1	Changes in Physical Properties of 4C Pyrrhotite (Fe ₇ S ₈) across the 32 K Besnus Transition
2	Revision 1
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12	

13 Abstract

14 Pyrrhotite Fe₇S₈ is a common sulfide mineral in the Earth's crust and mantle, as well as in a 15 range of meteorites, and is of interest to a wide variety of disciplines, including economic geology, 16 geophysics, and material science. The 4C variety of pyrrhotite shows a dramatic change in 17 magnetic properties at $T \approx 30$ K, known as the Besnus transition. Although this transition is 18 frequently used to detect pyrrhotite in geologic samples, the underlying mechanism driving the 19 transition has not yet been identified. This study presents a high-resolution view of the changes in 20 heat capacity, magnetic, and electronic properties of a natural single crystal of nearly pure, 21 monoclinic 4C pyrrhotite across the Besnus transition. Contrary to previous studies, all of these 22 properties show clear evidence of the Besnus transition, specific heat in particular revealing a clear 23 transition at 32 K, apparently of second order nature. Small-angle neutron scattering data are also 24 presented, demonstrating an unusual change in short-range magnetic scattering at the transition. *Corresponding author: Michael Volk, Institute for Rock Magnetism, Department of Earth

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25 Furthermore, a magnetic field dependence of the transition temperature can be seen in both induced 26 magnetization and electrical resistivity. These new observations help narrow the possible nature 27 of the phase transition, clearly showing that interactions between intergrown coexisting 4C and 28 5C* superstructures, as suggested in some literature, are not necessary for the Besnus transition. 29 In fact, the changes seen here in both the specific heat and the electronic transport properties are 30 considerably larger than those seen in samples with intergrown superstructures. To further 31 constrain the mechanism underlying the Besnus transition, we identify five separate potential 32 models and evaluate them within the context of existing observations, thereby proposing 33 experimental approaches that may help resolve ongoing ambiguities.

Keywords: Besnus transition, Phase transition, magnetism, heat capacity, resistivity, Sulfide,
Pyrrhotite,

36 1 INTRODUCTION

37 Iron mono-sulfides, or pyrrhotites, (Fe_{1-x}S, $0 \le x \le 0.125$), are commonly occurring iron sulfide 38 minerals that are important to a wide range of scientific disciplines. In the Earth and planetary 39 sciences, 4C pyrrhotite (Fe₇ S_8) acts as a recognizable ferrimagnetic phase, and is found in many 40 magmatic and metamorphic rocks (Dunlop and Özdemir, 1997), in claystones (Aubourg and Pozzi, 41 2010), and in limestones (*Muttoni*, 1995). In terrestrial rocks, the remanent magnetization carried 42 by this phase has been used to study polarity reversals of the Earth's magnetic field (e.g., 43 *Ouidelleur et al.*, 1992), to examine the nature of magnetization in guartzite clasts from the 44 Precambrian Jack Hills of Western Australia (Weiss et al., 2015), and as the basis for a crustal geo-45 thermometer (Aubourg and Pozzi, 2010) or barometer (Gilder et al., 2011). Furthermore, 4C 46 pyrrhotite is a major magnetic constituent of several classes of meteorites, most notably in Martian 47 meteorites (*Rochette et al.*, 2001). Since pyrrhotite is thermodynamically stable on the surface of

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48 Mars, it may also contribute to the Martian magnetic anomalies (Fegley Jr. et al., 1995). Moreover, 49 4C pyrrhotite can be found in chondritic meteorites, the most primitive and oldest class of 50 meteorites (Herndon et al., 1975, Zhang et al., 2008), and, as such, may possibly record the 51 magnetic fields in the early solar system, as well as transient fields associated with impact events 52 (Muxworthy et al., 2017). In general, pyrrhotites are also the most common iron sulfide minerals 53 in certain intrusions, such as the 1.1 Ga Duluth Complex, which hosts economically important 54 quantities of Cu, Ni and Pt group elements. Understanding the oxidation kinetics of pyrrhotite 55 minerals in mine tailings is of critical importance for avoiding water quality degradation due to 56 acid mine drainage (Lapakko, 2015). Further, pyrrhotite commonly occurs as inclusions in 57 diamonds, and may provide information about sulfur cycling within the lower mantle (*Sharp*, 58 1966, Gilder et al. 2011). Within the material science community, pyrrhotites have been studied 59 as potential candidate materials for anodes in lithium-ion batteries (Zhang et al., 2015), and can 60 be used to sequester arsenic contamination in the environment (*Cantu et al.*, 2016).

61 Given its importance to geoscience and material science, proper characterization of pyrrhotite 62 within natural and synthetic samples is critical, yet frustratingly difficult. As discussed in more 63 detail below, iron monosulfides can display a wide range of crystallographic symmetries and 64 superstructures based on their composition and growth history, and frequently occur as finely 65 intergrown polycrystalline assemblages. Thus, reflected light and electron microscopy methods 66 are often combined with electron and neutron diffraction to properly determine the structure of 67 small iron monosulfide grains. However, magnetic methods are one of the few techniques that 68 allow researchers to characterize the presence of pyrrhotite in bulk samples. In particular, 69 observations of the low-temperature magnetic transition at approximately 30 K (the "Besnus" 70 transition) of 4C pyrrhotite have been used with increasing frequency as a diagnostic indicator for

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the presence of pyrrhotite (*Dunlop and Özdemir, 1997*). Low temperature superconducting
quantum interference device (SQUID) magnetometers are typically sensitive enough to detect 4C
pyrrhotite in natural samples at parts per million concentrations (*Rochette et al.* 1990).

74 Despite the utility of low temperature magnetic methods in detecting the presence of pyrrhotite 75 in natural and synthetic samples, there is remarkably little consensus on the mechanism that 76 underlies this dramatic change in magnetic properties. The aim of this study is to provide a high-77 resolution view of the physical (heat capacity and resistivity) and magnetic properties (remanence 78 and high- and low-field induced magnetization) across the Besnus transition in a phase-pure 79 specimen of 4C pyrrhotite. Small angle neutron scattering is also presented to help determine the 80 length scales over which the changes in magnetic properties take place. This new data provides 81 the necessary evidence to differentiate between competing models that seek to explain this 82 phenomenon, and will enable researchers to more accurately interpret its magnetic recording on 83 Earth and other planetary bodies.

84

1.1 SUPERSTRUCTURES

85 Pyrrhotite occurs in many polytypes that are often intergrown or show incommensurate 86 chemical composition. Similar to troilite (FeS), these polytypes crystallize in a NiAs-like structure 87 with alternating filled and vacant Fe layers oriented parallel to the *ab* plane containing variably 88 ordered vacancies (Morimoto et al., 1970). Within each Fe-layer the spins are coupled 89 ferromagnetically. However, adjacent layers are coupled antiferromagnetically via superexchange 90 through the S atoms, which in turn forms distorted S octahedra around the Fe atoms (Levinson and 91 *Treves*, 1968). Ordering of the vacancies gives rise to superstructures, which are referred to using 92 an NC notation, where N gives the number of stacked NiAs cells along the crystallographic c-axis. 93 Four commensurate hexagonal varieties, FeS (2C), Fe₉S₁₀ (5C), Fe₁₀S₁₁ (11C) and Fe₁₁S₁₂ (6C),

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and one monoclinic $Fe_7S_8(4C)$ are stable at ambient conditions (*Nakazawa and Morimoto*, 1970; *Morimoto et al.*, 1970). The ordering of the vacancies in hexagonal polytypes creates a zero net magnetization, and thus hexagonal pyrrhotites are non-ferro/ferri-magnetic in the bulk, and incapable of retaining a spontaneous magnetization.

98 On the other hand, ordering of vacancies in the most Fe deficient variety of pyrrhotite, Fe₇S₈ 99 (4C), creates a small lattice distortion ($\beta \neq 90$), which lowers the mineral's symmetry to monoclinic 100 (Powell et al. 2004), and is often referred to as 'pseudo-hexagonal'. Furthermore, the ordered 101 vacancies in 4C pyrrhotite produce a small net magnetization, which gives rise to ferrimagnetism 102 up to its Néel temperature of 598 K (Powell et al., 2004). Within the 4C cell, the layers can be 103 labeled as *FAFBFCFD*, where F denotes a filled Fe layer and A-D are four distinct vacancy layers. 104 Each layer with vacancies contains four unique Fe sites (Bertaut, 1953; Tokonami et al., 1972). Of 105 these Fe sites, two have no vacancies above or below, while the Fe atoms in the filled (F) layers 106 are adjacent to either one or two vacancies (*Ericsson et al.*, 1994). Superexchange coupling 107 through the S atoms is stronger than the direct Fe-Fe coupling (Levinson and Treves, 1968). 108 Consequently, sites that do not have vacancies above or below should be largely unaffected by 109 changes in vacancy ordering (Levinson and Treves, 1968).

110 **1.2**

1.2 PHASE TRANSITION

111 At $T_{\text{Bes}} \approx 30$ K, 4C pyrrhotite undergoes a magnetic phase transition first described by *Besnus* 112 *and Meyer* (1964). The Besnus transition is commonly used to identify pyrrhotite in rocks 113 (*Rochette et al.*, 2011). Similar to the Morin transition in hematite (Fe₂O₃) or the Verwey transition 114 in magnetite (Fe₃O₄), the transition is marked by a loss of remanent magnetization on cooling, and 115 a drastic change in fundamental magnetic properties (*Fillion and Rochette*, 1988; *Dekkers et al.*, 116 1989; *Rochette et al.*, 1990, Dunlop and Özdemir, 1997). Furthermore, the anisotropy constants This is a preprint, the final version is subject to change, of the American Mineralogist (MSA) Cite as Authors (Year) Title. American Mineralogist, in press. DOI: https://doi.org/10.2138/am-2018-6514 Preprint submitted to American Mineralogist February 15, 2018

and the direction of the magnetic easy axis change during cooling (*Mikami et al.*, 1959; *Bin and Pauthenet*, 1963; *Sato et al.*, 1964; *Rochette et al.*, 1990). At room temperature (T_R) the easy axis lies within the basal (*a-b*) plane of the crystal but this progressively rotates out of the basal plane towards the *c* axis upon cooling below 205 K (*Bin and Pauthenet*, 1963, Powell et al. 2004). While a self-reversal of the magnetization has been observed in very large pyrrhotite crystals (*Fillion and Rochette*, 1988), the direction of the remanence vector at room temperature, remains constant upon

123 cycling through T_{Bes} (*Feinberg et al.*, 2015).

124 While the Besnus transition has been known to our community for more than 50 years, the 125 mechanism driving the transition is not well understood. Two prevailing explanations exist. 126 Wolfers et al. (2011) measured neutron diffraction and magnetic torque on twinned single crystals. 127 The magnetic torque shows a six-fold symmetry above $T_{\text{Bes.}}$ which changes to a four-fold 128 symmetry below the transition. Based on the change in symmetry, *Wolfers et al.* (2011) propose a 129 crystallographic transition from the monoclinic F2/d ($T > T_{\text{Bes}}$) to a triclinic F-1 ($T < T_{\text{Bes}}$) structure. 130 This mechanism is similar to the change from cubic to monoclinic symmetry that occurs within 131 magnetite upon cooling through the Verwey transition at $T \approx 110$ K. The second explanation was 132 proposed by Charilaou et al. (2015) and later expanded upon by Koulialias et al. (2016, 2018). 133 Here, the change in magnetic properties is explained by changing magnetic interactions between 134 two different coexisting superstructures (4C and 5C*). The proposed 5C* phase in *Charilaou et* 135 al. (2015) possesses a different vacancy ordering, but the same overall chemical composition as 136 the 4C structure. Their arguments against a crystallographic phase transition as the basis of the 137 Besnus transition are based in large part on the absence of a transition in the heat capacity $C_{\rm p}(T)$, 138 and the presence of a second inflection in hysteresis loops at low temperatures ($T \le 200$ K). Clearly, 139 this explanation relies on the existence of two distinct coexisting phases of $Fe_{1-x}S$.

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140 In a previous study, we measured major hysteresis loops in 72 directions within the basal plane 141 of an oriented single crystal of phase-pure 4C pyrrhotite (Volk et al., 2016). The chemical 142 composition and crystallographic orientation of the crystal were determined by Energy-dispersive 143 X-ray spectroscopy (EDX) and Electron backscatter diffraction (EBSD). The experiments were 144 conducted at 21 temperatures spanning the Besnus transition (20-50 K). Several magnetic 145 properties showed a six-fold symmetry above T_{Bes} , including the second hysteresis inflection 146 described by Koulialias et al. (2016). However, we found that these magnetic symmetries change 147 from six-fold ($T > T_{Bes}$) to four-fold ($T < T_{Bes}$), consistent with the behavior originally observed 148 by Wolfers et al., (2011). Furthermore, the second hysteresis inflection disappeared below $T_{\text{Bes.}}$ To 149 better differentiate between the competing models, we present here a detailed study of the 150 structural, magnetic and electronic properties of a second Fe₇S₈ single crystal subsampled from 151 the same piece used in Volk et al. (2016) from Morro Velho, Brazil, provided by the Munich 152 Mineralogical State Collection (Mineralogische Staatssammlung München).

153 **RESULTS AND DISCUSSION** 2

154 The chemical composition of the sample was determined by electron microprobe analysis with 155 a JEOL JXA-8900R operated with a beam energy of 20 keV, a current of 20 nA and diameter of 156 5 micron. Detection limits ranged from 0.020 weight percent for Fe K α to 0.022 weight percent 157 for S K α and analytical sensitivity (at the 99% confidence level) ranged from 0.157 percent relative for Fe K α to 0.406 percent relative for S K α . Elements were acquired using analyzing crystals 158 159 LiFH for Fe-ka, and PETJ for S-ka. The standards were Fe for Fe Ka, and Pyrite, FeS₂ for S Ka. 160 The specimen's bulk chemical composition, determined from 12 measurements, is Fe_{6.78±0.12}S₈ 161 (Fe_{0.85}S, Fe = 45.9 at. %). According to the phase diagram of *Nakazawa and Morimoto* (1971), the 162 sample's measured composition is within in the stability field of 4C pyrrhotite. Significantly, no

163 impurities, such as Ni, were detected in any of the 12 measurements (within the limits stated164 above).

165 In naturally occurring rocks, pyrrhotites are often an intergrown mixture of both the hexagonal 166 and monoclinic phases (Arnold, 1967). In our case, it is particularly important to detect the 167 presence of any such hexagonal intergrowths, since one of the proposed mechanisms for the 168 Besnus transition requires the presence of both 4C and 5C* superstructures (*Charilaou et al.*, 2015; 169 Koulialias et al., 2016). Non-magnetic hexagonal phases transform into a ferrimagnetic 170 phasebetween 200°C-265°C depending on their iron content (Schwarz and Vaughan, 1972). This 171 transformation, due to a rearrangement of vacancies, is identified as a sudden rise in magnetic 172 susceptibility, known as the λ -transition (*Schwarz and Vaughan*, 1972). We measured magnetic 173 susceptibility as a function of temperature using a Geophysica Kappabridge KLY-2 (300 Am⁻¹ 174 field at 920 Hz), which is able to detect even small hexagonal impurities. The specimen shows a 175 single Néel temperature (318°C, 591 K) close to the literature value (320°C (Dunlop and Özdemir, 176 (1997)) with no sign of a lambda transition, as shown in Figure S1. This conclusion of the absence 177 of any significant amount of 5C* phase is supported by structure determinations, also, as discussed 178 below.

179 2.1 POWDER X-RAY DIFFRACTION

To detect possible inclusions (*e.g.*, pyrite FeS₂) in our pyrrhotite sample, or pyrrhotite polytypes (*e.g.*, the 5C* phase discussed above), room temperature powder X-ray diffraction (PXRD) data were acquired with a PANalytical X'Pert PRO X-ray diffractometer equipped with a Co source (0.17909 nm) and an X'Celerator detector. The diffraction patterns were collected from 10° - 90° scattering angle in 0.0167° steps with an effective dwell time of 100 s per step.

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185 Figure 1 shows the PXRD pattern along with the result of Rietveld refinement with a single 186 phase with a 4C stacking sequence. The software Rietica was used to fit the data using the 187 structural model by *Powell et al.* (2004). The fitted cell parameters of the monoclinic space group 188 C2/c are a = 1.1915(1) nm, b = 0.68698(6) nm, c = 1.2907(1) nm and $\beta = 118.005(1)^{\circ}$, with quality 189 parameters of $R_{bragg} = 2.87$ %, $R_P = 2.24$ % and $R_{WP} = 3.51$ %. The lattice parameters agree well 190 with the values for the 4C phase obtained by Charilaou et al. (2015). Furthermore, the composition 191 obtained by PXRD refinement is $Fe_{6.76}S_8$ ($Fe_{0.83}S$), in excellent agreement with the measured 192 chemical composition from electron microprobe analysis. The PXRD results thus indicate that the 193 crystal is pure monoclinic 4C pyrrhotite without any detectable contamination from other minerals 194 (e.g., oxides) or vacancy superstructures. It is important to note that the hexagonal $(5C^*)$ phase, 195 which is clearly present in the PXRD data presented by Charilaou et al. (2015) (see Fig. 1 inset), 196 is not present here, within the limits of detection.

197

2.2 SMALL ANGLE NEUTRON SCATTERING

198 Small-angle neutron scattering (SANS) from a 4C pyrrhotite single crystal was used to 199 investigate potential structural and/or magnetic changes across the Besnus transition. This 200 powerful method (Muhlbauer et al., 2018), which has not yet been applied to pyrrhotite, provides 201 length-scale-dependent information on structural and magnetic inhomogeneities across a wide 202 range. In essence, the neutron scattering cross-section, $d\Sigma/d\Omega$, is measured as a function of the 203 scattering wavevector, O, revealing information on a length scale $2\pi/O$. Importantly, in the 204 unpolarized case used here, the inhomogeneities that generate the scattering can be structural (e.g., 205 long-range defects, twins, twin boundaries) or magnetic (e.g., domains and domain walls, short-206 range magnetic fluctuations, etc.). The SANS measurements presented were taken on the NG7 30 207 m SANS beamline at the NIST Center for Neutron Research, at 6 temperatures between 5 and 60

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208	K, i.e., around the Besnus transition. At each temperature, 2D area scans, or Q_x - Q_y intensity maps
209	(see Fig. 2a), were collected in two frames (one at 0.004 Å ⁻¹ < Q < 0.025 Å ⁻¹ and one at 0.02 Å ⁻¹
210	< Q < 0.15 Å ⁻¹), resulting in a total covered Q range from 0.004 to 0.15 Å ⁻¹ . We thus probe real-
211	space length scales $(2\pi/Q)$ from 4 to 160 nm. In such scans the neutron beam is perpendicular to
212	the plane of the detector, the dark region in the center of Fig. 2a is a beam-stop covering the
213	unscattered beam, and displacement outwards from the origin corresponds to increasing Q.

214 Fig. 2 shows the SANS results obtained with a sister specimen to the main crystal studied in 215 this work, oriented with the *c*-axis approximately parallel to the neutron beam. Panel (a) shows an 216 example 2D area scan (i.e., a Q_x - Q_y intensity map), at 5 K, out to Q_x/Q_y values of 0.02 Å⁻¹, i.e., 217 probing real space length scales above 30 nm. The six-fold in-plane symmetry is obvious in the 218 2D scan, and, due to the use of unpolarized neutrons, could arise from structural effects (such as 219 long-range (> 30 nm) defects aligned with high-symmetry in-plane directions) or from magnetic 220 domain formation along specific axes (note that the data were recorded here after nominal zero 221 field cooling). Of high interest in terms of the aforementioned potential explanations for the Besnus 222 transition, the observed six-fold pattern did not alter significantly across the 32 K transition 223 temperature. The angular scan in Fig. 2b, which simply shows the cross-section summed over the small range 0.005 Å⁻¹ < O < 0.009 Å⁻¹ (to improve statistics), was in fact averaged over all 224 225 temperatures between 5 and 60 K, no significant changes in the angular dependence occurring at 226 any T. Six-fold periodicity is thus clear, but, remarkably, it is not T-dependent across the Besnus 227 transition. If a crystal symmetry change occurs at 32 K, this indicates that the long-range (> 30 228 nm) structural defects or magnetic domains producing the scattering seen here do not change 229 significantly in response.

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230 There are, however, some subtle changes in the absolute scattering intensity. To illustrate this, 231 Fig. 2c plots $d\Sigma/d\Omega$ vs. the magnitude of O, acquired by circular averaging of 2D maps of the type 232 shown in Fig. 2(a); the example shown here is at 45 K. As is often the case, $d\Sigma/d\Omega(Q)$ can be fit 233 (solid line through the points) as the sum of two terms: A Porod term (blue dashed line) and a 234 Lorentzian term (red dashed line). The generalized Porod term has the form $d\Sigma/d\Omega = (d\Sigma/d\Omega)_P/O^n$, 235 where the exponent *n* provides details on the morphology of the scattering inhomogeneities, and 236 the prefactor $(d\Sigma/d\Omega)_{\rm P}$ parameterizes the strength of the Porod intensity. In our case n = 3.3 (i.e., 237 somewhat less than 4), as shown in Fig. 2c, indicating that the scattering inhomogeneities (long-238 range structural defects or magnetic domains) do not have completely smooth surfaces. A value of 239 3.3, however, not far from the generic 4.0 for smooth 3D objects, indicates that the scattering is 240 not occurring from clearly one- or two-dimensional objects, such as filaments or lamellae. The Lorentzian term has the form $d\Sigma/d\Omega = (d\Sigma/d\Omega)_L/(Q^2 + \kappa^2)$, where the prefactor $(d\Sigma/d\Omega)_L$ 241 parameterizes the strength of the Lorentzian intensity, and κ is a constant. Such scattering is most 242 commonly associated with short-range (nm-scale) magnetic fluctuations, where κ becomes $1/\xi$, 243 244 the magnetic correlation length. Standard behavior in a ferrimagnet would involve an order-245 parameter-like turn on of the Porod scattering from magnetic domains below the magnetic ordering 246 temperature, along with a peak in Lorentzian scattering at the ordering temperature due to critical 247 scattering. At $T \le 60$ K, far below the ordering temperature, Lorentzian scattering should thus be 248 small, and, if present at all, be decreasing on cooling. Not only does Fig. 2c show that the 249 Lorentzian scattering here is non-negligible, but Fig. 2d shows that it has an unusual T dependence. 250 Specifically, $(d\Sigma/d\Omega)_{\rm L}$ (right axis, red points) decreases on cooling from 60 to 35 K, as might be 251 expected, but then grows on further cooling to 5 K. This is unusual in a long-range-ordered ferro-252 or ferri-magnet far below its ordering temperature. The Besnus transition thus manifests itself in

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253	some change in short-range (< 10 nm) spin fluctuations/ordering in 4C pyrrhotite, but not
254	obviously so in the angular dependence of long-range (> 30 nm) scattering. We note that the T
255	dependence of $(d\Sigma/d\Omega)_P$ (left axis, black points in Fig. 2d) may provide some evidence of a plateau
256	around the Besnus transition, although much better statistics would be required to reinforce this.
257	Future SANS measurements to probe these findings in more detail could employ a much wider T
258	range (going above the Néel temperature), or polarized neutrons, to better separate magnetic and
259	structural scattering. Both of these approaches are certainly possible.

260 2.3 HEAT CAPACITY

261 As discussed above in the Introduction, the existence of an anomaly in specific heat capacity 262 $(C_{\rm P})$ at the Besnus transition is important in pyrrhotite, but is not clear from prior experimental 263 work. Specifically, the work of *Grønvold et al.* (1959) indicated a small anomaly around 32 K, 264 while the work on material with mixed 4C/5C* superstructures by Charilaou et al. (2015) revealed 265 no such anomaly. In general, first-order phase transitions involve latent heat and result in 266 discontinuities in $C_{P}(T)$, while second-order transitions involve no latent heat and result in peaks 267 in $C_{\rm P}(T)$. To probe these issues in our 4C pyrrhotite single crystal, $C_{\rm P}(T)$ was measured in a 268 Physical Property Measurement System (PPMS) from Quantum Design Inc., from 1.8 K to room 269 temperature. Relaxation calorimetry was employed, using temperature pulses amounting to 2 % of 270 the sample temperature. The crystal was attached to the measurement platform using Apiezon N 271 grease, and the heat capacity of the platform and grease were subsequently subtracted. For all 272 measurements the ratio of sample to addenda heat capacity was maintained at greater than 2.0, 273 safely above the recommended minimum of 0.5 from *Lashlev et al.* (2003). Similarly, the thermal 274 coupling parameter between the sample and calorimeter was > 96%, well above the minimum 275 Quantum Design recommendation of 90%.

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276 Fig. 3 shows our $C_{\rm P}(T)$ results along with those obtained from *Grønvold et al.* (1959), 277 Kobayashi et al. (1999), and Charilaou et al. (2015). Fig. 3a shows $C_{\rm P}(T)$ over the entire T range 278 studied (1.8 - 300 K), generally showing good agreement with previous studies over most of the T 279 range. Around the Besnus transition, however (inset to Fig. 3a), the results are substantially 280 different in the various studies. In our work, $C_P(T)$ exhibits a large anomaly at 31.7 K, larger than 281 the one seen by Grønvold et al. (1959) (in a synthetic sample), and in contrast to the lack of any 282 anomaly seen by Charilaou et al. (2015). In the latter work, multiple stacking sequences (i.e., 283 coexisting 4C and 5C* phases) were clearly observed in XRD. Here, however, we have a phase-284 pure 4C pyrrhotite crystal, which apparently results in a clear, well-defined transition in $C_{\rm P}(T)$, 285 taking the form of a classic " λ anomaly" expected at a second-order phase transition. (Note that 286 this is not to be confused with the 220 °C "λ-transition" in hexagonal pyrrhotite, as discussed 287 above). The apparent second-order nature of the transition is supported by the absence of any 288 detectable thermal hysteresis, both in relaxation calorimetry on cooling and warming, and by the 289 use of long-pulse measurements (Gillard et al. 2015). The striking differences between our phase-290 pure 4C natural single crystal and the 4C/5C* natural crystal of *Charilaou et al.* (2015), may well 291 result from the phase-purity with respect to the stacking sequences, and/or trace concentrations of 292 other metals (particularly nickel). It is interesting to note that the 5C polymorph measured later by 293 Grønvold et al. (1991) (not shown in Fig. 3) also shows no $C_{\rm P}(T)$ anomaly, suggesting that the 294 pure 4C structure may be the only form of pyrrhotite that displays a heat capacity anomaly at the 295 Besnus transition, and that any changes in properties across the Besnus transition are intrinsic to 296 4C pyrrhotite.

Subtracting a linear *T*-dependent background from our $C_P(T)$ data around the Besnus transition and then integrating the excess specific heat yields the entropy change across the Besnus transition. This gives 2.39 J mol⁻¹ K⁻¹, or, by mole of Fe, 0.34 J mol⁻¹ K⁻¹. For comparison, the entropy change across the Verwey transition of magnetite, a first-order structural transition, is approximately 5 J mol⁻¹ K⁻¹, or by mole of Fe, 1.33 J mol⁻¹ K⁻¹ (*Shepherd et al.* 1991). Furthermore, the Morin transition of hematite, a first-order spin-flop transition, has an entropy change of 0.43 J mol⁻¹ K⁻¹, or, per mole of Fe, 0.22 J mol⁻¹ K⁻¹ (*Pastor et al.* 2012). The entropy change across the Besnus transition in 4C pyrrhotite is thus significant.

305 Further analysis was performed on the low T specific heat (Fig. 3b), which can be used to 306 extract electronic and lattice contributions, in turn allowing for determination of the density-of-307 states at the Fermi level and the Debye temperature, respectively. (As discussed later in this work, 308 pyrrhotite is an electronic conductor). As shown in Fig. 3b, our $C_P(T)$ in this region is in excellent agreement with that of Kobayashi et al. (1999), but again rather different to the mixed phase 309 4C/5C* crystal of Charilaou et al. (2015). Our data were fit following the standard procedure, 310 using $C_{\rm P}(T) = \gamma T + \beta T^3 + \alpha T^5$, where the first term is the electronic contribution and the latter two 311 are from lattice dynamics. The T^3 term is the standard low-temperature expansion in the Debye 312 model, while the T^5 term is the next order term, as often needed to adequately describe 313 314 experimental data. Two fitting procedures were employed, the first considering all data up to 10 K, comparable to that used by Kobayashi et al. (1999). This resulted in a poor fit, with an 315 unphysically large αT^5 term. The second fitting approach acknowledges the clear upturn in the data 316 at $T^2 < 30 \text{ K}^2$, and fits only at $T^2 > 30 \text{ K}^2$, resulting in the good fit shown in Fig. 3b. Considering 317 the electronic contribution first, the results give $\gamma = 95$ mJ mol⁻¹ K⁻², a substantial value. 318 319 Theoretical values from the density-of-states at the Fermi level calculated from density functional theory suggest ~46.5 mJ mol⁻¹ K⁻², however (Shirai et al., 1996; Shimada, et al., 1998), 320 corresponding to only a modest effective mass enhancement of ~2, in good agreement with 321

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322 Shimada et al. (1998) and Kobayashi et al. (1999). With respect to lattice contributions, the 323 extracted β was used to determine the Debye temperature, Θ_D , yielding 278 K. This is in reasonable 324 agreement with the value of *Kobayashi et al.* (1999) ($\Theta_D = 313$ K), who, we note, did not include 325 the higher order αT^5 term used here. We add that in our case the αT^5 term is indeed almost an order 326 of magnitude smaller than the βT^3 term, even at the highest T, as would be expected. Finally, with 327 respect to the excess specific heat at $T^2 < 30 \text{ K}^2$ (i.e., below 5.5 K, as also seen in *Kobayashi et al.* 328 (1999)) we note only that upturns on $C_P(T)$ vs. T^2 plots at the lowest T are not at all uncommon in 329 solid-state systems, and can derive from Schottky anomalies, for example (Rosenberg, (1963); 330 Gopal, (1966)). The latter occur due to any mechanism that produces a manifold of low-lying 331 states, and can result from all manner of origins; further study at temperatures well below 1 K 332 could clarify this.

333 2.4 MAGNETIC PROPERTIES

334 2.4.1 REMANENT AND INDUCED MAGNETIZATION

335 In paleomagnetism, the Besnus transition is associated with a loss in remanent magnetization 336 on cooling. We measured the temperature dependence of a room temperature saturation isothermal 337 magnetization (RTSIRM) with a Quantum Design MPMS-XL (Magnetic Property Measurement System) on cooling and warming. The specimen acquired a RTSIRM in a 2.5 T magnetic field at 338 339 300 K applied along two arbitrarily chosen directions in the basal plane of the crystal (Φ_1, Φ_2). 340 both roughly perpendicular to the [001] c-axis. The temperature dependence of the RTSIRM was 341 measured twice, firstly spanning the full temperature range from 300-10 K (5K/min) and secondly 342 in high resolution (0.5 K/min) spanning the transition (20-50 K).

343 The cooling path (Fig. 4 insets) of the RTSIRM ($\Phi_1 = 1.5, \Phi_2 = 1.9 \text{ Am}^2/\text{kg}$) shows a gradual 344 decrease in magnetization to ≈ 45 % at 110-150 K, where it plateaus. Characteristic of the Besnus

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345 transition (Dekkers et al., 1989), the remanence drops sharply in the narrow temperature range of 346 35-25 K (fig. 4). Estimates of the transition from the maximum dM/dT, yield a slightly higher 347 temperature (34 K) than the peak in specific heat (31.7 K). While the shapes of the cooling paths 348 in Φ_1 and Φ_2 are comparable, the warming paths (path 4 in Fig. 4a,b)) are quite different in the 349 vicinity of T_{Bes} , when measured in high-resolution (0.5 K steps). The recovery of magnetization 350 by warming back through the Besnus transition is strongly dependent on the grain-size of the 351 sample and should be low for large grains (Dekkers et al., 1989). Along Φ_1 (Fig. 4a) a small portion 352 (\approx 5%) of the RTSIRM is recovered on warming from 10 K back to 40 K. In contrast, the 353 measurement along direction Φ_2 (Fig. 4b) shows a zig-zag behavior at T_{Bes} , with a local minimum 354 centered exactly at the specific heat peak. Furthermore, Φ_2 shows little remanence recovery. 355 Finally, the remaining remanence decays further with warming and only 20 ± 4 % of the initial 356 RTSIRM remains at room temperature.

357 Dekkers et al. (1989) showed that cycling to 4.2 K changes the median destructive field (MDF) 358 of the sample. Similarly, the second high resolution cooling/warming cycle (path 3.4) shows that 359 the newly acquired RTSIRM(50 K) is slightly larger than the first one. This indicates that the 360 changes occurring across the Besnus transition may not be entirely reversible. Furthermore, the 361 changes in MDF are grain-size-dependent (Dekkers et al., 1989), with the biggest increases 362 occurring in larger particles. Thus, the configuration of domain walls in large, multi-domain grains 363 may be altered after cycling though the Besnus transition. A possible rearrangement of vacancies 364 at the transition could change the defect structure of the crystal and lead to different pinning of 365 domain walls and consequently a hardening of MDF and increased remanence ratio (Table 1).

The application of DC magnetic fields can modify the precise temperature of some magnetic transitions, or even influence crystallographic changes during transitions in magnetic materials.

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368	For example, in magnetite the orientation of the monoclinic <i>c</i> -axis below the Verwey transition
369	can be strongly controlled by the application of a DC magnetic field during cooling, where the
370	monoclinic c -axis is selected from the cubic a -axis that is closest to the orientation of the applied
371	field (Li, 1932, Kasama et al., 2010). To look for field-dependent effects at the Besnus transition,
372	we performed in-field cooling (FC) (5 K/min) in 0.5, 1, 1.5, 2 and 3 T (B1[001]) followed by zero-
373	field warming (5 K/min) of the low temperature remanence (LTSIRM) that was acquired in the
374	respective field at 10 K (Fig. 5). The saturation magnetization determined from field cooling
375	$(M(300K, 3T) = 20.2 \text{ Am}^2/\text{kg}$, see table 1) is well within the range of reported values (18.3-21.2)
376	Am ² /kg) (O'Reilly et al., 2000; Kind et al., 2013). As the sample cools in a field (Fig. 5a), the
377	magnetization increases to a local maximum at $T \approx 180$ K ($M(180$ K, 3 T) = 21.7 Am ² /kg). The
378	maximum is located close to where the magnetic anisotropy constant K_4 is also at a maximum and
379	where K_3 changes sign from positive to negative (<i>Bin and Pauthenet</i> , 1963).

380 Cooling in fields lower than ≈ 1.5 T causes a sharp increase of magnetization at the Besnus 381 transition. By comparison, in strong fields (B > 1.5 T), the Besnus transition is only visible as a 382 slight bump in the magnetization, which has been reported in several studies (*Besnus and Mever*, 383 1964; Dekkers et al., 1989; Kind et al., 2013; Charilaou et al., 2015; Koulialias et al., 2016). The 384 changing crystalline anisotropy causes the spins, which are lying in the basal plane at room 385 temperature, to rotate out of the plane with decreasing temperature (*Bin and Pauthenet*, 1963; 386 Powell et al. 2004). Thus, the increase in magnetization could be related to a change in spin 387 configuration at the transition.

We can compare the directional dependence of the magnetization with data from magnetic hysteresis loops (*Volk et al.*, 2016), which were measured as a function of orientation (Φ) and temperature on a single crystal sister specimen from the same sample studied here. Consequently,

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391 each magnetization value can be expressed as a function of field, orientation and temperature M 392 (B, T, Φ) and used to determine quasi "FC". Figure 5b shows the magnetization in a 500 mT field 393 in different crystallographic orientations as a function of temperature. Similar to the MPMS 394 measurements along Φ_1 and Φ_2 (Fig. 5a), the quasi FC data shows orientation-dependent 395 magnetization. However, these measurements show a clear directional dependence of the shape 396 and magnitude of the Besnus transition. Moreover, the standard deviation of the induced moment 397 is about 2.6 times greater for T = 50 K than for T = 20 K. The decrease in variability at lower 398 temperature suggests changes in magneto-crystalline anisotropy as the crystal is cooled through 399 the transition.

To check for possible thermal hysteresis, we repeated the FC experiment (500 mT, Φ_2) at a higher resolution (0.5 K). While hysteretic behavior is typical for first-order transitions, no such behavior is expected for second-order ones, as noted above. The fact that both cooling and warming curves are perfectly aligned, completely reversible, and show no hysteresis (inset in Fig. 5b), further underlines the second-order character of the Besnus transition, consistent with our specific heat data.

The remanence acquired in the field-cooling experiments at 10 K (LTSIRM, Fig. 5c) was measured on warming back to room temperature in zero field. The absolute magnetization is the same for all fields used, which drops by 90 % at T_{Bes} and further decays to only \approx 2% at room temperature. While the LTSIRM is about 7 times stronger than the RTSIRM (tab. 1) at low temperature, the two are comparable after a full cooling-warming circle. There was no field dependence observed in the warming of LTSIRM for pure 4C pyrrhotite.

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413 The $M(B,T,\Phi)$ data (Fig. 5b) from the hysteresis loops, as described earlier, allow for the 414 extraction of T_{Bes} as a function of Φ (in the basal plane) and applied field. We define $T_{\text{Bes}}(B, \Phi)$ as 415 the maximum change in magnetization $(\max(dM/dT))$ on cooling. $T_{\text{Bes}}(B, \Phi)$ is determined from 416 a cubic spline interpolation of the smoothed data. Figure 6a shows that the mean $T_{\text{Bes}}(\Phi)$ is field 417 dependent, and increases from ~ 30 to ~ 34 K as applied field increases from 0 to 1 T. The most 418 prominent change is visible in the first 100 mT (see inset in Fig. 6a), where T_{Bes} changes by $\approx 2 \text{ K}$ 419 from 30 K to 32 K. As the field is further increased, T_{Bes} rises linearly until it saturates at 750 mT 420 and 34 K. In our previous study (Volk et al., 2016), we were not able to explain how the coercivity 421 of remanence (B_{cr}) of the single crystal could be lower than its coercivity (B_{c}) for temperatures 422 < 30 K. However, the field dependence of the transition can explain this phenomenon. When $B_{\rm cr}$ 423 is measured in zero field, the crystal is above T_{Bes} and consequently the coercivity of remanence 424 is low. $B_{\rm c}$, on the other hand, is measured *in* a magnetic field, which raises $T_{\rm Bes}$. The measurement 425 is therefore done below the transition temperature. B_c is greatly enhanced for T < T_{Bes} which leads 426 to the unusual $B_{cr} < B_c$. Furthermore, the mean $T_{Bes}(B)$ shows a directional dependence (Fig. 6b). 427 Parallel to the crystallographic *a*-axes (determined using EBSD (see *Volk et al.*, 2016)), we find 428 the maximum T_{Bes} , while minima lie between axes, coinciding with orientations whose hysteresis 429 loops most strongly show a second inflection. Finally, combining the field and directional 430 dependence (Fig. 6c) shows the minimum (dark color) of T_{Bes} (30.1 K) located close to the 431 projection of the *c*-axis onto the basal plane with a transition temperature close to the peak in heat 432 capacity. Thus, the highest Besnus transition temperatures occur when a field is applied parallel to 433 one of the *a*-axis directions, while the lowest T_{Bes} occurs in between these directions and closest 434 to the projection of the *c*-axis. It appears as though the application of DC magnetic fields in This is a preprint, the final version is subject to change, of the American Mineralogist (MSA) Cite as Authors (Year) Title. American Mineralogist, in press. DOI: https://doi.org/10.2138/am-2018-6514 Preprint submitted to American Mineralogist February 15, 2018

between the *a*-axes, or even potentially along the c-axis, stabilizes the $T > T_{\text{Bes}}$ phase, requiring lower temperatures to ultimately bring about the Besnus transition.

437 2.4.2 ALTERNATING CURRENT SUSCEPTIBILITY

Measurements of low-field AC susceptibility as a function of temperature can provide 438 439 important information about the mobility of domain walls as a mineral's intrinsic magnetic 440 anisotropy changes, as the pinning energies of vacancies change, or as a mineral undergoes a 441 crystallographic transition. Low field in-phase (χ') and out of phase (χ'') AC susceptibility 442 between 10 K and 300 K (see table 2) was collected at 5 field amplitudes (0.02 - 0.3 mT, 5K/min) 443 on cooling and subsequent warming along the sample's [001] axis (χ_{\parallel}) as well as perpendicular to 444 it $(\chi_{\perp} \Phi_2)$. The AC susceptibility (Fig. 7) in the basal plane (χ_{\perp}) shows a gradual decline during 445 cooling until it drops sharply at \approx 34 K. There is a strong amplitude dependence in both the inphase (Fig. 7a) and the out-of-phase components (Fig. 7b) of χ_{\perp} . While the amplitude dependence 446 447 of χ_{\perp} is approximately constant down to 35 K, it increases dramatically at T_{Bes} and then drops to 448 close to zero below the transition. The measurement of χ_{\parallel} , on the other hand, shows a vastly 449 different behavior. Here, both χ_{\parallel} and its amplitude dependence decrease with increasing 450 temperature. The maximum of $\chi_{\parallel}(17-24 \times \chi_{\parallel}(T_{R}))$ is reached close to the Besnus transition. Similar 451 to $\chi_{\perp}, \chi_{\parallel}$ evidences a sharp change at $T_{\text{Bes.}}$ While the measurements in this study come from a 452 multidomain single crystal, Kind et al. (2013) measured similar changes in susceptibility for a 453 natural Fe₇S₈ powder. They interpret the decrease in susceptibility during cooling across the 454 Besnus transition as a reduction in domain wall mobility at lower temperatures, an explanation 455 which would only apply to predominantly MD samples. *Bezaeva et al.*, (2016) also reported a drop 456 in susceptibility across the Besnus transition for predominantly single-domain 4C pyrrhotite,

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457 which suggests instead that something about the intrinsic magnetic anisotropy energy of the 458 material has changed below $T_{\text{Bes.}}$

459 To gain a better understanding of how susceptibility changes at the transition, we also 460 measured the frequency dependence of AC susceptibility (χ_{\perp} and χ_{\parallel}) at high resolution (0.5 K 461 steps) for 7 frequencies (1 - 1000 Hz) spanning T_{Bes} (from 20-40 K). There is no frequency dependence in the in-phase data (Fig. 6a) for either χ_{\perp} or χ_{\parallel} . There is a slight frequency dependence 462 463 in the out-of-phase data (Fig. 6b), however, the main drop of χ_{\parallel} occurring at temperatures ≈ 1 K 464 warmer than χ_{\perp} in both in-phase and out-of-phase measurements. This shift is reproducible and is 465 likely not an instrumental artifact. Furthermore, close inspection of the data of Koulialias et al. 466 (2016) reveals a similar shift, which seems to decrease with increasing frequency.

467 A prominent secondary peak is visible at ≈ 32 K (see inset in Fig. 7a) in χ_{\perp} ' and less expressed 468 in χ_{\perp} ", while χ_{\parallel} does not show it at all. Similar features were reported by *Koulialias et al.* (2016) 469 in their natural single crystal and Kind et al. (2013) in Fe₇S₈ powder. To resolve this feature, data 470 must be collected at sufficiently high resolution (here, 0.5 K steps). Thus, the feature is not visible 471 in the lower resolution amplitude dependence measurements (2.5 K steps). It is noteworthy that 472 this secondary peak is located directly at the temperature of the specific heat peak discussed above. 473 This indicates that the peak is directly related to the transition. Consequently, a Besnus transition 474 temperature, determined from the decay of susceptibility may be overestimating $T_{\text{Bes.}}$

The strong anisotropy of susceptibility in 4C pyrrhotite is reflected in the vastly different susceptibilities that are measured parallel and perpendicular to [001]. At room temperature, χ_{\perp} is ~170 times larger than χ_{\parallel} . As temperatures approach T_{Bes} , however, χ_{\parallel} increases while χ_{\perp} decreases, which results in a linear decrease of their ratio ($\chi_{\perp}/\chi_{\parallel}$) when plotted logarithmically (Fig. 8). This Preprint submitted to American Mineralogist February 15, 2018

479	linearity is broken at T_{Bes} , where the ratio shows a discontinuity, again well-aligned with the
480	specific heat peak (see inset Fig. 7). We interpret this change a consequence of the rotation of the
481	easy axis of the crystal during cooling. Powell et al. (2004) showed that as Fe ₇ S ₈ is cooled, the
482	easy axis rotates out of the (001) plane, consistent with increasing susceptibility in [001]. Similarly,
483	the sudden increase in $\chi_{\perp} \chi_{\parallel}$ could be caused by an abrupt rotation of the easy axis towards the
484	basal plane and back out, towards [001].

Finally, below T_{Bes} , χ_{\perp} decreases to <1% of its room temperature value (Table 2); χ_{\parallel} on the other hand is only reduced to about 50%. Consequently, below the transition both susceptibilities are comparable ($\chi_{\perp} \approx 2 \times \chi_{\parallel}$). Thus, the anisotropy of the crystal is greatly reduced below T_{Bes} . The sample in study has a platelet shape, and therefore shape anisotropy may play a role. If one were to correct for the shape of our sample, the sense of the correction would lower the susceptibility within the basal plane, and therefore it is possible that susceptibility parallel to [001] could be even closer to, if not larger than, χ_{\perp} .

492 2.4

2.5 ELECTRICAL RESISTIVITY

To gain an understanding of the electronic property changes that take place across the Besnus transition, we performed temperature- and magnetic-field-dependent measurements of resistivity (ρ) from 4 to 300 K, using a home-built Janis cryostat with a 9 T superconducting magnet. Silver paste contacts were employed in a 4-wire van der Pauw configuration (*van der Pauw*, 1958), using a Lakeshore 370 13.7 Hz AC bridge to measure 4-terminal resistances. Checks for Ohmicity, contact resistance, and self-heating were made at 4 and 300 K.

499 As shown in Fig. 9a, the ρ in the basal plane of our 4C pyrrhotite single crystal exhibits clearly

500 metallic behavior $(d\rho/dT$ is metallic-like (i.e., positive) over almost the entire temperature range,

501 with finite apparent conductivity as $T \rightarrow 0$), with a large anomaly visible at 32 K, where the

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502 resistivity drops significantly. We note that the peak in the ratio of $\chi_{\parallel}/\chi_{\perp}$, this resistivity anomaly, 503 and the peak in specific heat are all located at the same temperature. This highlights the 504 reproducibility and precision of T_{Bes} in our measurements, in addition to its clear, recognizable 505 effect on diverse properties (magnetic, electronic, and thermodynamic). In contrast to the Verwey 506 transition in magnetite, which displays a sharp drop in resistivity on warming (e.g., Kuipers and 507 Brabers, 1979), the resistivity change here is continuous, again consistent with second-order 508 character of the Besnus transition (there is also no thermal hysteresis in $\rho(T)$ on heating/cooling). 509 The room temperature resistivity $(\rho_{300 \text{ K}})$ of the crystal measured here is 0.32 m Ω cm, which 510 compares well with the $\rho_{300 \text{ K}}$ of 0.41 m Ω cm originally obtained by *Besnus and Meyer* (1964). 511 being considerably lower than the 2.1 m Ω cm obtained by *Charilaou et al.* (2015). This 512 consistency with the original work, and the lower values than those obtained by *Charilaou et al.* 513 (2015), again highlight the quality of this phase-pure natural single crystal, and reinforce the clear 514 expression of the Besnus transition in this phase-pure sample. Consistent with this, the temperature 515 dependence of ρ is also stronger in this crystal. For example, the residual resistivity ratio (RRR), 516 defined as RRR = $\rho_{300 \text{ K}}/\rho_{5 \text{ K}}$, is 4.04 in this work, compared to ~2.9 in Besnus and Meyer (1964), 517 and 1.34 in Charilaou et al. (2015). This indicates not only lower defect densities in our phase-518 pure crystal, but also a stronger decrease in $\rho(T)$ below $T_{\text{Bes.}}$ With a RRR of above 4, and also 519 given the absolute magnitudes of the resistivities, this pyrrhotite crystal can be considered a "bad" 520 or "dirty" metal or semimetal. In terms of constraining the *origin* of the Besnus transition, we note 521 that any explanation for it must be able to explain a substantial *decrease* in resistivity on cooling. 522 Fig. 9b shows the magnetic field dependence of ρ vs. T around the Besnus transition, with the 523 field applied along a random direction in the basal plane. As the field increases from 0 T to 0.5 T,

524 progressive smearing of the sharp anomaly in ρ vs. T and a weak negative magnetoresistance (MR)

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525	is seen, meaning that the apparent T_{Bes} increases monotonically with field, consistent with
526	magnetization data shown in Fig. 6a. In this field orientation, there is no significant additional field
527	dependence above 0.5 T. This can be seen in the inset to Fig. 9c, where the magnetoresistance (MR
528	= ($\rho(H) - \rho(0)$) / $\rho(0)$) as a function of magnetic field at 32 K only shows negative MR out to ~1
529	T. Application of larger fields then causes an inversion in the slope of the MR and a slight increase
530	in magnitude, saturating at high fields. The main panel of Fig 9c shows MR(9 T) as a function of
531	temperature. Upon warming, positive MR at low T gives way to a slight negative MR around T_{Bes} ,
532	with the peak negative MR being ~ -2% at $T_{\text{Bes.}}$ Above the transition, the MR goes positive again,
533	and remains relatively constant past 60 K at about ~0.5%. In this field orientation, MR effects are
534	thus small, the typical growth of positive MR on cooling, as expected in a metal, being interrupted
535	at T_{Bes} with only a small negative MR effect. This small negative MR saturates quickly with
536	increasing field.

537 Fig. 9d displays the magnetic field dependence of ρ vs. T around the Besnus transition, this 538 time with the field applied out of the basal plane, i.e., parallel to the *c*-axis. In this case, a clear 539 decrease in the apparent transition temperature is observed as the out-of-plane magnetic field is 540 increased from 0 to 9 T. The apparent transition temperature shift is monotonic, with an overall 541 shift of ~ 2 K with 9 T of applied field, which is relatively small. The presence of an out-of-plane 542 field thus apparently favors the $T > T_{\text{Bes}}$ phase. Importantly, this is consistent with the remanence 543 and hysteresis data (§2.4.1) (Fig. 6c) in that when the applied field is along, or parallel to, the c-544 axis, the lowest T_{Bes} values are observed. Specifically, it appears that fields applied parallel to the 545 *c*-axis pull the spins out-of-plane near the transition, which allows the phase above the transition 546 to persist to lower T. One way to rationalize this is in terms of Zeeman energy stabilization. The T $> T_{\text{Bes}}$ phase has a larger component of magnetization along the c-axis than the $T < T_{\text{Bes}}$ phase (see 547

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548	<i>Charilaou et al.</i> (2015)), allowing the $T > T_{Bes}$ phase to be stabilized when a field is applied along
549	the c-axis, lowering T_{Bes} as larger fields are applied. Naturally, and as shown in the expanded view
550	in Fig. 9d, the lowering of the apparent transition temperature with out-of-plane field combines
551	with the positive $d\rho/dT$ to generate positive MR. This is shown as a function of magnetic field at
552	30 K in the inset to Fig. 9e, where positive MR up to 7.5 % is seen, not saturating even at 9 T. Fig.
553	9e shows that this positive MR is found at all temperatures between $5 - 60$ K when the field is
554	applied parallel to the <i>c</i> -axis. The positive MR effect associated with the Besnus transition turns
555	on near T_{Bes} , reaches a sharp peak, and then gradually decreases towards 5 K.

556 **3 DISCUSSION**

557 This study represents one of the most exhaustive studies of the Besnus transition to date, 558 performed on a high purity, phase-pure, 4C pyrrhotite single crystal specimen. While the precise 559 mechanism(s) driving the transition still remain elusive, it is possible to now better constrain 560 potential explanations by considering all of the phenomena that have been observed to accompany 561 the transition. Ultimately, a comprehensive theory for the Besnus transition must explain:

- 562 1. An entropy change at T_{Bes} resulting from an apparently second-order phase transition in 563 specific heat (this study, *Grønvold et al.*, 1959)
- Loss of remanent magnetization upon cooling (RTSIRM) and warming (LTSIRM) through
 the transition.
- 566 3. Changing rock magnetic properties (B_c, B_{cr}, M_{rs}) , while M_s remains constant
- 567 4. Irreversible changes after cycling through T_{Bes}, represented as increases in remanence and
 568 MDF (this study, *Dekkers et al., 1989*)
- 5. The disappearance of the second inflection phenomenon in hysteresis loops at T_{Bes} (appears at \approx 200K, aligned with crystal structure) (*Volk et al., 2016, Koulialias at al., 2016*)

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571	6. Symmetry changes of rock magnetic properties from 6-fold above the transition to 4-fold
572	below the transition (Wolfers et al., 2011, Volk et al., 2016, Koulialias et al., 2018)
573	7. Field dependence of $T_{\text{Bes}} \perp [001]$, as determined from magnetic data and resistivity
574	8. Field dependence of T_{Bes} in resistivity $ [001] $
575	9. Different magnetoresistive behavior when applied fields are within the basal plane or
576	parallel to the <i>c</i> -axis
577	10. A distinct decrease in resistivity on cooling through T_{Bes}
578	11. Splitting of Fe sites (4 unique \rightarrow 5 unique) below T_{Bes} , as determined by Mössbauer
579	spectroscopy (Jeandey et al., 1991, Oddou et al., 1992)
580	12. The anomaly seen in short-range magnetic scattering seen in SANS data.
581	Unfortunately, no single mechanism appears capable of simply explaining all of the observed
582	behaviors across the transition, although existing observations do allow us to rule out some

583 proposed mechanisms, which ultimately helps constrain the nature of the transition. All of the 584 mechanisms proposed thus far are discussed in turn below in addition to several new ones.

585 3.1 The role of interacting superstructures. The specimen studied here is purely 4C pyrrhotite 586 and contains no additional superstructures, as supported by powder X-ray diffraction, microprobe 587 analysis, heat capacity measurements, a single magnetic ordering temperature, the absence of a 588 magnetic λ -transition, the M_s value, and the RRR value. Nevertheless, we clearly observe the 589 Besnus transition in all physical properties studied. The specific heat shows a clear peak at ≈ 31.7 590 K, in agreement with the transition temperature determined from resistivity measurements to 591 within one kelvin at zero field. This agrees well with that of the synthetic specimen studied by 592 Grønvold et al. (1959). The transition temperatures in the magnetic susceptibility, RTSIRM, 593 LTSIRM, and FC curves (≈ 32 - 34 K) are slightly higher than the C_p peak, but agree well with

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594 most published data (Dekkers et al., 1989; Rochette et al., 1990; Fillion and Rochette, 1988; Kind 595 et al., 2013). On the other hand, the 26 K transitions described by Charilaou et al. (2015) and 596 Koulialias et al. (2016) deviate markedly. Baranov et al. (2014) showed that Ti substitution can 597 induce changes in vacancy ordering, which results in a change from 4C to 3C in Fe₇S₈. 598 Consequently, it seems possible that the 0.5 % Ni substitution in their sample may be responsible 599 for the additional 5C* structure, and may explain the shifted transition temperature and the 600 suppressed specific heat peak. More substantial Ni substitution results in an entirely new iron 601 sulfide mineral Smythite (Fe,Ni)₉S₁₁ with a trigonal structure, that is often found associated with 602 pyrrhotite. We conclude that the Besnus transition is intrinsic to the 4C phase, and is not caused 603 by interactions of different superstructures as proposed by *Koulialias et al.* (2016, 2018).

604 3.2 The role of twinning. Pyrrhotite is prone to the formation of stacking faults and twins when 605 there are local perturbations in the superstructure (Bin and Pauthenet, 1963). Thus, one could 606 imagine that even subtle changes to the unit cell across the Besnus transition may induce twinning 607 at the nanometer scale (Wolfers et al., 2011). The formation of such features could also reconfigure 608 the distribution of dislocations, vacancies, and pinning sites for domain walls within the material. 609 Consequently, the magnetic interactions between these twins would change, as well as the mobility 610 of domain walls, and thus, the bulk magnetic properties would also change. Nanometer-scale 611 twinning may explain the irreversible changes (e.g., enhanced remanence and coercivity) seen after 612 low temperature cycling. Twinning is also consistent with the length scale of the 613 structural/magnetic changes observed in our high Q SANS data (Figure 2(c,d)), but is somewhat 614 inconsistent with the significant decrease in resistance observed below the transition. The 615 decreasing resistivity can perhaps only be explained by a reduction in the number of twins. 616 Furthermore, twinning is not able to explain the splitting of the occupancy of Fe sites observed in the Mössbauer data. The orientation of twinning in other magnetic minerals, such as magnetite, can be influenced by applied fields, and if this were the case for pyrrhotite, it could explain the field dependence of the transition. Depending on the orientation of the twin boundaries, it is possible that the symmetry of the magnetic properties may change from 6-fold to 4-fold below the transition. As a mechanism, twinning could thus potentially explain many of the phenomena observed below the Besnus transition.

623 3.3 The role of crystallographic transitions. Wolfers et al. (2011) not only proposed a change 624 in twinning schemes, but a true crystallographic transition, with a lowering of the symmetry from 625 monoclinic to triclinic. This is consistent with the changes in the symmetry of rock magnetic 626 properties from 6-fold symmetry above T_{Bes} to 4-fold symmetry below T_{Bes} (Volk et al., 2016). 627 Confusingly, however, our SANS data show a clear six-fold symmetry, even at 5 K. While this is 628 not conclusive evidence against a crystallographic transition, it indicates that any crystallographic 629 changes, if they occur, could be weak. It is worth reiterating in this regard that the 4C monoclinic 630 superstructure at room temperature is only 0.26° different from a perfect hexagonal unit cell 631 (Wolfers et al., 2011). Hence, any additional subtle distortions may not be easily detectable, and 632 will require further high-resolution crystallographic study.

A crystallographic change, where the monoclinic unit cell is subtly distorted to form a different, also monoclinic unit cell is another possible mechanism. Such an *isostructural* or *isomorphic* transformation would be difficult to detect with SANS. However, such a transition could explain many of the phenomena described above. The changing crystal structure could explain the second-order transition indicated by specific heat, but only if the phase transition is truly second-order, or only very weakly first-order. A change in crystal structure would result in changes in the band structure of the material, which would affect electronic transport, *i.e.*, the

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640 resistivity and MR. Therefore, the differences in electronic transport (specifically $d\rho/dT$) above 641 and below T_{Bes} could potentially be explained. Furthermore, the newly created twins could change 642 the domain structure and pin magnetic domain walls at twin boundaries. Consequently, the 643 changing unit cell and rearrangement of domains would lead to a loss of the remanent 644 magnetization. Pinning of domain walls, on the other hand, would increase the coercivity of the 645 crystal. If the changes are small enough, the saturation magnetization should not be affected. 646 Finally, the generation of twins that accompanies an isostructural transformation could explain 647 irreversible changes on cycling through the transition.

648 3.4 The role of changing crystalline anisotropy. Several studies have investigated the 649 crystalline anisotropy of pyrrhotite (Mikami et al. 1959; Bin and Pauthenet, 1963; Sato et al., 650 1964). Bin and Pauthenet (1963) showed that the crystalline anisotropy of pyrrhotite changes upon 651 cooling. The K_4 anisotropy constant, which dominates, shows a maximum, while the K_3 constant 652 changes sign and becomes negative at ≈ 200 K. The changing anisotropy causes the spins to rotate 653 out of the basal plane, which was confirmed by neutron diffraction data by Powell et al. (2004). 654 These changes in magnetic anisotropy may be able to explain the Besnus transition. In this 655 scenario, the magnetic easy axis switches to a different crystallographic direction. Consequently, 656 the remanent magnetization would be lost and domain wall orientation as well as the fundamental 657 magnetic properties would change. However, the measurements of Bin and Pauthenet (1963) were 658 done on an obviously twinned crystal. Furthermore, the authors were not able to determine the anisotropy from measurements in the basal plane at temperatures < 60 K, as the magnetic easy axis 659 660 seemed to switch when a magnetic field is applied. As the second inflection in magnetic hysteresis 661 loops appears at ≈ 180 K, it is likely related to this presumed easy axis switching.

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On balance, it seems unlikely that the Besnus transition is solely caused by a changing anisotropy. Given that the low temperature behavior of the crystalline anisotropy is not well known, the main changes seem appear at 200 K. However, the remanence is unaffected, as are the magnetic properties. Furthermore, the model fails to explain the changes seen in resistivity, Mössbauer, and high Q (Lorentzian) SANS. On the other hand, it seems likely that the anisotropy would change as a byproduct of even a subtle crystallographic transition.

668 3.5 The role of exchange inversion within NiAs structures. Kittel (1960) predicted a 669 phenomenon called exchange inversion in materials with a NiAs crystal structure. Since then it 670 has been found in Mn₂Sb and other Mn-Sb compounds with substitutions (e.g. *Bither et al.*, 1962). 671 At the phase transition, the formerly ferromagnetic coupling between spins switches, and becomes 672 antiferromagnetic. The spins within each Fe plane in pyrrhotite couple ferromagnetically, while 673 the spins between layers couple antiferromagnetically. It seems possible for exchange inversion to 674 take place at T_{Bes} . This would change the coupling within a plane to become antiferromagnetic. 675 becoming ferromagnetic to the adjacent plane. The changing spin orientation would remove most 676 of the remanent magnetization, without affecting the bulk saturation magnetization. The field 677 dependence of the transition could also be explained, as strong magnetic fields could exert force 678 on the spins and hinder their reorganization. Furthermore, a reorientation of spins was found to 679 change electronic properties in a pyrrhotite analogue Fe₇Se₈ with a 4C NiAs structure (*Li et al.*, 680 2016). While being able to explain several phenomena, this scenario fails to explain the splitting 681 of the iron sites in Mössbauer spectroscopy. Furthermore, a spin reorganization should be visible 682 in neutron diffraction data. Powell et al. (2004) did not observe such a transition in their neutron 683 study. However, in their work only one scan was taken below $T_{\text{Bes.}}$

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684 3.6 Future research. In light of the above, this study clearly emphasizes the need for further 685 research along at least three lines: High-resolution structural characterization, laboratory synthesis 686 of single phase pyrrhotite, and linking single crystal observations to bulk sample behavior. The 687 observations in this study and previous work suggest that nanometer-scale twinning within 688 pyrrhotite across the Besnus transition may explain many of the magnetic and physical properties 689 observations. However, it is not clear which mechanism(s) may be driving the twinning. To 690 distinguish between subtle crystallographic transitions or fundamental changes in the exchange 691 coupling, further research is required. Wolfers et al. (2011) proposed that fine low temperature X-692 ray diffraction experiments could be used to obtain a detailed description of any crystallographic 693 change near the 32 K transition. If such an approach is meant to differentiate between monoclinic-694 triclinic and isostructural transitions, then these diffraction experiments may only be possible using 695 synchrotron facilities. Similarly, synchrotron-based X-ray absorption near-edge structure 696 (XANES) analyses may shed light on the changes in Fe site occupancy across the transition 697 originally observed with Mössbauer spectroscopy. Finally, powder neutron diffraction, similar to 698 Powell et al. (2004) at temperatures spanning the transition would further show changes in spin 699 structure as well as possible structural changes.

Magnetic methods may also help constrain the importance of twinning in mixed phase (hexagonal and monoclinic) crystals. If nanometer-scale twinning does occur across the Besnus transition, then vacancies may reorder on warming, and the distribution of pyrrhotite superstructures at room temperature may change after cycling through the Besnus transition. If so, then there should be observable changes in the nature of the lambda transition near 220°C in susceptibility measurements. Given the field dependence of the Besnus transition observed here, it is possible that fields applied along different crystallographic directions may lead to different ratios of superstructures at room temperature. However, if no lambda transition is observed near
220°C after cycling through the Besnus transition, then perhaps the role of twinning may need to
be re-evaluated.

710 4 IMPLICATIONS

711 Many open questions also remain regarding the influence of the Besnus transition on the full-712 vector remanence held by pyrrhotite in natural samples. While the magnitude of magnetization 713 decreases after cycling through the Besnus transition, the direction of a laboratory-induced 714 isothermal remanent magnetization in a single crystal of pyrrhotite is stable (Feinberg et al., 2015). 715 This observation may have important implications for paleomagnetic studies of terrestrial and 716 meteoritic samples containing pyrrhotite. Similar to the ways in which low-temperature treatments 717 that cycle through the Verwey transition in magnetite are used to reduce the remanence held by 718 large multidomain grains and give a more unfettered view of the remanence held by well-behaved, 719 single-domain-like grains, cycling through the Besnus transition may be useful isolating high 720 coercivity remanence in pyrrhotite. Single-domain-like grains also express a Besnus transition, but 721 no observations exist regarding the directional stability of their magnetization across the transition.

The high purity of the 4C single crystal sample in this study is particularly unusual but was essential for demonstrating the intrinsic nature of the Besnus transition and its directional dependence. However, this sample is limited in mass, and, given the numerous ongoing questions about pyrrhotite, there is a clear need to develop dependable synthesis methods (metastability is often a problem) for pyrrhotite that control the resulting superstructures. Many existing synthesis methods are time consuming, require repeated annealing steps (which only partially control the ratio of resulting superstructures), and are limited to only producing fine grained samples. Methods This is a preprint, the final version is subject to change, of the American Mineralogist (MSA) Cite as Authors (Year) Title. American Mineralogist, in press. DOI: https://doi.org/10.2138/am-2018-6514 Preprint submitted to American Mineralogist February 15, 2018

capable of producing large single crystals of well-characterized pyrrhotite would be especiallybeneficial to future research.

731 In addition, due to our improved ability to detect pyrrhotite in natural samples, it is becoming 732 increasingly important to identify how this mineral influences or records numerous geophysical, 733 environmental, and planetary science processes. Similarly, the mineral and its superstructures are 734 important to the materials science community. A deeper understanding of the nature of the Besnus 735 transition will only improve our interpretations regarding pyrrhotite, and we hope this study makes 736 a convincing case for the need for further research that combines high-resolution structural 737 characterization with carefully controlled rock magnetic experiments on natural and synthetic bulk 738 samples.

739 5 ACKNOWLEDGMENTS

740 We would like to thank Anette von der Handt for acquiring electron microprobe data and 741 Jeanette Voelz for help in measuring powder X-ray diffraction data. Furthermore, we would like 742 to thank Rupert Hochleitner and the Bavarian Mineralogical State Collection for providing the 743 sample material and for their kind support. EM, BV, MM and CL acknowledge support from the 744 DOE through the UMN Center for Quantum Materials under DE-FG02-06ER46275 and DE-SC-745 0016371. We acknowledge the support of the National Institute of Standards and Technology, US 746 Department of Commerce, in providing the neutron facilities used in this work; the assistance of John Barker is gratefully acknowledged in that regard. The Institute for Rock Magnetism is a U.S. 747 748 National Multiuser Facility supported through the NSF-EAR Instrumentation and Sciences 749 Program and by funding from the University of Minnesota. This is IRM publication (#1710).

751 6 REFERENCES

- Arnold, R.G. (1967) Range in composition and structure of 82 natural terrestrial pyrrhotites.
- 753 The Canadian Mineralogist, 9, 31–50.
- Aubourg, C., and Pozzi, J.-P. (2010) Toward a new <250°C pyrrhotite-magnetite
- 755 geothermometer for claystones. Earth and Planetary Science Letters, 294, 47–57.
- 756 Baranov, N.V., Ibrahim, P.N.G., Selezneva, N.V., Kazantsev, V.A., Volegov, A.S., and
- 757 Shishkin, D.A. (2014) Crystal structure, phase transitions and magnetic properties of pyrrhotite-
- type compounds Fe7–xTixS8. Physica B: Condensed Matter, 449, 229–235.
- Bertaut, E.F. (1953) Contribution à l'étude des structures lacunaires: la pyrrhotine. Acta
 Crystallographica, 6, 557–561.
- Besnus, M.J., and Meyer, A.J. (1964) Nouvelles données expérimentales sur le magnétisme de
 la pyrrhotine naturelle. Proc. Int. Conf. Mag., 20, 507–511.
- 763 Bezaeva, N.S., Chareev, D.A., Rochette, P., Kars, M., Gattacceca, J., Feinberg, J.M., Sadykov,
- 764 R.A., Kuzina, D.M., and Axenov, S.N. (2016) Magnetic characterization of non-ideal single-
- domain monoclinic pyrrhotite and its demagnetization under hydrostatic pressure up to 2 GPa with
- implications for impact demagnetization. Physics of the Earth and Planetary Interiors, 257, 79–90.
- Bin, M., and Pauthenet, R. (1963) Magnetic Anisotropy in Pyrrhotite. Journal of Applied
 Physics, 34, 1161–1162.
- Bither, T.A., Walter, P.H.L., Cloud, W.H., Swoboda, T.J., and Bierstedt, P.E. (1962) New
 Modified Mn2Sb Compositions Showing Exchange Inversion. Journal of Applied Physics, 33,
 1346–1347.

- Cantu, J., Gonzalez, L.E., Goodship, J., Contreras, M., Joseph, M., Garza, C., Eubanks, T.M.,
- and Parsons, J.G. (2016) Removal of arsenic from water using synthetic Fe 7 S 8 nanoparticles.
- Chemical Engineering Journal, 290, 428–437.
- 775 Charilaou, M., Kind, J., Koulialias, D., Weidler, P.G., Mensing, C., Loffler, J.F., and Gehring,
- A.U. (2015) Magneto-electronic coupling in modulated defect-structures of natural Fe1-xS.
- Journal of Applied Physics, 118, 083903–6.
- 778 Dekkers, M.J. (1988) Magnetic properties of natural pyrrhotite Part I: Behaviour of initial
- susceptibility and saturation-magnetization-related rock-magnetic parameters in a grain-size
- dependent framework. Physics of the Earth and Planetary Interiors, 52, 376–393.
- 781 Dekkers, M.J., Mattéi, J.-L., Fillion, G., and Rochette, P. (1989) Grain-size dependence of the
- magnetic behavior of pyrrhotite during its low-temperature transition at 34 K. Geophysical
 Research Letters, 16, 855–858.
- Dunlop, D.J. (2002) Theory and application of the Day plot (Mrs/ Msversus Hcr/ Hc) 1. Theoretical curves and tests using titanomagnetite data. Journal of Geophysical Research: Solid Earth, 107, 2056.
- Dunlop, D.J., and Özdemir, Ö. (1997) Rock Magnetism. Cambridge University Press,
 Cambridge.
- Erd, R.C., Evans, H.T., and Richter, D.H. (1957) Smythite, a New Iron Sulfide, and Associated
 Pyrrhotite From Indiana. American Mineralogist, 42, 309–333.
- Fricsson, T., Amcoff, Ö., and Nordblad, P. (1994) Vacancy ordering in Fe7Se8-Fe7S8 solid
 solutions studied by Mössbauer, X-ray and magnetization techniques. Hyperfine Interactions, 90,
 515–520.

- Fegley, B., Jr., Lodders, K., Treiman, A.H., and Klingelhöfer, G. (1995) The Rate of Pyrite
- 795 Decomposition on the Surface of Venus. Icarus, 115, 159–180.
- Feinberg, J.M., Solheid, P.A., Swanson-Hysell, N.L., Jackson, M.J., and Bowles, J.A. (2015)
- 797 Full vector low-temperature magnetic measurements of geologic materials. Geochemistry,
- Geophysics, Geosystems, 16, 301–314.
- Fillion, G., and Rochette, P. (1988) The Low Temperature Transition In Monoclinic Pyrrhotite.
- Le Journal de Physique Colloques, 49, C8–907–C8–908.
- Gilder, S.A., Egli, R., Hochleitner, R., Roud, S.C., Volk, M.W.R., Le Goff, M., and de Wit,
- 802 M. (2011) Anatomy of a pressure-induced, ferromagnetic-to-paramagnetic transition in pyrrhotite:
- 803 Implications for the formation pressure of diamonds. Journal of Geophysical Research: Solid 804 Earth 116 P10101
- 804 Earth, 116, B10101.
- Gillard, T.M., Phelan, D., Leighton, C., and Bates, F.S. (2015) Determination of the Lamellae-
- 806 to-Disorder Heat of Transition in a Short Diblock Copolymer by Relaxation Calorimetry.
- 807 Macromolecules, 48, 4733–4741.
- E. S. R. Gopal (1966) Specific Heats at Low Temperatures, Plenum Press, New York
- Grønvold, F., Stølen, S., Labban, A.K., and Westrum, E.F. (1991) Thermodynamics of iron
- 810 sulfides I. Heat capacity and thermodynamic properties of Fe9S10 at temperatures from 5 K to 740
- 811 K. The Journal of Chemical Thermodynamics, 23, 261–272.
- 812 Grønvold, F., Westrum, E.F., and Chou, C. (1959) Heat Capacities and Thermodynamic
- 813 Properties of the Pyrrhotites FeS and Fe0.877S from 5 to 350°K. The Journal of Chemical Physics,
- 814 30, 528–531.

- 815 Halgedahl, S.L., and Fuller, M.D. (1981) The dependence of magnetic domain structure upon
- 816 magnetization state in polycrystalline pyrrhotite. Physics of the Earth and Planetary Interiors, 26,
- 817 93–97.
- 818 Herndon, J.M., Rowe, M.W., Larson, E.E., and Watson, D.E. (1975) Origin of magnetite and
- 819 pyrrhotite in carbonaceous chondrites. Nature, 253, 516–518.
- 820 Jeandey, C., Oddou, J.L., Mattéi, J.-L., and Fillion, G. (1991) Mössbauer investigation of the
- pyrrhotite at low temperature. Solid State Communications, 78, 195–198.
- Kasama, T., Church, N.S., Feinberg, J.M., Dunin-Borkowski, R.E., and Harrison, R.J. (2010)
- 823 Direct observation of ferrimagnetic/ferroelastic domain interactions in magnetite below the
- 824 Verwey transition. Earth and Planetary Science Letters, 297, 10–17.
- 825 Kind, J., Garcia-Rubio, I., Charilaou, M., Nowaczyk, N.R., Loffler, J.F., and Gehring, A.U.
- 826 (2013) Domain-wall dynamics in 4C pyrrhotite at low temperature. Geophysical Journal
 827 International, 195, 192–199.
- Kittel, C. (1960) Model of Exchange-Inversion Magnetization. Physical Review, 120, 335–
 342.
- 830 Kobayashi, H., Nozue, T., Matsumura, T., Suzuki, T., and Kamimura, T. (1999) The low-
- temperature specific heat of FeS and M 0.875X (M = Fe, Co; X = S, Se) with a NiAs-like structure.
- Journal of Physics: Condensed Matter, 11, 8673–8679.
- 833 Koulialias, D., Charilaou, M., Schäublin, R., Mensing, C., Weidler, P.G., Loffler, J.F., and
- 834 Gehring, A.U. (2018) Ordered defects in Fe 1- xS generate additional magnetic anisotropy
- symmetries. Journal of Applied Physics, 123, 033902–11.

- Koulialias, D., Kind, J., Charilaou, M., Weidler, P.G., Loffler, J.F., and Gehring, A.U. (2015)
- 837 Variable defect structures cause the magnetic low-temperature transition in natural monoclinic
- 838 pyrrhotite. Geophysical Journal International, 204, 961–967.
- 839 Kuipers, A.J.M., and Brabers, V.A.M. (1979) Electrical transport in magnetite near the Verwey
- transition. Physical Review B, 20, 594–600.
- Lapakko, K.A. (2015) Preoperational assessment of solute release from waste rock at proposed
- 842 mining operations. Applied Geochemistry, 57, 106–124.
- Lashley, J.C., Hundley, M.F., Migliori, A., Sarrao, J.L., Pagliuso, P.G., Darling, T.W., Jaime,
- 844 M., Cooley, J.C., Hults, W.L., Morales, L., and others (2003) Critical examination of heat capacity
- 845 measurements made on a Quantum Design physical property measurement system. Cryogenics,
- 846 43, 369–378.
- Levinson, L.M., and Treves, D. (1968) Mössbauer study of the magnetic structure of Fe7S8.
 Journal of Physics and Chemistry of Solids.
- Li, C.H. (1932) Magnetic Properties of Magnetite Crystals at Low Temperature. Physical
 Review, 40, 1002–1012.
- Li, G., Zhang, B., Baluyan, T., Rao, J., Wu, J., Novakova, A.A., Rudolf, P., Blake, G.R., de
- 852 Groot, R.A., and Palstra, T.T.M. (2016) Metal-Insulator Transition Induced by Spin Reorientation
- 853 in Fe7Se8 Grain Boundaries. Inorganic Chemistry, 55, 12912–12922.
- Mikami, I., Hirone, T., Watanabe, H., Maeda, S., Adachi, K., and Yamada, M. (1959) On the
- 855 Magnetic Anisotropy of a Pyrrhotite Crystal. Journal of the Physical ..., 14, 1568–1572.
- 856 Morimoto, N., Nakazawa, H., Nishigucmi, K., and Tokonami, M. (1970) Pyrrhotites:
- 857 Stoichiometric Compounds with Composition Fen--1Sn ($n \ge 8$). Science, 168, 964–966.

- 858 Muhlbauer, S., Honecker, D., Perigo, E., Bergner, F., Disch, S., Heinemann, A., Erokhin, S.,
- 859 Berkov, D., Leighton, C., Eskildsen, M., and others (2018) Magnetic Small-Angle Neutron
- 860 Scattering. Reviews of Modern Physics, submitted.
- 861 Muttoni, G. (1995) "Wasp-waisted" hysteresis loops from a pyrrhotite and magnetite-bearing
- remagnetized Triassic limestone. Geophysical Research Letters, 22, 3167–3170.
- 863 Muxworthy, A.R., Bland, P.A., Davison, T.M., Moore, J., Collins, G.S., and Ciesla, F.J. (2017)
- 864 Evidence for an impact-induced magnetic fabric in Allende, and exogenous alternatives to the core
- dynamo theory for Allende magnetization. Meteoritics & Planetary Science, 112.
- 866 Nakazawa, H., and Morimoto, N. (1970) Pyrrhotite Phase Relations below 320°C. Proceedings
- 867 of the Japan Academy, 46, 678–683.
- 868 O'Reilly, W., Hoffmann, V., Chouker, A.C., Soffel, H.C., and Menyeh, A. (2000) Magnetic
- properties of synthetic analogues of pyrrhotite ore in the grain size range 1-24 µm. Geophysical
 Journal International, 142, 669–683.
- Oddou, J.L., Jeandey, C., Mattéi, J.-L., and Fillion, G. (1992) Mössbauer study of the lowtemperature transition in pyrrhotite. Journal of Magnetism and Magnetic Materials, 104-107,
 1987–1988.
- Pastor, J.M., Pérez-Landazábal, J.I., Gómez-Polo, C., Recarte, V., Larumbe, S., Santamarta,
- 875 R., Fernandes Silva, M., Gómez Pineda, E.A., Winkler Hechenleitner, A.A., and Lima, M.K.
- 876 (2012) Entropy change linked to the magnetic field induced Morin transition in Hematite
- 877 nanoparticles. Applied Physics Letters, 100, 063102.
- 878 Powell, A.V., Vaqueiro, P., Knight, K.S., and Chapon, L.C. (2004) Structure and magnetism
- 879 in synthetic pyrrhotite Fe 7 S 8: A powder neutron-diffraction study. Physical Review B.

- Quidelleur, X., Valet, J.-P., and Thouveny, N. (1992) Multicomponent magnetization in
 paleomagnetic records of reversals from continental sediments in Bolivia. Earth and Planetary
 Science Letters, 111, 23–39.
- 883 Rochette, P., Fillion, G., and Dekkers, M.J. (2011) Interpretation of Low-Temperature Data
- 884 Part IV: The Low-Temperature Magnetic Transition of Monoclinic Pyrrhotite. The IRM Quarterly,
- 885 21, 1–11.
- 886 Rochette, P., Fillion, G., Mattéi, J.-L., and Dekkers, M.J. (1990) Magnetic transition at 30–34
- 887 Kelvin in pyrrhotite: insight into a widespread occurrence of this mineral in rocks. Earth and
- 888 Planetary Science Letters, 98, 319–328.
- 889 Rochette, P., Lorand, J.P., Fillion, G., and Sautter, V. (2001) Pyrrhotite and the remanent
- 890 magnetization of SNC meteorites: a changing perspective on Martian magnetism. Earth and
 891 Planetary Science Letters, 190, 1–12.
- H. M. Rosenberg (1963) Low Temperature Solid State Physics, Oxford University Press, New
 York,
- Sato, K., Yamada, M., and Hirone, T. (1964) Megnetocrystalline Anisotropy of Pyrrhotite.
 Journal of the Physical Scociety of Japan, 19, 1592–1595.
- Schwarz, E.J., and Vaughan, D.J. (1972) Magnetic Phase Relations of Pyrrhotite. Journal of
- geomagnetism and geoelectricity, 24, 441–458.
- Sharp, W.E. (1966) Pyrrhotite: a Common Inclusion in South African Diamonds. Nature, 211,
 402–403.

- 900 Shepherd, J.P., Koenitzer, J.W., Aragón, R., Spal ek, J., and Honig, J.M. (1991) Heat capacity
- 901 and entropy of nonstoichiometric magnetite Fe3(1- δ)O4: The thermodynamic nature of the
- 902 Verwey transition. Physical Review B, 43, 8461–8471.
- 903 Shimada, K., Mizokawa, T., Mamiya, K., Saitoh, T., Fujimori, A., Ono, K., Kakizaki, A., Ishii,
- 904 T., Shirai, M., and Kamimura, T. (1998) Spin-integrated and spin-resolved photoemission study
- 905 of Fe chalcogenides. Physical Review B, 57, 8845–8853.
- 906 Shirai, M., Suzuki, N., and Motizuki, K. (1996) Electronic band structure and photoemission
- 907 spectra of Fe7S8. Journal of Electron Spectroscopy and Related Phenomena, 78, 95–98.
- 908 Tokonami, M., Nishiguchi, K., and Morimoto, N. (1972) Crystal structure of a monoclinic
- 909 pyrrhotite (Fe7S8). American Mineralogist, 57, 1066–1080.
- 910 van der Pauw, L.J. (1958) A method of measuring the resistivity and Hall coefficient on
- 911 lamellae of arbitrary shape, Philips Tech. Rev. 20, 220
- 912 Volk, M.W.R., Gilder, S.A., and Feinberg, J.M. (2016) Low-temperature magnetic properties
- 913 of monoclinic pyrrhotite with particular relevance to the Besnus transition. Geophysical Journal
- 914 International, 207, 1783–1795.
- 915 Weiss, B.P., Maloof, A.C., Tailby, N., Ramezani, J., Fu, R.R., Hanus, V., Trail, D., Bruce
- 916 Watson, E., Harrison, T.M., Bowring, S.A., and others (2015) Pervasive remagnetization of detrital
- 917 zircon host rocks in the Jack Hills, Western Australia and implications for records of the early
- 918 geodynamo. Earth and Planetary Science Letters, 430, 115–128.
- 919 Wolfers, P., Fillion, G., Ouladdiaf, B., Ballou, R., and Rochette, P. (2011) The Pyrrhotite 32
- 920 K Magnetic Transition. Solid State Phenomena, 170, 174–179.

- 921 Zhang, Kailong, Zhang, T., Liang, J., Zhu, Y., Lin, N., and Qian, Y. (2015) A potential
- 922 pyrrhotite (Fe 7 S 8) anode material for lithium storage. RSC Advances, 5, 14828–14831.
- 923 Zhang, Ke, Zheng, H., Wang, J., and Wang, R. (2008) Transmission electron microscopy on
- 924 iron monosulfide varieties from the Suizhou meteorite. Physics and Chemistry of Minerals, 35,
- 925 425–432.

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928 Figure 1: Powder X-ray diffraction (room temperature) of a sister specimen using Co 929 radiation. The red line shows the fit obtained with Rietveld refinement. The inset shows an 930 expanded view of the most prominent peak, in comparison with data from Charilaou et al. (2015) 931 in green. The misfit is between the refinement and data.

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933 Figure 2: SANS results obtained with a 4C pyrrhotite crystal. a) The low Q 2D scattering 934 pattern (Q_x-Q_y map) obtained at 5 K (the brightness/color of each pixel indicates the differential 935 scattering cross-section). b) The sum of scattering cross-sections over the indicated low Q range 936 $(0.005 \text{ Å}^{-1} < Q < 0.009 \text{ Å}^{-1})$ as a function of the in-plane angle, ϕ , as defined in panel (a). These 937 data were averaged over all 6 temperatures probed (from 60 to 5 K), due to the absence of any 938 significant temperature dependence of the angular dependence. c) A representative (T = 45 K) 939 scattering cross-section $(d\Sigma/d\Omega)$ vs. Q scan obtained by circular integration of the 2D scattering 940 patterns. The fit (solid blue line) is the sum of a Porod term (blue dashed line) and a Lorentzian 941 term (red dashed line), as discussed in the text. These two terms are defined by $d\Sigma/d\Omega =$

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- 942 $(d\Sigma/d\Omega)_{\rm P}/Q^n + (d\Sigma/d\Omega)_{\rm L}/(Q^2 + \kappa^2)$, where "P" and "L" subscripts are for Porod and Lorentzian, n
- 943 is the generalized Porod exponent, and κ is defined in the text. d) The temperature (T) dependence
- 944 of the Porod (black) and Lorentzian (red) cross-sections, .

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Figure 3: Heat capacity measurements obtained in this study along with those from previously published studies. (a) C_P as a function of temperature (T), with the inset focusing on the Besnus transition at 32 K. The solid lines are simple linear segments connecting data points. (b) C_P/T as a function of T² at low temperature. The solid black line is a fit to the data as described in the text.

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955 Figure 4: Zero field low-temperature cycling of a room temperature SIRM moment acquired 956 in a 2.5 T field in the basal (a-b) plane. a) and b) show different orientations of the single crystal. 957 The experiment was repeated twice for each orientation, where the first experiment was measured 958 using a larger temperature interval (every 5 K) than the second experiment (every 0.5 K). The 959 magnetization in the second experiment is usually higher than in the first, suggesting a slight 960 hardening after the first thermal cycling. Arrows show cooling or warming, numbers indicate the 961 measurement sequence. Dashed lines denote the temperature of the heat capacity peak. Insets in 962 both (a) and (b) are SIRM moments over the entire temperature range measured, with solid lines 963 being the cooling paths and dashed lines warming.

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Figure 5: Field cooling (FC) experiments in different fields a). FC in 500 mT was measured along two arbitrary, perpendicular directions in the basal plane (Φ_1, Φ_2). Crosses mark direction Φ_2 on cooling and warming (filled marker). b) shows M(B = 500 mT, T, Φ) determined from hysteresis loops (see text) at selected orientations within the basal plane of an oriented single crystal of Fe₂S₈. Inset in b) demonstrates the lack of hysteresis of a high-resolution (0.5 K) FC measurement (see a)) in a 500 mT field (Φ_1) on warming and cooling. c) shows warming of a lowtemperature SIRM acquired by cooling in field (a) to 10 K.

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974 Figure 6: Transition temperature calculated from the maximum of dM/dT of the positive 975 ascending branch of the hysteresis loop for all orientations and fields. a) Average transition 976 temperature with one standard deviation (shaded region) for all orientations as a function of 977 magnetic field. The inset shows a zoomed-in view of the first 100 mT b) Directional dependence 978 of the mean transition temperature over all applied fields. A data point's distance from the center 979 of the plot corresponds to the T_{Bes} for that orientation and varies between 29 and 35K. The shaded 980 region shows 1σ uncertainty. c) Contour plot for all fields (radius) and orientation (angle), dark = 981 low, light = high transition temperature. Hashed regions show the area where T_{Bes} could not be 982 calculated. Solid lines in b) and c) show the crystallographic axis as determined by EBSD. The 983 dashed line is the projection of the c-axis onto the plane.

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986 **Figure 7**: Mass-normalized in-phase (χ') and out-of-phase (χ'') susceptibility as a function of 987 temperature in different field amplitudes (a, b) (frequency = 100 Hz, marker every 4th 988 measurement) and frequencies (c, d) (amplitude = 0.30 mT). Circles show measurements within 989 the basal plane (χ_{\perp}) of the crystal, while triangles show measurements done along the 990 crystallographic c-axis (χ_{\parallel} , secondary y-axis). Dashed vertical lines show the maximum from the 991 heat capacity measurement. Inset shows a zoom into the secondary peak overlaid with specific 992 heat data (see. fig. 2) in arb. units. The background in specific heat was removed by polynomial 993 interpolation.

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Figure 8: Ratios of in-plane over out-of-plane susceptibilities $(\chi_{\perp}/\chi_{\parallel})$ as a function of temperature for different field amplitudes. The inset shows the ratio determined from highresolution AC susceptibility data (100 Hz, 0.3 mT). The dashed line shows the transition temperature determined from specific heat.

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1000 **Figure 9:** 4C pyrrhotite single crystal electrical resistivity measured via the van der Pauw 1001 method as a function of temperature and magnetic field. (a) a-b plane resistivity (ρ) as a function

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- 1002 of temperature (T) upon warming from 4 K to 300 K in zero field (B). (b-d) show expanded views
- 1003 of transport and magnetotransport properties in the T_{Bes} region. (b) ρ vs. T as a function of B (from
- $1004 \quad 0 0.5 \text{ T}$) applied in the basal plane. (c) Corresponding magnetoresistance (MR) as a function of
- 1005 T at 9 T. The inset shows the B dependence of the MR at 32 K. (d) ρ vs. T as a function of B (from
- 1006 0 9 T) applied along [001]. (e) Corresponding MR as a function of T at 9 T. The inset shows the
- 1007 B dependence of the MR at 30 K. Red crosses in (c) and (e) mark the measurement temperature
- 1008 for the MR(B) traces shown in the insets.

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1011 **Figure 10:** Schematic representation of exchange inversion. Arrows indicate the spins of the

1012 atoms.

- 1014 **Table 1**: Mean saturation magnetization ($M_s(\theta_1, \theta_2)$) determined from in-field cooling in a 3 T
- field and saturation remanence (M_{rs}) at 10 (LTSIRM) and 300 K (RTSIRM). 1015

	M _{rs} [Am ² /kg]	M _s [Am ² /kg]	M_{rs}/M_s
10 K	13.85	20.09	0.69
300 K	1.67	20.19	0.08

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- 1018 **Table 2:** Amplitude dependence of susceptibility above (300 K) and below (20 K) T_{Bes}.
- 1019 Absolute values of the susceptibility are given in mAm²kg⁻¹mT⁻¹

Amplitude in mT	0.2	0.5	1.0	2.0	3.0
χ (300 K)	1.55	1.68	1.82	2.04	2.15
χ _{II} (20 K)	0.97	0.97	0.95	0.96	0.96
χ _{II} (20 K / 300 K) [%]	62.26	57.72	52.09	47.22	44.74
χ⊥(300 K)	258.96	296.64	324.94	354.51	370.13
χ⊥(20 K)	1.84	1.82	1.84	1.83	1.83
χ _⊥ (20 K / 300 K) [%]	0.71	0.61	0.57	0.52	0.50
χ⊥/ χ (300 K)	166.86	176.60	178.43	174.15	172.55
χ⊥/ χ (20 K)	1.91	1.87	1.94	1.91	1.91

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