## Application of mineral equilibria to estimate fugacities of H<sub>2</sub>O, H<sub>2</sub>, and O<sub>2</sub> in mantle xenoliths from the southwestern U.S.A.

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## ABSTRACT

Small amounts of  $H_2O$ , on the order of tens to hundreds of parts per million, can significantly influence the physical properties of mantle rocks. Determining the  $H_2O$  contents of nominally anhydrous minerals (NAMs) is one relatively common technique that has been applied to estimate mantle  $H_2O$  contents. However, for many mantle NAMs, the relation between  $H_2O$  activity and  $H_2O$  content is not well known. Furthermore, certain mantle minerals may be prone to  $H_2O$  loss during emplacement on Earth's surface. The goal of this study is to apply mineral equilibria to estimate values of  $a_{H_2O}$  in rocks that originated below the Moho.

The chemical compositions of olivine + orthopyroxene + clinopyroxene + amphibole + spinel  $\pm$  garnet were used to estimate values of temperature (*T*), pressure (*P*),  $a_{H_2O}$ , hydrogen fugacity ( $f_{H_2}$ ), and oxygen fugacity ( $f_{O_2}$ ) in 11 amphibole-bearing mantle xenoliths from the southwestern U.S.A. Application of amphibole dehydration equilibria yields values of  $a_{H_2O}$  ranging from 0.05 to 0.26 for these 11 samples and the compositions of coexisting spinel + olivine + orthopyroxene yield  $\Delta \log f_{O_2}$  (FMQ) of -1 to +0.6. For nine of the samples, values of  $f_{H_2}$  were estimated using amphibole dehydrogenation equilibria, and these values of  $f_{H_2}$  ranged from 6 to 91 bars. Values of  $f_{H_2}$  and  $f_{O_2}$  were combined, using the relation  $2H_2O = 2H_2 + O_2$ , to estimate a second value of  $a_{H_2O}$  that ranged from 0.01 to 0.57 for these nine samples. Values of  $a_{H_2O}$ , estimated using these two methods on the same sample, generally agree to within 0.05. This agreement indicates that the amphibole in these samples has experienced little or no retrograde H-loss and that amphibole equilibria yields robust estimates of  $a_{H_2O}$  that, in these xenoliths, are generally <0.3, and are often 0.1 or less.

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