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Kinetics of antigorite dehydration: a trigger for lower-plane seismicity in subduction zones

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Abstract

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Antigorite dehydration experiments were performed under ambient pressure using a non-isothermal thermogravimetric analysis. Antigorite, with a grain size of 5-10 μm , was analyzed using heating rates of 10, 15, 20, and 25 $\text{K} \cdot \text{min}^{-1}$ at temperatures of up to 1260 K. The results show that the progress of the dehydration reaction varies with the heating rate, and the dehydration reaction of antigorite occurs within a temperature range of 800-1050 K. Several models were used to fit the dehydration results, and the double-Gaussian distribution activation energy model (2-DAEM) yielded the best fit to the experimental data. The dehydration kinetics of antigorite follow 2-DAEM, and there is a compensation effect between the pre-exponential factor and the average activation energy. The activation energy of the first step of antigorite dehydration stretches out a wide interval, the second step has a significantly higher activation energy, distributed on a narrower interval. We determined that the release rate of H_2O is 8.0×10^{-5} and $2.1 \times 10^{-3} m_{fluid}^3 m_{rock}^{-3} \text{ s}^{-1}$ at 893 K and 973 K, respectively, which are near the onset temperature for the isothermal dehydration reaction. Our results indicate that antigorite dehydration is fast enough to induce mechanical instabilities that may trigger seismicity in the lower plane of the double seismic zone.

52

53 **Introduction**

54 In subduction zones aqueous fluids released during the dehydration of hydrous
55 minerals leads to metasomatism and partial melting in the overlying mantle (e.g.
56 Scambelluri et al. 2001), which, in turn, induces magmatism (e.g. Hattori and Guillot
57 2003; Peacock and Wang 1999; Ulmer and Trommsdorff 1995) and triggers
58 earthquakes (e.g. Hacker et al. 2003; Jung et al. 2004; Okazaki and Hirth 2016;
59 Peacock 2001; Yamasaki and Seno 2003). In particular, earthquakes occurring at
60 intermediate depths (50-200 km) in double seismic zones (DSZ; Yamasaki and Seno
61 2003) comprising upper and lower Wadati-Benioff planes may be related to
62 dehydration (Ferrand et al. 2017; Hacker et al. 2003; Incel et al. 2017; Peacock 2001).
63 The dehydration process in subducting oceanic slabs induces several geophysical
64 anomalies including low seismic velocity (Ferrand et al. 2017; Hacker et al. 2003;
65 Jung et al. 2004; Peacock 2001; Yamasaki and Seno 2003) and high conductivity
66 anomalies (Soyer and Unsworth 2006; Wang et al. 2017; Worzewski et al. 2010).

67 Recent publications have directly addressed antigorite dehydration and its
68 implications for the seismicity of subducting mantle (e.g. Ferrand et al. 2017; Gasc et
69 al. 2017). Several experiments have shown that antigorite dehydration is not “seismic”
70 (Chernak and Hirth 2010, 2011; Gasc et al. 2017; Okazaki and Hirth 2016).
71 Particularly, it was found that, using a ratio of heating rate to strain rate that is typical
72 for actual subducting slabs (100-1000 K), antigorite samples deform aseismically
73 (Chernak and Hirth 2011; Gasc et al. 2017; 2011), whereas antigorite-olivine mixtures

74 undergo seismic events (Ferrand et al. 2017). Most likely, a dehydration-driven stress
75 transfer (DDST) triggers earthquakes in fresh peridotite at the tip of dehydrating faults
76 (Ferrand et al. 2017), as supported by recent field observations (Ferrand et al. 2018;
77 Plümper et al. 2016; Scambelluri et al. 2017). Ferrand et al. (2017) experimentally
78 demonstrated that intermediate-depth earthquakes, particularly those of the lower
79 Wadati-Benioff planes (LWBP), are triggered by the dehydration of very limited
80 amounts of antigorite, both for > 0 and < 0 volume changes, which means that these
81 earthquakes are not a result of aqueous fluid overpressure. Nonetheless, the role of
82 released aqueous fluid by antigorite still needs to be investigated.

83 Serpentine minerals are the most abundant hydrous minerals in altered
84 ultramafic rocks (Hyndman and Peacock 2003) and contain up to approximately 13 wt%
85 H₂O (Schmidt and Poli 1998). Serpentine has three main varieties: lizardite, chrysotile
86 and antigorite (Ulmer and Trommsdorff 1995). Antigorite is the only serpentine
87 mineral stable above temperatures of ~ 875 K in subduction zones (Reynard et al.
88 2007; Shao et al. 2014). Several studies have considered that the limit of the antigorite
89 stability field is near an isotherm that fits the lower plane earthquake hypocenters (e.g.
90 Abers et al. 2013; Peacock 2001; Yamasaki and Seno 2003), i.e., between 1 and 5 GPa.
91 Therefore, antigorite dehydration has been regarded as a candidate to explain the
92 seismicity of the lower plane of the double seismic zone (Dobson et al. 2002; Hilairet
93 et al. 2007; Omori et al. 2004; Peacock 2001).

94 The dehydration of serpentine minerals has been the subject of many previous
95 studies (Eggler and Ehmman 2010; Gualtieri et al. 2012; Trittschack and Grobéty 2012;

96 Viti 2010). Recently, synchrotron HP experiments with acoustic recordings (Ferrand
97 et al. 2017) have demonstrated that antigorite dehydration leads to peridotite
98 embrittlement through a stress transfer from dehydrating serpentinized faults to fresh
99 peridotite volumes. However, the kinetics of antigorite dehydration still requires
100 investigations. Previously, antigorite dehydration was investigated using
101 time-resolved X-ray diffraction (Chollet et al. 2011; Gualtieri et al. 2012; Perrillat et
102 al. 2005), high-temperature infrared microspectroscopy (Sawai et al. 2013) and
103 thermogravimetric analysis (Pérez-Rodríguez et al. 2005; Viti 2010; Weber and Greer
104 1965). These studies have led to quite different conclusions. Dehydration is controlled
105 by surface growth processes at the edges of grains, as shown by X-ray diffraction
106 (Perrillat et al. 2005); in contrast, the one-dimensional diffusion process matches well
107 with the experimental data using in situ high-temperature infrared microspectroscopy,
108 which enables the water species to be distinguished (Sawai et al. 2013). A recent
109 thermogravimetric analysis using three heating rates shows that the dehydration
110 kinetics follow a three-dimensional phase boundary reaction model (Balucan et al.
111 2011).

112 In addition, most previous studies of antigorite dehydration kinetics have used
113 isothermal methods. Isothermal methods have some disadvantages that can be
114 overcome by non-isothermal methods. For example, non-isothermal methods can
115 determine changes in the activation energy as the reaction progresses. Each change
116 indicates that either the contribution of an individual parallel step changes or that a
117 new reaction step contributes to the overall rate (Trittschack and Grobéty 2012;

118 Vyazovkin and Wight 1997; Wang et al. 2015).

119 In this work, our goal was to study the kinetics of antigorite dehydration
120 using non-isothermal experiments to determine the best kinetic model and therefore to
121 understand the constraints on the fluid production rate resulting from serpentine
122 dehydration, which, in turn, should impact intermediate-depth seismicity in
123 subduction zones.

124

125 **Experimental Procedure**

126 **Sample Preparation**

127 Antigorite starting material was collected from the Nagasaki metamorphic belt
128 in Japan. The pure antigorite crystals were picked to serve as samples. The chemical
129 composition of this material was analyzed using an Electron Probe MicroAnalyzer
130 (EPMA) (JXA-8230 at HeFei University of Technology). The chemical composition
131 of the sample is $Mg_{2.75}Fe_{0.09}Al_{0.06}Si_{2.03}O_5(OH)_4$ (see detail in **Table S1**), which was
132 also presented by Wang et al. (2017). The pure antigorite crystals were crushed and
133 ground to obtain powder. Various grain size fractions were separated from the
134 powdered material according to Stokes's law. Particles with diameters ranging from 5
135 to 10 μm were selected and stored above 423 K in a vacuum oven for one week,
136 eliminating absorbed molecular H_2O .

137 **Experimental methods**

138 **Dehydration experiments**

139 The dehydration experiments were performed using thermogravimetric (TG)

140 analysis with a Q600 SDT device at Peking University. Non-isothermal runs were
141 monitored to study antigorite dehydration with heating rates of 10, 15, 20, and 25
142 $\text{K} \cdot \text{min}^{-1}$. The powdered samples were placed in corundum sample holders. The
143 sample weights for the non-isothermal experiments were 5.055, 5.024, 5.059 and
144 5.072 mg, respectively. The samples were heated in an N₂ inert atmosphere up to a
145 temperature of 1260 K. The data were collected at an interval of 1 second.

146 **First-principle calculations**

147 Our first-principle calculations were performed with CASTEP (Clark 2005) codes
148 based on the plane wave basis set, the norm-conserving pseudopotentials (Hamann et
149 al. 1979; Lin et al. 1993) for electron-ion interaction and the local density
150 approximation (LDA) for the exchange-correlation interaction. The Brillouin zones
151 were sampled using the Γ -point for antigorite. The cut-off energy was set to 500 eV,
152 and the SCF tolerance was $5 \times 10^{-7} \text{ eV} \cdot \text{atom}^{-1}$.

153

154 **Kinetic models**

155 To describe the dehydration progress under non-isothermal conditions, the
156 mass fraction of the released volatiles (α) is defined as:

$$157 \quad \alpha = \frac{m_0 - m_T}{m_0 - m_f} \quad (1)$$

158 where m_0 , m_T and m_f are the initial mass, the mass at temperature T and the final
159 mass of the sample, respectively.

160 The general rate equation for a non-isotherm reaction is:

$$161 \quad \frac{d\alpha}{dT} = \frac{kf(\alpha)}{\beta} \quad (2)$$

162 where $f(\alpha)$ is a mechanism function, $\beta = \frac{dT}{dt}$ is the heating rate, k is the rate
163 constant defined as:

$$164 \quad k = A \exp\left(-\frac{E}{RT}\right) \quad (3)$$

165 Combining **Eq. 2** and **Eq. 3**, an integrated form of the general rate equation for a
166 non-isothermal reaction is shown below (Ozawa 1965; Wang et al. 2015).

$$167 \quad G(\alpha) = \int_0^\alpha \frac{d(\alpha)}{f(\alpha)} = \frac{A}{\beta} \int_0^T \exp\left(-\frac{E}{RT}\right) dT \quad (4)$$

168 where $G(\alpha)$ is a mechanism function with an integrated form, E is the apparent
169 activation energy, and R is the gas constant. The apparent activation energy is
170 calculated using the Flynn-Wall-Ozawa (FWO) method (Flynn and Wall 1967; Ozawa
171 1965, 1986), which is expressed as:

$$\ln\beta = \ln\frac{AE}{RG(\alpha)} - 5.3308 - 1.052\frac{E}{RT} \quad (5)$$

172

173 **Results and Discussion**

174 **Fig. 1** shows the thermogravimetric data for natural antigorite with heating rates
175 of 10, 15, 20, and 25 K · min⁻¹. The peak temperature of antigorite dehydration
176 increases with increasing heating rate. The dehydration trend of talc shows similar
177 behavior (Wang et al. 2015). Antigorite dehydration is observed within the
178 temperature range of 800 K to 1050 K. The average weight loss is as high as 12.75
179 wt%, which is in agreement with the theoretical structural water content of antigorite
180 of ~13 wt% (Mellini et al. 1987; Uehara 1998), and the heating rate does not affect
181 the mass loss.

182 The activation energy derived from the FWO method changes from 130 to 400

183 $\text{kJ} \cdot \text{mol}^{-1}$ as shown in **Fig. 2**. The apparent activation energy increases as the
184 reaction progresses in our experiments. We analyzed the non-isothermal kinetic data
185 using the n -th order reaction. The n -th order reaction integral forms are as follows
186 (Ozawa 1970):

$$G(\alpha) = \frac{(1 - (1 - \alpha)^{1-n})}{1 - n} \quad n \neq 1 \quad (6)$$

187 The fitting parameters and errors are listed in **Table 1**. The best-fit value of n
188 is 2.2. However, the 2.2-order model cannot fit the data well, indicating that our
189 experimental data cannot be interpreted by the n -th order reactions, particularly within
190 the temperature range of 800-950 K as shown in **Fig. 3**. A problem arises: the
191 apparent activation energy (E) is a fixed value in the n -th reactions, but still the
192 activation energy of antigorite dehydration determined via the FWO method
193 fluctuates substantially as shown in **Fig. 2**. Antigorite dehydration may consist of
194 several parallel reactions. Therefore, we attempt to find a new model to describe
195 antigorite dehydration using a double-Gaussian distribution activation energy model
196 (2-DAEM).

197 The distribution activation energy model (DAEM) assumes that the
198 dehydration mechanism uses an infinite number of irreversible, independent, parallel
199 reactions of different activation energies that reflect variations in the bond strengths of
200 the species (Cai et al. 2014; Navarro et al. 2008; Várhegyi et al. 2011; Zhang et al.
201 2014). The different activation energies are represented by a continuous distribution
202 function. The double-Gaussian distributed activation energy model (2-DAEM) is
203 based on the assumption that two sets of parallel reactions have different distributions

204 of activation energy (de Caprariis et al. 2012; 2015). Their equations can be written
 205 as:

$$206 \quad \begin{cases} \alpha = 1 - \int_0^{+\infty} \left(\exp \left(- \int_0^T \frac{A}{\beta} \exp \left(\frac{-E}{RT} \right) dT \right) \right) G(E) dE & n = 1 \quad (7) \\ \alpha = 1 - \int_0^{+\infty} (1 - n) \left(- \int_0^T \frac{A}{\beta} \exp \left(\frac{-E}{RT} \right) dT \right)^{\frac{1}{1-n}} G(E) dE & n \neq 1 \quad (8) \end{cases}$$

207 where

$$G(E) = \frac{\omega}{\sigma_1 \sqrt{2\pi}} \exp \left(\frac{-(E - E_1)^2}{2\sigma_1^2} \right) + \frac{1 - \omega}{\sigma_2 \sqrt{2\pi}} \exp \left(\frac{-(E - E_2)^2}{2\sigma_2^2} \right) \quad (9)$$

208 In the aforementioned equation, E is the activation energy, A is the pre-exponential
 209 factor, R is the gas constant, β is the heating rate, n is the reaction order, T is the
 210 absolute temperature, $G(E)$ is the distribution of activation energies, and ω is the
 211 weighting factor ranging from 0 to 1 that describes the percentage of water that is
 212 released during the primary or secondary dehydration steps. $\omega = 1$ if all the water is
 213 released during the first step, and $\omega = 0$ if it is completely produced during the
 214 secondary step. E_1 and E_2 are the mean activation energies of the first and second
 215 reactions, respectively, and σ_1 and σ_2 are their standard deviations. In this equation,
 216 the seven parameters to be estimated are n , ω , E_1 , E_2 , σ_1 , σ_2 and A . $G(E)$ satisfies
 217 the following condition:

$$\int_0^{+\infty} G(E) dE = 1 \quad (10)$$

218 $G(E)$ has many forms, such as a Gaussian, Weibull or Gamma distribution (Cai et al.
 219 2014).

220 In the antigorite lattice (Uehara 1998), two different environments exist for OH
 221 groups along the [010] axis (Capitani and Mellini 2004). Thus, the double-Gaussian

222 model (**Eq. 9**) is justified for $G(E)$. The O-H1 bond is surrounded by silica
223 tetrahedrons (shown in blue in **Fig. 4**). We calculated the O-H bond lengths for
224 antigorite using ab initio computational methods and found that the lengths of the
225 O-H1 and O-H2 bonds are 0.970-0.971 Å and 0.974-0.981 Å, respectively. This result
226 implies that the energies required to break these bonds are different, as shown by their
227 different apparent activation energies upon dehydration.

228 In this work, $n = 1$ was chosen, and A varied between 10^8 s^{-1} and 10^{16} s^{-1} .
229 To explore antigorite dehydration, a simulated annealing algorithm (van Laarhoven
230 and Aarts 1987) was used to determine ω , E_1 , E_2 , σ_1 , σ_2 and A . Eq. 7 is
231 numerically integrated with the inner integral computed using the Gaussian
232 quadrature and the outer using the Romberg integral algorithm. An objective function
233 Δ determines the accuracy of the six parameters. The lower the Δ , the more accurate
234 the six parameters. This function Δ can be written as follows:

$$\Delta = \sum_{i=1}^M \sum_{j=1}^N (\alpha_{tij} - \alpha_{eij})^2 \quad (11)$$

235 where α_{tij} is the α value computed using Eq. 7 when the six parameters are
236 known, α_{eij} is the α value measured via experiments, N is the number of data
237 points in each heating rate, and M is the sum of the heating rates. The six parameters
238 were determined using the simulated annealing algorithm and minimizing the
239 objective function value.

240 As previously mentioned, the pre-exponential factor is shared by all the
241 reactions in the adopted model. The pre-exponential factor is highly correlated with

242 the activation energy (Anthony et al. 1975; Du et al. 1991; Miura and Maki 1998),
243 which leads to an unreliable value for the pre-exponential factor from the fit. This
244 finding means that when the value of the pre-exponential factor is fitted together with
245 the other parameters, non-unique values are typically found because of the
246 compensation effect. In other words, different pairs of kinetic parameters provide an
247 equally good fit to the experimental data. Thus, the pre-exponential factor must be
248 fixed. We fixed the pre-exponential factor A within a range of $10^8 - 10^{16} \text{ s}^{-1}$, which
249 is consistent with the transition-state theory (Várhegyi et al. 2011; Zhang et al. 2014)
250 and yields the ranges of $10^8 - 10^{12} \text{ s}^{-1}$, $10^8 - 10^{13} \text{ s}^{-1}$, $10^8 - 10^{14} \text{ s}^{-1}$,
251 $10^8 - 10^{15} \text{ s}^{-1}$, and $10^8 - 10^{16} \text{ s}^{-1}$ for the pre-exponential factor, resulting from
252 the simulated annealing algorithm searching in some ranges. The six parameters can
253 be obtained using the simulated annealing algorithm when the pre-exponential factor
254 is within a fixed range and $n = 1$. The fitting parameters are listed in **Table 2**. $\log(A)$
255 is linearly correlated with the average activation energies (**Fig. 5**); $\log(A)$ increases
256 with increasing activation energy. The correlation coefficients between $\log(A)$ and the
257 average activation energies (E_1 and E_2) are 0.9906 and 0.9986, respectively, which
258 are very high values as expected. As shown in **Fig. 5**, multiple and interrelated values
259 of the kinetic parameters are allowed by the fitting algorithm. The higher the A value,
260 the higher the average activation energy. We show a clear relationship between the
261 pre-exponential factor and the average activation energy, termed the compensation
262 effect.

263 Since the atomic vibration frequency is typically on the order of $\sim 10^{13} \text{ Hz}$, it is

264 reasonable to fix the pre-exponential factor on the order of $\sim 10^{13} \text{ s}^{-1}$. We adopted the
265 fitting parameters (**Table 2**) to obtain the dehydration kinetics when the
266 pre-exponential factor is fixed on the order of $\sim 10^{13} \text{ s}^{-1}$. Using these fitting
267 parameters, we calculate the reaction progress at different heating rates and show
268 comparisons with the experimental data as shown in **Fig. 6**. It seems that the 2-DAEM
269 works well with a very small $\Delta = 0.031$ and a correlation coefficient of 0.999
270 between the experimental data and the fitting results. Such a good fit of the
271 experimental results means that the kinetics of antigorite dehydration are well
272 modeled via 2-DEAM.

273 Processes G_1 and G_2 are referred to as the slow and the fast steps of
274 dehydration, respectively. The structure of the antigorite crystal lattice (**Fig. 4**)
275 contains two different local environments for OH groups, and the O-H1 and O-H2
276 bonds have different lengths. In addition, the O-H3 bond is located at the boundary
277 between the O-H1 and O-H2 bonds. Thus, the O-H3 bond is expected to show
278 transitional behavior. The O-H1 bonds characterize 25.0% of the total OH groups;
279 considering the O-H3 bonds as O-H1 bonds, their fraction reaches 34.4%. The
280 weighting factor ω (**Eq. 7**) is equal to 0.336, which is between 25% and 34.4%,
281 implying that the O-H3 bond breaks during the transition between the slow and fast
282 steps.

283 To explain antigorite dehydration kinetics, two sets of parallel dehydration
284 reaction models are illustrated in **2-DAEM**. In our model, parallel dehydration
285 reactions including a slow and fast mechanism occur at the same time. The weighting

286 factor ω is 0.336, indicating that the fast step is dominant during the dehydration
287 processes. In a previous study, the dehydration kinetics were investigated using
288 time-resolved synchrotron X-ray diffraction in a closed system, showing that the
289 dehydration occurs during two steps (Chollet et al. 2011), which impacts the
290 seismicity (Ferrand et al. 2017).

291 During the first step (Chollet et al. 2011), according to backscattered electron images,
292 only ~35% of the antigorite produces olivine and talc-like (intermediate hydrous
293 phase; Perrillat et al. 2005), which is consistent with the weighting factor of 0.336 in
294 the present study.

295 The probability densities of the G_1 and G_2 processes are plotted as a function
296 of E (**Fig. 7**). The whole process follows a continuous distribution function. The
297 activation energies that correspond to the slow step of antigorite dehydration (E_1) are
298 widely distributed, yielding a mean activation energy of $268.1 \text{ kJ} \cdot \text{mol}^{-1}$, while the
299 energies corresponding to the fast step (E_2) are in a narrow area and have much higher
300 values, yielding a mean activation energy of $299.2 \text{ kJ} \cdot \text{mol}^{-1}$.

301 Recently, Sawai et al. (2013) studied antigorite dehydration using in situ
302 high-temperature infrared microspectroscopy under isothermal conditions and found
303 that the activation energies were $219 \text{ kJ} \cdot \text{mol}^{-1}$ for bond 1, $243 \text{ kJ} \cdot \text{mol}^{-1}$ for bond
304 2, and $256 \text{ kJ} \cdot \text{mol}^{-1}$ for bond 3. These authors concluded that antigorite
305 dehydration occurred via one-dimensional diffusion. Gualtieri et al. (2012) conducted
306 in situ X-ray powder diffraction to study antigorite dehydration kinetics and combined
307 the results with transmission electron microscopy observations. These authors showed

308 that the apparent activation energy of antigorite dehydration within a temperature
309 range of 885-981 K was $255 \text{ kJ} \cdot \text{mol}^{-1}$. Weber and Greer (1965) obtained an
310 activation energy of 443-643 $\text{kJ} \cdot \text{mol}^{-1}$ using the Arrhenius equation and the
311 relationship between the rate of reaction and the concentration. Our average activation
312 energies ($268.1 \text{ kJ} \cdot \text{mol}^{-1}$ and $299.2 \text{ kJ} \cdot \text{mol}^{-1}$) obtained from 2-DAEM are higher
313 than those reported by both Sawai et al. (2013) and Gualtieri et al. (2012) but lower
314 than those ($443\text{-}643 \text{ kJ} \cdot \text{mol}^{-1}$) reported by Weber and Greer (1965). In addition, the
315 activation energy from Weber and Greer (1965) is higher than the maximum value of
316 $400 \text{ kJ} \cdot \text{mol}^{-1}$ obtained from the FWO method in this study. In our model, the
317 pre-exponential factor may explain that the different activation energies as $\log(A)$ are
318 linearly correlated with the activation energy (**Fig. 5**). The pre-exponential factor may
319 cause substantial changes to the absolute value of the activation energy.

320

321 **Geophysical implications**

322 In many subduction zones, double seismic zones (DSZ) where earthquakes
323 occur at different depths have been revealed. Intermediate-depth earthquakes occur
324 along two dipping planes, that is, an upper plane and a lower plane (Hasegawa et al.
325 1978). The hypothesis of a close link between earthquakes and dehydration events at
326 intermediate depths has been proposed by many researchers (Chollet et al. 2011;
327 Hilairet et al. 2007; Omori et al. 2004; Reynard 2013; Sawai et al. 2013) and is
328 strongly supported by recent experimental results (Ferrand et al. 2017). The upper
329 plane is likely related to the dehydration of metabasalts (e.g. Incel et al. 2017); the

330 seismicity of the lower plane in the subducting serpentinized mantle would be because
331 of a dehydration-driven stress transfer from dehydrating antigorite along faults to
332 fresh peridotite volumes (Ferrand et al. 2017). Whether peridotite embrittlement
333 occurs or not depends on the competition between the dehydration kinetics and the
334 solid matrix deformational rate. Dehydration kinetics are important parameters, and if
335 they are faster than the rate of solid matrix deformation, seismic ruptures nucleate.

336 The fluid production rate is important, as a high rate may locally induce fluid
337 overpressure along faults, which could explain the observed interplane seismicity
338 (Kita and Katsumata 2015; Kita et al. 2010). The fluid production rate during
339 antigorite dehydration is calculated from kinetic data. To determine the fluid
340 production rate, we consider an isothermal model. Six parameters can be applied in
341 the isothermal model to predict the dehydration rate of antigorite. In equation (7), we
342 substitute βdt for dT ; thus, the isothermal model can be written as:

$$\begin{aligned} \alpha = 1 - \int_0^{+\infty} \exp\left(-\int_0^t A \exp\left(\frac{-E}{RT}\right) dt\right) \\ \times \left(\frac{\omega}{\sigma_1 \sqrt{2\pi}} \exp\left(\frac{-(E - E_1)^2}{2\sigma_1^2}\right) \right. \\ \left. + \frac{1 - \omega}{\sigma_2 \sqrt{2\pi}} \exp\left(\frac{-(E - E_2)^2}{2\sigma_2^2}\right)\right) dE \end{aligned} \quad (12)$$

343 We used the six parameters to compute the reaction progress (α) at different
344 temperatures according to **Eq. 12**. The results are shown in **Fig. 8**. The antigorite
345 dehydration rate increases with increasing temperature. In addition, at higher
346 temperatures, the rate scarcely varies.

347 Dehydration rates ($V_{1/2}$) can be obtained from the half-life time of the reaction
348 $t_{1/2}$ when half of the antigorite has disappeared (Chollet et al. 2011).

$$V_{1/2} = \frac{C_{H_2O}}{t_{1/2}} \times \frac{\rho}{\rho_{H_2O}} \quad (13)$$

349 where $t_{1/2}$ is the half-life time of the reaction when α is equal to 0.5, C_{H_2O} is the
350 water content of the mineral, ρ is its density 2.62 g/cm^3 (Bezacier et al. 2010) and
351 ρ_{H_2O} is the density of water. The reaction progress (α) can be calculated with time t
352 at a given temperature T . The half-life time $t_{1/2}$ can be obtained from the isothermal
353 model as shown in **Fig. 8**. The fluid production rate $V_{1/2}$ is 8.0×10^{-5} , $1.9 \times$
354 10^{-4} , 4.4×10^{-4} , 9.8×10^{-3} , and $2.1 \times 10^{-3} m_{fluid}^3 m_{rock}^{-3} s^{-1}$ for the temperatures
355 of 893 K, 913 K, 933 K, 953 K, and 973 K, respectively, near the onset temperature of
356 the isothermal dehydration reaction. This study strongly supports rapid dehydration,
357 which is a key parameter for dehydration-induced seismicity. Sudden stress transfers
358 are likely to occur in the vicinity of dehydrating serpentized faults, as recently
359 demonstrated in laboratory analogs (Ferrand et al. 2017). These rates are faster than
360 the viscous relaxation of antigorite (3×10^{-7} to $3 \times 10^{-12} s^{-1}$) reported by Hilairet
361 et al. (2007), indicating that antigorite dehydration leads to increased strain
362 localization that is likely to induce mechanical instabilities in serpentized peridotites.
363 Previous studies also have shown similar results. Chollet et al. (2011) measured the
364 kinetics of antigorite dehydration in situ under high-temperature and high-pressure
365 conditions using X-ray diffraction in a closed system and determined that the release
366 of aqueous fluid occurs at $10^{-4} m_{fluid}^3 m_{rock}^{-3} s^{-1}$. Perrillat et al. (2005) found that the
367 discharge rate is on the order of 10^{-6} to $10^{-8} m_{fluid}^3 m_{rock}^{-3} s^{-1}$ using a real-time

368 X-ray diffraction study. Previous studies have indicated that pressure has no
369 significant effect on kinetics (Perrillat et al. 2005; Rubie and Thompson 1985);
370 therefore, our results under a condition of ambient pressure can be extrapolated to
371 subduction zones. By combining previous results (Chollet et al. 2011; Hilairet et al.
372 2007; Perrillat et al. 2005; Sawai et al. 2013) and those of this study, we found that
373 antigorite dehydration rates are higher than the characteristic relaxation, which is
374 calculated as the inverse of the Maxwell relaxation time, and that brittle failure is
375 likely to occur within dehydrating serpentinized peridotites. Therefore, as described
376 by the dehydration-driven stress transfer model (Ferrand et al. 2017), antigorite
377 dehydration may trigger seismicity in the lower plane of the double seismic zone.

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608 **Table 1 Kinetic triplet and errors obtained using an n -th order reaction**

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SSA (s^{-1})	A (s^{-1})	n	E ($\text{kJ} \cdot \text{mol}^{-1}$)	Δ
$10^8 \sim 10^{12}$	$(3.44 \pm 0.17) \times 10^9$	2.2 ± 0.1	212.2 ± 10.6	0.22

$10^8 \sim 10^{13}$	$(6.41 \pm 0.32) \times 10^9$	2.2 ± 0.1	217.3 ± 10.9	0.23
$10^8 \sim 10^{14}$	$(4.54 \pm 0.23) \times 10^9$	2.2 ± 0.1	215.6 ± 10.8	0.24

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SSA is the search space of A (SSA). A , n and E are three constant parameters termed the kinetic triplet in the n -order model, which are determined from **Eq. 4** and **Eq. 6**. Δ is an objective function, which was defined to describe the error between experimental and fitting data. The lower the Δ , the more accurate the kinetic triplet.

Table 2 Kinetic parameters and errors obtained using 2-DAEM

SSA	A	E_1	σ_1	E_2	σ_2	ω	Δ
(s^{-1})	(s^{-1})	$(kJ \cdot mol^{-1})$	$(kJ \cdot mol^{-1})$	$(kJ \cdot mol^{-1})$	$(kJ \cdot mol^{-1})$		

$10^8 \sim 10^{12}$	$(1.58 \pm 0.16) \times 10^{11}$	227.0±4.5	28.0±5.6	251.4±5.0	3.1±0.2	0.336±0.017	0.040
$10^8 \sim 10^{13}$	$(2.08 \pm 0.21) \times 10^{12}$	249.7±5.0	34.6±6.9	272.2±5.4	4.8±0.2	0.336±0.017	0.033
$10^8 \sim 10^{14}$	$(4.66 \pm 0.47) \times 10^{13}$	268.1±5.4	29.8±6.0	299.2±6.0	8.9±0.5	0.336±0.017	0.031
$10^8 \sim 10^{15}$	$(4.42 \pm 0.44) \times 10^{14}$	298.2±6.0	37.6±7.5	312.3±6.2	10.4±0.5	0.336±0.017	0.030
$10^8 \sim 10^{16}$	$(3.97 \pm 0.40) \times 10^{15}$	307.1±6.1	44.4±8.9	330.9±6.6	8.6±0.4	0.336±0.017	0.029

649 SSA is the search space of A . A , ω , E_i , and σ_i are the parameters determined from
 650 **Eq. 7**. Δ is an objective function to determine the accuracy of the six parameters.
 651 The lower the Δ , the more accurate the six parameters.

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684 **Figure captions**

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687 **Figure 1 Reaction progress curves for antigorite with different heating rates.**

688

689 The red, green, blue and violet lines represent the reaction progress with heating rates
690 of 10, 15, 20 and 25 $\text{K} \cdot \text{min}^{-1}$, respectively. The numbers are the heating rates.

691

692

693 **Figure 2 Activation energy of antigorite as a function of the reaction progress**
694 **determined using the FWO method.**

695

696 The FWO method is a model-free method for calculation of activation energy in TG
697 (Eq. 5). FWO is Flynn-Wall-Ozawa.

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699

700 **Figure 3 Experimental and theoretical reaction progress curves of the**
701 **non-isothermal experiments for antigorite derived from the n -th order reaction.**

702

703 a, b, c and d represent the fitting results of 10, 15, 20, and 25 $\text{K} \cdot \text{min}^{-1}$ based on the
704 n -th order model, respectively. The red line and blue line represent the calculated data
705 and experimental data, respectively, of α as a function of temperature, with
706 $A = 3.44 \times 10^9 \text{ s}^{-1}$; $n = 2.2$ and $E = 212.2 \text{ kJ} \cdot \text{mol}^{-1}$.

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708

709 **Figure 4 Structure of the antigorite crystal lattice, showing different local**
710 **environments for OH groups.**

711

712 The red, blue, white and yellow balls represent oxygen, silicon, hydrogen and
713 magnesium atoms, respectively. The structure of the antigorite crystal lattice (Capitani
714 and Mellini 2004) contains two different local environments for OH groups, that is,
715 O-H1 and O-H2. The O-H1 and O-H2 bonds have different lengths. The O-H3 bond is
716 at the boundary between the O-H1 and O-H2 bonds. The lengths of O-H1 and O-H2
717 bonds were determined using first-principle calculations.

718

719 **Figure 5 A compensation effect between the pre-exponential factor (A) and the**
720 **average activation energy based on 2-DAEM.**

721

722 Logarithm of pre-exponential factor (A) as a function of activation energy (E_1 and E_2)
723 for antigorite dehydration. The red and blue squares represent data points for E_1 and
724 E_2

725 , respectively. The regression equations are $\log(A) = 0.0527 \times E_1 - 0.7634$ and
726 $\log(A) = 0.0559 \times E_2 - 2.9165$. The correlation coefficients between $\log(A)$ and the
727 average activation energies (E_1 and E_2) are 0.9906 and 0.9986, respectively.

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729

730 **Figure 6 Comparisons between the calculated data and experimental data of**
731 **antigorite according to 2-DAEM.**

732

733 a, b, c and d represent the fitting results of 10, 15, 20, and 25 K · min⁻¹ based on
734 2-DAEM, respectively. The red and blue lines represent the calculated and
735 experimental data, respectively, of α with temperature. α was calculated according
736 to **Eq. 7** at a given A , E_i , σ_i and ω are listed in the third row in Table 2 when Δ
737 has the minimum value of 0.014. The correlation coefficients for the heating rates of
738 10, 15, 20, and 25 Kmin⁻¹ are, respectively, 0.9996, 0.9995, 0.9994 and 0.9991.

739

740 **Figure 7 Distributions of the activation energies in 2-DAEM.**

741

742 The green, blue and red lines denote $G_1(E)$, $G_2(E)$ and $G(E)$, respectively, where

743 $G_1(E) = \frac{\omega}{\sigma_1\sqrt{2\pi}} \exp\left(\frac{-(E-E_1)^2}{2\sigma_1^2}\right)$ is the distribution of the activation energy under local

744 environments for the O-H1 bond and $G_2(E) = \frac{1-\omega}{\sigma_2\sqrt{2\pi}} \exp\left(\frac{-(E-E_2)^2}{2\sigma_2^2}\right)$ is the

745 distribution of the activation energy under local environments for the O-H2 bond.

746 $G(E)$ is the sum of $G_1(E)$ and $G_2(E)$.

747

748 **Figure 8 Isothermal dehydration kinetic models for antigorite at different**
749 **temperatures derived from the parameters of the non-isothermal experiments.**

750

751 These curves show the calculated reaction progress at certain temperatures according
752 to **Eq. 12**. The numbers are temperatures. α was calculated according to **Eq. 12** at a
753 given temperature with parameters A , E_i and σ_i chosen from the third row of Table 2.

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Fig. 1

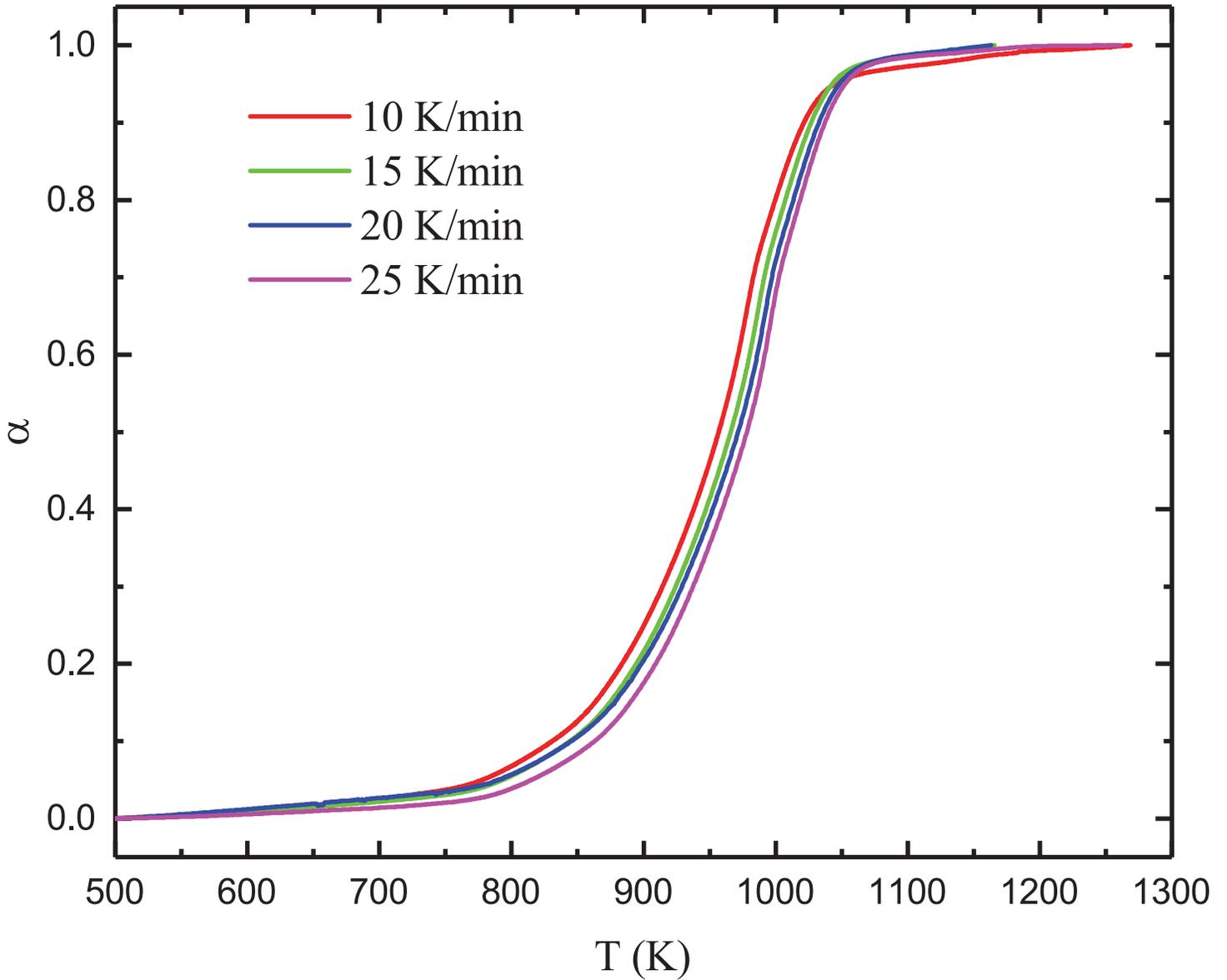


Fig. 2

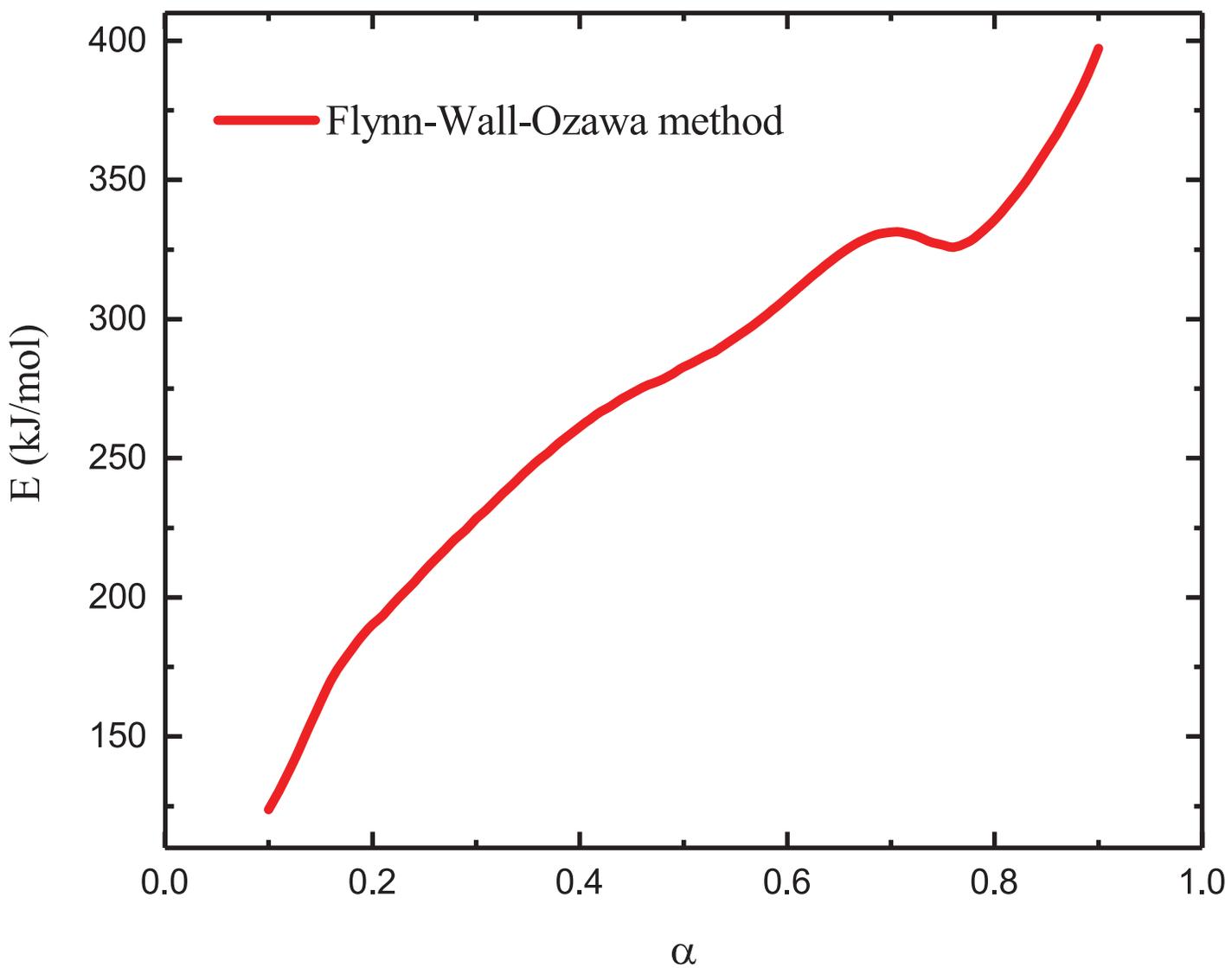


Fig. 3

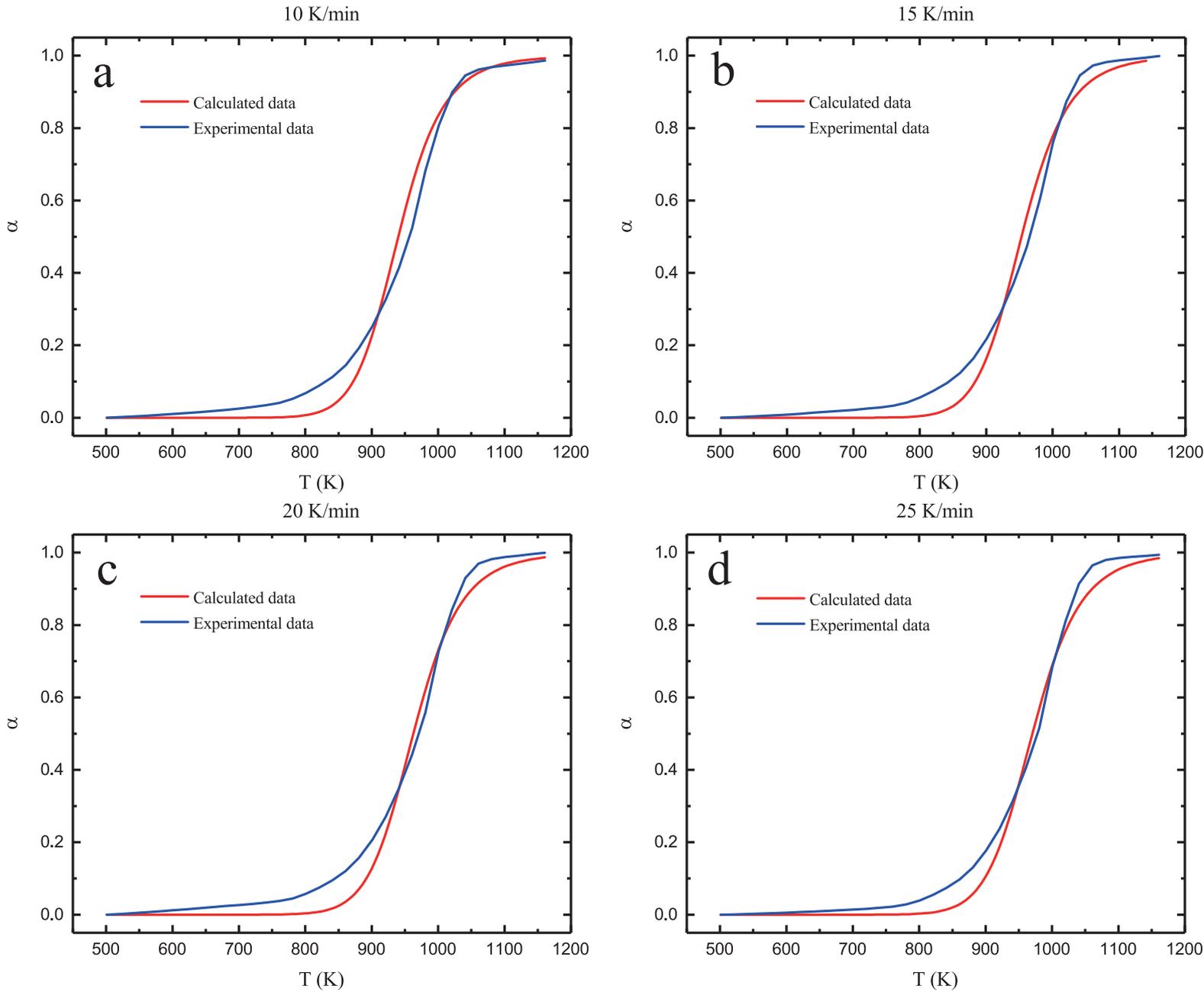


Fig. 4

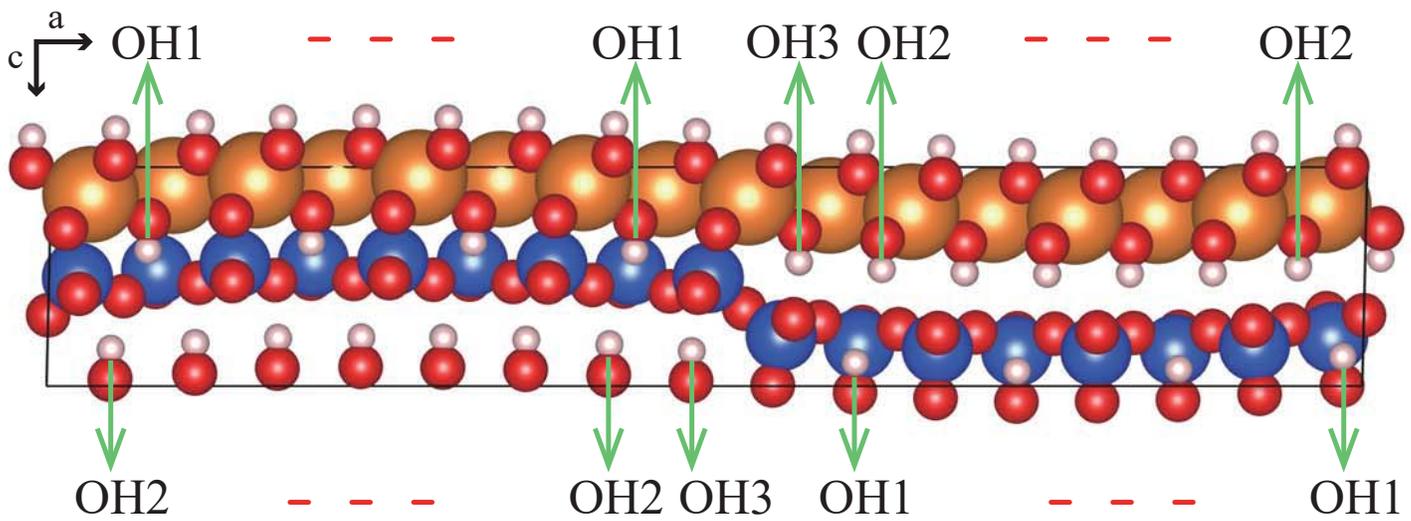


Fig. 5

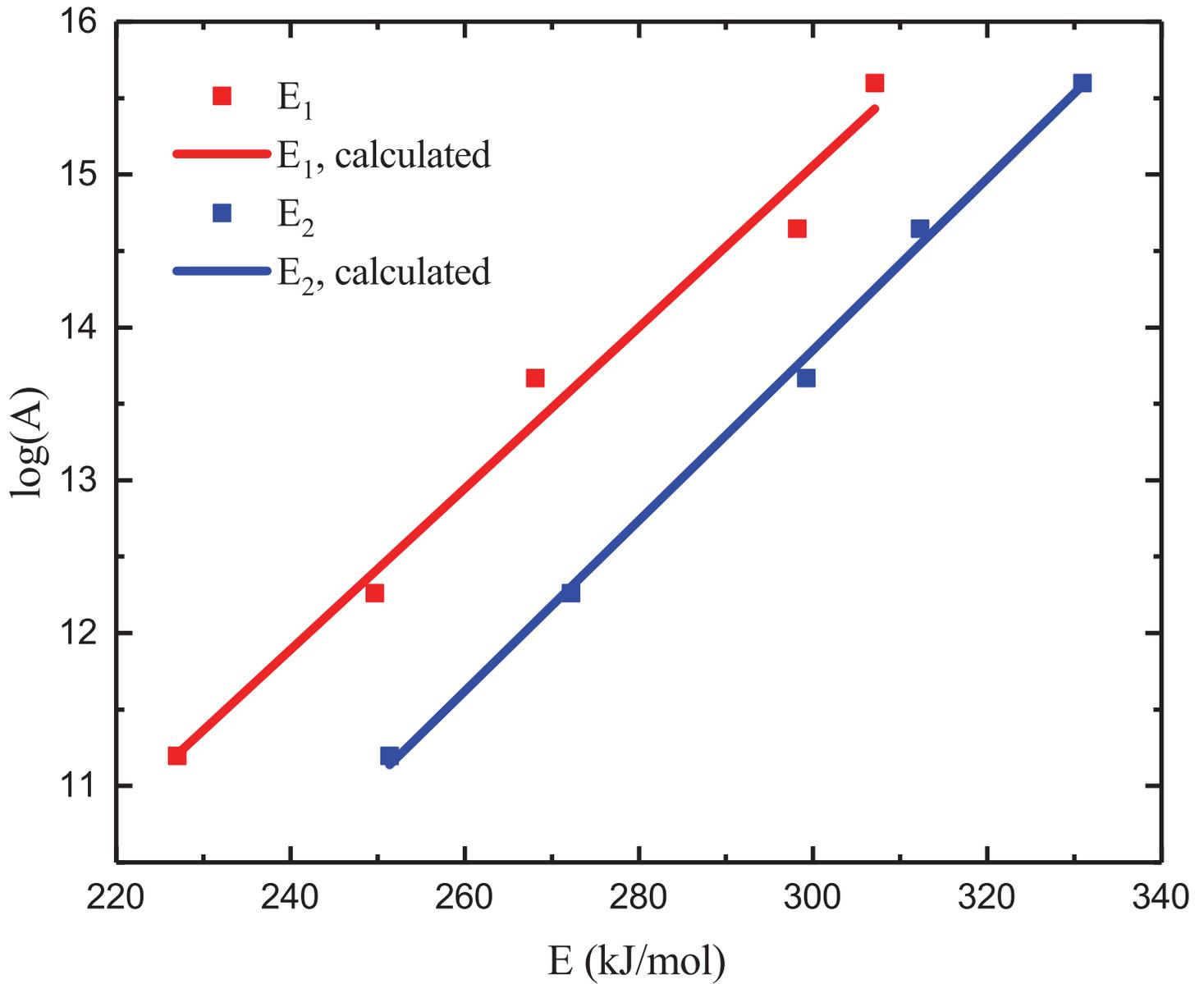


Fig. 6

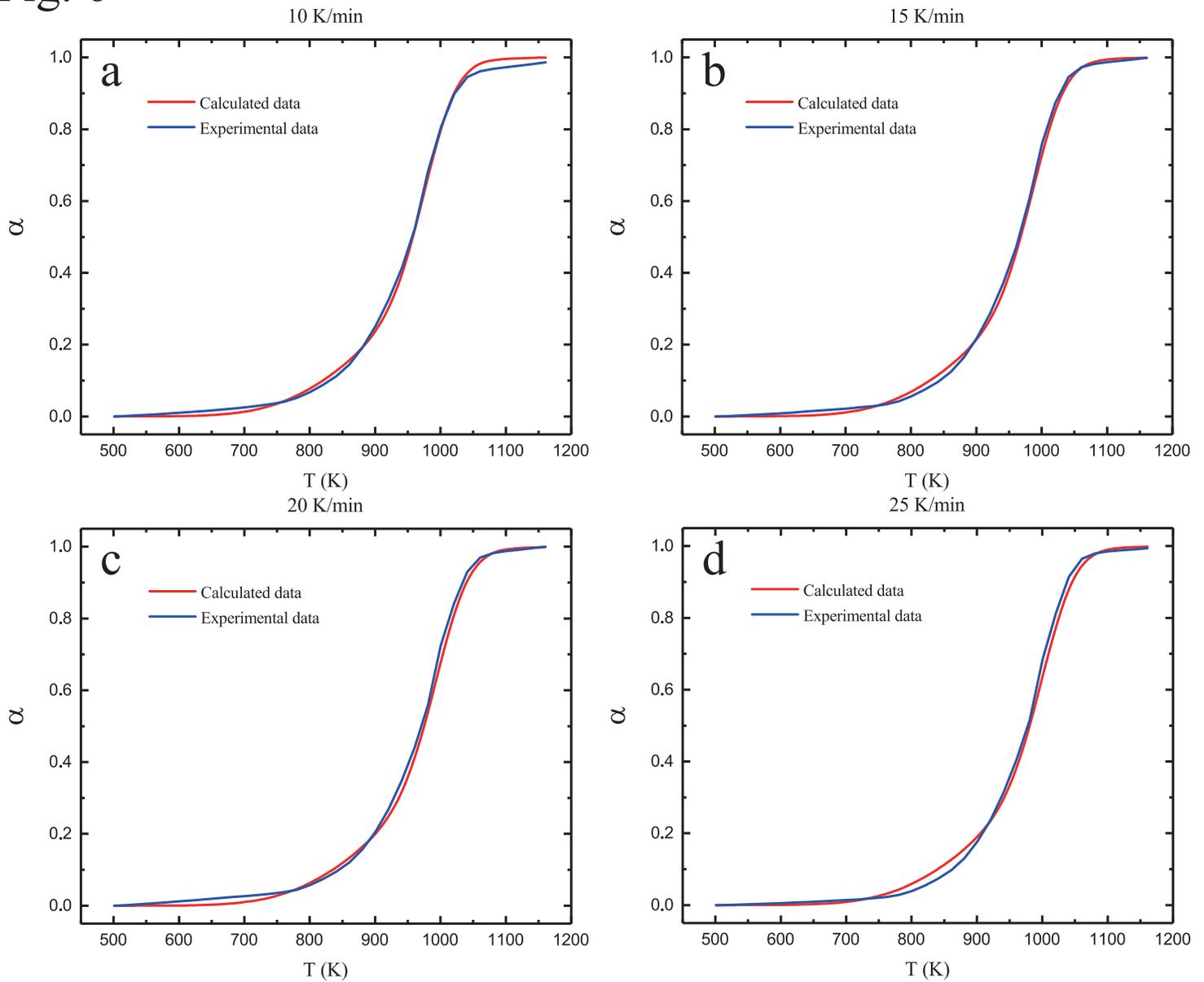


Fig. 7

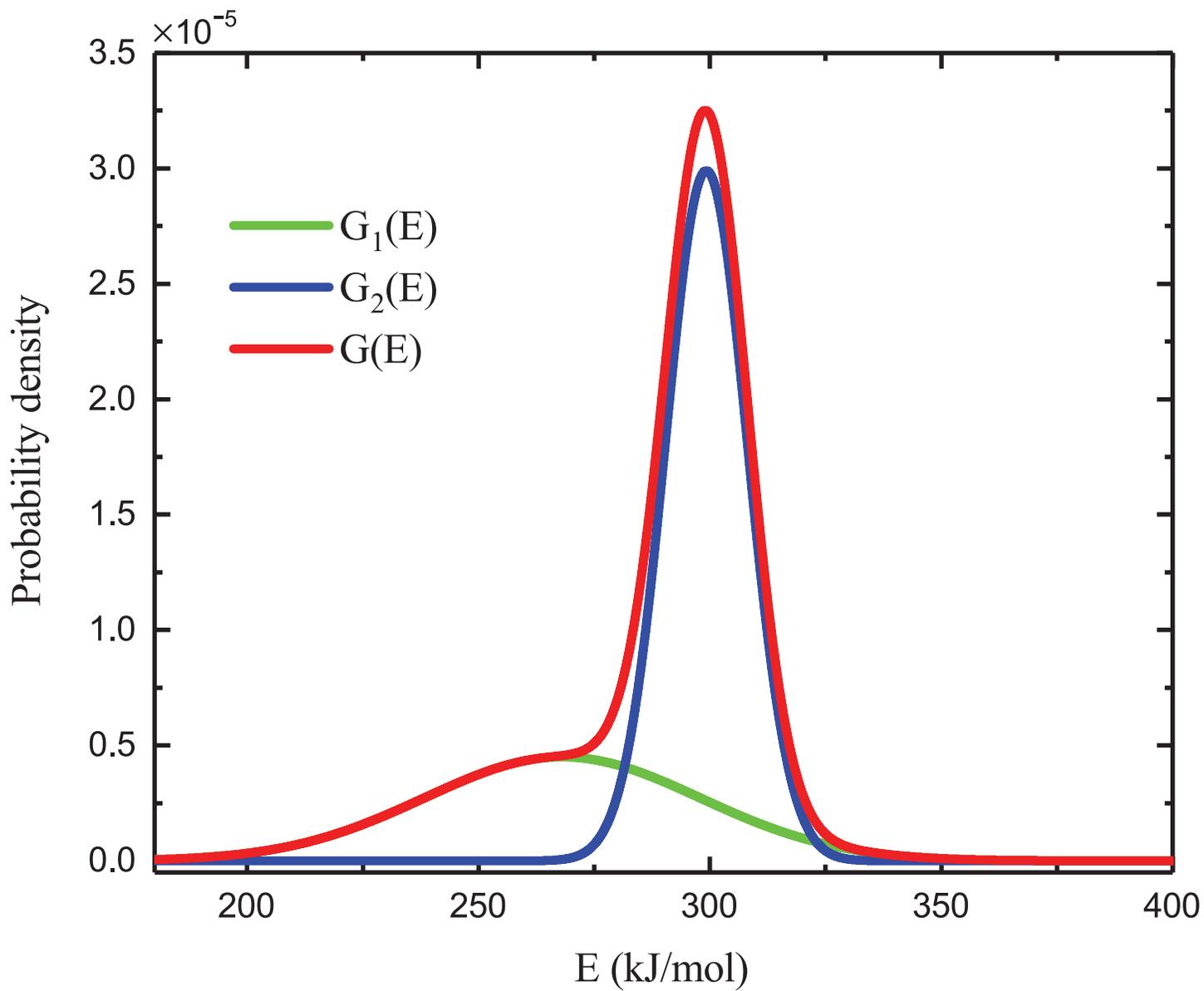


Fig. 8

