

1 **Revision 1**

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3 **Nitscheite,  $(\text{NH}_4)_2[(\text{UO}_2)_2(\text{SO}_4)_3(\text{H}_2\text{O})_2] \cdot 3\text{H}_2\text{O}$ , a new mineral with an unusual**  
4 **uranyl-sulfate sheet**

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17 **ABSTRACT**

18 Nitscheite (IMA2020-078),  $(\text{NH}_4)_2[(\text{UO}_2)_2(\text{SO}_4)_3(\text{H}_2\text{O})_2] \cdot 3\text{H}_2\text{O}$ , is a new mineral species from  
19 the Green Lizard mine, Red Canyon, San Juan County, Utah, U.S.A. It is a secondary phase  
20 found in association with chinleite-(Y), gypsum, pyrite, and Co-rich rietveldite. Nitscheite occurs  
21 in subparallel and divergent intergrowths of yellow prisms, up to about 0.3 mm in length.  
22 Crystals are elongated on [101] and exhibit the forms {100}, {010}, {001}, and {11-1}. The  
23 mineral is transparent with vitreous luster and very pale-yellow streak. It exhibits bright green

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24 fluorescence under a 405 nm laser. The Mohs hardness is ~2. The mineral has brittle tenacity,  
25 curved fracture, and one good cleavage on {010}. The measured density is 3.30(2) g·cm<sup>-3</sup>. The  
26 mineral is easily soluble in RT H<sub>2</sub>O. The mineral is optically biaxial (-),  $\alpha = 1.560(2)$ ,  $\beta =$   
27  $1.582(2)$ ,  $\gamma = 1.583(2)$  (white light);  $2V_{\text{meas}} = 17(1)^\circ$ ; no dispersion; orientation  $X = \mathbf{b}$ ,  $Z \approx [101]$ ;  
28 pleochroism  $X$  colourless,  $Y$  and  $Z$  yellow;  $X < Y \approx Z$ . Electron microprobe analysis provided the  
29 empirical formula  $(\text{NH}_4)_{1.99}\text{U}_{2.00}\text{S}_{3.00}\text{O}_{21}\text{H}_{10.01}$ . Nitscheite is monoclinic,  $P2_1/n$ ,  $a = 17.3982(4)$ ,  $b$   
30  $= 12.8552(3)$ ,  $c = 17.4054(12)$  Å,  $\beta = 96.649(7)^\circ$ ,  $V = 3866.7(3)$  Å<sup>3</sup>, and  $Z = 8$ . The structure ( $R_1$   
31  $= 0.0329$  for 4547  $I > 3\sigma I$  reflections) contains  $[(\text{UO}_2)_2(\text{SO}_4)_3(\text{H}_2\text{O})_2]^{2-}$  uranyl-sulfate sheets,  
32 which are unique among minerals, with NH<sub>4</sub> and H<sub>2</sub>O groups between the sheets.

33

34 *Keywords:* nitscheite; new mineral; uranyl sulfate sheet; crystal structure; Raman spectroscopy;  
35 Green Lizard mine, Red Canyon, Utah.

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## INTRODUCTION

38 Inactive uranium mines have proven to be fruitful underground “natural laboratories”. Since  
39 2012, the abandoned uranium mines in Red Canyon in southeastern Utah have provided more  
40 than 30 new uranyl minerals, predominantly uranyl sulfates, some of which possess bizarre  
41 structural features that have not been observed in laboratory synthetic experiments. The great  
42 diversity observed for uranyl sulfate minerals stems primarily from the large number of stable  
43 combinatorial linkages of uranyl pentagonal bipyramids and sulfate tetrahedra. The recent  
44 discoveries of novel structural topologies provide insight into the factors driving the  
45 crystallization of uranium minerals, in particular the relationship between mineral associations,  
46 local chemistry, and resulting structural topology. The formation of these mineral structures is

47 strongly affected by at least three parameters: pH, cation content, and water content (Plášil et al.  
48 2014). However, despite our growing knowledge of the crystal chemistry of uranyl sulfates, the  
49 cause of certain structural phenomena, manifested in particular in the minerals from Red Canyon,  
50 remains unknown. For instance, finite cluster topologies are relatively abundant among the  
51 sodium-uranyl-sulfate minerals, which occur commonly at the Blue Lizard mine, but do not  
52 occur in other mines nearby. Each new mineral has filled gaps in our understanding of how  
53 formation conditions influence the observed structural topologies, which is key to understanding  
54 the crystal-chemical nature of U-S systems, as well as to the entirety of uranyl mineralogy. The  
55 new uranyl sulfate nitscheite, described herein, possesses a type of uranyl sulfate sheet not  
56 previously observed in Nature.

57 Nitscheite is named in honor of German/American nuclear chemist Heino Nitsche (1949-  
58 2014) for his work on nuclear and radiochemistry of heavy elements, nuclear forensics, the  
59 chemistry of irradiated materials, and the confirmation of elements 114 (flerovium, Fl) and 117  
60 (Tennessine, Ts). Most recently, Nitsche was full professor in the Department of Chemistry at the  
61 University of California, Berkeley, a senior research scientist at Lawrence Berkeley National  
62 Laboratory (LBNL), and the founding director of LBNL's Glenn T. Seaborg Center. In 2014,  
63 Nitsche won the Hevesy Medal, the premier international award of excellence honoring  
64 outstanding achievements in radioanalytical and nuclear chemistry.

65 Although the K analogue of nitscheite has been synthesized (Korniyakov et al. 2020), we  
66 do not propose the use of a  $-(\text{NH}_4)$  suffix in the naming of nitscheite at this time because no K  
67 was detected in its composition. If the K analogue is found to occur naturally, we recommend  
68 that nitscheite be used as the rootname for both species with  $-(\text{NH}_4)$  and  $-(\text{K})$  suffixes added to  
69 distinguish them.

70 The new mineral and name were approved by the Commission on New Minerals,  
71 Nomenclature and Classification of the International Mineralogical Association (IMA 2020-078).  
72 The holotype specimen of nitscheite is deposited in the collections of the Natural History  
73 Museum of Los Angeles County, Los Angeles, California, USA, catalogue number 74163.

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#### OCCURRENCE

76 Nitscheite was found on specimens collected underground in the Green Lizard mine  
77 (37°34'37.10"N 110°17'52.80"W), Red Canyon, White Canyon District, San Juan County, Utah,  
78 USA. The mine is about 72 km west of the town of Blanding, Utah, and about 22 km southeast of  
79 Good Hope Bay on Lake Powell. It is located near the head of Low Canyon on the east side of  
80 Red Canyon, 2.1 km north of the Blue Lizard mine. The geology of the mine is similar to that of  
81 the Blue Lizard mine (Kampf et al. 2017b; Chenoweth 1993). The Green Lizard mine is also a  
82 type locality for shumwayite (Kampf et al. 2017a), greenlizardite (Kampf et al. 2018a), meitnerite  
83 (Kampf et al. 2018b), and straßmannite (Kampf et al. 2019).

84 Abundant secondary uranium mineralization in Red Canyon is associated with post-  
85 mining oxidation of asphaltum-rich sandstone beds laced with uraninite and sulfides in the damp  
86 underground environment. Nitscheite is a very rare mineral in the secondary mineral assemblages  
87 of the Green Lizard mine. It occurs with chinleite-(Y), gypsum, pyrite, and Co-rich rietveldite on  
88 matrix comprised mostly of subhedral to euhedral, equant quartz crystals that are recrystallized  
89 counterparts of the original grains of the sandstone.

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#### PHYSICAL AND OPTICAL PROPERTIES

92 Nitscheite occurs in subparallel and divergent intergrowths of yellow prisms, up to about 0.3 mm

93 in length (Fig. 1). Crystals are elongated on [101] and exhibit the forms {100}, {010}, {001}, and  
94 {11-1} (Fig. 2). [001/0-10/100] twinning by metric merohedry was identified based on the  
95 analysis of structure data. The mineral is transparent with vitreous luster and very pale-yellow  
96 streak. It exhibits bright green fluorescence under a 405 nm laser. It has a Mohs hardness of about  
97 2 based on scratch tests. The mineral has brittle tenacity, curved fracture, and one good cleavage  
98 on {010}. The density measured by flotation in a mixture of methylene iodide and toluene is  
99 3.30(2) g·cm<sup>-3</sup>. The density calculated using the empirical formula and single-crystal unit cell  
100 parameters is 3.278 g·cm<sup>-3</sup>. The mineral is easily soluble in H<sub>2</sub>O at room temperature.

101 Nitscheite is optically biaxial (-) with  $\alpha = 1.560(2)$ ,  $\beta = 1.582(2)$ ,  $\gamma = 1.583(2)$  measured  
102 in white light). The 2V measured using extinction data analyzed with EXCALIBRW (Gunter et  
103 al. 2004) is 17(1)°; the calculated 2V is 23.8°. No dispersion was observed. The optical  
104 orientation is  $X = \mathbf{b}$ ,  $Z \approx [101]$ . The mineral is pleochroic with X colorless, Y and Z yellow;  $X < Y$   
105  $\approx Z$ . The Gladstone–Dale compatibility,  $1 - (K_P/K_C)$ , (Mandarino 2007) is -0.010 (superior) based  
106 on the empirical formula using  $k(\text{UO}_3) = 0.118$ , as provided by Mandarino (1976).

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#### RAMAN SPECTROSCOPY

109 Raman spectroscopy was conducted on a Horiba XploRA PLUS using a 100× (0.9 NA)  
110 objective. The spectrum from 4000 to 60 cm<sup>-1</sup> obtained using a 532 nm diode laser, 50 μm slit,  
111 and 2400 gr/mm diffraction grating is shown in Figure 3. The spectrum from 2000 to 60 cm<sup>-1</sup>  
112 obtained using a 785 nm diode laser, 50 μm slit, and 1800 gr/mm diffraction grating is shown in  
113 Figure 4. The band assignments are based primarily upon those for uranyl sulfate minerals  
114 provided by Čejka (1999) and Plášil et al. (2010). All bands in the spectra were fit using pseudo-  
115 Voigt peak profiles.

116 A multitude of weak bands between  $\sim 2800$  and  $3600\text{ cm}^{-1}$  in the spectrum obtained using  
117 the 532 nm laser are assigned to  $\nu(\text{OH})$  and  $\nu(\text{NH})$  stretching vibrations of interlayer  $\text{NH}_4^+$   
118 groups. Using the empirically derived equation of Libowitzky (1999), the calculated distances of  
119 the corresponding hydrogen bonds range from  $\sim 3.3\text{ \AA}$  to  $\sim 2.6\text{ \AA}$ , in reasonable agreement with  
120 the O $\cdots$ O/N bond lengths determined from the structure refinement. Several very broad low  
121 intensity bands between  $\sim 2200$  and  $\sim 1800\text{ cm}^{-1}$  are probably overtones or combination bands. In  
122 both 532 and 785 nm spectra, no apparent band related to the  $\nu_2(\delta)$  bending vibrations of  $\text{H}_2\text{O}$  is  
123 present at approximately  $1600\text{ cm}^{-1}$ , which is not surprising considering the low sensitivity of  
124 Raman for the non-symmetrical vibrations. Due to the strong fluorescence observed with 532 nm  
125 laser illumination, we continue our discussion of band assignments using fittings obtained from  
126 the 785 nm spectrum.

127 The  $\nu_3(\text{SO}_4)$  antisymmetric stretching vibrations occur as weak bands at 1202, 1156,  
128 1134, and  $1102\text{ cm}^{-1}$ . Several weak to strong bands at 1047, 1038, 1030, 1020, 1007, 993, and  
129  $977\text{ cm}^{-1}$  are assignable to the  $\nu_1$  symmetric stretching vibration of  $\text{SO}_4$  groups. The presence of  
130 six symmetrically distinct  $\text{SO}_4$  tetrahedra in the structure of nitscheite leads to the multiple split  
131 bands in this region. The  $\nu_1(\text{UO}_2)^{2+}$  symmetric stretching vibration is present as a very strong  
132 band at  $856\text{ cm}^{-1}$ , with a weaker, overlapping shoulder at  $852\text{ cm}^{-1}$ . Bartlett and Cooney (1989)  
133 provided an empirical relationship to derive the approximate U–O $_{Ur}$  bond lengths from the band  
134 position assigned to the  $\text{UO}_2^{2+}$  stretching vibrations, which gives  $1.75\text{ \AA}$  ( $856\text{ cm}^{-1}$ ) and  $1.76\text{ \AA}$   
135 ( $852\text{ cm}^{-1}$ ), in excellent agreement with the average U1–O $_{Ur}$  bond length from the X-ray data:  
136  $1.763\text{ \AA}$ . At least five overlapping weak bands between  $\sim 650$  and  $600\text{ cm}^{-1}$  are attributable to the  
137  $\nu_4(\delta)(\text{SO}_4)$  bending vibrations, with centers at 650, 639, 629, 620, and  $614\text{ cm}^{-1}$ . Weak bands at  
138 457, 451, and  $443\text{ cm}^{-1}$  belong to the  $\nu_2(\delta)(\text{SO}_4)$  bending vibrations. A set of very broad and

139 very weak bands between  $\sim 400$  and  $350\text{ cm}^{-1}$  likely are due to out-of-plane bending vibrations of  
140 U–O<sub>eq</sub> bonds in the sheet. A complex group of bands between  $250\text{ cm}^{-1}$  and  $\sim 220\text{ cm}^{-1}$ , with  
141 centers at 279, 265, 260, 255, and  $228\text{ cm}^{-1}$  are attributable to the doubly degenerate  $\nu_2(\delta)$   
142 (UO<sub>2</sub>)<sup>2+</sup> bending vibrations and possibly to  $\nu$  (U–O<sub>eq</sub>) bending modes. The remaining bands  
143 below  $\sim 200\text{ cm}^{-1}$  arise due to unassigned phonon modes and molecular deformations.

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### CHEMICAL ANALYSIS

146 Electron probe microanalyses (5 points on 5 crystals) were performed at the University of Utah  
147 on a Cameca SX-50 electron microprobe with four wavelength dispersive spectrometers and  
148 using Probe for EPMA software. Analytical conditions were 15 kV accelerating voltage, 10 nA  
149 beam current, and 10  $\mu\text{m}$  beam diameter. Raw X-ray intensities were corrected for matrix effects  
150 with a  $\phi\rho(z)$  algorithm (Pouchou and Pichoir 1991). Because of the presence of substantial H<sub>2</sub>O,  
151 the formula concentration of oxygen was used in the matrix correction. A synthetic PC1 W/Si  
152 multilayer “crystal” was used for N analysis. Time-dependent, log-linear, corrections were  
153 applied to N (decreasing intensity) and U (increasing intensity). No other elements were detected.  
154 There was major beam damage. Because insufficient material is available for a direct  
155 determination of H<sub>2</sub>O, it has been calculated based upon the structure (U = 2 *apfu* and O = 21  
156 *apfu*). Analytical data are given in Table 1. The empirical formula is (NH<sub>4</sub>)<sub>1.99</sub>U<sub>2.00</sub>S<sub>3.00</sub>O<sub>21</sub>H<sub>10.00</sub>.  
157 The ideal formula is (NH<sub>4</sub>)<sub>2</sub>[(UO<sub>2</sub>)<sub>2</sub>(SO<sub>4</sub>)<sub>3</sub>(H<sub>2</sub>O)<sub>2</sub>]·3H<sub>2</sub>O, which requires (NH<sub>4</sub>)<sub>2</sub>O 5.46, UO<sub>3</sub>  
158 59.94, SO<sub>3</sub> 25.17, H<sub>2</sub>O 9.44, total 100 wt%.

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### X-RAY CRYSTALLOGRAPHY AND STRUCTURE DETERMINATION

161 Both powder and single-crystal X-ray studies were carried out using a Rigaku R-Axis Rapid II  
162 curved imaging plate microdiffractometer with monochromatized MoK $\alpha$  radiation. For the  
163 powder study, a Gandolfi-like motion on the  $\phi$  and  $\omega$  axes was used to randomize the sample,  
164 which consisted of several crystals. Observed  $d$  values and intensities were derived by profile  
165 fitting using JADE 2010 software (Materials Data, Inc. Livermore, CA). The observed powder  
166 diffraction pattern compares very well with the pattern calculated from the crystal structure; data  
167 are given in Supplemental<sup>1</sup> Table S1.

168 Crystals occur in subparallel intergrowths making the selection of a single crystal  
169 challenging. The crystal fragment used for the data collection comprised one major crystal and  
170 several subparallel satellite crystals. Initial indexing of single-crystal reflections suggested an  
171 orthorhombic unit cell:  $a = 11.5685(3)$ ,  $b = 12.8520(9)$ ,  $c = 12.9949(4)$  Å. This is comparable to  
172 the cells reported for synthetic  $(\text{NH}_4)_2(\text{UO}_2)_2(\text{SO}_4)_3 \cdot 5\text{H}_2\text{O}$  (Staritzky et al. 1956) and analogous  
173 phases such as  $\text{Rb}_2(\text{UO}_2)_2(\text{SO}_4)_3 \cdot 5\text{H}_2\text{O}$  (Serezhkina et al. 1990), which were purported to be  
174 orthorhombic with likely space group  $Pnma$  based on PXRD data, but without structure  
175 determination. Furthermore, the PXRD of nitscheite is a close match to those for these synthetic  
176 phases. We were unable to solve the structure using this cell and any orthorhombic space group.  
177 Ultimately, we solved the structure in space group  $P2_1/n$  using the larger ( $\times 2$ ) monoclinic cell  
178 reported (Table 2). The metrically orthorhombic cell noted above is a subcell of our monoclinic  
179 cell (see Fig. 5). Recently, Korniyakov et al. (2020) reported the structure of the synthetic K  
180 analogue of nitscheite,  $\alpha\text{-K}_2[(\text{UO}_2)_2(\text{SO}_4)_3(\text{H}_2\text{O})_2](\text{H}_2\text{O})_3$ , also with space group  $P2_1/n$  and with a  
181 comparable monoclinic cell.

182 The structure data for nitscheite were processed using the Rigaku CrystalClear software  
183 package, including the application of an empirical multi-scan absorption correction using



184 ABSCOR (Higashi, 2001). The structure was solved using SHELXT (Sheldrick, 2015a).  
185 Refinement proceeded by full-matrix least-squares on  $F^2$  using SHELXL-2016 (Sheldrick,  
186 2015b). Most atoms were located in the initial structure solution and the remaining non-hydrogen  
187 atoms were located in subsequent difference Fourier syntheses. At this stage, only U and S sites  
188 could be successfully refined with anisotropic displacement parameters,  $R_1$  (for 5006  $I > 2\sigma I$   
189 reflections) converged to a high value: 0.093, there were high positive and negative electron  
190 residuals (+5.29 and  $-4.88 e \cdot \text{\AA}^{-3}$ ), and there were some anomalous interatomic distances.  
191 Korniyakov et al. (2020) reported similar issues with their refinement of the structure of  $\alpha$ -  
192  $\text{K}_2[(\text{UO}_2)_2(\text{SO}_4)_3(\text{H}_2\text{O})_2](\text{H}_2\text{O})_3$ , and related them to the twinning due to pseudo-merohedry. The  
193 HKLF5 type of reflection file created by the utility in PLATON (Spek 2003) was found as the  
194 only way to handle the twinning successfully. Twinning in nitscheite is by reticular merohedry;  
195 the mirror in  $\{101\}$ , expressed by the matrix  $\begin{vmatrix} 0 & 0 & -1 \\ 0 & 1 & 0 \\ -1 & 0 & 0 \end{vmatrix}$  causes twinning of the  
196 nitscheite cell, which gives rise to an orthorhombic supercell, with  $a = 12.855 \text{ \AA}$ ,  $b = 23.141 \text{ \AA}$ ,  
197  $c = 25.996 \text{ \AA}$ ,  $\alpha = 89.98^\circ$ ,  $\beta = 90^\circ$ ,  $\gamma = 90^\circ$ , and  $V = 2 \cdot 3866.7 \text{ \AA}^3$ . The subsequent refinement in  
198 the JANA2006 program (Petříček et al. 2016), using the HKLF5 file and invoking the  
199 abovementioned twinning resolved all of the issues encountered in our initial structure  
200 refinement. All non-hydrogen atom sites were refined to full occupancies with anisotropic  
201 displacement parameters; however, difference Fourier syntheses were unsuccessful in locating  
202 hydrogen atom positions. The refined twin ratio equals to 0.7304(11)/0.2696(11). The refinement  
203 converged to  $R_1 = 3.29\%$  for 4547 reflections with  $I > 3\sigma(I)$ . Data collection and refinement  
204 details are given in Table 2, atom coordinates and displacement parameters can be found in the  
205 original CIF (as supplemental file<sup>1</sup>), selected bond distances in Table 3, and a bond-valence  
206 analysis in Table 4.

207

208

## DESCRIPTION OF THE STRUCTURE

209 Four U sites (U1, U2, U3, and U4) in the structure of nitscheite are each surrounded by seven O  
210 atoms forming squat  $UO_7$  pentagonal bipyramids. This is a typical coordination for  $U^{6+}$  in which  
211 the two short apical bonds of the bipyramid constitute the uranyl group (Burns, 2005). The two  
212 apical O atoms of the bipyramids ( $O_{Ur}$ ) form short bonds with the U, and this unit comprises the  
213  $UO_2^{2+}$  uranyl group. Five equatorial O atoms ( $O_{eq}$ ) complete the U coordination environment,  
214 some of which include O of  $H_2O$  groups with long U–O bond distances  $>2.4$  Å. The U1 site  
215 bonds with two  $H_2O$  groups (Ow3 and Ow4), U2 and U3 bond with a single  $H_2O$  group (Ow1  
216 and Ow2, respectively), and U4 forms no bonds with  $H_2O$  groups.

217 There are six S sites (S1, S2, S3, S4, S5, and S6) each centering an  $SO_4$  tetrahedron. The  
218  $SO_4$  tetrahedra share corners with the equatorial O atoms of the  $UO_7$  bipyramids to form a uranyl-  
219 sulfate sheet with the composition  $[(UO_2)_2(SO_4)_3(H_2O)_2]^{2-}$  (Fig. 5). Within this sheet, the U1  
220 bipyramid shares three of its  $O_{eq}$  corners with  $SO_4$  groups, the U2 and U3 bipyramids each share  
221 four of their  $O_{eq}$  corners with  $SO_4$  groups, and the U4 bipyramid shares all five of its  $O_{eq}$  corners  
222 with  $SO_4$  groups. The uranyl-sulfate sheet in the structure of nitscheite is unique in the mineral  
223 kingdom, but it has the same topology as the sheet in  $\alpha$ - $K_2[(UO_2)_2(SO_4)_3(H_2O)_2](H_2O)_3$  and  $\beta$ -  
224  $K_2[(UO_2)_2(SO_4)_3(H_2O)_2](H_2O)_3$  (Korniyakov et al. 2020), and it has similarities to those in several  
225 minerals and other synthetic phases (See Table 5 and Figure 5 in Lussier et al. 2016).

226 The interlayer region between the uranyl sulfate sheets (Fig. 6) contains four  $NH_4$  sites  
227 (N1, N2, N3, and N4) and six  $H_2O$  sites (OW5, OW6, OW7, OW8, OW9, and OW10). Each of  
228 the  $NH_4$  sites is eight-coordinated (for N–O  $< 3.44$  Å), linking to at least three O sites in the  
229 adjacent negatively charged sheets and providing essential charge balance.

230 It is noteworthy that the structural complexity (after Krivovichev 2012, 2013, 2014, 2018)  
231 of nitscheite,  $I_{G,\text{total}}$ , is very high, 2400.67 bits/cell (including a theoretical positions of the H  
232 atoms). Nitscheite is the second most complex uranyl sulfate mineral known, following closely  
233 behind natrozippeite, with 2528.63 bits/cell (Gurzhiy and Plášil 2019). Both minerals have unit-  
234 cell volumes greater than 3500 Å<sup>3</sup> and their structures are characterized by very large numbers of  
235 hydrogen bonds.

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### IMPLICATIONS

238 Recent investigations of synthetic and natural alkali and alkaline earth uranyl sulfate hydrates  
239 have shown that these phases adopt highly diverse bonding arrangements based on cluster, chain,  
240 sheet, and framework topologies (Gurzhiy and Plášil 2019; Tyumentseva et al. 2019; Korniyakov  
241 et al. 2020). This is partly due to the large degree of freedom of polymerization between U and  
242 SO<sub>4</sub> groups, as well as significant variability in coordination geometry of Na, K, and NH<sub>4</sub>  
243 cations, permitting a large number of crystal-chemically stable structures to form with only minor  
244 differences in U:S:Me:H<sub>2</sub>O content. Other factors such as pH and relative humidity influence the  
245 structure of these phases due to extensive variability in hydrogen bonding, while idiosyncrasies in  
246 associated minerals, along with sequential dissolution or crystallization on unique substrates has  
247 formed mineral structures that have not yet been reproduced under laboratory conditions.

248 For example, the mineral geschieberite, K<sub>2</sub>[(UO<sub>2</sub>)(SO<sub>4</sub>)<sub>2</sub>(H<sub>2</sub>O)](H<sub>2</sub>O) (Plášil et al. 2015),  
249 which has a higher H<sub>2</sub>O and S:U content than nitscheite, forms a distinct sheet topology due to  
250 the depolymerizing action of an equatorially bonded H<sub>2</sub>O group about the U atom in its structure.  
251 Such is also the case for the mineral beshtauite, (NH<sub>4</sub>)<sub>2</sub>(UO<sub>2</sub>)(SO<sub>4</sub>)<sub>2</sub>(H<sub>2</sub>O)<sub>2</sub> (Pekov et al. 2014),  
252 which contains a topology closely related to that of geschieberite – the only difference being an

253 alternating directionality of equatorial H<sub>2</sub>O groups in their sheets. The cause of this subtle  
254 topological “flip” is unknown and, although NH<sub>4</sub><sup>+</sup> and K<sup>+</sup> cations commonly substitute to form  
255 isomorphous minerals, we note that the potential K-analogue of nitscheite may adopt a similarly  
256 unique topological arrangement (Fig. 7). Nitscheite contains U sites with 0, 1, and 2 equatorially  
257 bonded H<sub>2</sub>O groups, arranged into the unique and novel topology described here.

258           It is unfortunate that the recent permanent sealing of all mines in Red Canyon and  
259 surrounding areas, enforced by the State of Utah and funded by the U.S. Department of Energy,  
260 has eliminated access to one of the most remarkable sources of uranyl sulfate minerals on Earth.  
261 The study of the diverse secondary mineralogy of Utah’s abandoned mines, which now must rely  
262 on previously collected samples, continues to yield invaluable insights into the environmental  
263 behavior of U in ways that only Nature can provide.

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#### ACKNOWLEDGEMENTS

266           Reviewers Sergey Krivovichev and Fernando Colombo are thanked for constructive  
267 comments, which improved the manuscript. We are grateful to retired miner Dan Shumway of  
268 Blanding, Utah, for advice and assistance in our collecting efforts in Red Canyon. Funding to JP  
269 was provided by the Czech Science Foundation (20-11949S). This study was also funded by the  
270 John Jago Trelawney Endowment to the Mineral Sciences Department of the Natural History  
271 Museum of Los Angeles County.

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358

359 **Endnote:**



360 <sup>1</sup>Deposit item AM-21-XXXXX, Supplemental tables and CIF. Deposit items are free to all  
361 readers and found on the MSA website, via the specific issue's Table of Contents (go to  
362 [http://www.minsocam.org/MSA/AmMin/TOC/2021/Xxx2021\\_data/Xxx2021\\_data.html](http://www.minsocam.org/MSA/AmMin/TOC/2021/Xxx2021_data/Xxx2021_data.html)).  
363

364

## FIGURE CAPTIONS

365 Figure 1. Nitscheite prisms with gypsum and pyrite on quartz. The field of view is 0.5 mm  
366 across.

367 Figure 2. Crystal drawing of nitscheite (clinographic projection in non-standard orientation, [101]  
368 vertical).

369 Figure 3. The Raman spectrum of nitscheite recorded with a 532 nm laser.

370 Figure 4. The fitted baseline-corrected Raman spectrum of nitscheite recorded with a 785 nm  
371 laser.

372 Figure 5. The  $[(\text{UO}_2)_2(\text{SO}_4)_3(\text{H}_2\text{O})_2]$  sheet in nitscheite viewed down [010].  $\text{UO}_7$  and  $\text{SO}_4$   
373 polyhedra are labelled. The unit cell outline is indicated by dashed black lines. The  
374 outline of the metrically orthorhombic subcell is shown by dashed red lines with cell edge  
375 lengths labeled. (color online).

376 Figure 6. The crystal structure of nitscheite viewed down [001]. N atoms of  $\text{NH}_4$  groups are small  
377 red balls and O atoms of interlayer  $\text{H}_2\text{O}$  groups are large white balls. The unit cell is  
378 indicated by dashed lines. (color online)

379 Figure 7. Cation topology of (a) nitscheite and (b) synthetic  $\alpha\text{-K}_2[(\text{UO}_2)_2(\text{SO}_4)_3(\text{H}_2\text{O})_2](\text{H}_2\text{O})_3$   
380 (Korniyakov et al. 2020). Interlayer cations (large balls: N-blue, K-magenta) are projected  
381 on U-S layers (small balls: U-yellow, S-blue). (color online)

382

383 Table 1. Chemical composition of nitscheite.

<b>Constituent</b>	<b>Mean</b>	<b>Range</b>	<b>Stand. Dev.</b>	<b>Standard</b>
(NH <sub>4</sub> ) <sub>2</sub> O	5.42	4.18–6.32	0.89	syn. Cr <sub>2</sub> N
UO <sub>3</sub>	59.75	57.33–62.09	1.70	syn. UO <sub>2</sub>
SO <sub>3</sub>	25.12	24.34–25.76	0.72	celestine
H <sub>2</sub> O*	9.41			
<b>Total</b>	<b>99.70</b>			

384 \* based upon the structure (U = 2 *apfu* and O = 21 *apfu*).

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385

386 Table 2. Data collection and structure refinement details for nitscheite.  
387

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388	Diffractometer	Rigaku R-Axis Rapid II
389	X-ray radiation/power	MoK $\alpha$ ( $\lambda = 0.71075 \text{ \AA}$ )/50 kV, 40 mA
390	Temperature	293(2) K
391	Structural Formula	(NH <sub>4</sub> ) <sub>2</sub> [(UO <sub>2</sub> ) <sub>2</sub> (SO <sub>4</sub> ) <sub>3</sub> (H <sub>2</sub> O) <sub>2</sub> ]·3H <sub>2</sub> O (H atoms not located)
392	Space group	<i>P</i> 2 <sub>1</sub> / <i>n</i>
393	Unit cell dimensions	<i>a</i> = 17.3982(4) $\text{\AA}$
394		<i>b</i> = 12.8552(3) $\text{\AA}$
395		<i>c</i> = 17.4054(12) $\text{\AA}$
396		$\beta = 96.649(7)^\circ$
397	<i>V</i>	3866.7(3) $\text{\AA}^3$
398	<i>Z</i>	8
399	Density (for above formula)	3.251 g·cm <sup>-3</sup>
400	Absorption coefficient	17.157 mm <sup>-1</sup>
401	<i>F</i> (000)	3392
402	Crystal size	170 × 25 × 20 $\mu\text{m}$
403	$\theta$ range	3.13 to 25.04°
404	Index ranges	$-20 \leq h \leq 20, 0 \leq k \leq 15, 0 \leq l \leq 20$
405	Reflections collected/unique	35158/6491; $R_{\text{int}} = 0.081$
406	Reflections with $I > 3\sigma I$	4547
407	Completeness to $\theta = 25.04^\circ$	95.2%
408	Refinement method	Full-matrix least-squares on $F^2$
409	Parameters/restraints	506/0
410	GoF	1.181
411	Refined twin ratio	0.7304(11)/0.2696(11)
412	Final <i>R</i> indices [ $F > 4\sigma(F)$ ]	$R_1 = 0.0329, wR_2 = 0.0681$
413	<i>R</i> indices (all data)	$R_1 = 0.0532, wR_2 = 0.0752$
414	Weighting scheme, weights	Weighting scheme based on measured s.u.'s; $w = 1/(\sigma^2(I) +$
415		$0.0004I^2)$
416	Largest diff. peak/hole	+2.58/−1.32 $e \cdot \text{\AA}^{-3}$

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417

418 Table 3. Selected bond distances (Å) for nitscheite.  
419

420	N1–O17	2.803(15)	N2–O9	2.775(16)	N3–O13	2.976(18)	N4–O8	2.946(14)
421	N1–O18	2.884(15)	N2–O10	2.817(16)	N3–OW10	2.98(2)	N4–O21	2.987(17)
422	N1–O5	2.935(15)	N2–O13	2.980(16)	N3–O3	2.991(16)	N4–O23	2.995(14)
423	N1–O21	3.018(15)	N2–O1	3.005(16)	N3–O16	3.014(16)	N4–O5	3.003(18)
424	N1–O30	3.042(16)	N2–O26	3.051(17)	N3–OW6	3.03(2)	N4–OW5	3.06(2)
425	N1–OW8	3.171(18)	N2–OW7	3.101(18)	N3–O1	3.044(17)	N4–OW9	3.06(2)
426	N1–O27	3.279(16)	N2–O28	3.320(17)	N3–O2	3.072(15)	N4–O24	3.076(15)
427	N1–O25	3.305(16)	N2–O31	3.324(17)	N3–O14	3.173(16)	N4–O7	3.232(15)
428	<N1–O>	3.055	<N2–O>	3.047	<N3–O>	3.035	<N4–O>	3.045
429								
430	U1–O25	1.765(9)	U2–O27	1.751(10)	U3–O29	1.746(10)	U4–O31	1.750(10)
431	U1–O26	1.766(10)	U2–O28	1.755(9)	U3–O30	1.762(9)	U4–O32	1.771(9)
432	U1–O4	2.323(7)	U2–O6	2.335(11)	U3–O19	2.364(9)	U4–O15	2.337(8)
433	U1–O23	2.333(9)	U2–O12	2.367(7)	U3–O16	2.371(8)	U4–O8	2.374(7)
434	U1–O24	2.378(8)	U2–O3	2.372(8)	U3–O14	2.373(9)	U4–O11	2.375(9)
435	U1–OW4	2.423(9)	U2–O2	2.381(9)	U3–O22	2.377(9)	U4–O7	2.393(9)
436	U1–OW3	2.446(9)	U2–OW1	2.495(9)	U3–OW2	2.429(8)	U4–O20	2.398(8)
437	<U1–O <sub>ap</sub> >	1.766	<U2–O <sub>ap</sub> >	1.753	<U3–O <sub>ap</sub> >	1.754	<U4–O <sub>ap</sub> >	1.761
438	<U1–O <sub>eq</sub> >	2.381	<U2–O <sub>eq</sub> >	2.390	<U3–O <sub>eq</sub> >	2.383	<U4–O <sub>eq</sub> >	2.375
439								
440	S1–O1	1.462(10)	S2–O5	1.414(11)	S3–O9	1.449(11)		
441	S1–O2	1.474(9)	S2–O6	1.471(11)	S3–O10	1.456(9)		
442	S1–O3	1.483(9)	S2–O7	1.479(10)	S3–O11	1.482(10)		
443	S1–O4	1.488(8)	S2–O8	1.490(8)	S3–O12	1.501(8)		
444	<S1–O>	1.477	<S2–O>	1.464	<S3–O>	1.472		
445								
446	S4–O13	1.442(11)	S5–O17	1.438(11)	S6–O21	1.436(11)		
447	S4–O15	1.474(10)	S5–O18	1.465(10)	S6–O22	1.466(9)		
448	S4–O16	1.479(9)	S5–O19	1.496(9)	S6–O23	1.498(10)		
449	S4–O14	1.480(10)	S5–O20	1.507(10)	S6–O24	1.507(9)		
450	<S4–O>	1.469	<S5–O>	1.477	<S6–O>	1.477		

451

Table 4. Bond valence analysis for nitscheite. Values are expressed in valence units.\*

	N1	N2	N3	N4	U1	U2	U3	U4	S1	S2	S3	S4	S5	S6	sum
O1		0.12	0.11						1.54						1.77
O2			0.10			0.49			1.49						2.08
O3			0.13			0.50			1.46						2.09
O4					0.56				1.44						2.00
O5	0.15			0.12						1.74					2.01
O6						0.54				1.50					2.04
O7				0.07				0.48		1.47					2.02
O8				0.14				0.50		1.43					2.07
O9		0.23									1.59				1.82
O10		0.20									1.56				1.76
O11								0.50			1.46				1.96
O12						0.51					1.40				1.91
O13		0.13	0.13									1.62			1.88
O14			0.08				0.50					1.47			2.05
O15								0.54				1.49			2.03
O16			0.12				0.50					1.47			2.09
O17	0.21												1.63		1.84
O18	0.17												1.53		1.70
O19							0.51						1.41		1.92
O20								0.48					1.37		1.85
O21	0.12			0.13										1.64	1.89
O22							0.50							1.52	2.02
O23				0.13	0.55									1.41	2.09
O24				0.10	0.50									1.37	1.97
O25	0.05				1.81										1.86
O26		0.11			1.81										1.92
O27	0.06					1.87									1.93
O28		0.05				1.85									1.90
O29							1.89								1.89
O30	0.11						1.82								1.93
O31		0.05						1.87							1.92
O32								1.79							1.79
OW1						0.39									0.39
OW2							0.44								0.44
OW3					0.43										0.43
OW4					0.45										0.45
OW5				0.10											0.10
OW6			0.11												0.11
OW7		0.09													0.09
OW8	0.08														0.08
OW9				0.10											0.10
OW10			0.13												0.13
sum	0.94	0.99	0.91	0.89	6.09	6.15	6.16	6.15	5.93	6.15	6.02	6.05	5.94	5.95	

\*  $\text{NH}_4^+$ -O bond valence parameters from Garcia-Rodriguez et al. (2000);  $\text{U}^{+6}$ -O and  $\text{S}^{+6}$ -O

bond-valence parameters from Gagné and Hawthorne (2015). Hydrogen bond contributions are

not included.

Figure 1



Figure 2

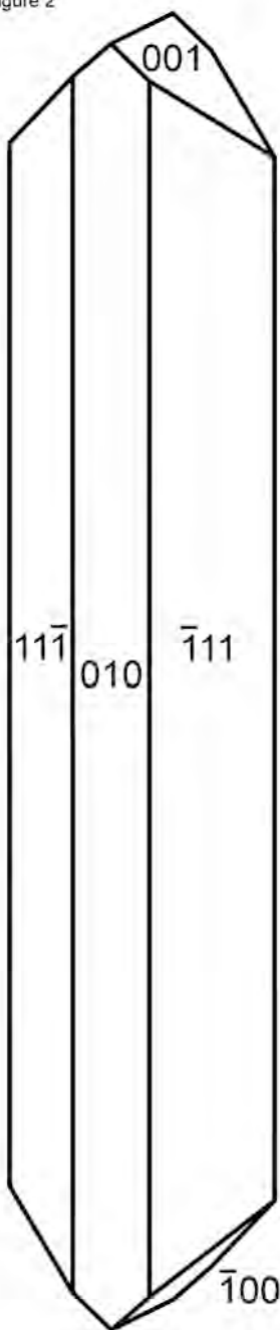




Figure 3

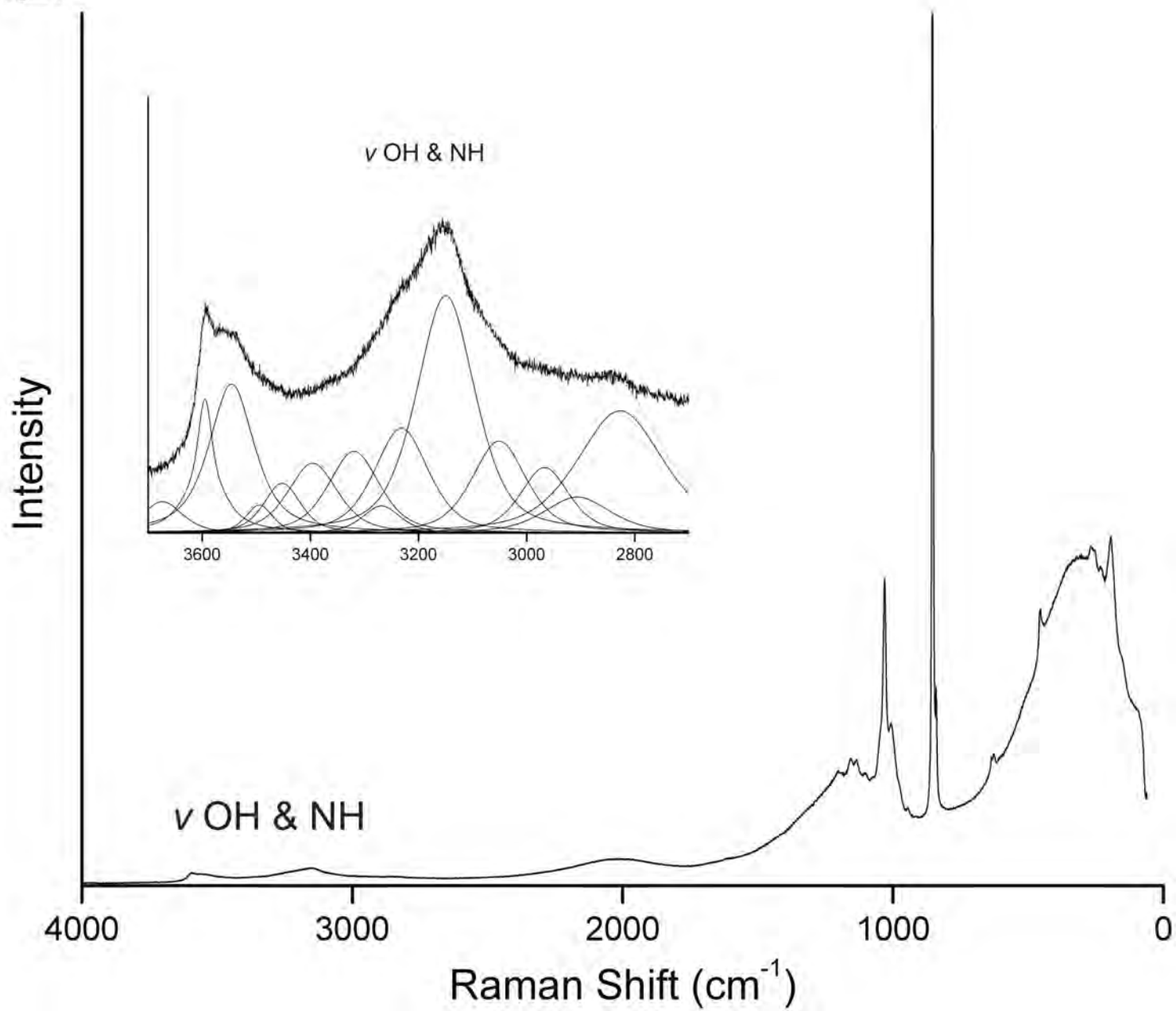


Figure 4

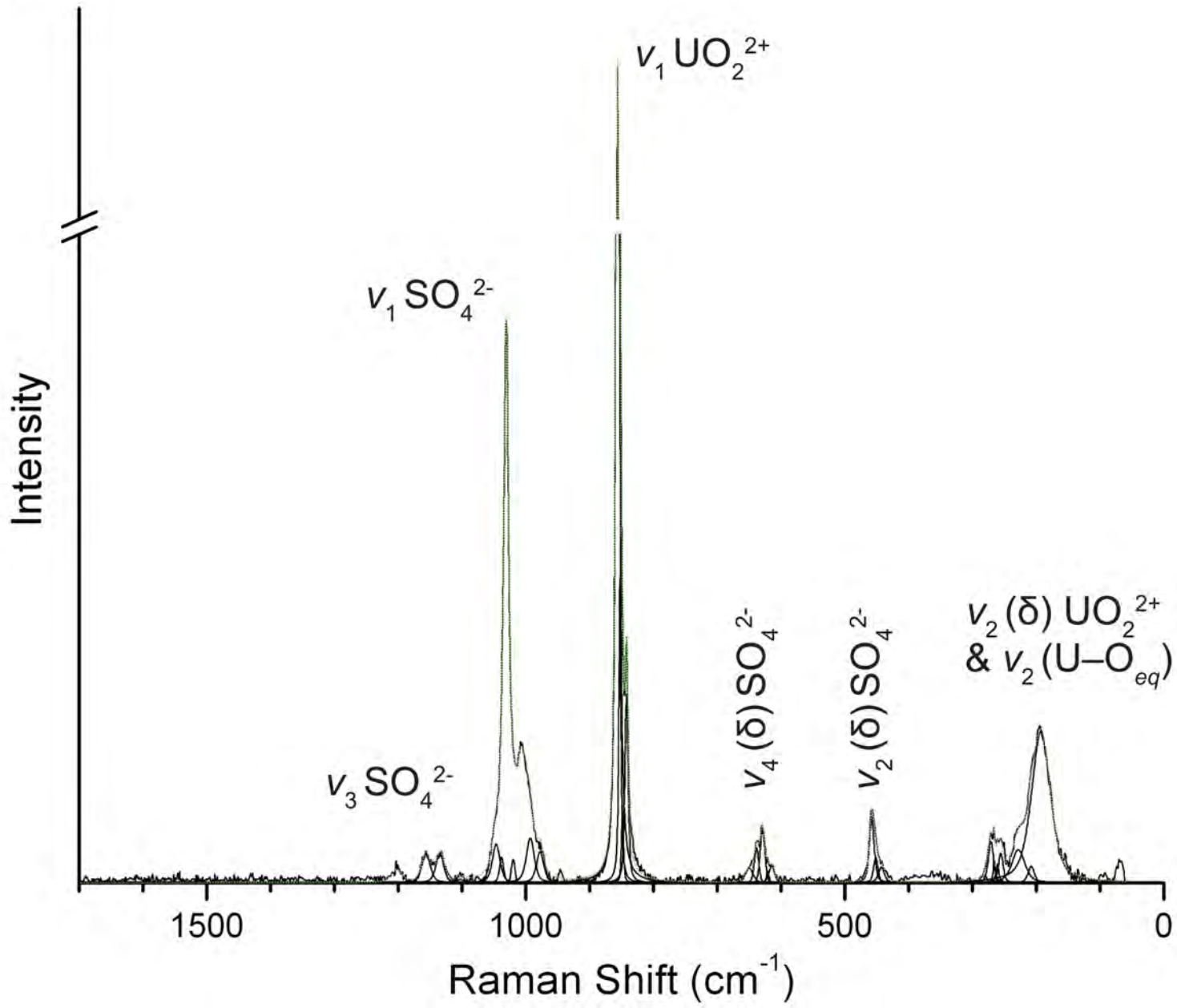


Figure 5

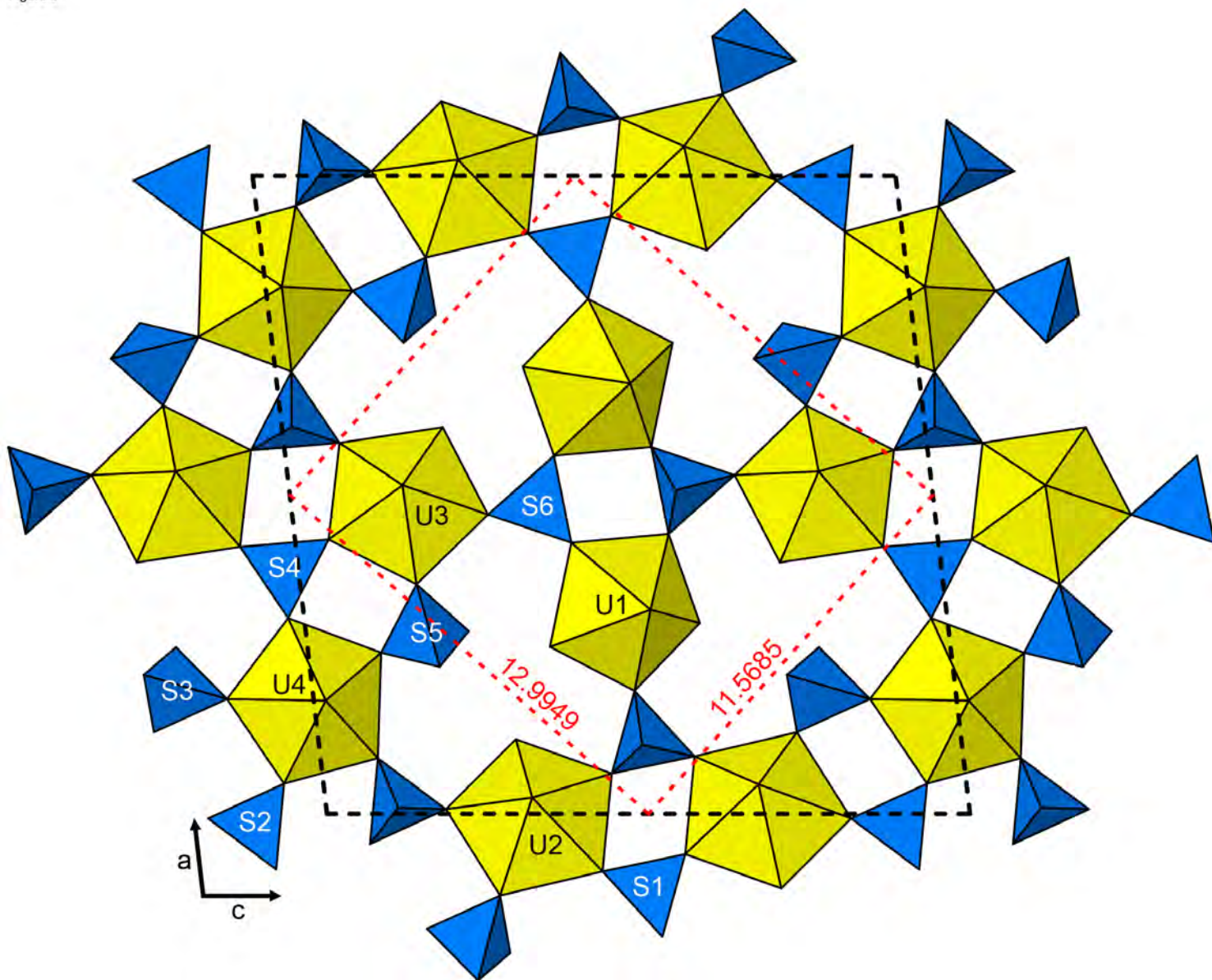


Figure 6

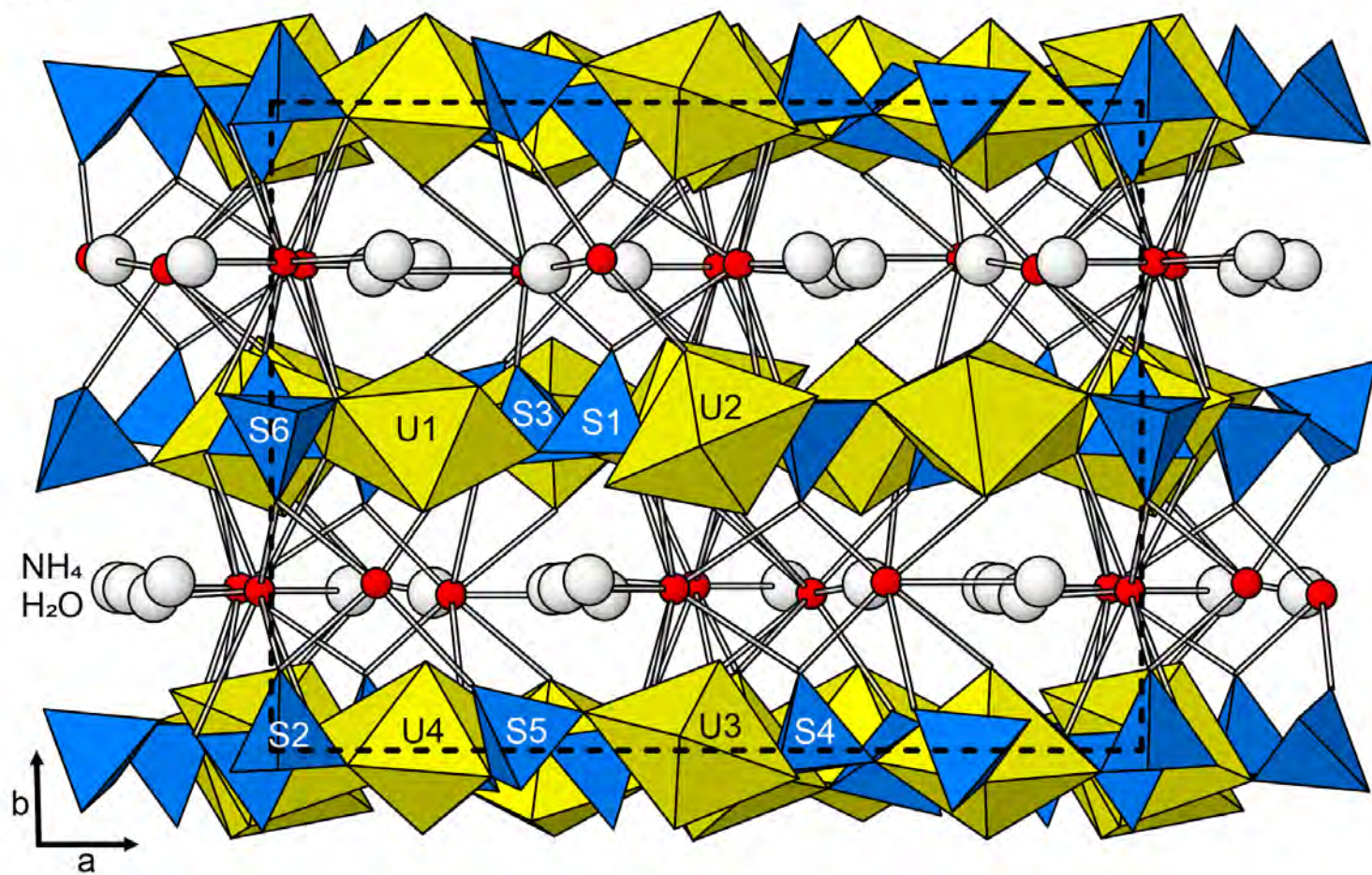


Figure 7

