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5	Collapsing minerals: Crackling noise of sandstone and coal, and
6	the predictability of mining accidents
7	XIANG JIANG ^{1,2} , DEYI JIANG ¹ , JIE CHEN ¹ , AND EKHARD K. H. SALJE ^{2,*}
8	¹ State Key Laboratory of Coal Mine Disaster Dynamics and Control, Chongqing University,
9	Chongqing 400044, People's Republic of China
10	² Department of Earth Sciences, University of Cambridge, Downing Street, Cambridge CB2 3EQ,
11	United Kingdom
12	*Corresponding authors: <u>ekhard@esc.cam.ac.uk</u>
13	
14	ABSTRACT
15	Mining accidents are sometimes preceded by high levels of crackling noise which
16	follow universal rules for the collapse of minerals. The archetypal test cases are
17	sandstone and coal. Their collapse mechanism is almost identical to earthquakes: the
18	crackling noise in large, porous samples follows a power law (Gutenberg-Richter)
19	distribution $P \sim E^{-\epsilon}$ with energy exponents ϵ for near critical stresses of $\epsilon = 1.55$ for dry
20	and wet sandstone, and $\varepsilon = 1.32$ for coal. The exponents of early stages are slightly
21	increased, 1.7 (sandstone) and 1.5 (coal), and appear to represent the collapse of

22	isolated, uncorrelated cavities. A significant increase of the acoustic emission, AE,
23	activity was observed close to the final failure event, which acts as 'warning signal' for
24	the impending major collapse. Waiting times between events also follow power law
25	distributions with exponents $2+\xi$ between 2 and 2.4. Aftershocks occur with
26	probabilities described by Omori coefficients p between 0.84 (sandstone) and 1 (coal).
27	The 'Båth's law' predicts that the ratio between the magnitude of the main event and
28	the largest aftershock is 1.2. Our experimental findings confirm this conjecture.
29	Our results imply that acoustic warning methods are often possible within the context
30	of mining safety measures but that it is not only the increase of crackling noise which
31	can be used as early warning signal but also the change of the energy distribution of the
32	crackling events.
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34	KEYWORDS: Sandstone, coal, crackling noise, failure and collapse event, precursor

35 effects.

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INTRODUCTION

38	Earthquakes and the collapse of porous materials are related phenomena deeply
39	connected by the emission of crackling noise (Baró et al. 2013; Salje and Dahmen 2014;
40	Sethna et al. 2001) where systems under slow perturbation respond through discrete
41	events, so-called 'jerks', with a huge variety of sizes and energies. The signatures of
42	seismic events in geophysics coincide with laboratory-scale experiments ('labquakes')
43	of compressed porous and fractured materials (Davidsen et al. 2007; Diodati et al.
44	1991; Hirata 1987; Kun et al. 2007, 2009; Lebyodkin et al. 2013; Nataf et al. 2014a;
45	Niccolini et al. 2009, 2010, 2011; Petri et al. 1994; Salje et al. 2013; Weiss and Miguel
46	2004) and have been simulated by numerical discrete element calculations of porous
47	materials (Kun et al. 2013, 2014). In laboratory experiments, external loading is applied
48	to the samples and the system's response is obtained by recording acoustic emission
49	(AE). Other recordings are the stepwise change of the macroscopic strain or the
50	emission of calorimetric heat jerk (Baró et al. 2014; Gallardo et al. 2010). All
51	recordings mimic earthquakes since the main stresses underlying tectonic quakes are
52	considered compressive and stationary (Main 1996). Baró et al. (2013) reported a very
53	complete parallel between the acoustic emissions produced by a porous material under
54	uniaxial compression and earthquakes. The same experimental techniques are widely
55	used for the investigation of device materials such as ferroelectric, ferromagnets, and
56	ferroelastics (Bolgár et al. 2016; Dul'kin et al. 2015; Guyot et al. 1988; Hoffmann et al.
57	2001; Salje et al. 2014, 2015; Skal's'kyi et al. 2009; Vives et al. 1994) with a significant

- increase of published data over recent years on the acoustic emission duringforce-induced changes of microstructures.
- 60

61 Previous experimental studies of labquakes have given some evidence that major 62 collapse events are preceded by increased precursor crackling noise, which may allow the prediction of failure during earthquakes. Equally importantly, predicting failure is 63 64 needed in the mining context. The main question is then: is it possible to predict a main 65 collapse event from pre-shocks before the failure event actually occurs? First observations of large-scale foreshock sequences go back to 1988 were a full sequence 66 67 was observed at the Chalfat earthquake by Smith and Priestley (1988) and large sequences of Californian earthquakes by Dodge et al. (1996). In each case the statistical 68 69 evidence was rather limited and related to technical issues of seismological 70 observations. More complete data from laboratory experiments, such as from 71 observations in porous goethite, FeO(OH), revealed two scenarios (Salje et al. 2013). 72 Samples with low porosity showed no evidence for any precursor effects and no 'early 73 warning' signal could be extracted from the compression noise. Samples with high 74 porosities (> 80%), on the other hand, did show precursor noise and opened the 75 possibility to use pico-seismic observations to predict the collapse of a goethite mine 76 (Salje et al. 2013). Subsequent work found no indication of increased crackling noise in 77 berlinite (Nataf et al. 2014b). Instead there was an indication that the exponent of 78 power-law energy distribution reduces near the critical point, in agreement with trends

in numerical simulations of collapse mechanisms by Kun et al. (2013).

80	No significant increase of the AE activity was observed in the first such study of
81	SiO_2 -based vycor (Salje et al. 2011) while an extensive study of other SiO_2 based
82	materials found only a very weak increase of precursor activity in sandstone (Nataf et
83	al. 2014a). No effect was found in charcoal (Ribeiro et al. 2015). These results
84	contradict the simulation results (Kun et al. 2013) where both a big increase of AE
85	activity and a change of the power law exponent were predicted. To test this scenario
86	we changed the experimental arrangement from very small load stresses and samples
87	(Salje et al. 2011) to one where much larger samples can be compressed under
88	enhanced forces. This novel experimental arrangement allows us to study
89	systematically the major collapse mechanisms in the mineralogical context.

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EXPERIMENTAL

The samples of sandstone and coal were collected from the Sichuan and Shanxi provinces of China. The samples were drilled with a high speed rotary saw from larger blocks, their shapes were cylindrical with 50 mm diameter and 100 mm length. The sides of the specimen were optically smooth, and the ends of the specimens were smotth within ± 0.02 mm (according to ISRM testing guidelines, Fairhurst and Hudson 1999). The density and porosity were determined by wax seal methods and mercury

- 98 intrusion analysis, respectively. They are very similar to those of smaller samples99 measured previously (Table 1)
- 100 Dry sandstone samples were heated to 110 for 24h, and cooled to room temperature
- 101 just before the compression experiment. Saturated sandstones were immersed in water
- 102 for 48h.
- 103
- 104 The compression experiment was performed using the loading equipment in Figure 1. 105 The slowly increasing load is provided by oil pouring into a container at a constant flow 106 rate, the weight of oil is then transferred to the lower tilting beam in Figure 1. The 107 samples were placed between the lower tilting beam and a static support. The 108 maximum load is 300 kN and the maximum vertical displacement is 5mm. The stress 109 rate was chosen to be $d\sigma/dt= 8.5$ kPa/s (1kN/min) for all samples.
- 110

Acoustic emission signals were measured during compression by two or more piezoelectric sensors (NANO-30 Physical Acoustics Company) fixed on the sample's round surface by rubber bands. The sensors were acoustically coupled to the sample by a thin layer of grease. The acoustic signal was pre-amplified (40 dB) and transferred to the AE analysis system (DISP from American Physical Acoustics Company). The threshold for detection was chosen as the signal of an empty experiment (45 dB).

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RESULTS

119	The energy of the AE signal, the AE activity (the number of AE hits per second), and
120	the cumulative AE activity are shown as function of the run time in Figure 2 for (a) dry
121	sandstone, (b) wet sandstone and (c) coal. Note the logarithmic scale for the event
122	energies which spans up to 4 decades. AE energies and activities are smaller in the early
123	stages of the compression experiment compared with the late stages. The AE emission
124	reaches a steady state in coal which is terminated before the final major failure event.
125	Some early signals come from the friction between sample's flat faces and the
126	compressive equipment, and the closure of some original micro-cracks in the samples.
127	The AE signals increase dramatically near the major failure event, most notably in
128	sandstone samples, where virtually no steady state emission occurs.
129	
130	The density and porosity of our sandstone samples are close to those of Nataf et al.
131	(2014a, see Table 1), although our failure stress is about 5 times higher than theirs (see
132	Table 2). This may be due to our faster loading rates (Table 2), and smaller levels of
133	stress concentrations in our cylindrical samples compared with the previously used
134	prismatic samples. Our sandstone is also particularly uniform with few preexisting
135	cracks which leads to fewer AE events during the middle stage (Figure 2). Coal
136	contains more original micro cracks than sandstone and more AE signals were recorded
137	before failure.

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ANALYSIS

140 The energy distribution (Gutenberg-Richter law)

141 The probability distribution function, PDF, of jerk energies (events) is P(E) and is 142 derived from the raw data by appropriate linear binning of the event energies. The 143 number of bins was chosen between 10^4 and 10^6 with little influence of the bin number 144 on the functional form of the PDF. Logarithmic binning did not change the results. 145 Figure 3 shows the energy distributions P(E) in log-log plots for dry and wet sandstone, 146 and coal. The histogram corresponds to the accumulation of signals along the whole 147 experiment. A good power law behavior is observed over more than three decades for 148 the early stages and late stages of each collapse sequence: that the distribution of energies follows a power law (Gutemberg-Richter law) $P(E) \sim E^{-\varepsilon}$ is a very good 149 150 approximation but with different exponents for each data set. A finer sequencing with 5 151 intervals in the early stages and 3 intervals for the late stages lead to the same result, 152 namely that the exponents show only two different numerical values with a larger value 153 at the early stages and the smaller value near the major failure event. In each case, the 154 distribution is described by

155
$$P(E)dE \sim \frac{E^{-\varepsilon}}{E_{min}^{1-\varepsilon}}dE \quad E > E_{min}$$
(1)

where E_{min} is a lower cutoff used for normalization. To examine the distribution in more detail we apply the Maximum Likelihood method (Clauset et al. 2009). This method avoids the construction of histograms and the choice of the number of bins. The analytical formula is

160
$$\epsilon(\mathbf{x}_{\min}) = 1 + n \left[\sum_{i=1}^{n} \ln \frac{\mathbf{x}_{i}}{\mathbf{x}_{\min}}\right]^{-1}$$
 (2)

161 where x_i , i = 1...n are the observed values of x such that $x_i \ge x_{min}$.

162 The standard error is

163
$$\sigma = \frac{\varepsilon(x_{\min}) - 1}{\sqrt{n}} + O(\frac{1}{n})$$
(3)

164 The results are as shown in the inserts of Figure 3, this analysis leads to a plateau which 165 defines the exponential parameter ε . Note that we define ε as a positive quantity, the 166 negative sign in the exponent then defines the negative slope of the power law curve. 167 As the number of data points at the late stages (blue curves in the inserts) is much 168 greater than in the early stages we find a much better defined plateau in this case with 169 well-defined exponents $\varepsilon \approx 1.55$ (sandstone) and $\varepsilon \approx 1.32$ (coal) (Figure 4). The 170 agreement between the two methods, namely direct binning and the Maximum 171 Likelihood, for the determination of ε is excellent.

172

173 **Distribution of waiting times**

174 Figure 5 shows the distribution of waiting times, defined as $\delta_i = t_i - t_{i-1}$, with j labeling 175 only the events with energy larger than a given threshold energy E_{min}^* . D(δ , E_{min}^*) is the 176 waiting time distribution function and indicates the probability to observe a waiting time δ for a threshold energy E_{min}^* . When plotting $D(\delta, E_{min}^*)$ as function of δ , we scale 177 178 the axes as $\langle r(E_{\min}^*) \rangle \delta$ and $D(\delta, E_{\min}^*) / \langle r(E_{\min}^*) \rangle$, where $\langle r(E_{\min}^*) \rangle$ is the mean 179 number of events per unit time with an energy $E > E_{min}^*$. This approach leads to a 180 collapse of all data in a single curve showing double power-law behavior with 181 exponents 1 - v for small arguments, and $2 + \xi$ for large arguments (we follow again the

182 convention to define the exponents as positive quantities and maintain the negative sign 183 to indicate the negative slope of the power law function). We have insufficient data to 184 determine (1 - v) while the results for $(2 + \xi) = 2.2 \pm 0.2$ are sufficiently well 185 constrained for all data sets. No significant change of $(2 + \xi)$ was found between dry 186 and water-saturated sandstone.

187

188 Aftershocks and Båth's law

189 The number of aftershocks (AS) is described by the Omori's law (Utsu et al. 1995) as 190 often employed for the analysis of earthquakes. It states that the number of AS decays 191 as a power-law after each mainshock (MS). We define MS as AE signals with energies E_{MS} between 10^k to 10^{k+1} aJ, with k = 1, 2, and 3. The sequence of AS is then continued 192 193 until another MS is found, which termines the AS sequence. We calculate the 194 aftershock rates as function of the lapse time since the MS. The results of dry sandstone, 195 wet sandstone, and coal are shown in Figure 6 (a)–(c). The power law dependence is 196 emphasized by the averaged slope with Omori exponents p=0.84 (sandstone) and 197 p=0.95 (coal). These values are typical for Omori sequences in natural earthquakes.

- 199 The Båth's law (Båth 1965; Console et al. 2003; Helmstetter and Sornette 2003) states
- 200 that the ratio of the energy magnitudes of a MS and a AS is, on average, near to 1.2.
- 201 This ratio is independent of the magnitude of the MS.

202 We plot in Figure 7 the normalized ratio between the magnitude of the mainshock MS

203 and its largest aftershock AS*

$$\Delta M = \log\left(\frac{E_{MS}}{E_{AS}^*}\right) \tag{4}$$

as function of the magnitude of MS. The ratio ΔM shows indeed a trend towards the

206 value of the Båth's law of 1.2 for large MS. The dependency of ΔM on the magnitude of

207 the main shock is well described by extended Debye model:

208
$$\Delta M = A_2 + \frac{(A_1 - A_2)}{1 + (\frac{E_{MS}}{E_0})^p}$$
(5)

209 $A_1=0.13$, $A_2=1.16$, $E_0=413.65$, p=1.21, and R-square = 0.97.

- 210
- 211 DISCUSSION
- 212 Our investigations represent a significant advancement with respect to previous 213 experimental studies of collapsing porous materials because we use large sample 214 volumes and high applied stresses. These conditions are closer to those in the context of 215 mining or the collapse of monuments than previous studies (Baró et al. 2013, 2014; 216 Castillo-Villa et al. 2013; Gallardo et al. 2010; Nataf et al. 2014a, 2014b; Salje et al. 217 2008, 2009, 2011, 2013; Soto-Parra et al. 2015). The large sample volume leads to a 218 new phenomenon which was not possible to observe in small sample, which is the 219 evolution of the event centers during the experiment. The results indicate that all 220 avalanche characteristics can be divided into two groups, namely those at the early 221 stages and those closer to the main failure event. Using eight AE detectors and standard

- localization software, we determine the distribution of event centers where AE jerksoriginate in the sample. These distributions are shown in Figure 8.
- 224

225	The initial stages are characterized by randomly distributed event centers with large
226	distances between them. The events are hence approximately randomly distributed in
227	time and space. We describe these events as 'uncorrelatated'. With increasing number
228	of defect centers we find an accumulation of event centers along the shear diagonal
229	which characterizes the final failure of the sample. The event centers are now very
230	densely located and spatially correlated. The energy exponents for the randomly
231	distributed events are generally larger than those of the correlated events although the
232	difference is small. The difference between the exponents given in Table 3 are not
233	visible in most other experiments while our high resolution investigations make it
234	possible to distinguish between exponents within error margins of $\delta \varepsilon = 0.1$.

235

The power law exponents near the major collapse event (1.54 and 1.32) are similar to those reported by Nataf et al. (2014a) who found exponents ε between 1.44 and 1.55 for three different kinds of sandstone. Furthermore, the exponents of stage 2 for coal (ε_2 =1.32) are also compatible with the statistical results come from ethanol-dampened charcoal (1.27-1.33) (Ribeiro et al. 2015). These observations suggest that our results are consistent with previous research while the exponents for stage 1 are not only higher than the value of the correlated stage, but also higher than the results

243	of previous research. A solution to this conundrum may stem from recent results of
244	Kun et al. (2013, 2014) who investigated cracking noise generated by failure in porous
245	materials by computer simulation. A most remarkable feature of their results is that the
246	value of their exponent $\boldsymbol{\epsilon}$ did not show fix-point behavior but smoothly decreased
247	during compression. The smallest value of $\boldsymbol{\epsilon}$ was found near the final collapse. This
248	simulation result is similar to the evolution of the exponent in this study. Our results
249	suggest that we do not have a smooth time dependence of $\boldsymbol{\epsilon}$ but a stepwise behavior
250	between two fix points, however. Only the second, 'correlated' fix point is close to the
251	mean field solution while the first fix point may relate to uncorrelated, isolated
252	collapses with a different collapse mechanism (Salje and Dahmen 2014). In the
253	simulations by Kun et al. (2013), spatial positions of bursts in their sample show
254	randomness in the early stage of compression with no memory, while damage bands are
255	formed at the later stage. This picture is confirmed by our result in Figure 8 where first
256	few AE centers appear in the top and bottom areas of sandstone sample (Figure 8a),
257	then some AE centers occur randomly (Figure 8b), and finally many AE centers form
258	the final crack (Figure 8c). This indicates that the lower energy exponent relates to
259	'criticality' near the final failure event and agrees with the observation that previous
260	measurements on small samples lead to the same energy exponent together with a much
261	reduced failure strength. This result can be understood if the sandstone samples of
262	Nataf et al. (2014a) were already weakened by internal cracks so that the physical
263	events leading to the crackling noise were intrinsically linked to correlated local failure

264 events rather than to localized collapse mechanisms as seen at the early stages during265 our studies.

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267	A crossover behavior of the avalanches size was found in the study of fiber bundles and
268	electrical fuse models (Pradhan et al. 2005). The avalanches size distributions near to
269	and far away from the catastrophic failure follow both power law behavior with
270	different exponents. Multiple fix-point behavior was also reported by Soto-Parra et al.
271	(2015) for the compression of martensitic porous Ti-Ni. The first series of events is
272	generated by de-twinning (ϵ =2) while later stages relate to fracture (ϵ =1.7). This
273	fracture exponent is close to the early stage, uncorrelated value in our study which
274	appears to indicate that collective shear plane failure does not occur in Ti-Ni alloys.

275

The waiting time distributions in Figure 5 are similar to those of previous investigations of less uniform and smaller samples. For dry and wet sandstone, the exponents $2+\xi=2.2$ and 2.4 are similar with the results for previous light-gray sandstone waiting time analysis ($2+\xi=2.0$) (Nataf et al. 2014a). Exponents of $2+\xi=2.45$ were reported for uniaxial compression of Vycor (Baró et al. 2013), a mesoporous silica ceramics. A reasonable agreement also exists with the power-law exponent $2+\xi=2.0$ from DEM simulation (Kun et al. 2013, 2014).

284	Our data show an excellent adherence of the collapse sequence of sandstone to the
285	Omori's law (p=0.84 in Figure 6). Baró et al. (2013) reported a stable Omori decay in
286	small samples with p=0.75 for Vycor, and p=0.78 for gray sandstone, and p=0.74 for
287	red and yellow sandstone (Nataf et al. 2014a). Coal shows an Omori behavior with
288	p=0.95, similar to charcoal (p=0.87) (Ribeiro et al. 2015). In contrast, recent reports
289	(Tsai et al. 2016) gave much higher values for broken bamboo chopstick and a bundle
290	of spaghetti with p=1.68 and p=3.53, respectively. The same tendency is seen for the
291	numerical values of ΔM . ΔM of charcoal was reported to be 1.2 which is identical with
292	our data. Breaking of bamboo chopsticks lead to $\Delta M = 1.7$ and that of a bundle of
293	spaghetti gave $\Delta M = 0.8$ (Tsai et al. 2016). For earthquakes, large fluctuations of ΔM
294	usually exist for different aftershock sequences (Hainzl et al. 2010).

295

296

IMPLICATION

297 The results of our research have two major implications. Firstly, it shows that the 298 prediction of collapse events is possible. The two indicators for an impending disaster, 299 such as a collapse of a mining shaft or the collapse of a building, are the increase of 300 acoustic emission of crackling noise and the change of its energy exponent. Both 301 indicators can be measured rather easily by highly sensitive microphones attached to 302 the material (sandstone sculptures, houses, coals seams etc). The signals need to be 303 processed as described in this paper, which can be easily done using a simple computer 304 device. The sensitivity of the methods depends largely on microstructural properties of

305	the minerals: uniform but porous sandstone is much more sensitive to 'early warning
306	noise' that coal seams which contains a multitude of local cracks.
307	Secondly, we now understand crackling noise much better because we developed a
308	much better measurement technique which allows us to investigate very large samples.
309	The analogy of crackling noise of porous minerals with earth quakes is now firmly
310	established with the surprising additional feature: the energy exponent (or b-value in
311	seismology) depends on the collectiveness of the collapse event. Local cavity collapse
312	displays higher energy exponents than collective movements such in crack propagation
313	and the formation of micro-faults. Only the latter is close to the theoretical mean field
314	value. This result will doubtlessly stimulate more theoretical research in this field.
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500 **LIST OF TABLE AND FIGURE CAPTIONS**

- 501 **TABLE 1.** The density, porosity of sandstone, coal samples and some previous studied
- 502 materials (Baró et al. 2013; Nataf et al. 2014a).
- 503
- 504 **TABLE 2.** The stress rate, $d\sigma/dt$, failure stress, σ_f , number of recorded AE signals, N,
- and the time span of our experiments and LGsan, Rsan, and Ysan in (Nataf et al.
- 506 2014a).
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- 508 **TABLE 3.** Exponents for statistical laws used in this and previous studies.
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510 **FIGURE 1.**Photograph representation of the compression arrangement. Oil is poured

511 into the hanging container (on the right) at a constant flow rate; the weight of oil (G) is

- transferred to the lower tilting beam. The samples were placed between the lower tilting
- 513 beam and a static support. (Color online.)
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- 516 and the accumulated number of events for (a) dry sandstone, (b) wet sandstone and (c)
- 517 coal samples. In all cases does the activity strongly increase prior to the main failure
- 518 event. This signal can be used as early warning sign for impending collapse, e.g. in coal
- 519 mines and buildings. (Color online.)
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522	and (c) coal samples in the different time windows of experiments. The inset shows the
523	ML-fitting exponent ε as a function of a lower threshold E_{\min} . (Color online.)
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527	analysis leads to a plateau that defines exponents. The exponents of dry (black symbols)
528	and wet sandstone (red symbols) are identical within experimental errors. (Color
529	online.)
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531	FIGURE 5. Distribution of waiting time for different values of E_{min}^* for (a) dry
532	sandstone, (b) wet sandstone and (c) coal samples. (Color online.)
533	
534	FIGURE 6. Rate of aftershocks per unit time, r, as a function of the time lapse to the
535	main shock for (a) dry sandstone, (b) wet sandstone and (c) coal samples. (Color
536	online.)
537	
538	FIGURE 7. The relationship between relative magnitude and mainshock energy for all
539	samples. Solid black line is fitting curve with an extended Debye model. (Color online.)
540	
541	FIGURE 8. Evolution of AE centers. (a) A few AE centers appear in the top and bottom
542	areas of sandstone sample from the friction between sample faces and the loading

- 543 device, (b) some AE centers occur randomly, (c) AE centers form the final crack, and
- 544 (d) photo of the cracked sample. (Color online.)

- 546
- 547

548 **TABLE 1.** The density, porosity of sandstone, coal samples and some previous studied

549 materials (Baró et al. 2013; Nataf et al. 2014a).

550

	Density Porosity		D.C	
	(g/cm ³)	(%)	Kelerence	
Sandstone	2.2	0.18	This study	
Coal	1.2	0.11	This study	
LGsan	2.3 0.13 (Nataf et a	(Nataf et al. 2014a)		
Rsan	2.2	2.2 0.17 (Nataf et al. 2014a) 2.3 0.17 (Nataf et al. 2014a)	(Nataf et al. 2014a)	
Ysan	2.3			
Vycor	ycor 1.5 0.40 (Baró et	(Baró et al. 2013)		

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553

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- and the time span of our experiments and LGsan, Rsan, and Ysan in (Nataf et al.
- 556 2014a).

	$d\sigma/dt$	$\sigma_{\!f}$		Т	Elastic modulus
	(kPa/s)	(MPa)	Ν	(s)	(GPa)
Dry sandstone	8.5	62.2	110906	7190	6.0
Wet sandstone	8.5	46.3	21827	5477	5.8
Coal	8.5	26.4	18968	3235	1.6
LGsan	2.9	13.4	21238		
Rsan	2.4	11.0	27271		3.2
Ysan	1.4	3.6	11058		

557

559

	ϵ_1/ϵ_2 or $<\epsilon>$	1 - v	$2 + \xi$	р	ΔΜ	Reference
Dry sandstone	1.77/1.53		2.2	0.84	1.2	This study
Wet sandstone	1.71/1.56		2.4	0.85	1.2	This study
Coal	1.51/1.32		2.0	0.95	1.2	This study
LGsan	<1.48>	0.86	2.0	0.78		(Nataf et al. 2014a)
Rsan	<1.55>	1.12	2.3	0.74		(Nataf et al. 2014a)
Ysan	<1.49>	1.01	1.8	0.74		(Nataf et al. 2014a)
Vycor	<1.39>	0.93	2.45	0.75		(Baró et al. 2013)
Charcoal	<1.3>			0.87	1.2	(Ribeiro et al. 2015)
Cu-Zn-Al alloys	<2.15>	0.9	2.2			(Baró et al. 2014)
Bamboo	<1.45>			1.68	1.7	(Tsai et al. 2016)
Spaghetti	<1.35>			3.53	0.8	(Tsai et al. 2016)
Simulation	<2.02>		2.0			(Kun et al. 2014)

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570 Oil is poured into the hanging container (on the right) at a constant flow rate; the weight

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599

598





603 **FIGURE 5.** Distribution of waiting time for different values of E_{min}^* for (a) dry

- 604 sandstone, (b) wet sandstone and (c) coal samples. (Color online.)
- 605



0.01

609

610

0.001

0.01

0.1

1

 $t-t_{MS}(s)$

(b)

10

100



614 main shock for (a) dry sandstone, (b) wet sandstone and (c) coal samples. (Color

- 615 online.)
- 616



618 **FIGURE 7.** The relationship between relative magnitude and mainshock energy for all

619 samples. Solid black line is fitting curve with an extended Debye model. (Color online.)

