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1	REVISION 2
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3 4	Computational study of pressure behaviour to 6 GPa of the $2M_1$ muscovite-paragonite series
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Abstract

22 23

The muscovite (Ms)-paragonite (Pg) series $[K_{1-x}Na_xAl_2(Si_{4-y}Al_y)O_{10}(OH)_2]$ is a group of micas with end members of Ms ($x = 0, y \approx 1$) and Pg ($x = 1, y \approx 1$). This mineral series is found in the Earth's crust and upper mantle. The series shows a wide immiscibility gap between the end members.

Density Functional Theory (DFT) is used to show the compression in five models of the $2M_1$ polytype Ms-Pg series to 6 GPa. Bulk moduli and cell-parameter moduli were obtained from a least square fitting of pressures and volumes to a thirdorder Birch-Murnaghan equation of state. Bulk-modulus values of the end members of the series agree with the range of experimental values. Bond lengths and atomic-group geometries were studied as a function of the pressure and composition of the series by determining the moduli. Compression mechanism has been determined.

The excess volumes, V^{ex} , were higher for the Na-rich members than for the K-35 rich members than the solid solution. V^{ex} follow a Redlich-Kister behaviour. The excess 36 37 free energy, G^{ex} , was calculated isobarically in a semiempirical way: the DFT excess volume data were calculated in one experimental model (A from Roux and Hovis, 38 1996) in a Redlich-Kister function. The G^{ex} as a function of the composition of the Ms-39 Pg join of the A model show two minima with constant composition to 0.75 GPa, 40 evolving to richer end member compositions at greater pressures. Therefore, the solvus 41 42 should increase the gap of immiscibility at high pressure.

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Keywords: Muscovite-paragonite series; DFT calculations; pressure behaviour of
atomic groups and crystal geometry; bulk modulus; cell-parameter moduli; bonds and
atomic-group moduli; compression mechanism; excess volume; excess free energy.

INTRODUCTION

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The muscovite (Ms)-paragonite (Pg) join is a series of dioctahedral 50 phyllosilicates of the true, white-mica group, with Na^+ and K^+ substitutions in the 51 interlayer, with a general formula $[K_{1,x}Na_xAl_2(Si_{4,y}Al_y)O_{10}(OH)_2]$. These minerals, 52 found in the Earth's crust and upper mantle, are among the most common and abundant 53 minerals in metamorphic rocks. Ms ($x = 0, y \approx 1$) and Pg ($x = 1, y \approx 1$) are the end 54 55 members of the series where an immiscibility gap exists (Guidotti et al., 1992). Natural samples show minor concentration of cations, which involve sites in the interlayer 56 (Ca²⁺, Ba²⁺, Sr²⁺, etc.), octahedra (Oc/VI) (Mg²⁺, Fe²⁺, Fe³⁺, Ti²⁺, and Mn²⁺) and 57 tetrahedra (T/IV) sheets. The effect of cation substitutions in the Oc, T sheets and 58 59 interlayer on the cell dimensions was studied by Guidotti et al. (1992). The Ms-Pg 60 series has the potential to be a geothermometer and a geobarometer (Guidotti et al., 61 1994; Blencoe et al., 1994).

62 Pressure behaviour of Ms was studied by different techniques, revealing a strong 63 anisotropy in the compressional behaviour (Comodi and Zanazzi, 1995). Table 1 64 presents the values of the bulk modulus of Ms and Pg with different experimental techniques. Bulk-modulus values of Ms are given from 49.0 GPa (Holland and Powell, 65 66 1998, 2011) to 61.4 GPa (Faust and Knittle, 1994). These variations in results may be 67 attributed to different samples, such as heterogeneities of composition, techniques, and 68 use of powder or single-crystal samples. In addition, in powder samples subject to high 69 pressure, preferred orientations of crystallites are possible. The bulk modulus of Pg was 70 also studied, showing more rigid values (65.5 GPa, Comodi and Zanazzi, 1997) than 71 Ms.

Ortega-Castro et al. (2010) by means of DFT calculations of Ms with increasing pressure to 6 GPa determined a bulk-modulus value of 60.1 GPa. They studied each atomic group as a function of pressure to yield the compression mechanism. This value is close to that of Faust and Knittle (1994). Teich-McGoldrick et al. (2012), by means of dynamical simulations, obtained a bulk modulus of Ms of 59.8 GPa (0.01 GPa of pressure and T=298 K). They found that the bulk moduli decreased with increasing temperature and increased with increasing pressure, as expected. From DFT-calculated

elastic constants, polycrystalline bulk moduli of the series have been found to beslightly larger than the known experimental values (Hernández-Haro et al. 2013).

Atomic-group variations as a function of pressure was also studied (Comodi and 81 Zanazzi, 1995, 1997, and Guidotti et al., 2000). The micro-environments crystal-82 83 chemical and structural properties evolve as a function of the concentration of one 84 cation in a solid solution, pressure, and temperature and they are related to the thermodynamic properties, stability, and evolution of the minerals. However, most of 85 the experiments use macroscopic and thus heterogeneous samples, giving average 86 values of the inter-atomic geometry and properties of the crystal. The use of crystal 87 models free of order/disorder, average compositional effects, heterogeneity, and 88 89 different experimental preparations can give values of inter-atomic geometry and properties that could be the key to accounting for the nano- and micro-behaviour, and 90 design of new materials. DFT studies, can offer a detailed explanation for the crystal-91 92 chemical behaviour of the solid solutions (Geiger, 2008) without the cited drawbacks of 93 the experimental methods.

Owing to the effects of pressure on the stability of Ms-Pg pairs and the K/Na 94 ratios of Na-saturated Ms, practical application of the Ms-Pg solvus thermometry is 95 restricted to equilibrated quasi-binary Ms-Pg pairs at pressures between ca. 0.2 and 0.8 96 97 GPa (Blencoe et al. 1994). The effect of pressure on the Ms-Pg solvus is manifested by the effect on the excess Gibbs free energy (G^{ex}) , which varies with the Margules' 98 coefficients (W_G) of excess volume (V^{ex}) (Ganguly, 2008, and references therein). There 99 are major disparities in the literature for experimental molar excess volume data of the 100 101 Ms-Pg system, making it difficult to constrain the functional form (quadratic, cubic, etc. polynomial) of V^{ex} (Eugster et al. 1972; Blencoe, 1977; Chatterjee, 1974; Flux and 102 103 Chatterjee, 1986, Chatterjee and Flux, 1986; Chatterjee and Froese, 1975; and 104 Chatterjee and Johannes, 1974). Because natural Ms-Pg micas are unstable below the solvus, experimental work in synthetic and natural micas only constrain the volumes in 105 the compositional ranges near the end members. Furthermore, V^{ex} of Ms-Pg micas are 106 small (commonly < 0.13 J/bar), so minor errors in the experimental molar volume affect 107 significantly the accuracy of the V^{ex} . Phase equilibrium data indicate a larger positive 108 V^{ex} than X-Ray measurements, particularly on micas synthesized at very high pressures 109 110 (Roux and Hovis, 1996). Because there is no such limitation in theoretical methods, the present study may constrain the functional form of V^{ex} of Ms-Pg micas. 111

We have previously studied the pressure behaviour of Ms (Ortega-Castro et al., 112 2010) and the elastic properties of the Ms-Pg series (Hernández-Haro et al., 2013) by 113 114 DFT methods. From the elastic constants, the calculated bulk moduli correspond to the 115 average polycrystalline bulk moduli from a Hill average (Hill, 1952) of the upper and 116 lower Hashin and Shtrikman limits (Hashin and Shtrikman, 1962a and b). In this article, we study the pressure behaviour of the series in a periodic single-crystal model, 117 118 increasing the pressure sequentially with the following aim: to ascertain theoretically how the crystal structure, geometry of the atomic groups, mixing behaviour, and 119 120 incompressibility moduli evolve with pressure and composition; and to clarify the influence of pressure on the Ms-Pg solvus. 121

METHODS

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125 Crystal model

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Each Ms-Pg series model is defined by the following: $[K_{1-x}Na_xAl_2(Si_3Al)O_{10}(OH)_2]$ (x 127 128 between 0 and 1). Crystal models used are similar to Hernández-Haro et al. (2013). The 129 models of the 2M₁ polytypes are monoclinic. Figure 1 shows the model of X_{Na} = Na/(Na+K)= 0.5 (atomic fraction of the mica). In Ms with $X_{Na} = 0$, cations in both 130 interlayers are K⁺ and they are increasingly substituted by Na⁺ to produce the five 131 132 studied compositions of the Ms-Pg series to $X_{Na} = 1$ for Pg. A supercell of 1x1x2 is used 133 for the $2M_1$ polytype $X_{\text{Na}} = 0$, 0.5, and 1. Interlayer cations with the same atomic 134 number are in opposite sides of the layer in both interlayers for $X_{\text{Na}} = 0.5$ (Fig. 1). Compositions with $X_{\text{Na}} = 0.25$ and 0.75 require a 2x1x2 supercell, with cations of the 135 136 same atomic number in opposite sides in both interlayers (Hernández-Haro et al. 2013).

137 Computational methods

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139 The Density Functional Theory (Hohenberg and Kohn 1964; Kohn and Sham 1965) 140 calculations were performed with the SIESTA2.0.2 code (Sánchez-Portal et al. 1997; 141 and Artacho et al. 1999; Soler et al., 2002). Periodic boundary conditions were applied 142 to the crystal cells. The atomic groups and cell parameters were optimized together in 143 the same simulation. Generalized gradient approximation (GGA) with the Perdew-144 Burke-Ernzerhof (Perdew et al. 1996) (PBE) correlation-exchange functional was used. A mesh cut-off of 500 Ry and 15 and 9 k-points in the Monkhorst Pack grid were found 145 146 to be adequate to produce rapid convergence in the 84 and 168 atom cells, respectively. 147 Our calculations were performed with numerical atomic orbitals, and double- ζ plus 148 polarization basis sets. Troullier-Martins norm-conserving pseudopotentials (Troullier and Martins 1991) were used. Muscovite- $2M_1$ is monoclinic with a C2/c space group. 149 150 All calculations were performed by fixing the desired pressure, and then optimizing cell 151 parameters and atomic positions to reach convergence thresholds of forces and stresses (0.005 eV/Å and 0.002 GPa, respectively). The applied pressures ranged from -1 to 6 152

GPa (negative values are tractions) in increments of 0.25 GPa between -1 and 1 GPa,and 0.5 GPa between 1 and 6 GPa.

The pseudopotentials and basis sets used here were previously optimized for a 155 set of dioctahedral 2:1 phyllosilicates (Ortega-Castro et al. 2008, 2009). In a previous 156 157 paper, the high-pressure behaviour of Ms was calculated in the local density 158 approximation (LDA) with the Ceperley-Alder exchange-correlation functional 159 (Ceperley and Alder 1980) and the GGA-PBE approach. The results showed that the GGA yielded cell parameters closer to the experimental data than LDA (Ortega-Castro 160 161 et al. 2010; Hernández-Haro et al. 2013). In addition, White et al. (2009) demonstrated 162 that the GGA functional yielded low energies and accurate geometries in kaolinite. 163 Tunega et al. (2012) analysed the role of dispersion corrections in DFT calculations in 164 talc, pyrophyllite, kaolinite and lizardite, employing different functionals with and 165 without dispersion forces, and found that using the dispersion corrections produced 166 better results, especially for c axis values. In those systems, where no cation 167 substitutions were present, the atoms across the interlayer were linked mainly by van 168 der Waals forces, whereas in the mica systems, where cations are present in the 169 interlayer, the most important forces are the Coulomb forces, well described by DFT 170 methods.

171 Equation of state

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The bulk modulus, B_0 , at P=0 GPa was calculated by fitting a Birch-Murnhagan (BM) equation of state to volumes and pressures [Birch, 1947, EOSFIT5.2 (Angel, 2000, 2001)]. The incompressibility of the cell-parameter linear moduli was also calculated by using a cubic crystal with a cell parameter equal to one of the monoclinic cell parameters, and fitting with the aid of EOSFIT 5.2 code (Angel 2000 and 2001). Most of the B_0 values were calculated from a third-order Birch-Murnaghan equation fitting (BM3).

180 Solid solutions

In a solid solution, the value of any cell parameter of any member, *i*, of the solution can
be expressed with respect to the end members by an ideal mixing law, which, for any
cell parameter, is defined by Vegard's law:

$$a_i = \sum_j X_{ij} a_j \tag{1}$$

where X_{ij} is the molar fraction of the *j* end member of the solid solution in any *i* member of solid solution. In our solid solution the numbers of end-members are two. A similar law is given for the volume of the crystal cell:

189
$$V_i = \sum_j X_{ij} V_j \tag{2}$$

190 Non-ideality can be expressed by mixing excess terms (Waldbaum and Thompson1968), which for the case of two components in the solid solution can be written as:

192
$$\Delta V_i^{ex} = W X_{ij} (1 - X_{ij})$$
(3a)

193
$$\Delta V_i^{ex} = X_{ij} (1 - X_{ij}) [W_1 (1 - X_{ij}) + W_2 X_{ij}]$$
(3b)

194

W are the Margules' coefficients, where W at Eq. (3a) is considered a symmetric excess mixing term of a regular mixing model. In the Eq. (3b), W_1 and W_2 are asymmetric excess mixing terms (Waldbaum and Thompson 1968, Hernández-Haro et al. 2013) belonging to a subregular mixing model; W_1 is associated with the increasing concentration cation considered in the function, and W_2 is associated with the other cation.

The systematization of the bulk modulus as a function of the substitution-cation in the solid solution is established from Eq. (2) by calculating the derivative with respect to pressure, resulting in (Takahashi et al. 1970):

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$$\frac{V_i}{B_{0i}} = \sum_j X_{ij} \frac{V_j}{B_{0j}}$$
(4)

where B_{0i} and B_{0j} are the bulk modulus of any member and one of the end members of the solid solution, respectively. This equation can be extrapolated to any modulus of the cell parameters, and to any other bond or atomic group (*u*) of the crystal:

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$$\frac{u_i}{B_{u0i}} = \sum_j X_{ij} \frac{u_j}{B_{u0j}}$$
(5)

209 Excess free-energy

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The Gibbs free energy of any member, *i*, of a solid solution, at a temperature (T)
and pressure (P) is given by (Saxena et al. 1993, Fabrichnaya et al. 2004):

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$$\Delta G_i^{mix}(T,P) = \sum_{j=1}^n X_{ij} G_j(T,P) + RT \left(\sum_{j=1}^n X_{ij} \ln X_{ij} \right) + \Delta G^{ex}(T,P)$$
(6)

where *n* is the number of end members and G_j is the free energy of any end member forming the solid solution in an ideal mixing equation. The second term is the configurational entropy of the solid solution terms, and the last term is the excess free energy.

The excess of free energy is given by a Redlich-Kister approach (Redlich and Kister 1948, Ganguly 2008) by a function:

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$$\Delta G^{ex}(T,P) = X_{ij} (1 - X_{ij}) \left[A + B (1 - 2X_{ij}) + C (1 - 2X_{ij})^2 \right]$$
(7)

Where the *A*, *B* and *C* coefficients are divided in enthalpy, entropy and volume contributions, such as: $A=A_H+A_sT+A_VP$, and with *B* and *C* being coefficients of the same partition. Chatterjee and Flux (1986) and Roux and Hovis (1996) used this equation to study the Ms-Pg system.

Although the free energy and the dependence with temperature of the free energy can be calculated theoretically from zero point energy (ZPE) and thermal effects, the number of atoms of our crystal models requires substantial computational resources. For these reasons, excess free energy is calculated isobarically by a classical and semiempirical approach, adding our computational excess free energy coefficients (Eq. 7) depending on the pressure to the experimental enthalpy and entropy coefficients of Eq. (7).

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RESULTS AND DISCUSSION

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235 Crystal structure at room pressure

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237 Cell parameters and volume variation as a function of the Na⁺ content was studied in Hernández-Haro et al. (2013). The main results of that study are summarized and used 238 239 here (Table 2) for consistency and for comparison purposes. Cell parameters a, b, 240 $0.5csin\beta$ and volumes in Ms and Pg at a pressure of 0 GPa were in agreement with the 241 experimental and average. Tetrahedral cation-oxygen bond averages (<T-O>) and octahedral aluminium –oxygen linkages averages (${}^{VI}Al^{3+}$ - O>) were in agreement 242 with the experimental values (<1%). Si and Al^{3+} tetrahedra, Al^{3+} octahedra volumes, 243 and the Oc-, T-sheet, and interlayer thicknesses (Table 2) yielded values close to the 244 245 experimental results. Average distances of interlayer cations with the further (outer) and 246 closer (inner) basal oxygen atoms (< X...O_{outer/inner}>, $X = Na^+$ or K⁺) were consistent with the experimental data (Table 2). The behaviour of these distances with respect to $X_{\rm Na}$ 247 was approximately linear, decreasing with the Na⁺ content. Tetrahedral rotation, α , 248 linearly increased with X_{Na} (Table 2), the average values and slope of $<\alpha>_{\text{Na}}$ being 249 higher than $\langle \alpha \rangle_{\rm K}$. This effect may be related to the highest charge/radius ratio of Na⁺ 250 with respect to K^+ , which increases the tetrahedral rotation (Muñoz-Santiburcio et al. 251 252 2011).

The calculations were performed at 0 K and 0 GPa without any calculation of ZPE and thermal effects, although the experimental data are from samples at room temperature and pressure. The agreement may be related to error cancellation, owing to the basis set size, pseudopotentials, correlation-exchange functional parameters, and the remaining approximations. Nonetheless, the PBE correlation-exchange functional is a nonempirical GGA functional, which yields reliable results (Goerigk & Grimme, 2011), especially for cell volumes (Kurth et al. 1999).

260

261 Crystal structure at high pressure

262 Ms

The *a* axis values of Ms decrease as a function of pressure, being approximately linear 263 from -1 GPa (Fig. 2a). They are compared with the experimental values in Fig. 2a. The 264 265 calculated value of the *a* axis incompressibility modulus is 490.6 GPa (Table 3), which 266 is larger than the experimental value of 337.8 GPa (Comodi and Zanazzi, 1995), and 267 400 GPa (Curetti et al., 2006). In this case, no temperature correction was applied to our 268 value, which could yield smaller values of the modulus of the axis, and possibly be 269 closer to the experimental value. In addition, the minerals of Comodi et al. 2002, Curetti 270 et al. (2006), and Gatta et al. (2010) showed phengitic substitutions that could justify the 271 differences with respect to our values. The value of the b axis as a function of pressure 272 is close to the experimental values (Fig. 2a). The calculated incompressibility modulus 273 (413.7 GPa, Table 3) of the b axis differs from the experimental values of 295.0 GPa 274 (Comodi and Zanazzi, 1995) and 342 GPa (Curetti et al. 2006).

Values of $0.5csin\beta$ as a function of pressure show a small nonlinear behaviour (Fig. 2b), and values of compressibility show a quasi-parallel behaviour with the experimental values. However, experimental and computational values slightly diverge at high pressures. The incompressibility modulus of $0.5csin\beta$ shows a sharp decrease with respect to the *a* and *b* axes, 79.9 GPa (Table 3), which agree with the experimental value (88.2 GPa, Comodi and Zanazzi, 1995). This sharp decrease is related to the weak Coulomb forces in the interlayer, which soften the incompressibility modulus.

The differences between the calculated incompressibility moduli of *a* and *b* and experimental values may be related to three factors: 1) approximations of the DFT methods; 2) the EoS fitting procedure; and 3) a temperature correction is not applied to these calculated values. The temperature correction would surely soften the moduli, thereby making the computational values closer to the experimental values.

287 The volume as a function of pressure also agrees with the experimental values 288 (Fig 2b). The calculated bulk modulus shows values of 60.0 GPa and 55.6 GPa (with 289 temperature correction from Comodi et al. 2002), which agree with the range of 290 experimental values (Table 1). The value of the first derivative of the bulk modulus with 291 respect to pressure, B_0 , is 7.9 (Table S1), which may be consistent with the 292 experimental value of phengite of 6.97 from Curetti et al. (2006), and Gatta et al. 293 (2010), and 6.9 of Faust and Knittle (1994). The closer values between the bulk moduli 294 and the $0.5csin\beta$ moduli indicate that the c cell parameter modulus determines the bulk

295 modulus. Thus, when pressure is applied to the crystal, the stiffness of the entire 296 structure is dominated by the softest direction of the crystal, and this is where the 297 interlayer is located in the phyllosilicates.

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299 Na-rich Ms

The natural sample with a known experimental bulk modulus has $X_{\text{Na}} = 0.37$ (Comodi and Zanazzi, 1995, 1997), which is between the $X_{\text{Na}} = 0.25 - 0.5$ models. Values of *a* and *b* of the $X_{\text{Na}} = 0.25$ member as a function of pressure are in agreement with known experimental values (computational and experimental values of *a*, *b*, $0.5csin\beta$, and *volume* are seen in Figures S1a and S1b).

305 The values of a and b moduli of $X_{\text{Na}} = 0.25$ decrease as a function of the Na⁺ content in the series (Table 3). Natural Na-rich Ms shows values (Comodi and Zanazzi, 306 1995) lower than both calculated values, following the same trend as in Ms (Table 3). 307 308 An additional softening is expected with a temperature correction. In general, the values 309 of B'_{a0} for $X_{Na} = 0.25$ and = 0.50 are the lowest of the series (Table S1). The value of $0.5csin\beta$ modulus increases as a function of the Na⁺ content (84.5 and 95.9 GPa, for X_{Na}) 310 =0.25 and 0.50, respectively), owing to the smaller interlayer as a consequence of 311 shorter ionic radius of the Na^+ , the repulsive forces of the negative charged basal O 312 atoms of both layers becoming greater. The experimental value of the c modulus for Na-313 rich Ms ($X_{Na} = 0.37$) is 98.9 GPa (Comodi and Zanazzi, 1995), which is close to our 314 results for $X_{\text{Na}} = 0.5$. 315

Our calculated bulk moduli for X_{Na} is equal to 0.25 and 0.50 (60.3 and 62.7 GPa, respectively). The former is slightly higher than the Ms bulk modulus (Table 3). Both values are higher than the experimental bulk modulus of the Na-rich Ms sample, which is 60 GPa (Table 1 and 3). The temperature corrections (Comodi et al. 2002) yield values smaller than the experimental value (Table 3). Values of B_0' of both models are similar to the value for Ms, indicating a similar variation rate for the resistance to the increasing hydrostatic pressure for this series of minerals (Table S1).

323 *Pg*

The values of B_{a0} and B_{b0} for $X_{Na} = 0.75$ and Pg ($X_{Na}=1$) are lower than previous 324 models, (Table 3), reaching values of 284.4 and 246.3 GPa, respectively. The 284.4 325 326 value agrees with the experimental value of Pg ($X_{Na} = 0.88$, Comodi and Zanazzi 1997). 327 Values of B_{0}^{\prime} of cell axes are close to the above models but, in general, they are the 328 largest in the series (Table S1). The value of $B_{0.5csin\beta}$ is the largest in the series, reaching 128.4 GPa for Pg. The experimental value (Table 3) is between the values of $X_{\text{Na}} = 0.75$ 329 330 and Pg. This large resistance of Pg to the compression along [001] may be a consequence of the shortest distances between the layers because of the smallest ionic 331 332 radius of the full substitutions of K^+ by Na^+ .

The bulk moduli of both samples are 65.1 and 65.2 GPa, which are equal to the experimental value of Comodi and Zanazzi (1997). By using a temperature correction (Comodi et al. 2002), these values are approximately 3 GPa lower than the experimental values.

The equilibrium volumes are generally calculated by minimization of the electronic energy, forces and stresses at 0 K. Small errors in the calculated 0 GPa and 0 Kelvin volumes would be magnified when calculating the bulk modulus; if V_0 is overestimated, B_0 is underestimated (Kurth et al. 1999). Nonetheless, we obtained consistent results, and where the errors seem to be larger, as in the cell parameter incompressibility moduli, they affect the bulk moduli minimally.

343 When the bulk moduli of polycrystalline samples of the series are determined 344 from the computational elastic constants from the average (Hill, 1952) of the upper and 345 lower limits of the Hashin-Shtrikman approach (Hashin and Shtrikman, 1962) different results are found: 68.4 and 68.8 GPa for Ms and Pg, respectively (with a minimum at 346 347 the middle content of Na^+ (Hernández-Haro et al. 2013). However, when temperature corrections were empirically introduced, values are closer to the experimental results. 348 349 Variations might be related to the approaches intrinsic to DFT calculations, the Hashin-350 Shtrikman model (Hashin and Shtrikman 1962) and the Hill average (Hill, 1952).

351

352 **Compression process**

354 *Atomic groups compression*

In general, the hydrostatic pressure compresses the entire crystal structure, the bonds, bond angles, atomic groups and polyhedra. Incompressibility moduli of bonds and atomic groups are calculated as the inverse of the compressibility values of the average deformations of bonds and atomic groups and fitted by the least square method to first, second and thrird order polynomia.

360 Our calculated Si/Al - Obasal/apical average distances decrease with pressure 361 (average values as a function of pressure are in Figure S2), and have very high 362 incompressibility moduli (Table 4). The Si - O_{apical} bonds are shorter than the Si - O_{basal} bonds. The ^{IV}Al - O bond distances are larger than the Si - O bonds, and the moduli 363 smaller than for Si – O bonds (-33%). Overall, Al - O bonds show a similar behaviour 364 as Si - O bonds. The Si/Al – O bond lengths as a function of pressure were studied by 365 Comodi and Zanazzi (1995, 1997) and Guidotti et al. (2000), who found small negative 366 367 compressibilities. Positive compressibilities were found in our work with high incompressibility moduli, distinguishing Si/Al – O_{basal} or O_{apical} bonds. It is well known 368 that Si –O bonds are mostly incompressible (Levien and Prewwit, 1981; and Hazen and 369 370 Finger, 1977) although, in some cases, this is dependent on the pressure range (Ralph 371 and Ghose, 1980; and Hugh-Jones and Angel, 1994). Our high values of 372 incompressibility in the Si/Al - O bonds may be considered the limit where a small compressibility could change from positive to negative values with small changes in the 373 374 DFT electronic parameters.

The calculated <OH> bond distances are around 0.974 Å, which agrees with the 375 experimental values of 0.998 Å (Mookherjee et al. 2002 in a $2M_1$ phengite). They are 376 377 constants in the Ms-Pg series model (Hernández-Haro et al. 2013). When pressure is applied to the crystal, OH bonds enlarge, and a negative compressibility occurs, 378 decreasing linearly with the Na⁺ content from $X_{Na} = 0.25$. The negative compressibility 379 380 was observed in Ms by vibrational spectroscopy as a "bond softening" (Butler and 381 Frost, 2006). From IR spectra, Williams et al. (2012) found that hydrogen bonding 382 increases with pressure.

Volumes of the Si tetrahedra and their trend in the series agree with the experimental values (Table 4). Volumes decline with increasing pressure, whereas the experimental Si tetrahedron volumes show a minor increase for Pg, which is consistent

with the behaviour of Si/Al - O bonds. Volumes of Al³⁺ tetrahedra are larger and incompressibility moduli are smaller than the Si tetrahedra and show the same trend (Table 4).

The tetrahedral-sheet thickness (T-thick) values also agree with the experimental values (Table 4), decreasing with increasing pressure, similar to the experimental results; values of their moduli are higher than the experimental values.

392 Volumes of the octahedral polyhedra agree with the experimental values in Ms, 393 but, in Pg, calculated volumes are slightly higher than the experimental values (Table 394 4), which decrease with pressure, in agreement with experimental values. The modulus of the octahedral volumes decreases with the Na⁺ content in the series, and they are 395 higher than the experimental values. Octahedral-sheet thickness in Ms shows values 396 similar to the experimental values (Table 4). Calculated incompressibility moduli are 397 398 larger than the experimental values but they show the same behaviour with increasing 399 Na⁺ content. The octahedral-sheet thickness is much more compressible than the 400 tetrahedral sheet.

The α angle increases with pressure. Calculated values at room pressure agree 401 402 with the experimental observation (Table 2), and $\alpha_{\rm K}$ values (Fig. 3a) and slopes as a function of X_{Na} are smaller than α_{Na} . When pressure is higher, interlayer cations are 403 introduced into the ring cavity, but because the ionic radius of Na^+ is smaller than K^+ . 404 tetrahedra must change their rotation to allow a better fit of the smaller Na⁺. Because the 405 K^{+} cavities are located near the Na⁺ cavities, the former change their tetrahedral rotation 406 407 as a consequence of the rotation of the latter, and this effect is propagated along the 408 layer, causing a stabilizing/destabilizing effect in the cavities. At high pressures, this 409 effect is magnified because the increasing value of α as a function of pressure, $\alpha_{K/Na}$ show negative compressibility. The incompressibility moduli of both α_K and α_{Na} are 410 negative and decline as a function of the Na⁺ content in the series (Fig. 3b, Table 4). 411 412 This effect is more pronounced in $\alpha_{\rm K}$.

413 The $K^+/Na^+ \cdots O_{outer/inner}$ bond distances yield moduli between 64 and 336 GPa 414 (Table 4). The moduli of $K^+/Na^+ \cdots O_{outer}$ are larger than $K^+/Na^+ \cdots O_{inner}$, and $Na^+ \cdots$ 415 O_{outer} modulus in Pg is the largest. Although the *outer* distances might be expected to be 416 more compressible than the *inner* distances because they are larger, the *inner* O atoms

are more greatly affected by the Coulomb field of interlayer cations than the *outer* O,

- 418 consequently showing greater compressibility. In general, $Na^+ \cdots O_{outer/inner}$ shows lower
- 419 incompressibility moduli than $K^+ \cdots O_{outer/inner}$, with the exception of Na $\cdots O_{outer}$ of Pg.

The key structural feature in phyllosilicates is the interlayer, which decreases as a function of Na⁺ content (Fig. 3c, Tables 1 and 4), but the slopes of the functions *interlayer thickness=f(X_{Na})* are smaller at increasing pressure. This behaviour is due to the small ionic radius of the Na⁺, which reduces the interlayer with its increasing of Na content. In the Pg end members the basal O atoms of both layers are sufficiently close and repulsion is larger. Thus the values at different pressures of the interlayer thicknesses are closer together at the Pg side than at the Ms side of Figure 3c.

427 The interlayer thickness incompressibility moduli are the smallest of the 428 structure and agree with the experimental data (Table 4). In Ms, Williams et al. (2012) 429 found two possible mechanisms of compression in the interlayer, below and above the 430 range of 5-8 GPa. The interlayer-thickness/interlayer-thickness-modulus relationship 431 [Eq. (5)] of the series decreases with X_{Na} (Fig. 3d), and is in agreement with the volume/bulk modulus relationship [Eq. (4)] of the series (Fig. 3e). This behaviour 432 433 shows asymmetric mixing terms, indicating different mixing behaviour for the Ms and 434 Pg sides of the Ms-Pg join.

There is disagreement between the calculated and experimental moduli of some structural features of the crystal but not with the interlayer moduli, indicating that the key structural feature determining the bulk modulus is the interlayer.

438 *Compression mechanism*

This series shows great anisotropy, with the bulk moduli of the *a* and *b* axes being 6.1-2.2 (Ms-Pg) times larger than $B_{0.5csin\beta}$. When pressure is applied, the interlayer decreases substantially, interlayer cations penetrate into the tetrahedral ring cavities, and changes accordingly. In addition, the K⁺/Na⁺···O_{inner/outer} distances decrease (more for K⁺/Na⁺···O_{inner} than K⁺/Na⁺···O_{outer}). In the layer, the Oc sheet compresses to a larger extent than the T sheets, but to a much lesser extent than the interlayer. Pg members of the series show more resistance to hydrostatic pressure than Ms members of the series.

447 **Solid solutions**

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449 Excess Volume of the Ms-Pg join

Chatterjee and Flux (1986) and Roux and Hovis (1996) used a Redlich-Kister mixing 450 equation [Eq. (7)] to account for the G^{ex} of micas in the Ms-Pg join. In Eq (7) V^{ex} was 451 452 reduced to a subregular model [Eq (3b)] because the coefficient dependent on the volume, C_{ν} , was set to 0. Hernández-Haro et al. (2013) showed that the mixing volume 453 of micas in the Ms-Pg join at 0 GPa followed a regular mixing law with small positive 454 mixing terms. Our choice of a regular mixing volume equation instead of a subregular 455 456 model equation was based on the smaller standard deviations obtained by the quadratic 457 polynomial fitting relative to those obtained by the cubic polynomial. As shown below, the calculated excess volumes of micas in the Ms-Pg join indicate that neither the 458 459 regular nor the subregular models are appropriate.

The excess volume of mixing on the Ms-Pg join is (Blencoe, 1977):

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$$V_{Pi}^{ex} = V_{Pi} - \sum_{i=1}^{2} X_{ii} V_{Pi} \qquad (16)$$

where V_{Pi} are the calculated volumes of any member of the solid solution, V_{Pi} are the 462 volumes of the end members of the Ms-Pg join at the temperature and pressure of 463 interest. For comparison to previous studies, we report the pressure in kbar and X_{ii} is the 464 atomic fraction of K in the mica $(V^{ex}-X_K)$ instead of the Na atomic fraction $(V^{ex}-X_{Na})$ 465 which we have previously used. V^{ex} is calculated at 0, 2.5, 5.0 7.5 15, 30 and 60 kbar. 466 467 V^{ex} at 0 and 60 kbar is depicted in Figs. 4a-b (in Figure S3 pressures of 7.5 and 15 kbar are shown). In general, the V^{ex} shows higher values for Na-rich than for K-rich 468 members, and V^{ex} shows relative minima for pressures ≤ 7.5 kbar, and maxima for 469 pressures ≥ 15 kbar. The largest V^{ex} for the richest Pg compositions may be related to the 470 highest repulsions between the layers as a consequence of the smallest distances 471 472 between them, because of the lower ionic radius of Na⁺. The minima at the approximate center K⁺ concentration at the low pressure Ms-Pg joins may arise from the symmetrical 473 cation configurations imposed in our computational models. This result occurs from a 474 balance of empty space in the K^+ and Na^+ cavities with the models being more 475 476 consistent with the ideal mixing behaviour than its lower and upper concentration solid

477 solution members. At high pressure, V^{ex} of the Pg-rich members are the largest in the 478 series.

The high asymmetry indicates that the V^{ex} of the Ms-Pg solid solution could be 479 better fitted to a Redlich-Kister model instead of subregular models, as determined by 480 481 Blencoe (1977) for the 1M and 2M₁ polytypes, Chatterjee and Froese (1975) for the 2M₁ 482 polytype for the end members of the solid solution, and Eugster et al. (1972) for the 1M483 polytype. Furthermore, in contrast to previous studies (Chatterjee and Flux 1986; and Roux and Hovis, 1996), the Redlich-Kister fitting of DFT V^{ex} requires $C_V \neq 0$, implying 484 that a fourth degree polynomial is required to account for the variation of V^{ex} on the Ms-485 Pg join (Fig. 4a and d). However, the limited data points in the $V^{ex}-X_{k}$ functions do not 486 487 allow fitting except by a quartic polynomial interpolation of the DFT data. The A_V, B_V 488 and C_V terms in Eq. (7) obtained with this interpolation are given in Table 5. The DFT 489 A_V and B_V values are similar to those reported in previous experimental and 490 thermodynamic studies using the Redlich-Kister model (Chatterjee and Flux 1986; 491 Roux and Hovis, 1996), and those derived from Margules' parameters of the subregular 492 model for 1*M* polytypes (Bencloe, 1977; Chatterjee and Froese, 1975) (cf. Table 5). 493 DFT results show the high asymmetry of V^{ex} and its strong pressure dependence. The V^{ex} data were obtained from the DFT calculations at 0 K, but the volumes at different 494 495 temperatures were corrected by thermal expansion coefficients (Holland and Powell, 496 1998). The A_V , B_V and C_V are given in Table 5 at different temperatures (298 - 993 K). 497 A_V and C_V present qualitatively distinct behaviour for low and high pressures (Table 5). 498 In contrast, the B_V does not show any clear systematic variation with pressure and is 499 always the smallest of the three coefficients.

500 *Excess Free energy*

Chatteriee and Flux (1986) and Roux and Hovis (1996) computed the G^{ex} , and the 501 502 corresponding enthalpy, entropy and volume excess coefficients using a Redlich-Kister 503 mixing model [Eq. (7)]. Roux and Hovis (1996) presented four models for the G^{ex} , 504 depending on the number of variables, measurements, and samples introduced in the 505 fitting. In their model A, they included all variables, measurements, and samples. Here, we compute G^{ex} in a semiempirical way, using in Eq. (7) the Roux and Hovis' (1996) 506 507 coefficients for excess enthalpy and excess entropy and our DFT coefficients for V^{ex} . Figure 5 shows the computed G^{ex} as a function of $X_{\rm K}$ for the A model of Roux and 508

Hovis (1996) at constant temperature (693 K) and selected pressures to 60 kbar. The 509 $G^{ex}-X_K$ functions for Model A (Fig. 5) have two minima at all pressures, one for the Pg-510 511 rich micas and the other for the Ms-rich micas. These minima are related to the bimodal 512 compositions of the Pg-Ms solvus at different pressures. The maxima between both 513 minima is approximately located at the center-composition phase, increasing their 514 relative values with respect to the minima at increasing pressure. The composition of 515 the minima are nearly constant to 7.5 kbar, and are increasingly displaced towards the 516 end-member components at higher pressures, indicating that the solvus and miscibility 517 gap in the Ms-Pg join widen with increasing pressure. This behaviour is consistent with the observed evolution of the Pg-Ms solvus with pressure (Roux and Hovis, 1996). 518

The behaviour of the model with our V^{ex} coefficients is similar to the behaviour found by Roux and Hovis' (1996), but giving a quantitative change, especially at high pressure, where more pure phases at high pressure are expected to be produced, widening the gap of immiscibility of the join.

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IMPLICATIONS

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526 In this work, we compute for the first time the volume behaviour of this solid solution 527 using an atomistic approach based on first principle computational methods. We demonstrate that this approach predicts accurately the volume and compressional 528 529 behaviour of the solid solution and compares well with macroscopic experimentally 530 derived parameters. Future work will apply first-principle modelling to more complex 531 solid solutions of white micas, to make possible the study of the behaviour of these 532 material in the Earth's interior with high precision. Results will include the magnitudes 533 of solid solutions, structure, and elasticity as a function of pressure at the atomic level. 534 We further show that macroscopic parameters derived from first principles can be 535 combined with thermodynamic data in a semi-empirical way to better understand the volume behaviour of the solid solution with pressure. This will allow the precise 536 537 determination of the solvus at high pressure with special applications to subduction 538 zone metamorphism.

In this work, the evolution of the solvus has been presented only at one temperature, because the dependence with temperature is obtained from experimental data. However, with less complex systems, where the free energy can be easily calculated, the solvus will be able to be calculated, which opens an important computational field to the study of solid solutions.

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556 **References**

557

Angel R.J. (2000) Equations of state. In R.M. Hazen and R.T. Downs (Eds.), Highpressure and high-temperature crystal chemistry. Reviews in Mineralogy and
Geochemistry, 41, 35-60.

561

562 Angel, R.J. (2001), http://www.geol.vt.edu/rja/

563

Artacho, E., Sánchez-Portal, D., Ordejón, P., García, A., and Soler, J.M. (1999) Linearscaling ab-initio calculations for large and complex systems. Physica Status Solidi (b),
215, 809–817.

567

568 Birch, F. (1947) Finite elastic strain of cubic crystal. Physical Review, 71, 809 – 824.

569

570 Blencoe, J.G., (1977), Molal volumes of synthetic paragonite-muscovite micas. 571 American Mineralogist, 62, 1200-1215.

572

Blencoe, J.G., Guidotti, C.V., and Sassi, F.P (1994) The paragonite-muscovite solvus:II.
Numerical geothermometers for natural, quasibinary paragonite-muscovite pairs.
Geochimica et Cosmochimica Acta, 58, 2277-2288.

576

- 577 Brigatti, M.F., Frigieri, P., and Poppi, L. (1998) Crystal chemistry of Mg-, Fe-bearing 578 muscovites-2*M*₁. American Mineralogist, 83, 775–785.
- 579
- 580 Burnham, C.W. and Radoslovich, E.W. (1964) Crystal structure of coexisting 581 muscovite and paragonite. Carnegie Institute of Washington Year Books 63, 232-236.

582

583 Butler, I.S. and Frost, R.L. (2006) An overview of the high-pressure vibrational spectra 584 of clays and related minerals. Applied Spectroscopy Reviews, 41, 449-471.

This is a preprint, the final version is subject to change, of the American Mineralogist (MSA) Cite as Authors (Year) Title. American Mineralogist, in press. (DOI will not work until issue is live.) DOI: http://dx.doi.org/10.2138/am-2016-5407 Catti, M., Ferraris, G., and Ivaldi, G. (1989) Thermal strain analysis in the crystal 586 structure of muscovite $2M_1$ at 700 °C. European Journal of Mineralogy, 1, 625-632. 587 588 589 Catti, M, Ferraris, G., Hull, S., and Pavese, A. (1994) Powder neutron diffraction study of 2M₁ muscovite at room pressure and at 2 GPa. European Journal of Mineralogy, 6, 590 591 171-178. 592 593 Ceperley, D.M., and Alder, B.J. (1980) Ground state of the electron gas by a stochastic 594 method. Physical Review Letters, 45, 566-569. 595 Chatterjee, N.D., (1974) X-ray powder pattern and molar volume of synthetic 2M-596 597 paragonite refinement. Contributions to Mineralogy and Petrology, 43, 25-28. 598 Chatteriee, N.D., and Flux, S. (1986) Thermodynamic mixing properties of muscovite-599 600 paragonite crystalline solutions at high temperatures and pressures, and their geological 601 applications. Journal of Petrology, 27, 677-693. 602 603 Chatterjee, N.D., and Froese, E. (1975) Themodynamic study of pseudobinary join 604 muscovite-paragonite in system KAlSi₃O₈-NaAlSi₃O₈-Al₂O₃-SiO₂-H₂O. American 605 Mineralogist, 60, 985-993. 606 Chatterjee, N.D. and Johannes, W. (1974) Thermal stability and standard 607 608 thermodynamics properties of synthetic $2M_1$ Muscovite, KAl₂[AlSi₃O₁₀(OH)₂]. 609 Contributions to Mineralogy and Petrology, 48, 80-114. 610 Comodi, P., and Zanazzi, P.F. (1995) High-pressure structural study of muscovite. 611 612 Physics and Chemistry of Minerals, 22, 170–177. 613 614 (1997) Pressure dependence of structural parameters of paragonite. Physics and Chemistry of Minerals, 24, 274-280. 615 616

Comodi, P., Gatta, G.D., Zanazzi, P.F., Levy, D., and Crichton, W. (2002) Thermal 617 equations of state of dioctahedral micas on the join muscovite-paragonite. Physics and 618 Chemistry of Minerals, 29, 538-544. 619 620 Curetti, N., Levy, D., Pavese, A., and Ivaldi, G. (2006) Elastic properties and stability of 621 coexisting 3T and $2M_1$ phengite polytipes. Physics and Chemistry of Minerals, 32, 670– 622 678. 623 624 625 Eugster, H.P., Waldbaum, D.R., Thompson, J.B., Bence, A.E., and Albee, A.L. (1972) The two-Phase region and excess mixing properties of paragonite-muscovite crystalline 626 627 solutions. Journal of Petrology, 13, 147-179. 628 Fabrichnaya, O., Saxena, S.K., Richet, P., and Westrum, E.F. (Eds.) (2004) 629 630 Thermodynamics data, models, and phase diagrams in multicomponets oxide systems. 631 Springer-Verlag, Berlin. 632 Faust, J., and Knittle, E. (1994) The equation of state, amorphization and high pressure 633 634 phase diagram of muscovite. Journal of Geophysical Research, 99, 19785–19792. 635 Flux, S., and Chatterjee, N.D. (1986) Experimental reversal of the Na-K exchange 636 reaction between muscovite-paragonite crystalline solutions and a 2 molal aqueous 637 638 (Na,K)Cl fluid. Journal of Petrology, 27, 665-676. 639 Ganguly, J. (2008) Thermodynamics in Earth and Planetary Sciences. Heidelberg, 640 641 Springer. 642 Gatta, G.D., Rotiroti, N., Lotti, P., Pavese, A., and Curetti, N. (2010) Structural 643 644 evolution of a $2M_1$ phengite mica up to 11 Gpa: in situ single crystal X-ray diffraction study. Physics and Chemistry of Minerals, 37, 581-591. 645 646 647 Geiger, C. A. (2008) Silicate garnet: A micro to macroscopic (re)view. American 648 Mineralogist, 93, 360 - 372.

This is a preprint, the final version is subject to change, of the American Mineralogist (MSA) Cite as Authors (Year) Title. American Mineralogist, in press. (DOI will not work until issue is live.) DOI: http://dx.doi.org/10.2138/am-2016-5407 649 650 Goerigk, L. and Grimme, S. (2011) A thorough benchmark of density functional methods for general main group thermochemistry, kinetics, and noncovalent 651 652 interactions. Physics Chemical Chemical Physics, 13, 6670-6688. 653 654 Guggenheim, S., Chang, Y.-H., and Koster van Groos, A.F. (1987) Muscovite dehydroxylation: High-temperature studies. American Mineralogist, 72, 537-550. 655 656 Guidotti C.V., Mazzoli C., Sassi F.P., and Blencoe, J.G. (1992) Compositional controls 657 on the cell dimensions of $2M_1$ muscovite and paragonite. European Journal of 658 Mineralogy, 4, 283-297. 659 660 Guidotti, C.V., Sassi F.P., Blencoe, J.G., and Selverstone, J. (1994) The paragonite-661 muscovite solvus: I. P-T-X limits derived from the Na-K compositions of natural, 662 quasibinary paragonite-muscovite pairs. Geochimica et Cosmochimica Acta, 58, 2269-663 664 2275. 665 Guidotti, C.V., Sassi, F.P., Comodi, P., Zazzi, P.F. and Blencoe, G. (2000) The 666 contrasting responses of muscovite and paragonite to increasing pressure: petrological 667 668 implications. The Canadian Mineralogist, 38, 707 – 712. 669 670 Hazen, R.M. and Finger, L.W. (1977) Compressibility and structure of Angra dos Reiss fassaite. Carnegie Institution of Washington Year Book, 76, 512-515. 671 672 Hashin, Z., and Shtrikman, S. (1962a) On some variational principles in anisotropic and 673 non-homogeneus elasticity. Journal of Mechanical Physics of Solids, 10, 335–342. 674 675 676 - (1962b) A variational approach to the theory of the elastic behaviour of polycrystals. Journal of Mechanical Physics of Solids, 10, 343 – 352. 677 678

Hill, R. (1952) The elastic behaviour of a crystalline aggregate. Proceedings of thePhysical Society of London A, 65, 349-354.

Hernández-Haro, N., Ortega-Castro, J., Pérez del Valle, C., Muñoz-Santiburcio, D.,
Sainz-Díaz, C.I., and Hernández-Laguna, A. (2013) Computational study of the elastic

behaviour of the $2M_1$ muscovite-paragonite series. American Mineralogist, 86, 651 - 665.

686

Hohenberg, P., and Kohn, W. (1964) Inhomogeneous electron gas. Physical Review B,
136, 864–871.

689

Holland, T.J.B., and Powell, R. (1998) An internally consistent thermodynamic data set
for phases of petrological interest. Journal of Metamorphic Geology, 16, 309 – 343.

692

(2011) An improved and extended internally consistent thermodynamic dataset
for phases of petrological interest, involving a new equation of state for solids.
Journal of Metamorphic Geology, 29, 333-383.

696

Hugh-Jones, D.A., and Angel, R.J. (1994) A compressional study of MgSiO₃
orthoenstatite up to 8.5 GPa. American Mineralogist, 79, 405-410.

699

Kohn, W., and Sham, L.J. (1965) Self-consistent equations including exchange andcorrelation effects. Physical Review, 140, A1133–A1138.

702

Kurht S., Perdew, J.P. and Blaha, P. (1999) Molecular and solid state tests of density
functionals approximations: LSD, GGAs, and Meta-GGAs. International Journal of
Quantum Chemistry, 75, 889-909.

706

Levien, L., and Prewitt, C.T. (1981) High-pressure study of diopside, American
Mineralogist, 66, 315-323.

709

Lin, C.Y., and Bailey, S.W. (1984) The crystal Structure of Paragonite $2M_1$. American Mineralogist, 69, 122–127.

- 713 Mookherjee, M., Redfern, S.A.T., and Zhang, M. (2001) Thermal response of structure
- and hydroxyl ion of phengite- $2M_1$: an in situ neutron diffraction and FTIR study.
- European Journal of Mineralogy, 13, 545 555.

716

Mookherjee, M., and Redfern, S.A.T. (2002) A high-temperature Fourier transform infrared study of the interlayer and Si-O-stretching region in phengite- $2M_1$. Clay Minerals, 37, 323-336.

720

- 721 Muñoz-Santiburcio, D., Ortega-Castro, J., Huertas, F.J., and Hernández-Laguna, A.
- (2011) Influence of the exchangeable cation on the adsorption of 2-nitro-1-propanol on
 smectite surface models. Chemical Physics Letters, 515, 49–55.

724

- 725 Ortega-Castro, J., Hernández-Haro, N., Hernández-Laguna, A., and Sainz-Díaz, C.I.
- 726 (2008) DFT calculation of crystallographic properties dioctahedral 2:1 phyllosilicates.
- 727 Clay Minerals, 43, 351–361.

728

Ortega-Castro, J., Hernández-Haro, N., Muñoz-Santiburcio, D., Hernández-Laguna, A.,
and Sainz-Díaz, C.I. (2009) Crystal structure and hydroxyl group vibrational
frequencies of phyllosilicates by DFT methods. Journal of Molecular Structure
THEOCHEM, 912, 82–87, DOI:10.1016/j.theochem.2009.02.013.

733

Ortega-Castro, J., Hernández-Haro, N., Timón, V., Sainz-Díaz, C.I. and HernándezLaguna, A. (2010) High-pressure behaviour of 2*M*₁ Muscovite. American Mineralogist,
95, 249–259. DOI: 10.2138/am.2010.3035.

737

Perdew, J.P., Burke, K., and Ernzerhof, M. (1996) Generalized gradient approximation
made simple. Physical Review Letters, 77, 3865–3868.

740

Prencipe, M., Scanavino, I, Nestola, F., Merlini, M. Civalleri, B., Bruno, M., and
Dovesi, R. (2011) High-pressure thermos-elastic properties of beryl (Al₄Be₆Si₁₂O₃₆)
from *ab initio* calculations, and observations about the source of thermal expansion.
Physics and Chemistry of Minerals, 38, 223-239.

- Ralph, R.L., and Ghose, S. (1980) Enstatite, Mg₂Si₂O₆: Compressibility and crystal
- 747 structure at 21 kbar (abs.). EoS, 61, 409.

748

Redlich, O., and Kister, T. (1948) Algebraic representation of thermodynamic
properties and the classification of solutions. Industrial and Engineering Chemistry, 40,
345-348.

752

753 Rothbauer, R. (1971) Untersuchung eines $2M_1$ -muskovits mit neutronenstrahlen, Neues 754 Jahrbuch für Mineralogie Monatshefte, 1971, 143-154.

755

Ruox, J., and Hovis, G.L. (1996) Thermodynamic mixing model for muscoviteparagonite solutions based on solutions calorimetric and phase equilibrium data. Journal
of Petrology, 57, 1241-1254.

759

- Sánchez-Portal, D., Ordejón, P., Artacho, E., and Soler, J.M. (1997) Density-functional
- method for very large systems with LCAO basis sets. International Journal of QuantumChemistry, 65, 453–461.

763

Saxena, S.K., Chatterjee, N., Fei, Y., and Shen, G. (1993) Thermodynamic data on
Oxides and silicates. Springer-Verlag, Berlin.

766

Sekine, T., Rubin, A.M., and Ahrens, T.J. (1991) Shock wave equation of state ofmuscovite. Journal of Geophysical Research, 96, 19675–19680.

769

- 770 Smyth, J.R., Jacobsen, S.D., Swope, R.J., Angel, R.J., Arlt, T., Domanik, K., and 771 Holloway, R. (2000) Crystal structures and compressibilities of synthetic $2M_1$ and 3T
- phengite micas. European Journal of Mineralogy, 12, 955-963.

773

- Soler, J.M., Artacho, E., Gale, J.D., García, A., Junquera, J., Ordejón, P., and Sánchez-
- Portal, D. (2002) The SIESTA method for ab-initio order-N materials simulation.
- Journal of Physics: Condensed Matter, 14, 2745-2779.

Takahashi, T., and Liu, L. (1970) Compression of ferromagnesian garnets and the effect 778 of solid solutions on the bulk modulus. Journal of Geophysical Research, 75, 5757-779 780 5766. 781 Teich-McGoldrick, S.L., Greathouse, J.A., and Cygan R.T. (2012) Molecular dynamic 782 simulations of structural and mechanical properties of muscovite: pressure and 783 temperature effects. Journal of Physical Chemistry C, 116, 15099-15107. 784 785 Troullier, N., and Martins, J.L. (1991) Efficient pseudopotentials for plane-wave 786 787 calculations. Physical Review B, 43, 1993-2006. 788 Tunega, D., Bučo, T., and Zaoui, A. (2012) Assessment of ten DFT methods in 789 790 predicting structures of sheet silicates: Importance of dispersion corrections. Journal of 791 Chemical Physics, 137, 114105. 792 Vaughan, M.T., and Guggenheim, S. (1986) Elasticity of muscovite and its relationship 793 794 to crystal structure. Journal of Geophysics Research, 91, 4657-4664. 795 796 Waldbaum, D.R., and Thomson, J.B. (Jr.) (1968) Mixing properties of sanidine crystalline solutions. II Calculation based volume data. American Mineralogist, 53, 797 798 2000-2017. 799

White, C.E., Provis, J.L., Riley, D.P., Kearley, G.J., and van Deventer, J.S.J. (2009)
What is the structure of Kaolinite? Reconciling Theory and Experiment. Journal of
Physical Chemistry B, 113, 6756-6765.

803

Williams, Q., Knittle, E., Scott, H.P., and Liu, Z. (2012) The high-pressure behavior of
micas: Vibrational spectra of muscovite, biotite, and phlogopite to 30 GPa. American
Mineralogist, 97, 241–252.

808 Captions for the figures

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Figure 1.- Crystal structure of the $X_{Na} = 0.5$ model of the Ms – Pg series. Sky- and navy-blue polyhedrons represent the Al³⁺ and Si⁴⁺ cation polyhedrons, respectively; the O, H, Na, and K atoms are represented by red, white, yellow, and purple spheres, respectively.

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Figure 2.- (a) and (c) variation of *a* and *b* [, (b) and (d) $0.5csin\beta$ (Å) (0.5csinbe or 0.5csb in the insets) and *volume* (*V*, Å³) of Ms and Pg, respectively, as a function of pressure. In the insets: *j*CZ1995 means values of parameter *j* from Comodi and Zanazzi (1995); *j*298Co2002 means values at 298K from Comodi et al. (2002); *j*Gat2010 means from Gatta et al. (2010); and *j*20MRZ2001 means values at 20 °C from Mookherjee et al. (2001); where *j* is either *a*, *b*, 0.5csinbe/0.5csb or volume (Vol).

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Figure 3.- (a) tetrahedral rotation (α , in degrees) in the Na⁺ and K⁺ cavities at 0, 3, and 6 GPa; (b) incompressibility modulus×(-1) (GPa) of tetrahedral rotation, α , in the Na⁺ and K⁺ cavities; (c) interlayer thickness (Å) at 0, 3, and 6 GPa; (d) Interlayer thickness/interlayer thickness modulus (Å/GPa); and (e) volume/bulk modulus (Å³/GPa) as a function of Na⁺ content in the series. Insets: least-square fitting equations, correlation coefficients (R or R²), standard deviations (SD), and Margules' coefficients (W_i).

829

Figure 4.- Excess volume V^{ex} (J/bar) at different pressures. (a) 0, and (b) 60 kbar as a function of X_K . Values in the insets are from interpolated quadratic polynomial, and A_V , B_V and C_V are the coefficients of Eq. (7) for the V^{ex} , at 298 K. Volumes have been corrected as a function of temperature (Holland and Powell, 1998).

Figure 5.- Excess free energy G^{ex} (KJ/mol) at different pressures. Coefficients of

enthalpy and entropy of Eq. (7) was taken from the Model A of Roux and Hovis (1996).

837 Model A is calculated at 693 K.

838

839 Captions for the supplementary material figures

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Figure S1.- (a) variation of *a* and *b*, (b) $0.5csin\beta$ (Å) (0.5csinbe or 0.5csb in the insets) and *volume* (*V* or *Vol*, Å³), of $X_{Na} = 0.25$ (Na-Ms), respectively, as a function of pressure. In the insets: *j*CZ1995 means values of parameter *j* from Comodi and Zanazzi (1995); where *j* is either *a*, *b*, 0.5csinbe/0.5csb or *volume*.

845

Figure S2.- Variation of Si/^{IV}Al – $O_{b(basal)/a(apical)}$ average distances (Å) as a function of pressure (GPa). Ms/Pg/NaMsT*i*OCZ95/97 mean tetrahedral cation O bond (T*i*-O) distances from Ms, Pg or Na-rich Ms from Comodi and Zanazzi (1995) or (1997),

849 where i means site 1 or 2.

- **Figure S3.** Excess volume V^{ex} (J/bar) at different pressures as a function of X_K . (a) 7.5
- kbar, and (b) 15 kbar. Values in the insets are from interpolated quadratic polynomial,
- and A_V , B_V and C_V are the coefficients of Eq. (7) for the V^{ex} , at 298 K. Volumes have
- been corrected as a function of temperature (Holland and Powell, 1998).

Table 1.- Bulk modulus (B_0 in GPa) and their derivatives with respect to pressure [B_0 ', B_0 '' (GPa⁻¹) first and second derivative of B_0 with respect to pressure in brackets] values for Ms and Pg from different experimental and computational studies.

Ref.	Ms	Pg
Vaughan et al. (1986)	58.2 ^a	
Sekine et al. (1991)	52 ^b (3.2)	
Faust and Knittle(1994)	61.4 ^{cd} (6.9)	
Catti et al. (1994)	$56^{e}(4)$	
Comodi & Zanazzi (1995, 1997)	$56.0^{c,f}$, $60.0^{c,f,g}$	65.5 ^{c,f,h}
Smyth et al., (2000) Phengite	57 ^c (9.2)	
Comodi et al. (2002)	57.0 ^{c,d}	59.9 ^{c,d}
Curetti et al. (2006), Phengite	57.3 ^{c,d} (6.97)	
Gatta et al. (2010)	57.3 ^{c,f} (6.97)	
Phengite		
Ortega-Castro et al. (2010)	60.1 ⁱ (7.3) , 55.7 ^j	
Holland & Powell (2011)	49.0 (4.15,-0.085)	51.5 (6.51, -0.126)
Teich-McGoldrick et al. (2012)	59.8 ^k	
Hernández-Haro et al. (2013)	$68.4^{\rm l}, 64.0^{\rm j}$	68.8 ¹ , 65.5 ^j

^a Brillouin scattering.

^b Shock wave.

^c X-Ray diffraction.

^d Neutron diffraction.

^e Powdered sample.

^f Monocrystal sample.

^g Na-rich Ms, Na 0.37 per formula unit (pfu).

^h Na 0.88 pfu.

ⁱ DFT.

^j An estimate of the bulk modulus at 298 K from Comodi et al. (2002).

^k Molecular dynamics simulations.

¹ DFT and Hashin-Shtrikman (1962)-Hill (1952) approaches.

Table 2.	Calculated and experimental	structural parameters of	of Ms-Pg series	(lengths in Å	A, angles in degrees,	and volumes in Å	X^{3}). $X_{Na} = Na$	1 ^{+/} (Na ^{+.}	$+K^+$)
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Parameter ^a	Experimental ^b	Experimental ^c	X _{Na} =0.0	0.25	0.5	0.75	1.0
a	5.174-5.226	5.128-5.135	5.187	5.176	5.160	5.152	5.123
b	8.976-9.074	8.898-8.906	9.006	8.994	8.975	8.932	8.895
С	19.875-20.097	19.287-19.384	20.148	19.917	19.653	19.378	19.113
β	95.59-95.84	94.35-94.6	95.44	95.55	95.55	94.08	94.02
Volume	926-945.4	877.51-883.6	936.999	922.829	906.007	889.581	868.767
T-O	1.64	1.653	1.651, 1.757	1.649,1.758	1.675,1.761	1.653, 1.763	1.655, 1.766
M-O	1.927-1.94	1.908-2.221	1.934	1.922	1.931	1.920	1.926
М-ОН	1.911	1.896	1.918	1.918	1.912	1.911	1.903
KO _{outer}	3.272-3.373	-	3.427	3.402	3.384	3.372	-
KO _{inner}	2.832-2.934	-	2.759	2.721	2.680	2.624	-
Δ_{K}	0.426-0.509	-	0.668	0.681	0.704	0.748	-

Na…O _{outer}	-	3.370-3.374	-	3.438	3.407	3.352	3.357
Na…O _{inner}	-	2.624-2.642	-	2.680	2.618	2.514	2.474
$\Delta_{ m Na}$	-	0.746-0.732	-	0.758	0.789	0.838	0.883
Tetrahedral rotation α_K , $\alpha_{Na}(^{o})$	10.3 ^d -11.3	16,16.2 - 19	14.6	15.1, 15.8	15.5 , 16.7	16.5, 17.5	18.8
V(T) Si,Al (Å ³)		2.32	2.273, 2.774	2.274, 2.784	2.282, 2.794	2.283, 2.810	2.298, 2.824
V(Oc) Al (Å ³)		9.10	9.386	9.361	9.345	9.339	9.283
Tetrahedral-sheet thickness (Å)	2.262 ^d	2.243	2.277	2.265	2.271	2.248	2.237
Octahedral-sheet thickness (\AA)	2.083 ^d	2.078-2.085	2.093	2.090	2.089	2.115	2.144
Interlayer thickness (Å)	3.375 ^d	3.053-3.090	3.361	3.273	3.128	3.021	2.867

^a V(T) means volume of the tetrahedral sheet, V(Oc) is the volume of octahedral sheet; $\tau = O_{apical}$ bond angles ; Tetrahedral rotation α (°) = $\frac{1}{2} \sum_{i=1,6} [120^\circ - (O_b - O_b)_i]/6$; Tetrahedral-sheet thickness (Å) = [(mean z coordinate of the "upper" basal O) – (mean z coordinate of the "lower" apical O)]*csinβ*; octahedral-sheet thickness (Å) = [(mean z coordinate of "upper" O or OH in the octahedral sheet) – (mean z coordinate of the "lower" of O or OH in the octahedral sheet) – (mean z coordinate of "lower" O or OH in the octahedral sheet)]*csinβ*.

^b Burnham and Radoslovish (1964) $K_{0.66}Na_{0.34}Al_2(AlSi_3)O_{10}(OH)_2$; Rothbauer (1971) $K_{0.85}Na_{0.1}(Al_{1.81}Fe^{2^+}_{0.14}Mg_{0.12})(Al_{0.9}Si_{3.1})O_{9.8}(OH)_2$; Guggenheim et al. (1987) $K_{0.93}Na_{0.08}(Al_{1.83}Fe_{0.16}Ma_{0.01})(Al_{0.90}Si_{3.10})O_{10}(OH)_{1.83}F_{0.17}$ and $K_{1.00}Na_{0.03}Ca_{0.01}(Al_{1.93}Fe_{0.01}Mn_{0.01})(Al_{0.91}Si_{3.09})O_{10}(OH)_{1.88}F_{0.12}$; Catti et al. (1989) $K_{0.86}Na_{0.11}(Al_{1.93}Fe_{0.07}Mg_{0.02})(Al_{0.92}Si_{3.08})O_{10}(OH)_2$; Catti et al. (1994) $K_{0.90}Na_{0.07}(Al_{1.63}Fe_{0.23}Mg_{0.16}Ti_{0.03})(Al_{0.80}Si_{3.20})O_{10}(OH)_2$; Guidotti et al. (1992); Brigatti et al. (1998) Different compositions; Mookherjee and Redfern (2002) $K_{0.95}Na_{0.05}(Al_{0.76}Fe_{0.14}Mg_{0.10})(Al_{0.75}Si_{3.25})O_{10}(OH_{1.96}F_{0.04})$

^c Comodi and Zanazzi (1997) Na_{0.88}K_{0.10}Ca_{0.01}Ba_{0.01}(Al_{1.97}Ti_{0.007}Fe_{0.01}Mn_{0.006})Si_{3.01}Al_{0.99}O₁₀(OH)₂: Guidotti et al. (1992); Lin & Baley, (1984) K_{0.042}Na_{0.916}Ca_{0.018}□_{0.024}(Al_{1.990}Fe_{0.028}Mg_{0.013}Ti_{0.003})(Al_{1.061}Si_{2.939})O₁₀(OH)₂: Guidotti et al. (1992); Lin & Baley, (1984) K_{0.042}Na_{0.916}Ca_{0.018}□_{0.024}(Al_{1.990}Fe_{0.028}Mg_{0.013}Ti_{0.003})(Al_{1.061}Si_{2.939})O₁₀(OH)₂: Guidotti et al. (1992); Lin & Baley, (1984) K_{0.042}Na_{0.916}Ca_{0.018}□_{0.024}(Al_{1.990}Fe_{0.028}Mg_{0.013}Ti_{0.003})(Al_{1.061}Si_{2.939})O₁₀(OH)₂: Guidotti et al. (1992); Lin & Baley, (1984) K_{0.042}Na_{0.916}Ca_{0.018}□_{0.024}(Al_{1.990}Fe_{0.028}Mg_{0.013}Ti_{0.003})(Al_{1.061}Si_{2.939})O₁₀(OH)₂: Guidotti et al. (1992); Lin & Baley, (1984) K_{0.042}Na_{0.916}Ca_{0.018}□_{0.024}(Al_{1.990}Fe_{0.028}Mg_{0.013}Ti_{0.003})(Al_{1.061}Si_{2.939})O₁₀(OH)₂: Guidotti et al. (1992); Lin & Baley, (1984) K_{0.042}Na_{0.916}Ca_{0.018}□_{0.024}(Al_{1.990}Fe_{0.028}Mg_{0.013}Ti_{0.003})(Al_{1.061}Si_{2.939})O₁₀(OH)₂: Guidotti et al. (1992); Lin & Baley, (1984) K_{0.042}Na_{0.916}Ca_{0.018}□_{0.024}(Al_{1.990}Fe_{0.028}Mg_{0.013}Ti_{0.003})(Al_{1.061}Si_{2.939})O₁₀(OH)₂: Guidotti et al. (1992); Lin & Baley, (1984) K_{0.042}Na_{0.916}Ca_{0.018}□_{0.024}(Al_{1.990}Fe_{0.028}Mg_{0.013}Ti_{0.003})(Al_{1.061}Si_{2.939})O₁₀(OH)₂: Guidotti et al. (1992); Lin & Baley, (1984) K_{0.042}Na_{0.916}Ca_{0.018}□_{0.024}(Al_{1.990}Fe_{0.028}Mg_{0.013}Ti_{0.003})(Al_{1.061}Si_{2.939})O₁₀(OH)₂: Guidotti et al. (1992); Lin & Baley, (1984) K_{0.042}Na_{0.916}Ca_{0.018}□_{0.024}(Al_{1.990}Fe_{0.028}Mg_{0.013}Ti_{0.003})(Al_{1.061}Si_{0.018}Ka_{0.018})(Al_{0.018}Ka_{0.018}

 $^{d} \text{ Comodi and Zanazzi (1995) } K_{0.9} Na_{0.07} Ba_{0.01} \square_{0.02} (Al_{1.84} Ti_{0.04} Fe_{0.07} Mg_{0.04}) (Al_{0.98} Si_{3.02}) O_{10} (OH)_2$

		<i>x</i> = 0.0	<i>x</i> = 0.25	<i>x</i> = 0.5	<i>x</i> = 0.75	<i>x</i> = 1.0
B _a	(0 GPa)	490.6	444.9	369.8	324.6	284.4
	(6 GPa)	526.0	452.7	386.0	356.4	319.84
	(exp)	337.8 ^a , 400 ^b		342.5 [°]		285.7 ^d
B_b	(0 GPa)	413.7 ^e	365.8	322.1	276.5	246.3
	(6 GPa)	437.7	385.0	349.1	304.7	284.1
	(exp)	295.0 ^a , 342 ^b	287.	3°		277.8 ^d
$B_{0.5csin\beta}$	(0 GPa)	79.9	84.5	95.9	115.8	128.4
	(6 GPa)	111.7	122.9	137.9	151.2	166.8
	(exp)	88.2 ^a ,86.9 ^{b,f}	98.9	c,f		120.4 ^{d,f}
B_0	(0 GPa)	60.0	60.3	62.7	65.1	65.2
	(0 GPa, 298K)	55.6	56.2	58.9	61.6	62.0
В	(6 GPa)	93.8	94.4	97.1	97.5	98.2
	(6 GPa, 298K)	103.12	105.3	106.1	101.1	103.6
B_0	(exp)	56.0 ^{a,d}	60.0	c,d		65.0 ^d

Table 3.- Bulk moduli (B_0 in GPa) and cell parameter moduli (B_a , B_b , and $B_{0.5csin\beta}$ in GPa) at 0 and 6 Gpa of the Ms-Pg series. Temperature correction for bulk moduli comes from Comodi et al. (2002), considering, in an initial approach, linearity from 0K.

 a From $K_{0.9}Na_{0.05}Ba_{0.01}$ $_{0.02}(Al_{1.84}Ti_{0.04}Fe_{0.07}Mg_{0.04})(Si_{3.02}Al_{0.98})O_{10}(OH)_2$ Comodi and Zanazzi (1995)

^b From $K_{0.98}Na_{0.02}(Al_{1.55}Mg_{0.24}Fe_{0.21}Ti_{0.02})(Si_{3.38}Al_{0.62})O_{10}(OH)_2$ Curetti et al. (2006).

^c From K_{0.6}Na_{0.37 0.03}(Al_{1.84}Ti_{0.02}Fe_{0.10}Mg_{0.06})(Si_{3.03}Al_{0.97})O₁₀(OH)₂ Comodi and Zanazzi (1995)

 d From $K_{0.10}Na_{0.88}Ca_{0.01}Ba_{0.01}(Al_{1.97}Ti_{0.007}Fe_{0.01}Mn_{0.002}Mg_{0.006})(Si_{3.01}Al_{0.99})O_{10}(OH)_{2}$ Comodi and Zanazzi (1997)

^e Birch-Murnagham second-order equation

 $^{\rm f}c$ axe.

Group	Ms < <i>i</i> > _{0,6}	Pg < <i>i</i> > _{0,6}	Mod. Ms	Mod. Pg
	0 GPa 6 GPa	0 GPa 6 GPa		
Si-O _b	1.653, 1.640	1.659, 1.651	775.2	1802.8
Si-O _a	1.646, 1.637	1.647, 1.639	1173.7	1000.0
^{IV} Al-O _b	1.754, 1.737	1.768, 1.754	520.8	751.9
^{IV} Al-O _a	1.763, 1.750	1.760, 1.748	781.2	925.9
T – O exp	1.646, 1.64 ^a	1.652, 1.66 ^b		
KO	2.760, 2.611	2.680, 2.551 ^c	74.8	88.3 ^c
KO	3.442, 3.356	3.310, 3.256 ^c	156.2	290.0 ^c
Na…O	2.706, 2.535 ^d	2.535, 2.370	64 ^d	64.7
Na…O _{outer}	3.395, 3.282 ^d	3.412, 3.370	109 ^d	336.7
V(Td)Si ^e	2.273, 2.221	2.297, 2.265	262	463
V(T)Si V(T)Si exp	2.28, 2.27 ^a	2.31, 2.35 ^b		
V(T)Al ^e	2.774, 2.697	2.824, 2.760	210.1	267.4
V(Oc) ^e	9.388, 9.212	9.283, 9.065	318	224
V(Oc) exp	9.31,9.04 ^a	9.10, 8.6 ^b	105 ^f	65.9 ^f
$\alpha_{_{\rm K}}$	14.6, 15.2	16.5, 18.4 ^c	- 92	- 40 ^c
$\alpha_{_{Na}}$	15.8, 17.6 ^d	18.9, 21.3	- 47 ^d	- 33.5
T-thick ^e	2.277, 2.262 ^a	2.237, 2.226 ^b	917	1828.1
T-thick exp	2.262, 2.237	2.243, 2.25	249	
Oc-thick ^e	2.093, 2.069	2.144, 2.107	483	248.7
Oc-thick exp	2.083, 2.047 ^a	2.085, 1.98 ^b	159 ^f	78^{f}
Interlayer-thick ^e	3.361, 2.954	2.867, 2.632	32	51.1
Interlayer-thick exp	3.375, 3.128 ^a	3.090, 2.81 ^b	37.6 ^f	44.7 ^f

Table 4.- Values of average ($\langle i \rangle$) bond distances (Å), volumes (Å³) tetrahedral distortion angle (α^{o}), and atomic groups at a pressure of 0 and 6 GPa, incompressibility moduli of the average variable ($\langle i \rangle$) (GPa) of Ms (Mod. Ms) and Pg (Mod. Pg); a and b subindexes under O mean apical and basal oxygens, respectively.

^a Ms at 0.0001 and 2.8 GPa, Comodi and Zanazzi (1995)

^b Pg at 0.0001 and 4.05 GPa, Comodi and Zanazzi (1997)

^c Values $X_{\text{Na}} = 0.75$

^d Values $X_{\text{Na}} = 0.25$

 e V(T)Si/Al = Volume of the Si/Al tetrahedra, V(Oc) = Volume of the octahedral, T/Oc-thick = Tetrahedral/Octahedral-sheet thickness, Interlayer-thick = Interlayer thickness

^f Incompressibility modulis estimated from the values of Comodi and Zanazzi (1995) and (1997)

Table 5.- A_V , B_V and C_V coefficients (J/bar) of Eq. (7) at different pressures (kbar) and temperatures (K).

Mod	P (kbar)	T (K)	A_V	B_V	C_V
\mathbf{B}^{a}	8	677-943	0.2760	-0.0372	-
C&Fr ^b	1-7	848-1008	0.1044	-0.5606	-
$C\&F^{c}$	5-7	723-893	0.2359	-0.084	-
A $(R\&H)^d$	1.0-9.0	853-973	0.350	0.037	-
This work ^e	0.0	298	0.190	0.070	0.315
"	0.0	693	0.196	0.071	0.320
"	0.0	793	0.198	0.072	0.322
"	0.0	893	0.200	0.072	0.323
"	0.0	993	0.202	0.073	0.325
"	2.5	298	0.181	0.050	0.390
"	5.0	298	0.192	0.056	0.390
"	7.5	298	0.188	0.047	0.332
"	15.0	298	0.217	0.119	0.139
"	30.0	298	0.226	0.092	0.161
"	60.0	298	0.213	0.065	0.079

^a Blencoe (1977), from his Margules' coefficients.

^b Chatterjee and Froese (1975), from their Margules' coefficients.

^c Chatterjee and Flux (1986).

^d Model A of Roux and Hovis (1996).

^e Volumes have been corrected as a function of temperature (Holland and Powell, 1998).



Figure 1



Figure 2a















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