1 **Revision 1** 2 3 Some thermodynamic properties of larnite (β -Ca₂SiO₄) constrained by high T/P4 experiment and/or theoretical simulation 5 ZHIHUA XIONG,^{1,2} XI LIU,^{1,2,*} SEAN R. SHIEH,³ SICHENG WANG,^{1,2} LINLIN 6 CHANG,⁴ JUNJIE TANG,^{1,2} XINGUO HONG,⁵ ZHIGANG ZHANG,⁶ AND 7 **HEJING WANG^{1,2}** 8 9 ¹ The Key Laboratory of Orogenic Belts and Crustal Evolution, Ministry of Education of 10 11 China, Beijing 100871, China ² School of Earth and Space Sciences, Peking University, Beijing 100871, China 12 ³ Department of Earth Sciences, University of Western Ontario, London, Ontario, N6A 13 14 5B7, Canada ⁴ College of Earth Science, University of Chinese Academy of Sciences, Beijing 100049, 15 16 China ⁵ Mineral Physics Institute, State University of New York, Stony Brook, New York 17 11974, USA 18 ⁶ Key Laboratory of Earth and Planetary Physics, Institute of Geology and Geophysics, 19 20 Chinese Academy of Sciences, Beijing 100029, China 21 22 *Corresponding Author: Xi Liu 23 Email: xi.liu@pku.edu.cn Tel: 86-10-6275-3585 Fax: 86-10-6275-2996

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25 Running Title: THERMODYNAMIC PROPERTIES OF LARNITE

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ABSTRACT

29 Pure larnite (β -Ca₂SiO₄; Lrn) was synthesized at 6 GPa and 1473 K for 6 hours by 30 using a cubic press, its thermal expansivity was investigated up to 923 K by using an X-31 ray powder diffraction technique (ambient P), and its compressibility was investigated up 32 to ~16 GPa by using a diamond-anvil cell coupled with synchrotron X-ray radiation 33 (ambient T). Its volumetric thermal expansion coefficient (α_V) and isothermal bulk modulus (K_T) were constrained as $\alpha_V = 4.24(4) \times 10^{-5}/K$ and $K_T = 103(2)$ GPa (the first 34 pressure derivative K_T' obtained as 5.4(4)), respectively. Its compressibility was further 35 36 studied with the CASTEP code using density functional theory and planewave pseudopotential technique. We obtained the K_T values as 123(3) GPa (LDA; high 37 boundary) and 92(2) GPa (GGA; low boundary), with the values of the K_T as 4.4(9) and 38 39 4.9(5), respectively. The phonon dispersions and vibrational density of states (VDoS) of 40 Lrn were simulated using density functional perturbation theory, and the VDoS was 41 combined with a quasi-harmonic approximation to compute the isobaric heat capacity (C_P) and standard vibrational entropy (S_{298}^0), yielding $C_P = 212.1(1) - 9.69(5) \times 10^2 T^{0.5}$ – 42 $4.1(3) \times 10^{6} T^{2} + 5.20(7) \times 10^{8} T^{3}$ J mol⁻¹ K⁻¹ for the T range of ~298-1000 K and $S_{298}^{0} =$ 43 129.8(13) J mol⁻¹ K⁻¹. The microscopic and macroscopic thermal Grüneisen parameters of 44 45 Lrn at 298 K were calculated to be 0.75(6) and 1.80(4), respectively. 46 47 **Keywords:** β -Ca₂SiO₄, compressibility, entropy, heat capacity, larnite, thermal

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expansivity, thermal Grüneisen parameter, thermodynamic property

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INTRODUCTION

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53 Natural larnite (Lrn; β -Ca₂SiO₄; space group $P2_1/n$ with Z = 4) was firstly 54 reported in the Larne district of Great Britain (Tilley 1929), and soon documented on the 55 island of Muck in Scotland (Tilley 1947) and the Tokatoka region in New Zealand 56 (Mason 1957). It usually locates in the contact zone between an igneous rock like dolerite 57 or andesite, and a calcitic rock such as chalk or limestone (Tilley 1929; Mason 1957). Lrn 58 commonly coexists with spurrite, melilite, merwinite and spinel, and readily transforms 59 to the olivine-structured Ca₂SiO₄ phase (the γ -Ca₂SiO₄ phase) if it is shocked or heated at 60 certain temperature (Tilley 1929). According to Bowen (1940) and Tilley (1951), Lrn 61 generally represents the high-T stage of the progressive metamorphism (or decarbonation) 62 of the siliceous limestones.

63 This conventional field occurrence of Lrn was recently supplemented with 64 another completely different appearance, tiny mineral inclusions, usually coexisting with 65 the walstromite-structured CaSiO₃ (Wal) and titanite-structured CaSi₂O₅ (Ttn), in 66 diamonds possibly originating from the lower mantle (Joswig et al. 1999; Nasdala et al. 67 2003; Brenker et al. 2005). This new appearance argues that some portions of the deep 68 interior of the Earth are significantly Ca-richer than the normal pyrolitic mantle 69 (Ringwood 1975). According to the results from some high-P experimental studies (e.g., 70 Green and Ringwood 1967; Takahashi 1986; Irifune 1994), the Ca-rich silicate in the 71 normal mantle normally changes from clinopyroxene, majoritic garnet to calcium 72 perovskite (CaPv; compositionally close to CaSiO₃) as P increases from the surface to the

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core-mantle boundary of the Earth. In order to stabilize Lrn, Wal and Ttn, significantly
higher levels of CaO should be introduced into the mantle, which has been presumably
accomplished via deep subduction of the CaO-rich continental crust material (Irifune et al.
1994; Liu et al. 2012a).

77 It is well known that Lrn is not a stable phase at ambient P (Fig. 1a). From low T 78 to high T, the composition of Ca_2SiO_4 can form many polymorphs, namely γ -Ca₂SiO₄, 79 α'_{I} -Ca₂SiO₄, α'_{H} -Ca₂SiO₄ and α -Ca₂SiO₄ (e.g., Barnes et al. 1980; Remy et al. 1995, 80 1997a, 1997b; Yamnova et al. 2011). Lrn never forms during heating, but appears as a 81 metastable phase in the stability field of the γ -Ca₂SiO₄ phase during cooling. 82 Consequently, it was usually synthesized by heating the Ca_2SiO_4 composition in the 83 stability field of the phase α'_{L} -Ca₂SiO₄, α'_{H} -Ca₂SiO₄ or α -Ca₂SiO₄, and then quickly 84 quenching into water (Remy et al. 1997a, 1997b; Fukuda et al. 1997), a process which 85 might introduce different amounts of defects in the products (e.g., twins are common in 86 Lrn; Groves 1983; Kim et al. 1992). In order to preserve the Lrn structure, some 87 "stabilizer" such as P₂O₅ (Saalfeld 1975), B₂O₃ (Remy et al. 1997a), vacancy of either Ca 88 (Rodrigues 2003) or Si (Yannaquis and Guinier 1959), or other species (Lai et al. 1992) 89 was usually introduced into the system.

According to the phase diagram of the composition Ca_2SiO_4 at high-*P* and high-*T* conditions (Fig. 1b), Lrn has a vast stability field. At the low *P* side, Lrn transforms to the γ -Ca₂SiO₄ phase at ~2 GPa and room *T* (Hanic et al. 1987; Remy et al. 1995, 1997b; Reynard et al. 1997), and the invariant point of β -Ca₂SiO₄ + γ -Ca₂SiO₄ + α'_L -Ca₂SiO₄ locates at ~0.34 GPa and 711 K (Hanic et al. 1987). At the high *P* side, the relationship

95	between Lrn and the α'_L -Ca ₂ SiO ₄ phase (Wang and W	Veidner 1994) or K ₂ NiF ₄ -structured	
96	Ca ₂ SiO ₄ phase (Liu 1978) is still poorly constrained (Fig. 1b).		
97	Lrn also participates in some other geologically	y important reactions such as	
98	3Wal = Ttn + Lrn	(1)	
99	and		
100	Ttn + Lrn = 3CaPv	(2),	
101	with the former taking place at ~ 8 GPa and the latter	r at ~12 GPa (Kanzaki et al. 1991;	
102	Wang and Weidner 1994; Akaogi et al. 2004; Sueda	a et al. 2006). To understand these	

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103 reactions, especially the one related to the CaPv, which is the dominant Ca-bearing phase 104 in the lower mantle of the Earth (Mao et al. 1977; Irifune et al. 1989) but unfortunately 105 unquenchable to room P for further characterizations such as single crystal 106 crystallographic study and calorimetric investigation (Liu and Ringwood 1975; Mao et al. 107 1977; Irifune 1994), it is necessary to investigate the thermodynamic properties of Lrn as 108 extensively and accurately as possibly. These thermodynamic properties of Lrn then can 109 be combined with the thermodynamic properties of Ttn to rigorously constrain the 110 thermodynamic properties of CaPv via an accurate experimental determination of the P-T111 conditions of equation (2). The thermal expansivity of Lrn has been investigated by 112 different techniques, but the volumetric thermal expansion coefficient at 298 K (α_{298}) varies from $\sim 1.80 \times 10^{-5}$ to 4.25×10^{-5} K⁻¹ (e.g., Remy et al. 1997b; Swamy and 113 114 Dubrovinsky 1997). The isothermal bulk modulus (K_T) of Lrn has also been investigated (Remy et al. 1997b; Swamy and Dubrovinsky 1997), but its value at 298 K varies from 115 ~119 to 166(15) GPa (the first pressure derivative of K_T , K_T , set as 4). Heat capacity 116 measurements have been performed from 52.66 to 296.48 K (Todd 1951) and from 406.0

118to 964.6 K (Coughlin and O'Brien 1957). With an empirical sum equation of Debye and119Einstein functions to extrapolate the heat capacity data down to zero K, an S_{298}^0 value as120127.6(8) J mol⁻¹ K⁻¹ has been derived, which has never been verified by any means so far.121It thus appears that a substantial amount of work has to be conducted in order to fully122understand the thermodynamic features of Lrn.123In this study, we synthesized pure Lrn in its *P-T* stability field by using a massive124cubic press, and determined its thermal expansivity by using an in situ high-*T* powder X-

126 radiation combined with a diamond-anvil cell (DAC; ambient *T*). Its compressibility was

ray diffraction method (ambient P), and its compressibility by using a synchrotron X-ray

127 further studied with the CASTEP code using density functional theory and planewave

128 pseudopotential technique. Using density functional perturbation theory, its phonon

129 dispersions and vibrational density of states (VDoS) were simulated, and the VDoS was

130 combined with a quasi-harmonic approximation to compute the isobaric heat capacity (C_P)

131 and vibrational entropy (S). Finally, the microscopic and macroscopic thermal Grüneisen

132 parameters of Lrn at the ambient *P*-*T* condition were calculated.

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EXPERIMENTAL AND SIMULATING METHODS

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137 The polycrystalline Lrn sample used in our thermal expansion and compression 138 experiments was synthesized with the CS-IV 6×14 MN cubic press at the High-Pressure 139 Laboratory of Peking University (Liu et al. 2012b). In the high-*P* synthesizing 140 experiments, we used the assembly BJC-1 (Liu et al. 2012c), with its pressure calibrated

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141 at ambient T by using the high-P phase transitions of Bi (I-II transition at 2.55 GPa, and 142 II-III transition at 2.69 GPa) and Ba (I-II transition at 5.5 GPa). The temperature in the 143 high-P synthesizing experiments was measured and controlled by employing a Pt₉₄Rh₆-144 Pt₇₀Rh₃₀ thermocouple (type B), with any potential pressure effect on the e.m.f. ignored. 145 We prepared the starting material for the high-P synthesizing experiments with the 146 following steps: we firstly mixed under acetone the powders of SiO₂ and CaCO₃ in a 147 molar ratio of 1:2, which were pretreated at 1 atm and 573 K for 72 hours; we secondly 148 pressed this mixture into a pellet and degassed it at 1 atm and 1273 K for about 48 hours; 149 we thirdly crushed finely the pellet under acetone into a powder, which was later stored at 150 383 K in a drying oven. The starting material for the high-P synthesizing experiments 151 was sealed into an hBN capsule. The synthesizing conditions were 6 GPa and 1473 K 152 with a heating time of 6 hours. In total we conducted five high-P synthesizing 153 experiments under nominally identical physical conditions, and obtained four coherent 154 samples and one dust-like sample. It appeared that some unknown factor played a role in 155 the high-*P* synthesizing process.

Some portions of the coherent samples from the high-*P* synthesizing experiments were processed and examined by a scanning electron microscope (Quanta 650 FEG), a confocal micro-Raman system (Renishaw system RM-1000) and an electron microprobe (JEOL JXA-8100), and confirmed to be pure Lrn. Other portions of these samples were slowly ground down to a fine powder, which was used in our thermal expansion and compression experiments.

162 The thermal expansion experiments at ambient *P* were carried out with an X'Pert 163 Pro MPD diffractometer system equipped with an Anton Paar HTK-1200N oven. A

164 Eurotherm temperature controller (Eurotherm 2604) collected with a type S thermocouple, 165 which was checked against the melting point of NaCl, was used to control the 166 temperature. Other details of the X'Pert Pro MPD diffractometer system included a Cu 167 target, operation voltage of 40 kV and current of 40 mA. The whole experimental system 168 was carefully checked up to 1273 K by using quartz as an internal standard (Hu et al. 169 2011). Following the heating and data-collection procedures reported in Liu et al. (2010), 170 we carried out the high-T experiments up to 1023 K. At every studied T, X-ray data were 171 collected from 10 to $70^{\circ}2\theta$, with a scanning step length of $0.017^{\circ}2\theta$ and a scanning time 172 of 10 seconds for each scanning step. As demonstrated by He et al. (2011) and Wang et al. 173 (2012), unit-cell parameters with high accuracy were readily extracted by a full powder 174 X-ray pattern refinement with the MDI's program Jade 6.0 (Material Data, Inc.). 175 The high-P angle dispersive X-ray diffraction experiments (ambient T) were

176 conducted with a symmetrical DAC at the beamline X17C, National Synchrotron Light 177 Source, Brookhaven National Laboratory. The experimental techniques were generally 178 identical to those used in our previous studies (e.g., Liu et al. 2011; He et al. 2012; Xiong 179 et al. 2015): a T301 stainless steel plate was used as the gasket, a 4:1 methanol-ethanol 180 mixture as the pressure medium, and a couple of ruby chips used as the pressure marker 181 (the ruby fluorescence method; Mao et al. 1978). The incident synchrotron radiation beam had a wavelength of 0.409929 Å, and its size was $\sim 25 \times 20 \ \mu m^2$. Each X-ray 182 183 diffraction pattern at certain P was collected for 60 minutes using an online CCD detector, 184 and later processed to generate the conventional one-dimension profile using the Fit2D 185 program (Hammersley 1996). Subsequently, the positions of the diffraction peaks were

determined with the Jandel Scientific PeakFit V4.11 program, and the unit-cellparameters were obtained by using the UnitcellWin program.

188 The first-principles simulations carried out to investigate the compression 189 behavior of Lrn were completed with the CASTEP code using density functional theory 190 (Hohenberg and Kohn 1964; Kohn and Sham 1965) and planewave pseudopotential 191 technique (Payne et al. 1992). We treated the exchange-correlation interaction by both the 192 local density approximation (LDA) (Ceperley and Alder 1980; Perdew and Zunger 1981) 193 and generalized gradient approximation (GGA) with the Perdew-Burker-Ernzerhof functional (Perdew et al. 1996), and used a convergence criterion of 10^{-6} eV/atom on the 194 195 total energy in the self-consistent field calculations. We employed a planewave basis set 196 with a cutoff of 900 eV to expand the electronic wave functions, and a norm-conserving 197 pseudopotential to model the ion-electron interaction (Lin et al. 1993; Lee 1995). We 198 sampled the irreducible Brillouin zone with a $4 \times 3 \times 2$ Monkhorst-Pack grid (Monkhorst 199 and Pack 1976). The effects of using larger cutoff and k point mesh on the calculated 200 properties were found to be insignificant. The computation cell contained four Ca₂SiO₄ 201 molecules (28 atoms), with the initial structure model from Tsurumi et al. (1994). The 202 equilibrium lattice parameters and internal coordinates at different pressures were 203 optimized by minimizing the Hellmann-Feynman force on the atoms and simultaneously 204 matching the stress on the unit cell to the target stress. These theoretical techniques were 205 used in our previous studies targeting the structures and thermodynamics of some silicate 206 minerals (e.g., Deng et al. 2010, 2011; Chang et al. 2013).

Based on the optimized structure with the LDA method, the phonon dispersions
and VDoS of Lrn were calculated by diagonalizing the dynamical matrix whose elements

were obtained using density functional perturbation theory (Baroni et al. 2001; Refson et
al. 2006). The q-vector grid spacing for interpolation was 0.05 $Å^{-1}$, which represented the
average distance between the Monkhorst-Pack q-points used in the dynamical matrix
calculations. The phonon dispersions were obtained at the high symmetry points (Z, G, Y,
A, B, D, E, C). The coordinates of these points on the surface of the Brillouin zone were
$Z = (0 \ 0 \ \frac{1}{2}), G = (0 \ 0 \ 0), Y = (0 \ \frac{1}{2} \ 0), A = (-\frac{1}{2} \ \frac{1}{2} \ 0), B = (-\frac{1}{2} \ 0 \ 0), D = (-\frac{1}{2} \ 0 \ \frac{1}{2}), E = (-\frac$
$\frac{1}{2}$ $\frac{1}{2}$ and C = (0 $\frac{1}{2}$ $\frac{1}{2}$).
RESULT AND DISCUSSION
High-P synthesizing of pure Lrn

221 The BSE images taken from the polished surfaces of the coherent samples suggest 222 two coexisting phases: the predominant phase has a volume percentage higher than $\sim 98\%$, 223 appears relatively dark, and attains a grain size of $\sim 10-50 \,\mu\text{m}$; the trace phase appears 224 relatively bright and has a grain size of $\sim 1-2 \,\mu m$. 10 electron microprobe analyses were 225 conducted on 10 arbitrarily selected grains of the predominant phase, and their average is 226 $Ca_{2,02(2)}Si_{0,99(1)}O_4$, suggesting a homogeneous composition closely matching the 227 stoichiometry of Lrn. 10 Raman spectra were collected on 10 arbitrarily selected grains of 228 the predominant phase, and they are similar, and in excellent agreement with the Raman 229 data of Lrn reported by Piriou and McMillan (1983) and Remy et al. (1997a). It follows 230 that the predominant phase in the coherent samples is Lrn indeed. On the other hand, the 231 trace phase is mostly likely a reaction residue (CaO) left from the formation of Lrn.

232 As suggested by our preliminary Raman spectroscopic data, the dust-like sample 233 is distinctly dominated by the γ -Ca₂SiO₄ phase, which could be produced during 234 temperature quenching or pressure release. The Ca_2SiO_4 system is infamous in its back-235 transformation (either partial or complete) of its high-T polymorphs at room P (Fig. 1a), 236 and the operating factors include the maximum synthesizing T, the cooling kinetics, the 237 grain size, and the release of strains and charge repulsions in the β -Ca₂SiO₄ phase (Remy 238 et al. 1997a). Our observation made in this study similarly suggests that the back-239 transformation process is not always escapable for the samples synthesized at high P, in 240 agreement with Remy et al. (1995) and Reynard et al. (1997). Anyhow, the first attempt 241 in synthesizing a large amount of pure Lrn in its P-T stability field (Fig. 1b) has been at 242 least partially successful, in spite of the complication caused by the back-transformation

243 process.

It is worthy to point out that the univariant curve of $\beta = \alpha'_L$ determined by 244 245 Klement and Cohen (1974) and Hanic et al. (1987) may be reasonably extrapolated to ~ 6 246 GPa (Fig. 1b). Assuming there was no any back-transformation from the α'_L phase to the 247 Lrn phase in our synthetic samples, the P-T condition (6 GPa and 1473 K) of our 248 synthesizing experiments does support this extrapolation. In contrast, the P-T condition (9) 249 GPa and 1350 K) for the phase transition from the Lrn phase to the α'_L phase constrained 250 by the in situ synchrotron X-ray experiments of Wang and Weidner (1994) does not. 251 More high-*P* experimental investigation on this issue seems necessary.

254 High-T X-ray diffraction experiments were conducted up to 1023 K at ambient 255 pressure (Fig. 2). We found that all major X-ray diffraction peaks observed at 256 temperatures below 973 K could be attributed to the Lrn phase (Jost et al. 1977; 257 Yamnova et al. 2011), and the new peaks appearing from 973 K on could be attributed to 258 the α'_L phase (ideally Ca₂SiO₄; Saalfeld 1975). As shown in Fig. 1(a), the stable form of 259 Ca_2SiO_4 at relatively low T is the y phase; its formation from the metastable Lrn phase 260 was not observed in our experiments though. This phase transition is thus generally 261 sluggish, in good agreement with the observation made by Remy et al. (1997a). The 262 mechanism in this phase transition has been regarded as complicated and not fully 263 understood, with the affecting factors presumably including the synthesizing condition, 264 impurity and grain size of the used Lrn material, and the details of the heating process 265 (Remy et al. 1997a; references therein).

266 The phase transition T from the metastable Lrn to the α'_{L} phase on heating was 267 located between 923 and 973 K by our experiments (Fig. 2). In contrast, it was 268 constrained between 960(5) to 988(5) K by in situ high-T Raman spectroscopy (Remy et 269 al. 1997a) and between 984(10) to 1005(10) K by in situ high-T powder X-ray diffraction 270 (Remy et al. 1997b). Taking into account the differences in the used materials and 271 heating processes in these studies, the agreement is good.

272 The variation of the unit-cell parameters of the Lrn phase with T as determined in this study is shown in Fig. 3, and it is almost linear for the investigated T interval. This 273 274 result is in general agreement with Fukuda et al. (1997), but contradicts Remy et al. 275 (1997b): the data from Remy et al. (1997b) clearly demonstrated non-linear correlations 276 between the *a*-axis and *T* (Fig. 3a), between the *b*-axis and *T* (Fig. 3b), between the *c*-axis

277 and T (Fig. 3c), and between the β and T (Fig. 3e). As T increases from ~300 to 923 K, 278 our data suggested that the *a*-axis increases by 0.61(1)%, the *b*-axis by 0.73(1)%, the *c*-279 axis by 1.12(1)%, and the volume by 2.62(1)% (Table 1); the ratio of the relative thermal 280 expansions along the a-, b- and c-axes is 1:1.10:1.84, suggesting that the c-axis is much 281 more expandable than the other two axes. In contrast, the data from Fukuda et al. (1997) 282 suggested that the *a*-axis increases by 0.61%, the *b*-axis by 0.65%, the *c*-axis by 1.02%, 283 and the volume by 2.45% for a similar T interval (from 298 to 918 K). It follows that the 284 volumetric thermal expansion constrained by Fukuda et al. (1997) should be generally 285 smaller than that defined by this study.

286 If the phase transition T from the metastable Lrn to the α'_L phase on heating is assumed to be ~973 K, the extrapolated unit-cell parameters of the Lrn phase at this T are 287 then a = 5.5462 Å, b = 6.8058 Å, c = 9.4156 Å, V = 354.87 Å³, and $\beta = 93.345^{\circ}$. 288 Compared to the unit-cell parameters of the α'_L phase at 973 K (Table 1), which can be 289 translated as a = 5.532 Å, b = 6.864 Å, c = 9.443 Å, V = 358.53 Å³, and $\beta = 90^{\circ}$ (Saalfeld 290 291 1975), the relative changes in the unit cells accompanying this phase transition are 292 suddenly a reduction in the *a*-axis (~0.26%), some expansions of the *b*-axis (~0.86%), *c*-293 axis (~0.29%) and volume (~1.03%), and a huge jump of the β angle. It concludes that 294 this phase transition is indeed of first-order character, in agreement with Barnes et al. 295 (1980), Remy et al. (1997a, 1997b), and Liu et al. (2002).

In order to derive the volumetric thermal expansion coefficient, the following equations have been fitted with our T-V data

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$$V_T = V_0 \exp[\int_0^T \alpha_T dT]$$
(3)

299 and

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$$\alpha_T = a_0 + a_1 T + a_2 T^{-2}$$
(4),

where V_T , V_0 and a_T are the high-*T* volume, room-*T* volume and volumetric thermal expansion coefficient at temperature *T*, respectively. a_0 , a_1 and a_2 are the constants obtained in fitting the experimental *T*-*V* data. Due to the generally linear relationship observed between the volume and *T* (Fig. 3), a_1 and a_2 should be close to 0, so that their values have been fixed as zero in our data-fitting process. The obtained parameter is then $a_0 = 4.24(1) \times 10^{-5} \text{ K}^{-1}$.

Similarly, replacing the volume data in the equations (3) and (4) with the axial dimensions, the axial thermal expansion coefficients can be obtained. Our experimental data yield $a_0 = 1.00(1) \times 10^{-5} \text{ K}^{-1}$ for the *a*-axis, $a_0 = 1.21(1) \times 10^{-5} \text{ K}^{-1}$ for the *b*-axis and a_0 $= 1.77(2) \times 10^{-5} \text{ K}^{-1}$ for the *c*-axis, which are in excellent agreement with those from Remy et al. (1997b; $1.00(10) \times 10^{-5} \text{ K}^{-1}$, $1.25(10) \times 10^{-5} \text{ K}^{-1}$ and $1.79(15) \times 10^{-5} \text{ K}^{-1}$, respectively).

312 The volumetric thermal expansion coefficient of the Lrn phase determined in this study is compared to the results from other studies in Table 2. With the only exception of 313 314 Fukuda et al. (1997), all other experimental studies have obtained remarkably similar values: for example, they vary slightly from 3.97×10^{-5} K⁻¹ to $4.25(14) \times 10^{-5}$ K⁻¹ when T is 315 ~298 K. The reason why Fukuda et al. (1997) does not agree to other experimental 316 317 studies is presently unknown. On the other hand, the values obtained with a quasi-318 harmonic lattice dynamic calculation done by Swamy and Dubrovinsky (1997) and with a 319 thermodynamic optimization method done by Holland and Powell (1998) are 320 significantly smaller. The value constrained with the thermodynamic optimization 321 method done by Swamy and Dubrovinsky (1997), however, is in extraordinarily good 322 agreement with the values from most experimental studies.

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324 Isothermal bulk modulus

High-*P* compression experiments at ambient *T* were conducted up to ~16 GPa for the Lrn. No phase transition was observed, which is in good agreement with Reynard et al. (1997) and Remy et al. (1997b). The X-ray diffraction patterns obtained at *P* higher than ~10 GPa did not show any apparent peak-broadening, which was merely a fortuitous situation; according to Klotz et al. (2009), our pressure medium could only maintain a hydrostatic pressure environment up to about 10 GPa.

The unit-cell parameters of the Lrn at different pressures are summarized in Table 332 3 and shown in Fig. 5. All unit-cell parameters vary non-linearly with *P* for the 333 investigated *P* range. As *P* increases from 1 atm to 16.25 GPa, the *a*-axis decreases from 334 5.510(1) to 5.297(4) Å (by 3.87(1)%), the *b*-axis decreases from 6.753(2) to 6.529(1) Å 335 (by 3.32(1)%), the *c*-axis decreases from 9.316(1) to 8.931(2) Å (by 4.14(1)%), the 336 volume decreases from 345.6(1) to 307.7(4) Å (by 10.9(2)%), and the β -angle increases 337 from 94.48(1) to 94.82(2) Å (by 0.36(1)%).

In order to determine the isothermal bulk modulus, the third-order Birch-Murnaghan equation of state (BM-EoS; Birch 1947) has been fitted with the P-V data of the Lrn by a least-squares method:

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$$P = 3K_T f_E (1 + 2f_E)^{\frac{5}{2}} [1 + \frac{3}{2}(K_T - 4)f_E]$$
(5)

where *P* is the pressure, K_T the isothermal bulk modulus, K_T the first pressure derivative of K_T , and f_E the Eulerian definition of finite strain, which is $[(V_0/V)^{2/3}-1]/2$, respectively. In the Eulerian definition of finite strain, V_0 is the volume at zero pressure whereas *V* is

345	the volume at high pressure. When K_T is set as 4, the isothermal bulk modulus (K_T) of
346	Lrn is 110.8(8) GPa and the zero-pressure volume is 345.5(1) Å ³ . If K_T' is not fixed, the
347	results of our best data-fitting are $K_T = 103(2)$ GPa, $K'_T = 5.4(4)$ and $V_0 = 345.6(1)$ Å ³ .
348	The quality of the derived third-order BM-EoS for the Lrn can be evaluated by
349	using the f_E -F plot (Fig. 6); F is defined as $F \equiv P/[3f_E(1+2f_E)^{5/2}]$. Using F, the third-order

350 BM-EoS can be rewritten as:

351
$$F = K_T + 3/2K_T(K_T - 4)f_E$$
(6)

so that the slope of the line defined by the experimental data should be equal to $3/2K_T(K_T'-4)$, and the intercept value is the isothermal bulk modulus. Accordingly, a slope of zero means $K_T'=4$, a negative slope $K_T'<4$, and a positive slope $K_T'>4$. Fig. 3 clearly suggests that the K_T' of the Lrn is larger than 4, in agreement with our *P-V* datafitting results detailed in the previous paragraph.

A linearized third-order Birch-Murnaghan equation of state (Angel 2000) was used to obtain the parameters of the equations of state for the crystallographic axes, yielding: $a_0 = 5.510(1)$ Å, $K_{T-a} = 99(5)$ GPa and $K'_{T-a} = 5.8(10)$ for the *a*-axis, $b_0 =$ 6.753(3) Å, $K_{T-b} = 137(7)$ GPa and $K'_{T-b} = 3.0(1)$ for the *b*-axis, and $c_0 = 9.316(1)$ Å, K_{T-c} = 89(2) GPa and $K'_{T-c} = 5.9(4)$ for the *c*-axis. The quality of the derived third-order Birch-Murnaghan equation of state for the axes of the Lrn has been evaluated by using the *f_E-F* plot as well (Fig. 6).

The geometry parameters of the Lrn phase at zero pressure obtained from the first-principles simulation are summarized in Table 4. As usual, the unit-cell parameters a, b and c calculated by the GGA method are larger (0.63 to 1.16%) whereas those

367 calculated by the LDA method are smaller (-1.75 to -2.16%), compared to the 368 experimental determinations by Tsurumi et al. (1994). Additionally, the unit-cell volume 369 calculated by the GGA method is larger by 2.48% whereas that calculated by the LDA 370 method is smaller by -5.89%. Summarily, both methods reproduce well the unit-cell 371 parameters of the Lrn phase.

372 The unit-cell parameters of the Lrn phase at high pressures predicted by our first-373 principles simulation are shown along with the experimentally-obtained data in Fig. 5. 374 Compared to the experimentally-determined curves, the curves obtained with the GGA 375 method have slightly steeper slopes whereas those obtained with the LDA method have 376 more gentle slopes, which is normal. In general, the GGA method should yield a lower 377 K_T and the LDA method should yield a higher K_T , with the experimentally-constrained K_T 378 falling in between (Deng et al. 2011). The K_T values obtained by fitting equation (5) with the *P*-*V* data essentially meet this expectation: $K_T = 92(2)$ GPa ($K_T = 4.9(5)$, $V_0 = 355.0(1)$ 379 Å³; GGA) and $K_T = 123(3)$ GPa ($K_T = 4.4(9)$, $V_0 = 325.94(9)$ Å³; LDA). When K_T is set 380 as 4, $K_T = 96(1)$ GPa ($V_0 = 354.8(1)$ Å³) and $K_T = 124(1)$ GPa ($V_0 = 325.9(6)$ Å³) are 381 382 derived from the data predicted with the GGA method and LDA method, respectively. It 383 concludes that our experimentally-constrained K_T value should be a good approximation 384 to the real K_T value of the Lrn phase.

The K_T values of the Lrn phase from different studies are compared in Table 5, with the most of them ranging from ~96 to 124 GPa (with the assumption $K_T' = 4$). Clearly, the result from Remy et al. (1997b) is the one at odd, which can be adequately explained by their used pressure medium of silicon oil. According to Klotz et al. (2009), silicon oil can maintain a hydrostatic pressure environment at P < 3 GPa only, suggesting the high-*P* data in Remy et al. (1997b; 4.6, 10 and 14.7 GPa) were severely affected by
deviatoric stress.

392

393 Heat capacity and entropy

394 We calculated the phonon dispersions and VDoS of the Lrn phase by the first-395 principles method using density functional perturbation theory; the optimized crystal 396 structure with the LDA method was used in the calculation since the geometries of the Ca 397 polyhedra and Si tetrahedra have been slightly better reproduced with the LDA method 398 (Table 4). The dynamical matrices were computed at 22 wave (q) vectors in the Brillouin 399 zone of the primitive cell, and interpolated to obtain the bulk phonon dispersions. Fig. 7 400 shows the predicted dispersion curves along several symmetry directions and the VDoS. 401 The results of the phonon spectra were used to compute the internal energy (E) and 402 isochoric heat capacity (C_V) as functions of temperature. The temperature dependence of 403 the E was obtained by the following equation,

404
$$E(T) = E_{tot} + E_{ZP} + \int \frac{h\omega}{\exp(\frac{h\omega}{kT}) - 1} F(\omega) d\omega$$
(7),

with E_{tot} representing the total electronic energy at 0 K, E_{ZP} the zero point vibrational energy, *h* the Planck's constant, *k* the Boltzmann's constant, and $F(\omega)$ the vibrational density of states. We evaluated the E_{ZP} in equation (7) using the following equation,

408
$$E_{ZP} = \frac{1}{2} \int F(\omega) h \omega d\omega$$
 (8),

409 Furthermore, we approximated the lattice contribution to the C_V with the following 410 equation:

411
$$C_{V}(T) = k \int \frac{\left(\frac{h\omega}{kT}\right)^{2} \exp\left(\frac{h\omega}{kT}\right)}{\left[\left(\frac{h\omega}{kT}\right)^{2} - 1\right]^{2}} F(\omega) d\omega$$
(9).

412 The C_V result calculated in this way is shown in Table 6.

413 We calculated the isobaric heat capacity (C_P) by adding an anharmonic effect to 414 the C_V obtained from the above calculations, using the following equation:

415
$$C_P = C_V + \alpha_T^2 K_T V_T T$$
 (10),

416 where α_T , K_T , and V_T were the thermal expansivity, isothermal bulk modulus, and volume 417 at 1 atm and *T* K, respectively. For the Lrn phase, $\alpha_T = 4.24(4) \times 10^{-5}$ K⁻¹ and $K_T = 103(2)$ 418 GPa ($K_T' = 5.4(4)$). The temperature dependence of K_T has not been experimentally 419 determined, thus assumed as zero in our calculation. V_T at *T* K was calculated with the 420 following equation:

421
$$V_T = V_{298} \exp(\int_{298}^T \alpha_T dT)$$
 (11),

where $V_{298} = 51.88 \text{ cm}^3/\text{mol}$ (Remy et al. 1997b). The calculated C_P values are listed in Table 6, along with the experimental results from Todd (1951) and from Coughlin and O'Brien (1957). Errors of our C_P values were propagated from the uncertainty in the thermal expansion coefficients and isothermal bulk modulus, and they decrease from $\sim 6\%$ at ~ 10 K to $\sim 0.06\%$ at ~ 202 K, and then increase to $\sim 0.21\%$ at 1000 K. The calculated C_P data were then expressed using the following polynomial of temperature,

428
$$C_P = \mathbf{k}_0 + \mathbf{k}_1 T^{0.5} + \mathbf{k}_2 T^2 + \mathbf{k}_3 T^3 + \mathbf{k}_4 T + \mathbf{k}_5 T^2 + \mathbf{k}_6 T^3$$
(12),

429 where C_P was in J mol⁻¹ K⁻¹ and *T* in K. In the fitting procedure, the data were divided 430 into three different *T* intervals, 10-50, 50-293, and 293-1000 K. The C_P equation 431 coefficients for the three *T* intervals are summarized in Table 7.

434 964.6 K). For the T range of 52.66-296.48 K, our C_P data show excellent agreement with

435 those from Todd (1951), with the relative difference quickly decreasing from $\sim 13\%$ at

~51 K to ~1% at ~303 K (Table 6). For the T range of 406.0-964.6 K, our C_P data are also 436

437 in good agreement with those from Coughlin and O'Brien (1957), with the relative

438 difference varying from $\sim 2.0\%$ (859 K) to $\sim 1.5\%$ (404 K).

439 The obtained C_P values have been applied to the calculation of the vibrational 440 entropy at T K using the following equation

441
$$S_T^0 = \int_0^T \frac{C_P}{T} dT$$
 (13).

The vibrational entropy at 298 K (S_{298}^0) calculated from the C_P values in the T range of 0 442 to 298 K obtained in this study is 129.8(13) J mol⁻¹ K⁻¹, in good agreement with the only 443 444 available value derived from the existing low-T heat capacity measurements, 127.6(8) J mol⁻¹ K⁻¹ (Todd 1951). 445

446

432

447 Grüneisen parameter

448 The Grüneisen parameter is often used to describe the pressure dependences of 449 some thermodynamic, elastic and transport properties which are crucial in the 450 investigation of the deep Earth. It is dimensionless, and varies slowly with P and T, and 451 therefore appears very attractive to geophysicists.

452 The mode Grüneisen parameters (y_i) of the Lrn phase (298 K) have been 453 calculated according to the following equation

454
$$\gamma_i = -\frac{\partial \ln \nu_i}{\partial \ln V} = \frac{K_T}{\nu_{0i}} \left(\frac{\partial \nu_i}{\partial P}\right)_T$$
(14),

455 where v_{0i} and v_i are the frequencies of the vibration mode *i* of the Lrn phase at room *P* 456 and high P, respectively (at ambient T). With the high-P Raman data for the Lrn phase 457 from Reynard et al. (1997) and our bulk modulus determined in this study, we calculate 458 the value of the γ_i of the 20 Raman vibration modes, and find that it varies from ~0 (peaks 152 and 520 cm⁻¹) to \sim 1.90 (peak 255 cm⁻¹), with an arithmetic average of \sim 0.69(52). The 459 460 details are summarized in Fig. 9. As expected (Gillet et al. 1991; Fujimori et al. 2002), 461 the γ_i values for the lattice modes are generally larger than those for the internal modes of 462 the SiO₄ tetrahedra.

463 The mode Grüneisen parameters can be cast into a microscopic thermal Grüneisen 464 parameter ($\gamma_{th,1}$) with the following average scheme,

465
$$\gamma_{th,1} = \frac{\sum_{i} C_{V,i} \cdot \gamma_i}{\sum_{i} C_{V,i}}$$
(15),

466 where $C_{V,i}$ is the contribution of the vibration mode *i* to the isochoric heat capacity. $C_{V,i}$ 467 is estimated on the basis of a harmonic Einstein oscillator,

468
$$C_{V,i} = k \left(\frac{hv_i}{kT}\right)^2 \exp\left(\frac{hv_i}{kT}\right) / \left[\exp\left(\frac{hv_i}{kT}\right) - 1\right]^2$$
(16)

469 where *h* is the Planck's constant, and *k* the Boltzmann's constant. The $\gamma_{th,1}$ of the Lrn 470 phase attains 0.77(8), about 12% larger than the arithmetic average of the mode 471 Grüneisen parameters.

472 Alternatively, a macroscopic thermal Grüneisen parameter ($\gamma_{th,2}$) can be calculated 473 from the following equation

474
$$\gamma_{th,2} = \frac{\alpha K_T V_T}{C_V}$$
(17)

With all the values of the variables on the right-hand side of equation (17) constrained in this study, we obtained $\gamma_{th,2} = 1.80(4)$ for the Lrn phase at the *P*-*T* condition of 298 K and 1 atm.

478 Clearly there is a sharp difference (relatively $\sim 133\%$) between the microscopic 479 isothermal Grüneisen parameter $y_{th,1}$ and the macroscopic isothermal Grüneisen parameter $\gamma_{th,2}$ for the Lrn phase, which should by no means exist under ideal 480 481 circumstance. In the cases of other minerals, the observed difference usually appears up 482 to ~ 25% only (Hofmeister and Mao 2002; Tang et al. 2014). Several factors may account 483 for this prominent difference. Firstly, the pressure medium (a 16:4:1 methanol-ethanol-484 water mixture) used by Reynard et al. (1997) was not able to maintain a fully hydrostatic pressure environment at P > -10 GPa (Klotz et al. 2009), which should have led to some 485 486 overestimate in the P measurements (He et al. 2004). Indeed, the v_i -P curves of the Lrn 487 phase shown by Reynard et al. (1997; their Fig. 7) bended towards the P-axis. This 488 situation was exacerbated by the fact that they only used the data in the P range of 6-16 489 GPa to derive the P dependence of the Raman frequency, which was thus possibly 490 substantially underestimated. Consequently, all the mode Grüneisen parameters should 491 have been strongly underestimated according to equation (14). Secondly, the factor-group 492 analysis (Piriou and McMillan 1983; Remy et al. 1997a) predicted 84 normal modes for 493 the Lrn phase, with 42 Raman-active modes, 39 infrared-active modes and 3 acoustic 494 modes. Since the number of the Raman peaks observed in Reynard et al. (1997) was 495 limited to 20 only, it is thus possible that the microscopic isothermal Grüneisen parameter 496 $y_{th,1}$ obtained here might have been strongly biased. Thirdly, the polyhedral bulk moduli

497 of the Ca polyhedra and the Si tetrahedron in the Lrn phase should be significantly 498 different (Hazen and Finger 1979), a different conversion and average scheme employing 499 the polyhedral bulk moduli, as the one proposed by Hofmeister and Mao (2002), might be 500 more appropriate than the equations (14) and (15). Fourthly, the assumption of a 501 harmonic Einstein oscillator (equation (16)) might be a poor approximation to the actual 502 chemical bonds in the Lrn phase, and the contribution of an anharmonic component may 503 be significant (Fujimori et al. 2002). Whatever the actual reasons are, more spectroscopic 504 investigation on the Lrn phase is deemed necessary in order to solve the discrepancy 505 between the microscopic isothermal Grüneisen parameter $\gamma_{th,1}$ and the macroscopic 506 isothermal Grüneisen parameter $\gamma_{th,2}$. 507 508 509 **IMPLICATIONS** 510 511 Pure Lrn of large quantity has been successfully synthesized for the first time in 512 513 514

Pure Lrn of large quantity has been successfully synthesized for the first time in its *P*-*T* stability field by using high-*P* experimental technique. Its thermal expansivity has been accurately constrained by some high-*T* X-ray powder diffraction experiments (ambient *P*), and the obtained volumetric thermal expansion coefficient is not very different from that determined with the Lrn metastably formed at ambient *P* (Table 2). In comparison, its experimentally-constrained bulk modulus, verified by supplementary first-principles simulation, is much lower than that obtained from the metastably formed material (Remy et al. 1997b), by ~38% (Table 5). However, the DAC experiments in Remy et al. (1997b) used silicon oil as the pressure medium, which should have resulted

520 in presumably significant overestimation in the P values (Klotz et al. 2009) and 521 consequently a nominally larger bulk modulus (He et al. 2004). The discrepancy in the 522 bulk moduli is thus highly possibly irrelevant to any possible difference in the used 523 materials. It follows that the difference in the two materials synthesized in the stability 524 field of Lrn or metastably generated at ambient P is too small to have any significant 525 impact on the volumetric thermal expansion coefficients (presumably, the bulk moduli as 526 well). Indeed, both materials were shown as heavily twined (Groves 1983; Kim et al. 527 1992; Wang and Weidner 1994; Remy et al. 1995).

528 Furthermore, the defects (such as twins) and the small amounts of impurities in 529 the Lrn phase (such as 0.32% iron and aluminium oxides and 0.14% magnesia: Todd 530 1951) have insignificant influence on the heat capacity and vibrational entropy. Our first-531 principles simulation, combined with the newly experimentally-determined volumetric 532 thermal expansion coefficient and bulk modulus, provides an independent examination on 533 the results from the heat capacity measurements carried out with the metastably formed 534 material (Todd 1951; Coughlin and O'Brien 1957). The agreement between the results 535 from these two completely different lines is good (Fig. 8); for example, the relative 536 difference between the values of the standard entropy is $\sim 1.6\%$ only.

Lrn has been discovered as tiny inclusions in diamonds originating from the deep interior of the Earth (Joswig et al. 1999; Nasdala et al. 2003; Brenker et al. 2005), which provides a means to explore the *P-T* conditions of the diamond formation (or the *P-T* condition for trapping the Lrn). Accurate estimates of the volumetric thermal expansion coefficient, bulk modulus and Grüneisen parameter are critical in constraining the remnant pressures of the Lrn inclusions (Ye et al. 2001; Gillet et al. 2002; Joswig et al. 543 2003; Nasdala et al. 2003). With the formation T of the diamond host independently 544 estimated from some geothermometer or assumed from some typical geotherm, the 545 formation P of the diamond host can be accurately constrained (Barron 2005). According 546 to Kagi et al. (2009), the coexisting of Lrn with other calcium-rich silicates such as Wal 547 and Ttn can be extremely useful, and capable to constrain the formation P without any 548 extra estimate of the formation T, provided the volumetric thermal expansion coefficient, 549 bulk modulus and Grüneisen parameter of the coexisting Wal or Ttn known well. 550 One exercise has been carried out to constrain the *P*-*T* condition for trapping the

Lrn inclusions by the diamonds from the Kankan district of Guinea (Fig. 10). Joswig et al. (1999) reported the unit-cell parameters of five Lrn inclusions coexisting with the Ttn phase in four diamonds, which, combined with our third order BM-EoS for Lrn, lead to a remnant pressure of 1.3(3) GPa. Using equation (1) in Kagi et al. (2009), and with the thermal expansion and compressibility data for diamond used by them and those for Lrn from this study, the *P-T* trajectory for the diamonds to trap the Lrn inclusions has been calculated as

558
$$P = 0.0057T + 0.17$$
 (18),

where *P* is in GPa, and *T* in K (the thick curve in Fig. 10). Clearly, this *P*-*T* trajectory locates in the stability field of diamond. Its intersection with the *P*-*T* curve for equation (1) indicates a minimum *P* value of ~9.9 GPa as the inclusion-trapping pressure. Furthermore, the *T* that the Lrn and Ttn inclusions and the diamond host once experienced is of typical mantle character, suggesting that some portions of the deep mantle have distinct CaO-rich composition but similar *T*, compared to the normal mantle.

565	The discovery of Lrn and other calcium-rich silicates in some diamonds argues
566	that some portions of the deep interior of the Earth have compositions substantially
567	different to the normal pyrolitic mantle (Ringwood 1975; Joswig et al. 1999; Brenker et
568	al. 2005; Zedgenizov et al. 2014). The enriching process of CaO to concentration levels
569	high enough to stabilize these calcium-rich silicates might have been achieved via deep
570	subduction of the CaO-rich continental crust material into the mantle (Irifune et al. 1994;
571	Liu et al. 2012a). How this material interacts with the regular mantle, how the calcium-
572	rich silicates such as Lrn survive the interaction, what new phases form from the
573	reactions at the P-T conditions of the mantle are largely unknown (Joachim et al. 2011,
574	2012). It is believed that the thermodynamic properties of the Lrn phase constrained in
575	this study will shed lights on understanding these possible reactions though.
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577	
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579	
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- 829

830 Figure Captions

831

832 **FIGURE 1.** Phase diagram of composition Ca_2SiO_4 : (a) Phase diagram at 1 atm (T in K); 833 (b) Phase diagram at high-P and high-T conditions (HP-HT). The phase sequence in 834 (a) is from Remy et al. (1997a). The solid univariant curves of $\gamma = \beta$, $\gamma = \alpha'_L$ and $\beta = \beta$ α'_{L} in (b) are from Remy et al. (1995). The solid circle in (b) represents the P-T 835 836 condition of our Lrn-synthesizing experiments, which locates in the stability field of the Lrn phase, but near the high *P*-*T* extension of the $\beta = \alpha'_L$ univariant curve. In 837 838 addition, Liu (1978) observed a K_2NiF_4 -structured Ca₂SiO₄ phase at ~24 GPa and 839 1273 K, and Wang and Weidner (1994) detected the phase transition between the β 840 phase and the α'_L phase at ~9 GPa and 1350 K. 841 842 FIGURE 2. XRD patterns collected at (a) 300 K, (b) 923 K, and (c) 973 K (ambient P). 843 The material underwent the phase transition from the β phase to the α'_L phase at a 844 temperature between 923 and 973 K. A few weak peaks, as indicated by the 845 asterisks, do not belong to the Lrn phase, but are attributable to the γ phase, which 846 might have been generated during the sample-grinding process. The relative 847 intensities of these weak peaks did not change as the experimental T increased, 848 suggesting that the volume proportion of the γ phase did not increase and the 849 potential phase transition from the metastable Lrn to the γ phase did not actually 850 take place at low temperatures (Fig. 1a). 851

9/16

852	FIGURE 3. Variation of the unit-cell parameters of Lrn with T (ambient P): (a) the a -
853	axis; (b) the <i>b</i> -axis; (c) the <i>c</i> -axis; (d) the volume; (e) the angle β . Note that lengths
854	of the error bars are generally equal to or smaller than the symbols.
855	
856	FIGURE 4. XRD patterns of Lrn collected at 3.80, 7.61, 12.30 and 16.25 GPa (ambient
857	<i>T</i>). Note that peak broadening at $P > 10$ GPa was not obvious.
858	
859	FIGURE 5. Variation of the unit-cell parameters of Lrn with P (ambient T for the
860	experimental data whereas zero T for the simulated data): (a) the <i>a</i> -axis; (b) the <i>b</i> -
861	axis; (c) the <i>c</i> -axis; (d) the volume; (e) the angle β . Note that lengths of the error
862	bars of the experimental data are generally equal to or smaller than the symbols.
863	The curves, both solid and broken, are drawn according to the derived third order
864	BM-EoS.
865	
866	FIGURE 6. Eulerian strain-normalized pressure (f_E-F) plots based on the BM-EoS
867	(experimental data only): (a) is for the <i>a</i> -axis; (b) is for the <i>b</i> -axis; (c) is for the <i>c</i> -
868	axis; (d) is for the volume. Estimated standard deviations were calculated following
869	the method in Heinz and Jeanloz (1984). The solid lines are the weighted linear fits
870	through the data.

871

872 **FIGURE 7.** Phonon dispersions and VDoS of Lrn (LDA).

- FIGURE 8. Isobaric heat capacity of Lrn. Our C_V results from the first-principles method were combined with our experimentally-determined volumetric thermal expansion coefficient and isothermal bulk modulus to derive the C_P values.
- 877
- FIGURE 9. Isothermal mode Grüneisen parameters (γ_i) of Lrn (ambient *T*). High-*P* Raman data used in the calculation are from Reynard et al. (1997). Although a full assignment of the Raman-active vibrations of Lrn has not been available, it is generally appropriate to assign the peaks with $v_i > -500$ cm⁻¹ as the internal modes of the SiO₄ tetrahedra and the peaks at lower frequencies as the lattice modes (Remy et al. 1997a).
- 884

885 **FIGURE 10.** *P*-*T* trajectory for trapping the Lrn inclusion by the diamonds from the 886 Kankan district of Guinea (thick curve). The P-T profiles of the cold slab and hot 887 slab are sketched from Thompson (1992), and the *P*-*T* profile for the normal mantle 888 follows the 1600 K adiabat for the simplified pyrolite composition in Jackson 889 (1998). Since the Lrn inclusion always coexists with the Ttn phase (Joswig et al. 890 1999), the phase boundaries between Wal and Lrn + Ttn (equation 1), and between 891 Lrn + Ttn and CaPv (equation 2) are plotted to show the stability field of the phase 892 assemblage Lrn + Ttn (Gasparik et al. 1994; Sueda et al. 2006). The well known 893 phase boundary of graphite and diamond is also sketched.



P(GPa)

Figure 1 of Xiong et al. (2015)



Figure 2 of Xiong et al. (2015)





Figure 4 of Xiong et al. (2015)



Figure 5 of Xiong et al. (2015)



Figure 6 of Xiong et al. (2015)



Figure 7 of Xiong et al. (2015)



Figure 8 of Xiong et al. (2015)



Figure 9 of Xiong et al. (2015)





_					
a (Å)	b (Å)	c (Å)	β (°)	$V(\text{\AA}^3)$	
5.5096(4)*	6.7518(4)	9.3059(8)	94.575(7)	345.07(19)	
5.5101(2)	6.7529(2)	9.3091(4)	94.533(4)	345.31(11)	
5.5126(2)	6.7567(2)	9.3178(4)	94.447(3)	346.02(9)	
5.5157(2)	6.7587(3)	9.3223(5)	94.369(4)	346.52(13)	
5.5183(2)	6.7649(2)	9.3324(3)	94.289(3)	347.41(9)	
5.5214(2)	6.7681(3)	9.3377(5)	94.191(5)	348.01(13)	
5.5246(3)	6.7726(3)	9.3476(7)	94.117(6)	348.85(17)	
5.5278(3)	6.7781(3)	9.3575(7)	94.025(6)	349.74(15)	
5.5303(3)	6.7801(4)	9.3643(8)	93.934(7)	350.31(17)	
5.5332(2)	6.7851(3)	9.3731(6)	93.818(5)	351.11(14)	
5.5351(2)	6.7887(3)	9.3826(6)	93.731(4)	351.81(11)	
5.5379(2)	6.7946(2)	9.3912(5)	93.626(4)	352.67(9)	
5.5407(4)	6.7992(5)	9.4031(9)	93.502(6)	353.57(15)	
5.5431(2)	6.8009(4)	9.4098(8)	93.379(6)	354.11(12)	
5.532(3)	20.591(16)	9.443(6)	-	1075.6(12)	
5.764(4)	21.019(11)	9.855(4)	-	1193.9(11)	
*Number in the parentheses represents one standard deviation in the rightmost digit.					
	a (Å) 5.5096(4)* 5.5101(2) 5.5126(2) 5.5157(2) 5.5183(2) 5.5214(2) 5.5246(3) 5.5278(3) 5.5303(3) 5.5332(2) 5.5351(2) 5.5351(2) 5.5379(2) 5.5407(4) 5.5431(2) 5.532(3) 5.764(4) in the parenthe	a (Å) b (Å)5.5096(4)*6.7518(4)5.5101(2)6.7529(2)5.5126(2)6.7567(2)5.5157(2)6.7587(3)5.5183(2)6.7649(2)5.5214(2)6.7681(3)5.5278(3)6.7726(3)5.5278(3)6.7781(3)5.5332(2)6.7851(3)5.5351(2)6.7887(3)5.5379(2)6.7946(2)5.5407(4)6.7992(5)5.532(3)20.591(16)5.764(4)21.019(11)	a (Å) b (Å) c (Å) $5.5096(4)^*$ $6.7518(4)$ $9.3059(8)$ $5.5101(2)$ $6.7529(2)$ $9.3091(4)$ $5.5126(2)$ $6.7567(2)$ $9.3178(4)$ $5.5126(2)$ $6.7567(2)$ $9.3178(4)$ $5.5157(2)$ $6.7587(3)$ $9.3223(5)$ $5.5183(2)$ $6.7649(2)$ $9.3324(3)$ $5.5214(2)$ $6.7649(2)$ $9.3324(3)$ $5.5214(2)$ $6.7681(3)$ $9.3377(5)$ $5.5246(3)$ $6.7726(3)$ $9.3476(7)$ $5.5278(3)$ $6.7781(3)$ $9.3575(7)$ $5.5303(3)$ $6.7801(4)$ $9.3643(8)$ $5.5332(2)$ $6.7851(3)$ $9.3731(6)$ $5.5351(2)$ $6.7946(2)$ $9.3912(5)$ $5.5407(4)$ $6.7992(5)$ $9.4031(9)$ $5.532(3)$ $20.591(16)$ $9.443(6)$ $5.764(4)$ $21.019(11)$ $9.855(4)$	a (Å) b (Å) c (Å) β (°)5.5096(4)*6.7518(4)9.3059(8)94.575(7)5.5101(2)6.7529(2)9.3091(4)94.533(4)5.5126(2)6.7567(2)9.3178(4)94.447(3)5.5157(2)6.7587(3)9.3223(5)94.369(4)5.5183(2)6.7649(2)9.3324(3)94.289(3)5.5214(2)6.7681(3)9.3377(5)94.191(5)5.5246(3)6.7726(3)9.3476(7)94.117(6)5.5278(3)6.7781(3)9.3575(7)94.025(6)5.5303(3)6.7801(4)9.3643(8)93.934(7)5.5332(2)6.7851(3)9.3731(6)93.818(5)5.5351(2)6.7887(3)9.3826(6)93.731(4)5.5379(2)6.7946(2)9.3912(5)93.626(4)5.5407(4)6.7992(5)9.4031(9)93.502(6)5.532(3)20.591(16)9.443(6)-5.764(4)21.019(11)9.855(4)-	

TABLE 1. Unit-cell parameters of Lrn and α'_L versus *T* (ambient *P*)

TABLE 2. Volumetric thermal expansion coefficients of Lrn (ambient P)^{*}

T range (K)	$a_0 (10^{-5})$	$a_1 (10^{-9})$	Data source	Note
300-923	4.24(4)	-	This study	powder XRD
298-984	4.25(14)	-	Remy et al. (1997b)	powder XRD
298-918	2.87(15)	35(4)	Fukuda et al. (1997)	powder XRD
293, 993 [†]	3.97	-	Barnes et al. (1980)	powder XRD
298	1.80	-	Swamy and Dubrovinsky (1997)	QHLDC [‡]
298-923	2.38(3)	17.8(9)	Holland and Powell (1998)	TOE§
-	4.11(0)	15.4(0)	Swamy and Dubrovinsky (1997)	TOE§

 $a_T = a_0 + a_1(T-T_0) + a_2(T-T_0)^{-2}$, where a_T and a_0 are in K⁻¹, a_1 is in K⁻², and a_2 (always zero in the case of Lrn) is in K. T_0 is set as 298 K.

[†]The reported unit-cell parameters at 293 and 993 K were used in the calculation.

[‡]Quasi-Harmonic Lattice Dynamic Calculation done with the GEMIN algorithm.

[§]Thermodynamically Optimized Estimate.

TABLE 3. Unit-cell parameters of Lrn versus P (ambient T)

P (GPa)	a (Å)	<i>b</i> (Å)	c (Å)	β (°)	$V(\text{\AA}^3)$
0.0001(1)	5.510(1)*	6.753(2)	9.316(1)	94.48(1)	345.6(1)
3.80(10) [†]	5.446(6)	6.685(6)	9.197(2)	94.57(2)	333.7(1)
5.09 (10)	5.428(5)	6.666(6)	9.163(2)	94.59(4)	330.5(3)
6.60(10)	5.406(3)	6.647(6)	9.125(2)	94.61(2)	326.8(1)
7.61(10)	5.394(7)	6.637(7)	9.101(6)	94.63(3)	324.7(2)
8.68(10)	5.380(6)	6.626(5)	9.078(2)	94.66(2)	322.5(1)
9.45(10)	5.371(5)	6.615(1)	9.061(2)	94.67(2)	320.9(4)
10.39(10)	5.362(6)	6.605(4)	9.041(5)	94.69(5)	319.1(1)
11.35(10)	5.351(4)	6.591(7)	9.021(2)	94.72(2)	317.1(2)
12.30(10)	5.341(5)	6.577(6)	9.003(2)	94.73(2)	315.2(1)
13.49(10)	5.329(4)	6.556(1)	8.980(8)	94.75(3)	312.7(3)
14.33(10)	5.319(4)	6.549(1)	8.965(2)	94.78(2)	311.2(1)
15.35(10)	5.308(2)	6.538(4)	8.947(1)	94.80(1)	309.4(1)
16.25(10)	5.297(4)	6.529(1)	8.931(2)	94.82(2)	307.7(4)

*The number in the parentheses represents one standard deviation in the rightmost digit.

[†]Uncertainty in the *P* measurement at high *P* assumed to be 0.1 GPa.

Parameters	Experimental*	Calculate	Calculated by first-principles simulation			
		LDA	R. D. (%) [†]	GGA	R. D. (%) [†]	
<i>a</i> (Å)	5.5041	5.4080	-1.75	5.5678	1.16	
<i>b</i> (Å)	6.7622	6.6310	-1.94	6.8103	0.71	
c(Å)	9.3281	9.1270	-2.16	9.3873	0.63	
$\beta(^{\circ})$	94.172	95.433	1.34	94.466	0.31	
$V(\text{\AA}^3)$	346.27	325.87	-5.89	354.87	2.48	
Ca polyhedron						
Cal-Ol(Å)	2.719	2.682	-1.02	2.752	1.2	
Ca1-O2(Å)	2.671	2.672	0.04	2.685	0.55	
Ca1-O3(Å)	2.459	2.456	-0.09	2.451	-0.32	
Ca1-O4(Å)	2.508	2.482	-1.01	2.513	0.22	
Ca2-O1(Å)	2.255	2.238	-0.77	2.280	1.12	
Ca2-O2(Å)	2.432	2.418	-0.57	2.464	1.31	
Ca2-O3(Å)	2.871	2.856	-0.53	2.907	1.26	
Ca2-O4(Å)	2.735	2.737	0.04	2.774	1.41	
Si tetrahedron						
Si-O1(Å)	1.598	1.562	-2.23	1.617	1.21	
Si-O2(Å)	1.533	1.512	-1.33	1.553	1.34	
Si-O3(Å)	1.470	1.457	-0.89	1.486	1.10	
Si-O4(Å)	1.672	1.660	-0.69	1.692	1.22	
01- Si-O2(°)	109.051	108.092	-0.88	109.891	0.77	
O1-Si-O3(°)	114.907	114.187	-0.62	115.101	0.17	
01-Si-O4(°)	104.195	103.581	-0.59	105.373	1.13	
O2-Si-O3(°)	112.622	112.701	0.07	113.343	0.64	
O2-Si-O4(°)	99.790	99.840	0.05	99.900	0.11	
$O3-Si-O4(^{\circ})$	114.932	114.161	-0.67	114.403	-0.46	

TABLE 4. Comparison between experimental and energy-optimized crystallographic data of Lrn

K_T (GPa)	$K_{T}^{'}$	Method	Data source
114(1)	4*	Powder XRD	This study
103(2)	5.4(4)	Powder XRD	This study
166(15)	4*	Powder XRD	Remy et al. (1997b)
124(1)	4*	First-principles simulation (LDA)	This study
123(3)	4.4(9)	First-principles simulation (LDA)	This study
96(1)	4*	First-principles simulation (GGA)	This study
92(2)	4.9(5)	First-principles simulation (GGA)	This study
119	4*	$QHLDAC^\dagger$	Swamy and Dubrovinsky (1997)
119	4*	Thermodynamic optimization	Swamy and Dubrovinsky (1997)
120	4*	Thermodynamic optimization	Holland and Powell (1998)

TABLE 5. Isothermal bulk modulus (K_T) of Lrn (ambient P)

 K_T' is set as 4.

 $^{\dagger}\textsc{Quasi-Harmonic}$ Lattice Dynamic Calculation done with the GEMIN algorithm.

<i>T</i> (K)	Calculated by First-principles*		$Experimental^{\dagger}$	Experimental [‡]
	C_V	C_P	C_P	C_P
10	0.06	0.17(1)		
20	1.05	1.26(1)		
30	4.08	4.38(1)		
41	9.16	9.56(1)		
51	15.72	16.23(2)	14.36	-
101	52.91	53.93(4)	53.05	-
152	81.59	83.12(6)	81.75	-
202	101.69	103.73(6)	102.60	-
253	116.37	118.92(9)	117.73	-
303	127.46	130.52(11)	129.29	-
354	135.99	139.56(13)		-
404	142.61	147.69(15)		145.50
455	147.80	151.35(16)		151.91
505	151.91	157.01(19)		157.20
556	155.20	160.81(21)		161.59
606	157.86	163.98(22)		165.25
657	160.03	166.67(24)		168.34
707	161.83	168.97(26)		170.98
758	163.32	170.98(28)		173.24
808	164.58	172.75(30)		175.21
859	165.65	174.32(32)		177.94
909	166.56	175.74(34)		178.46
950	167.19	176.79(35)		179.56
1000	167.89	177.99(37)		180.79
[*] This work.				
[†] Todd (1951).				

TABLE 6. Heat capacity $(J \text{ mol}^{-1} \text{ K}^{-1})$ of Lrn at selected T

[‡]Coughlin and O'Brien (1957).

TABLE 7. Coefficients of the C_P polynomials of Lrn (ambient P)^{*}

<i>T</i> = 10-50 K	<i>T</i> = 50-293 K	<i>T</i> = 293-1000 K			
$k_0 = 8.0(2)E-1^{\dagger}$	$k_0 = -3.56(7)E1$	$k_0 = 2.121(1)E2$			
$k_4 = -1.6(2)E-1$	$k_4 = 1.13(1)E0$	$k_1 = -9.69(5)E2$			
$k_5 = 9.2(3)E-3$	$k_5 = -2.77(9)E-3$	$k_2 = -4.1(3)E6$			
	$k_6 = 2.7(1)E-6$	$k_3 = 5.20(7)E8$			
$C_P = \mathbf{k}_0 + \mathbf{k}_1 T^{0.5} + \mathbf{k}_2 T^2 + \mathbf{k}_3 T^3 + \mathbf{k}_4 T + \mathbf{k}_5 T^2 + \mathbf{k}_6 T^3 $ (J mol ⁻¹ K ⁻¹).					

 $C_P = k_0 + k_1 T^{0.0} + k_2 T^2 + k_3 T^3$ [†]E-n represents ×10⁻ⁿ.