

1 **Revision 2**

2 **Representative size distributions of framboidal, euhedral and sunflower pyrite**
3 **from high-resolution X-ray tomography and scanning electron microscopy**
4 **analyses**

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16 **Abstract**

17 A statistical procedure designed to obtain representative size distributions for different
18 morphologies and arrangements of pyrite is described here. This statistical procedure is
19 applied to data acquired during scanning electron microscopy (SEM) and high-
20 resolution X-ray tomography (micro-CT) analyses. The statistical procedure was tested
21 in methane-derived carbonate pipes recovered in the Gulf of Cadiz. These samples
22 contain abundant pyrite together with pseudomorph iron oxyhydroxides showing
23 multiple morphologies including euhedral crystals, framboids and sunflowers
24 (framboidal core with outer crystals).

25 The SEM analysis consisted in the establishment of independent populations of pyrite
26 and iron oxyhydroxides grouped by morphology and arrangement and the determination
27 of its size distributions. Micro-CT analysis included a determination of the 3D volume
28 of pyrite from the density difference between pyrite and the rest of mineral forming the
29 samples. The use of the micro-CT technique implies that minerals with similar
30 attenuation coefficients than those of pyrite are scarce or not present in the studied
31 samples. A filtering process was applied to the 3D volume. This filtering process
32 consisted of the selection of objects with corrected sphericity greater than 0.80, discrete
33 compactness greater than 0.60, elongation and flatness of the circumscribed 3D
34 ellipsoid less than 1.80 and the sum of the elongation and flatness less than 3. Objects
35 with shapes similar to those expected in pyrite (spheroidal and regular shapes) were
36 selected with this filtering process. The optimal mixture of lognormal size distributions
37 was obtained applying statistical techniques to the entire size distribution represented by
38 the filtered objects. The correspondence between size distributions obtained during the
39 SEM and the micro-CT analyses was done by matching statistical parameters and using
40 3D renderings. The representative size distributions of pyrite as determined by the
41 proposed 3D processing methodology can be used to accurately quantify the paleo-
42 environmental conditions of pyrite formation, which would solve some of the
43 limitations resulting from analyses based on 2D images.

44 **Keywords:** pyrite, size distributions, framboids, euhedrae, sunflowers, micro-CT

45

Introduction

46 Framboids are defined as microscopic, spheroidal clusters of densely packed
47 homogeneous microcrystals (Ohfuji and Rickard 2005). The most common
48 mineralogical components of framboids are iron sulfides, such as pyrite. Under
49 sedimentary conditions they are formed in the vicinity of the sediment-water interface

50 (Wilkin et al. 1996). They are also very frequent in pyrite formed within methane-
51 derived carbonates in marine cold-seeps (Peckmann and Thiel 2004). In these
52 environments, framboidal pyrite is formed by the reaction between H₂S, produced by
53 the activity of sulfate reduction bacteria (SRB), and detrital Fe minerals. In the SRB-
54 mediated environments, euhedral crystals and other morphologies of pyrite usually
55 occur in close conjunction with framboidal pyrite (e.g. Merinero et al. 2009; Cavalazzi
56 et al. 2012; Wang et al. 2013; Gallego-Torres et al. 2015). The coexistence of
57 framboidal and euhedral pyrite has been attributed to the early formation of framboidal
58 pyrite under conditions of FeS₂ supersaturation and the subsequent decrease in FeS₂
59 saturation by the exhaustion of sources of highly reactive Fe or the decrease in SRB
60 activity (Cavalazzi et al. 2012). This decrease in FeS₂ saturation promotes the growth of
61 microcrystals to form euhedral crystals instead of framboidal pyrite (Butler and Rickard
62 2000) or the overgrowth and recrystallization of framboidal pyrite that evolves into
63 euhedral pyrite (Sawlowicz 1993). The existence of other morphologies of pyrite such
64 as sunflower pyrite (framboidal core with outer crystals) has been explained as the
65 results of the evolution from framboidal to euhedral pyrite (e.g. Merinero et al. 2008)
66 although the role of sunflowers in the whole process has not been detailed. Few works
67 refer to pyrite with internal framboidal texture and external regrowth as sunflower pyrite
68 (e.g. Large et al. 1999). Wilkin et al. (1996) described the infilling of interstitial voids in
69 framboidal pyrite leading to the “welded sphere” morphology previous to the formation
70 of the external overgrowth. Authors also use the terms “framboidal pyrite with
71 overgrowths” (Ding et al. 2014) or “overgrowth crystals upon framboids” (Wei et al.
72 2012). In all cases, this morphology represented the continuation of the growth of
73 framboids without incorporating new microcrystals into the framboidal texture but

74 instead with the overgrown of prismatic pyrite under conditions other than
75 supersaturation (Suits and Wilkin 1998).

76 The diameter of authigenic framboidal pyrite ranges generally from 5 to 20 μm , and the
77 size of the component microcrystals is usually less than 2 μm , although framboids up to
78 250 μm have been reported (Ohfuji and Rickard 2005). Based on the work of Wilkin et
79 al. (1996), the size distribution of small framboidal pyrite has been used to infer the
80 palaeoconditions (e.g., Wignall et al. 2005; Bond and Wignall 2010; Liao et al. 2010;
81 Wang et al. 2013; Guan et al. 2014; Tian et al. 2014; Takahashi et al. 2015; Wei et al.
82 2015). Moreover, the size distribution of framboidal and euhedral pyrite can be used to
83 establish the history and conditions of the nucleation and growth of pyrite (e.g. Pósfai et
84 al. 2001; Merinero et al. 2009). Up to now, the size distribution of microscopic pyrite
85 has been determined from direct observation of thin sections under optical microscopy
86 (e.g. Ding et al. 2014; Blood and Lash 2015) and from 2D photomicrographs obtained
87 through SEM (e.g. Wei et al. 2012, 2015; Guan et al. 2014). These techniques have
88 some limitations, such as the lack of representativeness, the underestimation of the
89 small sizes and the imprecision of measurement in random cuts through the microscopic
90 pyrite, resulting in size distributions with possibly significant errors. One still
91 outstanding issue to follow up on is whether the smaller framboids are actually
92 underestimated, in terms of their number in population distributions. Recently,
93 representative size distributions of microscopic pyrite occurring in shales have been
94 calculated from 3D representations obtained using micro-CT (Cárdenes et al. 2016).

95 The goal of this paper is to develop a statistical procedure that obtains representative
96 size distributions of pyrite grouped by morphology and arrangement based on SEM and
97 micro-CT analysis of geologic samples. This statistical procedure was tested here in
98 methane-derived carbonate pipes that contain abundant pyrite and pseudomorph iron

99 oxyhydroxides. Pyrite and iron oxyhydroxides show different morphologies and
100 arrangement including framboids, euhedrae and sunflowers.

101 **Experimental and analytical methods**

102 **Samples**

103 The carbonate pipes studied here were collected on the Iberian continental margin of the
104 Gulf of Cadiz during the oceanographic cruises Anastasya 2001, ARSA 0309 and
105 INDEMARES/CHICA 0211, aboard the research vessel Cornide de Saavedra. More
106 than 400 carbonate pipes were recovered from the seafloor during these cruises at
107 depths ranging from 700 to 1100 m using rectangular, benthic-type dredges (Díaz-del-
108 Río et al. 2003, Rueda et al. 2012; Vila et al. 2013). Mineralogy and textures of the
109 carbonate pipes were previously studied showing the presence of pyrite and several
110 unidentified iron oxyhydroxide phases (Díaz-del-Río et al. 2003, Merinero et al. 2008).

111 Four samples were selected based on the morphologies of pyrite and iron
112 oxyhydroxides. For two of them, the outer and inner parts of the sample were studied
113 separately (Fig. 1). The nomenclature of the samples used in this study is summarized in
114 Table 1.

115 **Scanning Electron Microscopy (SEM) analysis**

116 Thin sections (transversal and longitudinal views) of the carbonate pipes were studied in
117 detail under optical microscopy and SEM. The SEM analysis determined the size
118 distribution of pyrite and iron oxyhydroxides grouped by morphology and arrangement.
119 Observations and photomicrographs were performed in back scattered imaging mode
120 with a JEOL Superprobe JXA-8900M at the National Center for Electron Microscopy
121 (Complutense University of Madrid) with a voltage of 20 kV and a current intensity of
122 50 nA. Sizes were calculated from photomicrographs obtained with the JEOL

123 Superprobe following the method used by Merinero et al. (2009) and using ImageJ
124 1.50b software (Schneider et al. 2012).

125 **High resolution X-ray computed tomography (micro-CT) analysis**

126 Micro-CT is a 3D visualization and measurement technique in which the object being
127 studied is represented as a discretized 3D volume of the local linear X-ray attenuation
128 coefficient. From this 3D volume it is possible to estimate the size distribution of
129 minerals with X-ray linear attenuation coefficients different than those of their matrix
130 and of surrounding minerals (Cnudde and Boone 2013). Here we use this technique to
131 estimate the entire size distribution of the pyrite and iron oxyhydroxides.

132 We obtained cylindrical samples 4 mm in diameter and 10 mm in length from each
133 sample for analysis with micro-CT. These cylindrical samples were scanned with the
134 HECTOR setup (Masschaele et al. 2013), located at the Center for X-ray Tomography
135 at Ghent University (www.ugct.ugent.be). A high voltage of 130 kV and a tube power
136 of 10 W were used for the measurements, with a 1 mm aluminum filter in order to
137 reduce the beam-hardening effect (Cnudde and Boone 2013). More than 1800
138 projections of each sample were obtained during a complete rotation along a vertical
139 axis, achieving a voxel size of 2.5 μm . The projections obtained by micro-CT were
140 reconstructed using Octopus software (Vlassenbroeck et al. 2007). The reconstructed
141 images were analyzed with ImageJ 1.50b software (Schneider et al. 2012). Firstly,
142 reconstructed images were converted from 16 to 8 bits per pixel and filtered using the
143 median filter function of the ImageJ plugin 3D Fast Filters (Ollion et al. 2013).
144 Subsequently, individual mineral objects representing pyrite or iron oxyhydroxides
145 were 3D-segmented according to their grey-scale values using the ImageJ plugin
146 RoiManager3D 2.0 (Ollion et al. 2013). For each object, we measured the following
147 parameters: volume (μm^3), surface (μm^2) and 3D ellipsoid circumscribed to the object

148 (major, intermediate and minor radii in μm). The longest diameter of the 3D ellipsoid
149 was considered as the size of the analyzed object. For each object the following
150 parameters were estimated: classical sphericity ψ (Wadell 1935), corrected sphericity ψ_c
151 (Lindblad 2005) and discrete compactness C_d (Bribiesca 2008):

$$152 \quad \Psi = \frac{\pi^{\frac{1}{3}}(6V_p)^{\frac{2}{3}}}{A_p}; \Psi_C = \frac{\pi^{\frac{1}{3}}(6V_p)^{\frac{2}{3}}}{A_C}; C_d = \frac{V_p - \frac{A_p}{6}}{V_p - (\sqrt[3]{V_p})^2}$$

153 Where V_p and A_p are the measured volume and surface of the object and A_c is the
154 corrected surface estimated from the algorithm developed by Lindblad (2005).
155 Elongation and flatness of the circumscribed 3D ellipsoid were calculated from the ratio
156 between the major and the intermediate radii and the intermediate and minor radii
157 respectively. Finally, 3D renderings of individual objects were performed using the
158 plugin RoiManager3D 2.0 (Ollion et al. 2013).

159 A filtering process was applied to the segmented 3D objects to select exclusively
160 regular and spheroidal objects. The filtering process was divided into 4 steps (Table 2).
161 The first step consisted of the adjustment to the theoretical lognormal distribution in the
162 smaller sizes. During steps 2, 3 and 4, objects with different shapes than those expected
163 for pyrite and iron oxyhydroxides (framboidal, euhedral or sunflower) were removed on
164 the basis of to their corrected sphericity, discrete compactness, elongation and flatness
165 of the circumscribed 3D ellipsoid (Table 2).

166 **Statistical analysis**

167 Data management and analysis were performed using the statistical software R v. 3.2.3
168 (R_Core_Team 2015). Lognormality of the size populations obtained from the SEM
169 analysis was assessed by quantile-quantile plots and applying the Shapiro-Wilk test of
170 normality (Royston 1995) to the natural logarithm of the size. Mixture of lognormal
171 populations was analyzed from the entire size distribution obtained from the micro-CT

172 analysis and the four-step filtering process. The Mclust function of the R package
173 mclust v. 4.4 (Fraley and Raftery 2002; Fraley et al. 2012) was used to estimate the
174 optimum mixture modelling of lognormal populations fitted via an expectation-
175 maximization algorithm including Bayesian regularization and dimension reduction.
176 The Mclust function gives the parameters of the independent lognormal populations
177 (mean, standard deviation and proportion in the entire size distribution) together with
178 the probability that each individual object belongs to one or another population. The
179 percentage in volume was obtained multiplying this probability by the individual
180 volume of each object. The minimum and maximum of the independent lognormal
181 populations were estimated from the size of those objects with probability of belonging
182 to that population greater than 0.10.

183 **Results**

184 Carbonate pipes consist of microcrystalline carbonate (Fe-rich dolomite, ankerite and
185 Mg-rich calcite) containing well-preserved detrital grains of quartz and foraminiferal
186 shells and minor phyllosilicates, k-feldspars, apatite and Fe-Ti-oxides. Pyrite and
187 pseudomorph iron oxyhydroxides are commonly found in the void spaces resulting from
188 primary and secondary porosity. The main morphologies of pyrite and iron
189 oxyhydroxides are framboids, euhedrae and sunflowers (Fig. 2). In the Cornide sample,
190 pyrite is totally transformed into iron oxyhydroxides. In the Hespérides, Almazán outer
191 and inner samples, pyrite is partially transformed into iron oxyhydroxides. In the
192 DIASOM outer and inner samples, pyrite is preserved and transformation into iron
193 oxyhydroxides is not observed.

194 **SEM analysis**

195 The focus of the SEM analysis of the carbonate pipes was on the morphology and
196 arrangement of the iron minerals, which allowed us to establish independent

197 populations based on these features. Groups of specimens with spheroidal shape were
198 also considered as independent populations. Size distribution was calculated for each
199 population and represented in histograms.

200 In the Cornide sample, iron oxyhydroxides are disseminated in the matrix and
201 associated with the infilling of small fractures (Fig. 3a). Disseminated iron
202 oxyhydroxides show euhedral morphologies (Figs. 3a and 3b). In the fractures, iron
203 oxyhydroxides have sunflower morphologies (Figs. 3a and 3c). The larger objects
204 preserve the features of the sunflower morphology (Fig. 3c). The smaller objects show
205 euhedral shapes, preserving some internal features of the sunflower morphology and
206 showing dark inclusions (Fig. 3c). Groups of several objects with sunflower
207 morphologies are common and some of them have regular shapes (Fig. 3b).

208 In the Hespérides sample, pyrite and iron oxyhydroxides occur in association with the
209 infilling of small fractures, disseminated in the matrix and inside foraminiferal shells
210 (Fig. 3d). Pyrite and iron oxyhydroxides with euhedral, framboidal and sunflower
211 morphologies are observed. Two types of euhedral morphologies can be distinguished:
212 disseminated in the matrix (euhedral-1, Figs. 3d and 3e) and in close proximity to
213 framboids (euhedral-2, Figs. 3d and 3f). Disseminated euhedral crystals are of smaller
214 size than those in close proximity to framboids. Framboidal morphologies are observed
215 inside foraminiferal shells (framboidal-1, Fig. 3d) and fractures (framboidal-2, Fig. 3f),
216 although infillings of foraminiferal shells are very scarce. Sunflower morphologies as
217 individual or groups of specimens only occur in association with fractures (Fig. 3e).

218 The Almazán outer sample is characterized by the presence of individual euhedral
219 crystals of pyrite disseminated in the matrix (euhedral-1, Figs. 3g and 3h) and groups of
220 euhedral crystals infilling foraminiferal shells and small holes in the matrix (euhedral-2,
221 Figs. 3h and 3i). Disseminated euhedral crystals are of smaller size than those forming

222 groups. Individual framboids associated with the infilling of small fractures
223 (framboidal-1, Figs. 3g and 3h) and groups of framboids formed inside foraminiferal
224 shells (framboidal-2, Figs. 3g and 3i) are also common in this sample.

225 The Almazán inner sample is characterized by the existence of numerous holes and
226 unfilled fractures and the absence of foraminiferal shells. Iron minerals are observed
227 around these holes and present framboidal and euhedral morphologies (Figs. 4a, 4b and
228 4c). The formation of groups of framboids is very common in this sample (Fig. 4c).

229 In the DIASOM outer sample, pyrite is associated with the infilling of fractures and
230 foraminiferal shells (Fig. 4d). Pyrite associated with the infilling of fractures shows both
231 framboidal (Fig. 4e) and euhedral morphology (euhedral-1, Figs. 4d and 4f). Pyrite
232 inside foraminiferal shells shows euhedral morphology (euhedral-2, Fig. 4f). In this
233 sample, groups of framboids and foraminiferal shells with the chambers completely
234 infilled are very common and therefore considered as an individual population of
235 spheroidal shape (Figs. 4d and 4e).

236 The DIASOM inner sample is characterized by an abundance of pyrite infilling
237 fractures and forming concretions of multiple specimens (Fig. 4g). Sunflower pyrite is
238 the main morphology in this sample in close proximity to euhedral crystals with internal
239 sunflower features (Figs. 4h and 4i). Foraminiferal shells are also abundant and are
240 completely infilled with concretions of pyrite with internal framboidal features (Fig. 4i).
241 Individual euhedral crystals are also present disseminated in the matrix (Fig. 4g).

242 The mean, standard deviation, minimum and maximum of the size of each individual
243 population considered in the SEM analysis are included in Table 3, and histograms are
244 plotted in the Fig. 5. Most of the histograms are mono-modal and have a lognormal
245 shape (Fig. 5). The theoretical lognormal distribution fits on the mode in each
246 population (Fig. 5). The quantile-quantile plots and the Shapiro-Wilk test applied to the

247 natural logarithm of the size are also included in Fig. 5. Deviations from lognormality
248 are observed in some populations of the smaller or larger sizes, although most of the
249 populations can be considered lognormal.

250 **Micro-CT and statistical analysis**

251 The entire size distribution containing the mixture of independent populations of pyrite
252 and iron oxyhydroxides was obtained during the micro-CT analysis and the four-step
253 filtering process. The total number, mean, standard deviation, minimum and maximum
254 sizes of the objects discarded in each step of the filtering process are summarized in
255 Table 4. The histograms of the entire size distribution for each sample are represented in
256 Fig. 6. The presence of multiple peaks in the histograms corresponding to multiple
257 modes is the most significant feature. These peaks are more appreciable for the smaller
258 sizes.

259 The optimal number of independent lognormal subpopulations forming the entire size
260 distribution was obtained using the Mclust function. The mean, standard deviation,
261 minimum, maximum, percentages in volume and number of objects for each
262 independent lognormal subpopulation are summarized in Table 3. The probability
263 density functions of each lognormal subpopulation obtained with Mclust are represented
264 superposed over the histograms of the entire size distribution (Fig. 6). The modes of
265 most of these lognormal subpopulations fit on the peaks obtained in the histograms (Fig.
266 6).

267 **Discussion**

268 **Filtering process of the 3D objects and separation into independent subpopulations**

269 Although micro-CT analysis provides statistically representative size distributions of
270 minerals with different attenuation coefficients than the rest, there are several

271 limitations that should be taken into account when it is applied to samples with
272 abundant pyrite, and that the statistical procedure described here solves or minimize its
273 impact:

- 274 • The presence of minerals with attenuation coefficients similar to those of pyrite
275 or iron oxyhydroxides
- 276 • The presence of noise, affecting mostly to objects of smaller size
- 277 • The existence of groups of pyrite and iron oxyhydroxides not representing
278 individual framboids, euhedrae or sunflowers
- 279 • The mixture of independent populations of pyrite and iron oxyhydroxides in the
280 entire size distribution

281 **Selection based on attenuation coefficients.** The calculated linear attenuation
282 coefficients for pyrite and iron oxyhydroxides are very different from those of the other
283 main minerals in carbonate pipes, such as Fe-rich dolomite, ankerite, Mg-rich calcite
284 and quartz (Fig. 7). This allows for easy differentiation of pyrite and iron oxyhydroxides
285 from the carbonate matrix and clasts contained in the carbonate matrix. The linear
286 attenuation coefficients for most of the accessory minerals—such as phyllosilicates
287 (biotite-muscovite), k-feldspars and apatite—are also very different from those of pyrite
288 and iron oxyhydroxides (Fig. 7). However, the linear attenuation coefficient of the Fe-
289 Ti-oxides (also present in the carbonate matrix as clasts, though in less abundance) is
290 similar to those of pyrite and iron oxyhydroxides. The irregular shape of the Fe-Ti-
291 oxides is different from the spheroidal or regular shape of pyrite and iron
292 oxyhydroxides. Because of the non-spheroidal shape of Fe-Ti-oxides, most of them
293 were discarded during the four-step filtering process. Given the high number of objects
294 obtained representing pyrite and iron oxyhydroxides (15000 objects on average), the
295 inclusion of a few objects representing spheroidal Fe-Ti-oxides in the entire size

296 distribution does not alter the statistical representativeness of the results obtained using
297 micro-CT.

298 **Noise filtering.** Although the application of appropriate filtering to the micro-
299 CT reconstructed images may reduce noise (Brabant et al. 2011), some will always
300 persist. It is known that a significant number of smaller-size objects found in a sample
301 can be noise (Cnudde and Boone 2013). Assuming that the size distributions of pyrite
302 and iron oxyhydroxides were lognormal or quasi-lognormal, during the first step of the
303 filtering process the entire size distribution was adjusted to the theoretical lognormal
304 distribution at the smaller range of the size distribution, where noise is more of an issue.

305 **Groups and selection of spheroidal objects.** Micro-CT is effective to obtain
306 significant size distributions when pyrite occurs as individual specimens (Cárdenes et
307 al. 2016). However, when pyrite also forms groups it is impossible to separate
308 individual objects from groups without additional manipulation of the images (Higgins
309 2000). In the carbonate pipes that we studied, pyrite usually forms multiframeboidal
310 textures and polyframeboids (Merinero et al. 2008, 2009). Therefore, it is necessary to
311 select only individual objects to obtain the real size distribution. Here, the selection of
312 individual objects was done according to the sphericity, compactness, elongation and
313 flatness of the objects, discarding objects with shapes to different than those expected
314 (Table 2).

315 Classical sphericity is the most important shape factor for defining the roundness of an
316 object. It varies from 0 to 1, where the maximum correspond to a perfect sphere and
317 ranges from 0.8 to 1 for regular polyhedrons, such as cubes (0.806), regular octahedrons
318 (0.846), dodecahedrons (0.910) or icosahedrons (0.939) the most common forms that
319 pyrite takes. Therefore, objects representing frameboids (spheroidal packing of
320 microcrystals), euhedrae (isometric shapes) and sunflowers (frameboidal core with outer

321 crystals) should have a sphericity of greater than 0.8. When minerals are represented by
322 discrete objects, the problem of classical sphericity is to estimate the surface area
323 accurately. Classical sphericity uses the number of foreground voxels as surface area,
324 but this measure is only effective for large objects (Lindblad 2005). Corrected sphericity
325 uses the corrected surface based on the surface area estimation developed by Lindblad
326 (2005) for discrete 3D objects and reduces the error for sphericity estimated in small
327 objects. Therefore, during the second step of the filtering process (Table 2), objects with
328 corrected sphericity less than 0.80 were discarded, which corresponded to large objects
329 with mean size ranging from 48.95 to 103.6 μm (Table 4). With this criterion, large
330 irregular groups of pyrite such as polyframboids were removed from the sample.

331 Corrected sphericity is sensitive to rotations and translations, but invariant with respect
332 to the existence of holes in the object (Montero and Bribiesca 2009). An alternative
333 shape descriptor for roundness is discrete compactness, an intrinsic characteristic of the
334 shape of objects defined as the ratio between volume and area. Discrete compactness is
335 invariant with respect to rotations and translations, but sensitive to the existence of
336 holes; and objects with the same discrete compactness have similar shapes (Bribiesca
337 2008). Cubes and spheres have the greatest discrete compactness (1 and 0.99,
338 respectively), and regular octahedrons and dodecahedrons usually have a discrete
339 compactness greater than 0.60 (Bribiesca 2000). Therefore, objects with discrete
340 compactness greater than 0.60 were selected during the third step (Table 2), including
341 the expected shapes of pyrite. The mean size of the objects discarded in this step varied
342 from 8 to 11.50 μm (Table 4), an indication that the objects discarded in this step tended
343 to be of smaller size, where corrected sphericity is a shape descriptor of limited
344 usefulness.

345 Despite of the selection of the second and third steps (Table 2), some elongated and
346 flattened objects with spheroidal shape remained in the sample. These objects have
347 shapes different than those expected and probably represent groups of a few of
348 individuals. The filter applied during the fourth step (Table 2) discarded objects with
349 mean sizes from 12.43 to 18.29 μm , tending to confirm this interpretation.

350 After applying these filters, spheroidal objects corresponding to groups of framboids,
351 euhedrae or sunflowers or even to the infill of spheroidal holes in the carbonate matrix
352 and foraminifera shells were selected (Figs. 8a and 8b). For this reason, the size
353 distributions of populations of spheroidal groups were estimated during the SEM
354 analysis.

355 **Mixture of independent populations.** The entire size distribution obtained after
356 applying the four-step filtering process contains the mixture of independent populations
357 corresponding to framboids, euhedrae or sunflowers. Moreover, spheroidal groups were
358 also included in the entire size distribution. According to the law of proportionate
359 effects, the growth rate of natural minerals is independent of their size and therefore
360 crystal sizes follow lognormal distributions (Kile and Eberl 2003). Studies on the size
361 distributions of frambooidal and euhedral pyrite have shown that these have lognormal or
362 quasi-lognormal shapes (Wilkin et al. 1996; Pósfai et al. 2001; Merinero et al. 2009).
363 The SEM analysis of the studied carbonate pipes shows that populations of pyrite and
364 iron oxyhydroxides by morphology and arrangement follow lognormal or quasi-
365 lognormal distributions. Therefore, it can be assumed that the entire size distribution is
366 made up of a mixture of several independent lognormal populations. With this criterion,
367 the entire size distribution was separated into independent lognormal subpopulations
368 using the Mclust function.

369 **Correspondence between populations of the SEM and micro-CT analyses**

370 During the SEM analysis, independent populations of pyrite and iron oxyhydroxides
371 were characterized according to their morphology and arrangement. The size
372 distributions of these lognormal or quasi-lognormal populations were estimated. The
373 micro-CT analysis and the subsequent filtering process provided the entire size
374 distribution containing the mixture of several independent and lognormal populations.
375 These lognormal populations were characterized using the Mclust function.
376 Correspondence between SEM and micro-CT populations was made based on the main
377 parameters of the size distributions such as mean, standard deviation, minimum and
378 maximum (Table 3). Micro-CT images and 3D renderings showing the arrangement and
379 shape of the objects were also used to establish this correspondence (Fig. 8). The
380 proposed correspondence between SEM and micro-CT populations can be observed in
381 the rows of Table 3. In general, micro-CT populations have higher mean size than the
382 corresponding SEM populations. For the standard deviation, there is no clear variation
383 trend, although SEM populations of framboids have higher standard deviation than the
384 corresponding micro-CT population.

385 These differences in the mean size can be explained by the techniques used in our study.
386 During the SEM phase of the study, thin sections of the samples were used representing
387 a random cut through objects that were not uniform in size. This introduces a
388 measurement error that some authors estimate to be less than 10% (Wilkin et al. 1996;
389 Cashman and Marsh 1988), although the error rate increases for objects of larger size
390 due to the higher probability of large particle overlap in thin cross-sections (Castro et al.
391 2003). The resolution of micro-CT limits the minimum measurable size of the 3D
392 objects to 4.33 μm . This also explains the smaller mean size of the SEM populations,
393 mainly for the euhedral populations that have a smaller mean size.

394 Differences in the standard deviation can be explained by the representativeness of the
395 micro-CT populations. SEM study is limited by the number of measured objects.
396 Although measures of position such as the mean and the median can be estimated
397 accurately with some hundreds of SEM observations (Wilkin et al. 1996), dispersion
398 measures such as the standard deviation are more sensitive to the number of
399 observations. Given the high number of objects measured during the micro-CT analysis
400 (15000 on average) for each independent population, the estimation obtained for the
401 standard deviation of the real populations can be considered as very reliable.

402 Pyrite and iron oxyhydroxides in the Cornide sample, and the Almazán inner and
403 DIASOM inner samples show two different morphologies that have their corresponding
404 populations in the micro-CT populations (Figs. 8c and 8d). In these samples, a third
405 micro-CT population is obtained, related to spheroidal groups of sunflowers or
406 framboids. Evidence for this can be found in the morphology of the 3D renderings for
407 these populations (Fig. 8a). The filtering process applied to the entire size distribution
408 allows differentiation of irregular and non-spheroidal objects, but some small groups of
409 two or three objects with spheroidal shapes cannot be filtered. Moreover, SEM
410 populations of spheroidal objects follow lognormal or quasi-lognormal distributions and
411 this explains why the same kinds of populations under micro-CT analysis are also
412 lognormal.

413 In all the samples studied, euhedral crystals have the smallest mean size (Table 3). The
414 SEM populations of euhedrae in the Almazán outer sample and the first population of
415 euhedrae in the DIASOM outer sample do not have a corresponding micro-CT
416 population (Table 3). The size of the euhedrae in these populations is near or below the
417 minimum size detected with micro-CT (4.33 μm), which explain this result. Moreover,
418 the euhedral-2 and framboidal-2 populations in the Almazán outer sample form groups.

419 These objects can be separated into individual objects in the SEM study, but in the
420 micro-CT analysis they are considered as groups of objects (Fig. 8b). This explains why
421 they do not have a corresponding micro-CT population. In addition, the spherical
422 groups of euhedrae and framboids are part of the micro-CT population and correspond
423 to the population with the highest mean size.

424 The infilling of foraminiferal shells is very frequent in the DIASOM outer sample.
425 Some of the chambers of the foraminiferal shells have a spheroidal shape and therefore
426 they are not discriminated during the filtering process (Fig. 8b). This explains why the
427 third micro-CT population is related to the infilling of foraminiferal shells.

428 **Implications**

429 This study finds a statistical procedure to obtain representative size distributions for
430 different morphologies and arrangements of pyrite from SEM and micro-CT data.

431 The statistical procedure consisted in a first analysis using SEM data to obtain size
432 distributions grouped by morphology and arrangement followed by a second analysis
433 performed using micro-CT data, obtaining the size distributions of spheroidal and
434 regular 3D objects. These size distributions were statistically analyzed to obtain the
435 representative size distributions for each individual morphology and arrangement of
436 pyrite.

437 Size distributions of pyrite were mainly used to infer paleo-environmental conditions of
438 formation of framboids found in shales or limestones. Therefore, the statistical
439 procedure described here can be applied to these sedimentary rocks containing
440 framboidal pyrite. The high difference between the attenuation coefficients of the matrix
441 of shales or limestones and pyrite, allows the use of micro-CT analysis to obtain the 3D

442 volume of pyrite, on the assumption that minerals with similar attenuation coefficients
443 than pyrite are scarce or not present in the studied rocks.

444 The problem of micro-CT analysis is that all occurrences of pyrite were considered in
445 the 3D volume. The filtering process designed to select spherical and regular objects
446 and the subsequent determination of the optimal mixture of lognormal size distributions
447 allows representative size distributions to be obtained.

448 Previous studies determined size distributions of pyrite measuring the framboids in a 2D
449 surface and extrapolating the results to the whole sample. These size distributions of
450 pyrite had some limitations such as the lack of representativeness, the underestimation
451 of the small sizes and the imprecision of measurement in random cuts through the
452 microscopic pyrite, resulting in size distributions and inferences with possibly
453 significant errors. The micro-CT analysis solves part of these problems obtaining
454 representative size distributions and considering 3D morphologies rather than random
455 cuts, but continues underestimating the lower sizes. The use of voxel sizes below 2.5
456 μm in future works will reduce this underestimation.

457 In the studied carbonate pipes, the mean, standard deviation and maximum sizes of most
458 of the populations of framboids indicate formation within the first few centimeters of
459 sediment below the oxic water-column. The population framboidal-1 occurring inside
460 foraminiferal shells in the Hespérides sample has a size distribution that represents
461 redox conditions near the threshold between oxic/dyoxic and euxinic regimes.
462 Foraminiferal shells constitute isolated microenvironments where bacterial sulfate
463 reduction takes place and therefore where framboidal pyrite forms and grows.
464 Therefore, it is expected that foraminiferal shells represent different redox conditions
465 than other open spaces where framboidal pyrite formed.

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624

625 **Tables**

626

627 **Table 1.** Nomenclature of the samples used in this study, the complete sample
 628 identification, cruise and physiographic unit where samples were recovered.

Sample	Sample identification	Cruise	Physiographic unit
Cornide	ANAS-2001-DA02-MSN-N0016	Anastasya 2001	Cornide mud-volcano
Hespérides	CHICA0211-DA37-MSN-N0106	Chica 0211	Hespérides mud-volcano
Almazán outer	CHICA0211-BT26-MSN-N0107	Chica 0211	Almazán mud-volcano
Almazán inner	CHICA0211-BT26-MSN-N0107	Chica 0211	Almazán mud-volcano
DIASOM outer	ARSA0309-L34-MSN-N0101	Arsa 0309	DIASOM Field
DIASOM inner	ARSA0309-L34-MSN-N0101	Arsa 0309	DIASOM Field

629

630 **Table 2.** Four-step filtering process applied to the 3D objects obtained during the
 631 micro-CT analysis

Step	Filtering criteria	Objects filtered
Step 1	Adjust to the theoretical lognormal distribution in the smaller sizes	Noise
Step 2	Corrected sphericity less than 0.80	Non-spheroidal objects
Step 3	Discrete compactness less than 0.60	Non-spheroidal objects
Step 4	Elongation and flatness of the circumscribed 3D ellipsoid greater than 1.80 and the sum of the elongation and flatness greater than 3	Elongated and flattened objects

632

633

634 **Table 3.** Summary of the independent populations observed in the SEM analysis (left)
 635 and correspondence in the same row with lognormal populations obtained in the micro-
 636 CT analysis (right) for each sample. s.d. = standard deviation; min = minimum; max =
 637 maximum; %n = percentage in number of objects; %V = percentage in volume. n_{SEM}
 638 and $n_{micro-CT}$ are the number of objects measured in each analysis.

CORNIDE $n_{SEM}=1639$ $n_{micro-CT}=34014$													
		mean	s.d.	min	max	%n %V		mean	s.d.	min	max		
SEM	Euhedral	8.55	2.61	4.29	14.87	MCT	P1	36.31	4.92	9.32	2.22	4.33	14.61
	Sunflower	12.03	4.06	5.91	32.45		P2	39.15	25.53	16.18	4.24	8.89	33.25
	Sub-spherical groups	23.78	8.54	8.12	42.76		P3	24.53	69.56	24.91	9.06	12.02	147.57
							Total	100	100	15.15	9.19	4.33	147.57
HESPÉRIDES $n_{SEM}=597$ $n_{micro-CT}=6874$													
		mean	s.d.	min	max	%n %V		mean	s.d.	min	max		
SEM	Euhedral-1	4.96	1.61	2.80	10.32	MCT	P1	4.39	0.05	4.39	0.63	4.33	6.69
	Framboidal-1	6.93	2.55	4.30	11.92		P2	16.62	2.42	7.54	1.14	6.69	10.17
	Euhedral-2	8.38	1.26	5.66	11.17		P3	37.45	13.08	11.25	1.60	7.89	15.23
	Framboidal-2	12.05	3.29	5.04	19.03		P4	27.30	26.33	16.12	2.29	12.04	22.74
	Sunflower	18.41	13.46	3.76	50.46		P5	12.26	33.01	24.29	3.46	18.35	36.31
	Sub-spherical groups	22.95	9.59	10.03	40.54		P6	1.99	25.10	40.45	5.76	30.65	78.81
							Total	100	100	13.95	6.94	4.33	78.81
ALMAZÁN OUTER $n_{SEM}=507$ $n_{micro-CT}=24906$													
		mean	s.d.	min	max	%n %V		mean	s.d.	min	max		
SEM	Euhedral-1	2.96	0.46	2.14	3.99	MCT	P1	73.97	25.40	9.06	3.51	4.33	17.69
	Euhedral-2	4.64	0.39	4.13	5.21								
	Framboidal-1	7.55	1.97	2.47	13.82								
	Framboidal-2	11.61	3.80	5.46	18.00								
	Sub-spherical groups	22.85	16.07	4.30	101.28								
					Total	100	100	15.07	9.39	4.33	148.90		
ALMAZÁN INNER $n_{SEM}=500$ $n_{micro-CT}=9005$													
		mean	s.d.	min	max	%n %V		mean	s.d.	min	max		
SEM	Euhedral	4.57	1.77	2.20	10.52	MCT	P1	31.70	4.29	7.59	2.30	4.33	11.89
	Framboidal	10.67	2.42	7.57	13.83		P2	32.28	14.98	11.98	3.20	6.43	21.70
	Sub-spherical groups	17.77	10.46	7.58	52.99		P3	36.01	80.73	18.25	9.37	8.42	103.82
							Total	100	100	12.95	7.57	4.33	103.82
DIASOM OUTER $n_{SEM}=592$ $n_{micro-CT}=999$													
		mean	s.d.	min	max	%n %V		mean	s.d.	min	max		
SEM	Euhedral-1	4.65	0.36	3.91	5.36	MCT	P1	7.68	0.25	5.50	1.49	4.33	7.62
	Euhedral-2	5.87	2.24	3.01	11.38								
	Framboidal	9.40	5.31	2.67	44.35								
	Sub-spherical groups	19.65	7.49	9.12	31.60								
	Foraminifera	24.99	4.67	16.69	32.24								
					Total	100	100	15.59	9.53	4.33	81.73		
DIASOM INNER $n_{SEM}=1790$ $n_{micro-CT}=15650$													
		mean	s.d.	min	max	%n %V		mean	s.d.	min	max		
SEM	Euhedral	7.39	0.90	4.03	10.56	MCT	P1	16.20	1.82	6.47	2.12	4.33	9.51
	Sunflower	12.24	3.18	6.15	21.21		P2	63.22	37.67	11.65	3.82	5.57	26.77
	Sub-spherical groups	18.50	4.04	9.09	24.61		P3	20.58	60.52	20.54	6.74	12.16	76.45
							Total	100	100	12.65	6.37	4.33	76.45

639

640 **Table 4.** Summary of the four-step filtering process applied to 3D objects obtained
 641 during the micro-CT analysis. For each sample and step, the number (n), mean, standard
 642 deviation (s.d.), minimum (min) and maximum (max) sizes of the discarded objects
 643 were included.

	CORNIDE					ALMAZÁN OUTER					DIASOM OUTER				
	n	mean	s.d.	min	max	n	mean	s.d.	min	max	n	mean	s.d.	min	max
Initial	65693	17.98	22.72	2.5	933.09	35037	15.6	14.05	0	259.98	1400	15.62	11.69	2.5	108.39
Step 1	14322	5.39	0	4.33	7.38	2910	5.24	0	4.33	7.15	63	5.19	0	4.33	7.15
Step 2	3865	83.72	66.8	19.61	933.09	1964	63.45	31.85	20.02	259.98	65	48.95	18.39	25.54	108.39
Step 3	801	11.48	4.77	7.15	27.51	1587	8.01	3.12	5.95	16.61	40	8.4	3.69	5.95	15.88
Step 4	12601	18.29	15.35	5.57	116.17	3670	15.29	14.95	5.57	115.84	233	14.92	13.18	5.57	71
	HESPÉRIDES					ALMAZÁN INNER					DIASOM INNER				
	n	mean	s.d.	min	max	n	mean	s.d.	min	max	n	mean	s.d.	min	max
Initial	10235	17.41	23.08	2.5	895.7	13399	12.88	11.84	2.5	302.24	31894	14.13	16.27	2.5	771.33
Step 1	1510	5.15	0	4.33	7.15	1672	5.13	0	4.33	7.15	4808	5.26	0	4.33	7.26
Step 2	344	103.6	70.24	24.8	895.7	1258	56.2	35.15	18.31	302.24	3204	50.39	37.76	11.12	771.33
Step 3	49	8.59	3.63	7.14	23.3	458	9.3	4.32	5.95	22.99	2028	11.35	4.14	5.95	30.95
Step 4	1458	15.26	15.38	5.57	108.51	1006	12.43	12.59	5.57	98.59	6204	13.10	10.48	5.57	92.24

644

645

646 **Figure Captions**

647 **Figure 1.** Photographs and transversal and longitudinal sections of the carbonate pipes
648 from the Gulf of Cadiz studied here: **(a)** and **(b)** Cornide; **(c)** and **(d)** Hespérides; **(e)** and
649 **(f)** Almazán; **(g)** and **(h)** DIASOM.

650 **Figure 2.** Thin-section back-scatter SEM photomicrographs showing the three
651 morphologies of pyrite and iron oxyhydroxides in the studied samples: **(a)** framboids,
652 **(b)** euhedrae **(c)** sunflowers.

653 **Figure 3.** Thin-section back-scatter SEM photomicrographs showing the morphologies
654 and arrangements of pyrite and iron oxyhydroxides in: **(a)**, **(b)** and **(c)** Cornide; **(d)**, **(e)**
655 and **(f)** Hespérides; **(g)**, **(h)** and **(i)** Almazán Outer.

656 **Figure 4.** Thin-section back-scatter SEM photomicrographs showing the morphologies
657 and arrangements of pyrite and iron oxyhydroxides in: **(a)**, **(b)** and **(c)** Almazán Inner;
658 **(d)**, **(e)** and **(f)** DIASOM Outer; **(g)**, **(h)** and **(i)** DIASOM Inner.

659 **Figure 5.** Histograms and quantile-quantile plots of independent populations observed
660 in the SEM analysis. The theoretical lognormal distribution was included in the
661 histograms and the result of the Shapiro-Wilk test (Royston, 1995) in the quantile-
662 quantile plot. **(a)** euhedrae and **(b)** spheroidal groups in the Cornide sample; **(c)**
663 framboids and **(d)** spheroidal groups of framboids in the Hespérides sample; **(e)**
664 sunflowers and **(f)** spheroidal groups of sunflowers in the DIASOM Inner sample.

665 **Figure 6.** Histograms of the entire size distribution obtained during the micro-CT
666 analysis for each sample. The probability density functions of the lognormal
667 subpopulations obtained during the mixture analysis were plotted in different colors.

668 **Figure 7.** Attenuation coefficient for pyrite, iron oxyhydroxides and main (dashed) and
669 accessory (dotted) minerals in the samples studied, in relation to the energy of the X-ray

670 beam. Values are calculated taking into account the chemical composition of the
671 minerals. Data obtained from the Physical Measurement Laboratory, National Institute
672 of Standards and Technology (physics.nist.gov).

673 **Figure 8.** High-resolution X-ray tomography photomicrographs and 3D renderings
674 produced with the plugin RoiManager3D 2.0 for ImageJ: (a) and (b) spheroidal groups
675 of framboids in the Cornide sample; (c) and (d) spheroidal foraminifera infilling in the
676 DIASOM Outer sample; (e) and (f) euhedrae in the DIASOM Inner sample; (g) and (h)
677 sunflowers in the Hespérides sample. In red objects selected within the threshold of the
678 grey scale; in yellow, objects selected after applying the filtering process and
679 corresponding to the shape and size of the 3D renderings.

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