

1 Revision 2

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3 **Structural and spectroscopic study of the kieserite-dwornikite solid solution series,**
4 **(Mg,Ni)SO₄·H₂O, at ambient and low temperatures, with cosmochemical implications**
5 **for icy moons and Mars**

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15

16 **Abstract**

17

18 The investigation of the presence and role of sulfates in our solar system receives
19 growing attention, as these compounds play a crucial role in the water budget of planets such
20 as Mars and significantly influence melting equilibria on the icy moons of Saturn and Jupiter,
21 leading to the formation of subsurface oceans and even cryovolcanism. Despite the dominant
22 presence of higher sulfate hydrates such as epsomite, MgSO₄·7H₂O, and mirabilite,
23 Na₂SO₄·10H₂O, on these moons' surfaces, it is not excluded that lower-hydrated sulfates,
24 such as kieserite, MgSO₄·H₂O, are also present, forming from higher hydrates under pressures
25 relevant to the mantle of the icy moons. Given the composition of the soluble fraction in C1
26 and C2 chondritic meteorites, which are high in Ni content and also considered to represent
27 the composition of the rocky cores of the Jovian icy moons, the actual compositions of
28 potentially present monohydrate sulfates likely lie at intermediate values along the solid
29 solution series between kieserite and transition-metal kieserite-group endmembers,

30 incorporating Ni in particular. Moderate Ni contents are also probable in kieserite on Mars
31 due to the planet's long-term accumulation of meteoritic nickel, although likely to a much
32 lesser extent than Fe.

33 Structural and spectroscopic differences between the pure Mg- and Ni-endmembers
34 have been previously documented in the literature, but no detailed crystal chemical and
35 spectroscopic investigation along the Mg-Ni solid solution has been done yet. The present
36 work proves the existence of a continuous (Mg,Ni)SO₄·H₂O solid solution series for the first
37 time. It provides a detailed insight into the changes in lattice parameters, structural details and
38 positions of prominent bands in infrared (transmission, attenuated total reflectance, diffuse
39 reflectance) and Raman spectra in synthetic samples as the Ni/Mg ratio progresses, at both
40 ambient as well as low temperatures relevant for the icy moons and Mars. UV-Vis-NIR
41 crystal field spectra of the Ni endmember also help to elucidate the influence of Ni²⁺-related
42 bands on the overtone- and combination modes.

43 The (Mg,Ni)SO₄·H₂O solid solution series shows Vegard-type behavior, i.e. lattice
44 parameters as well as spectral band positions change along linear trends with increasing Ni
45 content. Infrared spectra reveal significant changes in the wavenumber positions of prominent
46 bands, depending on the Ni/Mg ratio. We show that also the temperature during measurement
47 has an influence on band position, mainly in case of H₂O-related bands. The changes
48 observed for several absorption features in the IR spectra enable rough estimation of the
49 Ni/Mg ratio in the monohydrate sulfate, which is applicable to present and future remote
50 sensing data, as well as in-situ measurements on Mars or the icy moons. The spectral features
51 most diagnostic of composition are the vibrational stretching modes of the H₂O molecule and
52 a band unique to kieserite-group compounds at around 900 cm⁻¹ in the IR spectra, as well as
53 the pronounced ν₃ and ν₁ sulfate stretching modes visible in Raman spectra.

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55 **Keywords:** Kieserite-dwornikite solid solution, icy satellites, Mars, crystal chemistry, FTIR
56 spectroscopy, Raman spectroscopy, UV-Vis-NIR spectroscopy

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61 Introduction

63 The past few decades of planetary exploration uncovered the presence of considerable
64 amounts of sulfates in our solar system. From the surface (and subsurface) of Mars to the icy
65 moons of Jupiter and Saturn and even veinlets in chondritic meteorites, sulfates of Mg, Fe,
66 Mn and Ni represent an important constituent of planetary and meteoritic material
67 (Frederiksson and Kerridge 1988; Burgess et al. 1991; Kargel 1991). The existence of sulfate
68 deposits containing kieserite within layered terrains on Mars is now well established (e.g.,
69 Bishop et al. 2009; Noel et al. 2015; Mangold et al. 2008; Lichtenberg et al. 2010), most
70 probably as a solid solution between kieserite ($\text{MgSO}_4 \cdot \text{H}_2\text{O}$) and szomolnokite ($\text{FeSO}_4 \cdot \text{H}_2\text{O}$).
71 Minor Ni-contents are suggested by Papike et al. (2007), who note that Martian kieserite, via
72 partial Fe-, Ni-, or Co-incorporation, could be an indicator for meteorite impacts as relevant
73 processes in the formation and redistribution of sulfates on Mars. Moreover, Knauth et al.
74 (2005) suggest that the observed sulfate-cemented sedimentary structures could be ascribed to
75 impact surges from a large nickel-iron meteorite.

76 Hydrated Mg-sulfates as alteration products are known from C1 and C2 type
77 carbonaceous chondrites, making up to 73% by mass of their soluble fraction (Frederiksson
78 and Kerridge 1988; Burgess et al. 1991). The identification of higher-hydrated Na- and Mg-
79 sulfates on the surface of Jupiter's moons Ganymede and Europa is another important
80 discovery (McCord et al. 2001; Dalton et al. 2005, 2012). The formation of $\text{MgSO}_4 \cdot \text{H}_2\text{O}$ -
81 Na_2SO_4 low-temperature eutectic brines is believed to be essential for the existence of crustal

82 rifting, cryovolcanism (both active and extinct), or the presence of subsurface oceans on
83 Jupiter's and Saturn's icy satellites (Kargel 1991; McCord et al. 2001). Such oceans could
84 theoretically sustain or have sustained extraterrestrial life in the past (Kargel et al. 2000;
85 McKinnon and Zolensky 2003; Spencer et al. 2009 and citations therein; Solomonidou et al.
86 2011). Active cryovolcanism, owing its existence to tidal heating, is now well documented
87 from the south-polar region of Saturn's moon Enceladus (Spencer et al. 2009). All of these
88 sulfates could have Ni substitution in their structures, and if they are monohydrated, their
89 composition could be identified with the research we present in this work.

90 The high sulfate hydration states interpreted to be present on the surface of icy moons
91 could be modified by interaction with cosmic radiation. For example, epsomite decomposes to
92 kieserite (Jamieson et al. 2014), the thermodynamic stability of which is strongly enhanced by
93 the surrounding vacuum (Zolotov and Shock 2001). The models of fractional eutectic melting
94 indicate the evolution of a Mg-sulfate precipitate ('salty ice') in the high-pressure ice mantle
95 region of the large icy satellites, potentially consisting also of lower sulfate hydrates (Kargel
96 1991; Nakamura and Ohtani 2011). Due to its higher buoyancy, this salty ice mixture is
97 believed to ascend upwards, supplying soluble salts to the overlying ocean (Journaux et al.
98 2017), thus providing nutrients for the potential evolution of extraterrestrial life (Vance et al.
99 2014). The basic stability of kieserite and isostructural sulfate monohydrates of transition
100 metals in such high-pressure environments, albeit modified by displacive second-order phase
101 transitions, is supported by recent investigations (Meusbürger et al. 2018, 2019; Ende et al.
102 2019a,b).

103 The rocky cores of the icy moons of the gas giants, from which the salts now present in
104 their oceans and upper ice layers originated, probably correspond in composition to C1 or C2
105 chondritic material, containing sulfates of Mg, Mn and Ni in the soluble fraction (Burgess et
106 al. 1991; Kargel 1991). As such, the actual composition of monohydrated sulfate, should it be
107 present, will likely represent an intermediate value between kieserite ($\text{MgSO}_4 \cdot \text{H}_2\text{O}$) and its

108 isostructural transition-metal bearing compounds, in particular with Ni- and Mn-contents
109 (Frederiksson and Kerridge 1988, Anders and Grevese 1989).

110 Despite the knowledge of significant changes in lattice parameters and IR spectra
111 between kieserite and the Ni-endmember dwornikite (Wildner and Giester 1991; Stoilova and
112 Lutz 2002; Stoilova 2003), no detailed investigation of the structural and spectroscopic
113 behavior along the Mg-Ni solid solution series was done so far. A structural and crystal
114 chemical study was recently published by Bechtold and Wildner (2016) for the kieserite-
115 cobaltkieserite series, followed by a thorough examination including spectroscopic results of
116 the kieserite-szomolnokite solid solution, $\text{Mg}_{1-x}\text{Fe}_x\text{SO}_4\cdot\text{H}_2\text{O}$, by Talla and Wildner (2019).

117 In this paper, we present the results of detailed structural, crystal chemical, Fourier
118 Transform Infrared (FTIR) and Raman-spectroscopic investigations on synthetically prepared
119 members along the binary kieserite-dwornikite solid solution, $\text{Mg}_{1-x}\text{Ni}_x\text{SO}_4\cdot\text{H}_2\text{O}$, at ambient
120 and lower temperatures, as relevant for surface conditions of Mars and icy satellites of Jupiter
121 and Saturn, complemented by ultraviolet-visible-near infrared (UV-Vis-NIR) crystal field
122 spectra of Ni^{2+} in dwornikite. The presented data may prove valuable in further refining the
123 chemical composition of kieserite present on extraterrestrial bodies in our solar system based
124 on available remote sensing data in the mid-infrared range, as well as providing a fundamental
125 dataset for interpretation of future data of in-situ measurements by landers and rovers.

126 Note that throughout the present paper we use the name “dwornikite”, originally
127 describing a mineral with approximate composition $(\text{Ni}_{0.9}\text{Fe}_{0.1})\text{SO}_4\cdot\text{H}_2\text{O}$ (Milton et al. 1982),
128 to refer to the pure $\text{NiSO}_4\cdot\text{H}_2\text{O}$ endmember.

129

130 **Experimental**

131

132 **Sample preparation**

133 The evaporation technique was used to prepare powder sample material for the IR and
134 Raman spectroscopic studies. Epsomite ($\text{MgSO}_4\cdot 7\text{H}_2\text{O}$) was mixed with $\text{NiSO}_4\cdot 6\text{H}_2\text{O}$ in the

135 desired Mg/Ni molar ratio, amounting to 6 g in total (both chemicals were of analytical
136 grade). The reagents were dissolved in a mixture of 80 ml of doubly distilled H₂O and 10 ml
137 95% H₂SO₄ in a beaker. The solution was evaporated at 70 °C for 10 days, yielding a crust of
138 pale green (Mg,Ni)SO₄·H₂O at the bottom of the vessel, covered by the remaining
139 concentrated acid. In addition, the following hydrothermal technique was used to synthesize
140 coarse-grained samples, enabling selection of individual single crystals for X-ray diffraction
141 and low-temperature Raman investigations. Teflon-lined stainless steel autoclaves with an
142 inner volume of ~2 cm³ were filled with a mixture of epsomite (MgSO₄·7H₂O) and
143 NiSO₄·6H₂O amounting to 0.4 g in total with a pre-set Mg/Ni molar ratio. The solvent
144 consisted of 0.7 ml distilled H₂O and 2.2 ml concentrated H₂SO₄ (w = 0.95). All reagents
145 were of analytical grade. An isothermal temperature run at 220 °C for a period of 14 days
146 followed by cooling to room temperature within 4 days was used. Initial problems such as Ni
147 preferably remaining in solution, synthesized crystals of inadequate size and polysynthetic
148 twinning inherent to tightly sealed autoclaves (samples Ni5a to Ni40a), were largely resolved
149 by applying the ‘leaking vessel approach’. A small leak between the Teflon reaction vessel
150 and its lid was allowed by only gently tightening the lid upon assembly. Through this leak,
151 H₂O slowly evaporated leading to slow precipitation of the monohydrate sulfates as an
152 aggregate of single crystals including individuals with sufficient quality and size.

153 Subsequent treatment of the product was the same for single crystals and powder
154 material as follows. After decanting the fluid and mechanical removal of the solids from the
155 reaction vessel, the product was rinsed twice with distilled H₂O, taking advantage of the
156 sulfate’s relatively slow dissolution rate. Thereafter, the crystals were washed using 98%
157 ethanol to remove the H₂O from the previous step, and the product was dried in an oven at 65
158 °C overnight, then stored in airtight vials.

159

160 **Chemical analyses and powder X-ray diffraction**

161 Wet-chemical analyses were done on a part of each sample batch at the Masaryk
162 University in Brno, Czech Republic. An amount of ~0.5 g of ground sample material was
163 dissolved in boiling HNO₃ (w = 0.65). Contents of Mg and Ni were both determined
164 spectrophotometrically (instrument Solaar M5 - TJA solutions, measurement time 4 s per
165 element) with an analytical error of 0.005 wt% for Mg and 0.002 wt% for Ni.

166 A further part of each sample batch was examined using powder X-ray diffraction, in
167 order to confirm expected lattice parameter shifts along the solid solution and the mono-phase
168 character of each batch. The material was pressed onto an Si-holder and measured on a
169 Philips PW 3710 diffractometer (measuring interval 5-120° 2 θ in 0.2° increments, measuring
170 time of 15 s per step). Following the phase identification with the program EVA2013,
171 Rietveld refinements were performed with the Bruker program TOPAS3.

172

173 **Single crystal X-ray diffraction**

174 Single crystals suitable for X-ray diffraction measurements were selected by choosing
175 individuals with homogeneous extinction under crossed polars. The crystal structures of ten
176 representatives along the Mg_{1-x}Ni_xSO₄·H₂O solid solution series, including the dwornikite
177 endmember, were determined from diffraction data measured at 295 K with graphite-
178 monochromatized MoK α -radiation on a Nonius Kappa CCD diffractometer equipped with a
179 0.3 mm monocapillary X-ray optics collimator. Complete Ewald spheres up to 2 θ = 80° were
180 each collected in several sets of ϕ - and ω -scans with 2° rotation per CCD frame, at a crystal to
181 detector distance of 30 mm. The integration and correction of the intensity data, including an
182 absorption correction by multi-frame scaling and the refinement of lattice parameters, were
183 done with the Nonius program DENZO-SMN.

184 Temperature-dependent X-ray data collections of the dwornikite endmember were
185 performed between +40 and -160 °C in steps of 40 K on a Bruker ApexII diffractometer
186 equipped with a CCD area detector and an Incoatec Microfocus Source I μ S (30 W, multilayer

187 mirror, Mo-K α radiation), in a dry stream of nitrogen (Cryostream 800, Oxford Cryosystems).
188 Several sets of ϕ - and ω -scans with 2° scan width were measured at a crystal-detector distance
189 of 40 mm up to 80° 2 θ full sphere. Absorption was corrected by evaluation of multi-scans.
190 For the sake of data consistency, the lattice parameters extracted from the ApexII
191 temperature-dependent measurements were corrected in such a way that the interpolated
192 values at 20 °C match those obtained from the Nonius Kappa CCD room-temperature
193 measurement of the very same dwornikite single crystal.

194 All structure refinements were performed on F^2 with SHELXL-97 (Sheldrick 2008) in
195 the ‘traditional’ non-reduced kieserite cell setting. Scattering curves for neutral atoms were
196 used. The Mg/Ni ratios of the particular single crystals studied were extracted as a refined
197 variable in the respective structure refinement runs.

198 CIF data for room temperature structures included in Tables 2 and 3 (plus CIF of
199 endmember kieserite from Bechtold and Wildner 2016) and of temperature-dependent
200 structure investigations included in Tables 4 and 5 have been deposited with the submission at
201 the MSA website.

202

203 **IR spectroscopy**

204 The general procedure followed that for the comparable investigations on the kieserite-
205 szomolnokite solid solution by Talla and Wildner (2019). A part of each sample batch was
206 ground by hand in an agate mortar. For our purposes, the obtained sample powder with the
207 size of the largest crystallites not exceeding 100 μ m did not have to be sieved or treated in any
208 additional way, since grain size influences only band amplitudes but not their wavenumber
209 position (Jamieson et al. 2014), because only the latter parameter is dependent on chemical
210 composition. FTIR (Fourier Transform InfraRed) measurements were conducted in
211 transmission, ATR (Attenuated Total Reflectance) and DRIFT (Diffuse Reflectance Infrared
212 Fourier Transform) modes were done by means of the Bruker Tensor 27 FTIR spectrometer

213 (Globar light source, KBr beam splitter, DTGS detector). For transmission measurements, the
214 powdered sample material was diluted in KBr at a weight ratio of 1:300 and pressed into
215 pellets under a vacuum press. ATR measurements were done using the mountable Bruker
216 ATR unit by pressing the sample powder against the diamond surface. For DRIFT
217 measurements, the powder was pressed into the appropriate sample holder of a Perkin–Elmer
218 DRIFT unit. Finely ground MgO was used as the reflectivity standard for the background
219 measurements. All reflectance data were converted via the Kubelka-Munk equation
220 implemented in the Bruker OPUS software. The full wavenumber range (7000–400 cm⁻¹) was
221 investigated in all analytical modes. The instrumental spectral resolution was 4 cm⁻¹. Each
222 measurement was averaged from 50 individual scans to reduce noise. In the scope of the
223 DRIFT analyses, each sample was measured in pure state, as well as diluted with KBr in a
224 ratio of 1:20, to enhance the absorption contribution to the resulting spectrum.

225 Low-temperature FTIR transmission measurements were done on KBr micropellets (2
226 mm aperture diameter in a steel gasket, sample dilution ratio 1:300) using a Linkam FTIR600
227 cooling stage equipped with KBr windows. Measurements were done in 40 K intervals
228 ranging from 313 K (40 °C) down to 113 K (–160 °C). Some additional measurements at 93
229 K (–180 °C) were done occasionally.

230 Our usage of wavenumbers ν (in cm⁻¹) as the principal spectral unit in text and figures,
231 as opposed to wavelengths λ (in μm), is justified by its direct proportionality to the energy of
232 the incident photons. The corresponding wavelengths can be easily obtained via the formula
233 $\lambda(\mu\text{m}) = 10000/\nu(\text{cm}^{-1})$.

234

235 **Raman spectroscopy**

236 Raman spectra were measured on a Horiba Jobin Yvon LabRam–HR spectrometer,
237 equipped with an Olympus BX41 optical microscope. The diffraction grating with 1800
238 grooves/mm and the 633 nm laser were chosen. The system is equipped with an Si-based,

239 Peltier-cooled charge-coupled device (CCD) detector. A 100× objective (NA = 0.55) was
240 used for all room-temperature measurements. The wavenumber accuracy was better than 0.5
241 cm^{-1} , and the spectral resolution was determined to be $\sim 0.3 \text{ cm}^{-1}$. Room-temperature Raman
242 spectra were acquired in tandem with DRIFT measurements (see above), taking advantage of
243 the powdered material for IR spectroscopy already being contained within a convenient
244 sample holder. All Raman spectra were acquired between 100–4000 cm^{-1} shift, using multi-
245 window scans with a counting time of 50 s per window and repeating every scan twice to
246 eliminate spikes.

247 Low-temperature Raman measurements were conducted on single crystals using the
248 Linkam FTIR 600 cooling stage. Despite varying and unknown crystal orientation, the use of
249 single crystals with significantly better scattering, allowing shorter measuring times (5 s per
250 window) was essential to mitigate ice buildup on the sample surface during cooling. The
251 crystals were placed on the Ag-block of the cooling stage covered with a thermal conducting
252 fluid. The bulky stage necessitated the use of a long-distance objective with a 50×
253 magnification (focal distance 10.6 mm), with the same setup of the Raman instrument
254 otherwise. The spectra were acquired in two separate spectral regions, from 100 to 1600 and
255 from 2800 to 4000 cm^{-1} shift. Most acquisition temperatures matched those used during FTIR
256 temperature-dependent investigations to allow for direct comparison.

257 Band positions for both FTIR and Raman spectra were obtained by fitting the spectra
258 with Voigt-shaped band profiles with the program Peakfit (Jandel Scientific, version 4.0).
259 Initial estimates of the number of bands and their distribution in IR and Raman spectra were
260 based on the investigations done on kieserite by previous authors (Stoilova and Lutz 1998;
261 Lane 2007; Lane et al. 2015 for FTIR spectroscopy; and Chio et al. 2007 and citations therein,
262 for Raman spectra). Linear regression analysis of FTIR and Raman data was accomplished
263 using the inbuilt statistical module of the program SigmaPlot 13.

264

265 **UV-Vis-NIR spectroscopy and crystal field calculations**

266 An unpolarized optical absorption spectrum of a single crystal of $\text{NiSO}_4 \cdot \text{H}_2\text{O}$ was
267 measured at room temperature in the spectral range $32000\text{--}5000\text{ cm}^{-1}$ on a mirror-optics
268 microscope IRscope-II, attached to a Bruker IFS66v/S FT-spectrometer, using a measuring
269 spot of $165\text{ }\mu\text{m}$. A KBr beam splitter and appropriate combinations of light sources (Xe- or
270 W-lamp) and detectors (GaP-, Si-, Ge-diodes) were used to cover the spectral range. The final
271 spectrum is combined from three partial spectra (UV and Vis segments: $32000\text{--}10000\text{ cm}^{-1}$,
272 both at 40 cm^{-1} spectral resolution and averaged from 512 scans; NIR: $10000\text{--}5000\text{ cm}^{-1}$ at 20
273 cm^{-1} spectral resolution, 256 scans). The UV and NIR spectral segments were each aligned in
274 absorbance to match with the unadjusted Vis spectrum at 20000 and 10000 cm^{-1} , respectively,
275 and then converted to the linear absorption coefficient α .

276 Crystal field (CF) calculations were performed with the HCFLDN2 module of the
277 computer program package by Y.Y. Yeung (Chang et al. 1994), based both on classical CF
278 parameters assuming a pseudotetragonal field (according to the octahedral $[4+2]$
279 coordination) as well as in the framework of the semi-empirical Superposition Model (SM) of
280 crystal fields. In the latter case, the reference metal–ligand distance R_0 was set to 2.06 \AA , the
281 mean $\langle\text{Ni–O}\rangle$ bond length in dwornikite, and some parameters were fixed to reduce the
282 number of variables (namely $t_2 = 3$, i.e., the ‘electrostatic’ value, and Racah $C/B = 4.2$, as
283 obtained in the classical approach). For further details concerning the calculation procedures
284 the reader is referred to comparable approaches in Wildner et al. (2013) and respective
285 references therein.

286

287

Results

288

289 **Sample chemistry and powder X-ray diffraction**

290 The wet chemical analyses yielded the actual Mg/Ni ratio in the synthesized bulk
291 products in their respective final state, i.e., after the purification process by water and ethanol,
292 described in the ‘Experimental’ section. They revealed systematic differences between the
293 pre-adjusted and the actual Mg/Ni ratio in the product. As is evident from Fig. 1, the synthesis
294 approach greatly influences the extent of discrepancy between pre-set and experimental x_{Ni}
295 values. Samples prepared by the evaporation technique show a convex trend with relative Ni
296 depletion peaking in intermediate members of the solid solution. The few finely crystalline
297 samples prepared in a fully sealed hydrothermal autoclave at 210 °C show pronounced Ni-
298 deficiency with the major part of Ni^{2+} remaining in solution. Details of the chemical
299 composition of samples used for FTIR and Raman spectroscopy are presented in Table 1. The
300 material obtained by the hydrothermal ‘leaking vessel approach’ yielded larger crystal
301 individuals. Because this material was intended for single-crystal measurements, the Mg/Ni
302 ratios of individual hand-picked crystals were extracted as a variable parameter in the
303 respective structure refinements.

304 Powder X-ray diffraction measurements and Rietveld refinements confirmed that the
305 batches consist nearly exclusively of kieserite-group $(\text{Mg,Ni})\text{SO}_4 \cdot \text{H}_2\text{O}$ phases; occasional
306 traces of nickelhexahydrate were found. The diffraction patterns showed no signs of peak
307 splitting or abnormal broadening, thus confirming compositional homogeneity and the mono-
308 phase character of the obtained kieserite-group solid solutions.

309

310 **Crystal structures**

311 Crystal data and details of the data collections and structure refinements of four selected
312 representatives of the $\text{Mg}_{1-x}\text{Ni}_x\text{SO}_4 \cdot \text{H}_2\text{O}$ solid solution, including the dwornikite endmember,
313 are summarized in Table 2, respective final atomic coordinates and displacement parameters
314 are listed in Table 3. Corresponding temperature-dependent data of the dwornikite
315 endmember are compiled for three selected temperatures in Tables 4 and 5. Room-

316 temperature data for endmember kieserite used throughout the present paper are given in
317 Bechtold and Wildner (2016), temperature-dependent data of $\text{MgSO}_4 \cdot \text{H}_2\text{O}$ are found in Talla
318 and Wildner (2019).

319 The structural behavior of the kieserite-dwornikite solid solution series,
320 $\text{Mg}_{1-x}\text{Ni}_x\text{SO}_4 \cdot \text{H}_2\text{O}$, as derived from single-crystal X-ray diffraction measurements, is
321 illustrated in Figures 2–4, and relevant crystal chemical data of four selected representatives
322 including the dwornikite endmember are given in Table 6. Respective details for endmember
323 kieserite from Bechtold and Wildner (2016) are included for easy reference.

324 The kieserite structure type consists of kinked chains of $\text{O3}(\equiv\text{H}_2\text{O})$ -corner-sharing
325 $\text{MgO}_4(\text{H}_2\text{O})_2$ octahedra, clearly elongated along the octahedral water-water axis, with
326 adjacent octahedra being further intra-linked by nearly regular SO_4 tetrahedra via common O2
327 corners. These octahedral-tetrahedral chains are aligned parallel to the c -axis and interlinked
328 to a framework structure by sharing the polyhedral O1 corners as well as by moderately
329 strong hydrogen bonds $\text{O3} \cdots \text{O2}$.

330 The present X-ray measurements and structure refinements do not indicate any
331 octahedral Mg-Ni cation ordering, domain formation or related effects. Figure 2 shows the
332 variation of the lattice parameters between kieserite and dwornikite. It is evident that
333 complete miscibility exists and that the solid solution series behaves according to Vegard's
334 law (Vegard 1921), with all lattice parameters changing in a linear way across the entire
335 composition. The cell volume V (-13.0 \AA^3), angle β (-0.25°), and the lattice parameters a , b
336 and c (-0.08 , -0.03 and -0.18 \AA , respectively) all decrease with Ni uptake. This is mainly
337 driven by the clear decrease of the average metal-oxygen (Me-O) bond length and of the
338 octahedral volume upon Ni intake, as shown in Figs. 3a and 4a. The strongest reduction in
339 individual bond lengths occurs for the Me-O3 bonds with the major vector component
340 parallel to the c -axis, thus resulting in its strong decrease. The Me-O1 and Me-O2 bonds
341 change much less, Me-O1 even increases slightly (Figs. 3a, Table 6). Overall, the octahedral

342 shape is preserved as a clear [4+2] coordination in dwornikite. In contrast to the octahedron,
343 the tetrahedral SO₄ group slightly expands with increasing Ni-content, but tetrahedral bond
344 length and also angular changes (Fig. 3d) rarely exceed a 3 σ limit. The donor–acceptor
345 distance of the medium-strength hydrogen bond O3–H \cdots O2 shortens significantly at higher Ni
346 contents (Fig. 4c).

347 Further factors contributing to the shrinkage of the unit cell with increasing Ni-content
348 are polyhedral rotations and tiltings, evidenced by a significant decrease of the two Me–O–S
349 angles shown in Fig. 4b, whereas the chain angle Me–O3–Me marginally increases. As a
350 consequence, the rather rigid SO₄ tetrahedron rotates by 3.2° around its twofold axis (also
351 compare Fig. 5 in Talla and Wildner 2019).

352 As expected, when temperature is reduced, the cell volume as well as β , a and c
353 decrease, but the b axis lengthens for dwornikite (Fig. 5), as also found for kieserite (Talla and
354 Wildner 2019). Also the mean Me–O bond lengths (Fig. 6a) and octahedral volumes (Fig. 5c)
355 decrease, mainly by reducing the longest Me–O3 bond, and, thus, octahedral distortion is
356 reduced too. S–O bond lengths and the tetrahedral volume show an artificial increase upon
357 cooling (Fig. 6d, 5c) due to changes in thermal motion: a ‘simple rigid bond’ correction
358 according to Downs et al. (1992) reveals practically constant <S–O> bond lengths along the
359 full temperature range. The comparatively short O3–H \cdots O2 hydrogen bond in dwornikite
360 further shortens, but by only about half the extent compared to kieserite (Talla and Wildner
361 2019). A roughly analogous situation is also found for the polyhedra-linking Me–O–S angles
362 in dwornikite which further decrease upon cooling, but less so than in kieserite (and
363 szomolnokite, Talla and Wildner 2019).

364

365 IR spectra

366 The IR spectra of (Mg,Ni)SO₄·H₂O feature several clearly discernible absorption
367 features (Fig. 7). In the FTIR spectra measured in transmission mode, the H₂O absorption

368 region is dominated by a prominent band at $3182\text{--}3057\text{ cm}^{-1}$ ($3.14\text{--}3.27\text{ }\mu\text{m}$), representing the
369 symmetric stretching vibration $\nu_{1(\text{H}_2\text{O})}$ of the H_2O molecule. A broad shoulder at higher
370 wavenumbers ($3367\text{--}3262\text{ cm}^{-1}$ / $2.97\text{--}3.07\text{ }\mu\text{m}$) represents $\nu_{3(\text{H}_2\text{O})}$, the antisymmetric H_2O
371 stretching mode. The band positions of both H_2O stretching vibrations decrease significantly
372 in their wavenumber position upon Ni intake (Fig. 9a, Table 8), whereas the H_2O bending
373 vibration $\nu_{2(\text{H}_2\text{O})}$ remains practically stable ($1525\text{--}1522\text{ cm}^{-1}$ / $6.56\text{--}6.57\text{ }\mu\text{m}$). The same
374 applies for the position of the strongest band in a major band group corresponding to the
375 antisymmetric $\nu_{3(\text{SO}_4)}$ stretching mode of the sulfate tetrahedra (a band group centered at
376 around 1150 cm^{-1} / $8.70\text{ }\mu\text{m}$) and the IR-forbidden symmetric stretching mode $\nu_{1(\text{SO}_4)}$, visible
377 as a weak but well-defined band at around 1030 cm^{-1} ($9.71\text{ }\mu\text{m}$). A pronounced absorption
378 phenomenon, even considered as a potential diagnostic feature for kieserite-group compounds
379 by Lane et al. (2007, 2015), henceforth labeled ‘Peak 900 cm^{-1} ’, occurs at wavenumbers from
380 884 to 941 cm^{-1} ($11.31\text{--}10.63\text{ }\mu\text{m}$) with increasing Ni content. Lastly, a band group at 630 cm^{-1}
381 ($15.87\text{ }\mu\text{m}$) is assigned to the tetrahedral $\nu_{4(\text{SO}_4)}$ bending modes.

382 It must be noted that the absorption band shape and even position vary between
383 individual FTIR measuring modes (Fig. 7, Table 8). As example, the symmetric stretching
384 vibration of the H_2O molecule occurs at 3182 cm^{-1} in kieserite measured in transmission
385 mode, whereas it is centered at 3165 cm^{-1} in ATR mode and as high as 3203 cm^{-1} using the
386 DRIFT technique on the very same sample. Nevertheless, the overall spectral appearance is
387 comparable, with the absorptions related to the H_2O bending vibration and the H_2O
388 combination modes being enhanced in DRIFT spectra (Fig. 7, Fig. 8, respectively), especially
389 on pure sample material. Contrary to this specific benefit of using undiluted material,
390 significant features such as the ‘Reststrahlen band’ (labeled ‘R’ in Fig. 7), a pronounced
391 absorption at around 1300 cm^{-1} , obscure the expected intrinsic sulfate vibrations between
392 1300 and 370 cm^{-1} . The spectral position of these features, however, is also dependent on the

393 sample composition, e.g., the Reststrahlenband decreases in position from 1360 to 1286 cm⁻¹
394 (7.35–7.77 μm) between kieserite and dwornikite.

395 The enhanced amplitude of absorption bands assigned to H₂O combination modes in the
396 4400–5200 cm⁻¹ (2.27–1.92 μm) spectral region in diffuse reflectance spectra of undiluted
397 sample material allows for their meaningful evaluation. Three distinct bands occur in this
398 region at 4688, 4845 and 5087 cm⁻¹ (2.13, 2.06 and 1.97 μm, respectively) in kieserite. As the
399 Ni content increases, both peripheral bands diverge from the original central band position for
400 kieserite at 4845 cm⁻¹ to 4566 and 5110 cm⁻¹ (2.19 and 1.96 μm, respectively), thus enhancing
401 the band separation with increasing Ni (Fig. 8ab), while the central band itself shifts down to
402 4735 cm⁻¹ (2.11 μm).

403 In many cases, the observed changes in spectral band position are rather limited, such as
404 for the sulfate vibrations with ≤ 30 cm⁻¹ in all three measuring modes, and that of the H₂O
405 bending mode, with a minimal decrease of only 3 cm⁻¹ from kieserite to the dwornikite
406 endmember.

407 Pronounced linear correlation trends can be observed for the H₂O stretching vibrations,
408 where the well-resolved ν_{1(H₂O)} band changes in position from 3182 to 3057 cm⁻¹ and the
409 ν_{3(H₂O)} shoulder decreases in wavenumber from 3367 to 3262 cm⁻¹ between kieserite and
410 dwornikite. A somewhat smaller but still prominent linear change can be seen for the apparent
411 kieserite-group ‘diagnostic’ absorption band (Lane et al. 2007, 2015), increasing from 884 in
412 kieserite to 941 cm⁻¹ with increasing Ni content (Fig. 9a).

413 The variation in the wavenumber position of the relevant spectral absorption
414 phenomena in the scope of the individual measurement modes, as well as the respective linear
415 regression coefficients are summarized in Table 8.

416 The results of FTIR measurements at low temperature, depicted for the Mg-Ni solid
417 solution series in Fig. 10a, show significant changes in the wavenumber position of the H₂O
418 stretching vibrations, which decrease significantly upon cooling. On the contrary, the

419 ‘diagnostic’ band at $\sim 900\text{ cm}^{-1}$ increases in wavenumber. The vibrations of the sulfate
420 tetrahedra remain nearly unaffected throughout the entire temperature range, as is also true for
421 the H_2O bending vibration. Reduced fitting accuracy obscures any observable temperature-
422 related trend for the $\nu_{3(\text{H}_2\text{O})}$ band (Fig. 10a).

423

424 **Raman spectra**

425 The Raman spectra consist of numerous narrow bands in the $100\text{--}1600\text{ cm}^{-1}$ shift
426 region, with much better resolution compared to FTIR spectra (Fig. 9 vs. 11a). The most
427 prominent band is the symmetric stretching vibration $\nu_{1(\text{SO}_4)}$ of the sulfate tetrahedron at 1042
428 and 1020 cm^{-1} shift in kieserite and dwornikite, respectively. Several relevant spectral features
429 occur at lower shift values, the most prominent situated at ~ 220 , ~ 430 and a doublet at ~ 630
430 cm^{-1} shift. Weak bands occur in the $1100\text{--}1600\text{ cm}^{-1}$ spectral region. With the exception of the
431 somewhat broader peak at roughly 1500 cm^{-1} (H_2O bending vibration), these bands are
432 assigned to the split $\nu_{3(\text{SO}_4)}$ vibration (Stoilova and Lutz 1998; Chio et al. 2007). Stretching
433 vibrations of the H_2O molecule are visible in the expected spectral region with the
434 antisymmetric $\nu_{3(\text{H}_2\text{O})}$ contribution present as a mere poorly defined shoulder of the $\nu_{1(\text{H}_2\text{O})}$
435 band, which decreases from 3179 down to 3071 cm^{-1} with increasing Ni-content (Fig. 12a,
436 Table 9). The position of the H_2O bending mode remains constant in analogy to FTIR results,
437 while most sulfate-related bands show a slight decrease upon Ni uptake with the exception of
438 the two peaks assigned to the $\nu_{2(\text{SO}_4)}$ bending mode in the $\sim 400\text{--}500\text{ cm}^{-1}$ shift spectral region
439 (Fig. 12a). The position of the band at $\sim 220\text{ cm}^{-1}$ shift assigned by Chio et al. (2007) to a
440 translational mode of H_2O and the MeO_6 octahedra increases significantly upon Ni uptake, as
441 is also observed for the band assigned to ‘lattice modes’ at $\sim 120\text{ cm}^{-1}$ shift.

442 Linear changes in band position can also be seen in the Raman spectra acquired across
443 the $(\text{Mg,Ni})\text{SO}_4\cdot\text{H}_2\text{O}$ solid solution, as could be expected from the FTIR results (Table 9).

444 Fig. 12a illustrates the situation in detail, while Table 9 gives the respective coefficients of the
445 linear regression between the band position and x_{Ni} .

446 Raman measurements at low temperature show a significant temperature dependence of
447 the spectral position of certain modes, in analogy to the results of low-temperature FTIR
448 measurements. The very good resolution and low FWHM of the bands in Raman spectra
449 allow to accurately examine even the subtle positional changes of sulfate modes, which are
450 not so apparent in IR spectra (Table 9). In some cases, unfortunately, the amplitude of some
451 of the monitored Raman bands is largely suppressed due to an unfavorable respective crystal
452 orientation, causing increased errors upon fitting.

453 Despite this, a composition-independent systematic shift of the band position with
454 temperature is seen mainly for the $\nu_{1(\text{H}_2\text{O})}$ band, which decreases in wavenumber by about 30
455 cm^{-1} across the full temperature range from +40 to -196°C , and represents the strongest
456 observable change among all examined bands. As a shoulder of the well-defined $\nu_{1(\text{H}_2\text{O})}$ band,
457 the $\nu_{3(\text{H}_2\text{O})}$ band is generally poorly resolved, thus impeding any precise fit. The wavenumber
458 change with temperature for bands related to sulfate vibrations is, as seen also in FTIR
459 spectra, quite small, in general never exceeding 10 cm^{-1} in the examined temperature range
460 (Table 9). A systematic increase in wavenumber is observed at low temperature for the
461 lowest-energetic bands, representing Mg(Ni)–O stretching modes according to Chio et al.
462 (2007). For all observed bands, the rates at which the band positions change with temperature
463 remain constant across the entire examined temperature range, allowing to determine a mean
464 shift value per 1°C for the individual bands, included in Table 9.

465

466 **UV-Vis-NIR spectroscopy and crystal field calculations**

467 The optical absorption spectrum of dwornikite is shown in Fig. 13. Three intense crystal
468 field absorption bands are observed around 8100 , 13450 and 24600 cm^{-1} , which are assigned
469 to spin-allowed electronic transitions from the $^3\text{A}_{2g}(\text{F})$ ground state to the $^3\text{T}_{2g}(\text{F})$, $^3\text{T}_{1g}(\text{F})$ and

470 ${}^3T_{1g}(P)$ levels of Ni^{2+} ($3d^8$ electron configuration) in octahedral (O_h) symmetry. The latter two
471 bands are modified by spin-forbidden states, i.e., ${}^1E_g(D)$ within the high-wavenumber slope of
472 the mid-energy band, and ${}^1T_{2g}(D)$ near the onset of the high-energy band; the structures below
473 6600 cm^{-1} originate from vibrational combination and overtone modes of the H_2O molecule.
474 Evidently, no significant splitting of the first and third spin-allowed bands is observed, either
475 by the naked eye or via peak-fitting analyses, as would be expected from the geometric [4+2]
476 elongation of the NiO_6 octahedron (Table 6, Fig. 3a). A splitting might be postulated for the
477 middle spin-allowed band (with the main component at 13450 cm^{-1}), but due to spin-orbit
478 mixing and intensity stealing by the spin-forbidden ${}^1E_g(D)$ level, the position and intensity
479 contribution of this second split component of ${}^3T_{1g}(F)$ can only be roughly estimated (~ 14800
480 cm^{-1}). The refined crystal field and interelectronic repulsion parameters resulting from these
481 assignments are summarized in Table 10, and respective observed and calculated energy
482 levels (from the ‘classical’ tetragonal CF approach) are included in Fig. 13.

483

484

Discussion

485

Crystal structural evolution and crystal chemistry

487 The present crystal structure investigations corroborate the existence of a continuous
488 solid solution series, $Mg_{1-x}Ni_xSO_4 \cdot H_2O$, between kieserite, $MgSO_4 \cdot H_2O$, and isotypic
489 dwornikite, $NiSO_4 \cdot H_2O$. In view of previous crystal structure investigations of the two
490 endmembers (for kieserite: Bechtold and Wildner 2016; Hawthorne et al. 1987; for
491 dwornikite: Wildner and Giester 1991), the respective solid solution series behaves according
492 to ‘basic’ expectations. In particular, all lattice parameters and crystal chemical data exhibit
493 Vegard-type behavior within the limits of error, undergoing linear changes with progressing
494 Mg/Ni ratio. This trend serves as basis for any further theoretical calculations of interest, and
495 the significant linear changes between the endmembers importantly allow, amongst others, to

496 infer the Mg/Ni ratio in binary kieserite-group samples of unknown composition from X-ray
497 data. Vegard-type behavior indicates one-site ideal mixing of the endmembers (Powell and
498 Holland 1993), as is also the case with the recently investigated kieserite-szomolnokite solid
499 solution (Talla and Wildner 2019). Enthalpy, entropy, molar volume, as well as heat capacity
500 and thermal expansion can be determined by the linear combination of these parameters for
501 both endmembers in their respective formula ratio in the sample (van Hinsberg et al. 2005a,b).

502 Likewise, the crystal chemical expectations are fulfilled in the sense that the
503 replacement of the larger Mg cation ($r_{\text{Mg}} = 0.720 \text{ \AA}$) by the smaller Ni^{2+} ($r_{\text{Ni}^{2+}} = 0.690 \text{ \AA}$; all
504 radii from Shannon 1976) leads to a decrease of the Me–O2 and especially Me–O3 bond
505 lengths as well as of the average $\langle \text{Me–O} \rangle$ distance, and consequently to a reduction of the
506 respective octahedral and the unit-cell volume (Figs. 2b, 3a and 4a). Due to a slight increase
507 of the Me–O1 bonds with Ni-content, a characteristic [4+2] coordination of the MeO_6
508 polyhedron is also realized in dwornikite (as in kieserite and most other kieserite-group
509 sulfates, e.g. cobaltkieserite, Bechtold and Wildner 2016) rather than tending towards a
510 [2+2+2]-type coordination, such as for the FeO_6 polyhedron in szomolnokite (Talla and
511 Wildner 2019). In the latter case of the Mg-Fe solid solution ($r_{\text{Fe}^{2+}} = 0.780 \text{ \AA}$), the cell
512 parameters (except for the *b*-axis) and all Me–O bond lengths significantly increase with Fe
513 content.

514 In contrast to the octahedron, the tetrahedral SO_4 group very slightly expands with
515 increasing Ni-content (Figs. 3c, 4a), a behavior analogously observed in the respective Mg-Fe
516 and Mg-Co solid solutions (Talla and Wildner 2019; Bechtold and Wildner 2016). The
517 donor–acceptor distance of the medium-strength hydrogen bond $\text{O3–H}\cdots\text{O2}$ shortens
518 significantly at higher Ni contents (Fig. 4c), with important impact on the vibrational spectra
519 discussed below.

520 The unit-cell contraction is not only driven by reduction of the Ni–O compared to the
521 Mg–O bond lengths (Fig. 3a), but further intensified by octahedral-tetrahedral tiltings via

522 reduced Me–O–S angles with increasing x_{Ni} (Fig. 4b), meaning that the contraction exceeds
523 the one expected from the mere differences in the Me^{2+} ionic radii. This finding parallels
524 those observed for the kieserite-cobaltkieserite (Bechtold and Wildner 2016) and kieserite-
525 szomolnokite (Talla and Wildner 2019) solid solutions, and has been attributed by these
526 authors to the absence (for Mg), respectively the presence (for Co, Fe; here: Ni), of 3d
527 orbitals; hence, for detailed discussions on this subject, which fully applies to the present case
528 of the replacement of Mg by Ni, the reader is referred to those papers (and references therein).
529 Similarly, the apparent difference in the character of the octahedral bond length distortion,
530 i.e., a clear [4+2]-type elongation in kieserite, dwornikite and all other kieserite-group
531 endmember compounds except szomolnokite, which displays a tendency towards a [2+2+2]
532 distortion, has been amply addressed by Talla and Wildner (2019), linking it with the 3d-
533 electron configuration of the particular Me^{2+} cation.

534

535 **IR, Raman and crystal field spectra**

536 **IR spectra at ambient conditions.** The linear Vegard-type behavior observed in the
537 structural data for the $\text{Mg}_{1-x}\text{Ni}_x\text{SO}_4 \cdot \text{H}_2\text{O}$ solid solution series (Figs. 2–4) is also reflected in
538 the results of FTIR and Raman spectroscopic measurements. With wavenumber units in use,
539 linear shifts of the IR spectral bands positions across the kieserite-dwornikite solid solution
540 series (Fig. 9a, Table 8) are recognized, in principle allowing to deduce the respective Mg/Ni
541 ratio solely from spectroscopic data.

542 The evidence that the wavenumbers of both stretching modes of the H_2O molecule,
543 $\nu_{1(\text{H}_2\text{O})}$ and $\nu_{3(\text{H}_2\text{O})}$, decrease upon Ni uptake is in agreement with the single-crystal X-ray
544 results (Fig 4c, Table 6), where a decrease in the hydrogen bond length between the O3 donor
545 and the O2 acceptor oxygen with increasing x_{Ni} can be correlated with the observed band
546 behavior according to well-established trends (Libowitzky 1999). The opposite behavior
547 occurs in the kieserite-szomolnokite solid solution, where Fe contents lead to the lengthening

548 of the O3···O2 donor-acceptor distance, resulting in an increase in wavenumber observable
549 mainly for the $\nu_{1(\text{H}_2\text{O})}$ stretching mode (Talla and Wildner 2019, Fig. 9b). The more or less
550 constant position of the bending vibration of the H₂O molecule independent of the Ni content
551 indicates a rigid character of the H₂O molecule, despite the widening of the acceptor–donor–
552 acceptor angle O2–O3–O2 from kieserite to dwornikite (136.7–140.5°). Both the symmetric
553 $\nu_{1(\text{SO}_4)}$ vibration and the three bands assigned to the split $\nu_{3(\text{SO}_4)}$ mode (Chio et al. 2007)
554 slightly decrease in wavenumber position with increasing x_{Ni} , in accord with the observed
555 slight relaxation of the SO₄²⁻ tetrahedron reflected in the elongation of S–O bond lengths
556 towards dwornikite (Figs. 3c, 4a, Table 6). Of the three bands assigned to the $\nu_{3(\text{SO}_4)}$ vibration,
557 only the major one is depicted and its position followed in Fig. 9a and Table 8. In contrast to
558 the diverging trends found for the hydrogen bonding schemes and O–H stretching vibrations,
559 the structural and spectroscopic sulfate behavior upon Ni uptake is closely comparable with
560 the situation recently investigated (Talla and Wildner 2019) for Fe uptake in the kieserite–
561 szomolnokite series (Fig. 9a,b). The prominent absorption band at ~900 cm⁻¹ deemed as
562 ‘diagnostic’ by Lane et al. (2015) shows a rather pronounced wavenumber increase towards
563 Ni-rich compositions, as opposed to its behavior upon Fe uptake (Fig. 9a,b). This band is
564 usually assigned to a librational mode of the H₂O molecule (e.g. Lane 2007; Lane et al. 2015),
565 which is supported by its contradicting behavior in case of the Mg-Fe and Mg-Ni solid
566 solutions, such as in the case of the H₂O stretching modes. Alternatively, considering the low-
567 temperature behavior described below, the band may be also attributed to a combination mode
568 of the sulfate and octahedral vibration modes (Talla and Wildner 2019). A group of bands at
569 ~630 cm⁻¹ (Fig. 7), assigned by Lane et al. (2015) to the $\nu_{4(\text{SO}_4)}$ vibration, can hardly be
570 exploited for cosmochemical considerations, as the spectral region not only consists of
571 numerous peaks with varying FWHM, but also is not clearly discernible from the signals of
572 other sulfates (Cloutis et al. 2006) and also of atmospheric CO₂. Therefore, the IR spectral
573 region between 700 and 300 cm⁻¹ (14.3–27.0 μm) was not studied in further detail.

574 Contrary, the 4400–5200 cm^{-1} (2.27–1.92 μm) spectral region, where combination
575 modes of the H_2O stretching and bending vibrations induce absorptions, is considered by
576 many authors to be an important spectral signature in reflectance spectra, allowing not only to
577 discern kieserite from other sulfate hydrates, in which these bands occur at higher
578 wavenumbers (Mangold et al. 2008; Noel et al. 2015), but also to roughly infer either its Fe
579 content (Cloutis et al. 2006; Bishop et al. 2009; Liu et al. 2016, Talla and Wildner 2019) or Ni
580 content (this work), due to the opposite behavior of the respective group of bands in the two
581 binary sulfate monohydrate solid solutions. The H_2O combination modes are most apparent in
582 diffuse reflectance spectra measured on undiluted sample material, whereas they are largely
583 suppressed and too faint to be tracked effectively in the other measuring modes. The kieserite
584 endmember shows a typical set of three bands at 4688, 4845 and 5087 cm^{-1} (2.13, 2.06, 1.96
585 μm) (Fig. 8ab, Table 8). The two lower-energetic components decrease significantly to 4566
586 and 4735 cm^{-1} (2.19 and 2.11 μm), respectively, in the dwornikite endmember, whereas the
587 peak at the highest wavenumber increases to 5110 cm^{-1} (Fig. 8a). In case of Fe-rich solid
588 solutions and pure szomolnokite, the peripheral bands show opposite behavior, converging to
589 the central one while largely maintaining the overall spectral position of the triplet. The
590 highest-energetic band is always rather weak, even creating the misleading impression of a
591 single broad absorption in the case of szomolnokite, since it occurs but as a weak shoulder in
592 the bandgroup (Talla and Wildner 2019).

593 The afore-mentioned impact of the sample dilution on DRIFT spectra and their artifacts
594 has important implications for orbiter measurements, where IR spectra reflected by ‘fluffy’
595 kieserite aggregates (more transmission) would somewhat differ in shape and band position
596 from signals acquired on compact kieserite masses or crusts due to the different
597 reflection/transmission ratio of the particular material. Disregard of these issues could
598 complicate the assessment of the properties and composition of the measured monohydrate

599 sulfate. Contrary, the grain size seems to have little effect on band *position* in IR reflectance
600 spectra (Jamieson et al. 2014; Pitman et al. 2014).

601 According to their divergent behavior in the IR spectra, the H₂O-related bands allow to
602 easily discern between Fe or Ni enrichment in the ideal case of a purely binary
603 Mg_{1-x}Me_xSO₄·H₂O solid solution at ambient temperature (Table 8). In case of Ni contents, the
604 wavenumbers of the H₂O stretching modes $\nu_{1(\text{H}_2\text{O})}$ and $\nu_{3(\text{H}_2\text{O})}$ are always lower than in
605 endmember kieserite. Conversely, respective Fe enrichment is indicated by a higher
606 wavenumber mainly of the $\nu_{1(\text{H}_2\text{O})}$ stretching vibration compared to pure kieserite (Talla and
607 Wildner 2019, Fig. 9a,b). In addition, the H₂O combination mode region centered at ~4900
608 cm⁻¹ (2.04 μm) in DRIFT spectra is a further promising candidate for the assessment of Ni or
609 Fe contents, due to the opposite behavior of the peripheral bands (Fig. 8, Table 8). Likewise,
610 the band at ~900 cm⁻¹ (11.1 μm) is another candidate allowing to discern between Fe or Ni
611 contents, since it increases in wavenumber by 57 cm⁻¹ in case of Ni uptake towards the
612 dwornikite endmember, but decreases by roughly 50 cm⁻¹ in szomolnokite in respect to the
613 kieserite endmember benchmark. The H₂O bending vibration as well as the sulfate modes all
614 show but minor changes in respect to the sample chemistry. These are comparable
615 irrespective of the transition element substituting Mg (Fig. 9 a,b).

616
617 **Temperature dependence of IR spectra.** The discussion so far was limited to room-
618 temperature data. However, temperature-related changes are to be expected, not only for data
619 from Mars (+20 to -120 °C at latitudes < 40°; Witzke et al. 1997), but especially considering
620 the situation on the icy moons of Jupiter and Saturn with mean equatorial surface temperature
621 in the range of 90–100 K (e.g., on Europa, Ashkenazy 2019). The structural changes upon
622 temperature decrease are rather similar for dwornikite (Tables 4 and 7), kieserite and
623 szomolnokite (Talla and Wildner 2019), without any indication of a phase transition in either
624 case. Examining and comparing the respective acquired low-temperature IR-spectra (Fig.

625 10a,b), two aspects are evident. Firstly, all vibrations related to the sulfate group are more or
626 less stable in their wavenumber position regardless of temperature across either solid solution,
627 confirming their rigid character, regardless of the element substituting Mg. The H₂O-related
628 symmetric stretching vibration shows a comparable decrease in wavenumber position for the
629 Mg-Ni and Mg-Fe solid solutions (Talla and Wildner 2019), irrespective of composition. The
630 band at ~900 cm⁻¹ in the Mg-Ni solid solution and at ~850 cm⁻¹ in the Mg-Fe series (Fig.
631 10a,b) shows a wavenumber increase by about 0.07 cm⁻¹/°C across the +40 °C to -180 °C
632 temperature range in both solid solutions. The parallel reaction to temperature decrease, along
633 with the inverted correlation trends upon Ni- or Fe- incorporation in accord with the behavior
634 of the H₂O stretching modes, strongly hint this band to be linked to an H₂O-related
635 (librational) mode.

636 Surprisingly, the H₂O bending vibration remains nearly unaffected upon cooling (Fig.
637 10a,b), indicating no major influence of temperature on the internal molecular structure of the
638 water molecule, thus corroborating its ‘rigid’ character. Accordingly, the acceptor-donor-
639 acceptor angle O2–O3–O2 is rather insensitive to temperature (-1.0° change along 200 K),
640 compared to the significant influence of the Mg/Ni ratio (-3° change).

641
642 **Raman spectra at ambient conditions.** Raman spectra acquired across the Mg-Ni
643 solid solution series also show systematic band shifts with increasing x_{Ni} (Figs. 11a and 12,
644 Table 9). The linear trends and wavenumber shifts of related vibrational phenomena observed
645 in Raman spectra basically correspond to their behavior in FTIR spectra.

646 As expected, the decrease in hydrogen bond length with increasing Ni content reported
647 above, leads to the decrease in the wavenumber of the H₂O symmetric stretching vibration
648 from 3178 to 3071 cm⁻¹ shift and that of the antisymmetric mode from 3388 to 3257 cm⁻¹. The
649 very good match between the wavenumber position observed in the IR transmission spectra
650 (Table 8) and Raman data (Table 9) underline the consistency of both sets of data. The

651 decrease in shift of the antisymmetric $\nu_{3(\text{H}_2\text{O})}$ vibration is well visible in case of the Mg-Ni
652 solid solution (Fig. 12a), in analogy to the FTIR spectra (Fig. 10a), despite its position as a
653 weak shoulder of the dominant $\nu_{1(\text{H}_2\text{O})}$ vibration. This is in sharp contrast to the kieserite-
654 szomolnokite solid solution (Fig. 12b), where no observable trend could be identified for the
655 $\nu_{3(\text{H}_2\text{O})}$ mode (Talla and Wildner 2019). This contrast can be attributed to the much stronger
656 change of the hydrogen bond length between the kieserite and dwornikite endmembers (–
657 0.068 Å) (Fig. 4c) compared to the respective difference between kieserite and szomolnokite
658 (+0.012 Å; Talla and Wildner 2019).

659 The nearly perfect match between the behavior of the sulfate group in the Mg-Ni and
660 Mg-Fe kieserite solid solution, including the absolute positions of the corresponding sulfate-
661 related bands (Fig. 12), is remarkable. This is in agreement with the relaxation of the sulfate
662 tetrahedron (Fig. 3cd, Table 6) both in case of Fe and Ni incorporation (Talla and Wildner
663 2019). On the other hand, the assignment of the band at 500 cm^{-1} to sulfate-related modes by
664 Chio et al. (2007) is disputable, since this band shows opposite behavior in both solid
665 solutions (Fig. 12).

666 The bands observed in the 100–250 cm^{-1} spectral region feature a component at ~220
667 cm^{-1} assigned by Chio et al. (2007) to a vibration involving the translation along the Fe–H₂O
668 bond in szomolnokite. However, its spectral position remains nearly unchanged regardless of
669 the Mg/Fe ratio (Fig. 12b). A much stronger variation (an increase in wavenumber) is found
670 for the corresponding band in case of Ni uptake (Fig. 12a, Table 9). Presuming the
671 involvement of the H₂O molecule in the given Raman mode, more pronounced changes are
672 indeed to be expected in case of the Mg-Ni solid solution, since, as was already mentioned,
673 the hydrogen bonding system undergoes much stronger changes. (Fig. 9a,b, Fig. 12a,b, Tables
674 8 and 9 and Fig. 4c, Table 6). For the lowest mode at ~130 cm^{-1} a contradicting behavior in
675 case of Ni and Fe incorporation is observed. While we are unable to ascertain the exact
676 character of this vibrational mode, this difference suggests the involvement of octahedral

677 modes, since these polyhedra also show opposite structural trends between the Mg-Ni solid
678 solution (decrease in polyhedral volume and Me–O₂,O₃ bond lengths) and the Mg-Fe series
679 (increase of polyhedral volume and all Me–O bonds).

680

681 **Raman spectra at low-temperature conditions.** Band shifts in low-temperature
682 Raman spectra show similar behavior to corresponding vibrational modes in the IR spectra
683 (Table 9), and comparison with previous data for the Mg-Fe solid solution (Talla and Wildner
684 2019) reveal that they are largely independent of the cation substituting for Mg. As expected,
685 the $\nu_{1(\text{H}_2\text{O})}$ vibration shows significant changes towards lower shift values upon cooling, while
686 the sulfate-related bands are rather stable. Better band resolution compared to IR spectra
687 allows to track even the subtle positional changes of these peaks. The highest-energy peak at
688 $\sim 1200 \text{ cm}^{-1}$ of the split $\nu_{3(\text{SO}_4)}$ vibration takes higher shift values, while a decrease in position
689 for the lowest-situated $\nu_{3(\text{SO}_4)}$ peak at $\sim 1100 \text{ cm}^{-1}$ shift can be observed (Table 9). The
690 symmetric stretching vibration $\nu_{1(\text{SO}_4)}$ shows a very minor decrease in wavenumber (Table 9).
691 The octahedra-related mode at $\sim 220 \text{ cm}^{-1}$ exhibits a notable increase in wavenumber at low
692 temperature, as does the lattice band at $\sim 130 \text{ cm}^{-1}$. The complex nature of this low-energy
693 lattice mode does not allow reliable band assignment without ab-initio calculations, beyond
694 the scope and aim of the present study.

695

696 **Use of vibrational spectra to estimate Ni- and Fe-contents of (extraterrestrial)**
697 **kieserite.** Linear trends, observed for changes in structural parameters across the kieserite-
698 dwornikite solid solution (Figs. 2-4), as well as in the position of IR spectral bands (when
699 expressed in wavenumber units) and Raman spectra, clearly correlate with the Mg/Ni ratio
700 (Figs. 9, 11). Aside from documenting in detail the behavior of kieserite-group compounds
701 throughout the Mg-Ni solid solution series relevant to the icy moons of Jupiter and Saturn, the
702 linearity of the data also allows its straightforward use as a standard of comparison in

703 evaluating IR-spectra acquired by orbiters, as well as Raman measurements which could be
704 conducted during eventual rover missions. In this regard, the use of wavelength units (in μm)
705 to express spectral band positions, as it is common in the cosmochemical community, would
706 seem to be more appropriate. However, our preference for – and the benefits of – the use of
707 wavenumber units (cm^{-1}) throughout this work for discussion and depiction of spectroscopic
708 data have been amply discussed in our previous paper (Talla and Wildner 2019).

709 In general, the evaluation of orbiter measurements in the VNIR-MIR spectral range
710 (visible to near/medium infrared) leading to the assessment of the mineral phases present on
711 the surface of celestial bodies and their composition, follows the ‘spectral unmixing’
712 approach, requiring reference endmember spectra, as is the case in the currently conducted
713 investigations of surface spectra from Mars (Cloutis et al. 2006; Combe et al. 2008; Mangold
714 et al. 2008; Bishop et al. 2009; Lichtenberg et al. 2010; Roach et al. 2010; Noel et al. 2015;
715 Liu et al. 2016). In brief, spectra from orbiter measurements are fitted by the least-squares
716 technique using reference endmember spectra to fit the measured signal after its correction for
717 the instrumental function (CRISM smile, etc.), atmospheric scattering (if relevant) and the
718 incidence angle using detailed knowledge of the local topography. While to this date, only
719 pure endmember kieserite spectra were used for this assessment, the discovery of linear trends
720 across the kieserite-dwornikite solid solution presented in this work as well as across the
721 kieserite-szomolnokite solid solution (Talla and Wildner 2019) allows to derive reference
722 spectra for monohydrate sulfates of an intermediate composition and, in principle, to obtain
723 additional information on the chemistry (Mg/Ni ratio in the present case) of the kieserite-
724 group sulfate monohydrate at hand, with semi-quantitative results at least. Two prominent
725 features in the IR spectra, which show a pronounced slope in their correlation with increasing
726 Ni content, are the symmetric and antisymmetric stretching vibration $\nu_{1(\text{H}_2\text{O})}$ and $\nu_{3(\text{H}_2\text{O})}$ of
727 H_2O , with a negative correlation to x_{Ni} , and the ‘diagnostic’ band at $\sim 900 \text{ cm}^{-1}$, the
728 wavenumber of which increases with Ni content in contrast to the situation in the Mg-Fe solid

729 solution (Figs. 9, 12b, Table 8). The fact that the antisymmetric stretching mode $\nu_{3(\text{H}_2\text{O})}$ of the
730 H_2O molecule shows a clear dependence on the Ni content can be regarded as an important
731 benchmark, allowing one to discern between Ni- and Fe-incorporation, since no such trend is
732 observed in case of Fe contents along the Mg-Fe solid solution (Talla and Wildner, 2019).
733 However, the two other promising features to infer the Ni content in the medium-infrared
734 region – $\nu_{1(\text{H}_2\text{O})}$ and the diagnostic band at $\sim 900 \text{ cm}^{-1}$, which show opposite behavior
735 compared to the Mg-Fe solid solution – are also the ones which show the strongest change in
736 wavenumber with decreasing temperature (Fig. 10). The same, considering the temperature
737 sensitivity of the fundamental H_2O bands, will apply to their combination modes in the ~ 4900
738 cm^{-1} region (Fig. 8ab), which, according to the results for kieserite of Jamieson et al. (2014),
739 split further apart with decreasing temperature. This behavior enhances the splitting of those
740 bands in addition to the effect of Ni incorporation, leading to a potential overestimation of x_{Ni} .
741 Despite being the exact opposite to the merging of these bands into a seemingly single broad
742 absorption with a shoulder upon Fe intake (Talla and Wildner 2019), the influence of
743 temperature complicates a reliable assessment of Ni contents in the examined material.

744 Anyway, the knowledge of surface temperatures (at least approximate values) during
745 remote measurements is a prerequisite to allow for meaningful quantitative comparisons,
746 otherwise only semi-quantitative information may be extracted. The rough estimation of
747 temperature is in part possible even from remote sensing spectra in themselves, because they
748 contain spectral regions which slightly shift according to the surface temperature, such as the
749 $\sim 5000 \text{ cm}^{-1}$ ($2 \mu\text{m}$) wavenumber region (Liu et al. 2016). Mid-infrared hyperspectral remote
750 sensing instruments (and bolometers) can provide radiance data. Because these follow Planck
751 behavior, they can easily be converted to emissivity by the application of a Planck function at
752 a given temperature, to provide accurate surface temperature values (for references see e.g.,
753 Ruff et al. 1997). In addition, any rovers, whether in the scope of present (Mars) or future
754 (e.g., the Jovian moons) missions, are typically equipped with thermometers, providing an

755 additional rough temperature value. The attempt to use sulfate-related bands for such
756 considerations is hampered by their wavenumber position being largely insensitive to
757 temperature changes in the IR spectra regardless of the substituting cation (Fe versus Ni), as
758 can be seen in (Fig. 10 a,b). This will likely prevent any exact observation of the subtle band
759 position changes also due to enhanced noise.

760 Additional sources of error arise from the use of different correction datasets preceding
761 the actual spectral unmixing procedure. Different versions of the atmospheric correction
762 model used in the evaluation of CRISM spectra from Mars may lead to significant changes in
763 the form of the resulting spectrum, even causing errors in the discrimination between kieserite
764 and szomolnokite in the same region of interest (Bishop et al. 2009, Noel et al. 2015).

765 A much more promising situation is to be expected considering the high resolution and
766 low FWHM of Raman bands, which aids in ascertaining their position changes from kieserite
767 to dwornikite with much higher precision when the Raman spectrometer is close to the target
768 (i.e, not from an orbital platform). While the superior spectral resolution allows efficient use
769 of the sulfate-related bands, such as the prominent $\nu_{1(\text{SO}_4)}$ band (Fig. 11), it is possible to
770 derive the *total* Mg/Me^{2+} ($\text{Me}^{2+} = \text{Fe}^{2+} + \text{Ni}^{2+}$) content only, because the sulfate modes behave
771 analogically in the case of Ni or Fe uptake, as opposed to the H_2O stretching vibrational
772 modes, that show opposite trends between Ni and Fe incorporation (Fig. 12a,b respectively).
773 These water bands are, however, weaker in the Raman spectra (Fig. 11) and are temperature-
774 sensitive.

775 Further complications arise if a ternary Mg/Fe/Ni solid solution is present. As described
776 above, the major difference in vibrational spectra of the Mg-Fe (Talla and Wildner 2019) and
777 Mg-Ni solid solution (this work) concerns the vibrational modes involving the H_2O molecule
778 (Fig. 9a,b). Structurally induced opposite trends are observed for both H_2O stretching modes.
779 In case of Fe uptake, mainly the symmetric $\nu_{1(\text{H}_2\text{O})}$ mode undergoes changes, shifting to higher
780 wavenumbers as the $\text{O3}\cdots\text{O2}$ donor–acceptor distance increases. This shift would be

781 dampened by the presence of Ni in a ternary solid solution. Fortunately, the antisymmetric
782 $\nu_{3(\text{H}_2\text{O})}$ mode remains constant in position regardless of the Mg/Fe ratio (at $\sim 3380 \text{ cm}^{-1}$ in
783 FTIR transmission mode), whereas a prominent decrease of its position occurs in case of Ni-
784 incorporation, as shown in Fig. 9a,b. This opens up the chance to discriminate between Ni and
785 Fe contents in kieserite: taking the position of the antisymmetric H_2O stretching mode $\nu_{3(\text{H}_2\text{O})}$
786 as reference to derive the Ni content, the Fe content may as well be estimated, based on the
787 difference between the position of the symmetric $\nu_{1(\text{H}_2\text{O})}$ stretching mode expected from the
788 deduced Ni formula content, and its actual position, expectably at somewhat higher
789 wavenumbers (indicative of Fe content).

790 The total content of both transition elements (Ni and Fe) can in theory be verified via
791 the sulfate-related bands in Raman spectra if these are available (preferably using the
792 strongest $\nu_{1(\text{SO}_4)}$ band), because their positions are rather insensitive to temperature (see above
793 and Talla and Wildner 2019). They undergo the same changes to lower shift values,
794 irrespective of the transition element present.

795
796 **Crystal field spectra.** The crystal field (CF) spectrum of dwornikite (Fig. 13) shows
797 that VNIR spectra from orbiter or rover measurements might be influenced by Ni contents of
798 monohydrate sulfates, in particular by the broad first spin-allowed ${}^3\text{A}_{2g}(\text{F}) \rightarrow {}^3\text{T}_{2g}(\text{F})$ band, at
799 ambient conditions centered at $\sim 8100 \text{ cm}^{-1}$ ($\sim 1.24 \mu\text{m}$), but extending from 6400 to 10400
800 cm^{-1} (1.57 to 0.95 μm). Besides, a Superposition Model (SM) calculation applying the
801 ambient SM parameters from Table 10 to the octahedral NiO_6 geometry at $-160 \text{ }^\circ\text{C}$ (Table 7)
802 reveals that temperature-dependent band shifts are expected to be negligible, i.e., 10 cm^{-1} for
803 ${}^3\text{T}_{2g}(\text{F})$ to at most 20 cm^{-1} for the two higher-energetic ${}^3\text{T}_{1g}$ bands.

804 In agreement with the absence of any perceivable band splitting, at least for the first and
805 third spin-allowed band (Fig. 13), the obtained CF and SM parameters do not properly
806 reproduce the structural [4+2] elongation of the NiO_6 octahedron. In fact, with $\text{Dt} = -10 \text{ cm}^{-1}$

807 even a faint polyhedral compression of the pseudotetragonal axis is indicated. This
808 discrepancy can be best explained by the higher CF strength of H₂O ligands forming the
809 elongated axis (compared to the oxygen ligands of the sulfate groups), in this way obviously
810 fully compensating the structural elongation. A similar but less pronounced partial
811 compensation effect was also found in the CF spectra of cobaltkieserite (Wildner 1996) and
812 szomolnokite (Talla and Wildner 2019). The same reasoning may also be applicable to some
813 of the extracted SM parameters, namely t_4 and \bar{B}_2 , which yielded quite unrealistic low values.
814 Albeit, a tendency towards t_4 values much lower than their electrostatic ideal value ($t_4 = 5$)
815 has been observed previously for divalent transition metal compounds, among them also
816 cobaltkieserite (Andrut et al. 2004).

817

818

Implications

819

820 The presented data for the kieserite-dwornikite solid solution can assist cosmochemical
821 investigations, given the presence of kieserite-group sulfates on the surface of Mars and
822 potentially on the Jovian moons, where conditions are favorable for their formation. The
823 linear character of the observed spectral and structural trends at room temperature and the
824 knowledge of their changes upon cooling provides a solid starting basis to draw conclusions
825 about the content of Ni in kieserite based on absorption band positions in IR and Raman
826 spectra. Even in the unfavorable case of a semi-quantitative approach (no knowledge of
827 temperature), zoning of Ni contents in kieserite-containing sediments can be monitored using
828 the presented data. Given the change in the spectral position of the antisymmetric H₂O
829 vibration occurring only upon Ni-incorporation, a good chance to reliably estimate Ni versus
830 Fe contents in a ternary Mg-Fe-Ni-kieserite solid solution presents itself as well, presuming at
831 least a rough knowledge of the surface temperature on the investigated planetary body.

832

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834

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842

843 **References**

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1007 **Table captions / Tables**

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1010 **Table 1.** Composition of the studied samples as shown in Fig. 1, determined by wet chemical
1011 analyses. Deviations of the actual sample Ni content $x_{\text{Ni}(\text{sample})}$ from the preset Mg/Ni ratio in
1012 the batch $x_{\text{Ni}(\text{preset})}$ are also given. The analytical error amounts to 0.005 wt% for Mg and 0.002
1013 wt% for Ni. Note that the Mg/Ni ratios of individual hand-picked crystals used for the single
1014 crystal X-ray diffraction studies (from ‘leaking vessel’ runs) were extracted as variable
1015 parameter in the respective structure refinement runs.

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Sample ID	Mg (wt%)	Ni (wt%)	$x_{\text{Ni}(\text{sample})}$	$x_{\text{Ni}(\text{preset})}$	$x_{\text{Ni}(\text{sample})}-x_{\text{Ni}(\text{preset})}$
Ni5a	17.60	0.68	0.016	0.050	-0.034
Ni10a	17.63	1.11	0.025	0.100	-0.075
Ni15a	17.36	1.61	0.037	0.150	-0.113
Ni20a	15.68	2.49	0.062	0.200	-0.138
Ni30a	14.72	3.98	0.101	0.300	-0.199
Ni40a	14.00	6.05	0.152	0.400	-0.248
Ni100a	0.11	27.29	0.990	1.00	0.001
Ni10aq	16.74	3.20	0.073	0.100	-0.027
Ni20aq	14.55	7.77	0.181	0.200	-0.019
Ni30aq	12.32	9.90	0.250	0.300	-0.050
Ni40aq	10.26	13.92	0.360	0.400	-0.040
Ni50aq	8.40	14.48	0.417	0.500	-0.083
Ni60aq	6.55	16.37	0.508	0.600	-0.092
Ni70aq	4.66	21.96	0.661	0.700	-0.039
Ni80aq	3.07	22.92	0.756	0.800	-0.044
Ni90aq	1.47	22.60	0.864	0.900	-0.036
Ni100aq	0.01	27.85	0.999	1.000	-0.001

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1028 **Table 2.** Crystal data and details of X-ray data collections and structure refinements for
 1029 selected representatives of the $\text{Mg}_{1-x}\text{Ni}_x(\text{SO}_4)\cdot\text{H}_2\text{O}$ solid solution series. Common data:
 1030 monoclinic, space group $C2/c$ (no. 15), $Z = 4$. Respective data for endmember kieserite are
 1031 reproduced from Bechtold and Wildner (2016).
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$\text{Mg}_{1-x}\text{Ni}_x(\text{SO}_4)\cdot\text{H}_2\text{O}$	$\text{Mg}_{1.00}\text{Ni}_{0.00}$	$\text{Mg}_{0.76}\text{Ni}_{0.24}$	$\text{Mg}_{0.48}\text{Ni}_{0.52}$	$\text{Mg}_{0.24}\text{Ni}_{0.76}$	$\text{Mg}_{0.00}\text{Ni}_{1.00}$
x_{Ni}	0.0	0.236(2)	0.525(2)	0.759(2)	1.0
a (Å)	6.910(1)	6.891(1)	6.859(1)	6.845(1)	6.829(1)
b (Å)	7.634(2)	7.622(1)	7.620(1)	7.614(1)	7.605(1)
c (Å)	7.643(2)	7.592(1)	7.545(1)	7.502(1)	7.463(1)
β (°)	118.00(1)	117.98(1)	117.84(1)	117.80(1)	117.75(1)
V (Å ³)	356.00(14)	352.14(9)	348.67(9)	345.84(9)	342.98(9)
μ (mm ⁻¹)	0.97	2.16	3.64	4.86	6.14
D_{calc} (gcm ⁻³)	2.582	2.763	2.980	3.159	3.346
Exposure time (s)/frame	140	180	60	40	40
CCD frames processed	508	497	505	531	503
Frame scale factors $_{\text{max, min}}$	1.11, 0.89	1.16, 0.77	1.21, 0.89	1.20, 0.72	1.18, 0.81
Total number of intensity data	12919	14225	16826	13763	11698
Total number of reflections	8865	8270	7928	8698	8191
Intensity data for unit cell	5101	7662	12062	6952	4887
Number of hkl's	3995	4052	4037	3995	3728
Unique hkl's	1089	1082	1071	1072	1056
R_i (%)	3.20	3.66	2.94	3.53	2.60
$F_o > 4\sigma(F_o)$	937	919	1005	1029	1019
Variables	39	42	42	42	40
wR2 [for all F_o^2] (%)	6.17	7.62	5.03	4.80	4.99
weighting parameters a, b^a	0.028, 0.20	0.021, 0.85	0.019, 0.30	0.024, 0.08	0.024, 0.52
R1 [for $F_o > 4\sigma(F_o)$] (%)	2.27	3.33	2.05	1.91	1.95
R1 [for all F_o] (%)	2.94	4.59	2.28	2.02	2.05
Goodness of fit	1.109	1.170	1.143	1.104	1.077
Extinction coefficient	0	0.002(2)	0.010(2)	0.016(2)	0.023(2)
$\Delta\rho_{\text{max, min}}$ (e ⁻ Å ⁻³)	0.47, -0.51	0.83, -0.44	1.02, -0.55	1.18, -0.77	0.78, -0.97

1033 ^a $w = 1 / [\sigma^2(F_o^2) + (a \times P)^2 + b \times P]$; $P = \{[\text{max of } (0 \text{ or } F_o^2)] + 2F_c^2\} / 3$

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1036 **Table 3.** Structure parameters at room temperature for selected representatives of the
 1037 $\text{Mg}_{1-x}\text{Ni}_x(\text{SO}_4)\cdot\text{H}_2\text{O}$ solid solution series. U_{ij} are given in pm². Wyckoff positions: Mg/Ni on
 1038 $4b$ (sym $\bar{1}$): $0, \frac{1}{2}, 0$ etc.; S and O3 on $4e$ (sym 2): $0, y, \frac{1}{4}$ etc., $U_{23} = U_{12} = 0$; O1, O2 and H on $8f$
 1039 (sym 1): x, y, z etc. Respective data for endmember kieserite are given in Bechtold and Wildner
 1040 (2016).

$\text{Mg}_{1-x}\text{Ni}_x(\text{SO}_4)\cdot\text{H}_2\text{O}$		$\text{Mg}_{0.76}\text{Ni}_{0.24}$	$\text{Mg}_{0.48}\text{Ni}_{0.52}$	$\text{Mg}_{0.24}\text{Ni}_{0.76}$	$\text{Mg}_{0.00}\text{Ni}_{1.00}$
Mg/Ni	x_{Ni}	0.236(2)	0.525(2)	0.759(2)	1.0
	U_{11}	85(2)	69(1)	67(1)	57(1)
	U_{22}	92(2)	86(1)	74(1)	69(1)
	U_{33}	103(2)	84(1)	72(1)	62(1)
	U_{23}	4(1)	3(1)	3(1)	2(1)
	U_{13}	37(2)	29(1)	31(1)	25(1)
	U_{12}	-4(2)	-3(1)	-3(1)	-3(1)
	U_{eq}	96(1)	82(1)	76(1)	64(1)
S	y	0.15535(6)	0.15612(3)	0.15643(3)	0.15664(4)
	U_{11}	71(2)	56(1)	53(1)	44(1)
	U_{22}	89(2)	76(1)	62(1)	59(1)
	U_{33}	111(2)	92(1)	77(1)	67(1)
	U_{13}	36(1)	26(1)	28(1)	23(1)
	U_{eq}	93(1)	78(1)	65(1)	58(1)
O1	x	0.17513(17)	0.17421(10)	0.17356(10)	0.17289(12)
	y	0.04528(15)	0.04590(10)	0.04637 (9)	0.04649(11)
	z	0.39678(17)	0.39953(10)	0.40142(10)	0.40331(12)
	U_{11}	100(4)	82(2)	78(2)	71(2)
	U_{22}	165(4)	160(3)	150(2)	142(3)
	U_{33}	178(5)	164(3)	150(2)	136(3)
	U_{23}	67(4)	68(2)	70(2)	69(2)
	U_{13}	51(3)	42(2)	43(2)	38(2)
	U_{12}	40(3)	37(2)	37(2)	37(2)
	U_{eq}	154(2)	142(1)	130(1)	121(1)
O2	x	0.09460(18)	0.09821(10)	0.10079(10)	0.10307(12)
	y	0.26806(14)	0.26909 (9)	0.26961 (8)	0.26994 (9)
	z	0.15065(17)	0.15246(10)	0.15372(10)	0.15466(11)
	U_{11}	133(4)	118(2)	127(2)	95(2)
	U_{22}	125(4)	117(2)	114(2)	95(2)
	U_{33}	186(5)	164(3)	170(3)	135(3)
	U_{23}	49(3)	43(2)	51(2)	43(2)
	U_{13}	94(4)	82(2)	94(2)	71(2)
	U_{12}	11(3)	14(2)	16(2)	14(2)
	U_{eq}	140(2)	127(1)	113(1)	101(1)
O3	y	0.63411(21)	0.63291(12)	0.63209(11)	0.63135(13)
	U_{11}	104(6)	98(3)	91(3)	79(3)
	U_{22}	140(6)	123(3)	109(3)	103(3)
	U_{33}	131(6)	108(3)	94(3)	82(3)
	U_{13}	52(5)	47(3)	45(2)	37(3)
	U_{eq}	126(3)	110(1)	97(1)	88(1)
H	x	0.092(4)	0.111(3)	0.111(3)	0.108(4)
	y	0.688(4)	0.700(3)	0.695(3)	0.694(3)
	z	0.286(4)	0.294(3)	0.291(3)	0.296(3)
	U_{iso}	262(68)	334(50)	196(40)	201(50)

1041 **Table 4.** Crystal data and details of selected temperature-dependent X-ray data collections
 1042 and structure refinements for dwornikite, Ni(SO₄)·H₂O. Common data: monoclinic, space
 1043 group *C2/c* (no. 15), *Z* = 4; each 930 CCD frames measured; exposure time/frame 10 s; 40
 1044 variables. Respective temperature-dependent data for endmember kieserite are given in Talla
 1045 and Wildner (2019), its room temperature data in Bechtold and Wildner (2016), and those for
 1046 dwornikite in Table 2.

Ni(SO ₄)·H ₂ O	0 °C	-80 °C	-160 °C
<i>a</i> (Å)	6.825(1)	6.812(1)	6.804(1)
<i>b</i> (Å)	7.607(1)	7.612(1)	7.618(1)
<i>c</i> (Å)	7.459(1)	7.444(1)	7.437(1)
β (°)	117.72(1)	117.63(1)	117.58(1)
<i>V</i> (Å ³)	342.79	341.98	341.66
μ (mm ⁻¹)	6.14	6.16	6.16
<i>D</i> _{calc} (gcm ⁻³)	3.348	3.356	3.359
Frame scale factors _{max, min}	0.75, 0.64	0.75, 0.66	0.75, 0.66
Total number of reflections	11574	11491	11115
Intensity data for unit cell	8253	8457	8497
Unique hkl's	1064	1063	1061
<i>R</i> _i (%)	2.40	2.32	2.35
<i>F</i> _o > 4σ(<i>F</i> _o)	1031	1036	1034
wR2 [for all <i>F</i> _o ²] (%)	4.11	4.20	4.22
weighting parameters <i>a</i> , <i>b</i> ^a	0.021, 0.34	0.021, 0.47	0.021, 0.53
R1 [for <i>F</i> _o > 4σ(<i>F</i> _o)] (%)	1.54	1.60	1.61
R1 [for all <i>F</i> _o] (%)	1.62	1.67	1.66
Goodness of fit	1.127	1.118	1.123
Extinction coefficient	0.025(1)	0.021(1)	0.019(1)
$\Delta\rho$ _{max, min} (e ⁻ Å ⁻³)	0.67, -0.68	0.80, -0.72	0.82, -0.67

1047 ^a $w = 1 / [\sigma^2(F_o^2) + (a \times P)^2 + b \times P]$; $P = \{[\max \text{ of } (0 \text{ or } F_o^2)] + 2F_c^2\} / 3$

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 1057 **Table 5.** Selected temperature-dependent structure parameters for dwornikite, Ni(SO₄)·H₂O.
 1058 *U*_{ij} are given in pm². Wyckoff positions: Ni on 4*b* (sym $\bar{1}$): 0, 1/2, 0 etc.; S and O3 on 4*e* (sym
 1059 2): 0, *y*, 1/4 etc., *U*₂₃ = *U*₁₂ = 0; O1, O2 and H on 8*f* (sym 1): *x*, *y*, *z* etc. Respective room-
 1060 temperature data are given in Table 3.

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Ni(SO ₄)·H ₂ O		0 °C	-80 °C	-160 °C
Ni	<i>U</i> ₁₁	52(1)	39(1)	27(1)
	<i>U</i> ₂₂	63(1)	47(1)	32(1)
	<i>U</i> ₃₃	55(1)	41(1)	28(1)
	<i>U</i> ₂₃	2(1)	2(1)	1(1)
	<i>U</i> ₁₃	22(1)	17(1)	13(1)
	<i>U</i> ₁₂	-2(1)	-1(1)	-1(1)
	<i>U</i> _{eq}	58(1)	43(1)	29(1)
S	<i>y</i>	0.15681(3)	0.15727(3)	0.15763(4)
	<i>U</i> ₁₁	40(1)	31(1)	23(1)
	<i>U</i> ₂₂	54(1)	42(1)	30(1)
	<i>U</i> ₃₃	60(1)	45(1)	32(1)
	<i>U</i> ₁₃	20(1)	16(1)	12(1)
	<i>U</i> _{eq}	53(1)	40(1)	28(1)
O1	<i>x</i>	0.17279(10)	0.17262(10)	0.17257(10)
	<i>y</i>	0.04668(9)	0.04732(9)	0.04772(9)
	<i>z</i>	0.40353(10)	0.40413(10)	0.40473(9)
	<i>U</i> ₁₁	65(2)	52(2)	39(2)
	<i>U</i> ₂₂	132(3)	100(2)	70(2)
	<i>U</i> ₃₃	125(2)	94(2)	67(2)
	<i>U</i> ₂₃	63(2)	46(2)	32(2)
	<i>U</i> ₁₃	36(2)	28(2)	20(2)
	<i>U</i> ₁₂	32(2)	25(2)	18(2)
<i>U</i> _{eq}	111(1)	84(1)	60(1)	
O2	<i>x</i>	0.10339(10)	0.10428(10)	0.10507(10)
	<i>y</i>	0.27006(8)	0.27059(8)	0.27108(8)
	<i>z</i>	0.15477(9)	0.15519(9)	0.15552(9)
	<i>U</i> ₁₁	89(2)	67(2)	51(2)
	<i>U</i> ₂₂	86(2)	66(2)	47(2)
	<i>U</i> ₃₃	122(2)	94(2)	67(2)
	<i>U</i> ₂₃	38(2)	26(2)	18(2)
	<i>U</i> ₁₃	66(2)	51(2)	38(2)
	<i>U</i> ₁₂	12(2)	7(2)	3(2)
<i>U</i> _{eq}	93(1)	70(1)	51(2)	
O3	<i>y</i>	0.63108(12)	0.63077(12)	0.63055(12)
	<i>U</i> ₁₁	76(3)	60(3)	42(3)
	<i>U</i> ₂₂	92(3)	70(3)	56(3)
	<i>U</i> ₃₃	76(3)	61(3)	50(3)
	<i>U</i> ₁₃	37(2)	29(2)	22(2)
	<i>U</i> _{eq}	81(1)	63(1)	49(1)
H	<i>x</i>	0.108(3)	0.109(3)	0.109(3)
	<i>y</i>	0.695(2)	0.695(3)	0.695(3)
	<i>z</i>	0.294(3)	0.294(3)	0.294(3)
	<i>U</i> _{iso}	197(43)	204(48)	200(50)

1063 **Table 6.** Survey of crystal chemical data for selected representatives of the
 1064 $Mg_{1-x}Ni_x(SO_4) \cdot H_2O$ solid solution series: bond lengths (Å) and angles (°), polyhedral volumes
 1065 (Å³), bond strengths (without H atoms; calculated according to Brese & O’Keeffe, 1991), and
 1066 polyhedral distortion parameters (Brown & Shannon, 1973; Robinson *et al.*, 1971; Griffen &
 1067 Ribbe, 1979). Data for endmember kieserite are reproduced from Bechtold and Wildner
 1068 (2016).

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$Mg_{1-x}Ni_x(SO_4) \cdot H_2O$	$Mg_{1.00}Ni_{0.00}$	$Mg_{0.76}Ni_{0.24}$	$Mg_{0.48}Ni_{0.52}$	$Mg_{0.24}Ni_{0.76}$	$Mg_{0.00}Ni_{1.00}$
x_{Ni}	0.0	0.236(2)	0.525(2)	0.759(2)	1
Me–O1 (2×)	2.0216(7)	2.0245(11)	2.0267(7)	2.0304(7)	2.0334(8)
Me–O2 (2×)	2.0414(7)	2.0393(11)	2.0370(7)	2.0349(6)	2.0327(7)
Me–O3 (2×)	2.1714(6)	2.1558(8)	2.1409(5)	2.1281(5)	2.1162(5)
<Me–O>	2.0782	2.0732	2.0682	2.0645	2.0608
$\Delta_{oct} \times 10^3$	1.023	0.802	0.622	0.476	0.362
Σ v.u. (Me)	2.15	2.12	2.08	2.04	2.01
O1–Me–O2 (2×) ^a	85.91(3)	85.77(5)	85.52(3)	85.35(3)	85.15(3)
O1–Me–O3 (2×) ^a	88.50(2)	87.98(3)	87.56(2)	87.18(2)	86.79(2)
O2–Me–O3 (2×) ^a	87.20(3)	87.17(5)	87.30(3)	87.28(3)	87.26(2)
σ_{oct}^2	9.75	10.90	12.11	13.44	15.03
V_{oct}	11.901(10)	11.814(11)	11.727(7)	11.660(7)	11.592(8)
S–O1 (2×)	1.4639(7)	1.4629(11)	1.4652(7)	1.4646(6)	1.4651(7)
S–O2 (2×)	1.4817(7)	1.4810(11)	1.4830(6)	1.4850(6)	1.4868(7)
<S–O>	1.4728	1.4720	1.4741	1.4748	1.4760
BLDP $\times 10^3$	6.98	7.10	6.99	7.99	8.49
Σ v.u. (S)	6.02	6.03	6.00	5.99	5.97
O1–S–O1’	109.94(6)	110.01(9)	110.05(6)	110.19(6)	110.26(7)
O1–S–O2 (2×)	108.59(4)	108.53(6)	108.59(4)	108.46(3)	108.43(4)
O1–S–O2’(2×)	110.41(4)	110.34(7)	110.29(4)	110.34(4)	110.28(4)
O2–S–O2’	108.89(6)	109.09(9)	109.03(6)	109.05(5)	109.16(6)
σ_{tetr}^2	0.77	0.74	0.68	0.85	0.84
V_{tetr}	1.639(2)	1.636(3)	1.643(2)	1.646(2)	1.650(2)
Me–O1–S	140.38(4)	139.44(7)	138.34(4)	137.62(4)	136.82(5)
Me–O2–S	134.31(4)	133.42(7)	132.42(4)	131.63(4)	130.89(4)
Me–O3–Me	123.27(5)	123.39(7)	123.53(4)	123.59(4)	123.67(5)
Σ v.u. (O1)	1.95	1.94	1.92	1.91	1.90
Σ v.u. (O2)	1.86	1.85	1.84	1.82	1.81
Σ v.u. (O3)	0.55	0.56	0.56	0.57	0.57
O3···O2	2.7449(8)	2.7273(13)	2.7052(8)	2.6913(8)	2.6774(9)
O3–H	0.81(2)	0.69(3)	0.84(2)	0.83(2)	0.81(2)
O2···H	2.01(2)	2.07(3)	1.92(2)	1.90(2)	1.92(2)
O3–H···O2	151(2)	158(3)	155(2)	159(2)	156(2)

1070 ^a plus corresponding obtuse angles

1071 **Table 7.** Survey of selected temperature-dependent crystal chemical data for dwornikite,
 1072 Ni(SO₄)·H₂O: bond lengths (Å) and angles (°), polyhedral volumes (Å³), bond strengths
 1073 (without H atoms, calculated according to Brese and O’Keeffe 1991), and polyhedral
 1074 distortion parameters (Brown and Shannon 1973; Robinson et al. 1971; Griffen and Ribbe
 1075 1979). Respective room-temperature data are given in Table 6, those for endmember kieserite
 1076 in Bechtold and Wildner (2016); temperature-dependent data for kieserite are given in Talla
 1077 and Wildner (2019).

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Ni(SO ₄)·H ₂ O	0 °C	−80 °C	−160 °C
Ni–O1 (2×)	2.0338(7)	2.0336(7)	2.0336(7)
Ni–O2 (2×)	2.0328(6)	2.0317(6)	2.0309(6)
Ni–O3 (2×)	2.1145(5)	2.1106(5)	2.1085(5)
<Ni–O>	2.0604	2.0586	2.0577
Δ _{oct} ×10 ³	0.345	0.319	0.305
Σ v.u. (Ni)	2.01	2.02	2.02
O1–Ni–O2 (2×) ^a	85.15(3)	85.15(3)	85.13(3)
O1–Ni–O3 (2×) ^a	86.78(2)	86.82(2)	86.81(2)
O2–Ni–O3 (2×) ^a	87.28(3)	87.31(3)	87.33(3)
σ _{oct} ²	15.01	14.86	14.92
V _{oct}	11.585(7)	11.557(7)	11.540(7)
S–O1 (2×)	1.4654(6)	1.4657(7)	1.4672(7)
S–O2 (2×)	1.4869(6)	1.4872(6)	1.4888(6)
<S–O>	1.47615	1.47645	1.4780
BLDP×10 ³	8.41	8.41	8.44
Σ v.u. (S)	5.97	5.96	5.94
O1–S–O1'	110.26(6)	110.36(6)	110.41(6)
O1–S–O2 (2×)	108.44(3)	108.45(4)	108.46(4)
O1–S–O2' (2×)	110.25(4)	110.24(4)	110.23(4)
O2–S–O2'	109.19(5)	109.10(5)	109.04(5)
σ _{tetr} ²	0.81	0.84	0.85
V _{tetr}	1.650(2)	1.651(2)	1.657(2)
Ni–O1–S	136.75(4)	136.63(4)	136.47(4)
Ni–O2–S	130.79(4)	130.57(4)	130.38(4)
Ni–O3–S	123.73(4)	123.72(4)	123.72(4)
Σ v.u. (O1)	1.89	1.89	1.89
Σ v.u. (O2)	1.81	1.81	1.80
Σ v.u. (O3)	0.58	0.58	0.59
O3···O2	2.6760(8)	2.6706(8)	2.6667(8)
O3–H	0.81(2)	0.82(2)	0.82(2)
O2···H	1.91(2)	1.90(2)	1.89(2)
O3–H···O2	156 (2)	157(2)	157(2)

1079 ^a plus corresponding obtuse angles

1080 **Table 8.** Wavenumber positions of relevant absorption phenomena in FTIR spectra in each
 1081 measuring mode for the kieserite and dwornikite endmembers, Mg/Ni(SO₄)·H₂O, at room
 1082 temperature. Linear regression coefficients of the correlation between the wavenumber
 1083 position of each absorption feature and the Ni-content (x_{Ni}) are listed. Positions of weak yet
 1084 important bands observed in the region of H₂O combination modes are given in the scope of
 1085 DRIFT measurements on undiluted sample material.

Transmission (1:300 sample dilution in KBr)				
Vibration	Wavenumber position (cm⁻¹)		Linear regression coefficients[†]	
	kieserite	dwornikite	a	b
Peak 900	884(1)	941(2)	52(3)	887(1)
ν ₁ (SO ₄)	1043(1)	1019(1)	-26(2)	1041(1)
ν ₃ (SO ₄) _{center}	1165(1)	1135(1)	-24(3)	1164(1)
ν ₂ (H ₂ O)	1525(4)	1522(2)	-4(1)	1526(1)
ν ₁ (H ₂ O)	3182(4)	3057(5)	-134(7)	3191(3)
ν ₃ (H ₂ O)	3367(12)	3262(21)	-117(12)	3364(5)
Attenuated Total Reflectance (ATR)				
Vibration	Wavenumber position (cm⁻¹)		Linear regression coefficients[†]	
	kieserite	dwornikite	a	b
Peak 900	866(3)	904(6)	35(4)	860(2)
ν ₂ (H ₂ O)	1520(1)	1517(2)	-3(4)	1520(1)
ν ₁ (H ₂ O)	3165(4)	3037(5)	-127(3)	3163(1)
ν ₃ (H ₂ O)	3345(8)	3292(22)	-37(6)	3347(3)
Diffuse Reflectance (1:20 sample dilution in KBr)				
Vibration	Wavenumber position (cm⁻¹)		Linear regression coefficients[†]	
	kieserite	dwornikite	a	b
Peak 900	886(3)	938(2)	53(2)	890.6(8)
ν ₁ (SO ₄)	1044(2)	1020(1)	-26.6(9)	1044(1)
ν ₃ (SO ₄) _{center}	1172(3)	1143(4)	-24(3)	1166(2)
ν ₂ (H ₂ O)	1523(1)	1521(1)	2.8(4)	1523(1)
ν ₁ (H ₂ O)	3180(2)	3062(8)	-121(5)	3188(3)
ν ₃ (H ₂ O)	3365(13)	3265(20)	-88(5)	3353(3)
Diffuse Reflectance (non-diluted sample)				
Vibration	Wavenumber position (cm⁻¹)		Linear regression coefficients[†]	
	kieserite	dwornikite	a	b
Peak 600	690(1)	696(2)	8(1)	687.4(6)
Reststrahlenband	1316(6)	1286(3)	-31(4)	1310(2)
ν ₂ (H ₂ O)	1526(3)	1525(1)	-1.4(9)	1527(1)
ν ₁ (H ₂ O)	3203(20)	3099(6)	-107(4)	3207(2)
ν ₃ (H ₂ O)	3412(23)	3314(19)	-129(48)	3432(22)
Peak 4700 cm ⁻¹	4688(4)	4566(4)	-123(4)	4690(2)
Peak 4850 cm ⁻¹	4845(3)	4735(5)	-92(6)	4830(3)
Peak 5090 cm ⁻¹	5087(10)	5110(25)	13(7)	5085(3)

1086 [†]Linear regression equation $y = a(x_{Ni}) + b$, where y denotes the wavenumber position of the
 1087 vibration mode
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1089 **Table 9.** Positions of relevant bands in Raman spectra for the kieserite and dwornikite
 1090 endmembers, Mg/Ni(SO₄)·H₂O. Linear regression coefficients of the correlation between the
 1091 band position and the Ni-content (x_{Ni}) are listed. Additionally, changes in the band position
 1092 with decreasing temperature are given for all relevant Raman bands.
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Peak assignment	Raman shift position at RT (cm ⁻¹)		Linear regression coefficients ^a		Band shift with temperature decrease (×10 ⁻² cm ⁻¹ / °C)
	kieserite	dwornikite	a	b	
Lattice modes	125.3(1)	136.7(8)	12.4(8)	124.3(4)	0.026
$\nu_{transl} Fe-H_2O$	219.2(1)	245.57(7)	26(1)	219.6(6)	0.013
$\nu_2(SO_4)_1$	432.2(1)	418.75(6)	-13.8(6)	432.6(3)	-0.004
$\nu_2(SO_4)_2$	502.1(1)	519.0(5)	16.8(6)	501.7(3)	0.019
$\nu_4(SO_4)_1$	628.7(4)	622.0(2)	-7.1(4)	628.2(2)	0.006
$\nu_4(SO_4)_2$	633.8(2)	629.7(1)	-3.5(3)	633.0(1)	-0.003
$\nu_1(SO_4)$	1041.8(1)	1020.88(1)	-23(1)	1042(1)	-0.003
$\nu_3(SO_4)_1$	1101.6(1)	1076.06(6)	-27(1)	1101(1)	-0.012
$\nu_3(SO_4)_2$	1118.9(1)	1093.87(4)	-27(1)	1119(1)	0.011
$\nu_3(SO_4)_3$	1216.7(1)	1197.69(7)	-20.4(9)	1216(1)	0.061
$\nu_2(H_2O)$	1508.3(1)	1501.2(8)	8(5)	1496(2)	0.018
$\nu_1(H_2O)$	3178(2)	3071(5)	-101(7)	3179(3)	-0.481
$\nu_3(H_2O)$	3388(4)	3257(25)	-82(9)	3361(4)	0.162 ^b

1094 ^aLinear regression equation $y = \mathbf{a}(x_{Ni}) + \mathbf{b}$, where y denotes the Raman shift position

1095 ^bhigh uncertainty

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1111 **Table 10.** Summary of results from classical crystal field (CF) and superposition model (SM)
1112 calculations for Ni²⁺ in dwornikite, based on observed transitions energies as indicated in Fig.
1113 13. Fixed SM parameters are $R_0 = 2.06 \text{ \AA}$, $t_2 = 3$. All given values are in cm^{-1} except the
1114 nephelauxetic ratio β . Racah B₀ is taken from Figgis and Hitchman (2000).

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Szomolnokite	SM (triclinic)	'classical' CF (tetragonal)
Dq _{cub}	–	810
Dq _{eq}	–	804
Dt	–	–10 [†]
Ds	–	–120
Racah B	920	920
Racah C	4.2B*	3860
β (B ₀ = 1042 cm^{-1})	0.88	0.88
Dq _{cub} (from s ₄)	813	810
\bar{B}_4	4920	–
t_4	0.8 [†]	–
\bar{B}_2	250 [†]	–
s ₄	7448	7424
s ₂	51	376

1116 * fixed

1117 † see text

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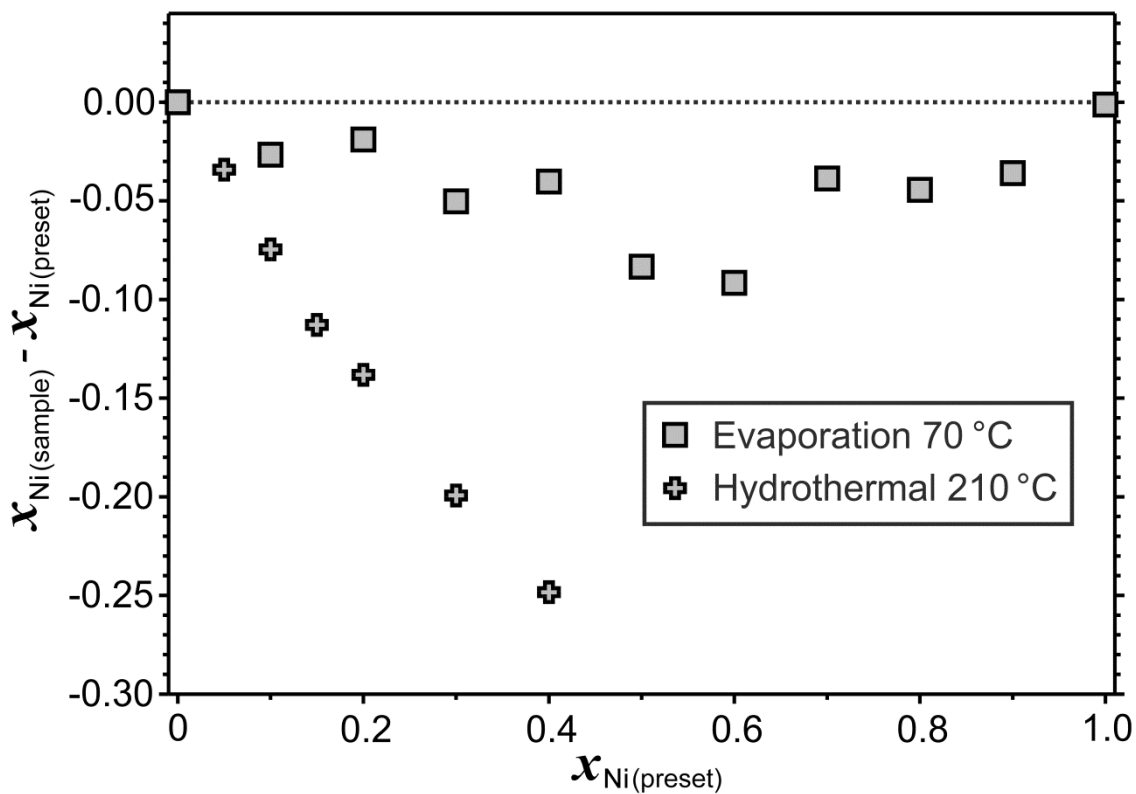
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1131 **Figures and Figure captions**

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1133 **Figure 1.** Deviations between actual Ni-content of the $\text{Mg}_{1-x}\text{Ni}_x(\text{SO}_4)\cdot\text{H}_2\text{O}$ solid solution
1134 samples used for spectroscopic investigations $x_{\text{Ni}(\text{sample})}$ as determined by the wet chemical
1135 analyses and the preset Mg/Ni ratio in the batch $x_{\text{Ni}(\text{preset})}$. Errors are equal or smaller than the
1136 symbol size. Note the strong Ni deficiency in the product with the hydrothermal technique in
1137 use (tightly sealed autoclave). This deviation led to the use the ‘leaking vessel approach’ for
1138 synthesis of single crystals (not depicted), with the actual x_{Ni} ratio refined from the single-
1139 crystal X-ray diffraction results.

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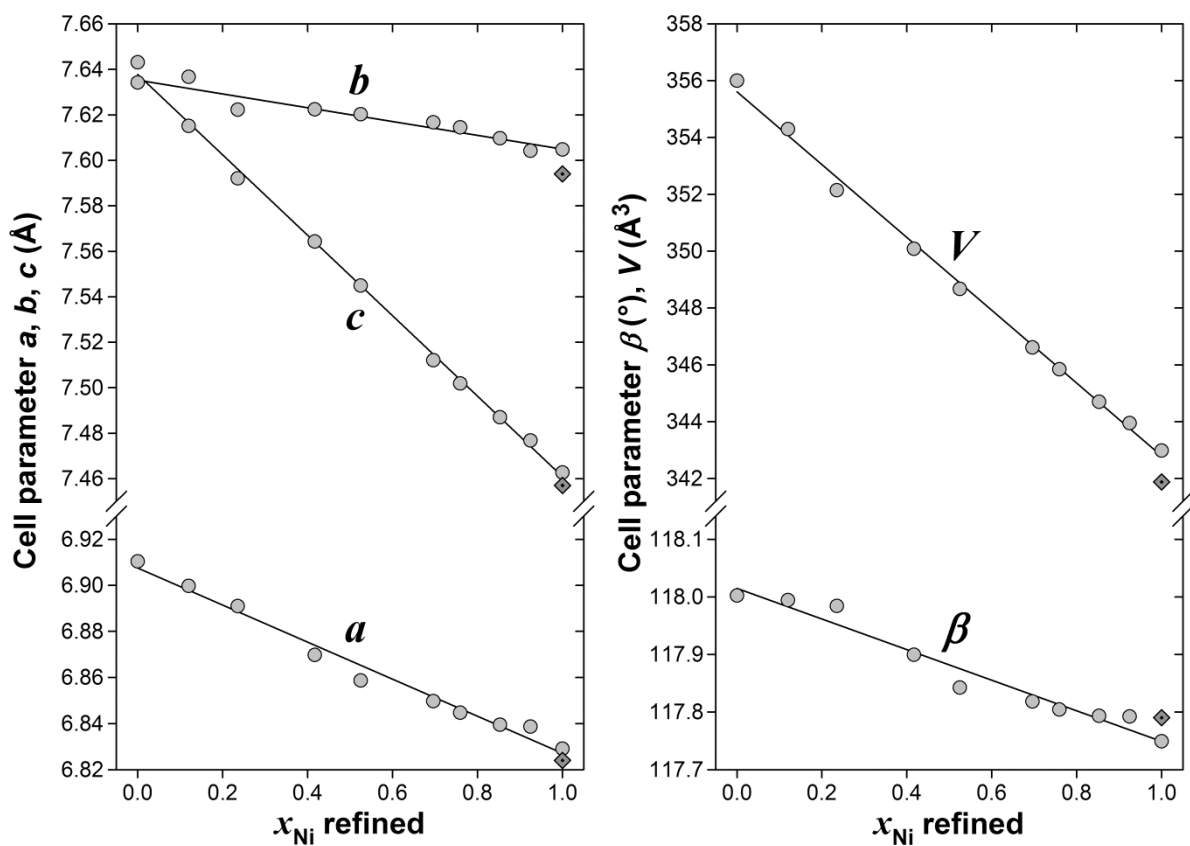
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1148 **Figure 2.** Variation of the lattice parameters (a) a , b , c and (b) β and V along the
1149 $\text{Mg}_{1-x}\text{Ni}_x(\text{SO}_4)\cdot\text{H}_2\text{O}$ solid solution series with linear regression lines. Errors are equal or
1150 smaller than the symbol size. The data for endmember kieserite are taken from Bechtold and
1151 Wildner (2016). Previous data for dwornikite (Wildner and Giester 1991) are shown (without
1152 errors) as dotted diamond symbols.



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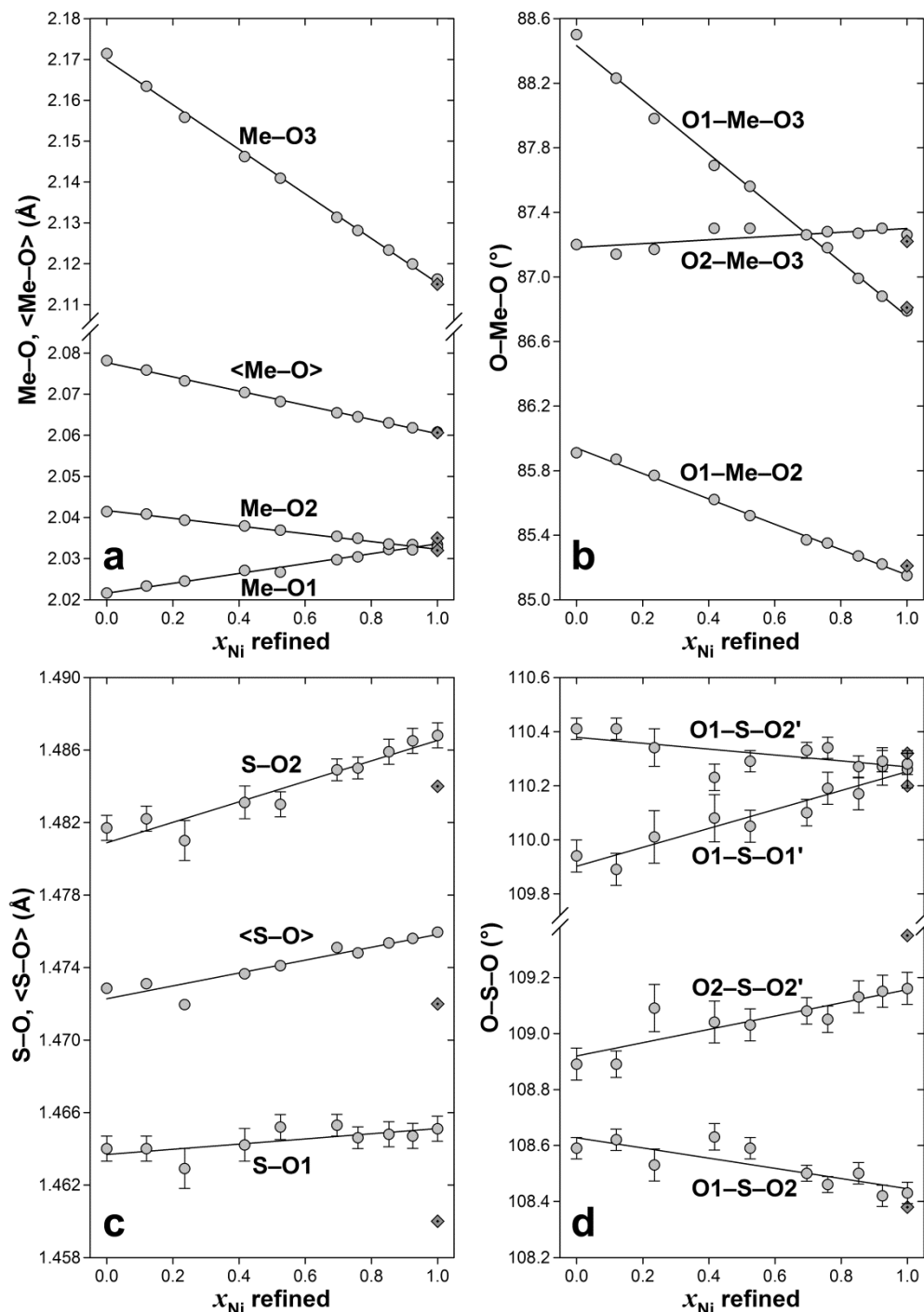
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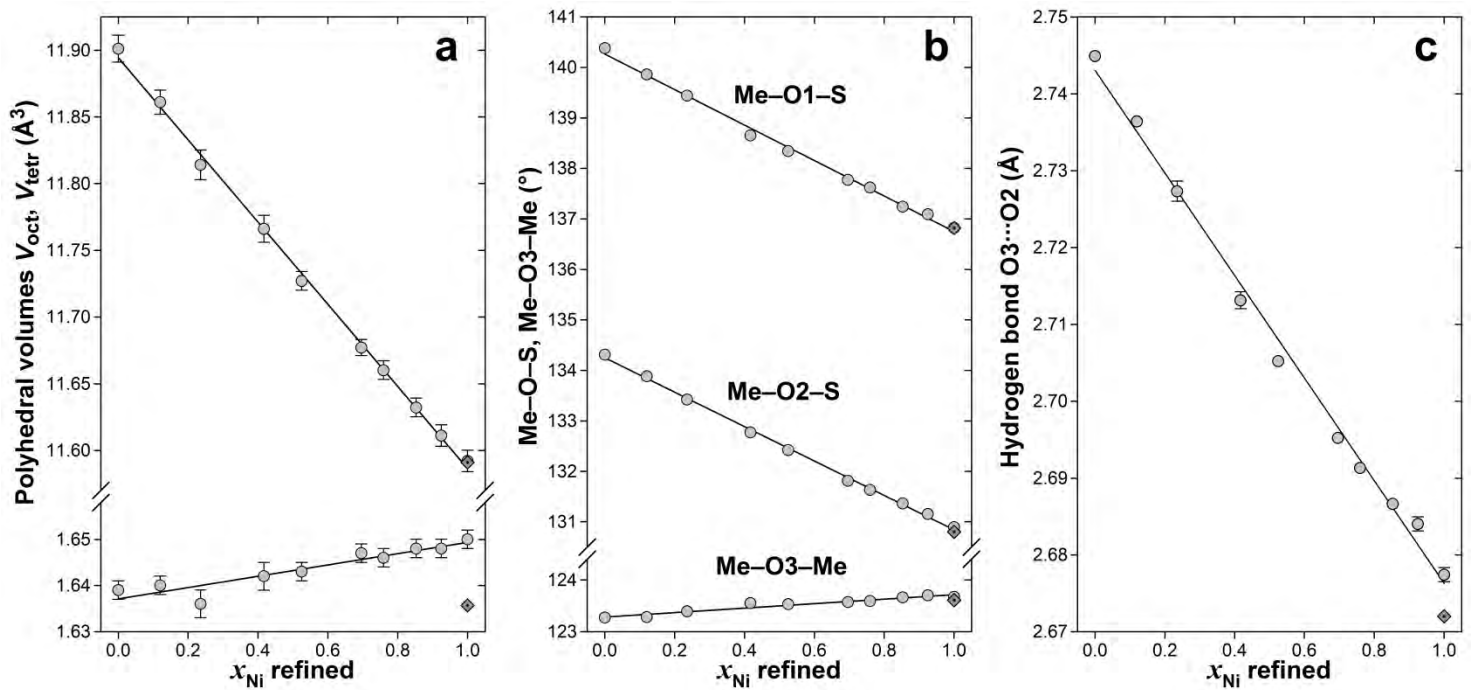
1163 **Figure 3.** Polyhedral geometries along the $\text{Mg}_{1-x}\text{Ni}_x(\text{SO}_4)\cdot\text{H}_2\text{O}$ solid solution series with
1164 linear regression lines; **(a)** octahedral Me–O bond lengths and **(b)** O–Me–O angles, **(c)**
1165 tetrahedral S–O bond lengths and **(d)** O–S–O angles, where Me represents the metal cation. If
1166 not indicated, errors are equal or smaller than the symbol size. The data for endmember
1167 kieserite are taken from Bechtold and Wildner (2016). Previous data for dwornikite (Wildner
1168 and Giester 1991) are shown (without errors) as dotted diamond symbols.



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1170 **Figure 4.** (a) Polyhedral volumes, (b) Me–O–S and Me–O–Me angles and (c) hydrogen bond
 1171 lengths along the $\text{Mg}_{1-x}\text{Ni}_x(\text{SO}_4)\cdot\text{H}_2\text{O}$ solid solution series with linear regression lines. For
 1172 (a), note the strongly different scales for the tetrahedral and octahedral volumes. If not
 1173 indicated, errors are equal or smaller than the symbol size. The data for endmember kieserite
 1174 are taken from Bechtold and Wildner (2016). Previous data for dwornikite (Wildner and
 1175 Giester 1991) are shown (without errors) as dotted diamond symbols.

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1195 **Figure 5.** Variation of the lattice parameters (a) a , b , c and (b) β and V , as well as (c)

1196 polyhedral volumes (V_{tetr} not corrected for thermal motion) for dwornikite within the

1197 investigated temperature range, with second-order regression lines. If not indicated, errors are

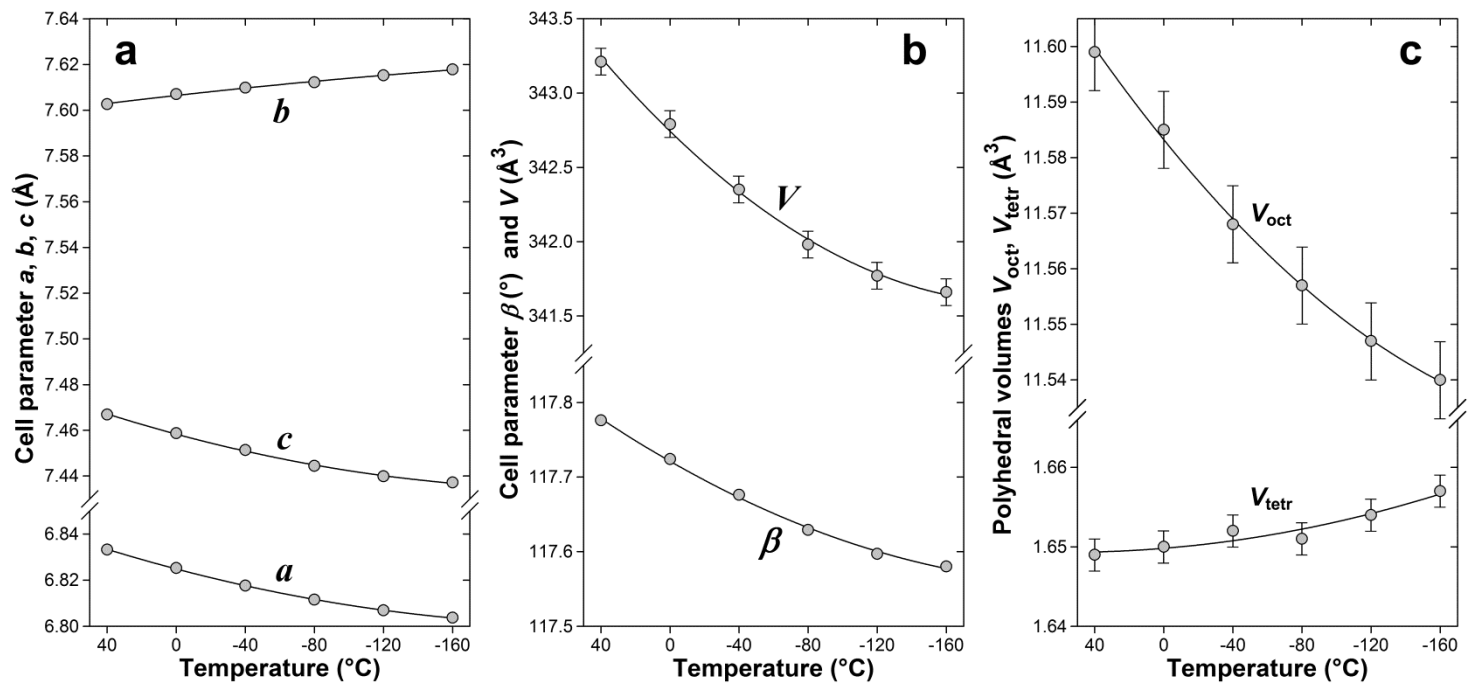
1198 equal or smaller than the symbol size. For (c), note the somewhat different scales for the

1199 tetrahedral and octahedral volumes.

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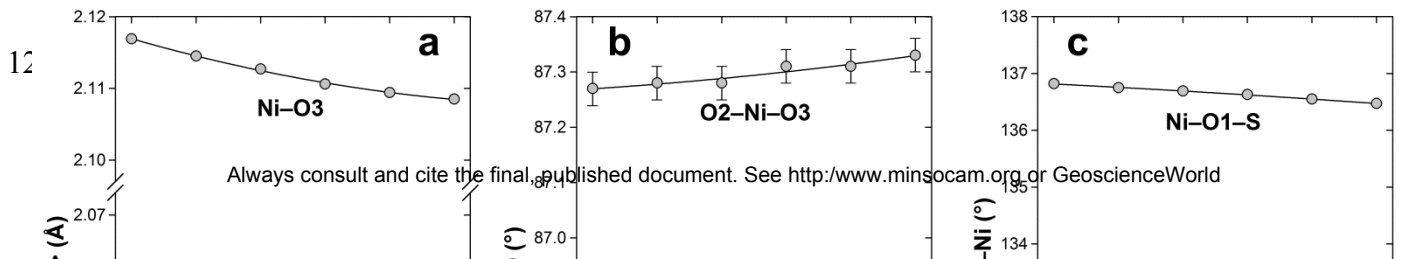
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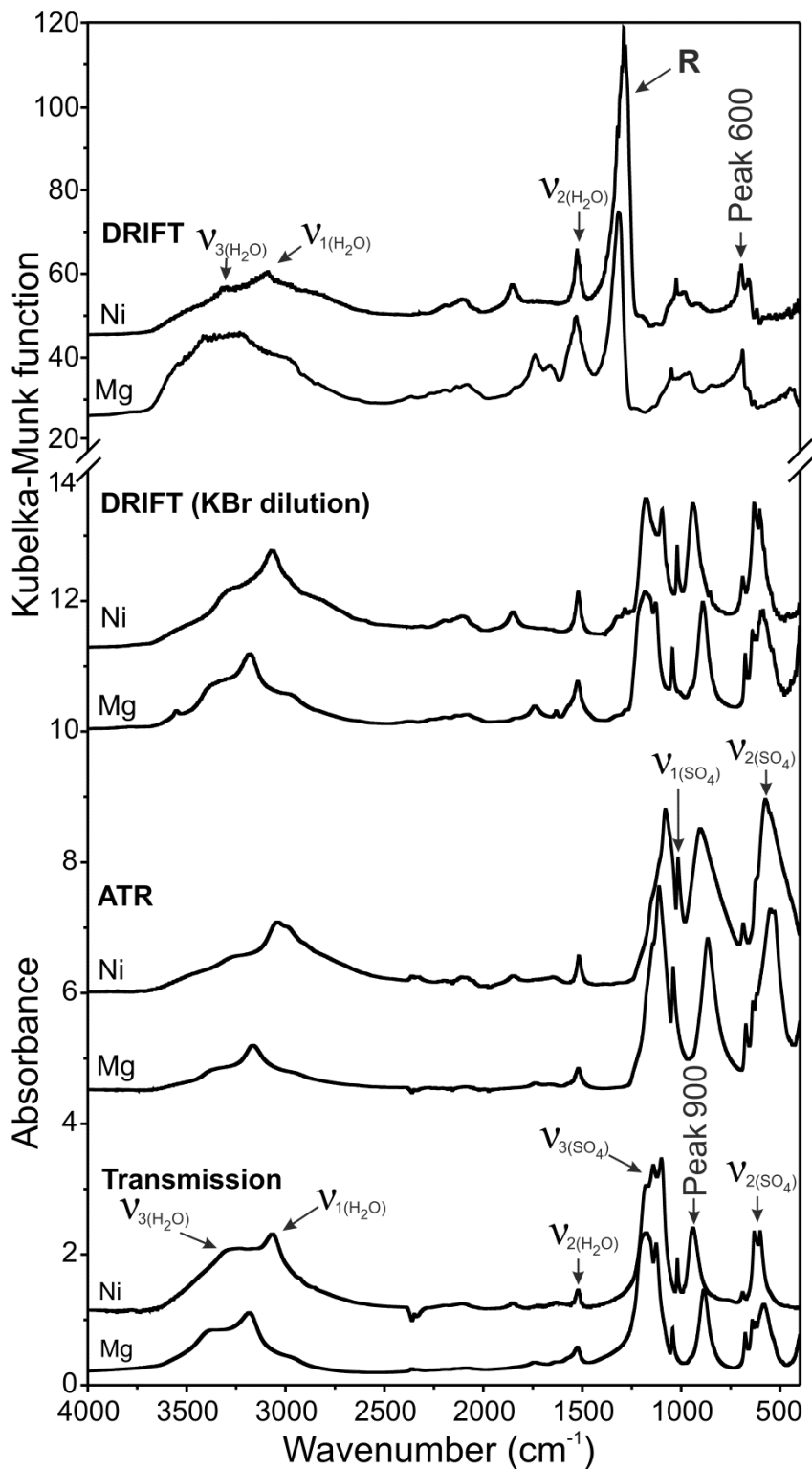
1221 **Figure 6.** Variation of the polyhedral geometries (a) octahedral Me–O bond lengths and (b)
 1222 O–Me–O angles, (d) tetrahedral S–O bond lengths (uncorrected for thermal motion) and (e)
 1223 O–S–O angles, as well as of (c) Ni–O–S and Ni–O–Ni angles and (f) the hydrogen bond
 1224 length in dwornikite within the investigated temperature range, with second-order regression
 1225 lines. If not indicated, errors are equal or smaller than the symbol size.

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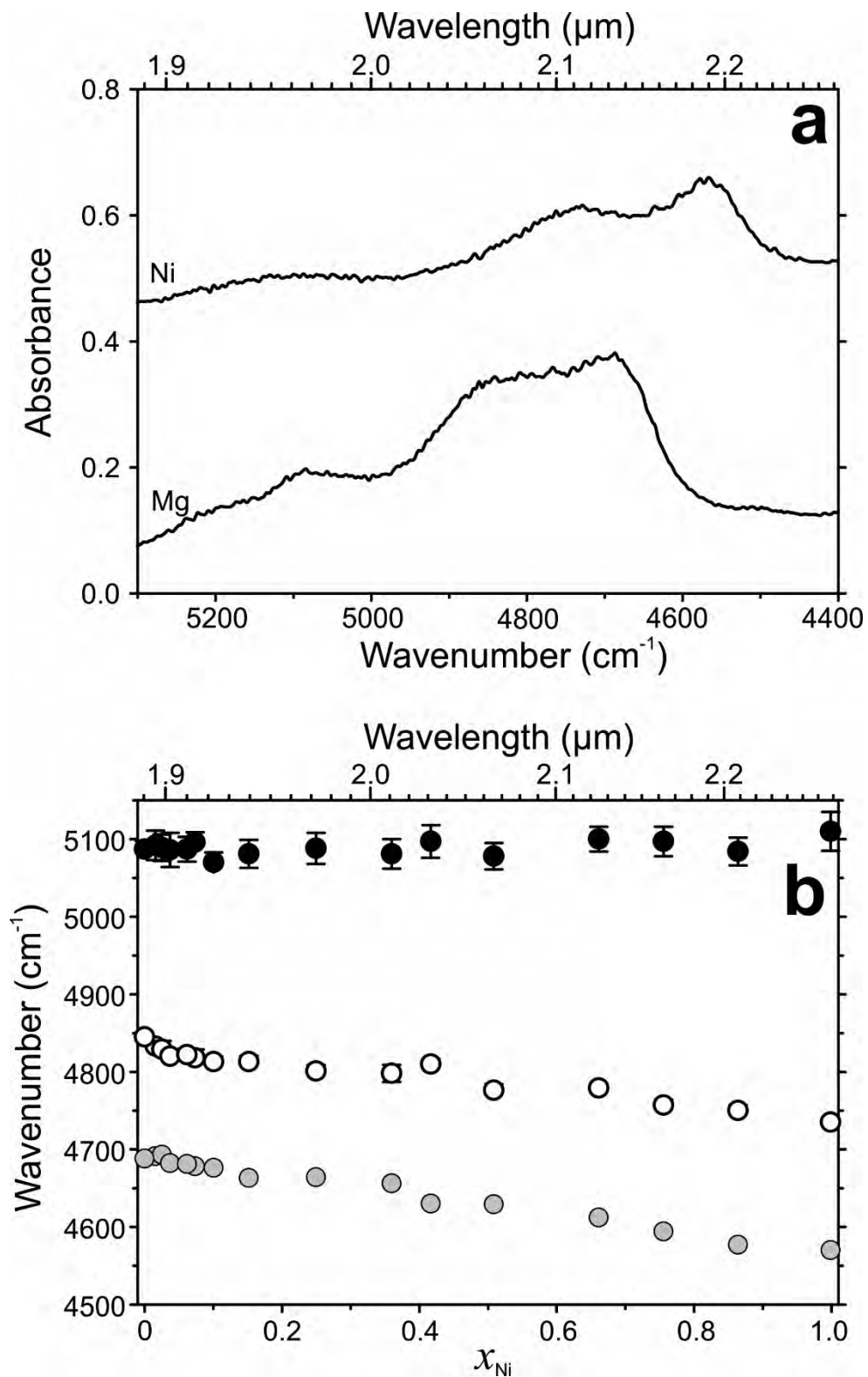
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1247 **Figure 7.** FTIR spectra of kieserite (Mg) and dwornikite (Ni) in each measuring mode. In
1248 addition to band position changes between the endmembers, note the differences in band
1249 positions and shapes between the various measuring modes. Data for kieserite are taken from
1250 Talla and Wildner (2019).



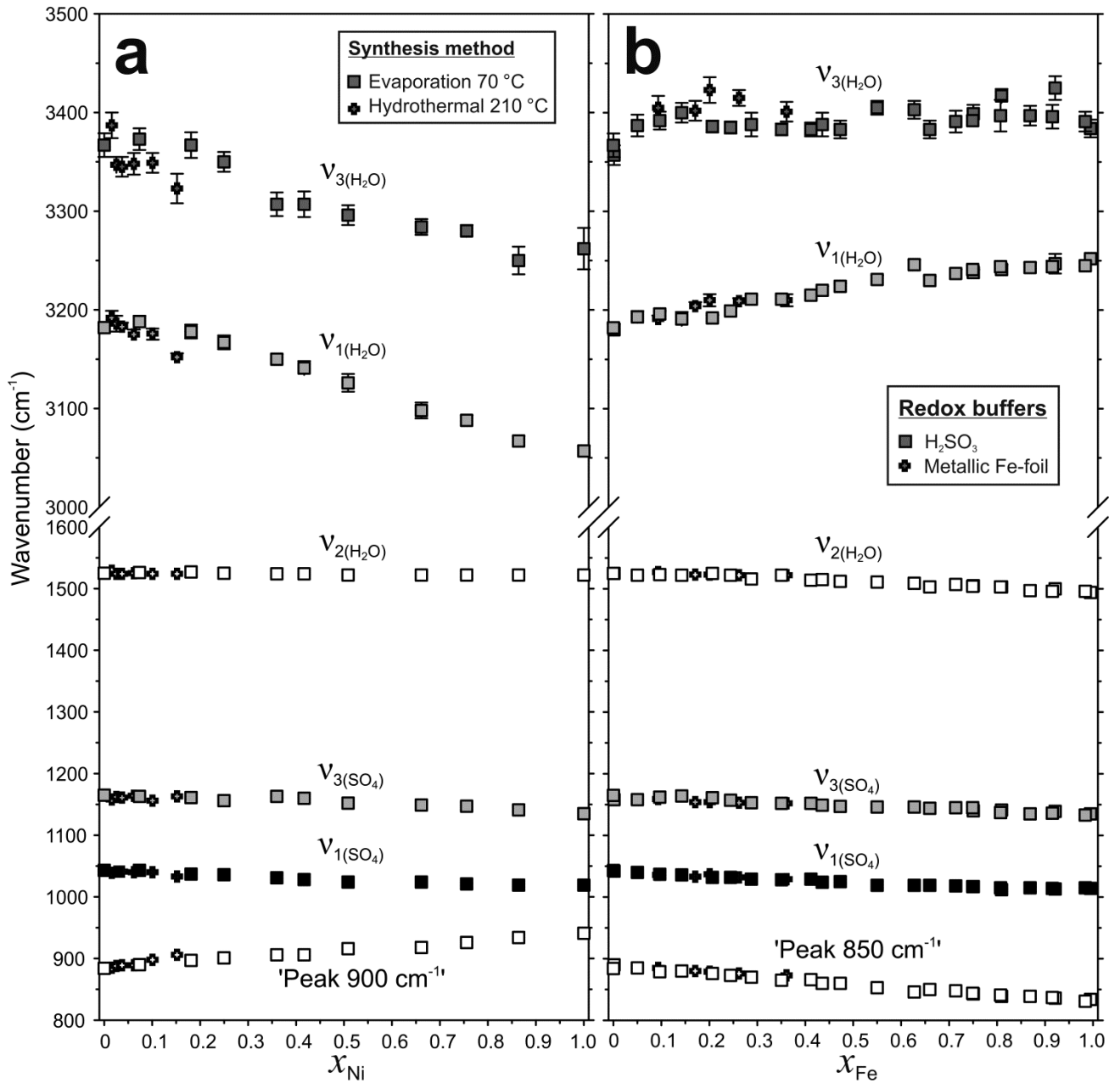
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1252 **Figure 8.** Spectral region with three diagnostic H₂O combination modes, (a) observed in
1253 DRIFT measurements of kieserite (from Talla and Wildner 2019) and dwornikite on undiluted
1254 sample material, (b) detailed plot of band position change for the examined bands across the

1255 kieserite-dwornikite solid solution series. Errors equal or smaller than the symbol size are not
1256 plotted.



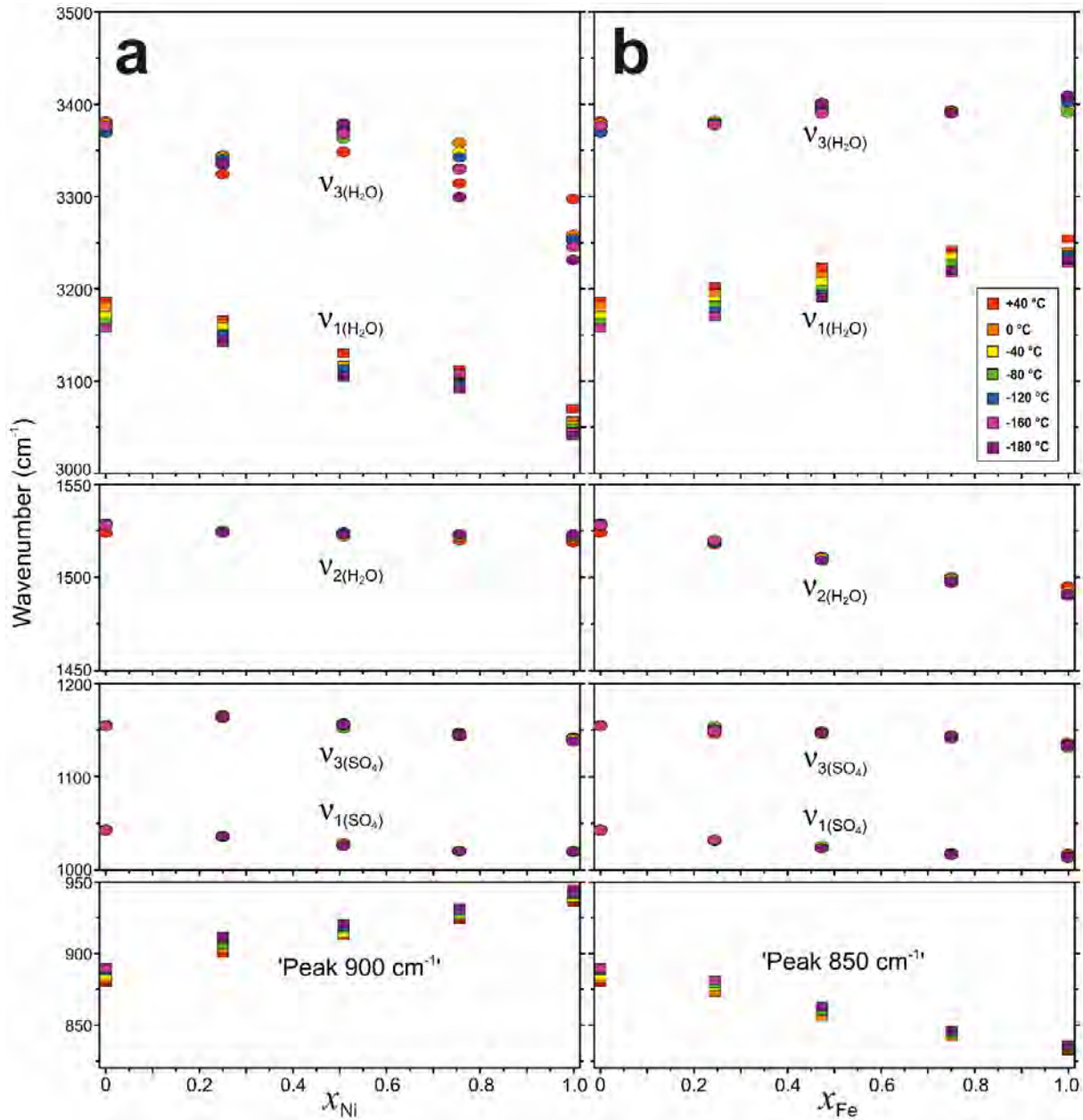
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1258 **Figure 9.** Dependence of the positions of relevant FTIR absorption features (transmission
1259 measurements) of the $\text{Mg}_{1-x}\text{Me}_x(\text{SO}_4)\cdot\text{H}_2\text{O}$ solid solutions (a) on the Ni-content and (b) on the

1260 Fe-content at room temperature (RT). Data for the Mg-Fe solid solution are taken from Talla
1261 and Wildner (2019). Note the opposite trends in the H₂O-related modes. Errors equal or
1262 smaller than the symbol size are not plotted.
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1268 **Figure 10.** Dependence of the FTIR absorption band positions of (a) the $\text{Mg}_{1-x}\text{Ni}_x(\text{SO}_4)\cdot\text{H}_2\text{O}$
1269 solid solution and (b) the $\text{Mg}_{1-x}\text{Fe}_x(\text{SO}_4)\cdot\text{H}_2\text{O}$ (b) on temperature, as seen in transmission
1270 mode. Data for the Mg-Fe solid solution are taken from Talla and Wildner (2019).



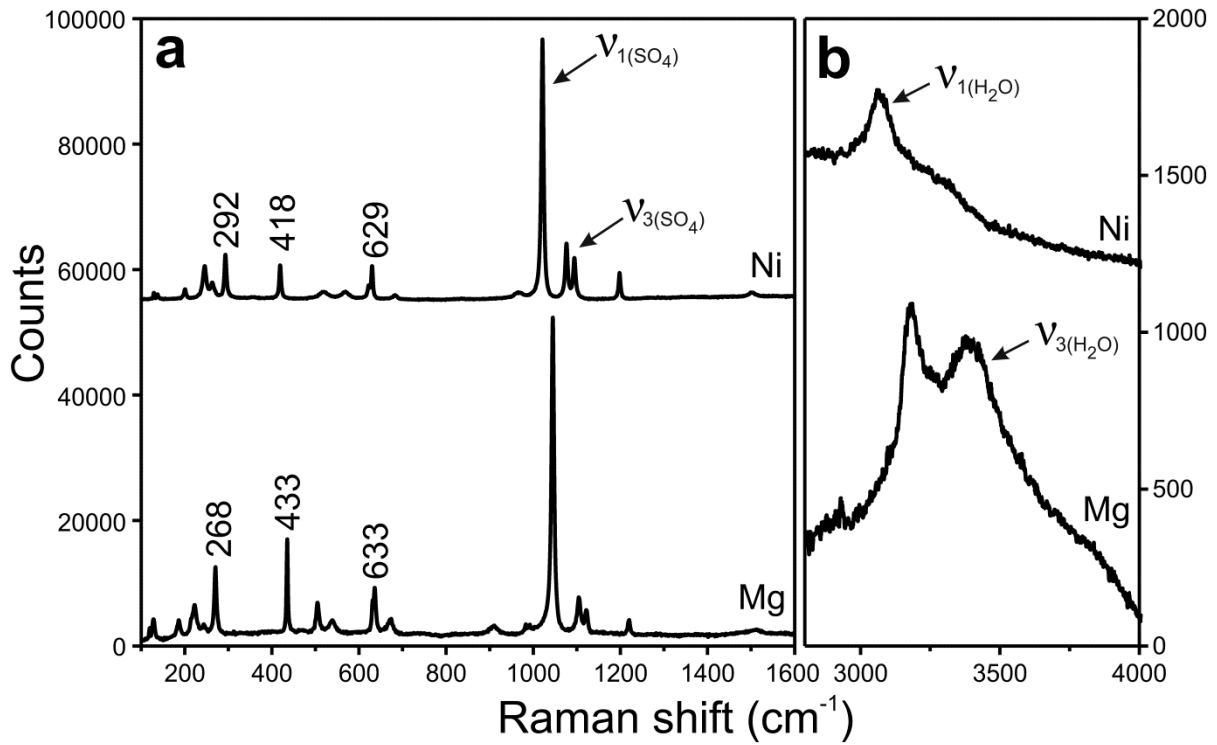
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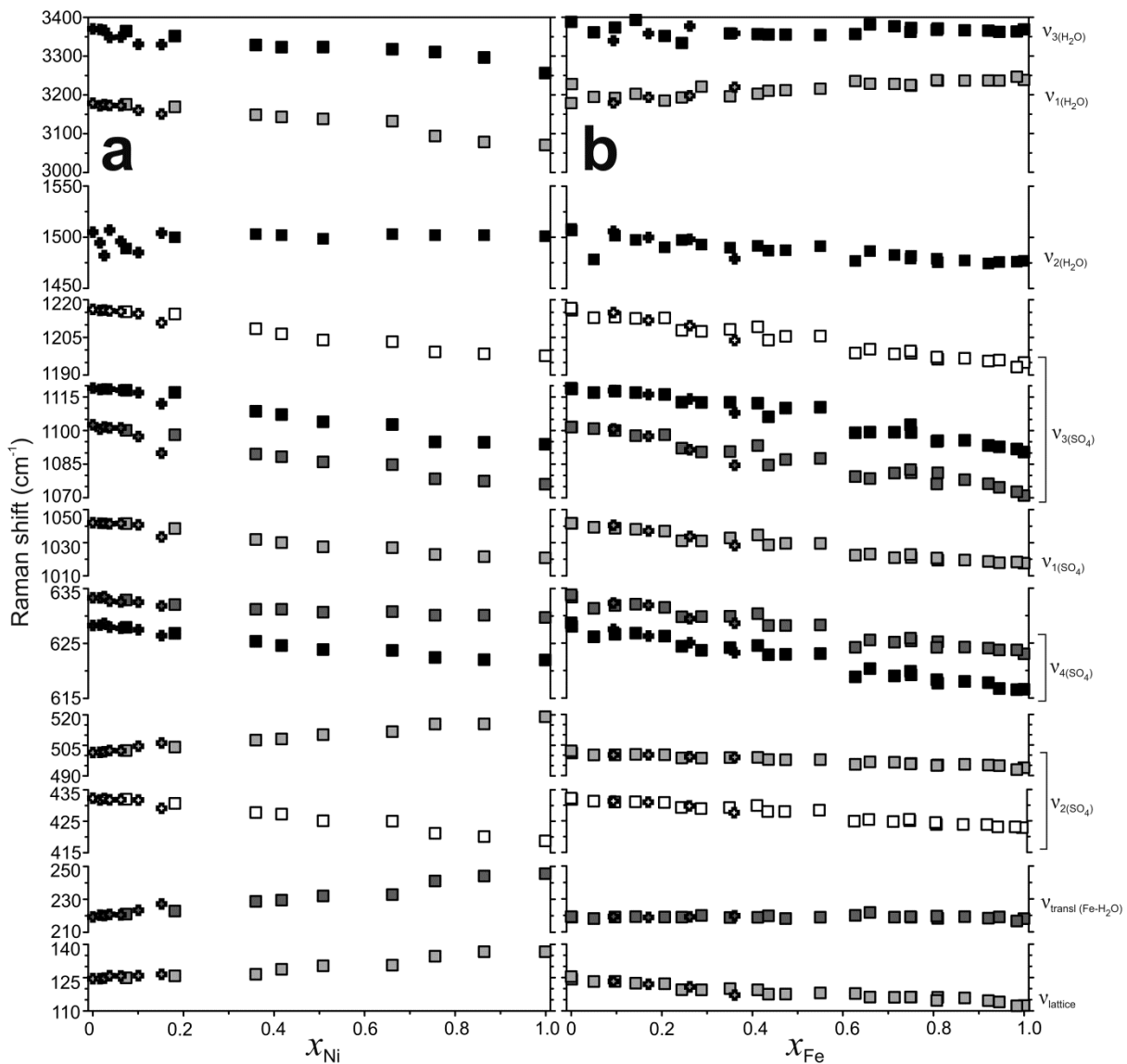
1275 **Figure 11.** Raman spectra of dwornikite (Ni) and kieserite (Mg) at room temperature in the
1276 region of (a) tetrahedral, octahedral and lattice modes, and (b) in the region of the O–H
1277 stretching modes. Data for kieserite are taken from Talla and Wildner (2019).



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1292 **Figure 12.** Dependence of the Raman band positions in the $\text{Mg}_{1-x}\text{Me}_x(\text{SO}_4) \cdot \text{H}_2\text{O}$ solid
1293 solution series on (a) the Ni content and on (b) the Fe content at room temperature. Errors are
1294 equal or smaller than the symbol size. Data for the Mg-Fe solid solution are taken from Talla
1295 and Wildner (2019). Cross symbols in (a) differentiate between samples obtained
1296 hydrothermally with a fully sealed autoclave from those prepared by the evaporation
1297 technique (squares), while crosses in (b) discern the use of a metallic Fe redox buffer from
1298 H_2SO_3 .



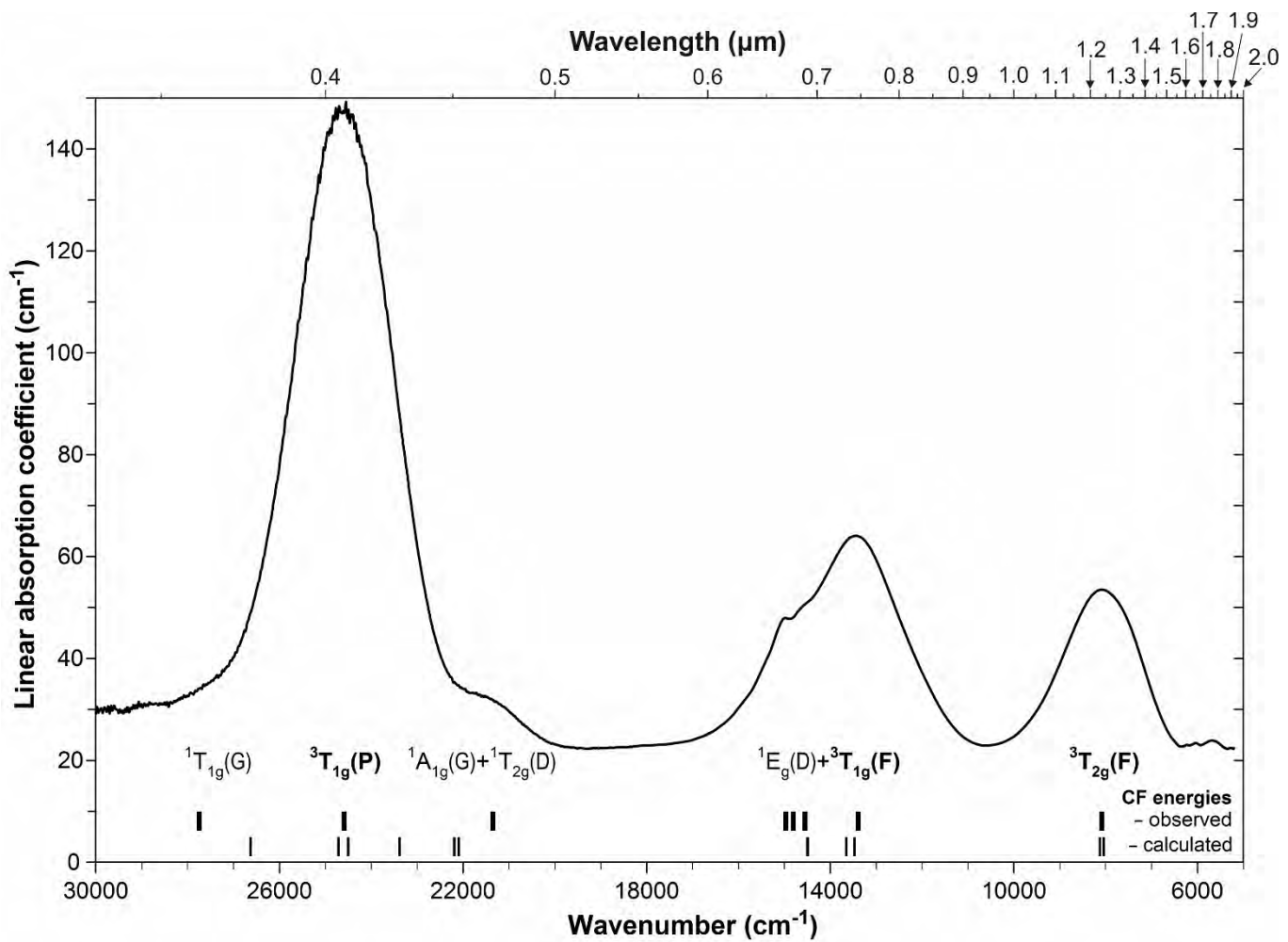
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Figure 13. UV-Vis-NIR absorption spectrum of dwornikite in the range from 30000–5000 cm^{-1} with observed (bold line marks) and calculated energy levels (thin line marks) and respective assignments for cubic symmetry.



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