1 Revision 1: High Pressure Behavior of the Polymorphs of FeOOH

2 Mary M. Reagan<sup>1</sup>, Arianna E. Gleason<sup>2</sup>, Luke Daemen<sup>3</sup>, Yuming Xiao<sup>4</sup>, Wendy L.
3 Mao<sup>1,5</sup>

<sup>1</sup>Department of Geological Sciences, Stanford University, Stanford, CA 94305 <sup>2</sup>Shock
and Detonation Physics, LANL, Los Alamos, NM 87545 <sup>3</sup> Spallation Neutron Source,
Oak Ridge National Laboratory, Oak Ridge, TN, 37830 <sup>4</sup>Advanced Photon Source,
Argonne National Laboratory, Lemont, IL, 60439 <sup>5</sup>Photon Science, SLAC National
Accelerator Laboratory, Menlo Park, CA 94025

9

### Abstract:

10 The high-pressure structural and electronic behavior of  $\alpha$ -,  $\beta$ -, and  $\gamma$ -FeOOH were studied 11 in-situ using a combination of synchrotron X-ray diffraction (XRD) and X-ray emission 12 spectroscopy (XES). We monitored  $\alpha$ -FeOOH by XES as a function of pressure up to 85 13 GPa and observed an electronic spin transition that began at approximately 50 GPa, 14 which is consistent with previous results. In the  $\gamma$ -FeOOH sample, we see the initiation of 15 a spin transition at 35 GPa that remains incomplete up to 65 GPa.  $\beta$ -FeOOH does not 16 show any indication of a spin transition up to 65 GPa. Analysis of the high-pressure XRD 17 data shows that neither  $\beta$ -FeOOH nor  $\gamma$ -FeOOH transform to new crystal structures, and 18 both amorphize above 20 GPa. Comparing our EOS results for the  $\beta$  and  $\gamma$  phases with 19 recently published data on the  $\alpha$  and  $\varepsilon$  phases, we found that  $\beta$ -FeOOH exhibits distinct 20 behavior from the other three polymorphs, as it is significantly less compressible and 21 does not undergo a spin transition. A systematic examination of these iron hydroxide 22 polymorphs as a function of pressure can provide insight into the relationship between electronic spin transitions and structural transitions in these OH- and Fe<sup>3+</sup> bearing phases 23

that may have implications on our understanding of the water content and oxidation stateof the mantle.

26

### Introduction:

27 Iron hydroxides, including FeOOH and its polymorphs, are common on the surface of the 28 Earth, where they are abundant in soils, aquifers and sediments. FeOOH has four 29 polymorphs, three of which are naturally occurring: goethite ( $\alpha$ -FeOOH), akaganeite ( $\beta$ -FeOOH), and lepidocrocite ( $\gamma$ -FeOOH), A fourth polymorph,  $\epsilon$ -FeOOH, can be 30 31 synthesized at high pressure (Bolotina et al., 2008, Gleason et al., 2008, Voigt et al., 32 1981). Goethite is the thermodynamically stable phase at ambient conditions, whereas 33 akaganeite is rare and forms in Cl-rich environments like hot brines and rust in marine 34 environments. Lepidocrocite occurs in rocks, soils and rust and is often an oxidation product of Fe<sup>2+</sup>. Their structures consist of corner-linked double bands of  $FeO_3(OH)_3$ 35 36 octahedra (Figure 1).  $\alpha$ -FeOOH has double bands of edge sharing octahedra that form 37 2x1 channels.  $\beta$ -FeOOH also has double bands; however, they are arranged in a circular 38 shape forming large 2x2 channels, which are stabilized by the presence of a variable 39 molecule or ions such as H<sub>2</sub>O, OH, Cl<sup>-</sup>, or NO<sub>3</sub>. The  $\gamma$ -FeOOH bands are connected via 40 hydrogen bonds (OH-O) and form corrugated 2d layers perpendicular to the b-axis. The 41 high-pressure  $\varepsilon$ -FeOOH phase is a slightly distorted rutile structure with corner-linked 42 single bands of edge-shared octahedra parallel to the c-axis (Otte et al., 2009). It is 43 isostructural with  $\delta$ -AlOOH, a phase that may transport hydrogen deep within the planet 44 and potentially down to the core-mantle boundary (Sano et al., 2004; Suzuki, 2010).

45 Based on crystal field theory, the distribution of electrons in the d orbitals in transition

46 metal complexes depends on the ligand field geometry and the metal d-electron 47 configurations. The spin state of iron is determined by the difference between the energy 48 levels of  $\Delta_c$  (the crystal field splitting parameter) and the spin pairing energy ( $\Lambda$ ). Under 49 ambient conditions,  $\Delta_c$  for most octahedrally coordinated iron compounds is lower than  $\Lambda$ . 50 This results in favorable energetic conditions where 3d electrons occupy both the  $t_{2g}$ 51 orbitals and the higher energy  $e_g$  orbitals with unpaired spins. With increasing pressure, 52 the crystal field splitting energy increases with respect to the spin pairing energy, this can 53 result in the spin pairing of the electrons in the lower energy  $t_{2g}$  orbital. Thus spin 54 transitions in iron in lower mantle minerals may result in significant changes in their 55 physical and chemical properties (Speziale et al., 2005)-- including their density, bulk 56 modulus, seismic velocities, electrical conductivity, radiative heat transfer, and element 57 partitioning (Lin et al., 2013; Lin et al., 2008; Lin et al., 2007; Speziale et al., 2005; 58 Stackhouse et al., 2007). Previous work (Gleason et al., 2013; Xu et al., 2013) showed 59 that spin transitions take place in  $\alpha$ -FeOOH and  $\epsilon$ -FeOOH. Gaining insight into the high 60 pressure behavior and spin state of the polymorphs of FeOOH can improve our 61 understanding of more complex hydrogen bearing compounds that may be common in 62 the Earth's deep interior (Williams et al., 2001).

63

#### **Experimental methods:**

### 64 Sample synthesis:

65 α, β and γ- FeOOH were synthesized using the methods in Schertmann (2000), with β-66 FeOOH was prepared following a slight modification. After dissolving 54.6 g 67 FeCl3.6H2O (0.2 mol) in 200 mL of deionized water in a 500 mL polyethylene bottle, a 68 solution of 6 g NaOH in 120 mL deionized water was added to the iron chloride solution.

69 A brown precipitate forms, which redissolves upon shaking. The resulting clear, brown 70 solution is allowed to age at room temperature for 3 days, at which point 40 mL of a 10 71 M solution of NaOH in deionized water is added rapidly with vigorous stirring. The 72 resulting solution is heated to 80 °C in the closed polyethylene bottle until a brown 73 precipitate starts appearing. This requires 6 to 7 days. After the solution becomes a 74 cloudy suspension, heating is continued for 4 days. The suspension of  $\beta$ -FeOOH is 75 difficult to filter by conventional means and the solid is separated from the supernatant 76 liquid by centrifugation. The solid is washed repeatedly by resuspension in 250 mL of 77 water, stirring for several minutes, and centrifugation. This washing operation is repeated 78 several times until the supernatant liquid no longer gives a positive test for chloride by 79 precipitation with a silver nitrate solution. The solid material is then dried in a vacuum 80 oven at 40 °C.

## 81 Chloride analysis:

82 The simplest method for chloride determination is Mohr's titration with aqueous silver 83 nitrate, using sodium chromate as an indicator (Skoog et al., 1996). The method requires 84 to bring  $\beta$ -FeOOH in solution. Hydrochloric acid dissolves  $\beta$ -FeOOH quite readily, but, 85 obviously, cannot be used for the sample preparation in the determination of chloride in 86  $\beta$ -FeOOH. Furthermore, any Broensted acid used to dissolve  $\beta$ -FeOOH cannot introduce 87 an anion that forms an insoluble salt with silver. Nitric acid would be ideal, but does not 88 dissolve  $\beta$ -FeOOH at or above room temperature. Sample preparation proceeded as 89 follows. A small amount of  $\beta$ -FeOOH (circa 50 mg) was placed in the teflon liner of a 23 90 mL Parr autoclave. Deionized water (8mL) was placed in the vessel, then 4 mL of 91 concentrated nitric acid. Four drops of triflic acid were added, after which the autoclave

92 was sealed and placed in an oven. The temperature was raised progressively to 180 °C 93 over a period of one day and the autoclave was kept at this temperature for two more 94 days. The autoclave was allowed to cool down to room temperature before opening. The 95 teflon liner contained a clear liquid with no visible traces of solid. This method of sample 96 preparation has the advantage of preventing any loss of chlorine in the form of HCl. The 97 clear solution was transferred to a 250 mL beaker, diluted with 100 mL of water, and pH 98 was raised to about 6 with the addition of small amounts of sodium bicarbonate. At this 99 pH, iron precipitates as iron hydroxide. This solid is filtered, washed carefully, and the 100 filtrate is then used for Mohr titration with a 20 mM aqueous solution of silver nitrate. 101 Three samples were prepared in this fashion for chloride determination. If a nominal 102 formula for  $\beta$ -FeOOH is assumed to be, Fe O (OH) x Cl (1-x), then the titration results 103 gave x = 0.84 + -0.015, which indicates that some 15% of the hydroxide ions are 104 replaced by chloride ions in the  $\beta$ -FeOOH structure. This is consistent with previous 105 observations.

## 106 Sample Preparation

107 Powdered samples of synthesized  $\alpha$ -FeOOH,  $\beta$ -FeOOH, and  $\gamma$ -FeOOH (Gerth, 1990; 108 Kosmulski et al., 2003; Wang et al., 2004) were loaded separately into a beryllium (for 109 XES) or tungsten and stainless steel (for XRD) gasket. For  $\alpha$ -FeOOH, a 100  $\mu$ m hole 110 served as the sample chamber while a sample chamber of 120  $\mu$ m was used for  $\beta$ -111 FeOOH, and  $\gamma$ -FeOOH. All samples were compressed between 300  $\mu$ m diamond culets in 112 a symmetric diamond anvil cell (DAC). No pressure-transmitting medium was used for 113 the XES experiments to maximize signal, and silicone oil was used for the XRD 114 experiments. Ruby and gold (Mao et al., 1978, Heinz et al., 1984) were used as pressure

115 calibrants for the XRD experiments, while ruby only was used for the XES experiments.

# 116 Synchrotron X-ray Experiments

117 Angle-dispersive powder XRD measurements were collected at beamline 12.2.2 of the 118 Advanced Light Source (ALS), Lawrence Berkeley National Laboratory (LBNL) where a 119 wavelength of 0.4959 Å x-rays with a beam spot of 10 x 10  $\mu$ m were used and a smaple 120 to detector distance of 354.41 mm, and at HP-CAT, beamline 16-BM-D at the Advanced 121 Photon Source (APS), Argonne National Laboratory (ANL) where a wavelength of 122 0.4246 Å x-rays with a beam spot size of 5 x 5  $\mu$ m were used and a sample to detector 123 distance of 320.66mm. XRD data for  $\beta$ -FeOOH was collected at APS, while data for  $\gamma$ -124 FeOOH was collected at ALS. All data were collected on a MAR345 image plate and cell 125 parameter refinements were carried out using the MAUD program (Lutterotti et al., 126 1997). The Fe K $\beta$  XES spectra of  $\alpha$ ,  $\beta$ , and  $\gamma$ -FeOOH were collected at HP-CAT, 127 beamline 16ID-D, APS, ANL. The incident X-ray energy were monochromatized using a 128 Si(111) double crystal monochromator and was centered at 11.3 keV, and the scans were 129 set relative to 7.058 keV with a range of -40 to +25 eV. Table 2 provides a summary of 130 the experimental techniques.

131

#### **Results:**

### 132 X-ray Diffraction:

The 2D patterns were integrated using Fit 2D (Hammersley, 1998) and Le Bail refinements of the integrated patterns were carried out to determined the unit cell volumes at each pressure using MAUD program (Lutterotti et al., 1997). Background fits were determined using a 3<sup>rd</sup> order polynomial that was manually set. Lattice parameters

137 were refined to a goodness of fit of 0.6 % ( $\beta$ ) and 0.5 % ( $\gamma$ ) as determined by MAUD. For 138  $\gamma$ -FeOOH, the (200) peak was used to constrain the *a* lattice parameter at each pressure 139 step. Analysis of the XRD data shows that neither  $\beta$ -FeOOH or  $\gamma$ -FeOOH undergo any 140 first order structural transitions over the pressure range studied, and both show evidence of amorphization above ~20 GPa (Figures 2, 3). We fit our pressure-volume data to a  $2^{nd}$ 141 142 order Birch-Murnaghan equation of state (EOS) (Birch, 1978), and found that  $\beta$ -FeOOH has a  $V_0$  of 336.6(3) Å<sup>3</sup>, and  $K_0$  value of 284(1) GPa. The  $\gamma$ -FeOOH phase has a  $V_0$  of 143 147.76(2) Å<sup>3</sup> and a  $K_0$  value is 104(1) GPa. These results are summarized in Table 1 and 144 145 plotted in Figures 4 and 5, where the  $\beta$  and  $\gamma$  phases are compared with recently 146 published work on the  $\alpha$ - (Xu et al., 2013) and  $\varepsilon$ - (Gleason et al., 2013) phases.

# 147 X-ray Emission Spectroscopy:

148 The spin state of iron is monitored through K $\beta$  XES, where a K shell electron core-hole is 149 created when an electron absorbs an X-ray photon and is emitted from the atom. This is 150 followed by a 3p electron falling to the K shell, and in transition metals (e.g., iron), an 151 exchange interaction occurs between the 3p core hole and the unpaired 3d shell when a 152 material is in high spin. The resultant emission spectrum consists of  $K\beta_{1,3}$  peak and the 153 smaller satellite peak, K $\beta$ ' that arises from the exchange interaction between 3p core hole 154 and the 3d shell. We monitored  $\alpha$ -FeOOH as a function of pressure up to 85 GPa and 155 observed that it undergoes a spin transition beginning at  $\sim$ 50 GPa. This is seen in the 156 evolution of the K $\beta$  emission spectra where a clear reduction of K $\beta$ ' was observed 157 (Figure 6). Each spectrum at a given pressure was summed and normalized using the 158 integrated absolute difference method (IAD) (Mattila et al., 2007). Our IAD values 159 indicate the high to low spin transition occurred over a range of 50-70 GPa (Figure 7).

160 This is in good agreement with a previous study that found a high to low spin cross over 161 above 45 GPa (Xu et al., 2013). In the  $\gamma$ -FeOOH sample, we see a spin transition 162 beginning at 35 GPa that remains incomplete to 65 GPa, which is evidenced by the partial 163 reduction of the satellite peak (Figure 8). The  $\beta$ -FeOOH sample does not show any 164 indication of a spin transition for pressures up to the highest pressure investigated, 65 165 GPa (Figure 9). We note that although both the  $\beta$  and the  $\gamma$  phases become amorphous at 166 pressures above 20 GPa, the study of the spin state provides valuable information about 167 the electronic configuration despite the loss of the long-range order.

#### 168 **Discussion**:

# 169 At ambient pressure, the polymorphs of FeOOH are made up of asymmetric octahedra of $Fe^{3+}_{HS}$ (O<sub>1</sub>-H)<sub>3</sub>(O<sub>2</sub><sup>...</sup>H)<sub>3</sub> with two distinct types of oxygen bonding, the shorter O<sub>1</sub>-H and 170 171 the longer $O_2^{\dots}H$ . These bond lengths vary as they are subjected to pressure, with the 172 length of the Fe-O<sub>1</sub> and Fe-O<sub>2</sub> bonds and the O<sub>1</sub> <sup>...</sup> O<sub>2</sub> distance affecting the spin state. 173 These previous studies (Xu et al., 2013) have also linked the electronic spin transition 174 with the symmetrization of the hydrogen bonds in $\alpha$ -FeOOH. Increasing pressure leads to 175 a shift in the proton position where the $O_1$ -H and $O_2$ <sup>...</sup>H bonds become more equal in 176 length. Single-crystal studies show that at 45 GPa, there is a large drop in volume that corresponds to a reduction of the $Fe^{3+}$ radius (Xu et al., 2013). The asymmetry of the 177 178 polyhedra is reduced as the Fe-O<sub>1</sub> and Fe-O<sub>2</sub> bonds are close to equal and the $O_1$ -Fe-O<sub>2</sub> 179 angle becomes 180°. α-FeOOH undergoes a spin transition at 47 GPa, with hydrogen 180 bond symmetrization occurring at pressures above 50 GPa. In this case, the spin 181 crossover leads to the hydrogen bond symmetrization (Xu et al., 2013). ε- FeOOH's

182	behavior shows an opposite order of events, with increasing pressure: the spin transition
183	starts at 49 GPa, while DFT calculations show that its hydrogen bonds symmetrize at a
184	lower pressure of 43 GPa. The hydrogen bond symmetrization results in the elongation of
185	the $O_1$ -H and the compression of $O_2^{\cdots}$ H, with hydrogen reaching an equidistant position
186	at ~1.2 Å and the $O_1$ -H <sup></sup> $O_2$ bond angle adjusting to 180°. (Gleason et al., 2013).
187	Interestingly, previous DFT calculations show that in $\gamma$ - FeOOH the O <sub>1</sub> -H remain
188	constant at ~ 1.2 Å, while at 15 GPa, the $O_2^{\cdots}H$ bond reduces from 3.1 Å to 2.9 Å. (Otte
189	et al., 2009). Its spin transition begins at higher pressure of 36 GPa where the sample is
190	amorphous and remains incomplete up to 65 GPa, the highest pressure we studied (Figure
191	8). Previous DFT calculations for $\beta$ -FeOOH, indicate that its O-H bonds remain nearly
192	constant with increasing pressure up to 20 GPa (Otte et al., 2009). The DFT calculations
193	for the $\beta$ and $\gamma$ phases only go up to 20 GPa, so it is not known how their bond lengths
194	are changing at higher pressures, or at the onset of the spin transition in $\gamma$ -FeOOH.

195 The  $\beta$  phase has a distinct behavior when compared to the other polymorphs of FeOOH. 196 It is the only polymorph that does not undergo an electronic spin transition over the range 197 of pressures studied. DFT calculations for  $\beta$ -FeOOH show that its oxygen-hydrogen 198 bonds remain constant with increasing pressure. It is significantly less compressible as 199 indicated by its higher  $K_0$  value compared to the other polymorphs (Table 1). Its 200 electronic configuration remaining in high spin maybe related to its incompressibility 201 thus leading to polyhedra remaining more asymmetric and less compressed. There is 202 small amount of Cl<sup>-</sup> present in the  $\beta$ -FeOOH that stabilizes the 2x2 channels in its 203 structure and is therefore less compressible allowing it to remain in the high spin state with increasing pressure.  $\beta$ -FeOOH and to a lesser extent  $\gamma$ -FeOOH also have a larger volume per formula unit when compared with  $\alpha$ - and  $\epsilon$ -FeOOH (Table 1, Figure 4). This may also contribute to why both  $\beta$ - and  $\gamma$ -FeOOH become amorphous at high pressure as they are energetically less stable than their denser polymorphs but the transition is hindered until they become mechanically unstable and amorphize. Future computational work to look at the hydrogen bonding as a function of pressure for  $\beta$ - and  $\gamma$ -FeOOH would provide further insight into the high-pressure behavior of these polymorphs.

211

#### **Implications:**

It has been proposed that the electronic transition  $inFe^{3+}$  and the symmetrization of the 212 213 hydrogen bond are closely connected for the FeOOH polymorphs (Gleason et al., 2013; 214 Xu et al., 2013). The spin transition is driven by the decreasing  $\text{Fe-O}_{1,2}$  bonds and the 215 change in the  $O_1 \, O_2$  distance, which leads to an environment that increases the crystal 216 field splitting parameter until it is more energetically favorable for the low spin state. In 217 full spin transitions, this change is often accompanied by the hydrogen bond 218 symmetrization. In both  $\alpha$ -FeOOH and  $\epsilon$ -FeOOH, a complete spin transition is coupled 219 with hydrogen bond symmetrization. In  $\gamma$ -FeOOH, there is only a partial shift in its 220 hydrogen atom, which maybe a reflection that the polymorph does not undergo a full spin 221 transition.  $\beta$ -FeOOH also shows this connection with its bond length remaining relatively 222 unchanged and a lack of a spin transition. Hydrogen bond symmetrization has also been 223 observed at high pressure in  $\alpha$ -AlOOH, which is isostructural with  $\alpha$ -FeOOH and is 224 present in subducting sedimentary rocks. Its hydrogen bonds become increasingly more 225 symmetric up through the studied range of 50 GPa (Friedrich et al., 2007). Other oxy-

226 hydroxide transition metal compounds may also undergo hydrogen bond symmetrization 227 when subjected to high pressure, and further work is needed to understand the possible 228 connection between hydrogen bond symmetrization and electronic structure changes. The 229 symmetrization of hydrogen bonds is expected to have an effect on the elasticity and 230 plasticity of hydrogen-bearing materials. Quantifying the stability of the hydrogen bonds in relation to hydrous Fe-bearing minerals with the potential for undergoing an electronic 231 232 transition can give insight into the behavior of hydrogen in an Fe-rich environment at 233 deep Earth conditions (Benoit et al., 1998; Gilli et al., 2009; Friedrich et al., 2007; 234 Holzapfel, 1972; Xu et al., 2013).

235 Iron is the most abundant transition metal in the deep Earth, and iron-bearing materials 236 may transition to a low spin state at shallower depths as their water content increases 237 (Frost & McCammon, 2008; Otte et al., 2009; Xu et al., 2013). The experimental work on 238 the polymorphs of FeOOH indicate that increasing pressure can have dramatic effects on 239 the electronic state and the crystal structure, including the nature of the hydrogen bonds. 240 This effect may be ubiquitous in transition metal bearing compounds that contain water-241 and relevant to the deep Earth. Ferric iron bearing silicates may play an important role in 242 transporting water into the deep mantle. Changes to their spin states and hydrogen bonds 243 might then have an influence on the water dynamics and balance within the deep Earth 244 (Williams and Hemley, 2001). This has implications for the oxidation state of the Earth's mantle, where the redox state controls  $Fe^{3+}$  and OH<sup>-</sup> content within minerals, as well as, 245 246 water partitioning in fluids/melts and minerals (Frost and McCammon, 2008).

247

#### **Acknowledgements:**

- 248 M.M. Reagan, A.E. Gleason, and W.L. Mao are supported by the NSF Geophysics
- 249 Program (EAR-1446969). We thank Paul Chow (APS), Alastair MacDowell (ALS), and
- 250 Jinyuan Yan (ALS) for their assistance with the synchrotron experiments, and Jinfu Shu
- 251 (Geophysical Lab) for help with sample loading. Portions of this work were performed at
- beamline 12.2.2 of ALS, LBNL. ALS is supported by the Director, Office of Science,
- 253 Office of Basic Energy Sciences (BES), of the U.S. Department of Energy (DOE) under
- 254 Contract No. DE-AC02-05CH11231. Portions of this work were also performed at
- 255 HPCAT (Sector 16), APS, ANL. HPCAT operations are supported by DOE-NNSA under
- Award No. DE-NA0001974 and DOE-BES under Award No. DE-FG02-99ER45775,
- 257 with partial instrumentation funding by NSF. APS is a DOE-BES User Facility operated
- for the DOE Office of Science by ANL under Contract No. DE-AC02-06CH11357.

## 259 **References:**

- Bolotina, N., Molchanov, V., Dyuzheva, T., Lityagina, L., and Bendeliani, N., (2008).
  Single- crystal structures of high-pressure phases FeOOH, FeOOD, and GaOOH.
  Crystallography Reports. 53 (6), 960–965.
- 263
- Benoit, M., Marx, D., and Parrinello, M. (1998). Tunnelling and zero-point motion in
   high-pressure ice. Nature, 392(6673), 258–261.
- Birch, F. (1978). Finite Strain Isotherm and Velocities for Single-Crystal and
  Polycrystalline NaCl at High Pressures and 300°K. Journal of Geophysical
  Research, 83(B3), 1257–1268.
- Friedrich, A., Haussül, E., Boehler, R., Morgenroth, W., Juarez-Arellano, E. a., and
  Winkler, B. (2007). Single-crystal structure refinement of diaspore at 50 GPa.
  American Mineralogist, 92(10), 1640–1644.
- Frost, D. J., and McCammon, C. a. (2008). The Redox State of Earth's Mantle. Annual
  Review of Earth and Planetary Sciences, 36(1), 389–420.
- Gerth, J. (1990). Unit-cell Dimensions of Pure and Trace Metal-associated Goethites.
  Geochimica et Cosmochimica Acta, 54(2), 363–371.

- Gilli, G. and Gilli, P. (2009) The Nature of the Hydrogen Bond, Oxford University Press,
   New York.
- Gleason, A. E., Jeanloz, R., and Kunz, M. (2008). Pressure-temperature stability studies
   of FeOOH using X-ray diffraction. American Mineralogist, 93(11-12), 1882–1885.
- Gleason, A. E., Quiroga, C. E., Suzuki, A., Pentcheva, R., and Mao, W. L. (2013).
  Symmetrization driven spin transition in ε-FeOOH at high pressure. Earth and
  Planetary Science Letters, 379, 49–55.
- Hammersley, A. P. (1998). Fit2D: V99.129 Reference Manual Version 3.1. Internal
  Report ESRF 98 HA01.
- Heinz, D., and Jeanloz, R. (1984). The equation of state of the gold calibration standard.
  Journal of Applied Physics, 55(June 1983), 885–893.
- Holzapfel, W. B. (1972). Symmetry of the hydrogen bonds in Ice VII. Journel of
  Chemical Physics, 56(2), 712–715.
- Kosmulski, M., Maczka, E., Jartych, E., and Rosenholm, J. B. (2003). Synthesis and
  characterization of goethite and goethite-hematite composite: Experimental study
  and literature survey. Advances in Colloid and Interface Science, 103(1), 57–76.
- Lin, J.F, Speziale, S., Mao, Z., and Marquardt, H. (2013). Effects of the Electronic Spin
  Transitions of Iron in Lower Mantle Minerals : Implications for Deep Mantle
  Geophysics and Geochemistry. Reviews of Geophysics, 51(2), 244–275.
- Lin, J.F., and Tsuchiya, T. (2008). Spin transition of iron in the Earth's lower mantle.
  Physics of the Earth and Planetary Interiors, 170(3-4), 248–259.
- Lin, J.F., Vankó, G., Jacobsen, S. D., Iota, V., Struzhkin, V. V., Prakapenka, V. B.,
  Kuznetsov, A., and Yoo, C.-S. (2007). Spin transition zone in Earth's lower mantle.
  Science, 317(5845), 1740–3.
- Lutterotti, L., Matthies, S., Wenk, H.-R., Schultz, A. S., and Richardson, J. W. (1997).
   Combined texture and structure analysis of deformed limestone from time-of-flight neutron diffraction spectra, Journel of Appied Physics, 81, 594–600.
- Mao, H. K., Xu, J., and Bell, P. M. (1978). High Pressure Physics: Sustained Static
   Generation of 1.36 to 1.72 Megabars. Science. 200(4346).
- Mattila, A., Pylkkänen, T., Rueff, J-P., Huotari, S., Vankó, G., Hanfland, M., Lehtinen,
   M., and Hämäläinen, K. (2007). Pressure induced magnetic transition in siderite
   FeCO3 studied by x-ray emission spectroscopy. Journal of Physics, 19(38), 386206.

308 309 310	Momma, K., and Izumi, F., (2008). VESTA: a three-dimensional visualization system for electronic and structural analysis. Journel of Appied Crystallography, 41, 653–658.386206.
311 312 313	Otte, K., Pentcheva, R., Schmahl, W. W., and Rustad, J. R. (2009). Pressure-induced structural and electronic transitions in FeOOH from first principles. Physical Review B, 80(20), 205116.
314 315 316 317	Sano, A., Ohtani, E., Kubo, T., and Funakoshi, K. I. (2004). In situ X-ray observation of decomposition of hydrous aluminum silicate AlSiO3OH and aluminum oxide hydroxide d-AlOOH at high pressure and temperature. Journal of Physics and Chemistry of Solids, 65, 1547–1554.
318 319	Skoog, D.A., West, D.M., and Holler, F.J., (1996). Fundamentals of Analytical Chemistry, 7th Edition, Thomson Learning, Inc, USA.
320 321 322	Speziale, S., Milner, a, Lee, V. E., Clark, S. M., Pasternak, M. P., and Jeanloz, R. (2005). Iron spin transition in Earth's mantle. Proceedings of the National Academy of Sciences of the United States of America, 102(50), 17918–22.
323 324 325	Stackhouse, S., Brodholt, J. P., and Price, G. D. (2007). Electronic spin transitions in iron-bearing MgSiO <sub>3</sub> perovskite. Earth and Planetary Science Letters, 253(1-2), 282–290.
326 327	Suzuki, A. (2010). High-pressure X-ray diffraction study of epsilon-FeOOH. Physics and Chemistry of Minerals, 37(3), 153–157.
328 329	Schwertmann, U., and R.M. Cornell, R.M., (2000). The Iron Oxides in the Laboratory, 2nd edition, Wiley-VCH, Weinheim.
330 331	Voigt, R. and Will, G. (1981). The system Fe2O3-H2O under high pressures. Neues Jahrbuch fur Mineralogie, 2, 89–96.
332 333 334	Wang, X., Chen, X., Gao, L., Zheng, H., Ji, M., Tang, C., Shen, T., and Zhang, Z., (2004). Synthesis of b -FeOOH and a -Fe2O3 nanorods and electrochemical properties of b -FeOOH, Journel of Material Chemistry, 905–907.
335 336	Williams, Q., and Hemley, R. (2001). Hydrogen in the deep Earth. Annual Review of Earth and Planetary Sciences, 29. 365–418.
337 338 339 340	Xu, W., Greenberg, E., Rozenberg, G. K., Pasternak, M. P., Bykova, E., Boffa-Ballaran, T., Dubrovinsky, L., Prakapenka, V., Hanfland, M., Vekliova, O. Simak, S., and Abrikosov, I. A. (2013). Pressure-Induced Hydrogen Bond Symmetrization in Iron Oxyhydroxide. Physical Review Letters, 111(17), 175501.

# 342 Tables:

Table 1. A comparison of unit cell volume, bulk modulus, and the spin transition
 pressure for the FeOOH polymorphs. The pressure-volume data for all the phases were fit
 to a 2<sup>nd</sup> order Birch-Murnaghan equation of state.

Phase	<i>V</i> <sub>0</sub> / f.u. (Å <sup>3</sup> )	<i>K</i> <sub>0</sub> (GPa)	Spin Transition Pressure (GPa)
α*	34.6 (3)	120 (3)	49
β	42.1 (3)	283.6 (11)	N/A
γ	36.9 (2)	103.5 (10)	36
£**	33.2 (5)	158 (5)	49

- 348 \* Xu et al., 2013
- 349 \*\*Gleason et al., 2008

# **Table 2**. A comparison of the experiments performed in this study.

# 362

	Technique (Beamline)	Specifications			
Phase		Gasket	Pressure Medium	X-ray wavelength / energy	Highest Pressure (GPa)
α	XES (16-IDD, APS)	Be	None	11.3keV	85
β	XES (16-IDD, APS)	Be	None	11.3keV	65
β	XRD (16-BM-D, APS; 12.2.2 ALS)	Stainless Steel	Silicone Oil	0.4133 Å 0.4246 Å	29
y	XES (16-IDD, APS)	Be	None	11.3keV	64
γ	XRD (12.2.2, ALS)	W	Silicone Oil	0.4246 Å	34

363

364

365 Figures:

This is a preprint, the final version is subject to change, of the American Mineralogist (MSA) Cite as Authors (Year) Title. American Mineralogist, in press. (DOI will not work until issue is live.) DOI: http://dx.doi.org/10.2138/am-2016-5449



**Figure 1.** The ambient pressure crystal structures of the FeOOH polymorphs showing

iron (gold), oxygen (red), hydrogen (light pink), chlorine (green), and FeO<sub>3</sub>(OH)<sub>3</sub>

375 octahedra (tan). (a)  $\alpha$ -FeOOH (goethite), (b)  $\beta$ -FeOOH (akaganeite), (c)  $\gamma$ -FeOOH

376 (lepidocrocite) and (d)  $\varepsilon$ -FeOOH. These structures were generated using VESTA

- 377 (Momma et al., 2008).
- 373





374

375

376 Figure 2. XRD spectra for  $\beta$ -FeOOH for selected pressures. The sample becomes

377 amorphous above ~17 GPa.  $\lambda = 0.4133$  Å 0.3 GPa- 15.1 GPa and 0.4246 Å above 17.3 378 GPa.

This is a preprint, the final version is subject to change, of the American Mineralogist (MSA) Cite as Authors (Year) Title. American Mineralogist, in press. (DOI will not work until issue is live.) DOI: http://dx.doi.org/10.2138/am-2016-5449



**Figure 3.** XRD spectra for  $\gamma$ -FeOOH for select pressures. It becomes amorphous above  $\sim 20$  GPa.  $\lambda = 0.4246$  Å.



**Figure 4.** EOS for the FeOOH polymorphs.  $\alpha$  and  $\varepsilon$  are plotted for comparison from previous published results (Xu et al., 2013; Gleason et al., 2008). Note the lower compressibility for  $\beta$ -FeOOH reflecting its higher *K* values. Both the  $\beta$  and  $\gamma$  phases become amorphous and their XRD spectra cannot be fit to a crystalline structure at higher pressures.

This is a preprint, the final version is subject to change, of the American Mineralogist (MSA) Cite as Authors (Year) Title. American Mineralogist, in press. (DOI will not work until issue is live.) DOI: http://dx.doi.org/10.2138/am-2016-5449



**Figure 5.** Volume per formula unit (f.u.) as a function of pressure.  $\beta$ –FeOOH has a significantly larger volume compared to the other three polymorphs.



**Figure 6**.  $K_{\beta}$  emission spectra as a function of pressure for  $\alpha$ -FeOOH.

393

391

394

395



**Figure 7**. Average spin number as a function of pressure for  $\alpha$ -FeOOH (green squares),  $\beta$ -FeOOH (black triangles),  $\gamma$ -FeOOH (blue diamonds) using the IAD method (Mattila et al., 2007). The low spin state for  $\alpha$ -FeOOH was used to calculate the IAD values for  $\beta$ -FeOOH and  $\gamma$ -FeOOH.  $\epsilon$ -FeOOH XES data from Gleason et al. (2013) was added (red circles). The spin transition for  $\alpha$ -FeOOH begins at ~50 GPa.





404 **Figure 8**.  $K_{\beta}$  emission spectra as a function of pressure for  $\gamma$ -FeOOH.

This is a preprint, the final version is subject to change, of the American Mineralogist (MSA) Cite as Authors (Year) Title. American Mineralogist, in press. (DOI will not work until issue is live.) DOI: http://dx.doi.org/10.2138/am-2016-5449



407

408 **Figure 9.**  $K_{\beta}$  emission spectra as a function of pressure for  $\beta$ -FeOOH.