2 Revised Manuscript (First Revision) 3	1	
 Defect Contributions to the Heat Capacities and Stabilities of Some Chain, Ring and Sheet Silicates, with Implications for Mantle Minerals Ring and Sheet Silicates, with Implications for Mantle Minerals H. Wayne Nesbitt¹, A.N. Cormack², Grant S. Henderson³ ¹Dept. of Earth Sciences, Univ. of Western Ontario, London On., N6A 5B7, Canada ²New York State College of Ceramics, Alfred Univ., Alfred, NY 14802, USA ³Dept. of Earth Science, Univ. of Toronto, Toronto On., M5S 3B1, Canada ³Dept. of Earth Science, Univ. of Toronto, Toronto On., M5S 3B1, Canada ⁶Corresponding Author: H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: <u>hwn@uwo.ca</u> 	2	Revised Manuscript (First Revision)
 4 Defect Contributions to the Heat Capacities and Stabilities of Some Chain, 5 Ring and Sheet Silicates, with Implications for Mantle Minerals 6 7 8 H. Wayne Nesbitt¹, A.N. Cormack², Grant S. Henderson³ 9 10 ¹Dept. of Earth Sciences, Univ. of Western Ontario, London On., N6A 5B7, Canada 12 ² New York State College of Ceramics, Alfred Univ., Alfred, NY 14802, USA 13 ³Dept. of Earth Science, Univ. of Toronto, Toronto On., M5S 3B1, Canada 14 15 16 17 18 19 20 21 Corresponding Author: H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: hwn@uwo.ca 	3	
 Ring and Sheet Silicates, with Implications for Mantle Minerals Ring and Sheet Silicates, with Implications for Mantle Minerals H. Wayne Nesbitt¹, A.N. Cormack², Grant S. Henderson³ ¹Dept. of Earth Sciences, Univ. of Western Ontario, London On., N6A 5B7, Canada ² New York State College of Ceramics, Alfred Univ., Alfred, NY 14802, USA ³Dept. of Earth Science, Univ. of Toronto, Toronto On., M5S 3B1, Canada ³Dept. of Earth Science, Univ. of Toronto, Toronto On., M5S 3B1, Canada ⁶Corresponding Author: H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: hwn@uwo.ca 	4	Defect Contributions to the Heat Capacities and Stabilities of Some Chain,
 6 7 8 H. Wayne Nesbitt¹, A.N. Cormack², Grant S. Henderson³ 9 10 ¹Dept. of Earth Sciences, Univ. of Western Ontario, London On., N6A 5B7, Canada ² New York State College of Ceramics, Alfred Univ., Alfred, NY 14802, USA ³Dept. of Earth Science, Univ. of Toronto, Toronto On., M5S 3B1, Canada ¹⁴ 15 16 17 18 19 20 21 Corresponding Author: H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: <u>hwn@uwo.ca</u> 	5	Ring and Sheet Silicates, with Implications for Mantle Minerals
 H. Wayne Nesbitt¹, A.N. Cormack², Grant S. Henderson³ H. Wayne Nesbitt¹, A.N. Cormack², Grant S. Henderson³ ¹Dept. of Earth Sciences, Univ. of Western Ontario, London On., N6A 5B7, Canada ² New York State College of Ceramics, Alfred Univ., Alfred, NY 14802, USA ³Dept. of Earth Science, Univ. of Toronto, Toronto On., M5S 3B1, Canada ³Dept. of Earth Science, Univ. of Toronto, Toronto On., M5S 3B1, Canada ⁶ Corresponding Author: H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: <u>hwn@uwo.ca</u> 	6	
 H. Wayne Nesbitt¹, A.N. Cormack², Grant S. Henderson³ ¹Dept. of Earth Sciences, Univ. of Western Ontario, London On., N6A 5B7, Canada ²New York State College of Ceramics, Alfred Univ., Alfred, NY 14802, USA ³Dept. of Earth Science, Univ. of Toronto, Toronto On., M5S 3B1, Canada ³Dept. of Earth Science, Univ. of Toronto, Toronto On., M5S 3B1, Canada ⁶Corresponding Author: H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: <u>hwn@uwo.ca</u> 	7	
 9 10 11 ¹Dept. of Earth Sciences, Univ. of Western Ontario, London On., N6A 5B7, Canada 12 ² New York State College of Ceramics, Alfred Univ., Alfred, NY 14802, USA 13 ³Dept. of Earth Science, Univ. of Toronto, Toronto On., M5S 3B1, Canada 14 15 16 17 18 19 20 21 Corresponding Author: H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: <u>hwn@uwo.ca</u> 	8	H. Wayne Nesbitt ¹ , A.N. Cormack ² , Grant S. Henderson ³
 ¹⁰ ¹Dept. of Earth Sciences, Univ. of Western Ontario, London On., N6A 5B7, Canada ² New York State College of Ceramics, Alfred Univ., Alfred, NY 14802, USA ³Dept. of Earth Science, Univ. of Toronto, Toronto On., M5S 3B1, Canada ¹⁴ ¹⁵ ¹⁶ ¹⁷ ¹⁸ ¹⁹ ²⁰ ²¹ Corresponding Author: H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: <u>hwn@uwo.ca</u> 	9	
 ¹Dept. of Earth Sciences, Univ. of Western Ontario, London On., N6A 5B7, Canada ² New York State College of Ceramics, Alfred Univ., Alfred, NY 14802, USA ³Dept. of Earth Science, Univ. of Toronto, Toronto On., M5S 3B1, Canada 14 15 16 17 18 19 20 20 21 Corresponding Author: H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: <u>hwn@uwo.ca</u>	10	
 ² New York State College of Ceramics, Alfred Univ., Alfred, NY 14802, USA ³ Dept. of Earth Science, Univ. of Toronto On., M5S 3B1, Canada 14 15 16 17 18 19 20 20 21 Corresponding Author: 18 22 23 24 Corresponding Author: 19 20 20 21 H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: <u>hwn@uwo.ca</u>	11	¹ Dept. of Earth Sciences, Univ. of Western Ontario, London On., N6A 5B7, Canada
 ³Dept. of Earth Science, Univ. of Toronto On., M5S 3B1, Canada General Science, Univ. of Toronto On., M5S 3B1, Canada Corresponding Author: H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: <u>hwn@uwo.ca</u> 	12	² New York State College of Ceramics, Alfred Univ., Alfred, NY 14802, USA
 14 15 16 17 18 19 20 21 Corresponding Author: H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: <u>hwn@uwo.ca</u> 	13	³ Dept. of Earth Science, Univ. of Toronto, Toronto On., M5S 3B1, Canada
 15 16 17 18 19 20 21 Corresponding Author: H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: <u>hwn@uwo.ca</u> 	14	
 16 17 18 19 20 21 Corresponding Author: H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: <u>hwn@uwo.ca</u> 23 24 	15	
 17 18 19 20 21 Corresponding Author: H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: <u>hwn@uwo.ca</u> 23 24 	16	
 18 19 20 21 Corresponding Author: H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: <u>hwn@uwo.ca</u> 23 24 	17	
 19 20 21 Corresponding Author: H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: <u>hwn@uwo.ca</u> 23 24 	18	
 20 21 Corresponding Author: 22 H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: <u>hwn@uwo.ca</u> 23 24 	19	
 21 Corresponding Author: 22 H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: <u>hwn@uwo.ca</u> 23 24 	20	
	21 22 23 24	Corresponding Author: H. Wayne Nesbitt, Tel: (519) 661-2100 ext. 83194, e-mail: <u>hwn@uwo.ca</u>

25

ABSTRACT

26 At temperatures less than \sim 1500K, previously published C_p data demonstrate that the heat 27 capacities of orthoenstatite, proto-enstatite, diopside and pseudowollastonite include primarily 28 Debye type vibrational and anharmonic contributions whereas the alkali chain, sheet and ring 29 silicates, Na₂SiO₃, Li₂SiO₃, K₂SiO₃ and Na₂Si₂O₅ include a third contribution. The third contribution to C_p arises from defect formation due to the mobility Na, K, Li, and O²⁻. The 30 31 contribution becomes apparent at temperatures above 700-800K for Na and K silicates, and above 32 900-1000K for Li metasilicate. With strong thermal agitation, alkali-non-bridging oxygen (NBO) 33 bonds are ruptured with the cations exiting their structural sites to occupy interstitial sites, thereby producing intrinsic Frenkel defects which contribute to the Cp of the alkali silicates. The 34 35 magnitudes of the C_p defect contributions correlate inversely with cation-oxygen bond strengths, as 36 measured by bond dissociation energies. K-O and Na-O bond strengths are weak (239 and 257 37 kJ/mol) and defect contributions are large for these alkali chain, ring and sheet silicates. The greater bond strength of Li-O (341 kJ/mol) correlates with a weaker defect contribution to the C_p 38 of Li₂SiO₃. Mg-O and Ca-O bonds are stronger still (394 and 464 kJ/mol) and no C_p defect 39 40 contributions are observed for the pyroxenes and pseudowollastonite up to ~ 1500 K.

Above ~800K a polymerization reaction occurs in Na₂SiO₃ which produces some Q³ species and free oxygen (O²⁻ or oxide ion). The polymerization reaction annihilates an oxygen structural site so that the O²⁻ produced must reside on non-structural sites thus producing intrinsic anionic defects. The same reactions likely occur in Na₂Si₂O₅ and K₂SiO₃. Raman spectra of Na₂SiO₃ indicate >10% of Na⁺ and ~1.7% of O²⁻ on interstitial sites at 1348K.

46 Ca- and Mg-bearing mantle minerals subjected to temperature greater than ~1500K
47 experience the destabilizing effects of disordering (Frenkel defect formation). The minerals may

respond either by changing their composition or by changing phase. An abundance of Ca and Na defects in pyroxenes, for example, likely promote production of new components (e.g., CaAl₂SiO₆, NaAlSi₂O₆) in pyroxenes. By their production, Ca and Na defect concentrations are reduced thereby stabilizing the phases. Mg-O bond dissociation and production of intrinsic Mg^{2+} and O^{2-} point defects within olivine likely destabilize it and promote the phase transition to wadsleyite at the base of the upper mantle.

- Keywords: heat capacity of silicate minerals, Frenkel defects in silicates, Cation Disorder, Silicate
 mineral stability, stability of mantle minerals
- 57

INTRODUCTION

Richet et al. (1996) compared the temperature dependence of the heat capacities (C_p) of the 58 isostructural chain silicates Li₂SiO₃(c) and Na₂SiO₃(c), and noted that C_p of Na₂SiO₃(c) was 59 60 consistently greater than that of Li₂SiO₃(c), regardless of temperature. The differing values below ~600K may be due to different Debye Temperatures (θ) of the phases because θ affects mostly the 61 low temperature C_p values of a phase. At higher temperatures, θ values have minimal effect on 62 63 heat capacity at constant volume (C_v), or on C_p , in that C_v approaches the universal value of 3nR 64 regardless of θ values. Richet et al. (1996) explored these differences at both low and high temperatures using sophisticated C_p models (e.g., Kieffer, 1982; 1980), which included 65 anharmonic effects, and they summarized the situation by stating that the excess C_p of Na₂SiO₃ 66 remained unexplained. We address this conundrum by evaluating previously published 67 contributions to C_p of eight chain, ring and sheet silicates, orthoenstatite (MgSiO₃), proto-enstatite 68 69 (MgSiO₃), diopside (MgCaSi₂O₆), pseudowollastonite (CaSiO₃), Na₂SiO₃, Na₂Si₂O₅ K₂SiO₃ and Li₂SiO₃. All crystals contain Q² species (silicate chains or rings) except Na₂Si₂O₅ which is a sheet 70 silicate consisting of O³ species (where O represents a Si tetrahedron and the superscript indicates 71

the number of bridging oxygen atoms bonded to the central Si atom of the tetrahedron).

73 Cations of some crystals, heated to high temperature, may become mobile through rupture of oxygen-cation bonds (e.g., Nesbitt et al., 2017; George et al., 1998) thereby contributing to 74 disorder by forming intrinsic defects (e.g., Frenkel defects). There is overwhelming evidence for 75 76 high temperature alkali and alkaline earth cation disorder in chain and ring silicates (Courtial et al., 77 2000; Thiéblot et al., 1999; Richet et al., 1998, 1996, 1993; George et al., 1998; Dimanov and Jaoul, 1998; Dimanov and Ingrin, 1995). George et al. (1998), for example, conducted a study of 78 Na₂SiO₃ and Li₂SiO₃ by collecting ⁷Li, ²³Na and ²⁹Si NMR spectra as a function of temperature. 79 80 The study indicated limited Li mobility at high temperature. For Na₂SiO₃, motional averaging commenced at ~700-800K and at ~900-1000K, Na⁺ was sufficiently mobile to behave in a liquid-81 like, disordered manner, residing on both structural and non-structural sites (George et al., 1998). 82 Recently, Nesbitt et al. (2017) emphasized that Na^+ and O^{2-} disorder occurred in Na₂SiO₃ crystals 83 84 at temperatures as low as 700-800K, and well before the premelting region of the crystal was 85 encountered (Fig. 1a).

An appreciation for the problem to be addressed is gained by plotting C_p of Na₂SiO₃ against 86 temperature (Fig. 1a). Three sets of C_p data are plotted and they are reasonably consistent 87 (addressed subsequently). An obvious aspect is the rapid increase in C_p within the premelting 88 89 region (e.g., Richet et al., 1994) beginning at ~1200K. A less obvious aspect is the almost linear increase in C_p between ~400K and ~1200K, the slope of which is appreciably steeper than 90 91 expected considering that the rule of Dulong and Petit (Slater, 1963) requires the heat capacity to reach a high temperature limit of 3nR (~149 Jmol⁻¹K⁻¹ for Na₂SiO₃). To emphasize this aspect, C_p 92 93 of Na₂SiO₃ has been normalized to the number of atoms 'n' in the formula unit (n=6 for Na₂SiO₃) to yield a "normalized C_p" (i.e., C_p/n). The normalized experimental data are plotted against 94

temperature in Figure 1b as is the heat capacity (C_v) calculated using the Debye Function where the Debye Temperature (θ) = 750K (Fig. 1b, dashed curve and Table 1). The Debye Function (i.e., C_v) reproduces reasonably well the normalized C_p data in the range 300K to 600K but C_v deviates from experiment between ~700K and 1200K. It may result from anharmonic contributions to C_p but Richet et al. (1996) concluded that C_v and anharmonic contributions alone cannot account for the discrepancy. We address this discrepancy and attempt to explain it.

101

THERMODYNAMICS ASPECTS

Heat capacity includes at least two major contributions between \sim 300K and the premelting region of a crystal. One is the heat capacity at constant volume (C_v or Debye Function) which derives from the 'Debye elastic continuum' (Ghose et al., 1991; Keifer, 1980; Debye, 1912; Born and von Kármán, 1912; Einstein, 1907). Anharmonic contributions constitute the second contribution arising from the nature of site potentials within a crystal, which give rise to thermal expansion. The two contributions allow calculation of C_p according to:

108
$$C_p = C_v + [\alpha_V^2 V/\beta]T$$
 (1)

109 where V is the volume of the phase, α_V is the volume coefficient of thermal expansion, β is the 110 compressibility and T is temperature. With V, α and β of a crystal known, the anharmonic term, 111 ($[\alpha_V^2 V/\beta]T$), can be evaluated and C_p calculated as a function of temperature (provided C_v is 112 known). Equation (1) can be simplified using a relationship elucidated by Grüneisen (1912; 1926):

113
$$\alpha_{\rm V} V/\beta = \gamma C_{\rm v}$$
(2)

114 where γ is the thermal Grüneisen parameter (Slater, 1963, p. 238). Substitution of Equation (2) into 115 Equation (1) and rearrangement yields:

116
$$C_{p} = (1 + \alpha_{V}\gamma T)C_{v}$$
(3)

117 Estimates of the Grüneisen parameter for chain silicates range from 0.88 to 1.3 (Table 2; see also

Angel and Jackson, 2002) with the average value being 1.1 ($n = 7, \sigma = 0.15$). The average is used in subsequent calculations. The γ parameter is largely independent of temperature (Slater, 1963, p. 219) and is similar for all silicates studied here (Angel and Jackson, 2002; Hofmeister and Mao, 2002; Ita and Stixrude, 1992; Jeanloz and Thompson, 1983).

The volume of crystalline Na₂SiO₃ is illustrated in Figure 2a where the solid circles represent volumes from 298 to ~1200K and the open circles represent measurements within the premelting region. A linear least squares best fit to the solid circles was performed and the fit is plotted on the figure. The slope, (dV/dT) is effectively constant at ~1.43x10⁻² (Å³/K). The coefficient of thermal expansion is defined as:

127
$$\alpha_{\rm V} = ({\rm dV/dT})_{\rm p}/{\rm V} \tag{4}$$

where 'p' denotes constant pressure. Because dV/dT is effectively constant, calculation of α_V requires only that it be divided by V at the temperature of interest. The resulting α_V values for Na₂SiO₃ are plotted on Figure 2b and α_V is effectively linear in T. The same procedure was used to obtain α_V for other crystals studied and the results are illustrated in Figures 2c to 2f.

132 Accurate calculation of C_v for silicate crystals is complex (e.g., Ghose et al., 1991; Keiffer, 1982; 1980) but as noted in the Introduction, reasonable estimates of C_v can be obtained using the 133 134 Debye Function or a modification of it. The Debye Function (e.g., Table 1) employs only one 135 adjustable parameter, the Debye Temperature (θ), which can be evaluated from elastic or other 136 properties of the crystal (Grimvall, 2001). In the approach used here, the Debye Temperature is treated as a fit parameter. To obtain approximate C_v values for a crystal, the Debye Temperature is 137 138 adjusted to provide a good fit to the heat capacity in a specified temperature range (here 300-600 K). The same θ value necessarily yields reasonable estimates of C_v at high temperature because the 139 140 Debye Function (C_v) approaches 3nR regardless of the value of θ . Treating θ as a fit parameter has 141 the advantage that the calculation of C_v may accommodate other moderate temperature lattice 142 contributions such as some optical modes (Ghose et al., 1991; Kieffer, 1980). Highly accurate 143 estimates of C_v are not required for this exercise, and the Debye Function is used to calculate C_v .

144

CALCULATED AND EXPERIMENTAL C_P DATA

145 **Orthoenstatite and Proto-enstatite**

Three sets of experimental C_p data are plotted on Figure 3a. The large open and shaded 146 147 circles represent, respectively, Cp of orthoenstatite and proto-enstatite. The values are derived from 148 relative enthalpies (heat contents) reported by Thiéblot et al. (1999). They measured the heat 149 evolved upon cooling a sample from a temperature T to a reference temperature ($T^* = 273$ K). The 150 heat evolved is the *difference* in enthalpy (H) at the two temperatures $(H_T - H_{T*})$. A finite 151 difference approach was used to estimate an average C_p (i.e., $C_p(av)$) over the temperature range of any two measurements. Subtracting two heat content measurements (one at T_2 and one at T_1) 152 yields $(H_{T2} - H_{T1})$ and dividing it by the difference in the two temperatures $(T_2 - T_1)$ gives $C_p(av)$: 153

154
$$C_p(av) = (H_{T2} - H_{T1})/(T_2 - T_1) = \Delta H/\Delta T$$
 (5)

The $C_p(av)$ value represents the average C_p over the interval T_1 to T_2 and may be plotted on Fig. 3 155 156 at $T(av) = (T_2-T_1)/2$. T₁ need not be T* and may be any temperature at which a heat content 157 measurement has been made. The original relative enthalpies of Thiéblot et al. (1999) for 158 orthoenstatite and proto-enstatite are reproduced in Table 3 and were used to calculate $C_p(av)$, 159 which were, in turn, normalized to 'n' to obtain $C_p(av)/n$ (n=5 for MgSiO₃). Heat capacities 160 derived from relative enthalpies commonly are of lower precision than those obtained by other 161 techniques due to subtraction of two large numbers. The resulting values are, however, devoid of 162 assumptions and propagation of errors, and may be more accurate than techniques which include 163 assumptions and propagation of errors. Use of C_p(av) to approximate C_p introduces minimal

uncertainty in C_p above about 700K, as now shown. The average value of T_2 - T_1 of Table 3 is ~100K so that the temperature range associated with each $C_p(av)$ is about ±50K. Uncertainty of this magnitude does not affect the conclusions drawn subsequently.

The small circles of Figure 3a represent experimental $C_p(av)/n$ values for a synthetic orthoenstatite. The large open and shaded circles represent $C_p(av)/n$ data for orthoenstatite and proto-enstatite, respectively. The small and large circles merge between ~450K and 800K, substantiating the accuracy of both sets of data. There is a group of measurements between 850K and 1000K (Fig. 3a, small circles) which is offset to slightly greater C_p values relative to the other data. These 'offset' data are not considered in the following discussion.

173 A Debye Temperature of 875K was used to calculate C_v for both enstatites and C_v is plotted 174 in Figure 3a as the dashed curve. The expression for α_V provided in Figure 2c and the Grüneisen parameter (=1.1) were employed to calculate C_p according to Equation (3). The result is illustrated 175 176 by the solid curve of Figure 3a. The solid curve reproduces (within experimental uncertainty) the 177 orthoenstatite and proto-enstatite experimental values up to ~ 1650 K and the difference between the solid curve and the dashed curve represents the anharmonic contribution to C_p. Clearly there are 178 179 only two major contributions to C_p of the enstatites between ~300K and the premelting region of 180 proto-enstatite. Moreover, the α_V and Grüneisen parameter used in the calculations are reasonable for both enstatites. Finally, the conformity of the calculated and experimental C_p values indicates 181 that the Debye Function provides reasonable estimates of C_v for the enstatites. 182

183 **Diopside**

184 Two sets of experimental C_p data for diopside (CaMgSi₂O₆) are plotted on Figure 3b and 185 they merge smoothly. With a Debye Temperature of 900K, the α_V expression of Figure 2b and a 186 Grüneisen parameter of 1.1, the C_p values for diopside were calculated (Eq. 3) and plotted as the solid curve of Figure 3b. The solid curve conforms closely to the experimental data up to ~1500K, beyond which premelting and Ca disorder commences (Dimanov and Jaoul, 1998; Richet et al., 1998; Dimanov and Ingrin, 1995; Richet et al., 1994; Richet and Fiquet, 1991). As for orthoenstatite and proto-enstatite, there are two major contributions to the C_p of diopside, C_v and an anharmonic contribution. Diopside is structurally different from the enstatite polymorphs but the same Grüneisen parameter (1.1) allows reasonable estimates of C_p for the three crystals.

193 **Pseudowollastonite**

194 The C_p/n (n= 5) values calculated for two sets of experimental data are plotted on Figure 195 3c. The C_v of pseudowollastonite (CaSiO₃) was calculated with θ = 900K and plotted on Figure 3c 196 (dashed curve). The expression for α_V (Fig. 2f) and a Grüneisen parameter = 1.1 were used to 197 calculate the effect of anharmonicity (Eq. 3) and the solid curve represents the result. The essential 198 agreement between experiment and calculation indicates that within the temperature range ~ 300 K 199 to ~1500K, its C_p includes only C_v and anharmonic contributions. The Grüneisen parameter for the 200 pyroxenes and this ring silicate are effectively the same, demonstrating that structural variability 201 has no appreciable effect on the parameter.

202 Li₂SiO₃

The crystal is an orthorhombic chain silicate with Si₂O₆ chains linked by LiO₄ polyhedra (Hesse, 1977). Experimental $C_p(av)/n$ values (n = 6) of Téqui et al. (1992) are plotted on Figure 3d (open circles). The C_v contribution was calculated with θ = 950K (Fig. 3d, dashed curve) and C_p was calculated (Eq. 3) with α_V given in Figure 2e and a Grüneisen parameter = 1.1 (Fig. 3d, solid curve). It conforms to the experimental data between ~500K and ~700K but systematically underestimates C_p at 900K and greater temperatures. There are three likely reasons for the discrepancy; the value for α_V or C_v is incorrect at high temperature, or γ is greater than 1.1. An 210 incorrect $\alpha_{\rm V}$ is unlikely in that it was accurately measured and behaves systematically with 211 temperature (Fig. 2e). That C_v would be incorrect at temperatures greater than ~900K also seems 212 unlikely because all C_v values tend toward the same 3nR limit (i.e., θ has little effect on C_v at high 213 temperature). This leaves the Grüneisen parameter. An extreme value of 1.3 was adopted for the 214 Grüneisen parameter (Table 2) and C_p was recalculated with the results plotted as a dotted curve 215 on Figure 3d. The C_p discrepancy is diminished but nevertheless remains at temperatures greater 216 than ~1000K. Noting that $\gamma = 1.1$ for the previously treated metasilicates, there is no reason to expect a different value for Li_2SiO_3 and we suggest that there is a third, weak contribution to C_p of 217 218 Li₂SiO₃ at high temperature. We return to this aspect subsequently.

219 Na₂SiO₃

220 Na₂SiO₃ is a chain silicate isostructural with Li₂SiO₃. The crystal undergoes a polymorphic 221 transition (from Cmc 2₁ symmetry to Pmc 2₁ symmetry) at ~850K involving Na-O polyhedra but the Q² chains are unaffected (Richet et al., 1996). Heat content data from three sources are plotted 222 on Figure 3e. Naylor (1945) used $T^* = 298.15$ whereas Richet et al. (1984) used $T^* = 273.15$ and 223 their different procedures resulted in a systematic difference in C_p(av). To standardize the data, 224 $C_p(av)$ was calculated for both sets of data and the difference between them was 15 Jmol⁻¹K⁻¹ at 225 226 ~1080K. The Richet et al. (1984) $C_p(av)$ values were increased by this amount making the two sets 227 of data remarkably consistent (Fig. 3e). They are also consistent with those reported by Kelley 228 (1960).

Na₂SiO₃ C_v values were obtained using $\theta = 750$ K (Fig. 3e, dashed curve) and C_p was calculated via Equation (3) using α_V provided in Figure 2b and a Grüneisen parameter = 1.1. (Fig. 3e, solid curve). The calculated and experimental C_p values are consistent between ~300 and ~700K but they diverge beyond ~700K. At ~1200K the C_p discrepancy is ~15% relative to the experimental value. There is no obvious effect on C_p at the polymorphic transition (~850K) as shown in Figures 2b or 3e so that any effect of the transition on C_p is within experimental uncertainty. It is concluded that there is a third contribution to the C_p of Na₂SiO₃ with onset at ~700K and this contribution increases systematically with temperature.

237 K₂SiO₃

238 K₂SiO₃, like pseudowollastonite, consists of three-membered rings (Si₃O₉), with potassium 239 polyhedra linking them to produce a monoclinic unit cell (Werthmann and Hoppe, 1981). The heat 240 contents of Beyer et al. (1979) were employed to calculate $C_p(av)/n$ values and these are plotted on 241 Figure 3f (crosses). The results mimic the experimental data for Na₂SiO₃ (Fig. 3f, circles). The 242 melting temperature of K_2SiO_3 is 1195K and the $C_p(av)/n$ experimental datum point at 1150K is 243 much greater than the value at 1100K, which may indicate a premelting region for K₂SiO₃. A θ_D = 244 750 was employed to calculate C_v of the crystal and it fits reasonably the $C_p(av)/n$ experimental 245 result at ~500K (Fig. 3f) but underestimates $C_p(av)/n$ at greater temperatures. No value for α_V was found for K_2SiO_3 so that C_p could not be calculated and the α_V value of Na_2SiO_3 was used as a 246 247 proxy, yielding the solid curve of Figure 3f. The agreement with the experimental data in the range ~450-700K indicates that the α_V proxy is reasonable but the calculated C_p/n values underestimate 248 249 experimental values at temperatures greater than ~700K. No reasonable value of α_V can reproduce the experimental data at high temperature and apparently there is a third contribution to C_p of 250 251 K₂SiO₃ which is similar in character and magnitude to that of the Na₂SiO₃ C_p discrepancy.

252

EXPLANATION FOR C_p DISCREPANCIES

253 Introduction

Free translation of atoms through crystals is unlikely and electronic contributions to C_p of insulators are similarly unimportant in the temperature range ~300K to 1500K (Slater, 1963; Borg and Dienes, 1992). Substitutional disorder resulting from exchange of cations on different structural sites is also impossible for Na₂SiO₃, K₂SiO₃, Na₂Si₂O₅ or Li₂SiO₃ unless Si sites are involved, which is highly unlikely. The C_p discrepancies observed for these alkali-bearing crystals therefore arise from other contributions. Heat capacity is the sum of numerous internal and configurational contributions (Grimvall, 2001):

261
$$C_p(crystal) = C_p(vibration) + C_p(defect) + C_p(rotation) + C_p(libration)$$
 (6)

262 Lattice vibrations and anharmonic contributions are here accounted for by Equation (3). There are 263 no structural changes to Li₂SiO₃ or K₂SiO₃ from 298K to their premelting regions, hence there is no contribution to their C_p resulting from polymorphism. Na₂SiO₃ undergoes a transition at ~850K 264 265 from orthorhombic (*Cmc* 2_1) to a lower symmetry, possibly *Pmc* 2_1 (Richet et. al., 1996). There is, however, no discontinuity evident in the C_p trend at this temperature (Fig. 3e) so that any resulting 266 267 C_p contribution is within the uncertainty associated with the measurements of C_p. The effects of dislocations on C_p are well known to experimentalists, and care special is taken to avoid samples 268 269 with high dislocation densities through sample selection and treatment (e.g., Richet et al., 1996; 270 Naylor, 1945). This leaves configurational contributions such as point defects and rotation/libration 271 as the most likely causes of the C_p discrepancies.

Whether intrinsic defect formation or rotation/libration contributions are responsible for the C_p discrepancies can be assessed by reference to crystalline Na₂Si₂O₅, a sheet silicate consisting of Q³ species. Naylor (1945) reports Na₂Si₂O₅ heat contents from which $C_p(av)/n$ values were calculated and the results are plotted on Figure 4 (filled circles). The $C_p(av)/n$ data for Na₂SiO₃ are also plotted (Fig. 4, open and shaded circles). The Na₂Si₂O₅ and Na₂SiO₃ experimental data are effectively coincident from ~300K to ~1000K and the Debye Function reproduces well the experimental C_p data for Na₂Si₂O₅ between 300 and 600K with $\theta_D = 750$ (Fig. 4, dashed curve). The anharmonic contribution to C_p of Na₂Si₂O₅ could not be calculated for lack of a α_V value and the Na₂SiO₃ value (Fig. 2b) was used as a proxy to obtain the solid curve of Figure 4. The C_p/n vs. T trend described by the Na₂Si₂O₅ experimental data is too steep to be accounted for by Equation (3) using a reasonable value for α_V . We conclude that Na₂Si₂O₅ displays a C_p discrepancy of the same magnitude as that for Na₂SiO₃.

The structures of the two crystals are different (chain vs sheet) and the Q^2 and Q^3 tetrahedra 284 of the respective crystals should make different rotational or librational contributions to C_p due to 285 their different linkages. There is, nevertheless a remarkable similarity in the C_p of the two crystals 286 287 and in their C_p discrepancies. We conclude that rotation and libration are unlikely to be responsible 288 for the similarities in their C_p discrepancies. Instead, the essentially coincidence of the Na₂Si₂O₅ and Na2SiO3 Cp trends relate to their common cation, and specifically to intrinsic cation defect 289 290 formation in both phases. Although rotation and libration occur in some silicates (Voronko et al., 291 2006; George et al., 1998; Ghose et al., 1986; Massa et al., 1983) defect formation is the cause of 292 C_p discrepancies in these crystals.

293 Cation Mobility, Disorder and Relationship to C_p

294 Li Mobility in Li₂SiO₃. There is experimental evidence for Li mobility in Li₂SiO₃ where 295 subjected to strong thermal agitation. George et al. (1998) conclude from their ⁷Li NMR results 296 that: "Li⁺ site hopping is clearly observed in Li₂SiO₃ by a partial averaging of the ⁷Li quadrupolar 297 peak shape, requiring exchange among a few, ordered orientations of LiO4 tetrahedra". Clearly, Li 298 is not sufficiently mobile, even at 1420K, to have a completely averaged, isotropic environment. 299 With Li-O bond dissociation at high temperature, the majority of mobile Li⁺ alight on previously 300 vacated structural sites and re-establish Li-O bonds, which results in no net change to C_p. The spectroscopic results do not preclude, however, a small fraction of mobile Li⁺ from having 301

alighted, fleetingly, on interstitial sites to form a few intrinsic (Frenkel) defects. The fractionshould increase with temperature as more Li become mobile.

Na Mobility in Na₂SiO₃. George et al. (1998) observe that Na is strongly mobile in Na₂SiO₃ at high temperatures. Quoting George et al. (1998): "From ²³Na NMR results, Na⁺ site hopping in Na₂SiO₃ appears to involve a more liquid-like behavior, implying exchange among many sites with different orientations in a disordered fashion". The disordering probably results from a fraction of mobile Na⁺ alighting on interstitial sites to create intrinsic Frenkel defect pairs, as observed for Ca²⁺ in diopside and pseudowollastonite (see next section). These defects increase the C_p of the crystal in proportion to the number of mobile ions at any instant.

Insight into the high temperature 'liquid-like' behaviour of Na⁺ in Na₂SiO₃ may be gleaned 311 by analogy with Na⁺ behaviour in Na-silicate glass. Although there is some ambiguity in 312 distinguishing different types of cationic sites in the glass, Cormack et al. (2002) nevertheless 313 314 distinguish 'cationic sites' from more energetic 'transitional sites', the distinction being dependent 315 on the residence time of the cation on each type of site. Consider, in an energetic sense, the 316 'cationic sites' of the glass to be analogous to the cationic structural sites of the crystal and the 317 'transitional sites' of the glass to be analogous to interstitial sites of the crystal. Cormack et al. (2002) note that Na⁺ of the glass mostly occupies 'cationic sites', but during transit from one 318 319 cationic site to another, the cation may occupy, fleetingly, 'transitional sites' (i.e., in Fig. 5 of Cormack et al., 2002; the Na⁺ atom in cation site 1 migrates to a transitional site with its 3rd step). 320 321 The potential energy minima of 'transitional sites' almost certainly are shallower than the minima of the 'cationic sites'. Occupation of 'transitional sites' consequently increases C_p of the glass by 322 323 an amount proportional to the number of transitional sites occupied at any instant, just as occupation of interstitial (defect) sites in the crystal increases its C_p in proportion to the number of 324

325 defects present at any instant.

326 Ca Mobility in Diopside and Pseudowollastonite. There is very strong experimental 327 evidence for Ca mobility in diopside and pseudowollastonite at temperatures greater than ~1500K 328 (Bouhfid et al., 2002; Dimanov and Jaul, 1998; Dimanov and Ingrin, 1995). Based on Rutherford 329 Backscattering analysis Dimanov and Ingrin (1995) observed a high temperature intrinsic diffusion 330 regime in diopside which "takes place near the onset of premelting as detected in calorimetric 331 measurements and can be interpreted in terms of enhanced formation of Frenkel defects". The 332 premelting region commences at 1500-1600K. These results were confirmed by Dimanov and Jaul 333 (1998) who also concluded that Ca self-diffusion may involve Ca-Frenkel point defects. The 334 Raman and conductivity experiments of Bouhfid et al. (2002) demonstrated the mobility of Ca in 335 both pseudowollastonite and gehlenite, where mobility of Ca in the two phases occurs at 336 temperatures greater than \sim 1500K. The Ca-Frenkel defects formed will contribute to the C_p of 337 these crystals but only at temperatures above ~1500K (i.e., in their premelting regions).

338 C_p Discrepancies and Cation-O Bond Strengths

339 The mobility of cations and the abundance of intrinsic cation defects increase with 340 temperature because thermal agitation promotes rupture of cation-oxygen bonds. There may be, 341 therefore, a relationship between cation-oxygen bond dissociation energies, cation mobility, defect 342 formation and Cp-discrepancies. K-O and Na-O bond dissociation energies are weak at 239 kJ/mol 343 and 257 kJ/mol respectively whereas the Li-O bond is somewhat stronger at 341 kJ/mol (Speight, 344 2005). The alkaline earth-oxygen bond dissociation energies are stronger still, ranging between 345 394 and 563 kJ/mol. From these values we conclude that dissociation of K-O and Na-O bonds 346 should occur at comparatively low temperatures, Li-O dissociation at somewhat greater 347 temperature and Mg-O and Ca-O bond dissociation should occur at still greater temperatures. The 348 Na and K metasilicate and disilicate C_p discrepancies commence at ~700-800K and their C_p 349 discrepancies are the most pronounced (Figs. 3, 4). The Li metasilicate C_p discrepancy (Fig. 3d) 350 commences at greater temperature (900-1000 K), and it is less pronounced than the discrepancies 351 for Na and K silicates (Figs. 3d 3e, 3f). Finally, no C_p discrepancies are apparent up to 1500K for 352 the alkaline earth metasilicates (Figs. 3a, 3b, 3c). Bond dissociation energies correlate well with the temperatures at which C_p discrepancies are initiated and with the strength of the discrepancies. 353 354 We propose that bond dissociation, resulting from strong thermal agitation, promotes cation mobility, cation (Frenkel) defect formation and production of the C_p discrepancies of the alkali 355 356 metasilicates (Figs. 3 and 4). The same order of bond strength (K₂O<Na₂O<Li₂O<MgO/CaO) is 357 obtained from Coulombic interaction considerations using effective ionic radii (Huheey et al., 358 1993, Chap. 4).

359 Reactions Producing Defects and C_p Discrepancies

360 **Cationic Defects.** We have concluded that Na-NBO bond dissociation is observed in 361 crystalline Na_2SiO_3 above ~700-800K (se also Nesbitt et al., 2017) and the reaction proceeds 362 according to:

363
$$[Na_2-Q^2]^0 \rightarrow [Na_1-Q^2]^{1-} + Na^+$$
 (7)

where $[Na_2-Q^2]^0$ is a Q^2 species with a Na atom bonded to each NBO, and $[Na_1-Q^2]^{1-}$ represents a Q² species where one NBO has no associated Na atom. With NBO-M bond dissociation, the cation (M) vacates its structural site and migrates to a non-structural (interstitial) site. The escape of Na⁺ from a structural site creates a vacancy, and together with the occupation of a non-structural site, creates a *Frenkel* defect pair. Creation of these defects increases the C_p of the crystal (i.e., the C_p discrepancy) and consequently increases the entropy of the crystal (Courtial et al., 2000; Richet et al., 1998; Dimanov and Ingrin, 1995). These contributions are *configurational* in nature, and are a major contributor to the C_p discrepancies observed in Figures 3d, 3e, 3f and 4. The cation defects must be located between the Q² chains thus weakening forces binding the chains together.

Anionic Defects. With formation of a $[Na_1-O^2]^{1-}$ moiety via Reaction (7), the tetrahedron 373 374 becomes 'unpinned' at its NBO⁻ apex (i.e., the apex with no associated Na⁺). Due to thermal agitation it may undergo librations different from, and more severe than, 'pinned' [Na₂-O²]⁰ 375 376 moieties. George et al. (1998) corroborate the effects of thermal agitation, noting that there is 377 "some kind of extensive, librational motion of SiO₄ tetrahedral" in crystalline Na₂SiO₃ at high temperature. As emphasized by Nesbitt et al. (2017), the $[Na_1-O^2]^{1-1}$ moiety produced by Reaction 378 (7) is a nucleophile which may attack Si centers of O^2 species resident on adjacent chains. The 379 reaction produces Q^3 species ($[Na_1-Q^3]^0$) according to: 380

381
$$[Na_1 - Q^2]^{1-} + [Na_2 - Q^2]^0 \rightarrow 2[Na_1 - Q^3]^0 + Na^+ + O^{2-}$$
(8)

where O^{2-} is free oxygen or the oxide ion (Nesbitt et al., 2017, their Rxn 2b). The presence of Q^{3-} 382 383 species in the Raman spectra of Na₂SiO₃ confirms that Reaction (8) occurs at temperatures as low as 774K (Richet et al., 1996). The O²⁻ produced must also form anionic Frenkel pairs because 384 Reaction (8) annihilates a tetrahedral oxygen site through formation of BO. Wherever O^{2-} alights it 385 will occupy an interstitial site. It may reside on sites with weak potential energy minima associated 386 with cationic polyhedra (located between Q^2 chains) or it may react with other moieties of the 387 388 crystal, as discussed in the next paragraph. The abundance of cationic Frenkel defects is at least 389 double the abundance of anionic Frenkel defects (see Eqns. 7 and 8)

The strong nucleophile, O^{2-} , is a highly reactive, strong nucleophile and some O^{2-} may attack Q species to produce Si-pentahedra (SiO₅) such as observed in alkali silicate glasses and melts (e.g., Berezhnoi and Boiko, 2005; Stebbins, 1991). The SiO₅ moiety then becomes the defect. At high temperatures the lifetimes of all Si tetrahedra are short (Farnan and Stebbins, 1990) as will be the lives of the SiO₅ moieties. The activation energy associated with formation of the SiO₅ moiety may inhibit its production at low temperatures. Nevertheless, the anionic defects O^{2-} or SiO₅ necessarily contribute to C_p(config.).

397 Defects in Other Alkali Silicate Crystals. The K-O bond dissociation energy is less than 398 that of Na-O so that reactions equivalent to (7) and (8) likely occur in crystalline K₂SiO₃ (and 399 K₂Si₂O₅). By analogy with Na-NBO bond dissociation in Na₂SiO₃, K-NBO dissociation should 400 occur at ~700-800K with intrinsic cationic point defects forming in the K-silicate crystals, thus 401 explaining the C_p deficiency observed in Figure 3f. The Li-O bond strength is greater than Na-O 402 bond strength so that Li cation point defect formation should be inhibited in Li₂SiO₃ relative to 403 Na₂SiO₃. The prediction holds in that Li₂SiO₃ displays a weaker C_p discrepancy than Na₂SiO₃ and 404 the discrepancy is restricted to higher temperatures, commencing at ~900-1000K (Fig. 3d). There 405 is no evidence that Reaction (2) occurs in Li₂SiO₃ (Richet et al., 1996), hence there is no evidence for appreciable O²⁻ disorder in Li₂SiO₃. Finally, based on Na disorder in Na₂SiO₃, Na-O bond 406 407 dissociation and cation defects formation probably occurs in Na₂Si₂O₅ above ~700K, accounting 408 for the C_p discrepancy observed in Figure 4.

409 Defect Abundances in Na₂SiO₃

410 Creation of Q^3 species via Reaction (2) annihilates Na and O structural sites and the 411 displaced atoms must reside on other sites, here considered 'defect' sites. Production of Q^1 species 412 creates new Na and O structural sites by establishment of additional Si-NBO-Na moieties (Nesbitt 413 et al., 2017, their Reaction 3). The 1348K Raman spectrum of Na₂SiO₃ collected by Richet et al. 414 (1996) was fit by Nesbitt et al. (2017) to obtain ~10% Q^1 species, ~68% Q^2 species and ~22% Q^3 415 species. These Q species abundances were used to calculate the abundance of O^{2-} , yielding ~1.7 416 mol% O^{2-} or X_{O2} = 0.017 (see Sawyer et al., 2015, Appendix A for calculation). Considering that

there are 3 moles of O per unit formula, there exists 0.051 (3x0.017) moles of O²⁻ in the crystal. 417 Reactions (7) and (8) combined indicate twice the moles of 'defect' Na (i.e., Na*) as O²⁻, requiring 418 419 0.102 moles of Na* in the crystal at 1348K. But there are two moles of Na per unit formula so that 420 $X_{Na*} = 0.051$. Richet et al. (1996) also collected and fit Na₂SiO₃ Raman spectra at 1217K and 965K. Q^1 species were absent and Q^3 species were present at ~10% and ~5% respectively (visual 421 estimates), indicating that $X_{Na*} \sim 0.025$ and $X_{O2-} \sim 0.016$ at 1217K, and $X_{Na*} \sim 0.012$ and $X_{Na*} \sim 0.012$ and $X_{Na*} \sim 0.012$ and $X_{Na*} \sim 0.012$ and $X_{Na*} \sim 0.012$ 422 423 ~ 0.008 at 965K. The O³ signal is undetectable in Raman spectra below 774K (Richet et al., 1996). the implication being that Na* and O²⁻ are negligible below this temperature. 424

425 Reaction (7) creates a cationic vacancy and interstitial site (i.e., a Frenkel pair) whereas 426 Reaction (8) annihilates a cationic site by a different mechanism (Nesbitt et al., 2017). Thus the 427 abundance of Frenkel pairs is that produced by Reaction (7), which represents half the X_{Na^*} 428 calculated in the previous paragraph. The values of X_{Na*} therefore were halved to obtain the 429 number of Frenkel pairs produced by Reaction (7) and these are plotted on Figure 5 (solid circles). 430 The equation appropriate to prediction of Frenkel pairs is (Borg and Dienes, 1992):

 $Ln(X_{Na^*}) = -E^*/(2RT)$ (3)

where E* is the energy required to create the vacancy site-interstitial site pair, R is the gas constant 432 433 and T is temperature (K). With $E^* = -84560$ J/mol, the calculated increase in X_{Na^*} as a function of 434 temperature is illustrated by the solid curve (Fig. 5), which agrees reasonably with the 435 experimental values at 965K, 1217K and 1348K (Fig. 5, solid circles). The calculation indicates that X_{Na*} ~0.001 at 700K. The calculation indicates that a C_p discrepancy for Na₂SiO₃ should be 436 437 vanishingly small up to ~700K (Fig. 3e). Beyond 700K Na* concentrations increase rapidly with temperatures, as does the C_p discrepancy of Na₂SiO₃. The calculations are consistent with the 438 Raman spectral results (Richet et al., 1996) and with C_p measurements (Fig. 3e). Raman spectra of 439

440 Li_2SiO_3 reveal no obvious Q³ species up to 1465K (Richet et al., 1996), indicating an absence of 441 O^2 . The very weak C_p discrepancy beginning at ~900-1000K (Fig. 3d) likely is attributed solely to 442 minor Li disorder (e.g., Frenkel pair formation). There are insufficient experimental data to 443 calculate cation defect abundances for the other alkali silicates.

444

IMPLICATIONS FOR MANTLE MINERALS

445 The observations and deductions concerning C_p indicate that alkali-O bond dissociation and 446 defect formation occur in alkali metasilicate and disilicate crystals due primarily to thermal agitation. Alkali-O bond dissociation energies are low compared with alkaline earth-O bond 447 448 dissociation energies, implying that greater thermal agitation is required for the latter bonds to be 449 ruptured. The very high temperatures of the mantle may be sufficient to cause alkaline earth-O bond dissociation and to produce defects in pyroxenes and olivines. Dimanov and Ingrin, (1995) 450 451 Richet et al. (1998), Dimanov and Jaoul (1998) and Bouhifd et al. (2002) documented the 452 enhanced mobility (disorder) of Ca in pseudowollastonite, diopside and gehlenite, where disorder 453 commences at ~1500-1600K in the first two phases and ~1800K for the last. In addition, Mg 454 disordering, hence Mg-O bond dissociation, probably occurs within the premelting region of 455 enstatite and forsterite (Richet et al., 1993; Thiéblot et al., 1999), and likely occurs within the 456 premelting regions of most minerals subjected to mantle temperatures (Richet et al., 1998; 1994; 457 Richet and Figuet, 1991). Where Ca or Mg disordering occurs in pyroxenes, reactions entirely 458 analogous to Reactions (7) and (8) may occur (with Ca or Mg substituted for 2Na) thus promoting 459 formation of intrinsic, Frenkel type, defects in the minerals. Two implications for mantle minerals 460 follow.

461 The onset of premelting of diopside and proto-enstatite occurs at ~1550K and ~1700K
462 respectively, resulting in cation disorder and formation of Frenkel type cationic defects (Thiéblot

463 et al., 1999; Richet et al. 1998; Dimanov and Ingrin 1995). As here emphasized, the defects tend to 464 destabilize the minerals by increasing their C_p, entropy and free energy values. In response, the 465 minerals may change structure or composition to produce a more stable phase. Structural changes 466 driven by defect formation may include creation of 'new' sites which incorporate previously 467 disordered cations. These new structural sites may be accompanied by other changes such as 468 incorporation of Al into octahedral or tetrahedral sites, which should stabilize the minerals at high 469 temperature (Al-O bond dissociation energy is ~512 kJ/mol). Through such changes, the 470 abundance of point defects should decrease, thereby stabilizing pyroxenes against the effects of 471 thermal agitation. The Ca-Tschermak (CaAl₂SiO₆) and NaAlSi₂O₆ components are likely examples 472 of pyroxene components produced through minimization of cationic defects. We suggest that 473 elimination of defects through changes to composition and structure play an important role in 474 stabilizing pyroxenes against extreme thermal agitation.

475 Richet et al. (1993) suggests that Mg undergoes disordering within the premelting region of 476 forsterite (~2050-2163 K) which spans the temperature range of the wadslevite stability field 477 (Frost, 2008; Katsura et al., 2004). Mg-O bond dissociation and production of Frenkel defects 478 should contribute to the instability of olivine and to formation of wadslevite by converting the highly energetic Mg Frenkel defects of olivine to more stable edge-sharing Mg²⁺ octahedral 479 480 structural sites of wadsleyite (Ashbrook et al., 2005). In addition, Mg-O bond dissociation necessarily produces a negatively charged $[O^0]^-$ moiety within the crystal (e.g., Nesbitt et al., 481 482 2017). The moiety is a strong nucleophile and due to thermal agitation it may encounter and react with adjacent Q^0 species to produce a [Si₂O₆] moiety and Q^{2-} according to: 483

$$484 2[Q^0]^- \to 2[Q^1] + O^{2-} (4)$$

485 The reaction is analogous to Reaction (4) and occurs in Mg_2SiO_4 glass and melt in which Q^1

486	species are observed (Nesbitt et al., 2015; Sen and Tangeman, 2008; Voronko et al., 2006). The
487	products of Reaction (4), Q^1 and O^{2-} , are essential constituents of wadsleyite, β -Mg ₂ SiO ₄ ,
488	(Ashbrook et al., 2005) and the reaction should promote its formation at the expense of olivine.
489	The presence of wadsleyite in the mantle may have as much to do with Mg-O bond dissociation
490	(through thermal instability) and production of Q^1 species and O^{2-} in olivine as it has to do with
491	pressure.

492

507

ACKNOWLEDGEMENTS

The authors acknowledge the logistical support provided by their associated Universities. We thank Dr. G.M. Bancroft for numerous fruitful discussions and encouragement. We gratefully acknowledge P. Richet, his colleagues and associates for their remarkably detailed measurements of the high temperature properties of alkali and alkaline earths silicates.

497 **References**

- Anderson, O.L. (1967) Equation for thermal expansivity in planetary interiors. Journal of
 Geophysical Research, 72, 3661–3668.
- Angel, R.J., and Jackson, J.M. (2002) Elasticity and equation of state of orthoenstatite, MgSiO₃.
 American Mineralogist, 87, 558-561.
- Ashbrook, S.E., Berry, A.J., Hibberson, W.O., Steuernagel, S., and Wimperis, S. (2005) High resolution ¹⁷O MAS NMR spectroscopy of forsterite (α-Mg₂SiO₄), wadsleyite (β-Mg₂SiO₄),
 and ringwoodite (γ-Mg₂SiO₄). American Mineralogist, 90, 1861-1870.
- 505 Berezhnoi, G.V., and Boiko, G.G. (2005) Defects and Oxygen Diffusion in Metasilicate Melts,
- 506 Molecular Dynamics Simulation, Glass Physics and Chemistry, 31, 145 -154.
- 508 potassium metasilicate and disilicate, United States Bureau of Mines, Report of Investigations,

Beyer, R.P., Ferrante, M.J., Brown, R.R., and Daut, G.E. (1979) Thermodynamic properties of

509 8410.

- 510 Borg, R.J., and Dienes, G.J. (1992) The Physical Chemistry of Solids. Academic Press Inc.,
 511 Boston, 584 p.
- 512 Born, M., and von Kármán, (1912) Über Schwingungen in Raumgittern. *Phys. Zeits.*, **13**, 297-309.
- 513 Bouhfid, M.A., Gruener, G., Mysen, B.O., and Richet, P. (2002) Premelting and calcium mobility
- 514 in gehlenite (Ca₂Al₂SiO₇) and pseudowollastonite (CaSiO₃). Physics and Chemistry of 515 Minerals, 29, 655-662.
- 516 Chopelas, A., (2000) Thermal expansivity of mantle relevant magnesium silicates derived from
 517 vibrational spectroscopy at high pressure. American Mineralogist, 85, 270-278.
- 518 Cormack, A.N., Du, J., and Zeitler, T.R. (2002) Alkali ion migration mechanisms in silicate
 519 glasses probed by molecular dynamics simulations. Physical Chemistry and Chemical
 520 Physics, 4, 3193-3197.
- 521 Courtial, P., Téqui, C. and Richet, P. (2000) Thermodynamics of diopside, anorthite,
 522 pseudowollastonite, CaMgGeO₄ olivine, and åkermanite up to near the melting point.
 523 Physics and Chemistry of Minerals, 27, 242-250.
- 524 Debye, P. (1912) Zur theorie der spezifischen warmen. Ann. d. Physik., **39**, 789-839.
- Dimanov, A., and Ingrin, J. (1995) Premelting and High-Temperature Diffusion of Ca in Synthetic
 Diopside: An Increase of the Cation Mobility. Physics and Chemistry of Minerals, 22, 437 442.
- 528 Dimanov, A., and Jaoul, O. (1998) Calcium self-diffusion in diopside at high temperature: 529 implications for transport properties. Physics and Chemistry of Minerals, 26, 116-127.
- 530 Einstein, A. (1907) Die plancksche theorie der strahlung und die theorie der spezifischen warmen.
- 531 Ann. d. Physik. 22, 180-190.

- Farnan, I., and Stebbins, J.F. (1990) High-temperature ²⁹Si NMR investigation of solid and molten
 silicates. Journal of the American Chemical Society, 112, 32-39.
- 534 Frost, D. (2008) The upper mantle and transition zone. Elements, 4, 171-176.
- 535 George, A.M., Richet, P., Stebbins, J.F. (1998) Cation dynamics and premelting in lithium 536 metasilicate (Li₂SiO₃) and sodium metasilicate (Na₂SiO₃): A high-temperature NMR study.
- 537 American Mineralogist, 83, 1277-1284.
- Ghose, S., Choudhury, N. Chaplot, S.L., and Rao, K.R. (1991) Phonon Density of States and
 Thermodynamic Properties of Minerals. Chapter 11, in Thermodynamic Data: Systematics
 and Estimation. (S.K. Saxena ed.), p. 283-314. Springer-Verlag, New York, 367 p.
- 541 Ghose, S., Schomaker, V., and McMullan, R.K. (1986) Enstatite, Mg₂Si₂0₆: A neutron diffraction
 542 refinement of the crystal structure and a rigid-body analysis of the thermal vibration.
 543 Zeitschrift für Kristallographie, 176, 159 -175.
- 544 Grimvall, G. (2001) Dependence of Thermodynamic Properties on Atomic Masses and Bonding in
- 545 Solids. Chapter 2, in Solid Solutions in Silicate and Oxide Systems, ed. C.A. Geiger, v3,
- 546 p11-36. Eötvös University Press, Budapest.
- 547 Grüneisen, E. (1912) Theorie des festen Zustandes einatomiger Elelmente. Ann. Phys. **39**, 257-306.
- 548 Grüneisen, E (1926) Zustand des festen Körpers. *Handbuch Phys.*, **1**, 1-52.
- 549 Hesse, K.F. (1977) Refinement of the crystal structure of lithium polysilicate. Acta
 550 Crystallographica B 33:901-902.
- Hofmeister, A.M., and Mao, H-K (2002) Redefinition of the mode Grüneisen parameter for
 polyatomic substances and thermodynamic implications. Proceedings of the National
 Academy of Sciences, 99, 559-564.
- 554 Huheey, J.E., Keiter, E.A., and Keiter, R.L. (1993) Inorganic Chemistry. Harper-Collins College

- 555 Publishers, New York, 964 p.
- Ita, J., and Stixrude, L. (1992) Petrology, Elasticity, and Composition of the Mantle Transition
 Zone. Journal of Geophysical Research, 97, 6849-6866.
- Jackson, J.M., Palko, J.W., Andrault, D., Sinogeikin, S.V., Lakshtanov, D.L., Wang, J., Bass, J.D.,
- and Zha, C-S. (2003) Thermal expansion of natural orthoenstatite to 1473 K. European
 Journal of Mineralogy, 15, 469-473.
- Jeanloz, R., and Thompson, A.B. (1983) Phase transitions and mantle discontinuities. Reviews of
 Geophysics and Space Physics, 21, 51-74.
- 563 Katsura, T., Yamada, H., Nishikawa, O., Song, M., Kubo, A., Shinmei, T., Yokoshi, S., Aizawa,
- 564 Y., Yoshino, T., Walter, M.J., Ito, E., and Funakoshi, K. (2004) Olivine-wadsleyite

transition in the system (Mg,Fe)₂SiO₄. Journal of Geophysical Research, 109, B02209 1-12.

- Kelley, K.K. (1960) Contribution to the data on theoretical metallurgy: XII. High-temperature heat
 content, heat-capacity and entropy data for the elements and inorganic compounds. Bureau
 of Mines Bulletin 584. United States Government Printing Office, Washington, 163 p.
- Kieffer, S.W. (1980) Thermodynamics and Lattice Vibrations of Minerals: Application to Chain
 and Sheet Silicates and Orthosilicates. Reviews of Geophysics and Space Physics. 18, 862886.
- Kieffer SW (1982) Thermodynamics and lattice vibrations of minerals: 5. Applications to phase
 equilibria, isotopic fractionation, and high-pressure thermodynamic properties. Reviews of
 Geophysics and Space Physics 20, 827-849.
- Krupka, K.M., Hemingway, B.S., Robie, R.A., and Kerrick, D.M. (1985) High-temperature heat
 capacities and derived thermodynamic properties of anthophyllite, diopside, dolomite,
 enstatite, bronzite, talc, tremolite, and wollastonite. American Mineralogist, 70, 261-271.

- Massa, N.E., Ullman, F.G., and Hardy, J.R. (1983) Interpretation of anomalies in the Raman
 spectrum of K₂Se0₄ in terms of oxygen sublattice disorder. Physical Review B, 27, 15231540.
- Naylor, B.F. (1945) High-temperature heat contents of sodium metasilicate and sodium disilicate.
 Journal of the American Chemical Society, 67, 466-467.
- Nesbitt, H.W., Bancroft, G.M., Henderson, G.S., Sawyer, R., Secco, R.A., (2015). Direct and
 Indirect Evidence for Free Oxygen (O²⁻) in MO-Silicate Glasses and Melts (M = Mg, Ca,
 Pb). Am. Mineral. 100, 2566-2578.
- Nesbitt, H.W., Bancroft, G.M., Henderson, G.S., Richet, P., and O'Shaughnessy, C. (2017)
 Melting, crystallization and the glass transition: toward a unified description for silicate
 phase transitions. American Mineralogist, 102, 412-420.
- 589 Pandolfo, F., Cámara, F., Domeneghetti, M.C., Alvaro, M., Nestola, F., Karato, S-I., and Amulele,
- 590 G. (2015) Volume thermal expansion along the jadeite–diopside join. Physics and 591 Chemistry of Minerals. 42, 1-14.
- 592 Pitzer, K.S., and Brewer, L. (1961) Thermodynamics (revised after Lewis, G.N., and Randall, M),
 593 McGraw-Hill Inc. New York, 723 p.
- Richet, P., and Fiquet, G. (1991) High-temperature heat capacity and premelting of minerals in the
 system MgO-CaO-Al₂O₃-SiO₂. Journal of Geophysical Research, 96, 445-456.
- Richet, P., Bottinga, Y., and Téqui, C. (1984) Heat capacity of sodium silicate liquids. Journal of
 the American Ceramics Society. 67, C6-C8.
- Richet, P., Ingrin, J., Mysen, B.O., Courtial, P., and Gillet, P. (1994) Premelting effects in
 minerals: an experimental study. Earth and Planetary Science Letters, 121, 589-600.
- 600 Richet, P., Leclerc, F., and Benoist, L. (1993) Melting of forsterite and spinel, with implications

- for the glass transition of Mg_2SiO_4 liquid. Geophysical Research Letters, 20, 1675-1678.
- Richet, P., Mysen, B.O., and Andrault, D. (1996) Melting and premelting of silicates: Raman
 spectroscopy and X-ray diffraction of Li₂SiO₃ and Na₂SiO₃. Physics and Chemistry of
 Minerals. 23, 157-172.
- Richet, P., Mysen, B.O., and Ingrin, J. (1998) High-temperature X-ray diffraction and Raman
 spectroscopy of diopside and pseudowollastonite. Physics and Chemistry of minerals, 25,
 401-414.
- 608 Sawyer, R., Nesbitt, H.W., Bancroft, G.M., Thibault, Y., and Secco, R.A. (2015) Spectroscopic
- studies of oxygen speciation in potassium silicate glasses and melts. Canadian Journal of
 Chemistry, 93, 60–73.
- Saxena, S.K., Chatterjee, N., Fei, Y., and Shen, G. et al., (1993) Thermodynamic Data of Oxides
 and Silicates. Springer-Verlag, Berlin, p. 386-387.
- Sen, S., and Tangeman, J. (2008) Evidence for anomalously large degree of polymerization in
 Mg₂SiO₄ glass and melt. American Mineralogist, 93, 946-949.
- 615 Slater, J.C. (1963) Introduction to Chemical Physics. McGraw-Hill Book Company Inc. (1st
 616 Paperback Edition), New York, 521p.
- Speight, J.G. (2005) Lange's Handbook of Chemistry 70th Anniversary Edition. McGraw Hill
 Standard Handbook, New York, Table 4.11, p. 4.41-4.51.
- 619 Stebbins, J.F. (1991) NMR evidence for five-coordinated silicon in silicate glass at atmospheric
 620 pressure. Nature, 351, 638-639.
- 621 Téqui, C., Grinspan, P., and Richet, P. (1992) Thermodynamic Properties of Alkali Silicates: Heat
- 622 Capacity of Li₂SiO₃ and Lithium-Bearing Melts. Journal of the American Ceramic Society, 75,
- 623 2601-2604.

Thiéblot, L., Téqui, C., and Richet, P. (1999) High-temperature heat capacity of grossular
(Ca₃Al₂Si₃O₁₂), enstatite (MgSiO₃), and titanite (CaTiSiO₅). American Mineralogist, 84, 848855.

- Voronko, Yu. K., Sobol, A.A., and Shukshin, V.E. (2006) Raman Spectra and Structure of Silicon–
 Oxygen Groups in Crystalline, Liquid, and Glassy Mg₂SiO₄. Inorganic Materials, 42, 981-988.
 Werthmann, R., and Hoppe, R. (1981) Über K₂SiO₃ das erste cyclotrisilicat eines alkalimetalls
 sowie Rb₂SiO₃, Cs₂SiO₃, Rb₂GeO₃ und Cs₂GeO₃. Revue de Chimie Minérale, 18, 593-607.
 Yang, H., and Ghose, S. (1994) Thermal expansion, Debye temperature and Grüneisen parameter
 in synthetic (Fe, Mg)SiO₃ orthopyroxenes. Physical Chemistry of Minerals, 20, 575-586.
- Zhao, Y., Schiferl, D., and Shankland, T.J. (1995) High P-T single-crystal x-ray diffraction study
 of thermoelasticity of MgSiO₃ orthopyroxene. Physics and Chemistry of Minerals, 22, 393–
 398.

636 Figure Captions

Fig. 1: (a) Illustrates the heat capacity (at constant pressure) of crystalline and liquid Na₂SiO₃ as a function of temperature. The vertical dotted line represents initiation of the premelting region of the crystal the dotted, curved arrow indicates the trend in C_p within the premelting region. The vertical solid arrow indicates the melting point. (b) Illustrates the normalized heat (C_p/n) of the crystal as a function of temperature (n = no. of atoms in the unit formula). The dashed curve represents C_v (Debye Function) calculated assuming the Debye Temperature = 750K and the dotted arrow indicates the premelting trend.

Fig. 2: Illustrates volumetric properties of some chain and ring silicates as a function of temperature. (a) volume of $Na_2SiO_3(c)$ as a function of temperature. The straight line is the least squares best fit to the solid circles. The arrow indicates onset of the premelting region. (b), (c), (d), 647 (e) and (f) Illustrates the volume coefficient of thermal expansion (α_V) of the indicated crystal as a 648 function of temperature. The solid curves are a least squares best fit to the solid circles. The best fit 649 equation and the source of the data are provided in each diagram.

650 Fig. 3: Illustrates the experimental and calculated 'normalized' C_p results for the indicated crystals 651 (normalized $C_p = C_p/n$ where n = no. of atoms per unit formula). The solid curve in each diagram 652 represents the C_p calculated from Equation (3). The dashed curves represent C_v calculated using 653 the Debye Temperature indicated in the diagram. The short arrow (labelled m.p. in some diagrams) 654 indicates the melting point. The dotted curves indicate the C_p trends within the premelting regions 655 of the crystals. See text for details. (a) C_p/n values for orthoenstatite (small and large circles) and 656 proto-enstatite (shaded circles); (b) C_p/n values for diopside.; (c) C_p/n values for 657 pseudowollastonite (p-CaSiO₃); (d) C_p/n for Li₂SiO₃(c). The dotted curve represents C_p/n values 658 calculated assuming that the Grüneisen parameter = 1.3. (e) C_p/n of $Na_2SiO_3(c)$; (f) C_p/n for 659 K₂SiO₃(c) is indicated by crosses and the C_p/n data for Na₂SiO₃ are indicated by open and shaded 660 circles.

Fig. 4: Illustrates the experimental C_p/n values for Na₂Si₂O₅ (solid circles). The experimental C_p/n values for Na₂SiO₃ are plotted for comparison (open and shaded circles). The dashed curve represents the Debye Function and the two dotted arrows indicate C_p/n trends within the premelting regions of Na₂Si₂O₅ and Na₂SiO₃.

Fig. 5: Defect concentrations of Na (X_{Na*}) determined from experiment (dots) and the calculated
Frenkel pair defect concentrations (solid curve) using the activation energy shown in the diagram.
Details are provided in the text.









Fig. 4



Fig. 5

Table 1: The Debye Function ¹				
θ/Τ	θ/T T (K)			
	750	J/mol ⁻¹ T ⁻¹		
0.10	0.0	24.89		
0.20	0.0	24.85		
0.30	0.0	24.81		
0.40	0.0	24.73		
0.50	0.0	24.60		
0.60	0.0	24.48		
0.70	0.0	24.31		
0.80	0.0	24.14		
0.90	0.0	23.93		
1.00	0.0	23.72		
1.50	0.0	22.34		
2.00	0.0	20.59		
2.50	0.0	18.62		
3.00	0.0	16.53		
3.50	0.0	14.48		
4.00	0.0	12.55		
5.00	0.0	9.20		
6.00	0.0	6.61		

1 Pitzer and Brewer (1961), Table A5-1, p. 660.

Phase	Composition γ Error		References			
Orthopyroxene	(Mg,Fe)SiO ₃	1.1	±0.1	Jeanloz and Thompson, 1983		
Orthopyroxene	FeSiO ₃	1.1	?	Anderson, 1967		
Orthopyroxene	(Mg,Fe)SiO ₃	0.87	±0.04	Yang and Ghose, 1994		
Orthopyroxene	(Mg,Fe)SiO ₃	1.28	±0.05	Chopelas, 2000		
Orthopyroxene	(Mg,Fe)SiO ₃	1.05	?	Zhao et al., 1995		
Orthopyroxene	MgSiO ₃	1.3	?	Hofmeister and Mao, 2002		
Clinopyroxene	Ca(MgFe)Si ₂ O ₆	1.1	±0.05	Jeanloz and Thompson, 1983		
Average:		1.1				
Standard Dev.		0.15				

Table 2: Gruneisen Parameter (γ) for Chain Silicates

I able J.	LIIStati	ie Op(av)	values	Calculated			linaipies
Exptl.	T_2	$H_{T2}-H_{T1}^{2}$	T ₁	$\Delta T = (T_2 - T_1)$	T(av) ³	Cp(av) ⁴	Cp(av)/n ⁵
Run	(K)	kJ/mol	(K)	(K)	(K)	Jmol ⁻¹ T ⁻¹	Jmol ⁻¹ T ⁻¹
Orthoenstatite							
	273.15	0.0					
EE12-B	415.6	12.64	273.15	142.45	344.38	88.73	17.75
EE1-B	513.7	22.87	415.6	98.1	464.65	104.28	20.86
EE11-B	514.1	22.73	415.6	98.5	464.85	102.44	20.49
EE4-B	638.5	36.54	514.1	124.4	576.30	111.01	22.20
EE2-B	696.6	43.3	638.5	58.1	667.55	116.35	23.27
EE20-B	724.1	46.75	638.5	85.6	681.30	119.28	23.86
EE5-B	771.7	52.16	696.6	75.1	734.15	117.98	23.60
EE13-B	849.6	61.7	771.7	77.9	810.65	122.46	24.49
EE6-B	907.9	68.89	849.6	58.3	878.75	123.33	24.67
EE8-B	968.4	76.51	907.9	60.5	938.15	125.95	25.19
EE9-B	1051.8	86.96	968.4	83.4	1010.10	125.30	25.06
EE10-B	1168.6	101.85	1051.8	116.8	1110.20	127.48	25.50
EE17	1225.4	109.51	1051.8	173.6	1138.60	129.90	25.98
Proto-en	statite						
EX17	1366	129.17					
EX11	1427	136.84	1366	61	1396.70	126.15	25.23
EX4	1481	143.47	1366	115	1423.65	124.67	24.93
EX2	1575	155.98	1481	94	1527.80	133.65	26.73
EX3	1683	169.66	1575	109	1628.85	126.08	25.22
EX26	1740	177.16	1683	57	1711.50	132.04	26.41
EX16	1745	177.83	1683	62	1714.00	132.20	26.44
EX5	1773	182.56	1683	90	1728.15	143.17	28.63
EX15	1787	184.77	1745	47	1763.60	160.55	32.11
EX10	1811	191.86	1745	67	1778.15	210.98	42.20

Table 3: Enstatite Cp(av) Values Calculated from Relative Enthalpies¹

1 Relative enthalpies (heat content) are from Thieblot et al. (1999)

2 Values may have large errors where ΔT values are less than 70 K.

3 $\Delta H = (H_{T2} - H_{T1}); \Delta T = (T_2 - T_1); T(av) = (T_2 + T_1)/2$

4 Cp(av) = $(H_{T2}-H_{T1})/(T_2-T_1) = \Delta H/\Delta T$. Cp(av) is plotted at T(av).

5 Cp/n = Cp normalized to 'n', the number of atoms in the unit formula