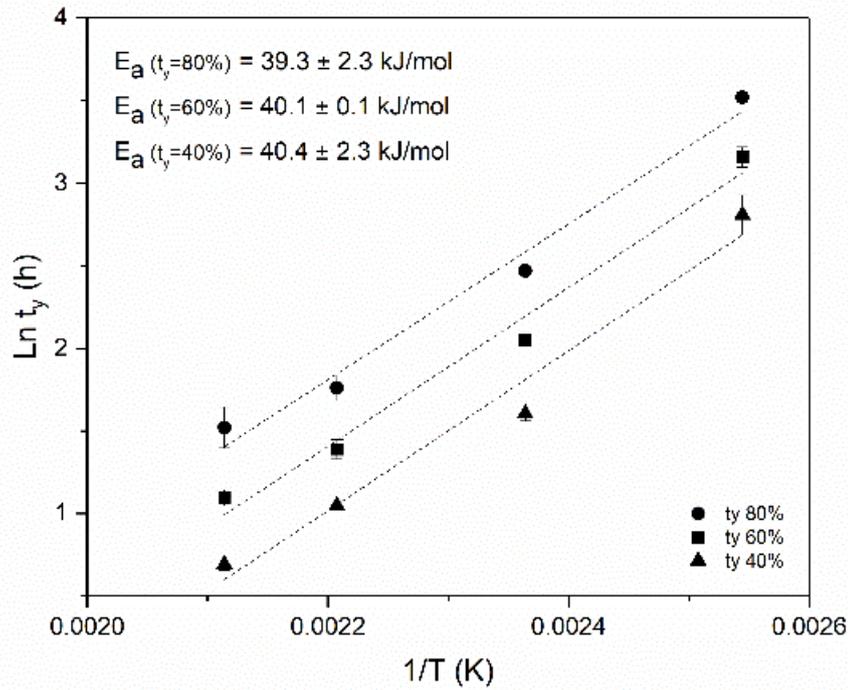
Fig. 7



846 Figure 7. Linear fitting of the rate constant K versus the reciprocal of temperature (expressed
847 in Kelvin).

848

Fig. 8



849 Figure 8: Experimental data of the time ($\ln t_y$) for several fractions ($Y=40, 60$ and 80%) of
850 transformed anhydrite versus $1/T$. The slope of the fitting straight lines corresponds to E_a/R .

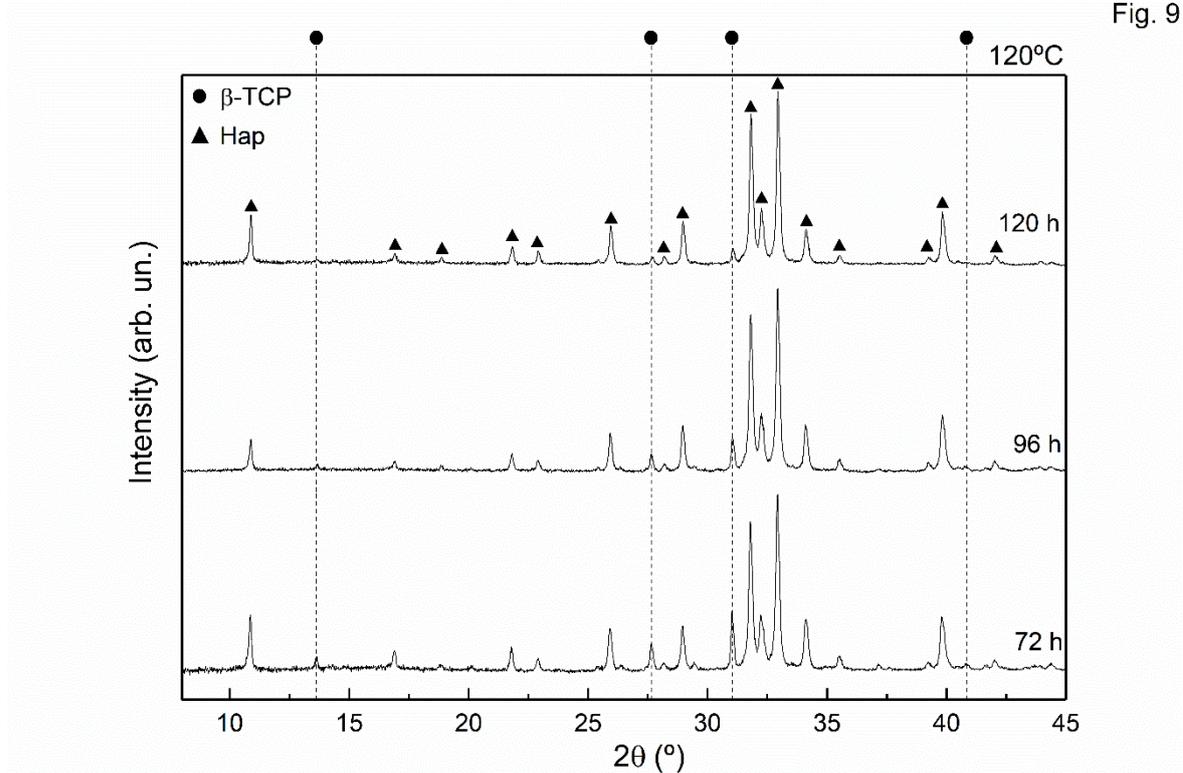
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852

853 **Supplementary**

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855



856

857 Figure 9: X-ray powder diffraction patterns showing the mineralogical evolution after
858 complete replacement of anhydrite at 120°C. The amount of β -TCP continues to decrease with
859 time.

860

861

120°C				150°C				180°C				200°C			
t (h)	% Anh	% β-TCP	% Hap	t (h)	% Anh	% β-TCP	% Hap	t (h)	% Anh	% β-TCP	% Hap	t (h)	% Anh	% β-TCP	% Hap
6	91.7		8.3	3	78.5	3.2	18.3	2	73.5	12.9	13.6	2	58.8	27.1	14.1
12	75	5.2	19.8	6	48.4	20.4	31.3	4	40.3	21.9	37.9	4	26.8	6.6	66.7
24	41.9	20.6	37.5	8	41	22.5	36.5	6	20	16.7	63.3	6	8.9	4.4	86.7
36	11.5	29.9	58.6	12	19.8	21.1	59.2	8	8.2	7.7	84.2	10	0	0	100
48	9.5	17.2	73.3	20	0	13.5	86.5	10	6.1	3.9	90				
72		20.6	79.4					12	4.1	2	93.9				
								14	0	0	100				

* R_{wp} is approximately 10% in all refinements.

862

863 Table 1. Calculated transformation (% calcium phosphates) calculated by Rietveld refinement of X-ray diffraction patterns during hydrothermal
 864 experiment.

865