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2 **Cobalt Mineral Ecology**

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12
13 **ABSTRACT**

14 Minerals containing cobalt as an essential element display systematic trends in their diversity
15 and distribution. We employ data for 66 approved Co mineral species (as tabulated by the
16 official mineral list of the International Mineralogical Association, ruff.info/ima, as of 1 March
17 2016), representing 3554 mineral species-locality pairs (mindat.org and other sources, as of 1
18 March 2016). We find that cobalt-containing mineral species, for which 20% are known at only
19 one locality and more than half are known from 5 or fewer localities, conform to a Large
20 Number of Rare Events (LNRE) distribution. Our model predicts that at least 81 Co minerals
21 exist in Earth's crust today, indicating that at least 15 species have yet to be discovered—a
22 minimum estimate because it assumes that new minerals will be found only using the same
23 methods as in the past. Numerous additional cobalt minerals likely await discovery using micro-
24 analytical methods.

25 Primary Co minerals include 26 species, most of which are hydrothermally deposited
26 chalcogenides. We identify 33 additional plausible as yet undiscovered primary cobalt

27 chalcogenide minerals, including 28 phases with spinel, nickeline, pyrite, and marcasite
28 structural topologies. All 40 secondary cobalt minerals are oxides, and 37 of these phases also
29 incorporate hydrogen. We tabulate an additional 117 plausible secondary Co minerals that are
30 related compositionally and/or structurally to known species. New cobalt minerals are likely to
31 be discovered in specimens collected at the 10 most prolific Co localities, all of which are
32 mining districts with hydrothermal Co mineralization and at least 10 different primary and
33 secondary Co species.

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35 **Keywords:** cobalt, mineral ecology, new minerals, statistical mineralogy, philosophy of
36 mineralogy, rarity, accumulation curves, LNRE distributions

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INTRODUCTION

41 Descriptive mineralogy has traditionally focused on new minerals as they are discovered by
42 systematic surveys or chance finds. Predictions of the numbers, nature, and localities of Earth's
43 as yet undiscovered crystalline phases have been less frequent. Applications of large
44 mineralogical data resources (Hazen 2014; LaFuente et al. 2015) coupled with statistical
45 methods from ecology and lexicology (Baayen 2001; Evert and Baroni 2008) are now leading to
46 predictions of how many minerals remain to be discovered in Earth's crust: Earth's "missing"
47 minerals (Hystad et al. 2015a, 2015b; Hazen et al. 2015a, 2015b, 2016; Grew et al. 2016a).

48 Here we apply the methods of mineral ecology to the minerals of cobalt, a redox-sensitive
49 first-row transition element that is of special interest because of its strategic importance
50 (National Research Council 2008; Orcutt 2011), as well as its critical roles in biology (Young
51 1979; Kobayashi and Shimizu 1999). Cobalt is a relatively common minor element in the crust
52 (Wedepohl 1995; Palme and Jones 2005), averaging ~20 ppm crustal abundance (Rudnick and
53 Gao 2005), with an estimated 29 ppm in the bulk continental crust and 10 ppm in the upper
54 continental crust (McLennan and Taylor 1999), and 48 ppm in ocean basalt (Mielke 1979).
55 Nevertheless, cobalt is an essential element in only 66 minerals, as recorded in ruff.info/ima as
56 of 1 March 2016 (Downs 2006). [Ruff.info/ima](http://ruff.info/ima) is a website created and maintained by the
57 Outreach Committee of the International Mineralogical Association (IMA); it is the official IMA
58 mineral list. The parsimony of cobalt minerals is a consequence of cobalt's charge and ionic
59 radius mimicking that of the more common mineral-forming elements Fe, Ni, and Mg, making it
60 easier for Co to incorporate as a minor element in abundant rock-forming minerals than to be
61 isolated into its own species. Thus it is possible to complete a comprehensive survey of cobalt

62 mineral species and their localities. A subsequent contribution will focus on the temporal
63 distribution and tectonic settings of Co minerals.

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THE MINERALS OF COBALT

66 The 66 minerals of cobalt (Table 1) all occur in crustal environments. We are not aware of
67 any cobalt mineral reported exclusively from mantle rocks or from extraterrestrial environments;
68 indeed the only reported extraterrestrial occurrences are three minerals in meteorites (Rubin
69 (1997a, b; Hazen et al. 2015a). Cobaltite (CoAsS) and safflorite (CoAs₂) have been reported only
70 from the “fossil” Brunflo, Sweden meteorite, an altered H4-H5 chondrite, inferred to have fallen
71 in calcareous mud at 460 to 470 Ma; the cobalt minerals formed by terrestrial alteration of pre-
72 existing meteoritic minerals (Nyström and Wickman 1991). This leaves wairauite (CoFe) as the
73 only cobalt mineral having a definitive meteoritic as well as terrestrial parentage. Hua et al.
74 (1995) inferred that a single grain in the Ningqiang carbonaceous chondrite formed at low
75 temperatures under relatively oxidizing conditions; wairauite was also reported from the CV3
76 chondrites Efremovka and Allende (Krot et al. 1999).

77 Terrestrial cobalt minerals can be divided into two groups. Primary phases (26 species) form
78 principally by direct crystallization in vein deposits by hydrothermal processes. Several of these
79 minerals, including cobaltite (CoAsS), linnaeite (Co₃S₄), safflorite (CoAs₂), and skutterudite
80 (CoAs_{3-x}), are found at hundreds of localities (Table 2) and constitute major ores of cobalt. By
81 contrast, most secondary cobalt minerals (40 species) result from alteration of other Co-bearing
82 phases, principally by oxidative weathering, hydration, and/or other forms of alteration.
83 Additional occurrences of secondary Co minerals arise through alteration of other minerals, most
84 commonly Ni or Cu chalcogenides, that incorporate minor amounts of cobalt. Sphero-cobaltite

85 (CoCO₃) is a special case, possibly occurring as both a primary phase crystalized from
86 carbonate-rich solutions and as an alteration phase derived from primary Co minerals. However,
87 most specimens occur as crystalline or botryoidal coatings on other phases, based on photos on
88 mindat.org and other sources. Therefore, we include spherocobaltite in the list of secondary
89 minerals.

90 Apart from cobalt itself, twenty-five of the 72 essential chemical elements found in minerals
91 overall (rruff.info/ima; Hazen et al. 2015a) are also essential constituents of cobalt minerals. The
92 most frequently encountered elements (Table 3) are oxygen (in 40 species), hydrogen (37),
93 arsenic (29), and sulfur (22). Next in abundance are metals Fe (8), Ni (7), Cu (6), and Ca (5),
94 while chalcogenides Se and Sb are both represented by 5 species. Forty-six of the mineral-
95 forming elements occur in no known cobalt-bearing minerals; these elements include B, Ba, K,
96 Ti, W, and Zn (all of which are represented in synthetic Co phases; Tables 4 and 5). Note that all
97 but two of the 66 known Co-bearing minerals incorporate O, S, As, and/or Se [the exceptions
98 being kiefertite (CoSb₃) and mattagamite (CoTe₂)]. Fifteen Co minerals incorporate only two
99 different chemical elements [Co plus S (4 species), As (5), Se (3), Sb (1), Te (1), and Fe (1)],
100 whereas asbolane [Mn⁴⁺(O,OH)₂·(Co,Ni,Mg,Ca)_x(OH)_{2x}·nH₂O], bouazzerite
101 [Bi₆(Mg,Co)₁₁Fe₁₄(AsO₄)₁₈O₁₂(OH)₄·86H₂O], and julienite [Na₂Co(SCN)₄·8H₂O] are the most
102 chemically complex Co minerals, each with seven different essential elements. The average for
103 26 primary Co minerals is 2.5 essential elements per species, whereas the 40 secondary phases
104 incorporate an average of 4.8 elements per mineral. The average number of essential elements in
105 66 cobalt minerals is 3.9 elements per species.

106 Coexisting elements in primary cobalt minerals differ markedly from those in secondary
107 minerals in several ways. Oxygen is present in all 40 secondary Co minerals, but in none of the

108 26 primary Co species (with the possible exception of spherocobaltite, which may occasionally
109 occur as a primary phase). All 37 known Co minerals with hydrogen are secondary; only
110 cochromite (CoCr_2O_4), petewilliamsite $[(\text{Ni},\text{Co})_2(\text{As}_2\text{O}_7)]$, and spherocobaltite are anhydrous
111 secondary minerals.

112

113 **COBALT MINERAL ECOLOGY**

114 Hystad et al. (2015a) reported that the relationship between the diversity and distribution of
115 mineral species is similar to the frequency distribution of words in a book: typically a few words
116 and phrases occur many times, but most words and phrases are used only once or twice. These
117 rare words and phrases can thus be analyzed to identify the genre and authorship of an unsigned
118 text, and they conform to a Large Number of Rare Event (LNRE) frequency distribution (Baayen
119 2001; Evert and Baroni 2008).

120 We modeled the frequency distribution of cobalt minerals based on the numbers of known
121 localities for each of the 66 approved cobalt minerals. The easiest approach to estimating the
122 number of localities for each species is to interrogate the crowd-sourced data resource
123 mindat.org, which tabulates locality information for every mineral species. This approach was
124 used in the initial mineral ecology studies (Hazen et al. 2015a, 2015b; Hystad et al. 2015a,
125 2015b). For cobalt minerals, we initially analyzed 3662 raw species-locality data from
126 mindat.org (Supplementary Table 1), with 13 species recorded at only one locality, an additional
127 8 species at exactly 2 localities, and 4 species at exactly 3 localities. By contrast, more than half
128 of these species-locality data relate to 3 minerals—cobaltite (762 localities in mindat.org),
129 erythrite (742), and skutterudite (474). The 13 most common cobalt minerals account for 92% of
130 all species-locality data. This distribution of species among localities, with a few common

131 species and many more rare ones, is typical of all minerals, as well as for various subsets of
132 minerals (Hazen et al. 2015a, 2016; Hystad et al. 2015b; Grew et al. 2016b). These data fit to
133 both a finite Zipf-Mandelbrot (fZM) and a Generalized Inverse Gauss-Poisson (GIGP) model
134 (Hystad et al. 2015a; Supplement Text 1). The GIGP parameters for Co minerals ($\gamma =$
135 0.2853056 ; $B = 0.03610296$; $C = 0.2432292$; P -value = 0.81; Hystad et al. 2015a) facilitate
136 modeling of the Co-mineral accumulation curve (Supplemental Figure 1), with a prediction of a
137 total of at least 87 cobalt minerals. By this estimate, at least 21 Co-bearing minerals occur on
138 Earth but have not yet been described.

139 Grew et al. (2016b) noted that uncritical use of locality data from mindat.org can lead to
140 errors for two reasons. In some cases mindat.org does not list all known localities, leading to
141 underestimates of the number of localities. Thus, for example, mindat.org lists only 2 localities
142 (in Canada and Finland) for mattagamite (CoTe_2); however, a systematic search of Georef
143 reveals another report of an Australian occurrence, for a corrected total of 3 localities. In other
144 instances mindat.org lists too many localities because many of the distinct “localities” in
145 mindat.org refer to specimens from different specific outcrops, veins, or quarries within a single
146 ore body or mineralized zone, and thus should be treated as only 1 locality. Thus, the 9
147 mindat.org localities for cobaltlotharmeyerite [$\text{CaCo}_2(\text{AsO}_4)_2 \cdot 2\text{H}_2\text{O}$] in fact represent only two
148 mineralization events—the Bou Azer District in Morocco, and the Schneeberg District of
149 Saxony, Germany.

150 A relatively fast and straightforward approach to improved locality counts is to eliminate the
151 geographically redundant mindat.org localities, while adding missing localities cited in the
152 *Handbook of Mineralogy* but not in mindat.org. The resulting 3525 locality counts
153 (Supplementary Table 1) include 13 species from only one locality, an additional 10 species

154 from exactly 2 localities, and 6 species from exactly 3 localities. These data fit to both fZM and
155 GIGP models. The GIGP parameters ($\gamma = -0.6325573$; $B = 0.02307183$; $C = 2.189657$; P -value =
156 0.38; Hystad et al. 2015a) lead to a prediction of at least 80 cobalt minerals (see Supplemental
157 Figure 2). This improved estimate of 14 “missing” Co-bearing minerals is significantly smaller
158 than that obtained from the uncritical use of mindat.org data noted above.

159 Grew et al. (2016a) argue that the most valid approach to determining the numbers of
160 localities for each species is an exhaustive search of the primary literature for all but the most
161 common species. (The exact locality counts for minerals known from more than 15 localities
162 have little effect on LNRE models, which depend on the rarest species.) Therefore, we searched
163 the primary literature for confirmation of every cobalt mineral locality for species with 50 or
164 fewer localities reported on mindat.org—a task that required examination of approximately 1000
165 primary references. In the process we discovered 41 new localities distributed among 19 rare
166 cobalt minerals, while eliminating 12 unconfirmed species-locality data. The resulting total of
167 3554 mineral species-locality data (Table 1) includes 13 species from only one locality, with an
168 additional 8 species from exactly 2 localities, and 7 species from exactly 3 localities. These data
169 fit to both fZM and GIGP models. The GIGP parameters ($\gamma = -0.4938359$; $B = 0.03418968$; $C =$
170 0.701821 ; P -value = 0.96; Hystad et al. 2015a) lead to a prediction of at least 81 cobalt minerals
171 (Figure 1). This model, and its estimate of 15 “missing” Co-bearing minerals, is nearly identical
172 to that obtained by the simpler procedure of modifying mindat.org locality counts and adding
173 additional localities from the *Handbook of Mineralogy*. However, it remains to be seen whether
174 the simpler protocols are adequate for all mineral groups.

175 Note that our prediction of 15 Co minerals yet to be discovered is a minimum estimate
176 because it assumes that cobalt minerals will continue to be discovered employing the same

177 techniques as in the past. However, we suggest that a significant number of Co minerals have yet
178 to be discovered by application of microanalytical techniques. Presently, an unusually high
179 percentage of known Co minerals (55 of 66, or 83%) can be identified by their distinct color
180 and/or morphology in hand specimen or with a hand lens. The relatively small percentage (16%)
181 of Co minerals discovered by microanalytical techniques contrasts with minerals of many other
182 elements, including Na and Te, for which as many as 50% of species have been discovered
183 through microanalysis (Hazen et al. 2015b). Therefore, we suggest that dozens of additional Co
184 minerals await discovery through micro-Raman spectroscopy, electron microprobe analysis,
185 and/or electron microscopy—species that would not be reflected in our accumulation curve
186 (Figure 1b).

187

188 **THE “MISSING” MINERALS OF COBALT**

189 The 66 known minerals of cobalt represent a small fraction of the thousands of known
190 inorganic Co compounds [Mellor 1935; International Crystal Structure Database (<http://icsd.fiz->
191 [karlsruhe.de](http://icsd.fiz-))]. Hazen et al. (2015a) suggested that every chemical element has the potential to
192 form at least 1000 mineral species; however, relatively rare elements tend to form fewer mineral
193 species because of the special geochemical conditions required to concentrate those elements
194 (Christy 2015; Hazen et al. 2015b). Accordingly, Table 4 records 33 primary and Table 5 lists
195 117 secondary mineral-like synthetic Co compounds that have not been reported in nature, but
196 which might occur in the inventory of as yet undiscovered cobalt minerals on Earth (or on other
197 highly differentiated planets).

198 As a rare element that mimics the crystal chemistry of several other more common transition
199 elements, including Fe, Ni, and Cu, most cobalt in Earth’s crust is probably distributed as a

200 minor element in more common phases, most notably pyrite and other sulfides, but also oxides
201 (notably spinel; Chauviré et al. 2015; Lenaz et al. 2015) and silicates, including pyroxene and
202 olivine (White et al. 1971; Morimoto et al. 1974; Mantovani et al. 2014). As a consequence,
203 cobalt displays far fewer species than might be predicted from its crustal abundance (Christy
204 2015; Hazen et al. 2015). Only when Co is locally concentrated do phases with essential Co
205 crystallize.

206 Synthetic cobalt phases point to plausible as yet undiscovered minerals. Table 4 lists 33 Co
207 chalcogenides that possess the same structural topologies and incorporate the same elements as
208 known minerals. These phases include 9 marcasite-type, 8 pyrite-type, 7 nickeline-type, and 4
209 spinel-type compounds—all of which are structures represented by common cobalt minerals.
210 These synthetic phases also incorporate the same transition elements—Cr, Cu, Fe, and Ni—that
211 are found in primary Co minerals.

212 Secondary cobalt minerals are compositionally more diverse, though all incorporate oxygen,
213 and all but two minerals incorporate hydrogen as an essential element. These compositional
214 characteristics are reflected in our list of 117 plausible Co minerals that are known as synthetic
215 compounds but have not yet been discovered in nature (Table 5). These cobalt compounds
216 include CoO, well-known as a synthetic phase with the rocksalt structure (e.g., Roth 1958;
217 Sumino et al. 1980; Armentrout et al. 2013), as well as two other oxides (spinel-type Co_3O_4 and
218 alumina-type Co_2O_3) and cobaltous hydroxide $[\text{Co}^{2+}(\text{OH})_2]$ with the brucite structure. We also
219 tabulate 11 plausible cobalt double oxides, all of which incorporate either Ca or a second
220 transition element.

221 Only one cobalt silicate mineral has been approved by the IMA: oursinite
222 $[\text{Co}(\text{UO}_2)_2(\text{SiO}_3\text{OH})_2 \cdot 6\text{H}_2\text{O}]$, known from a single locality in the Democratic Republic of the

223 Congo. A second potential cobalt silicate mineral is the Co-dominant analogue of staurolite,
224 “lusakite,” which was reported from near Lusaka, Zambia (Skerl et al. 1934); cobaltian staurolite
225 has been described from Lusaka and Samos, Greece (Čech et al. 1981; Bringhurst and Griffen
226 1986; Taran et al. 2009). Phillips and Griffen (1986) synthesized the series staurolite-“lusakite”
227 over the entire range of Co/(Co+Fe) from 0 to 1. In addition, several other synthetic cobalt
228 silicates are known, notably cobaltian tourmaline (Taran et al. 1993; Setkova et al. 2009) and
229 isomorphs of spinel, olivine, garnet, and pyroxene (e.g., Kohn and Eckart 1962; Geller 1963,
230 White et al. 1971; Morimoto et al. 1974). Cobalt often substitutes as a trace element for Mg-Fe in
231 natural analogues of these and other structures, but has not yet been found as the dominant
232 cation. For example, from 1 to 7 wt% CoO has been found in högbomite, zincohögbomite,
233 gahnite, ilmenite, magnetite, rhodonite and Ni-dominant chlorite (Čech et al. 1976, 1981; Silaev
234 and Yanulova 1993; Taran et al. 1989; Feenstra 1997; Feenstra et al. 2003).

235 Only one cobalt halide has been documented, the chloride leverettite $[\text{Cu}_3\text{CoCl}_2(\text{OH})_6]$, which
236 is known from a single locality. Many additional synthetic fluorides and chlorides are known,
237 most of which are hydrous compounds (Table 5).

238 Twenty-nine of the 40 known secondary cobalt minerals are sulfates, arsenates, or selenates.
239 Table 5 records 38 additional examples known only as synthetic compounds. The majority of
240 these synthetics, as with known minerals, are hydrous. All of these phases also incorporate other
241 essential elements that occur in Co minerals, but we also list phases with potassium and
242 barium—elements that are well documented in synthetic cobalt compounds but not yet known in
243 natural examples.

244 Two secondary hydrous carbonates have been described—the rare minerals comblainite
245 $[\text{Ni}_6\text{Co}^{3+}_2\text{CO}_3(\text{OH})_{16}\cdot 4\text{H}_2\text{O}]$ and kolwezite $[\text{CuCoCO}_3(\text{OH})_2]$. At least five additional mineral-
246 like hydrous cobalt carbonates are known as synthetic compounds.

247 Only one cobalt phosphate is known in nature: pakhomovskiyite $[\text{Co}_3(\text{PO}_4)_2\cdot 8\text{H}_2\text{O}]$, found
248 only on the Kola Peninsula of Russia. Nevertheless, numerous synthetic Co phosphates have been
249 described (36 of which appear in Table 5) and they point to the potential diversity if a deposit
250 enriched in Co + P were to be discovered. In addition to 9 hydrous Co phosphates, compounds
251 with Na, K, Mg, Ba, and several transition metals demonstrate the potential richness of these as
252 yet undiscovered phases.

253 A number of synthetic cobalt compounds incorporate elements that have not yet been found to
254 coexist with Co in any mineral species, but are plausible as candidates for as yet undiscovered
255 natural phases. Most notably, these elements include potassium, which is known in dozens of
256 synthetic cobalt halides, sulfates, selenates, carbonates, and phosphates. Other likely candidate
257 elements in Co minerals include boron, barium, zinc, chlorine, and titanium, all of which are
258 known in mineral-like synthetic Co compounds.

259 Note that we do not include Co^{4+} compounds, which have not yet been found in nature, but
260 have been synthesized in a variety of phases, including cobalt dioxide, CoO_2 , and a variety of
261 “percobaltites” such as Na_2CoO_3 , BaCo_2O_5 , and MgCoO_3 . The absence of natural Co^{4+} minerals
262 is likely due to the fact that tetravalent cobalt requires redox conditions that are unlikely in any
263 crustal environment, much less in Earth’s mantle.

264 This survey of plausible as yet undiscovered cobalt minerals points to strategies for the
265 discovery of new Co phases. In particular, the 10 localities with 10 or more Co mineral species
266 (Table 2) are especially rich in unusual species; they include type localities for 26 of 66 approved

267 cobalt minerals. Close examinations of specimens from those localities, several of which are
268 flooded or otherwise no longer accessible, are likely to yield additional cobalt minerals.

269

270 **IMPLICATIONS: MINERAL ECOLOGY OF A TRACE ELEMENT**

271 Previous studies in mineral ecology have focused on elements, including beryllium, boron,
272 and carbon (Grew and Hazen 2014; Grew et al. 2016a, Hazen et al. 2016), that have unique
273 crystal chemical characteristics and thus tend to concentrate in their own minerals. Cobalt, by
274 contrast, is the first element to be studied with this methodology that usually mimics more
275 common elements—most notably magnesium, iron, and nickel—rather than form its own
276 minerals. For example, ocean basalt, with total volume $\sim 3 \times 10^9 \text{ km}^3$ and density $\sim 3 \text{ gm/cm}^3$ (and
277 therefore total mass $\sim 9 \times 10^{24} \text{ gm}$), has an average Co content of 48 ppm, with up to 150 ppm in
278 olivine (Herzberg et al. 2016). Ocean basalt thus holds $\sim 4 \times 10^{14}$ metric tons of cobalt—more
279 than 50,000,000 times greater than the 2016 United States Geological Survey estimate of world
280 cobalt ore reserves ($\sim 7.1 \times 10^6$ tons; <http://minerals.usgs.gov>). Even the most common primary
281 Co minerals, such as cobaltite and skutterudite, are volumetrically insignificant by comparison.
282 One might expect, therefore, that cobalt minerals display a different frequency distribution than
283 minerals of more common elements, or of rare elements that are not readily incorporated as
284 minor constituents of common rock-forming minerals.

285 Our results suggest that, on the contrary, the minerals of cobalt conform to a robust LNRE
286 distribution that facilitates predictions of the numbers of as yet undiscovered Co mineral species,
287 just as an LNRE distribution does for all minerals in Earth's crust (Hazen et al. 2015a). We
288 conclude that other subsets of minerals grouped by both major and minor elements could also

289 conform to a LNRE distribution: the viability of LNRE modeling for boron minerals is currently
290 under study (Grew et al. 2016a).

291 More general questions concern the underlying reasons for this fascinating statistical feature
292 of minerals on Earth, and whether similar distributions occur on other worlds. In particular, the
293 majority of rare cobalt minerals are secondary phases that arise from near-surface oxidation;
294 these minerals are thus indirect consequences of biological processes. Studies in progress will
295 explore whether the observed LNRE distributions of Earth's minerals are uniquely a
296 consequence of rare biologically-mediated species and, consequently, if LNRE distributions are
297 an unambiguous mineralogical signature of life on Earth-like planets.

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299

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- 427

428 **Table 3.** Coexisting essential elements in cobalt minerals, including elements in the 26 primary
429 versus 40 secondary species. Numbers for 25 elements are based on mineral species and
430 chemical formulas in ruff.info/ima as of 1 March 2016.

431

432	Element	# of Co Minerals	Primary	Secondary
433	O	40	0	40
434	H (All with O)	37	0	37
435	As	29	10	19
436	S	22	12	10
437	Fe	8	3	5
438	Ni	7	2	5
439	Cu	6	2	4
440	Ca	5	0	5
441	Se	5	4	1
442	Sb	5	5	0
443	C	4	0	4
444	Bi	4	0	4
445	Na	2	0	2
446	U	3	0	3
447	Mg	2	0	2
448	Pb	2	0	2
449	N	1	0	1
450	Al	1	0	1
451	Si	1	0	1
452	Cl	1	0	1
453	P	1	0	1
454	V	1	0	1
455	Cr	1	0	1
456	Mn	1	0	1
457	Te	1	1	0

458

459 **TABLE 4.** Plausible as yet undescribed primary cobalt chalcogenide minerals, based on synthetic Co-
 460 bearing phases tabulated in the International Crystal Structure database that incorporate from 2 to 4
 461 different chemical elements and possess structural topologies known from other primary Co minerals
 462 (Table 1).

463	<hr/>		
464	Formula	Structure Type	Related known minerals
465	<hr/>		
466	<i>Sulfides</i>		
467	NiCo ₂ S ₄	Spinel	Co ₃ S ₄ (linnaeite); CuCo ₂ S ₄ (carrollite);
468	CoCr ₂ S ₄	Spinel	CoNi ₂ S ₄ (siegenite)
469	CoFeS ₂	Nickeline	CoS (jaipurite)
470	(CoCu)(S ₂) ₂	Pyrite	CoS ₂ (cattierite)
471	(CoNi)(S ₂) ₂	Pyrite	
472	(CoFe)(S ₂) ₂	Pyrite	
473	CoSeS	Pyrite	
474	Co ₈ FeS ₈	Pentlandite	Co ₉ S ₈ (cobaltpentlandite)
475			
476	<i>Arsenides</i>		
477	CoAs ₂	Arsenopyrite	(Co,Fe)AsS (glaucodot)
478	CoAsSe	Pyrite	CoSe ₂ (trogtalite)
479	CoAs ₂	Marcasite	
480	CoAsSe	Marcasite	
481	Co ₂ AsSb	Nickeline	CoAs (langisite)
482	CoNiAs ₂	Nickeline	
483	Co _{0.5} Ni _{0.5} As ₂	Marcasite	
484	Co _{0.5} Fe _{0.5} As ₂	Marcasite	
485			
486	<i>Selenides</i>		
487	Co ₉ Se ₈	Pentlandite	
488	CoSe ₂	Marcasite	
489	CoSbSe	Marcasite	
490	CoCr ₂ Se ₄	Spinel	Co ₃ Se ₄ (bornhardtite);
491	Co ₂ NiSe ₄	Spinel	(Co,Cu,Ni) ₃ Se ₄ (tyrrellite)
492	CoFe ₂ Se ₄	Wilkmanite	

493	$\text{CoNi}(\text{Se}_2)_2$	Pyrite	
494	CoNiSe_2	Nickeline	CoSe (frieboldite)
495			
496	<i>Antimonides</i>		
497	CoSb	Nickeline	
498	CoSb_2	Marcasite	
499	$\text{Co}_{0.5}\text{Fe}_{0.5}\text{Sb}_2$	Marcasite	
500	CoSbTe	Marcasite	
501	CoSb_2	Monoclinic CoSb_2	
502			
503	<i>Tellurides</i>		
504	CoTe	Nickeline	
505	CoNiTe_2	Nickeline	
506	CoTe_2	Pyrite	
507	$\text{CoNi}(\text{Te}_2)_2$	Pyrite	
508			
509			
510			

511

512 **TABLE 5.** Plausible as yet undescribed secondary cobalt minerals, based on synthetic Co-bearing phases
513 tabulated in the International Crystal Structure database that incorporate oxygen and from 1 to 5
514 additional chemical elements (Table 1).

515

516 **Formula** **Structure Type**

517

518 Co Single Oxides/Hydroxides

519 CoO Halite
520 CoO₂ CdI₂
521 Co₃O₄ Spinel
522 Co(OH)₂ Brucite

523 Double Oxides

524 NiCo₂O₄ Spinel
525 CoFe₂O₄ Spinel
526 MnCo₂O₄ Spinel
527 CoCu₂O₃
528 Ca₂CoO₃
529 CaCo₂O₄ Harmunite
530 CoMnO₃
531 CoTiO₃ Ilmenite
532 CoCrO₄ CrVO₄
533 CoMoO₄
534 CoWO₄

535 Silicates/Aluminates

536 Co₂SiO₄ Spinel
537 CoAl₂O₄ Spinel
538 Co₂SiO₄ Olivine
539 MgCoSiO₄ Olivine
540 Ca₂Co(Si₂O₇) Akermanite
541 Co₃Al₂Si₃O₁₂ Garnet
542 CaCo(Si₂O₆) Diopside
543 Na₂Co(Si₄O₁₀) Layer silicate

544 Fluorides

545	CoF_2	
546	$\text{CoF}_2 \cdot \text{H}_2\text{O}$	
547	$\text{Co}_2\text{OF}_2 \cdot \text{H}_2\text{O}$	
548	K_2CoF_4	
549	$\text{KCoF}_3 \cdot \text{H}_2\text{O}$	
550	$\text{NaCoF}_3 \cdot \text{H}_2\text{O}$	
551	$\text{AlCoF}_5 \cdot 7\text{H}_2\text{O}$	
552	$\text{CoF}_3 \cdot n\text{H}_2\text{O}$	
553	<i>Chlorides</i>	
554	CoCl_2	
555	$\text{CoCl}_2 \cdot \text{H}_2\text{O}$	
556	$\text{Co}(\text{OH})\text{Cl}$	
557	$\text{Co}_2(\text{OH})_3\text{Cl}$	
558	$\text{MgCoCl}_4 \cdot 8\text{H}_2\text{O}$	
559	$\text{Zn}_2\text{CoCl}_6 \cdot 12\text{H}_2\text{O}$	
560	$\text{ZnCoCl}_4 \cdot 6\text{H}_2\text{O}$	
561	<i>Sulfates</i>	
562	CoSO_4	
563	$\text{Co}_3(\text{SO}_4)_2(\text{OH})_2 \cdot 2\text{H}_2\text{O}$	
564	$\text{Co}_5(\text{SO}_4)_2(\text{OH})_6 \cdot 4\text{H}_2\text{O}$	
565	$\text{NaCo}(\text{SO}_4)\text{F}$	
566	$\text{NaCo}_2(\text{SO}_4)_2(\text{H}_2\text{O})(\text{OH})$	
567	$\text{K}_2\text{Co}_2(\text{SO}_4)_3$	
568	$\text{K}_2\text{Co}(\text{SO}_4)_2 \cdot 2\text{H}_2\text{O}$	
569	$\text{K}_2\text{Co}(\text{SO}_4)_2 \cdot 6\text{H}_2\text{O}$	
570	$\text{CoNi}(\text{SO}_4)_2 \cdot 12\text{H}_2\text{O}$	
571	$\text{Co}(\text{UO}_2)(\text{SO}_4)_2 \cdot 5\text{H}_2\text{O}$	
572	<i>Arsenates</i>	
573	$\text{Co}_3(\text{AsO}_4)_2$	
574	$\text{Co}_2\text{As}_2\text{O}_5$	
575	$\text{Co}_2\text{As}_2\text{O}_7$	Thortveitite
576	$\text{Co}_2(\text{AsO}_4)(\text{OH})$	Andalusite
577	$\text{Co}_2(\text{As}_2\text{O}_7) \cdot 2\text{H}_2\text{O}$	
578	$\text{Co}_2(\text{AsO}_4)\text{Cl}$	

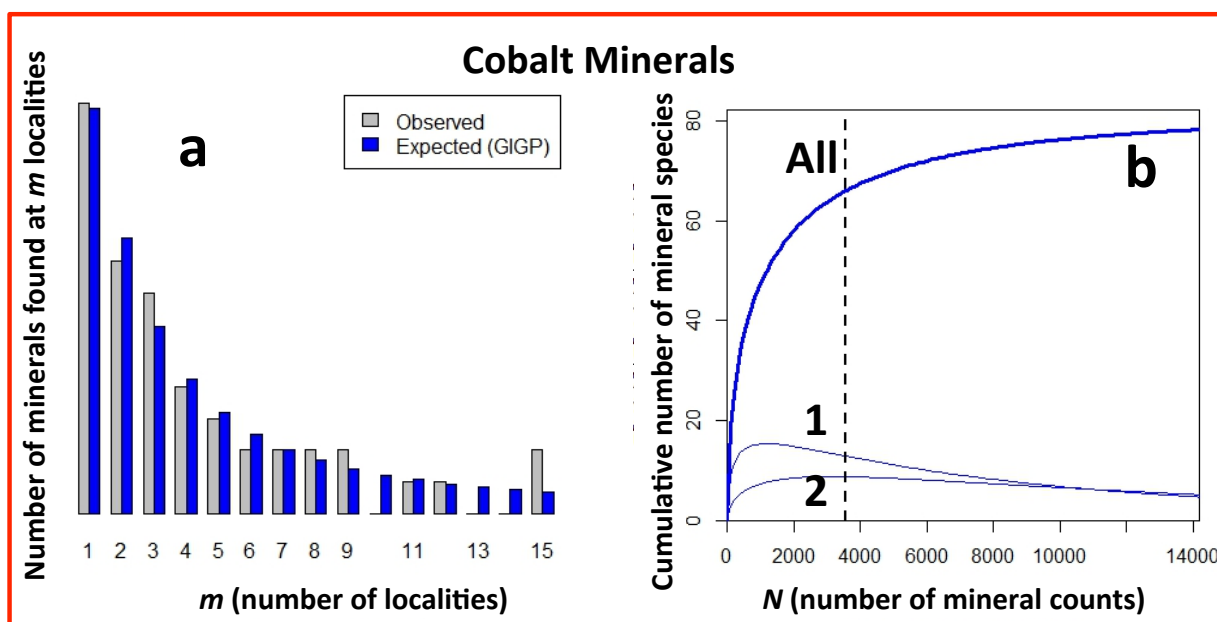
579	$\text{BaCo}_2(\text{AsO}_4)_2$	
580	$\text{Co}(\text{UO}_2)_2(\text{AsO}_4)_2 \cdot 12\text{H}_2\text{O}$	
581	<i>Selenates</i>	
582	$\text{Co}(\text{SeO}_4) \cdot \text{H}_2\text{O}$	
583	$\text{Co}(\text{SeO}_4) \cdot 5\text{H}_2\text{O}$	
584	$\text{Co}(\text{SeO}_3) \cdot \text{H}_2\text{O}$	
585	$\text{Co}(\text{HSeO}_3)_2 \cdot 2\text{H}_2\text{O}$	
586	$\text{Co}(\text{HSeO}_3)_2 \cdot 4\text{H}_2\text{O}$	
587	$\text{Co}_3(\text{SeO}_3)_3 \cdot \text{H}_2\text{O}$	
588	$\text{Co}_5(\text{SeO}_3)_4\text{Cl}_2$	
589	$\text{Co}(\text{HSeO}_3)\text{Cl} \cdot 2\text{H}_2\text{O}$	
590	$\text{Co}(\text{HSeO}_3)\text{Cl} \cdot 3\text{H}_2\text{O}$	
591	$\text{Na}_2\text{Co}(\text{SeO}_4)_2 \cdot 2\text{H}_2\text{O}$	Fairfieldite
592	$\text{K}_2\text{Co}_2(\text{SeO}_3)_3$	
593	$\text{K}_2\text{Co}_2(\text{SeO}_3)_2(\text{OH})$	
594	$\text{K}_2\text{Co}_2(\text{SeO}_3)_3 \cdot 2\text{H}_2\text{O}$	
595	$\text{K}_2\text{Co}(\text{SeO}_4)_2 \cdot 2\text{H}_2\text{O}$	
596	$\text{K}_2\text{Co}(\text{SeO}_4)_2 \cdot 6\text{H}_2\text{O}$	
597	$\text{Ca}_3\text{Co}(\text{SeO}_3)_4$	
598	$\text{BaCo}(\text{SeO}_3)_2$	
599	$\text{BaCo}_2(\text{SeO}_3)_3 \cdot 3\text{H}_2\text{O}$	
600	$\text{Co}_3\text{Fe}_2(\text{SeO}_3)_6 \cdot 2\text{H}_2\text{O}$	
601	$\text{Co}_2(\text{UO}_2)_3(\text{SeO}_4)_5 \cdot 16\text{H}_2\text{O}$	
602	<i>Carbonates</i>	
603	$\text{Co}_2(\text{CO}_3)(\text{OH})_2$	
604	$\text{K}_2\text{Co}(\text{CO}_3)_2(\text{H}_2\text{O})_4$	
605	$\text{Na}_2\text{Co}(\text{CO}_3)_2(\text{H}_2\text{O})_4$	
606	$\text{Ba}_2\text{Co}(\text{CO}_3)_2\text{F}_2$	
607	$\text{Na}_3\text{Co}(\text{CO}_3)_2\text{Cl}$	
608	<i>Phosphates</i>	
609	CoPO_4	
610	$\text{Co}_3(\text{PO}_4)_2$	Chopinite
611	$\text{Co}_3(\text{PO}_4)_2(\text{H}_2\text{O})$	
612	$\text{Co}_3(\text{PO}_4)_2(\text{H}_2\text{O})_6$	

613	$\text{Co}_3(\text{HPO}_4)_2(\text{OH})_2$	Lazulite
614	$\text{Co}_3(\text{PO}_4)_2(\text{H}_2\text{O})_4$	
615	$\text{Co}_2(\text{PO}_4)(\text{OH})$	
616	$\text{Co}_2(\text{PO}_4)_2(\text{OH})_4$	
617	$\text{Co}_5(\text{PO}_4)_2(\text{OH})_4$	
618	$\text{Co}(\text{HPO}_4)(\text{H}_2\text{O})$	
619	$\text{Co}(\text{H}_2\text{PO}_4)_2(\text{H}_2\text{O})_2$	
620	$\text{Co}_2(\text{PO}_4)\text{Cl}$	
621	$\text{NaCo}(\text{PO}_4)$	Olivine
622	$\text{NaCo}_4(\text{PO}_4)_3$	
623	$\text{NaCo}_3(\text{PO}_4)_2(\text{OH})\cdot\text{H}_2\text{O}$	
624	$\text{Na}_2\text{Co}(\text{PO}_4)\text{F}$	
625	$\text{KCo}(\text{PO}_4)$	Andalusite
626	$\text{K}_2\text{Co}(\text{PO}_4)_3$	
627	$\text{KCo}_4(\text{PO}_4)_3$	
628	$\text{KCoAl}(\text{PO}_4)_2$	
629	$\text{KCo}_3\text{Fe}(\text{PO}_4)_3$	
630	$\text{KCo}_2(\text{PO}_4)_2\cdot\text{H}_2\text{O}$	
631	$\text{KCo}(\text{PO}_4)\cdot\text{H}_2\text{O}$	
632	$\text{Mg}_3\text{Co}_3(\text{PO}_4)_4$	
633	$\text{BaCo}_2(\text{PO}_4)_2$	
634	$\text{Co}(\text{TiO})_2(\text{PO}_4)$	
635	$\text{Zn}_2\text{Co}(\text{PO}_4)_2$	
636	$\text{Zn}_3\text{Co}_3(\text{PO}_4)_4$	
637	$\text{ZnCo}(\text{PO}_4)(\text{OH})$	
638	$\text{CoZn}_2(\text{PO}_4)_2\cdot 4\text{H}_2\text{O}$	
639	$\text{CoFe}(\text{PO}_4)\text{O}$	Sillimanite
640	$\text{Co}_2\text{Fe}(\text{PO}_4)_2$	
641	$\text{Co}_3\text{Cr}_4(\text{PO}_4)_6$	
642	$\text{FeCo}(\text{PO}_4)\text{F}$	
643	$\text{Co}_3(\text{BO}_3)(\text{PO}_4)$	
644	$\text{Co}(\text{UO}_2)_2(\text{PO}_4)_2\cdot 10\text{H}_2\text{O}$	

645

646

647 **Figure 1.** (a) Frequency spectrum analysis of 66 cobalt-bearing minerals, with 3554 individual
648 mineral-locality data (from mindat.org as of 1 March 2016), employed a Generalized Inverse
649 Gauss-Poisson (GIGP) function to model the number of mineral species for minerals found at
650 exactly 1 to 5 localities (Hystad et al. 2015a). (b) This model facilitates the prediction of the
651 mineral species accumulation curve (upper curve, “All”), which plots the number of expected Co
652 mineral species (y -axis) as additional mineral species/locality data (x -axis) are discovered. The
653 vertical dashed line indicates data recorded as of 1 March 2016 in mindat.org, as well as locality
654 data from the *Handbook of Mineralogy* (Anthony et al. 2003) and systematic searching under
655 each mineral name in Georef. The model also predicts the varying numbers of mineral species
656 known from exactly one locality (curve 1) or from exactly two localities (curve 2). Note that the
657 number of mineral species from only one locality is now decreasing, whereas the number from
658 two localities is now increasing, but is almost at a maximum and will eventually decrease. We
659 predict that the number of minerals known from two localities will surpass those from one
660 locality when the number of species-locality data exceeds ~9000.



661

Table 1: IMA recognized primary and secondary minerals of cobalt, with numbers of recorded occurrences in parentheses (see text), chemical formulas, paragenetic modes, and selected mineral localities (see Table 2 for key to localities; the type locality is underscored).

# Name (# Localities)	Formula	Paragenetic Mode	Select Localities (see Table 2 for key)
<u>Primary Co Minerals</u>			
1 Alloclasite (37)	CoAsS	Low-T hydrothermal	1,3,4,6,9,10,20,24,27, <u>28</u> ,51
2 Bornhardite (3)	Co ²⁺ Co ³⁺ ₂ Se ₄	Hydrothermal	<u>31</u>
3 Carrollite (127)	CuCo ₂ S ₄	Hydrothermal vein deposits	1,5,7,10,12,13,17,18,26,28, <u>32</u> ,35,42,47,50
4 Cattierite (31)	CoS ₂	Carbonate-hosted	2, <u>5</u> ,8,12,13,17,18,19,20
5 Clinosafflorite (19)	CoAs ₂	Hydrothermal Co-Ni ore	1,3, <u>6</u> ,8,13
6 Cobaltite (762)	CoAsS	High-T hydrothermal	1,2,,3,4,6,7,8,10,12,13,15,16,17,18,19,20,21,22,23,25 26,27,28,30,31,34,35,40,41,43,44,45,46,47,50,52,56
7 Cobaltpentlandite (54)	Co ₉ S ₈	Hydrothermal	6,7,17,18,19,28, <u>38</u> ,40,46
8 Costibite (15)	CoSbS	Hydrothermal	4, <u>15</u> ,25
9 Ferroskutterudite (1)	(Fe,Co)As ₃	Hydrothermal	<u>41</u>
10 Freboldite (4)	CoSe	Hydrothermal; carbonate hosted	<u>31</u>
11 Glaucodot (108)	Co _{0.5} Fe _{0.5} AsS	High-T hydrothermal	1,2,4,10,16,25,27,28, <u>42</u> ,51
12 Jaipurite (4)	CoS	Hydrothermal ore	<u>43</u>
13 Kieftite (3)	CoSb ₃	Sulfide skarn	1, <u>7</u>
14 Langisite (5)	CoAs	Hydrothermal zone	<u>6</u>
15 Linnaeite (181)	Co ₃ S ₄	Hydrothermal veins	1,3,5,7,10,13,14,15,17,18,19,22,25,27,32, <u>35</u> ,40,41,50
16 Mattagamite (3)	CoTe ₂	Massive Te ore	19, <u>44</u>
17 Modderite (3)	CoAs	Heavy mineral concentrates	<u>45</u>
18 Oenite (1)	CoSbAs	Cu-Co skarns	<u>7</u>
19 Paracostibite (7)	CoSbS	Massive sulfide; hydrothermal	<u>47</u>

20 Safflorite (266)	CoAs ₂	Moderate-T hydrothermal	1,2,3,6,8,9,10,11,12,13,15,16,20,21,22,23,27,30,36 40,46,48,51
21 Siegenite (211)	CoNi ₂ S ₄	Hydrothermal sulfide vein	3,4,5,6,10,12,13,14,18,19,25,26,32
22 Skutterudite (474)	CoAs _{3-x}	Medium- to High-T veins	1,2,3,6,8,9,10,12,14,15,16,17,20,21,22,23,24,27,30 34,36,39,40,48,50,51
23 Trogtalite (8)	CoSe ₂	Hydrothermal ore	<u>31</u>
24 Tyrrellite (9)	(Co,Cu,Ni) ₃ Se ₄	Hydrothermal ore	<u>52</u>
25 Wairauite (8)	CoFe	Serpentinization	<u>65</u>
26 Willyamite (16)	CoSbS	Carbonate-hosted veins	4,7, <u>15</u> ,25

Secondary Co Minerals

27 Aplowite (5)	CoSO ₄ ·4H ₂ O	Efflorescences with sulfides	11, <u>29</u> ,46
28 Asbolane (122)	Mn ⁴⁺ (O,OH) ₂ ·(Co,Ni,Mg,Ca) _x (OH) _{2x} ·nH ₂ O	Weathering; ultramafic soils	1,2,3,4,5,9,10,11,12,16,21,23,24,26,36,39,42,48
29 Bieberite (53)	CoSO ₄ ·7H ₂ O	Oxidation of S-As ores	2,3,10,11,14,23,24, <u>30</u> ,39
30 Bouazzerite (1)	Bi ₆ (Mg,Co) ₁₁ Fe ₁₄ (AsO ₄) ₁₈ O ₁₂ (OH) ₄ ·86H ₂ O	Oxidation of ores	<u>1</u>
31 Burgessite (2)	Co ₂ (H ₂ O) ₄ (AsO ₃ OH) ₂ ·H ₂ O	Secondary weathering zone	<u>6</u>
32 Cobaltarthurite (3)	CoFe ⁺³ ₂ (AsO ₄) ₂ (OH) ₂ ·4H ₂ O	Weathered zone	1, <u>33</u>
33 Cobaltaustinite (5)	CaCoAsO ₄ (OH)	Oxidized zone of ores	1,2,11, <u>34</u>
34 Cobaltkieserite (2)	CoSO ₄ ·H ₂ O	Oxidized ore	3, <u>35</u>
35 Cobaltkoritnigite (17)	Co(AsO ₃ OH)·H ₂ O	Weathering of glaucodot	1,2,3,8,9,21, <u>36</u>
36 Cobaltlotharmeyerite (2)	CaCo ₂ (AsO ₄) ₂ ·2H ₂ O	Oxidized ore	1, <u>2</u>
37 Cobaltneustädtelite (2)	Bi ₂ Fe ³⁺ (Co,Fe ³⁺)(AsO ₄) ₂ (O,OH) ₄	Oxidized ore	<u>2</u>
38 Cobaltoblodite (1)	Na ₂ Co(SO ₄) ₂ ·4H ₂ O	Oxidized ore	<u>39</u>
39 Cobaltomenite (6)	CoSe ⁴⁺ O ₃ ·2H ₂ O	Oxidized ore	5, <u>37</u>
40 Cobalttsumcorite (1)	PbCo ₂ (AsO ₄) ₂ ·2H ₂ O	Oxidized ore	<u>2</u>
41 Cobaltzippeite (2)	Co(UO ₂) ₂ (SO ₄)O ₂ ·3.5H ₂ O	Oxidized ore	3, <u>39</u>

42 Cochromite (3)	CoCr_2O_4	Thermal metamorphism of chromite	<u>40</u>
43 Comblainite (4)	$\text{Ni}_6\text{Co}^{3+}_2\text{CO}_3(\text{OH})_{16}\cdot 4\text{H}_2\text{O}$	Oxidized ore	<u>5</u> ,31
44 Erythrite (742)	$\text{Co}_3(\text{AsO}_4)_2\cdot 8\text{H}_2\text{O}$	Oxidized ore	<u>1,2,3,4,5,6,7,8,9,10,11,13,14,15,16,18,20,21,22,23,24,25,26,27,29,30,31,33,34,35,36,39,42,46,48,49,50,51</u>
45 Hetrogenite (102)	$\text{Co}^{3+}\text{O}(\text{OH})$	Weathering	<u>1,2,4,5,6,8,9,16,21,23,24,26,33,42</u>
46 Hloušekite (1)	$(\text{Ni},\text{Co})\text{Cu}_4(\text{AsO}_4)_2(\text{AsO}_3\text{OH})_2$	Oxidized ore	<u>1</u>
47 Jamborite (17)	$\text{Ni}^{2+}_{1-x}\text{Co}^{3+}_x(\text{OH})_{2-x}(\text{SO}_4)_x\cdot n\text{H}_2\text{O}$ [$x \leq 1/3$; $n \leq (1-x)$]	Oxidized ore	<u>54</u>
48 Julienite (1)	$\text{Na}_2\text{Co}(\text{SCN})_4\cdot 8\text{H}_2\text{O}$	Incrustations on talc schist	<u>5</u>
49 Karpenkoite (1)	$\text{Co}_3(\text{V}_2\text{O}_7)(\text{OH})_2\cdot 2\text{H}_2\text{O}$	Oxidized ore	<u>55</u>
50 Kolwezite (6)	$\text{CuCoCO}_3(\text{OH})_2$	Oxidized ore	<u>5</u>
51 Leverettite (1)	$\text{Cu}_3\text{CoCl}_2(\text{OH})_6$	Oxidized ore	<u>56</u>
52 Metakirchheimerite (2)	$\text{Co}(\text{UO}_2)_2(\text{AsO}_4)_2\cdot 8\text{H}_2\text{O}$	Oxidized ore	<u>3,9</u>
53 Moorhouseite (12)	$\text{CoSO}_4\cdot 6\text{H}_2\text{O}$	Effluoresces with S	<u>1,2,3,9,11,14,29</u>
54 Neustädtelite (4)	$\text{Bi}_2\text{Fe}^{3+}(\text{Fe}^{3+},\text{Co})_2(\text{AsO}_4)_2(\text{O},\text{OH})_4$	Oxidized ore	<u>2,22</u>
55 Oursinite (1)	$\text{Co}(\text{UO}_2)_2(\text{SiO}_3\text{OH})_2\cdot 6\text{H}_2\text{O}$	Secondary mineral	<u>5</u>
56 Pakhomovskyite (1)	$\text{Co}_3(\text{PO}_4)_2\cdot 8\text{H}_2\text{O}$	Carbonatite alteration	<u>46</u>
57 Petewilliamsite (1)	$(\text{Ni},\text{Co})_{30}(\text{As}_2\text{O}_7)_{15}$	Oxidized ore	<u>48</u>
58 Pradetite (2)	$\text{CoCu}_4(\text{AsO}_4)_2(\text{AsO}_3\text{OH})_2\cdot 9\text{H}_2\text{O}$	Oxidized ore	<u>3,49</u>
59 Rappoldite (1)	$\text{PbCo}_2(\text{AsO}_4)_2\cdot 2\text{H}_2\text{O}$	Oxidized ore	<u>2</u>
60 Roselite (15)	$\text{Ca}_2\text{Co}(\text{AsO}_4)_2\cdot 2\text{H}_2\text{O}$	Oxidized ore	<u>1,2,4,8,9,34</u>
61 Roselite-beta (9)	$\text{Ca}_2\text{Co}(\text{AsO}_4)_2\cdot 2\text{H}_2\text{O}$	Oxidized ore	<u>1,2,4,8</u>
62 Schneebergite (3)	$\text{BiCo}_2(\text{AsO}_4)_2(\text{OH})\cdot \text{H}_2\text{O}$	Oxidized ore	<u>2</u>
63 Smolyaninovite (11)	$\text{Co}_3\text{Fe}^{3+}_2(\text{AsO}_4)_4\cdot 11\text{H}_2\text{O}$	Oxidized ores	<u>1,2,4,24,34,35,51</u>
64 Spherochalcite (21)**	CoCO_3	Carbonate alteration	<u>1,2,5,11,14,22,26,54</u>

65 Thérèsemagnanite (2)	$\text{Co}_6\text{SO}_4(\text{OH})_{10}\cdot 8\text{H}_2\text{O}$	Oxidized ore	<u>49</u>
66 Wupatkiite (7)	$\text{CoAl}_2(\text{SO}_4)_4\cdot 22\text{H}_2\text{O}$	Oxidized zone	<u>4,14</u>

*Locality data compiled from MinDat.org as of 1 March 2016.

**Sphero-cobaltite may also occur as a primary mineral.

Table 2. Mineral localities with the greatest diversity of cobalt minerals, number and identity of Co minerals, their lithological settings, principal Co-mineral-forming elements, and deposit age. Listed are all localities with at least 7 different Co mineral species, as well as additional localities that yielded the type specimen for each of the 66 known Co minerals (Table 1). The identification key to numbers for Co mineral species appears in Table 1; type minerals are designated by an underscored number. Numbers corresponding to the 26 primary Co minerals appear in boldface.

Locality	#Co Minerals	Lithological Context (key elements)
1. Bou Azer District, Tazenakht, Ouarzazate Province, Morocco	22 (1,3,5,6,11,13,15,20,22 ; 28,30,32,33,35,36,44,45,53,60,61,63,64)	Veins in serpentinized ophiolite + diabase intrusions (Co,Ni,Cu,Ag with As>S).
2. Schneeberg District, Erzgebirge, Saxony, Germany	21 (4,6,20,22 ; <u>28</u> ,29,33,35, <u>36,37,40</u> ,44,45,53,54,59,60,61,62,63,64)	Hydrothermal Ag-Bi-Co-Ni-U veins with As > S.
3. Jáchymov District, Karlovy Vary Region, Czech Republic	18 (1,5,6,11,15,20,21,22 ; 28,29,34,35,41,44, <u>46</u> ,52,53,58)	Classic 5-element (Ag-Co-Ni-Bi-U) deposit; hydrothermal veins with S and As.
4. Mount Isa, Cloncurry area, Queensland, Australia	14 (1,6,8,11,21,26 ; 28,44,45,60,61,63,64,66)	Fe oxide-Cu-Au deposit. (Co-Ni-As-S-Sb)
5. Katanga Copper Crescent, Katanga, Democratic Republic of Congo	13 (3,4,15,20 ; 28,39, <u>43</u> ,44,45, <u>48</u> , <u>50</u> , <u>55</u> ,64)	U with Co-Ni-Fe-S mineralization (656 Ma), followed by Co-Cu-S; no Ni-As (602 Ma) Carbonate hosted.
6. Cobalt area, Timiskaming District, Ontario, Canada	12 (1,5,6,7,11,14,20,21,22 ; <u>31</u> ,44,45)	Hydrothermal Ag-Co-Ni-Cu-As-S quartz-calcite veins
7. Tunaberg, Nyköping, Södermanland, Sweden	11 (3,6,7,8,13,15,18,20,26 ; 43)	Co-Cu sulfides in a skarn zone
8. Ore dumps, Richelsdorf Smelter, Hesse, Germany	10 (4,5,6,20,22 ; 35,44,45,60,61)	Veins of Bi-Co-Ni-As-S (Cu-Pb-Zn)
9. Wittichen, Baden-Württemberg, Germany	10 (1,20,22 ; 28,35,44,45, <u>52</u> ,53,60)	Ag-Bi-Co-Ni-U mineralization in granite
10. Siegerland, North Rhine-Westphalia, Germany	11 (1,3,6,11,15,20,21,22 ; 28,29,44)	F-Cu-Co-S-As mines
11. Lavrion District, Attiki Prefecture, Greece	9 (20 ; 27,28,29,33,44,50,53,64)	Cu-Zn-Pb-Ag ore related to emplacement of Miocene granodiorite into marbles and other

		Mesozoic metamorphic rocks.
12. North Bohemia Uranium District, Liberec Region, Czech Republic	8 (3,4,6,14,20,21,22; 28)	Co-As-S
13. Nordmark District, Värmland, Sweden	8 (3,4,5,6,15,20,21; 44)	Metamorphosed Fe-Mn deposits; Pb-Mn-As-Sb; skarn mineralization; dolomite lens between potassic metarhyolites and basic rocks; Co incidental
14. Coconino County, Arizona, USA	8 (15,21,22; 29,44,53,64,66)	U mineralization
15. Broken Hill, Yancowinna Co., New South Wales, Australia	7 (6,8,15,20,22,26; 44)	Ag-Pb-Zn deposit; Co incidental
16. Port Radium District, Great Bear Lake, Northwest Territories, Canada	7 (6,11,20,22; 28,44,45)	U-Bi-Ag-Ni-Co with As-S
17. Shilu Mine, Changjiang County, Hainan Province, China	7 (3,4,6,7,15,21,22)	
18. Zhongtiaoshan ore field, Yuanqu County, Yuncheng Prefecture, China	7 (3,4,6,7,15,21; 44)	
19. Kuusamo, Northern Finland Region, Finland	7 (4,6,7,12,15,16,21)	Hosted by Paleoproterozoic Kuusamo Schist Au-Co-S with Ag-Cu-Mo-Ni-REE-U
20. Clara Mine, Wolfach, Baden-Wurttemberg, Germany	7 (1,4,6,20,22; 35,44)	Barite-fluorite veins with Cu-Ag-Pb hosted by gneiss and Triassic sandstone. Incidental Co-As-S.
21. Annaberg District, Erzgebirge, Saxony, Germany	7 (6,20,22; 28,35,44,45)	Vein deposits with Ag-Co-Ni-U. [Co-As] mineralization in Siluro-Devonian sediments
22. Marienberg District, Erzgebirge, Saxony, Germany	7 (6,15,20,22; 44,54,64)	Vein deposits of Fe-Ag-Sn-U-F. Co-Ni-S-As mineralization incidental.
23. Kamsdorf, Saalfeld, Thuringia, Germany	7 (6,20,22; 28,29,44,45)	
24. Palhal Mine, Branca, Albergaria-a-	7 (1,22; 28,29,44,45,63)	

- Velha, Aveiro District, Portugal
25. Hnúšťa, Rimavská Sobota County, Banská Bystrica Region, Slovakia 7 (**3,8,11,15,21,26**; 44)
 26. Cerro Minado Mines, Cuesta Alta, Almería, Andalusia, Spain 7 (**3,6,21**; 28,44,45,64)
 27. San Carlos Mine, San Juan de Plan, Huesca, Aragón, Spain 7 (**1,6,11,15,20,22**; 44)
 28. Banat Mountains, Caras-Severin Co, Romania 5 (**1,3,6,7,11**)
 29. Walton Barite Mine, Hants County, Nova Scotia, Canada 3 (27,44,53) Barite, Pb-Zn-Cu; Co incidental
 30. Bieber, Hesse, Germany 5 (**6,20,22**; 29,44)
 31. Lautenthal, Harz, Lower Saxony, Germany 6 (**2,6,10,23**; 43,44) Veins with in Ag-Pb-Zn-Cu-Ba in greywacke, associated with diabase intrusions. (Co-Ni-S-Se)
 32. Sykesville, Carroll County, Maryland 3 (**3,15,21**)
 33. Dolores Prospect, Murcia, Spain 3 (32,44,45)
 34. Olary Province, South Australia, Australia 6 (**6,22**; 33,44,60,63) Calc-silicate hosted Cu-S/Co-As
 35. Bästnas mines, Riddarhyttan, Västmanland, Sweden 6 (**3,6,15**; 34,44,63)
 36. Schwarzenberg District, Erzgebirge, Saxony, Germany 5 (**20,22**; 28,35,44)
 37. Cerro de Cacheuta, Mendoza, Argentina 1 (39) Cu deposit in ultramafics; Co incidental
 38. Varislahti deposit, South Karelia, 1 (**7**) Cu-Ni-Co-Fe-S

Southern Finland, Finland

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| 39. White Canyon District,
San Juan County, Utah | 6 (21 ; 28,29, <u>39</u> , <u>41</u> ,44) | |
| 40. Barberton District, Mpumalanga
Province, South Africa | 6 (6,7,15,20,22 ; <u>42</u>) | |
| 41. Noril'sk, Taimyr Peninsula,
Eastern-Siberian Region, Russia | 4 (6,9,15,20) | |
| 42. Huasco, Atacama Region, Chile | 6 (3,6,11 ; 28,44,45) | Cu with Co-U veins in diorite. |
| 43. Jhunjhunu District, Rajasthan, India | 2 (6,12) | Cu-Co-S veins |
| 44. Mattagami, Nord-du-Québec,
Québec, Canada | 2 (6,16) | Co |
| 45. Witwatersrand field, Gauteng/
Mpumalanga Provinces, South Africa | 2 (6,17) | |
| 46. Kovdor/Khibiny massifs, Kola
Peninsula, Northern Region, Russia | 6 (6,7,20 ; 27,44, <u>56</u>) | |
| 47. Kenora District, Ontario, Canada | 3 (3,6,19) | Cu-Pb-Zn with Co-Ni-S-As-Sb |
| 48. Johanngeorgenstadt District,
Erzgebirge, Saxony, Germany | 5 (20,22 ; 28,44, <u>57</u>) | |
| 49. Cap Garonne Mine, Var, Provence-
Alpes-Côte d'Azur, France | 4 (44,53, <u>58,65</u>) | Co minerals all secondary |
| 50. Skutterud and nearby Mines,
Modum, Buskerud, Norway | 6 (3,6,11,15,22 ; 44) | Cu-Co-S vein deposits |
| 51. Cobalt deposits, Tuva Republic,
Eastern-Siberian Region, Russia | 6 (1,11,20,22 ; 44, <u>63</u>) | |

52. Beaverlodge Region, Saskatchewan, Canada	2 (6 , <u>24</u>)	Co-Ni-Co-S-As-Se hydrothermal veins
53. Red Hill, Wairau Valley, South Island, New Zealand	1 (<u>25</u>)	Serpentinized zone
54. Northern Apennines, Emila-Romagna, Italy	3 (15 ; <u>47</u> ,64)	
55. Yellow Cat District, Grand County, Utah	1 (<u>49</u>)	
56. El Tamarugal Province, Tarapaca Region, Chile	2 (6 ; <u>51</u>)	
