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2 **Incorporation of incompatible trace elements into molybdenite: Layered**

3 **PbS precipitates within molybdenite**

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## ABSTRACT

20 Trace elements in molybdenite can provide important information regarding the  
21 composition of ore-forming fluid and the evolution and genesis of ore deposits.  
22 However, the occurrence states and behavior of relatively incompatible trace elements  
23 (e.g., Pb and Os) in natural molybdenite remain ambiguous. Here, we report an  
24 abnormally high enrichment of Pb and layered PbS precipitate within molybdenite  
25 grains from the Huanglongpu carbonatite-hosted Mo-Pb deposit in the Qinling  
26 orogenic belt of Northern China. High-resolution transmission electron microscopy  
27 (HRTEM) and related nano-beam techniques were applied to characterize the  
28 occurrence states of Pb within molybdenite at the atomic scale. The results show that  
29 up to several weight percent of Pb can be incorporated into the molybdenite structure  
30 during initial crystallization, which can lead to the formation of screw dislocations  
31 and  $3R$ /disordered stacking of S-Mo-S sandwich layers. Observations using a  
32 scanning transmission electron microscope also reveal that Pb diffuses from the host  
33 molybdenite into the layered PbS precipitates under prolonged electron beam  
34 irradiation. Pb-bearing molybdenite tends to transform into a Pb-poor ordered  $2H_1$   
35 polytype upon Pb exsolution during cooling. Pb preferentially exsolves along the (001)  
36 plane of molybdenite and is stored in structural defects (e.g., dislocation loops) and  
37 grain boundaries, resulting in nano-scale Pb heterogeneities in molybdenite. Further  
38 coarsening of the exsolved Pb results in the formation of layered PbS precipitates  
39 along the (001) plane of molybdenite. This study provides an example of the  
40 consequences of the incorporation and exsolution of incompatible trace elements in  
41 molybdenite and demonstrates that careful mineralogical examination is required to  
42 interpret geochemical data obtained by in-situ analysis techniques.

43 **Keywords:** Molybdenite, layered PbS, polytype, incompatible trace elements

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## INTRODUCTION

46 Molybdenite ( $\text{MoS}_2$ ) is a layered-structure sulfide mineral with two hexagonally  
47 coordinated S layers and one sandwiched Mo layer. The sandwich-like [S-Mo-S]  
48 layers that stack along the *c*-axis are weakly bonded by van der Waals forces.  
49 Molybdenite hosts important economic resources of Mo and Re (Fleischer 1959;  
50 Frondel and Wickman 1970; Terada et al. 1971; Plotinskaya et al. 2018).  $^{187}\text{Re}$ - $^{187}\text{Os}$   
51 systematics in molybdenite are also widely applied to date geological events (Stein et  
52 al. 2001, 2003), and trace element (e.g., Pb, Bi, Re, and W) characteristics provide  
53 important insight into the multi-stage mineralization history of related deposits  
54 (Ciobanu et al. 2013; Barra et al. 2017).

55 Highly compatible trace elements (e.g., Re and W) are preferentially  
56 incorporated into the molybdenite structure due to their same valence and similar  
57 ionic radii (e.g.,  $\text{Re}^{4+}$ , 0.63 Å;  $\text{W}^{4+}$ , 0.66 Å) as  $\text{Mo}^{4+}$  (0.65 Å), and are commonly  
58 considered to be relatively immobile (Stein et al. 2003). In contrast, relatively  
59 incompatible trace elements (e.g., Pb, Bi, Te, and Au) commonly occur as  
60 independent mineral phases in molybdenite (Petruk 1965; Ciobanu et al. 2013; Pasava  
61 et al. 2016; Plotinskaya et al. 2018), have relatively higher mobilities (Stein et al.  
62 2003), and tend to form atomic clusters within structural defects and mm- to nm-sized  
63 exsolution/inclusion phases (Stein et al. 2003; Pasava et al. 2016).

64 An understanding of the mobility of relatively incompatible trace elements in  
65 molybdenite is critical to accurately interpreting the trace element data and retrieved  
66 isotopic ages. Previous studies have demonstrated the effectiveness of using  
67 transmission electron microscopy (TEM) to observe the behavior (e.g., occupancy and  
68 migration) of Au and Re in synthetic  $\text{MoS}_2$  under electron beam irradiation (Lin et al.

69 2014a, 2014b). However, the occurrence states and mobility of relatively  
70 incompatible trace elements in molybdenite with regards to different geological  
71 conditions have received scarce attention in the literature (Stein et al. 2003). The  
72 incorporation of trace elements may also influence the formation of polytypes (e.g.,  
73  $3R$  and  $2H_1$ ) in molybdenite by introducing lattice strain and screw dislocations in the  
74 stacked S-Mo-S sandwich layers (Newberry 1979a, 1979b); nevertheless, the  
75 relationship between incorporated trace elements and the formation of specific  
76 polytypes remains ambiguous (Fron del and Wickman 1970; McCandless et al. 1993;  
77 Newberry 1979a, 1979b; Voudouris et al. 2009; Grabezhev and Voudouris 2014;  
78 Plotinskaya et al. 2019).

79 Although relatively incompatible, Pb commonly occurs in relatively high  
80 concentrations (ppm to wt.% levels) in molybdenite (Stein et al. 1997; Du et al. 2004;  
81 Ciobanu et al. 2013; Pasava et al. 2016), which makes it an ideal candidate for  
82 investigating the behavior of relatively incompatible trace elements. In this study,  
83 high-resolution transmission electron microscopy (HRTEM) is used with a  
84 combination of nanobeam techniques to investigate the occurrence of Pb in  
85 molybdenite from the Huanglongpu Mo-Pb ore deposit. The results provide important  
86 insight into the behavior of Pb in molybdenite and the correlation between specific  
87 polytypes and incompatible trace elements. These findings also provide critical clues  
88 for interpreting the uncertainty and anomaly of geochemical data obtained by in-situ  
89 analysis techniques.

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## 91 **SAMPLING AND ANALYTICAL METHODS**

### 92 **Sampling**

93 The Qinling orogenic belt is an important Mo mineralization area in China with  
94 most Mo contained within the porphyry systems associated with Late Mesozoic  
95 granite (Li and Pirajno 2017). Mo mineralization is particularly notable in the  
96 Huanglongpu Mo-Pb deposit located in the eastern Qinling orogen belt, which is  
97 genetically related to Early Mesozoic carbonatite dykes and mantle-derived elemental  
98 resources including Mo, Pb, C, O, and S (Huang et al. 1985; Stein et al. 1997; Xu et al.  
99 2009; Song et al. 2015, 2016; Li and Pirajno 2017). The carbonatite veins consist of  
100 40–60 vol.% calcite, 30–50 vol.% quartz, ~5 vol.% microcline, ~4.5 vol.%  
101 barytocelestine, ~1 vol.% biotite, and significant quantities of molybdenite, galena,  
102 and pyrite (Stein et al. 1997).

103 Six molybdenite-bearing specimens were collected from the carbonatite veins in  
104 the Huanglongpu Mo-Pb deposit in mostly calcite aggregates (Figs. 1a, 1b). The  
105 samples contain two distinct molybdenite occurrences: (1) sparsely disseminated and  
106 speckled in the calcite and quartz (Fig. S1a) and (2) densely disseminated and  
107 distributed in small strips in the calcite aggregates (Fig. S1b). The samples from the  
108 carbonatite veins occur as aggregates of thin flakes ranging from 50 to 80  $\mu\text{m}$  in length  
109 and less than 10  $\mu\text{m}$  in thickness (Fig. 1b). Molybdenite has been shown to contain 1–3 wt.  
110 % Pb (Stein et al. 1997; Du et al. 2004). Two occurrences of Pb were identified in the  
111 studied samples including sub-micro-scale galena filling in the interstitial space of  
112 molybdenite clusters (Fig. 1c) and nanoscale layered PbS with an epitaxial relationship  
113 with the molybdenite host (Fig. 1d).

114

### 115 **Electron microprobe analysis**

116 The chemical composition of the Pb-bearing molybdenite samples was  
117 determined using a JEOL JXA-8230 electron probe microanalyzer (EPMA) under the

118 operating conditions of 20-kV accelerating voltage and 50-nA probe current using a  
119 1- $\mu\text{m}$  beam. The counting times were 20 s for Mo, S, and Fe, 40 s for Ca and Cu, and  
120 60 s for Re, Pb, Zr, As, and W. The analytical lines were  $\text{Re}M_{\alpha}$ ,  $\text{WM}_{\alpha}$ ,  $\text{Pb}L_{\alpha}$ ,  $\text{Mo}L_{\alpha}$ ,  
121  $\text{Zr}L_{\alpha}$ ,  $\text{As}L_{\alpha}$ ,  $\text{Ca}K_{\alpha}$ ,  $\text{Fe}K_{\alpha}$ ,  $\text{Cu}K_{\alpha}$ , and  $\text{SK}_{\alpha}$ . The standards for calibration included  $\text{MoS}_2$   
122 (Mo, S), rhenium (Re), copper (Cu), and tungsten (W) metals, galena (Pb), pyrite (Fe),  
123 zircon (Zr), Cr-diopside (Ca), and arsenopyrite (As).

124

### 125 **Micro-X-ray diffraction analysis**

126 Well-crystallized molybdenite flakes were embedded in epoxy resin and sliced  
127 into  $\sim 50\text{-nm}$ -thick foils using a Leica EM UC7 ultramicrotome equipped with a  
128 diamond knife. The cutting was performed perpendicular to the (001) plane of the  
129 molybdenite flakes and the obtained foils were put on carbon-coated Cu/Ni grids for  
130 TEM analysis. Micro-X-ray diffraction (micro-XRD) analysis was carried out on the  
131 exposed area of the sliced epoxy resin (Fig. 2a) using a Rigaku DMAX rapid  
132 diffraction system (Mo- $K\alpha$ ) at 50 kV and 90 mA. Micro-XRD patterns were collected  
133 using a 30- $\mu\text{m}$  X-ray beam collimator. The stage oscillated about the  $\omega$ -axis ( $2\theta = 2\text{--}$   
134  $35^\circ$ ) and spun about the  $\phi$ -axis to produce a maximum beam footprint of 90  $\mu\text{m}$  on the  
135 sample.

136

### 137 **Transmission electron microscopy analysis**

138 The  $\sim 50\text{-nm}$ -thick molybdenite slices were further analyzed by TEM after  
139 plasma cleaning. HRTEM and scanning transmission electron microscopy (STEM)  
140 analyses were carried out at 200-kV voltage using an FEI Talos F200S field-emission  
141 transmission electron microscope coupled with two SuperX high-resolution  
142 energy-dispersive spectrometer (EDS) detectors and a high-angle annular dark-field

143 (HAADF) detector. The spatial resolution of the STEM mode was  $\sim 0.16$  nm.  
144 Z-contrast images were acquired using a camera length of 160 mm to maximize the  
145 contrast variations between the different atoms in the HAADF images (Xu et al.  
146 2014b, 2014c, 2015).

147

## 148 **RESULTS**

### 149 **Phase identification and quantitative analysis**

150 The molybdenite polytypes were identified by referring to theoretical mixtures of  
151  $2H_1$  and  $3R$  polytypes calculated by Wickman and Smith (1970). In Fig. 2a, the sharp  
152 reflections at  $16.31^\circ$  (2.50 Å),  $17.93^\circ$  (2.27 Å),  $22.36^\circ$  (1.83 Å), and  $24.98^\circ$  (1.64 Å)  
153 in the micro-XRD pattern are attributed to the (102), (103), (105), and (106) planes of  
154 the  $2H_1$ -molybdenite (JCPDS card No. 73-1508 Mo-K $\alpha$ ), respectively. The broad  
155 peaks at  $17.41^\circ$  (2.34 Å),  $18.61^\circ$  (2.19 Å),  $21.64^\circ$  (1.89 Å), and  $23.32^\circ$  (1.76 Å)  
156 correspond to the (014), (015), (017), and (018) reflections of the  $3R$ -molybdenite  
157 (JCPDS card No. 75-0034 Mo-K $\alpha$ ), respectively. The unindexed reflections, e.g.,  
158  $12.16^\circ$  (3.35 Å) (Fig. 2a), are likely derived from the layered PbS precipitates that  
159 occur as a separate phase in the molybdenite. The relative intensities of the (103) and  
160 (105) reflections for the  $2H_1$  polytype and the (014), (015), (017), and (018)  
161 reflections for the  $3R$  polytype were chosen for quantitative analysis (Wickman and  
162 Smith 1970). The results suggest the co-existence of  $2H_1$  and  $3R$  polytypes in the  
163 investigated molybdenite with a calculated ratio of  $2H_1/3R \approx 9:1$ .

164

### 165 **Chemical composition of the Pb-bearing molybdenite**

166 Two occurrences of the Pb-bearing phase were identified in the backscattered  
167 electron (BSE) images (Figs. 1c, 1d). The galena inclusions are several micrometers

168 in length, >100 nm in thickness, and coexist with the molybdenite grains (Figs. 1c,  
169 S1d). The nano-scale Pb-bearing zones occur randomly within the molybdenite grains  
170 (Figs. 1d, S1f).

171 The EPMA results indicate a heterogeneous distribution of Pb in the molybdenite  
172 samples (Fig. 2b), reaching 0.59–1.33 wt.% in the Pb-bearing zones (Table 1). Owing  
173 to the large EPMA beam diameter (>1  $\mu\text{m}$ ), which is substantially larger than the  
174 nanoscale Pb-bearing zones, the Pb concentrations in these areas may be even higher  
175 (Figs.1d, S2). The Pb-poor molybdenite (Table 1) has close to stoichiometric Mo and  
176 S concentrations (60 wt.% Mo and 40 wt.% S). The EPMA mapping reveals that the  
177 Pb-bearing zones are commonly located in the cores of some molybdenite grains (Fig.  
178 2b). The HAADF-STEM images reveal ~1-nm-wide bright bands in the Pb-bearing  
179 zones (open triangles in Fig. 2c). The elemental maps using the Pb L-line indicate that  
180 the Pb enrichment coincides with the bright bands in the host molybdenite (right side  
181 of Fig. 2c). The X-ray EDS spectra show that the bright bands are primarily composed  
182 of Mo, S, and Pb (Fig. 2d). The bright bands have a higher peak Pb intensity than the  
183 neighboring host molybdenite (Figs. S2i, S2j), which also indicates the higher Pb  
184 contents (Fig. 2d). Because the concentrations of other elements (e.g., O, C, and W)  
185 are negligible, the Pb-rich band is likely Pb-sulfide, Pb-bearing molybdenite, or  
186 metallic Pb (Figs. 2, S2).

187

### 188 **Nano-scale features of layered PbS**

189 Z-contrast images with atomic resolution were obtained to identify the mineral  
190 phases in the Pb-rich bands (Fig. 3a). The bright spots in the Z-contrast image (Fig. 3a)  
191 correspond to the positions of the atom columns along the beam direction. The plots  
192 of molybdenite ( $\text{MoS}_2$ ) and PbS along the [010] zone axis are overlain on the



193 Z-contrast image. The intensity profile along the *c*-direction of the Z-contrast image  
194 shows that eight relatively weak Mo + S peaks are separated by two stronger Pb + S  
195 peaks (Fig. 3b). These features suggest that the sandwich-type MoS<sub>2</sub> sheets are  
196 separated by layered PbS sheets (Fig. 3a). The distances between the weak and strong  
197 peaks are ~0.62 nm and ~0.30 nm, corresponding to the fringe spacings along the  
198 *c*-axis of the molybdenite and layered PbS, respectively. The layered PbS has a  
199 similar fringe spacing to that of galena (i.e., ~0.30 nm), and can therefore be regarded  
200 as a distorted galena structure.

201

## 202 **Crystallographic characteristics of the Pb-bearing molybdenite**

203 The TEM image of the Pb-poor zone shows a smooth surface (Fig. 4a). The  
204 sharp and bright diffraction spots without specific extinction in the selected-area  
205 electron diffraction (SAED) pattern indicate well-crystallized 2H<sub>1</sub>-polytype  
206 molybdenite. In contrast, the surface of the Pb-bearing region (black open triangles in  
207 Figs. 4a and 4b) shows non-periodic features/contrast and its SAED pattern includes  
208 streaks (white arrows in Fig. 4a) between the *00l* and *h0l* diffraction spots. These  
209 streaks reflect highly disordered stacking sequences along the *c*-axis. The nano-beam  
210 electron diffraction (NBED) patterns further confirm that the Pb-poor area in the  
211 molybdenite belongs to an ordered 2H<sub>1</sub> polytype and the Pb-rich area has disordered  
212 layer stacking. A 3R polytype is also identified in the vicinity of the Pb-rich area (Fig.  
213 5). Nearly all of the investigated samples contain coexisting ordered (Figs. 4c, 4d) and  
214 disordered layer stacking domains (Figs. 4c, 4e), corresponding to the Pb-poor and  
215 Pb-bearing zones, respectively. The streaking diffraction observed in the SAED/FFT  
216 (Fast Fourier Transform) patterns of the Pb-bearing molybdenite reveals that the 3R  
217 polytype and disordered molybdenite usually occur in the same region (Figs. 4, 5),

218 corresponding to the broad reflections attributed to the 3R polytype in the micro-XRD  
219 pattern (Fig. 2a).

220 The HAADF-STEM image (Fig. 6a) shows 2, 4, and 6 layers of PbS pairs along  
221 the *c*-axis in the host molybdenite, corresponding to 1, 2, and 3 unit cells of layered  
222 PbS, respectively. In the STEM images, the [S-Mo-S] layers are completely (Fig. 6b,  
223 top) or partially (Fig. 6b, bottom) replaced by PbS. Pb-bearing domains also occur  
224 neighboring the Pb-poor domains and/or layered PbS in the host molybdenite (Fig.  
225 6c), corresponding to the broad light band in the low-magnification STEM image (Fig.  
226 7a). The enlarged area of Fig. 6b displays an intermediate state between the  
227 Pb-bearing molybdenite and 2-layer PbS (Fig. 6d). The mismatched structure results  
228 in the formation of crystal defects (e.g., dislocations) surrounding the layered PbS  
229 (open triangles in Fig. 6b). Pb was also observed to significantly diffuse under  
230 prolonged electron beam irradiation, further scavenging Pb from the host molybdenite  
231 into the layered precipitates (Fig. 7), as discussed in the following section.

232

233

## DISCUSSION

### 234 **Formation conditions for Pb-bearing phases in molybdenite**

235 The sub-micro-scale galena commonly occurs as >100-nm inclusions in the  
236 interstices of molybdenite flakes and does not show an orientation relationship with  
237 the neighboring molybdenite (Figs. 1c, S1d). This galena probably formed by direct  
238 precipitation from a hydrothermal fluid in the interstitial spaces of the growing  
239 molybdenite, in accordance with previous studies (Ciobanu et al. 2013; Pasava et al.  
240 2016). The HAADF-STEM images indicate that Pb ions can be incorporated into the  
241 molybdenite structure, as shown by the high-contrast regions in Fig. 6c. The  
242 incorporation of impurities in a growing mineral depends on the thermodynamic

243 conditions (e.g., temperature and pressure), crystallization rate, and similarity of the  
244 ionic properties (e.g., radius and charge) between the component ions and impurities  
245 (see Sunagawa 2005 and references therein). The prominently different ionic radii of  
246 Pb ( $\text{Pb}^{2+} = 1.29 \text{ \AA}$ ,  $\text{Pb}^{4+} = 0.98 \text{ \AA}$ ), S ( $\text{S}^{2-} = 1.84 \text{ \AA}$ ), and Mo ions ( $\text{Mo}^{4+} = 0.65 \text{ \AA}$ )  
247 suggest that Pb incorporation causes lattice compression or expansion and disordered  
248 stacking in molybdenite (Fig. 4e). On the other hand, impurities are more likely to be  
249 incorporated under relatively high temperatures and rapid crystallization rates. High  
250 crystallization temperatures have been shown to provide an additional driving force  
251 for Pb incorporation into the molybdenite structure (c.f., Sunagawa 2005 and  
252 references therein).

253 The Pb-bearing molybdenite aggregates are commonly infilled with calcite (Fig.  
254 1b) or closely coexist with calcite aggregates in the carbonatite dykes (Fig. S1c). The  
255 paragenesis suggests that the molybdenite formed simultaneously with or earlier than  
256 the calcite. Given that carbonatite dykes generally have high water contents (~15%),  
257 the melts that form from carbonatite dykes tend to have relatively low viscosities  
258 (Keppler 2003). In the Huanglongpu deposit, this low viscosity enabled the  
259 carbonatite melt to rapidly ascend to the near-surface (Xu et al. 2014a). The  
260 Pb-bearing molybdenite thus likely formed before or during melt ascent and rapidly  
261 cooled within the host carbonatite melt (Fig. 1), which is in accordance with the  
262 generally smaller grain size compared with porphyry-skarn hydrothermal systems in  
263 the East Qinling orogenic belt (Li and Pirajno 2017). In the present study, most of the  
264 Pb-bearing molybdenite occurs as cores within the molybdenite grains (Figs. 1d, 2b),  
265 suggesting that Pb was incorporated into the molybdenite at relatively high  
266 temperatures during initial crystallization. The high cooling rate also inhibited Pb  
267 migration out of molybdenite, resulting in a local enrichment of Pb up to several

268 weight percent. The relatively high crystallization temperature and rapid  
269 crystallization rate therefore provide a thermodynamic advantage for the formation of  
270 Pb-bearing molybdenite.

271

## 272 **Genesis of layered PbS in molybdenite**

273 Previous studies have suggested that layered PbS intergrown with coherent  
274 Bi-chalcogenides can form mixed-layer configurations (Cook et al. 2007a, 2007b;  
275 Ciobanu et al. 2009a). However, the layered PbS observed in the present study most  
276 likely formed by the exsolution of Pb from the Pb-bearing molybdenite. The STEM  
277 observations reveal that Pb diffusion occurs under electron beam irradiation with the  
278 gradual disappearance of the Pb-bearing region and formation of layered PbS in the  
279 molybdenite (Figs. 6c, 7). The observed 2–6 atomic-layer-thick PbS within the  
280 Pb-poor molybdenite (Figs. 6, S2) also suggests that the layered PbS likely formed by  
281 Pb diffusion during post-crystallization processes.

282 Minerals tend to exclude incompatible elements from their crystal lattice to  
283 minimize the Gibbs free energy in the mineral system during crystallization. As a  
284 result, incompatible elements generally occur in limited concentrations within the  
285 lattice of their host minerals (c.f., Sunagawa 2005 and references therein). In the  
286 present study, the dislocations (Fig. 6b) and disordered stacking (Fig. 4) induced by  
287 Pb incorporation significantly lowered the thermodynamical stability of the  
288 Pb-bearing molybdenite. As a result, the metastable Pb-bearing molybdenite solid  
289 solution (Fig. 6c) tended to decompose into layered PbS and Pb-poor molybdenite  
290 (Fig. 6d) by migration and local Pb enrichment (Fig. 7) after crystallization. Due to  
291 the anisotropy of the layered molybdenite structure, the linear thermal expansion  
292 coefficient of the *c*-direction is greater than that of the *a*- and *b*-directions

293 (El-Mahalawy and Evans 1976), which may affect the diffusion rates of Pb and result  
294 in the formation of layered-PbS precipitates along the (001) plane rather than the (*hk*0)  
295 planes. The exsolution of Pb along the (001) plane leads to a transformation of the  
296 metastable Pb-bearing S-Mo-S layers to layered PbS with a two-atomic-layer  
297 thickness (Fig. 6d).

298 The chemical composition and crystal structure results indicate that the 2–6  
299 layers of Pb-S pairs are very similar to those of galena. Due to the spatial constraint of  
300 the S-Mo-S sandwich structure, the layered PbS structure is significantly distorted  
301 compared with the standard galena. The Pb-S distance ( $\sim 3.0$  Å) of the layered PbS is  
302 similar to the S-Mo-S height ( $\sim 3.1$  Å), which results in an epitaxial relationship  
303 between the exsolved layered PbS and molybdenite matrix. These mixed-layer  
304 compounds of PbS and MoS<sub>2</sub> exhibit a combination of a pseudo-tetragonal (*Q*) with a  
305 pseudo-hexagonal (*H*) layer (Jaszczak et al. 2016). The *Q* (PbS-type) layer is a (100)  
306 slab of the PbS/NaCl archetype,  $2n$  ( $n \geq 1$ ) atomsthick, similar to previously reported  
307 complex cylindrite group crystals [(MX)<sub>1+y</sub>]<sub>m</sub>(TX<sub>2</sub>)<sub>n</sub> (with M = Sn, Pb, Sb, and Bi, T  
308 = Ti, V, Nb, and Mo, and X = S and Se) (Williams and Hyde 1988; Lin et al. 2008;  
309 Makovicky et al. 2008; Garbrecht et al. 2011; Jaszczak et al. 2016).

310 Based on the above discussions, the layered PbS in molybdenite was proposed to  
311 forms in the following scenario (Fig. 8). Up to several weight percent of Pb was  
312 incorporated into the structure of molybdenite during the initial stage of  
313 crystallization (Fig. 8a). The incorporation of Pb atoms caused lattice expansion and  
314 screw dislocations (Fig. 6b), and further led to 3R/disordered stacking (Figs. 4, 5). Pb  
315 migration and enrichment at the grain boundaries or defects led to the formation of  
316 local Pb-bearing and Pb-poor molybdenite domains (Fig. 8b). The exsolution of Pb  
317 from the sandwiched Pb-bearing molybdenite led to the formation of layered PbS

318 along the (001) plane of molybdenite (Fig. 8c). Due to space restrictions of the  
319 molybdenite unit cell, the newly formed PbS shows a periodic layered structure with  
320 two atomic layers along the *c*-direction (Fig. 8c). The host molybdenite also  
321 transformed from a Pb-bearing  $3R$ /disordered structure to a Pb-poor ordered  $2H_1$   
322 polytype. The incorporation of Pb and formation of nano-scale layered PbS by  
323 post-crystallization exsolution within molybdenite therefore led to the heterogeneous  
324 distribution of Pb in the molybdenite (Figs. 1d, 2b).

325

326

### IMPLICATIONS

327 Trace elements in molybdenite can provide important constraints on the  
328 evolution processes and dating of a wide range of geological events. The behavior of  
329 incompatible trace elements tends to be more complicated than that of compatible  
330 trace elements in molybdenite. A better understanding of the behavior of incompatible  
331 trace elements is therefore of high importance for interpreting the related geochemical  
332 data. Recent nanoSIMS evidence has demonstrated significant Pb trapping by  
333 molybdenite (Rollog et al. 2019, 2020). The present study also shows a heterogeneous  
334 distribution of typically incompatible Pb in natural molybdenite. These findings  
335 provide important insight into the behavior of other incompatible elements within host  
336 minerals to better interpret the related geological data. For example, the reliability of  
337  $^{187}\text{Re}$ - $^{187}\text{Os}$  isotopic data for molybdenite is based on the assumption that Re and Os  
338 are immobile and heterogeneously distributed within molybdenite (Stein et al. 2001,  
339 2003). Previous studies suggested that incompatible radiogenic  $^{187}\text{Os}$  is likely to be  
340 more mobile and stored in the micro-defects of molybdenite compared with parent  
341  $^{187}\text{Re}$

342 Re and may therefore become decoupled over time on the micrometer scale (Stein  
et al. 2001, 2003). This provides an explanation of why in-situ ICP-MS spot dating of

343 molybdenite consistently fails (Stein et al. 2003). The behavior of Pb in natural  
344 molybdenite can thus serve as a proxy for understanding the heterogeneous  
345 distribution and mobility of Os and other relatively incompatible trace elements in  
346 molybdenite.

347 Franckeite is a naturally occurring SnS<sub>2</sub>/PbS van der Waals heterostructure and is  
348 a promising material for emerging energy storage technology and near-infrared  
349 photodetectors (Molina-Mendoza et al. 2017; Vělický et al. 2017). Synthetic MoS<sub>2</sub>/  
350 PbS van der Waals heterostructures have been used as infrared memory devices  
351 (Wang et al. 2018). Natural Pb-bearing molybdenite (MoS<sub>2</sub>)/layered-PbS van der  
352 Waals heterostructures may therefore also have relevant properties that can serve as a  
353 natural model for the synthesis of layered MoS<sub>2</sub>/PbS heterostructures.

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363

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## REFERENCES

- 365 Barra, F., Deditius, A., Reich, M., Kilburn, M.R., Guagliardo, P., and Roberts, M.P.  
366 (2017) Dissecting the Re-Os molybdenite geochronometer. *Scientific Reports*, 7,  
367 16054.
- 368 Ciobanu, C.L., Pring, A., Cook, N.J., Self, P., Jefferson, D., Dima, G.I., and Melnikov,  
369 V. (2009) Chemical-structural modularity in the tetradymite group: A HRTEM  
370 study. *American Mineralogist*, 94, 517–534.
- 371 Ciobanu, C.L., Cook, N.J., Kelson, C.R., Guerin, R., Kalleske, N., and Danyushevsky,  
372 L. (2013) Trace element heterogeneity in molybdenite fingerprints stages of  
373 mineralization. *Chemical Geology*, 347, 175–189.
- 374 Cook, N.J., Ciobanu, C.L., Stanley, C.J., Paar, W.H., and Sundblad, K. (2007a)  
375 Compositional data for Pb-Bi tellurosulphides. *The Canadian Mineralogist*, 45,  
376 417–435.
- 377 Cook, N.J., Ciobanu, C.L., Wagner, T., and Stanley, C.J. (2007b) Minerals of the  
378 system Bi-Te-Se-S related to the tetradymite archetype: Review of classification  
379 and compositional variation. *The Canadian Mineralogist*, 45, 665–708.
- 380 Du, A., Wu, S., Sun, D., Wang, S., Qu, W., Markey, R., Stain, H., Morgan, J., and  
381 Malinovskiy, D. (2004) Preparation and Certification of Re-Os Dating Reference



- 382 Materials: Molybdenites HLP and JDC. Geostandards and Geoanalytical  
383 Research, 28, 41–52.
- 384 El-Mahalawy, S., and Evans, B. (1976) The thermal expansion of  $2H\text{-MoS}_2$ ,  
385  $2H\text{-MoSe}_2$  and  $2H\text{-WSe}_2$  between 20 and 800 °C . Journal of Applied  
386 Crystallography, 9, 403-406.
- 387 Fleischer, M. (1959) The geochemistry of rhenium, with special reference to its  
388 occurrence in molybdenite. Economic Geology, 54, 1406–1413.
- 389 Frondel, J.W., and Wickman, F.E. (1970) Molybdenite polytypes in theory and  
390 occurrence .II. Some naturally-occurring polytypes of molybdenite. American  
391 Mineralogist, 55, 1857-1875.
- 392 Grabezhev, A.I., and Voudouris, P.C. (2014) Rhenium distribution in molybdenite  
393 from the Vosnesensk porphyry Cu ± (Mo,Au) deposit (Southern Urals, Russia).  
394 The Canadian Mineralogist, 52, 671–686.
- 395 Garbrecht, M., Spiecker, E., Tillmann, K., and Jager, W. (2011) Quantitative atom  
396 column position analysis at the incommensurate interfaces of a  $(\text{PbS})_{1.14}\text{NbS}_2$   
397 misfit layered compound with aberration-corrected HRTEM. Ultramicroscopy,  
398 111, 245-250.
- 399 Huang D.H, Wang Y.C, Nie F.J, and Jiang X.J. (1985) A new type of molybdenum  
400 deposit: Geologic characteristics and metallogenic mechanism of Huanglongpu

- 401 carbonatite vein-type of molybdenum (lead) deposit, Shaanxi. *Acta Geologica*  
402 *Sinica*, 59, 241-257. (in Chinese with English abstract)
- 403 Jaszczak, J.A., Rumsey, M.S., Bindi, L., Hackney, S.A., Wise, M.A., Stanley, C.J.,  
404 and Spratt, J.(2016) Merelaniite,  $\text{Mo}_4\text{Pb}_4\text{VSbS}_{15}$ , a new molybdenum-essential  
405 member of the cylindrite group, from the merelani tanzanite deposit, Irelatema  
406 mountains, Manyara region, Tanzania. *Minerals*, 6, 115.
- 407 Keppler, H. (2003) Water solubility in carbonatite melts. *American Mineralogist*, 88,  
408 1822-1824.
- 409 Li, N. Pirajno, F. (2017) Early Mesozoic Mo mineralization in the Qinling Orogen:  
410 An overview. *Ore Geology Reviews*, 81, 431-450.
- 411 Lin, Q., Heideman, C.L., Nguyen, N., Zschack, P., Chiritiescu, C., Cahill, D.G., and  
412 Johnson, D.C. (2008) Designed synthesis of families of misfit-layered  
413 compounds. *European Journal of Inorganic Chemistry*, 15, 2382-2385.
- 414 Lin, Y.C., Dumcenccon, D.O., Huang, Y.S., and Suenaga, K. (2014a) Atomic  
415 mechanism of the semiconducting-to-metallic phase transition in single-layered  
416  $\text{MoS}_2$ . *Nature Nanotechnology*, 9, 391-396.
- 417 Lin, Y.C., Dumcenco, D.O., Komsa, H.P., Niimi, Y., Krasheninnikov, A.V., Huang,  
418 Y.S., and Suenaga, K. (2014b) Properties of individual dopant atoms in

- 419 single-layer MoS<sub>2</sub>: atomic structure, migration, and enhanced reactivity.  
420 *Advanced Materials*, 26, 2857-2861.
- 421 Makovicky, E., Petricek, V., Dusek, M., and Topa, D. (2008) Crystal structure of a  
422 synthetic tin-selenium representative of the cylindrite structure type. *American*  
423 *Mineralogist*, 93, 1787-1798.
- 424 McCandless, T.E., Ruiz, J., and Campbell, A.R. (1993) Rhenium behavior in  
425 molybdenite in hypogene and near-surface environments: Implications for Re-Os  
426 geochronometry. *Geochimica et Cosmochimica Acta*, 57, 889-905.
- 427 Molina-Mendoza, A.J., Giovanelli, E., Paz, W.S., Nino, M.A., Island, J.O., Evangelini,  
428 C., Aballe, L., Foerster, M., van der Zant, H.S.J., Rubio-Bollinger, G., Agrait, N.,  
429 Palacios, J.J., Perez, E.M., and Castellanos-Gomez, A. (2017) Franckeite as a  
430 naturally occurring van der Waals heterostructure. *Nature Communications*, 8,  
431 14409.
- 432 McCandless, T.E., Ruiz, J., and Campbell, A.R. (1993) Rhenium behavior in  
433 molybdenite in hypogene and near-surface environments: Implications for Re-Os  
434 geochronometry. *Geochimica et Cosmochimica Acta*, 57, 889–905.
- 435 Newberry, R.J.J. (1979a) Polytypism in molybdenite (I): a non-equilibrium  
436 impurity-induced phenomenon. *American Mineralogist*, 64, 758–769.

- 437 Newberry, R.J.J. (1979b) Polytypism molybdenite (II): relationships between  
438 polytypism, ore deposition/alteration stages and rhenium contents. American  
439 Mineralogist, 64, 768–775.
- 440 Pašava, J., Svojtka, M., Veselovský, F., Ďurišová, J., Ackerman, L., Pour, O., Drábek,  
441 M., Halodová, P., and Haluzová, E. (2016) Laser ablation ICPMS study of trace  
442 element chemistry in molybdenite coupled with scanning electron microscopy  
443 (SEM) — An important tool for identification of different types of  
444 mineralization. Ore Geology Reviews, 72, 874–895.
- 445 Pennycook, S.J. (2002) Structure determination through Z-contrast microscopy.  
446 Advances in Imaging Electron Physics, 123, 173-206.
- 447 Petruk, W. (1965) Mineralogy of the Mount Pleasant tin deposit, New Brunswick. In  
448 Canada Mines Branch Tech, pp. 1–35.
- 449 Plotinskaya, O.Y., Abramova, V.D., Groznova, E.O., Tesselina, S.G., Seltmann, R.,  
450 and Spratt, J. (2018) Trace-element geochemistry of molybdenite from porphyry  
451 Cu deposits of the Birgilda-Tomino ore cluster (South Urals, Russia).  
452 Mineralogical Magazine, 82, S281-S306.
- 453 Plotinskaya, O. Y., Shilovskikh, V. V., Najorka, J., Kovalchuk, E. V., Seltmann, R.,  
454 and Spratt, J. (2019) Grain-scale distribution of molybdenite polytypes versus  
455 rhenium contents:  $\mu$ XRD and EBSD data. Mineralogical Magazine, 83, 639-644.

- 456 Rollog, M., Cook, N.J., Guagliardo, P., Ehrig, K., Ciobanu, C.L., and Kilburn, M.  
457 (2019) Detection of Trace Elements/Isotopes in Olympic Dam Copper  
458 Concentrates by nanoSIMS. *Minerals*, 9, 336.
- 459 Rollog, M., Cook, N.J., Guagliardo, P., Ehrig, K., and Kilburn, M. (2020)  
460 Radionuclide distributions in Olympic Dam copper concentrates: The  
461 significance of minor hosts, incorporation mechanisms, and the role of mineral  
462 surfaces. *Minerals Engineering*, 148, 106176.
- 463 Song, W., Xu, C., Qi, L., Zhou, L., Wang, L., and Kynicky, J. (2015) Genesis of  
464 Si-rich carbonatites in Huanglongpu Mo deposit, Lesser Qinling orogen, China  
465 and significance for Mo mineralization. *Ore Geology Reviews*, 64, 756-765.
- 466 Song, W., Xu, C., Smith, M.P., Kynicky, J., Huang, K., Wei, C., Zhou, L., and Shu, Q.  
467 (2016) Origin of unusual HREE-Mo-rich carbonatites in the Qinling orogen,  
468 China. *Scientific Reports*, 6, 37377
- 469 Stein, H., Scherstén, A., Hannah, J., and Markey, R. (2003) Subgrain-scale  
470 decoupling of Re and <sup>187</sup>O<sub>s</sub> and assessment of laser ablation ICP-MS spot  
471 dating in molybdenite. *Geochimica et Cosmochimica Acta*, 67, 3673–3686.
- 472 Stein, H.J., Markey, R.J., Morgan, J.W., Du, A., and Sun, Y. (1997) Highly precise  
473 and accurate Re-Os ages for molybdenite from the East Qinling molybdenum  
474 belt, Shaanxi Province, China. *Economic Geology*, 92, 827–835.

- 475 Stein, H.J., Markey, R.J., Morgan, J.W., Hannah, J.L., and Schersten, A. (2001) The  
476 remarkable Re-Os chronometer in molybdenite: how and why it works. *Terra*  
477 *Nova*, 13, 479–486.
- 478 Sunagaw, I. (2005) *Crystals: growth, morphology, and perfection*, 308 p. Cambridge  
479 University Press, UK.
- 480 Terada, K., Osaki, S., Ishihara, S., and Kiba, T. (1971) Distribution of rhenium in  
481 molybdenites from Japan. *Geochemical Journal*, 4, 123–141.
- 482 Vělický, M., Toth, P.S., Rakowski, A.M., Rooney, A.P., Kozikov, A., Woods, C.R.,  
483 Mishchenko, A., Fumagalli, L., Yin, J., Zolyomi, V., Georgiou, T., Haigh, S.J.,  
484 Novoselov, K.S., and Dryfe, R.A.W. (2017) Exfoliation of natural van der Waals  
485 heterostructures to a single unit cell thickness. *Nature Communications*, 8,  
486 14410.
- 487 Voudouris, P.C., Melfos, V., Spry, P.G., Bindi, L., Kartal, T., Arikas, K., Moritz, R.,  
488 and Ortelli, M. (2009b) Rhenium-rich molybdenite and rheniite in the Pagoni  
489 Rachi Mo-Cu-Te-Ag-Au prospect, Northern Greece: Implications for the Re  
490 geochemistry of porphyry-style Cu-Mo and Mo mineralization. *The Canadian*  
491 *Mineralogist*, 47, 1013-1036.
- 492 Wang, Q.S., Wen, Y., Cai, K.M., Cheng, R.Q., Yin, L., Zhang, Y., Li, J., Wang, Z.X.,  
493 Wang, F., Wang, F.M., Shifa, T.A., Jiang, C., Yang, H., and He, J. (2018)

- 494 Nonvolatile infrared memory in MoS<sub>2</sub>/PbS van der Waals heterostructures.  
495 Science Advances, 4, eaap7916.
- 496 Williams, T.B., and Hyde, B.G. (1988) Electron-Microscopy of Cylindrite and  
497 Franckeite. Physics and Chemistry of Minerals, 15, 521-544.
- 498 Xu, C., Chakhmouradian, A.R., Taylor, R.N., Kynicky, J., Li, W.B., Song, W.L., and  
499 Fletcher, I.R. (2014a) Origin of carbonatites in the South Qinling orogen:  
500 Implications for crustal recycling and timing of collision between the South and  
501 North China Blocks. Geochimica et Cosmochimica Acta, 143, 189-206.
- 502 Xu, C., Song, W., Qi, L., and Wang, L. (2009) Geochemical characteristics and  
503 tectonic setting of ore-bearing carbonatites in Hunglongpu Mo ore field. Acta  
504 Petrologica Sinica, 25, 422-430. (in Chinese with English abstract)
- 505 Xu, H., Shen, Z., Konishi, H., and Luo, G. (2014b) Crystal structure of  
506 Guinier-Preston zones in orthopyroxene: Z-contrast imaging and ab initio study.  
507 American Mineralogist, 99, 2043–2048.
- 508 Xu, H., Shen, Z., and Konishi, H. (2014c) Si-magnetite nano-precipitates in silician  
509 magnetite from banded iron formation: Z-contrast imaging and ab initio study.  
510 American Mineralogist, 99, 2196–2202.
- 511 Xu, H., Shen, Z., and Konishi, H. (2015) Natural occurrence of monoclinic Fe<sub>3</sub>S<sub>4</sub>  
512 nano-precipitates in pyrrhotite from the Sudbury ore deposit: a Z-contrast

513 imaging and density functional theory study. *Mineralogical Magazine*, 79, 377–

514 385.

515

516 **Figure Captions:**

517 **FIGURE 1.** Photomicrographs (**a-b**) and backscattered electron (BSE) images (**c-d**)

518 showing the occurrences of molybdenite features in the carbonatite veins. (**a**)

519 Molybdenite aggregates of thin flakes in carbonatite veins. (**b**) Reflected light images

520 of dendritic molybdenite clusters (50–80  $\mu\text{m}$  in diameter) that closely coexist with

521 calcite. (**c**) Micrometer galena inclusion within molybdenite grain. (**d**) Nanoscale

522 Pb-bearing zones within molybdenite grain. Cal, calcite; Gn, galena; Mol,

523 molybdenite.

524

525 **FIGURE 2.** X-ray diffraction pattern and nano-scale features of Pb-bearing

526 molybdenite. (**a**) Micro-X-ray diffraction pattern of the  $2H_1$  and  $3R$  polytype

527 molybdenite. The X-ray diffraction pattern area (epoxy resin with embedded

528 molybdenite) was obtained from an ultrathin section for transmission electron

529 microscopy (schematic shown in upper right). (**b**) Heterogeneous Pb distribution

530 within molybdenite of the inserted area (top left), revealed by X-ray intensity

531 mapping. (**c**) High-angle annular dark-field scanning transmission electron

532 microscopy image shows  $\sim 1\text{-nm}$ -wide bright bands in the Pb-bearing zones of

533 molybdenite, as indicated by the open triangles (left) and the corresponding elemental

534 maps of the Pb-L line of the target Pb-bearing zones (right). (**d**) X-ray

535 energy-dispersive spectrum collected from the Pb-enriched area in (**c**). The Cu signal

536 is from the Cu grid.

537



538 **FIGURE 3.** Atom-scale features of the Pb-bearing zones in molybdenite. **(a)**  
539 Layered-PbS in the Pb-bearing zones in a high-angle annular dark-field scanning  
540 transmission electron microscopy (HAADF-STEM) image. The ball-and-stick model  
541 displays the corresponding structures of molybdenite (MoS<sub>2</sub>) and layered PbS along  
542 the [010] zone axis. The blue, green, and yellow spheres represent Pb, Mo, and S  
543 atoms, respectively. **(b)** Atom intensity distribution of the layered-PbS band and  
544 neighboring molybdenite sheets along the *c*-direction, as indicated by the red line. The  
545 intensity (*I*) of the atom spots is proportional to the atomic number (*Z*) and the  
546 number of atoms (*m*) along the beam direction (i.e.,  $I \sim mZ^n$ ) in the HAADF-STEM  
547 images (Pennycook 2002).

548

549 **FIGURE 4.** Z-contrast images with corresponding transmission electron microscope  
550 (TEM) images and selected-area electron diffraction (SAED) patterns collected from  
551 a Pb-poor zone (white dotted circle) and Pb-bearing zone (white solid circle) along  
552 the [010] zone axis within a single molybdenite crystal. **(a)** Low-magnification  
553 high-angle annular dark-field scanning transmission electron microscopy  
554 (HAADF-STEM), TEM images, and the corresponding SAED patterns of Pb-poor  
555 and Pb-bearing molybdenite. **(b)** HAADF-STEM image of another Pb-bearing  
556 molybdenite particle. **(c)** Enlarged HRTEM image from the white dotted square in **b**.  
557 **(d)** FFT patterns from the Pb-poor (white dotted square in **c**) and **(e)** Pb-bearing zones  
558 (white solid square in **c**). The black open triangles in **(a)** and **(b)** indicate the Pb-rich  
559 band.

560

561 **FIGURE 5.** **(a)** High-angle annular dark-field scanning transmission electron  
562 microscopy (HAADF-STEM) image of Pb-bearing molybdenite and nano-beam

563 electron diffraction (NBED) patterns (1–8) of different areas of the molybdenite. **(b-c)**  
564 Corresponding STEM–energy-dispersive spectroscopy element maps of S and Pb.  
565 The NBED patterns of the selected area (~10 nm per spot) in **(a)** show  
566 2*H*-molybdenite in the Pb-poor domain (1, 2), disordered stacking in the Pb-rich  
567 domain (3–5, 8), and the 3*R*-molybdenite near the Pb-rich band (6, 7).

568

569 **FIGURE 6.** Nano-scale and atomic-scale features of the layered PbS precipitates  
570 distributed along the *c*-axis of the host molybdenite. **(a)** 2, 4, and 6 layers of PbS  
571 distributed in Pb-bearing molybdenite. **(b)** High density of defects around the layered  
572 PbS, as indicated by the triangle arrows. **(c)** High-resolution high-angle annular  
573 dark-field scanning transmission electron microscopy (HAADF-STEM) image  
574 showing layered PbS stacked along the *c*-axis in the Pb-bearing molybdenite. **(d)**  
575 Atomic resolution HAADF-STEM image displaying the transformation of the  
576 S-Mo-S sandwich structure to an NaCl-type layered PbS. The ball-and-stick model  
577 displays the corresponding structures of molybdenite (MoS<sub>2</sub>) and layered PbS along  
578 the [010] zone-axis.

579

580 **FIGURE 7.** High-angle annular dark-field scanning transmission electron microscopy  
581 (HAADF-STEM) images and corresponding energy-dispersive spectroscopy (EDS)  
582 mapping and profile of the Pb-bearing molybdenite. **(a-b)** STEM images show a  
583 relatively broad light band **(a)** and narrow light band **(b)** after 20 min of electron  
584 beam irradiation within the host molybdenite. **(c)** STEM-EDS element maps of Mo  
585 and Pb in **(a)**. **(d)** EDS profile of the Pb-L spectrum shows the diffusion of Pb atoms  
586 during EDS mapping. EDS profiles 1 and 2 were collected from the top down in  
587 images **(a)** and **(b)**, respectively.

588

589 **FIGURE 8.** Schematic diagram showing the transformation of Pb-bearing  
590 molybdenite to layered PbS and Pb-poor molybdenite. The red, green, and yellow  
591 spheres represent Pb, Mo, and S atoms, respectively. The blue dotted circles in the  
592 [S-Mo-S] structure indicate S vacancies.

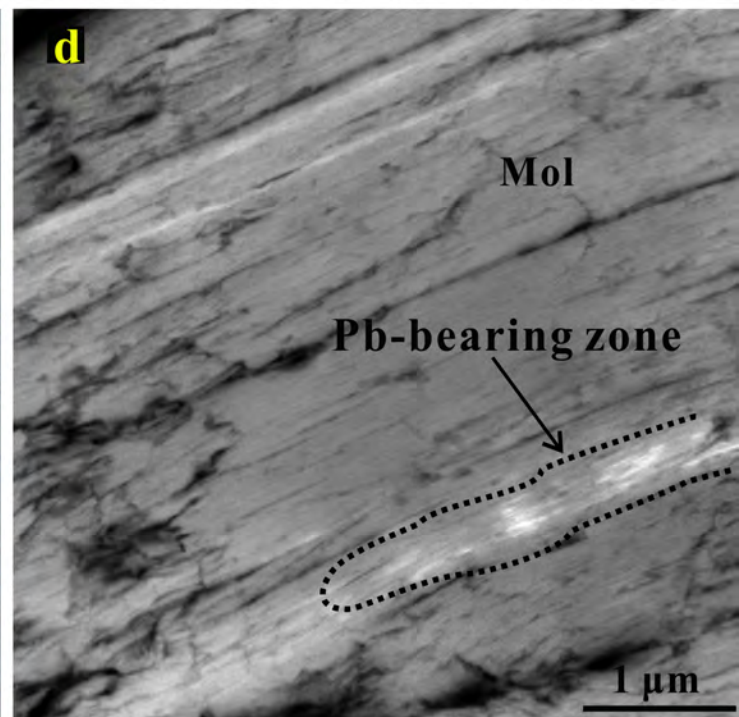
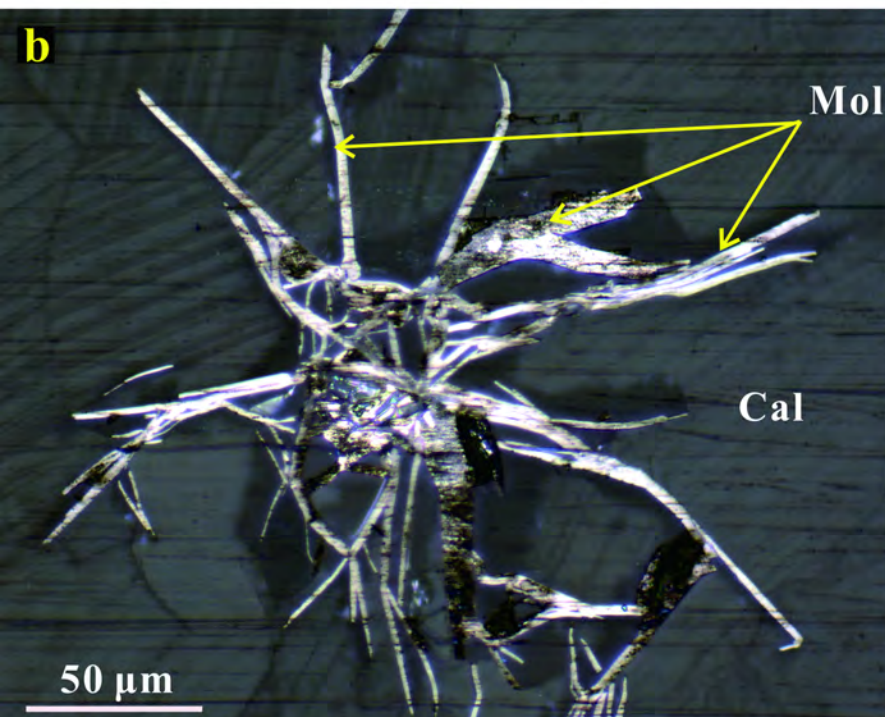
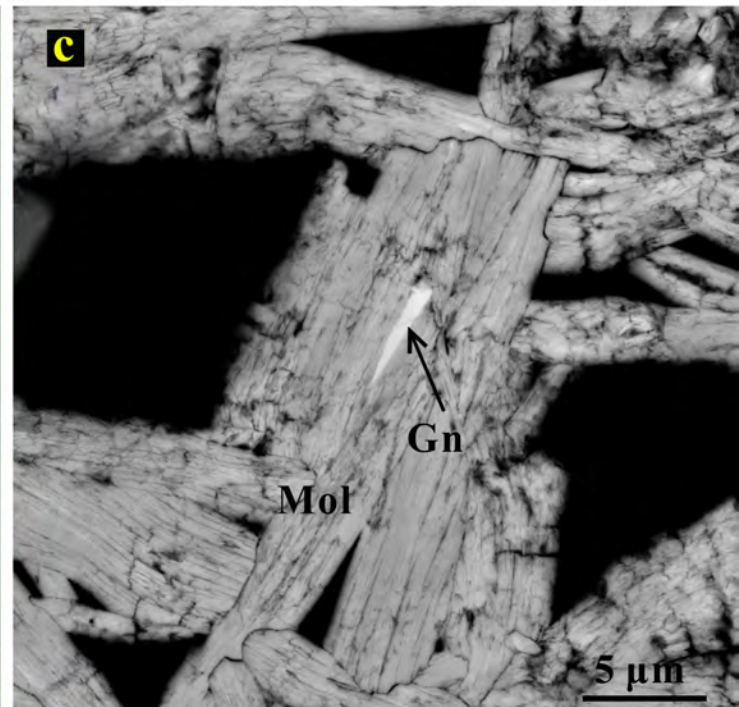
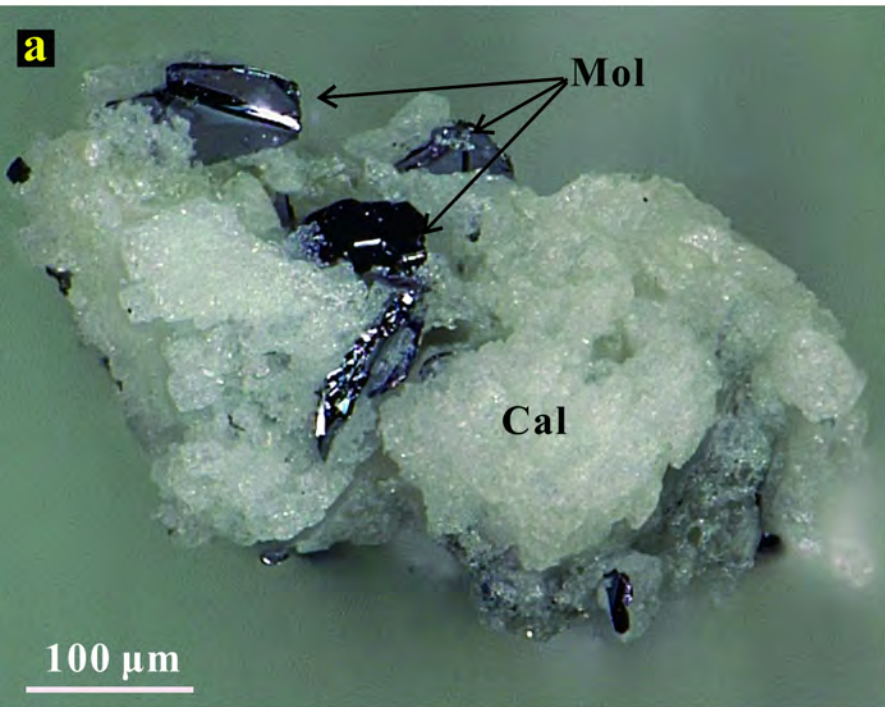
593

**Table 1. Representative composition of major elements of molybdenite from the Huanglongpu Mo-Pb ore deposit based on electron microprobe analysis.**

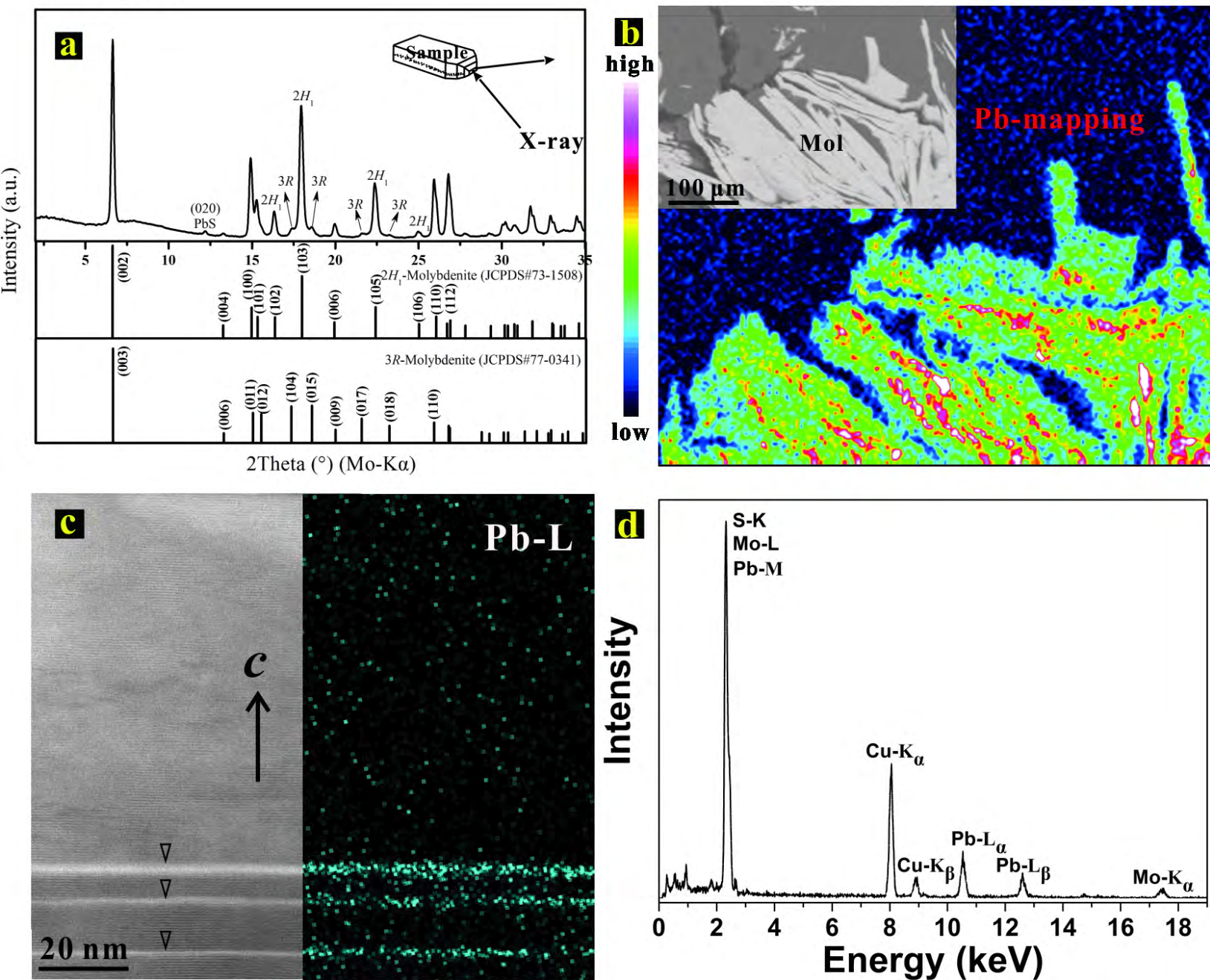
wt. %	Pb-poor zone						Pb-bearing zone			
	1	2	3	4	5	Avg.	1	2	3	Avg.
Mo	59.94	59.92	59.86	60.80	59.62	60.03	58.38	59.01	58.56	58.65
S	40.05	40.03	39.44	40.49	39.55	39.91	38.52	39.12	39.21	38.95
Fe	0.02	0.02	0.15	0.08	0.10	0.07	0.11	0.03	0.05	0.06
Ca	bdl	bdl	bdl	bdl	bdl	bdl	0.03	0.10	0.03	0.05
Re	0.04	0.03	0.05	0.03	0.03	0.04	0.05	0.09	0.06	0.07
Pb	bdl	bdl	bdl	bdl	bdl	bdl	1.33	0.75	0.59	0.89
Total	100.05	100.00	99.50	101.40	99.30	100.05	98.42	99.10	98.50	98.67
<b>Calculated formulas</b>										
Mo	0.9991	0.9991	0.9953	0.9974	0.9969	0.9977	0.9848	0.9885	0.9922	0.9887
S	2.0028	2.0025	1.9673	1.9929	1.9838	1.9898	1.9494	1.9661	1.9930	1.9699
Fe	0.0006	0.0006	0.0043	0.0023	0.0029	0.0020	0.0032	0.0009	0.0015	0.0017
Ca	-	-	-	-	-	-	0.0012	0.0040	0.0012	0.0020
Re	0.0003	0.0003	0.0004	0.0003	0.0003	0.0003	0.0004	0.0008	0.0005	0.0006
Pb	-	-	-	-	-	-	0.0104	0.0058	0.0046	0.0070

Note: Chemical formulas are calculated based on a normalized 100% for the total cations (i.e., Mo, Fe, Re, Ca, and Pb). As, Zr, Cu, and W are below the detection limit (bdl) in all of the spot analyses.

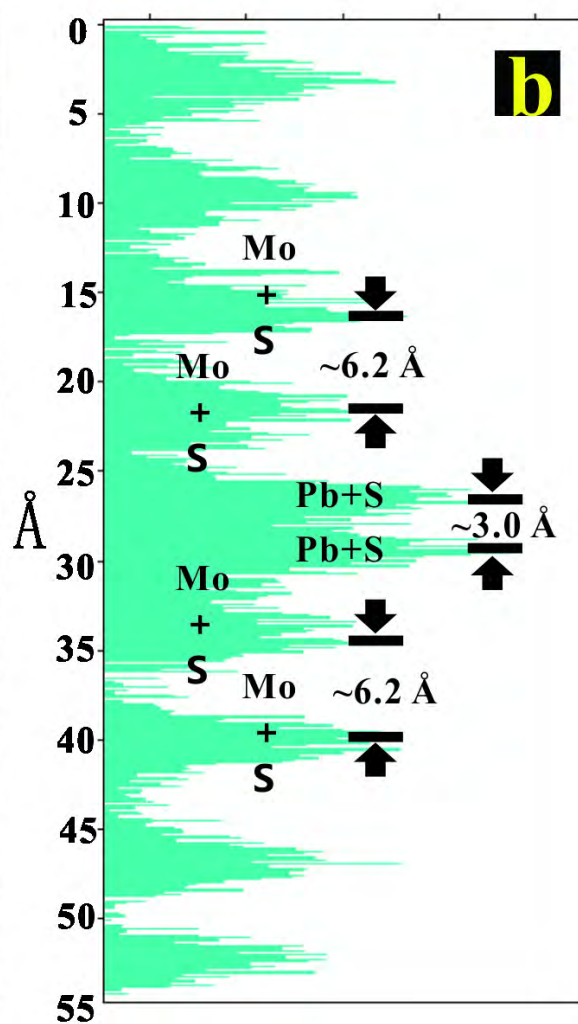
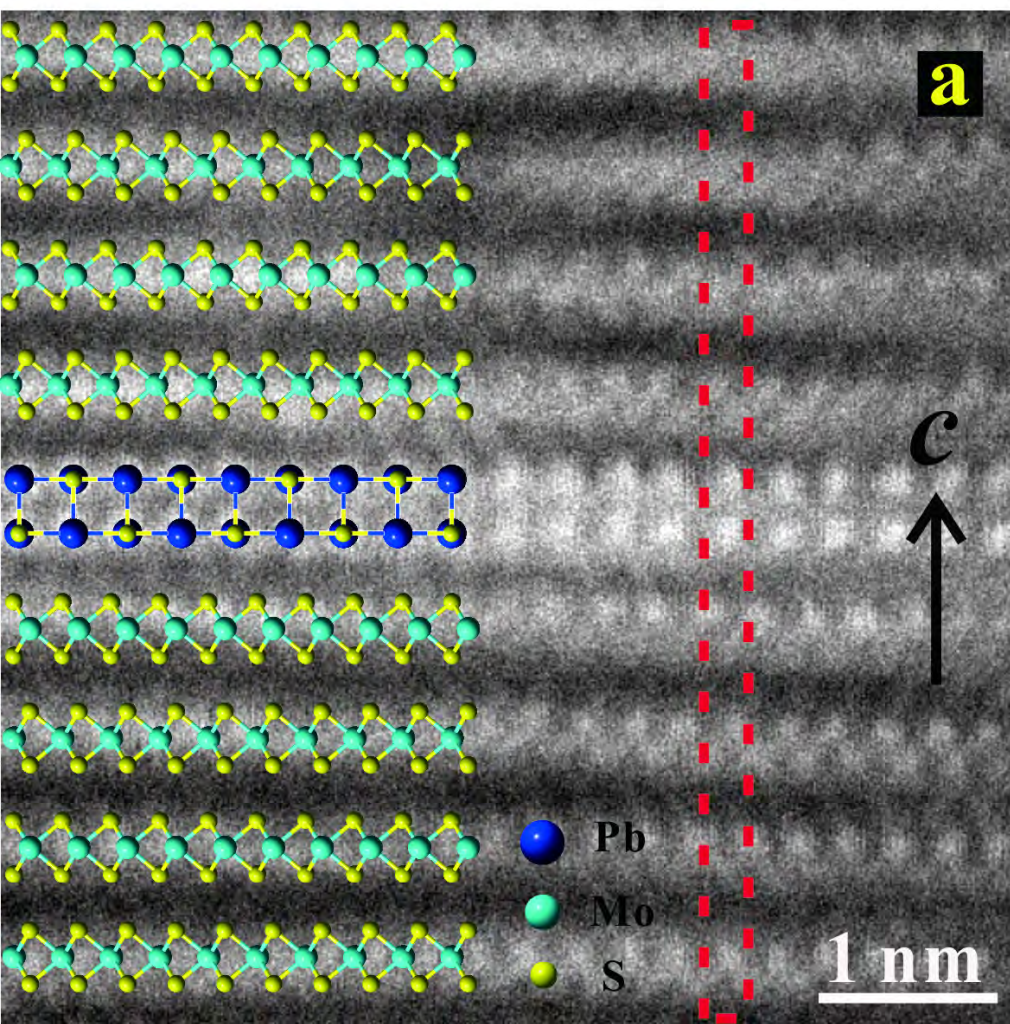
**FIGURE 1**

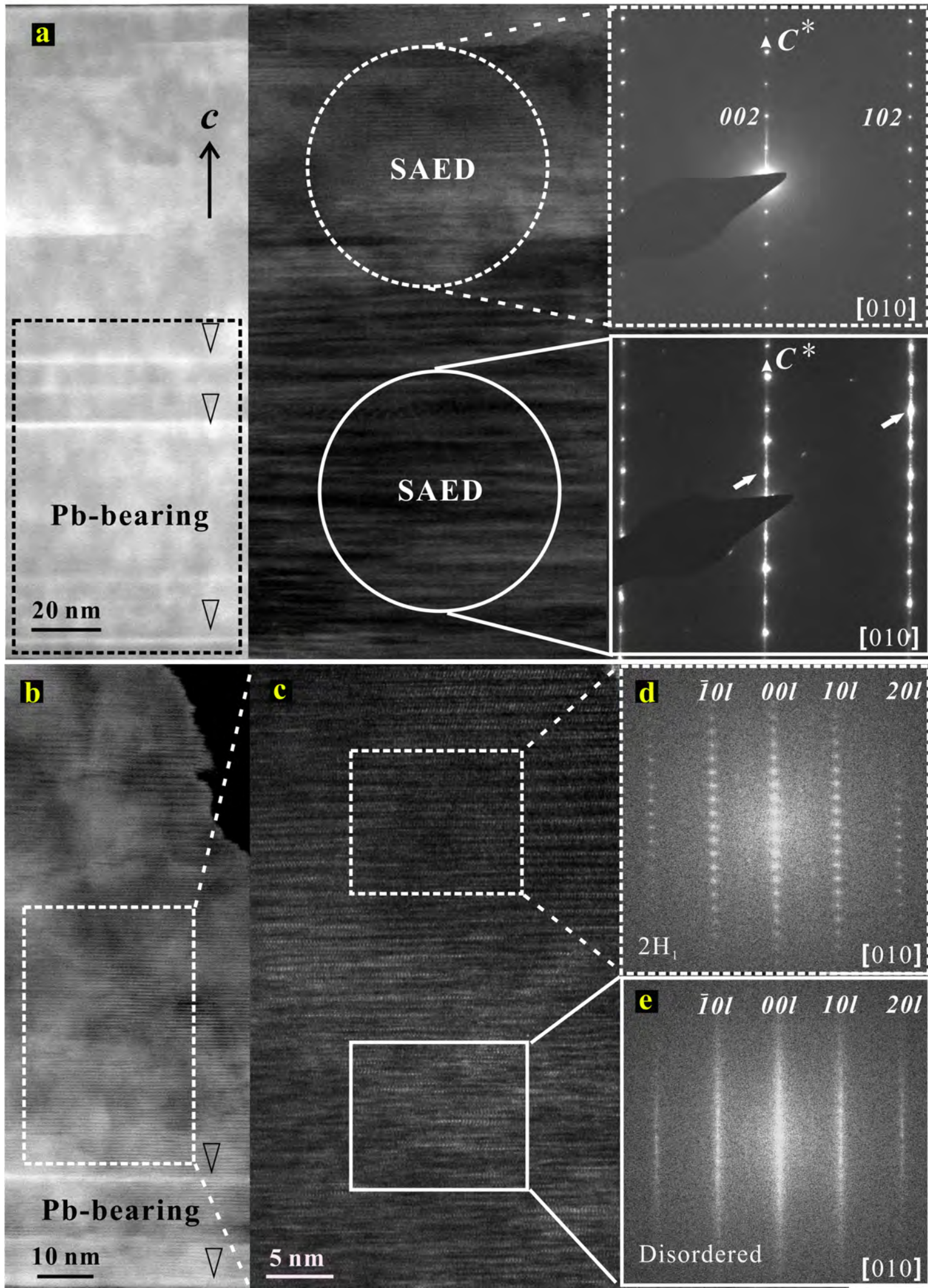


**FIGURE 2**



# FIGURE 3



**FIGURE 4**



**FIGURE 5**

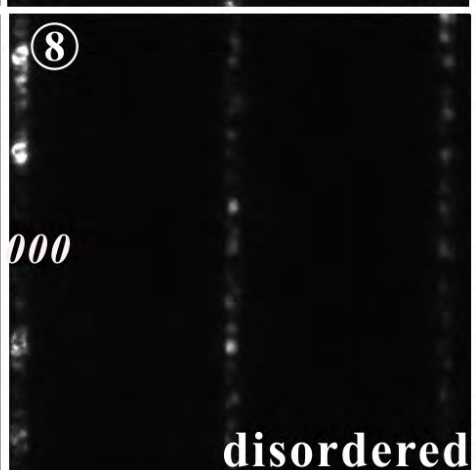
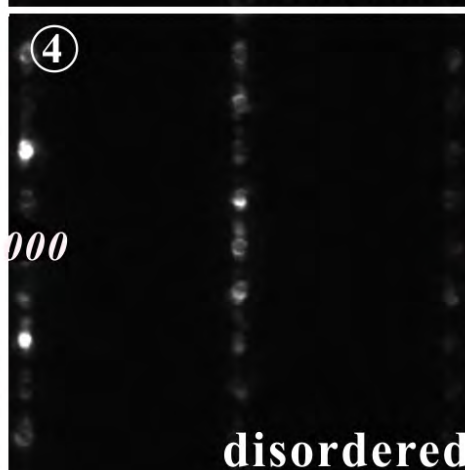
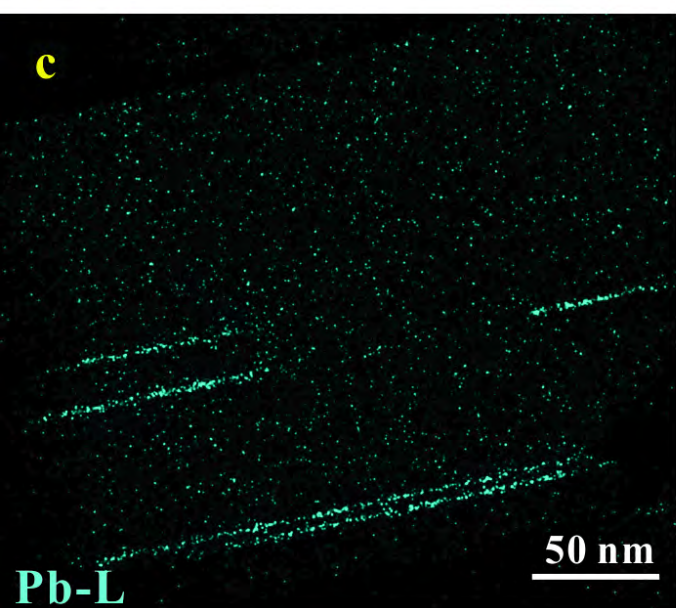
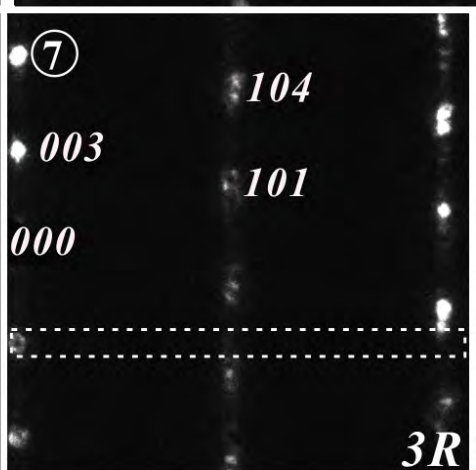
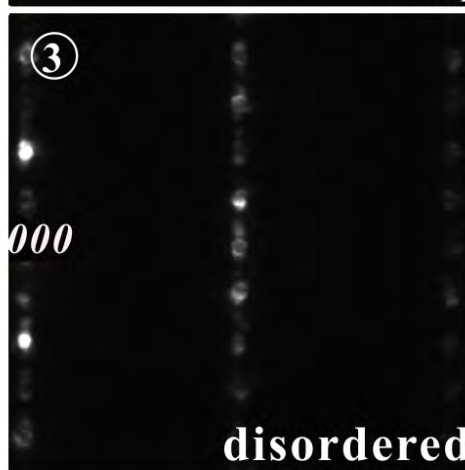
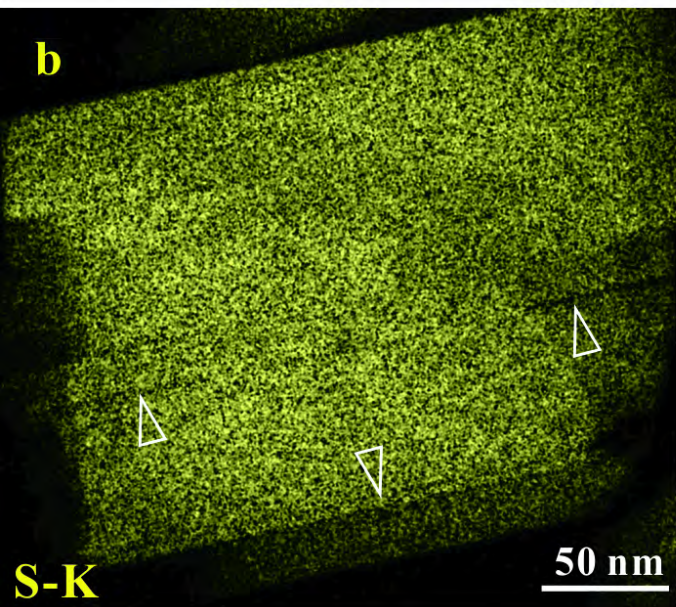
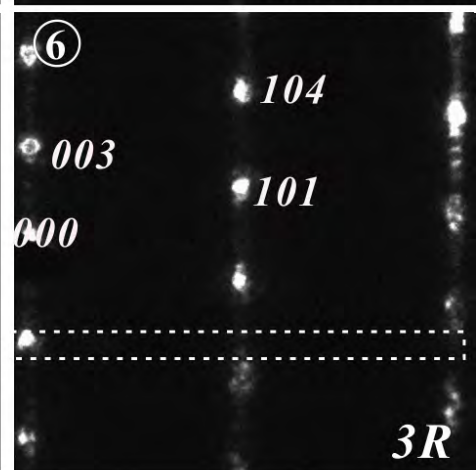
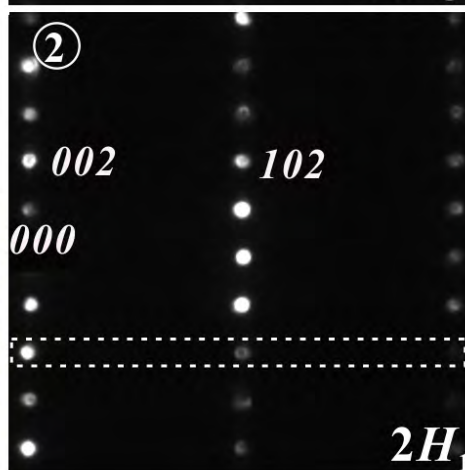
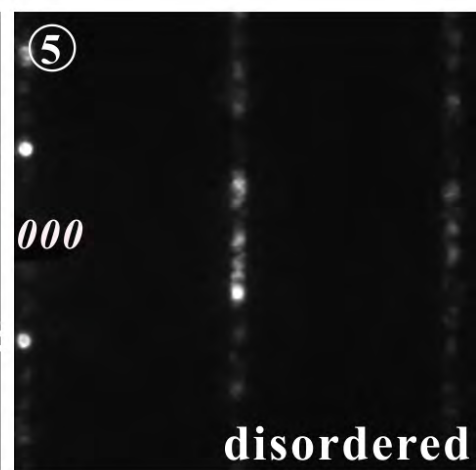
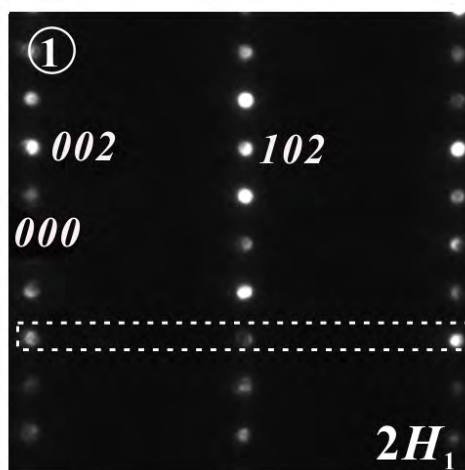
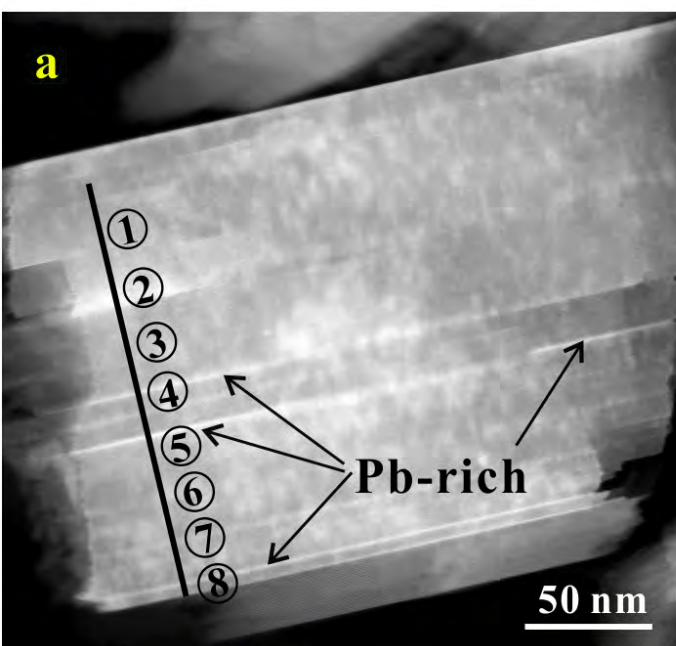
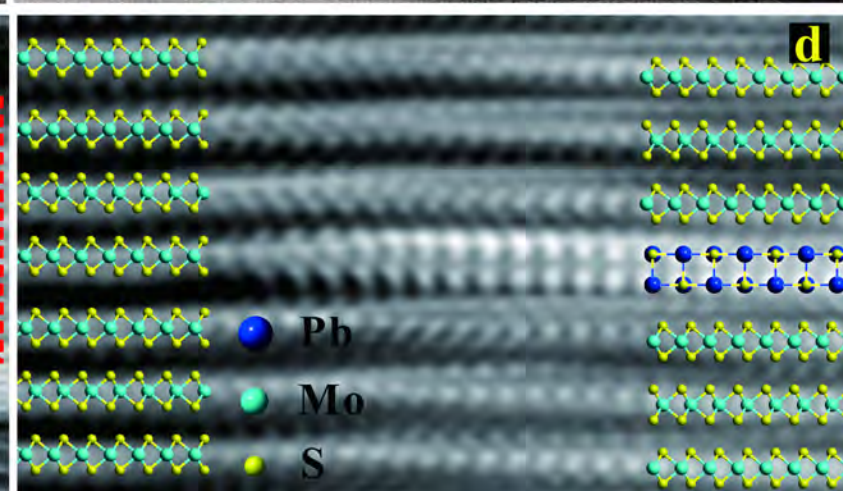
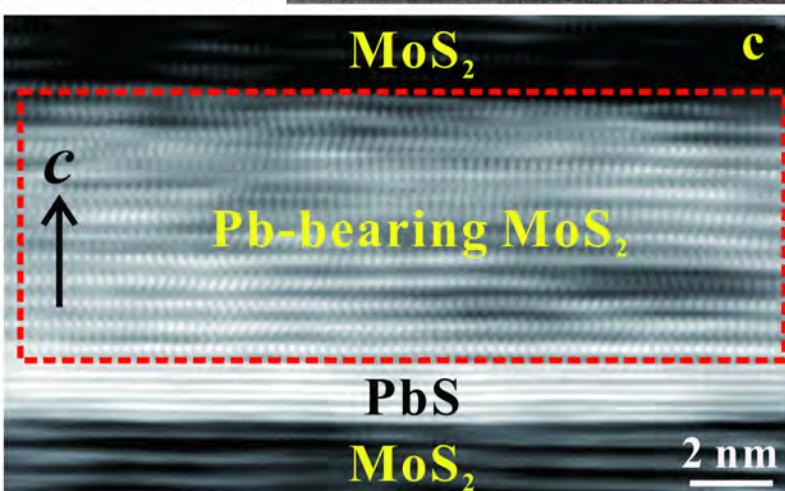
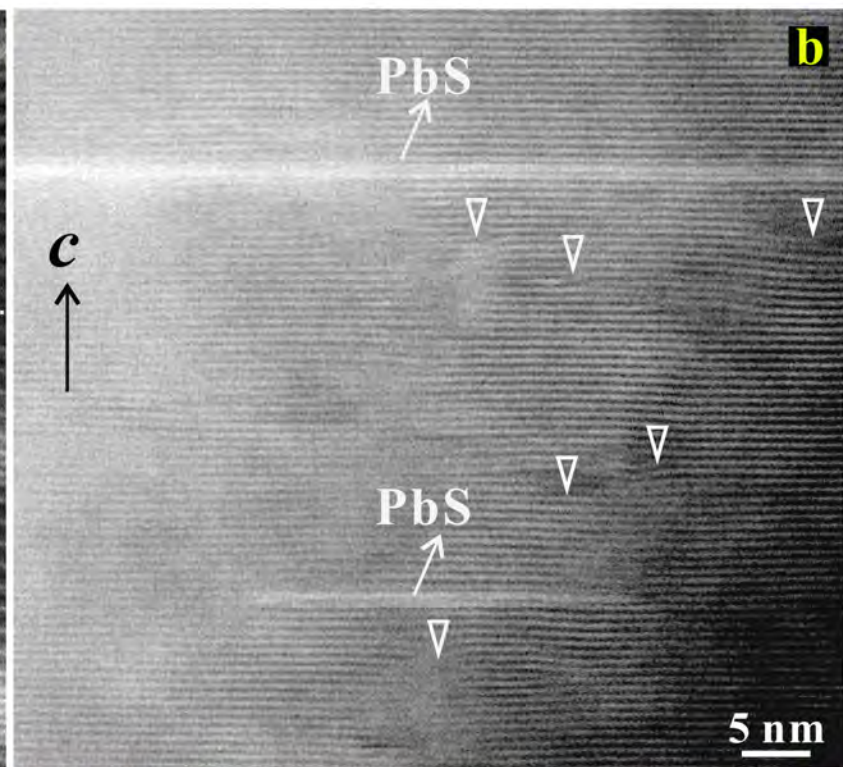
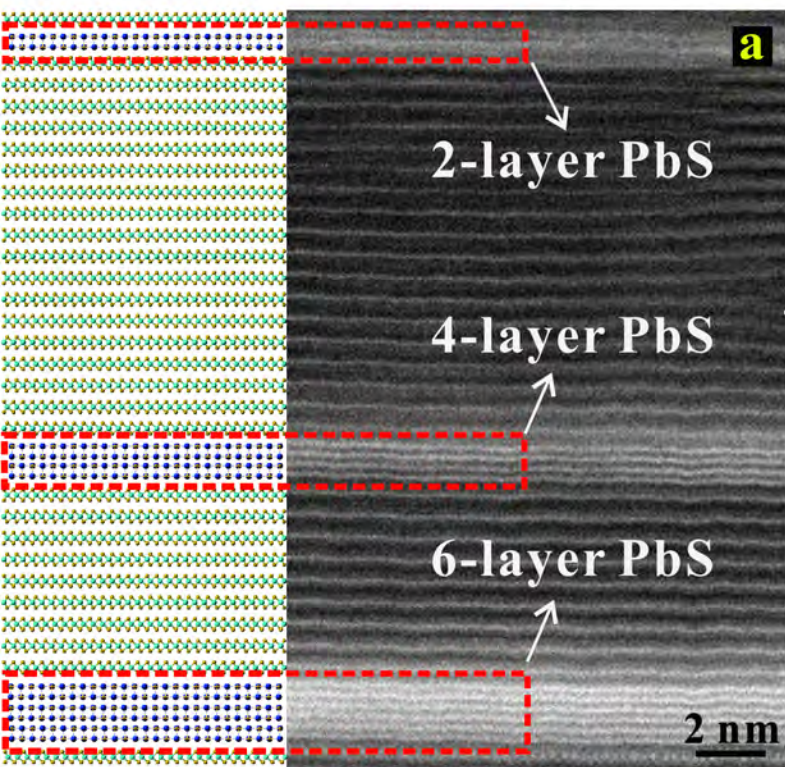
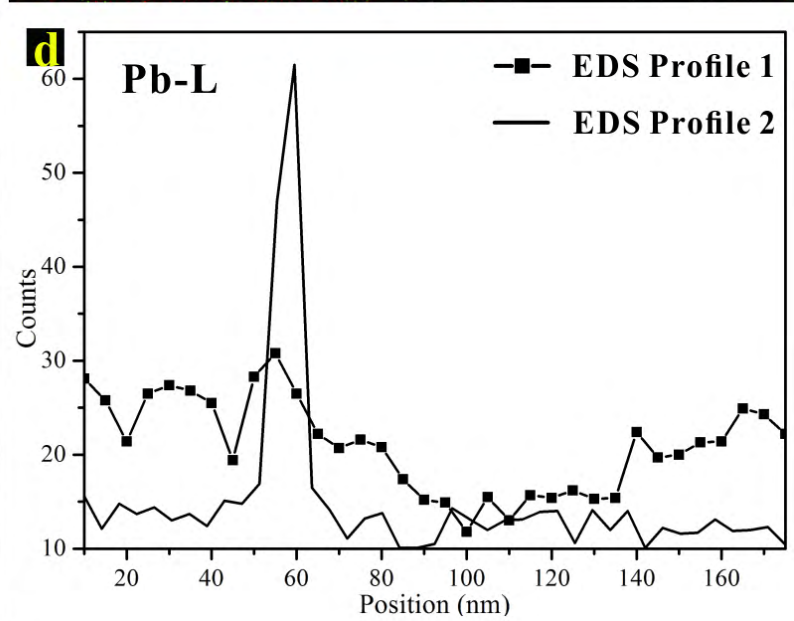
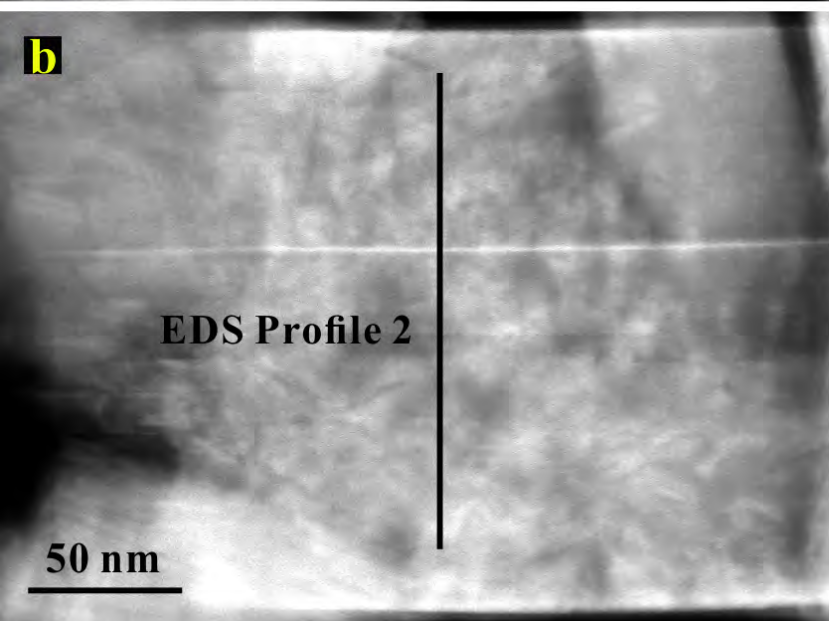
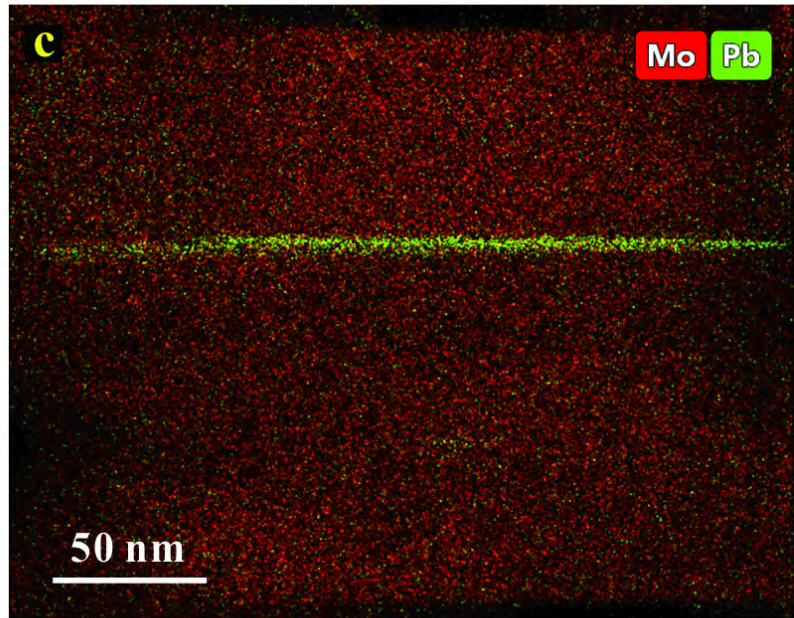
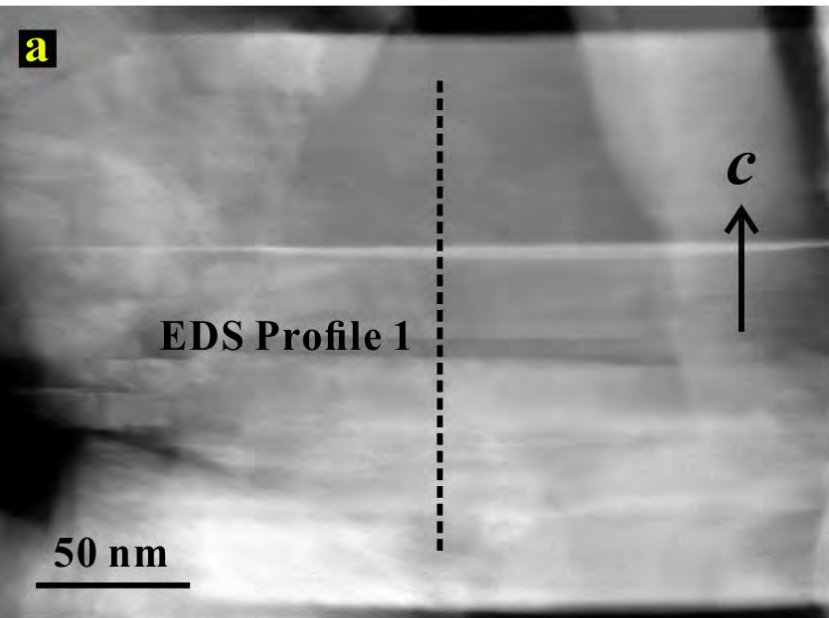


FIGURE 6



**FIGURE 7**



**FIGURE 8**

