

1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25

REVISION 3

^{222}Rn and ^{220}Rn emanations from powdered samples of samarskite as a function of annealing temperature

Dariusz Malczewski

Faculty of Earth Sciences, University of Silesia, Bedzinska 60, 41-200 Sosnowiec, Poland

e-mail: dariusz.malczewski@us.edu.pl

Maria Dziurawicz

Faculty of Earth Sciences, University of Silesia, Bedzinska 60, 41-200 Sosnowiec, Poland

e-mail: maria.dziurawicz@us.edu.pl

ABSTRACT

Emanation coefficients for radon (^{222}Rn) and thoron (^{220}Rn) were measured from fully metamict samarskite collected from Centennial Cone after 1 h and 24 h annealing in argon from 473 to 1373 K. For the 1 h annealing run, ^{222}Rn emanation coefficients ranged from 5×10^{-6} to 2.1×10^{-5} %, while ^{220}Rn coefficients varied from 6.3×10^{-3} to 2×10^{-2} %. For the 24 h annealing run, ^{222}Rn coefficients ranged from 5.8×10^{-6} to 2.3×10^{-5} %, while ^{220}Rn coefficients varied from 4.1×10^{-3} to 1.5×10^{-2} %. The ^{222}Rn and ^{220}Rn emanation coefficients versus annealing temperature data can be described by an exponentially decreasing sinusoidal function. Both ^{222}Rn and ^{220}Rn emanation coefficient values after annealing considerably exceeded those measured from an unheated powder reference sample and from the original samarskite sample.

Keywords: samarskite, radon emanations, thoron emanations, recrystallization, Centennial Cone, ^{222}Rn , ^{220}Rn

26

INTRODUCTION

27 Samarskite is a complex Nb-Ta-Ti-REE + Y-Ca-U-Th multiple oxide containing
28 uranium, thorium, iron and other elements and has always been found to be
29 completely metamict (Sugitani et al. 1985). Due to its chemical complexity and
30 metamictization, samarskite's chemical formula and crystal structure have not been
31 unambiguously characterized. The proposed structural formulae are: AB_2O_6 , $A_3B_5O_{16}$
32 and ABO_4 where A = REE, U, Th, Ca, Fe and Ti, and B = Nb, Ta and Ti (Komkov
33 1965; Graham and Thornber 1974; Ewing 1975; Lumpkin et al. 1988). The recently
34 suggested ABO_4 formula is based on microprobe analysis of 19 samarskite samples
35 after annealing at 800° C under hydrogen and on analysis of samarskite-(Yb) from
36 the Little Patsy pegmatite annealed under a weakly reducing atmosphere at
37 temperatures up to 1100° C (Warner and Ewing 1993; Simmons et al. 2006).

38 Radon isotopes ^{222}Rn ($T_{1/2} = 3.82$ d) and ^{220}Rn ($T_{1/2} = 55.6$ s) belong to the
39 ^{238}U and ^{232}Th decay series and occur as inert gases. As part of the uranium series,
40 ^{226}Ra decays by α emission ($E_\alpha = 4.77$ MeV) to form a ^{222}Rn nucleus with an energy
41 of 86 keV. Similarly, ^{224}Ra decays as part of thorium series by α emission ($E_\alpha = 5.67$
42 MeV) with a recoil energy of 103 keV for the daughter ^{220}Rn nucleus. Estimated direct
43 recoil lengths for ^{222}Rn and ^{220}Rn within solids typically range from 20-50 nm (Sakoda
44 et al. 2010; Ishimori et al. 2013). For example, the calculated recoil ranges for ^{222}Rn
45 and ^{220}Rn in quartz and zircon are 34 and 38 nm, and 23 and 26 nm, respectively
46 (Sakoda and Ishimori, 2017). These relatively short ranges mean that without internal
47 defects only radon atoms formed near the mineral surface can be detected as
48 emanations (Krupp et al. 2017 and references therein). Emanation coefficients help
49 characterize retention of radon isotopes within mineral matrices. These ratios
50 (reported as percentages) estimate the number of radon or thoron atoms released

51 from the mineral relative to the number of radon or thoron atoms produced by the
52 decay series occurring within the mineral (Semkow 1990; Morawska and Phillips
53 1993). Emanation coefficients for metamict minerals can be correlated with uranium
54 and thorium concentrations as well as with their spatial distributions, absorbed α -
55 doses, grain size and nuclear track annealing rates. Emanation coefficients may
56 reflect the extent of structural void space and cracks created by radiation damage
57 from progressive overlap of recoil nuclei cascades of ^{238}U , ^{232}Th and ^{235}U and their
58 daughter products.

59 Few studies have addressed radon emanations from metamict minerals and
60 only one study has considered radon and thoron emanations from a large sample of
61 samarskite (Malczewski and Dziurawicz 2015). This study analyzes a fragment of the
62 same massive, dark brown, fully metamict specimen of samarskite (SCC; Fig. 1)
63 collected from a granitic pegmatite in Centennial Cone, Jefferson County, Colorado
64 (USA). Table 1 lists basic characteristics of the SCC specimen. According to
65 nomenclature proposed for samarskite-group minerals, the SCC sample described
66 here categorizes as samarskite-(Y) (Hanson et al. 1999; Simmons et al. 2006). As
67 seen in Table 1, the uranium and thorium concentrations correspond to a calculated
68 total absorbed α -dose, D_T , of 6.5×10^{17} α -decay mg^{-1} . The α -decays from the ^{238}U
69 and ^{235}U series comprise the dominant contribution to the total α -dose of samarskite
70 from Centennial Cone. The ratio of α -doses from $D_{238} + D_{235}$ to D_{232} is about 27.

71 The aim of the study is to determine the relationship between ^{222}Rn and ^{220}Rn
72 emanations and annealing temperature for powdered samples of samarskite.
73 Additionally, this work aims to show that ^{222}Rn and ^{220}Rn emanations can be
74 correlated with the thermally induced transition from the low- to high-temperature
75 phase of samarskite. Results obtained are compared with emanation values from a

76 fragment of the original sample, the unannealed powdered reference sample, and
77 with values from powdered monazite, thorite, uraninite and zircon samples crushed to
78 comparable small grain size as reported from literature sources (Table 2).

79 MATERIALS AND METHODS

80 Concentrations of ^{238}U and ^{232}Th were determined from the intact SCC sample
81 using ^{214}Pb and ^{214}Bi (^{238}U), and ^{228}Ac (^{232}Th) gamma-ray activities. Activity
82 concentrations of ^{235}U assumed a natural abundance of $^{238}\text{U}/^{235}\text{U} = 137.88$. Gamma-
83 ray spectra were recorded using a GX3020 system consisting of a coaxial HPGe
84 detector (32% efficiency) in a lead and copper shield (60 mm) with a multichannel
85 buffer (InSpector 2000 DSP). The detector bias voltage was 4000 V and the energy
86 resolution was 0.8 keV at 122 keV and 1.7 keV at 1.33 MeV. Two software packages
87 were used for the efficiency calibration and the determination of radionuclides:
88 LabSOCS (Laboratory Sourceless Calibration Software) and Genie 2000 v.3.4.

89 After breaking the sample SCC into fragments, 20 pieces of about 2 g were
90 placed in quartz tubes, sealed under argon, and annealed for 1 h and 24 h in a muffle
91 furnace. The temperature program ran from 473 K to 1373 K in increments of 100 K.
92 The furnace stabilized each temperature step within $\pm 2^\circ$. After annealing, the
93 samples were quenched and mechanically ground to an average grain size fraction
94 of 5 μm using an agate ball mill. The original material, crushed to 5 μm grain size but
95 not annealed, is herein designated SRE. The grain sizes were determined for all
96 samples using scanning electron microscopy. Six months after annealing, the
97 powdered samples were placed in copper discs with an outer diameter of 5 cm, an
98 inner diameter of 2 cm, and a 0.15 cm deep groove. The discs were inserted into a
99 stainless steel cylinder ($\phi = 8$ cm, $h = 3$ cm) with two inlets on opposite sides. After
100 insertion, the lid was firmly tightened and inlets were connected to a desiccant and a

101 RAD7 inlet (Fig. 2). The RAD7 radon system (DurrIDGE Company, Inc.) was used to
102 measure ^{222}Rn and ^{220}Rn emanations. Detailed description of RAD7 electronics and
103 measurement configurations are provided by DurrIDGE Company, Inc. (2000) and in
104 Malczewski and DziurOWICZ (2015). The detector operates with a sensitivity of 4 Bq
105 m^{-3} , an upper linear detection limit of 800 kBq m^{-3} , and a manufacturer's calibration
106 accuracy of $\pm 5\%$. The drying unit remained open to the ambient air (open loop
107 mode). Measurement for a given powdered sample occurred over a 15 min cycle
108 repeated 10 times for a total run time of 150 min. Each measurement was conducted
109 independently using three RAD7 detectors operating within a temperature range of
110 20 - 23 $^{\circ}$ C and 4 - 8% internal humidity. The final results represent the average of
111 these three runs. Similar to previous reports (Malczewski and DziurOWICZ 2015;
112 Malczewski et al. 2018), the total emission rates for ^{222}Rn (E_{222}) and ^{220}Rn (E_{220}) from
113 the samples were calculated in atoms s^{-1} according to the following equations:

$$114 \quad E_{222} = \frac{C_{222} \cdot v}{6 \cdot 10^4 \cdot \lambda_{222}} \quad (1)$$

115 and

$$116 \quad E_{220} = \frac{1.28 \cdot C_{220} \cdot v}{6 \cdot 10^4 \cdot \lambda_{220}} \quad (2)$$

117 where C_{222} and C_{220} are respective ^{222}Rn and ^{220}Rn concentrations minus the
118 ambient concentrations (Bq m^{-3}) and v is the flow rate of 1 L min^{-1} . The terms λ_{222}
119 and λ_{220} are respective decay constants for ^{222}Rn and ^{220}Rn of 2.1×10^{-6} and 0.012
120 s^{-1} (Firestone 1996). Our experimental setup (Fig. 2) included a 20 s delay between
121 the emission and measurement by the RAD7 unit. The C_{220} term was therefore
122 multiplied by 1.28.

123 The ^{222}Rn and ^{220}Rn emanation coefficients (e_{222} and e_{220} , respectively) were
124 calculated as the ratio of ^{222}Rn and ^{220}Rn atoms emitted from a samarskite sample (s^{-1})
125 and the total amount of ^{222}Rn and ^{220}Rn produced within the sample. Coefficient
126 equations were as follows:

$$127 \quad e_{222} = \frac{E_{222}}{N_{222}} \quad (3)$$

128 and

$$129 \quad e_{220} = \frac{E_{220}}{N_{220}} \quad (4)$$

130 where N_{222} and N_{220} represent respective estimates of ^{222}Rn and ^{220}Rn nuclei in the
131 sample. Since γ -emitters in both uranium series ($^{234\text{m}}\text{Pa} \rightarrow ^{226}\text{Ra} \rightarrow ^{214}\text{Pb} \rightarrow ^{214}\text{Bi}$)
132 and thorium series ($^{228}\text{Ac} \rightarrow ^{224}\text{Ra} \rightarrow ^{212}\text{Pb} \rightarrow ^{212}\text{Bi} \rightarrow ^{208}\text{Tl}$) were in radioactive
133 equilibrium within the sample analyzed, we assumed that ^{222}Rn and ^{220}Rn values
134 equaled corresponding ^{238}U and ^{232}Th activity concentrations.

135 Powdered and annealed samarskite samples were also analyzed for their X-
136 ray diffraction (XRD) patterns using a Philips X'Pert diffractometer measuring $\text{CuK}\alpha$
137 radiation from the Θ - Θ system in scan mode with a 0.02° step size. Figures 3 and 4
138 show XRD patterns. These figures indicate recrystallization of the fully metamict
139 samarskite SCC begins at 673 K and 573 K for samples subjected to 1 h and 24 h
140 annealing, respectively. The positions of the main diffraction peaks are the same as
141 those reported by Sugitani et al. (1984) for a Kawabe samarskite (Japan) obtained
142 after annealing at 550, 650 and 950°C in a reducing H_2 atmosphere. Sugitani et al.
143 (1984) suggested that the original samarskite phase is formed only under reducing

144 conditions, whereas our results suggest that heating in an inert argon atmosphere
145 also restores metamict samarskite to its original crystalline state.

146 RESULTS

147 Table 3 lists the total emission rates and calculated ^{222}Rn and ^{220}Rn
148 emanation coefficients for the unannealed reference sample and samples of
149 samarskite after annealing in argon for 1h ($e_{1\text{h}222}$ and $e_{1\text{h}220}$). Table 4 lists the total
150 emission rates and ^{222}Rn and ^{220}Rn emanation coefficients ($e_{24\text{h}222}$ and $e_{24\text{h}220}$)
151 calculated for samples after 24 h annealing in argon.

152 *^{222}Rn emanation coefficients versus temperature after 1 h annealing*

153 As shown in Table 3 and Fig. 5, the unheated reference sample (SRE) gave
154 the lowest ^{222}Rn emanation value of $4.9 \times 10^{-6} \%$. For annealed samples, emanation
155 coefficients varied from $5 \times 10^{-6} \%$ to $2.1 \times 10^{-5} \%$ with an average (arithmetic mean)
156 value of $1.51 \times 10^{-5} \%$. Among annealed samples, those annealed at 973 K (S7) and
157 1073 K (S8) gave the lowest $e_{1\text{h}222}$ values of $5 \times 10^{-6} \%$ and $9.6 \times 10^{-6} \%$, whereas
158 those annealed at 673 K (S4) and 573 K (S3) gave the highest values of $2.1 \times 10^{-5} \%$
159 and $2 \times 10^{-5} \%$. Figure 5 shows $e_{1\text{h}222}$ initially increasing from 473 K to 673 K. At
160 annealing temperatures of 773-873 K, $e_{1\text{h}222}$ significantly decreases to a minimum
161 value observed at 973 K (S7). After this point, $e_{1\text{h}222}$ values gradually increase up to
162 1373 K (S11).

163 *^{220}Rn emanation coefficients versus temperature after 1h annealing*

164 Similar to the ^{222}Rn emanation data, the unheated reference sample (SRE)
165 gave the lowest observed ^{220}Rn emanation value of $3.8 \times 10^{-3} \%$ (Tab. 3 and Fig. 6).
166 For annealed samples, ^{220}Rn emanation coefficients ($e_{1\text{h}220}$) ranged from $6.3 \times 10^{-3} \%$
167 to $\sim 2 \times 10^{-2} \%$ with an average value of $1.27 \times 10^{-2} \%$. Samples annealed at 973 K

168 (S7) and 1173 K (S9) gave the lowest e_{1h220} values of $6.3 \times 10^{-3} \%$ and $8.4 \times 10^{-3} \%$,
169 whereas those annealed at 1373 K (S11) and 773 K (S5) gave the highest e_{1h220}
170 values of $2 \times 10^{-2} \%$ and $1.8 \times 10^{-2} \%$, respectively. As seen in Fig. 6, ^{220}Rn
171 emanation coefficients increase from 473 K to 773 K and then, similar to ^{222}Rn
172 emanation coefficients, rapidly decrease to a minimum at 973 K (S7). At higher
173 temperatures, ^{220}Rn emanation coefficients increase up to a maximum value
174 observed at 1373 K (S11). Variation in ^{220}Rn emanation coefficients with temperature
175 resemble those observed for ^{222}Rn emanation coefficients except for sample S11,
176 which gave the highest e_{1h220} value observed.

177 ***^{222}Rn emanation coefficients versus temperature after 24 h annealing***

178 As shown in Table 4 and Fig. 7, and similar to the ^{222}Rn emanation coefficients
179 after 1h annealing, all samarskite samples annealed for 24 h gave higher e_{24h222}
180 values than those of the unannealed reference sample ($e_{\text{SRE}222}$; Tab. 3). The ^{222}Rn
181 emanation coefficients for samples annealed for 24 h ranged from $5.8 \times 10^{-6} \%$ to 2.3
182 $\times 10^{-5} \%$ with an average value of $1.34 \times 10^{-5} \%$. Samples annealed at 473 K (S12)
183 and 1373 K (S21) gave the lowest e_{24h222} values of $5.8 \times 10^{-6} \%$ and $9 \times 10^{-6} \%$,
184 whereas those annealed at at 573 K (S13), 673 K (S14) and 773 K (S15) gave the
185 highest values of $2.3 \times 10^{-5} \%$ and $2 \times 10^{-5} \%$ (respectively). Figure 7 shows that the
186 e_{24h222} emanation coefficient increases fourfold from sample S12 (473 K) to sample
187 S13 (573 K) and noticeably decreases to a local minimum for sample S17 (973 K). In
188 contrast to the 1h annealing data, e_{24h222} values for samples annealed at the highest
189 temperatures decrease with increasing temperature.

190 ***^{220}Rn emanation coefficients versus temperature after 24 h annealing***

191 Similar to the ^{220}Rn emanation coefficients after 1 h annealing, the unheated
192 reference sample (SRE) also gave the lowest ^{220}Rn emanation value after 24 h
193 (Tables 3 and 4 and Fig. 8). For samarskite samples annealed for 24 h, ^{220}Rn
194 emanation coefficients ($e_{24\text{h}220}$) ranged from $4.1 \times 10^{-3} \%$ to $1.5 \times 10^{-2} \%$ with an
195 average value of $9.87 \times 10^{-3} \%$. Samples annealed at 473 K (S12) and 1073 K (S18)
196 gave the lowest $e_{24\text{h}220}$ values of $4.1 \times 10^{-3} \%$ and $5.7 \times 10^{-3} \%$, whereas samples
197 annealed at 773 K (S15) and 673 K (S14) gave the highest values of $\sim 1.5 \times 10^{-2} \%$
198 (for both). The ^{220}Rn emanation coefficients increase from 473 K to 773 K and then
199 significantly decrease to a second minimum at 1073 K (S18). The $e_{24\text{h}220}$ values
200 increase for samples annealed at 1173 K (S19) and 1273 K (S20), but the $e_{24\text{h}220}$
201 value for sample S21 annealed at 1373 K decrease to $7.4 \times 10^{-3} \%$. Variation in ^{220}Rn
202 emanation coefficients ($e_{24\text{h}220}$) with temperature resembles those observed for ^{222}Rn
203 emanation coefficients ($e_{24\text{h}222}$; Fig. 7) except in the case of sample S18 (1073 K),
204 which exhibited a distinct minimum (Fig. 8).

205 DISCUSSION

206 Samples of samarskite annealed for both 1 h and 24 h gave maximum ^{222}Rn
207 emanation coefficients ($e_{1\text{h}222}$ and $e_{24\text{h}222}$) within a 573 K to 773 K temperature range.
208 Both datasets also showed a clear minimum at 973 K (Figs. 5 and 7). Samples
209 annealed for 1 h above 973 K gave $e_{1\text{h}222}$ values that increase up to 1373 K whereas
210 those annealed for 24 h gave $e_{24\text{h}222}$ values that decrease at 1273 K and 1373 K.

211 Using the analogy of a damped sinusoidal vibration, the variation in all $e_{1\text{h}222}$,
212 $e_{24\text{h}222}$, $e_{1\text{h}220}$, and $e_{24\text{h}220}$ emanation coefficients with temperature, after annealing for
213 1 h and 24 h in argon, can be relatively well fitted using an exponentially decreasing
214 sinusoidal function of the form:

(5)

$$215 \quad e_i(\%) = e_{0i} + A_i \cdot \exp(-b_i \cdot T) \cdot \sin(c_i \cdot (T - T_{0i}))$$

216 where e_{0i} is the weighted average of e_i , A_i is the initial amplitude of the envelope, b_i is
217 the damping factor of emanation, and T_{0i} is the temperature at which e_i equals e_{0i} .
218 The term $c_i = \pi / \Delta T_i$, where ΔT_i is the periodicity of e_{0i} and T_{0i} . The index i ($i = 1, 2, 3,$
219 and 4) refers to e_{1h222} , e_{24h222} , e_{1h220} , and e_{24h220} , respectively. The parameters are
220 explained in Fig. 9, the values of the fitted parameters are listed in Table 5, and the
221 fitted curves are shown in Fig. 10.

222 The function graphs exhibit two temperature ranges, from 473 K to about 1000
223 K and from 1000 K to 1373 K, which generally coincide with the recrystallization path
224 proposed by Sugitani et al. (1984). Their research suggested that the low-
225 temperature orthorhombic samarskite phase formed after 16 h annealing in H_2 at 550°
226 C (823 K) and then on heating up to 950° C (1273 K), at which point a high-
227 temperature monoclinic phase is formed. A splitting of the most intense diffraction
228 peak ($2\theta \approx 30^\circ$) into a doublet accompanies the transition from the low- to high-
229 temperature phase. For the Kawabe samarkite, it occurred after heating at 650° C
230 (923 K), whereas in our study the same main diffraction peak ($2\theta \approx 30^\circ$) began
231 splitting at 973 K (Figs. 3 and 4). Warner and Ewing (1993) annealed samarskite
232 samples at 800 °C (1073 K) in an H_2 atmosphere for 4 h. Those annealed samples
233 were not completely crystalline, which agreed well with our results. Tomašić et al.
234 (2010) analysed a fully metamict samarskite from Beinmyr pegmatite (Norway)
235 annealed in air and Ar/H_2 atmosphere. The high-temperature samarskite phase was
236 not observed at temperatures lower than 800°C (1073 K). The TGA-DTA data for
237 samarskite from Beinmyr recorded from RT (298 K) to 1000°C (1273 K) showed a
238 strong endothermic peak occurring at about 300°C (573 K). This result coincides well

239 with the observation of the highest emanations of ^{222}Rn in a narrow temperature
240 range of 573 to 673 K (Figs. 5 and 7).

241 As shown in Table 5, the calculated b_i values after 24 h annealing, both for
242 $e_{24\text{hh}222}$ and $e_{24\text{h}220}$, are about one to two orders of magnitude higher than those for
243 $e_{1\text{h}222}$ and $e_{1\text{h}220}$ after 1 h annealing. Combined with X-ray patterns, this means that
244 the high-temperature and long-lasting (24 h in this case) annealing lead to the
245 formation of a stable, fully crystalline polymorph of samarskite. As a result, small
246 fluctuations for both $e_{24\text{hh}222}$ and $e_{24\text{h}220}$ emanation coefficients about their average
247 values are observed (Figs. 10 b and d). After 1 h of annealing, these fluctuations are
248 noticeably higher (Figs. 10 a and c).

249 As seen in Fig. 11, ratios of $e_{1\text{h}222}$ to the ^{222}Rn emanation coefficient of the
250 unheated sample ($e_{\text{SRE}222}$) ranged from 1 to 4 with an average 3.1. Ratios of $e_{1\text{h}222}$ to
251 the ^{222}Rn emanation coefficient reported for original SCC sample ($e_{\text{SCC}222}$; 1.28×10^{-6})
252 varied from 8 to 16 with an average value of 11.8. Ratios corresponding to ^{222}Rn
253 emanation coefficients after 24 h annealing gave similar values. As shown in Fig. 12,
254 the ratio of $e_{24\text{h}222}$ to the SRE reference sample emanation coefficient ranged from
255 about 1 to 5 with an average value of 2.7. Ratios of $e_{1\text{h}222}$ to the SCC ^{222}Rn
256 emanation coefficient varied from 5 to 18 with an average value of 10.4. Average
257 ^{222}Rn emanation coefficients for samarskite samples ground to a $\sim 5 \mu\text{m}$ grain size
258 fraction were about two orders of magnitude lower than those reported by Garver and
259 Baskaran (2004) for samples of monazite, zircon, thorite and uraninite, which were
260 crushed to the grain sizes of less than $63 \mu\text{m}$ and annealed at 873 K for 6 h (Tab. 2).
261 These samples gave ratios of the coefficients from ^{222}Rn emanations after annealing
262 at 873 K to the coefficients from unannealed samples from 0.25 to 0.57, with an
263 average value of 0.4. This value differs from the respective $e_{1\text{h}222}/e_{\text{SRE}222}$ and

264 e_{24h222}/e_{SRE222} values of 3.1 and 2.7 reported here. Our results show that annealed
265 samples of samarskite SCC always result in higher values of ^{222}Rn emanation than
266 the unannealed reference sample. A similar effect should be observed for other
267 metamict phases.

268 The 1.27×10^{-2} average value for ^{220}Rn emanation coefficients after 1 h
269 annealing (e_{1h220}) slightly exceeded the 9.87×10^{-3} average value calculated for
270 samarskite samples after 24 h annealing (e_{24h220}). Generally, ^{220}Rn emanation
271 coefficients from both 1 h and 24 h annealing exceed ^{222}Rn emanations (e_{1h222} and
272 e_{24h222}) by about three orders of magnitude. Figure 13 shows that ratios of e_{1h220} to
273 the ^{220}Rn emanation coefficient for the unannealed reference sample (e_{SRE220}) range
274 from about 2 to 5 with an average value of 3.3. Ratios of e_{1h220} to the original SCC
275 sample ^{220}Rn emanation coefficient ($e_{SCC220} = 3.8 \times 10^{-4} \%$) vary from 17 to 51 with an
276 average value of 33. After 24 h annealing (Fig. 14), ratios of e_{24h220} to the emanation
277 coefficient of the SRE reference sample range from about 1 to 4 with an average
278 value of 2.6. Ratios of e_{24h220} to SCC sample ^{220}Rn emanation coefficients vary from
279 11 to 40 with an average value of 26. From Figs. 11-14, it appears that the
280 emanation coefficients for both the ^{222}Rn and ^{220}Rn , annealed at temperatures from
281 473 K to 1373 K, are 3 times higher on average than these observed for the
282 unannealed sample. These figures also show that the emanation coefficients for both
283 ^{222}Rn and ^{220}Rn from samarskite SRE, crushed to 5 μm grain size, exceeded by only
284 one order of magnitude those observed for the intact 4 cm fragment, SCC. A similar
285 effect with reference to granite was reported by Amin and Rama (1986). There was
286 no observed significant difference between radon emanation coefficients from a
287 granite cube with a 30 cm edge and 1-2 mm granite grains.

288

IMPLICATIONS

289 The observed variations in the ^{222}Rn and ^{220}Rn emanation coefficients with
290 temperature for samples of fully metamict samarskite-(Y) ground to a 5 μm grain size
291 and annealed from 473 K to 1373 K coincide well with the structural conversion from
292 a low- to high-temperature samarskite phase reported in previous studies. The same
293 mineral species showed noticeably different emanation coefficients of radon and
294 thoron depending on the crystallographic system induced by annealing in an argon
295 atmosphere. The ^{222}Rn emanation coefficients obtained both for 1 h and 24 h
296 annealing were significantly lower than the values reported in the literature for
297 comparable metamict minerals. The results reported here indicate that samarskite-
298 (Y) behaves as a closed system for radon retention across a very broad temperature
299 range, from an untreated sample to a sample annealed at 1373 K, despite the high
300 concentration of uranium and unusual structural complexity.

301

ACKNOWLEDGEMENTS

302 This work was supported by the National Science Centre, Poland, through grants
303 Nos. 2014/15/B/ST10/04095 and 2018/29/B/ST10/01495. We thank two unknown
304 reviewers for critical comments that improved the final version of this paper.

305

REFERENCES CITED

306 Amin, B.S., Rama, (1986) Using radon as probe for investigating characteristic of
307 fractures in crystalline materials. Nuclear Instruments and Methods in Physics
308 Research B, 17, 527-529.

309 Bryant, B., McGrew, L.W., and Wobus, R.A. (1981) Geologic map of the Denver 1 x 2
310 degree quadrangle, north-central Colorado: U.S. Geological Survey Miscellaneous
311 Investigations Map I-1163, scale 1:250,000.

312 Durrige Company Inc. (2000) RAD7 Radon Detector. Part No. 1280, Bedford, MA.

- 313
314 Ewing, R.C. (1975) The crystal chemistry of complex niobium and tantalum oxides.
315 IV. The metamict state: Discussion. American Mineralogist, 60, 728-733.
- 316 Firestone, R.B. (Ed.) (1996) Table of isotopes. Wiley-Interscience, Lawrence
317 Berkeley National Laboratory.
- 318
- 319 Garver, E., and Baskaran, M. (2004) Effects of heating on the emanation rates of
320 radon-222 from a suite of natural minerals. Applied Radiation and Isotopes, 61, 1477-
321 1485.
- 322 Graham, J., and Thornber, M.R. (1974) The crystal chemistry of complex niobium
323 and tantalum oxides I. Structural classification of MO₂ phases. American
324 Mineralogist, 59, 1026-1039.
- 325 Hanson, S.L., Simmons, W.B., Falster, A.U., Foord, E.E., and Lichte, F.E. (1999)
326 Proposed nomenclature for samarskite-group minerals: new data on ishikawaite and
327 calciosamarskite. Mineralogical Magazine, 61(1), 27-36.
- 328 Ishimori, Y., Lange K., Martin, P., Mayya, Y.S., and Phaneuf (2013) Measurement
329 and calculation of radon releases from NORM residues. International Atomic Energy
330 Agency (IAEA). Technical Reports Series No. 474.
- 331 Komkov, A.I. (1965) Crystal structure and chemical constitution of samarskite.
332 Doklady Akademiia Nauk SSSR, Earth Science Section, 160, 127-129. (English
333 translation).
- 334 Krupp, K., Baskaran, M., and Brownlee, S.J. (2017) Radon emanation coefficients of
335 several minerals: How they vary with physical and mineralogical properties. American
336 Mineralogist, 102, 1375-1383.

- 337 Lumpkin, G.R., Ewing, R.C., and Eyal, Y. (1988) Preferential leaching and natural
338 annealing of alpha-recoil tracks in metamict betafite and samarskite. Journal of
339 Materials Research, 3(2), 357-368.
- 340 Malczewski, D., and Dziurawicz, M. (2015) ^{222}Rn and ^{220}Rn emanations as a
341 function of the absorbed α -doses from select metamict minerals. American
342 Mineralogist, 100, 1378-1385.
- 343 Malczewski, D., and Dziurawicz, M., Krzykowski, T., and Stryjewski, A. (2018) ^{222}Rn
344 and ^{220}Rn emanations from zircon crystals as a function of absorbed α -doses. The
345 Canadian Mineralogist, 56, 451-462.
- 346 Morawska, L., and Phillips, C.R. (1993) Dependence of the radon emanation
347 coefficient on radium distribution and internal structure of the material. Geochimica et
348 Cosmochimica Acta, 57, 1783-1797.
- 349 Sakoda, A., Hanamoto, K., Ishimori, Y., Kataoka, T., Kawabe, A., and Yamaoka, K.
350 (2010) First model of the effect of grain size on radon emanation. Applied Radiation
351 and Isotopes, 68, 1169-1172.
- 352 Sakoda, A., and Ishimori, Y. (2017) Mechanism and Modelling Approaches of Radon
353 Emanation for Natural Materials. Japanese Journal of Health Physics, 52(4), 296-
354 306.
- 355 Semkow, T.M. (1990) Recoil-emanation theory applied to radon release from mineral
356 grains. Geochimica et Cosmochimica Acta, 54, 425-440.
- 357 Simmons, W.B., Hanson, S.L., and Falster, A.U. (2006) Samarskite-(Yb): A new
358 species of the samarskite group from the Little Patsy Pegmatite, Jefferson County,
359 Colorado. The Canadian Mineralogist, 44, 1119-1125.

360 Sugitani, Y., Suzuki, Y., and Nagashima, K. (1984) Recovery of the original
361 samarskite structure in a reducing atmosphere. *American Mineralogist*, 69, 377-379.

362 Sugitani, Y., Suzuki, Y., and Nagashima, K. (1985) Polymorphism of samarskite and
363 its relationship to other structurally related Nb-Ta oxides with the α -PbO₂ structure.
364 *American Mineralogist*, 70, 856-866.

365 Tomašić, N., Gajović, A., Bermanec, V., Linarić, M.R., Su, D., and Škoda, R. (2010)
366 Preservation of the samarskite structure in a metamict ABO₄ mineral: a key to crystal
367 structure identification. *European Journal of Mineralogy*, 22, 435-442.

368 Warner, J.K., and Ewing, R.C. (1993) Crystal chemistry of samarskite. *American*
369 *Mineralogist*, 78, 419-424.

370 **Figure captions**

371 **FIGURE 1.**

372 Photo of fully metamict samarskite (SCC) from Centennial Cone, Colorado, USA. The
373 specimen is 4 cm in length. About half of the specimen was used in this study.

374 **FIGURE 2.**

375 Experimental setup for measuring ²²²Rn and ²²⁰Rn emanations from powdered
376 samarskite samples.

377 **FIGURE 3.**

378 X-ray diffraction (XRD) patterns for the unheated SRE samarkite sample and
379 samples after 1 h annealing under argon from 473 K (S2) to 1373 K (S11).

380 **FIGURE 4.**

381 X-ray diffraction (XRD) patterns for samarskite samples after 24 h annealing under
382 argon from 473 K (S12) to 1373 K (S21).

383 **FIGURE 5.**

384 ^{222}Rn emanation coefficients for unheated SRE samarskite sample ($e_{\text{SRE}222}$) and for
385 samples after 1 h annealing in argon ($e_{1\text{h}222}$) from 473 K to 1373 K.

386

387 **FIGURE 6.**

388 ^{220}Rn emanation coefficients for unheated SRE samarskite sample ($e_{\text{SRE}220}$) and for
389 samples after 1 h annealing in argon ($e_{1\text{h}220}$) from 473 K to 1373 K.

390 **FIGURE 7.**

391 ^{222}Rn emanation coefficients for unheated SRE samarskite sample ($e_{\text{SRE}222}$) and for
392 samples after 24 h annealing in argon ($e_{24\text{h}222}$) from 473 K to 1373 K.

393 **FIGURE 8.**

394 ^{220}Rn emanation coefficients for unheated SRE samarskite sample ($e_{\text{SRE}220}$) and for
395 samples after 24 h annealing in argon ($e_{24\text{h}220}$) from 473 K to 1373 K.

396 **FIGURE 9.**

397 The parameters used in this work to describe variations in the ^{222}Rn and ^{220}Rn
398 emanation coefficients with temperature, based on Eq. (5).

399 **FIGURE 10.**

400 Plots of **(a)** $e_{1\text{h}222}$, **(b)** $e_{24\text{h}222}$, **(c)** $e_{1\text{h}220}$, and **(d)** $e_{24\text{h}220}$ versus temperature. The solid
401 lines show the fit to the experimental data based on Eq. (5). Thin horizontal lines
402 represent the fitted values of e_{0i} .

403 **FIGURE 11.**

404 Plot of ratios of ^{222}Rn emanations from samarskite samples after 1 h annealing
405 (e_{1h222}) to ^{222}Rn emanation from unheated reference sample ($e_{\text{SRE}222}$) (circles) and
406 ratios of e_{1h222} to ^{222}Rn emanation from original sample ($e_{\text{SCC}222}$) (triangles) versus
407 temperature. Thin solid line shows average $e_{1h222} / e_{\text{SRE}222}$ ratio and thick solid line
408 shows average $e_{1h222} / e_{\text{SCC}222}$ ratio.

409

410 **FIGURE 12.**

411 Plot of ratios of ^{222}Rn emanations from samarskite samples after 24 h annealing
412 (e_{24h222}) to ^{222}Rn emanation from unheated reference sample ($e_{\text{SRE}222}$) (circles), and
413 ratios of e_{24h222} to ^{222}Rn emanation from original large sample $e_{\text{SCC}222}$ (triangles)
414 versus temperature. Thin solid line shows average $e_{24h222} / e_{\text{SRE}222}$ ratio and thick
415 solid line shows average $e_{24h222} / e_{\text{SCC}222}$ ratio.

416 **FIGURE 13.**

417 Plot of ratios of ^{220}Rn emanations from samarskite samples after 1 h annealing
418 (e_{1h220}) to ^{220}Rn emanation from unheated reference sample ($e_{\text{SRE}220}$) (squares) and
419 ratios of e_{1h220} to ^{220}Rn emanation from original sample ($e_{\text{SCC}220}$) (inverted triangles)
420 versus temperature. Thin solid line shows average $e_{1h220} / e_{\text{SRE}220}$ ratio and thick solid
421 line shows average $e_{1h220} / e_{\text{SCC}220}$ ratio.

422 **FIGURE 14.**

423 Plot of ratios of ^{220}Rn emanations from samarskite samples after 24 h annealing
424 (e_{24h220}) to ^{220}Rn emanation from unheated reference sample ($e_{\text{SRE}220}$) (squares) and
425 ratios of e_{24h220} to ^{220}Rn emanation from original sample ($e_{\text{SCC}220}$) (inverted triangles)
426 versus temperature. Thin solid line shows average $e_{24h220} / e_{\text{SRE}220}$ ratio and thick
427 solid line shows average $e_{24h220} / e_{\text{SCC}220}$ ratio.

Table 1

Age, basic chemical composition (wt.%), calculated absorbed α -doses and ^{222}Rn and ^{220}Rn emanation coefficients for the original intact samarskite sample (SCC).

Age	1550(150) Ma ^a
O	20.12(81)
Si	2.39(10)
Ca	2.0(1)
Ti	1.78(10)
Fe	2.83(15)
Y	7.17(32)
Nb	29.05(64)
Ta	5.19(54)
Pb	2.69(56)
Th	1.82(2)
U	10.54(40)
Calculated total dose (D_T) ^b (α -decay mg^{-1})	6.5×10^{17}
Calculated dose from ^{232}Th (D_{232}) (α -decay mg^{-1})	2.3×10^{16}
Calculated dose from ^{238}U (D_{238}) (α -decay mg^{-1})	5.77×10^{17}
Calculated dose from ^{235}U (D_{235}) (α -decay mg^{-1})	4.97×10^{16}
$e_{\text{SCC}222}$ ^c	$1.28 \times 10^{-6}\%$
$e_{\text{SCC}220}$	$3.8 \times 10^{-4}\%$

^a Bryant et al (1981). ^b Doses were calculated as: $D_{238} = 8 \times N_{238}(e^{\lambda_{238}t} - 1)$, $D_{235} = 7 \times N_{235}(e^{\lambda_{235}t} - 1)$, $D_{232} = 6 \times N_{232}(e^{\lambda_{232}t} - 1)$ and $D_T = D_{238} + D_{235} + D_{232}$. N_{238} , N_{235} and N_{232} are the present number of atoms of ^{238}U , ^{235}U and ^{232}Th per milligram, λ_{238} , λ_{235} and λ_{232} are the decay constants of ^{238}U , ^{235}U and ^{232}Th (respectively), and t is the geologic age. The absorbed ^{235}U α -doses were calculated assuming a natural atomic abundance of $^{238}\text{U}/^{235}\text{U} = 137.88$. ^c ^{222}Rn emanation coefficient (e_{222}) and ^{220}Rn emanation coefficient (e_{220})

Table 2

Previously reported ^{222}Rn emanation coefficients for metamict minerals crushed to grain size of less than $63\ \mu\text{m}$ and heated for 6 h (Garver and Baskaran 2004).

Sample	Heating temperature (K)	e_{222} (%)	e_{600}/e_{RT}
Monazite	RT	2.05	
	873	0.66	0.32
Zircon	RT	1.04	
	873	0.47	0.45
Uraninite	RT	0.53	
	873	0.30	0.57
Thorite	RT	5.38	
	873	1.34	0.25
Average			0.40

Table 3

Total emission rates and ^{222}Rn and ^{220}Rn emanation coefficients for the samarskite samples after 1 h annealing and for unheated reference sample (SRE).

Annealing temperature (K)	Sample	N_{222} (10^8 atoms)	E_{222} (atom s^{-1})	e_{1h222} (%) ^a	N_{220} (10^3 atoms)	E_{220} (atom s^{-1})	e_{1h220} (%) ^b
RT	SRE	7.26	35.6	4.90×10^{-6}	7.23	0.27	3.76×10^{-3}
473	S2	4.91	56.1	1.14×10^{-5}	4.89	0.46	9.46×10^{-3}
573	S3	2.93	58.9	2.01×10^{-5}	2.92	0.30	1.02×10^{-2}
673	S4	4.95	102	2.07×10^{-5}	4.93	0.83	1.69×10^{-2}
773	S5	4.91	93.3	1.90×10^{-5}	4.89	0.88	1.79×10^{-2}
873	S6	6.04	112	1.86×10^{-5}	6.02	0.82	1.36×10^{-2}
973	S7	6.96	35	5.04×10^{-6}	6.94	0.44	6.34×10^{-3}
1073	S8	5.69	54.9	9.65×10^{-6}	5.67	0.53	9.34×10^{-3}
1173	S9	4.75	65	1.37×10^{-5}	4.74	0.37	8.36×10^{-3}
1273	S10	5.29	83.5	1.58×10^{-5}	5.27	0.79	1.49×10^{-2}
1373	S11	4.98	87.2	1.75×10^{-5}	4.97	0.98	1.97×10^{-2}
Average				1.51×10^{-5}			1.27×10^{-2}

^a Estimated uncertainties $\Delta e_{222}/e_{222} \leq 14\%$. ^b Estimated uncertainties $\Delta e_{220}/e_{220} \leq 10\%$.

Table 4

Total emission rates and ^{222}Rn and ^{220}Rn emanation coefficients for samarskite samples after 24 h annealing.

Annealing temperature (K)	Sample	N_{222} (10^6 atoms)	E_{222} (atom s^{-1})	e_{24h222} (%) ^a	N_{220} (10^3 atoms)	E_{220} (atom s^{-1})	e_{24h220} (%) ^b
473	S12	6.81	39.7	5.82×10^{-6}	6.79	0.28	4.14×10^{-3}
573	S13	3.29	75.3	2.29×10^{-5}	3.28	0.45	1.37×10^{-2}
673	S14	3.49	69.2	1.98×10^{-5}	3.48	0.51	1.46×10^{-2}
773	S15	4.91	96.1	1.96×10^{-5}	4.89	0.74	1.52×10^{-2}
873	S16	5.21	60.1	1.16×10^{-5}	5.19	0.65	1.04×10^{-2}
973	S17	5.69	54.5	9.56×10^{-6}	5.67	0.50	8.83×10^{-3}
1073	S18	7.57	93.2	1.23×10^{-5}	7.54	0.43	5.74×10^{-3}
1173	S19	3.98	49.4	1.24×10^{-5}	3.96	0.37	9.20×10^{-3}
1273	S20	4.70	50.4	1.07×10^{-5}	4.68	0.45	9.51×10^{-3}
1373	S21	6.48	58.6	9.02×10^{-6}	6.46	0.48	7.40×10^{-3}
Average				1.34×10^{-5}			9.87×10^{-3}

^a Estimated uncertainties $\Delta e_{222}/e_{222} \leq 14\%$. ^b Estimated uncertainties $\Delta e_{220}/e_{220} \leq 10\%$.

Table 5

Fitted parameters for the function given by Eq. (5).

i	1	2	3	4
Parameter	e_{1h222}	e_{24h222}	e_{1h220}	e_{24h220}
e _{0i} (%)	1.37 x 10 ⁻⁵	1.17 x 10 ⁻⁵	1.24 x 10 ⁻²	9.1 x 10 ⁻³
A _i	1.48 x 10 ⁻⁵	3.27 x 10 ⁻⁴	6.1 x 10 ⁻³	6.9 x 10 ⁻²
b _i (K ⁻¹)	6.8 x 10 ⁻⁴	5.3 x 10 ⁻³	9.8 x 10 ⁻⁵	3.9 x 10 ⁻³
T _{0i} (K)	506	504	547	520
ΔT _i (K)	339	346	342	415
Adj. R ²	0.82	0.81	0.78	0.76

* Individual uncertainties estimated for parameters are ≤ 20%

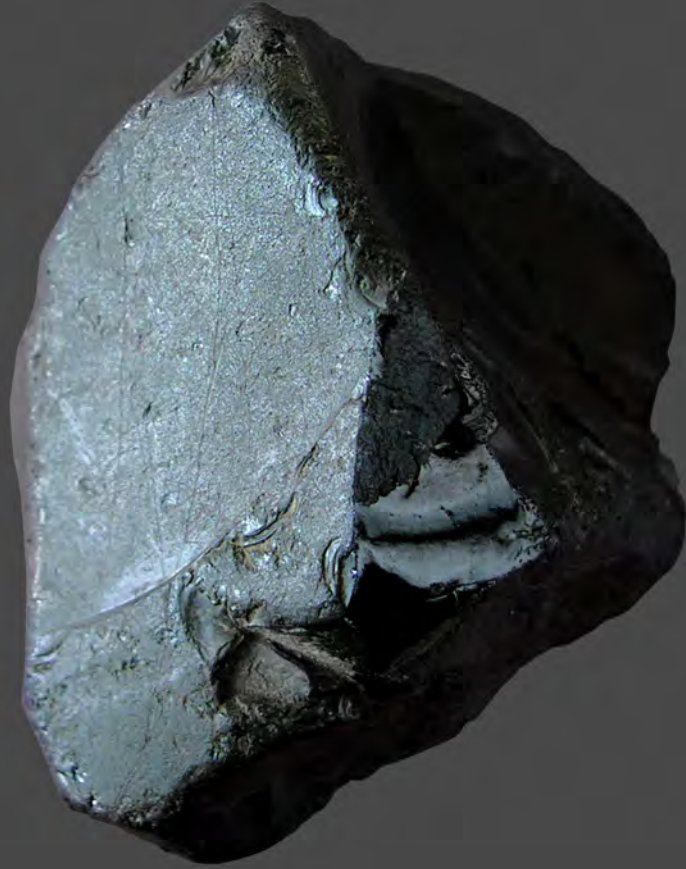


FIGURE 1

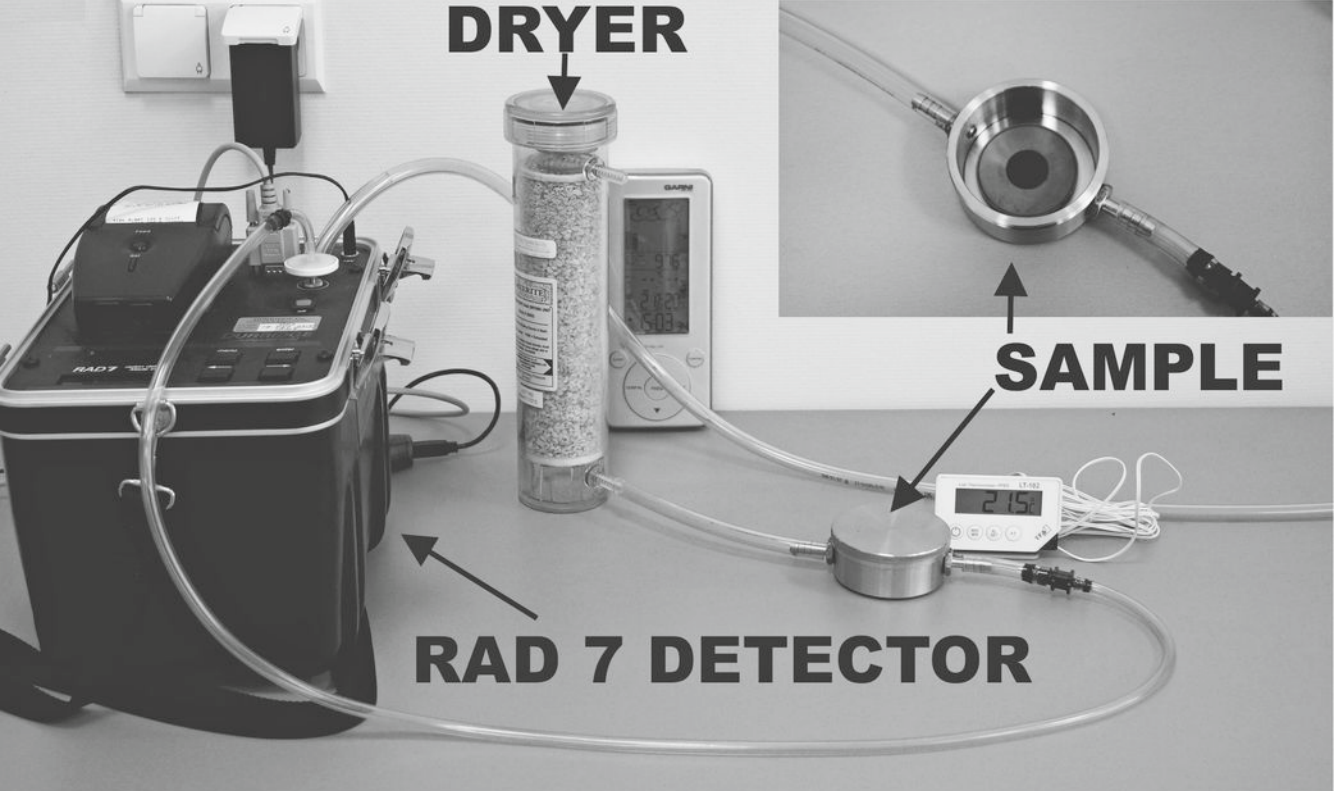


FIGURE 2

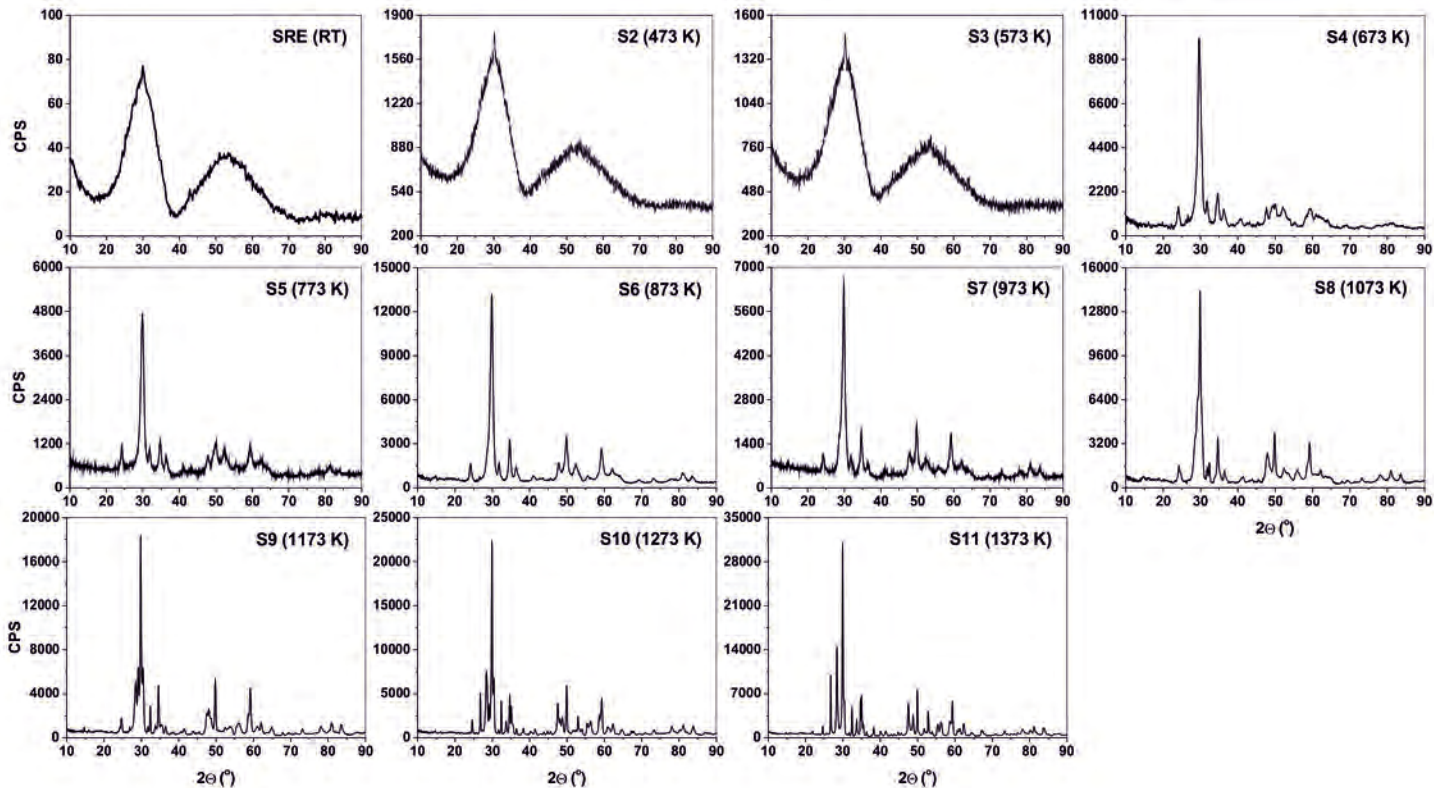


FIGURE 3

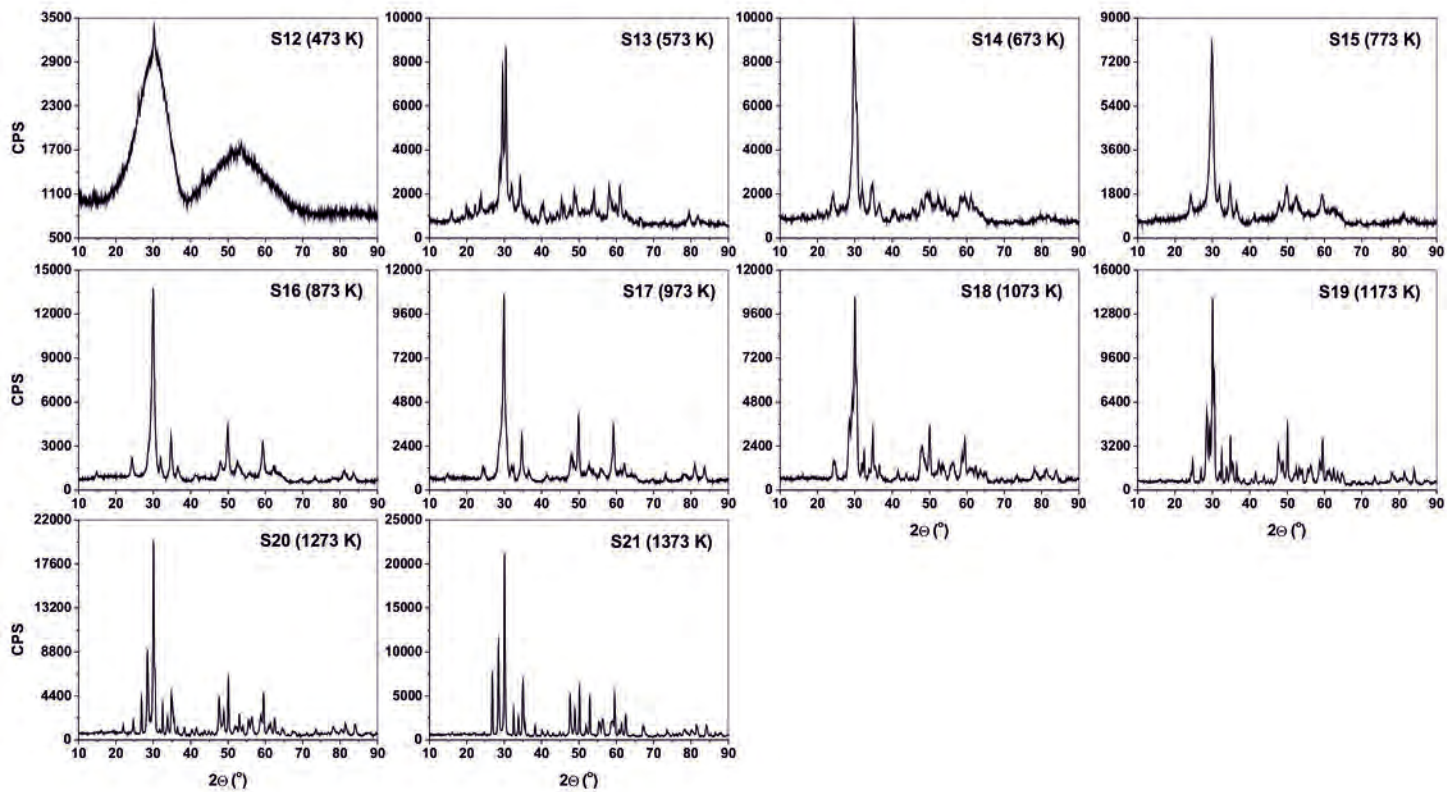


FIGURE 4

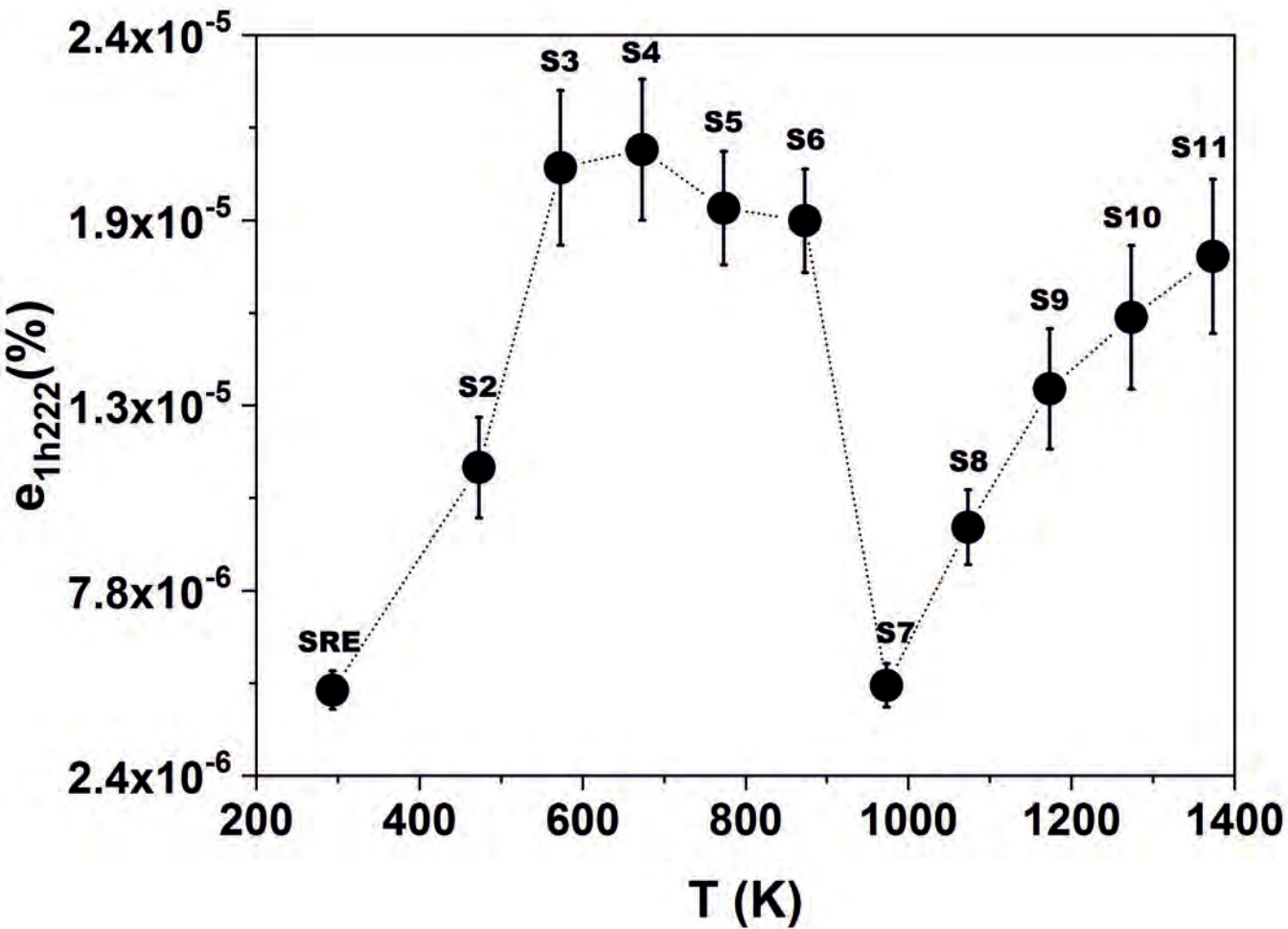


FIGURE 5

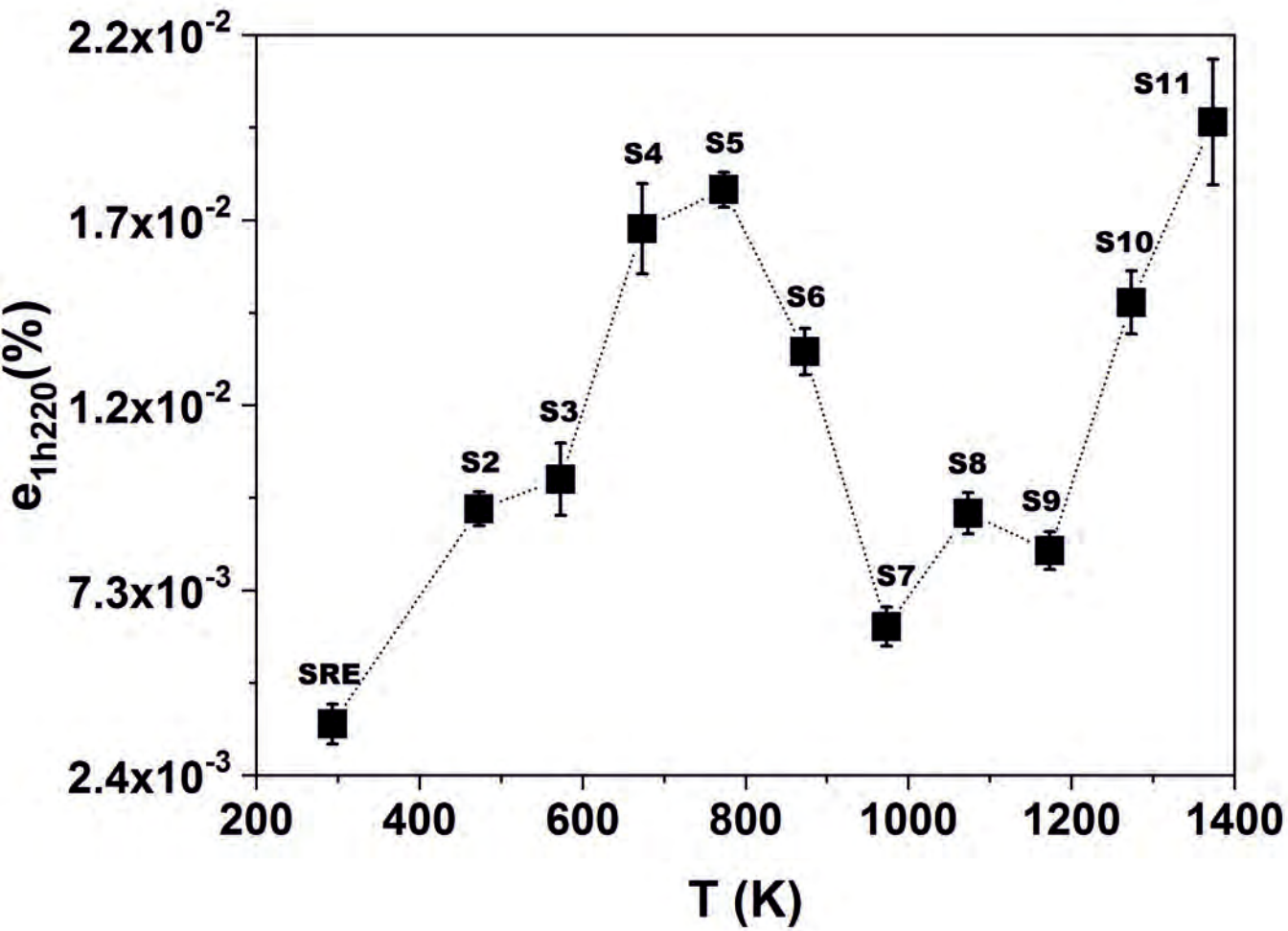


FIGURE 6

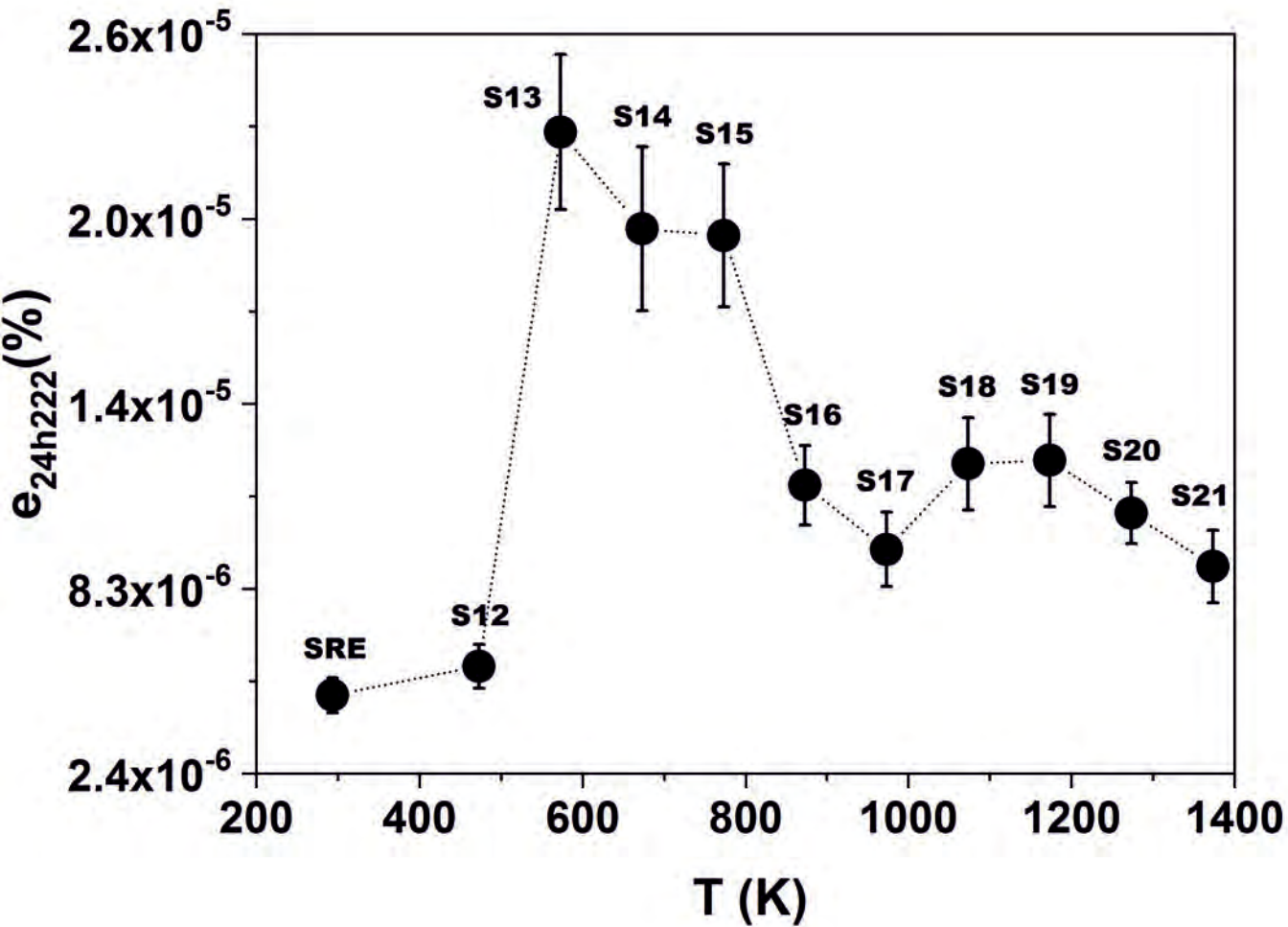


FIGURE 7

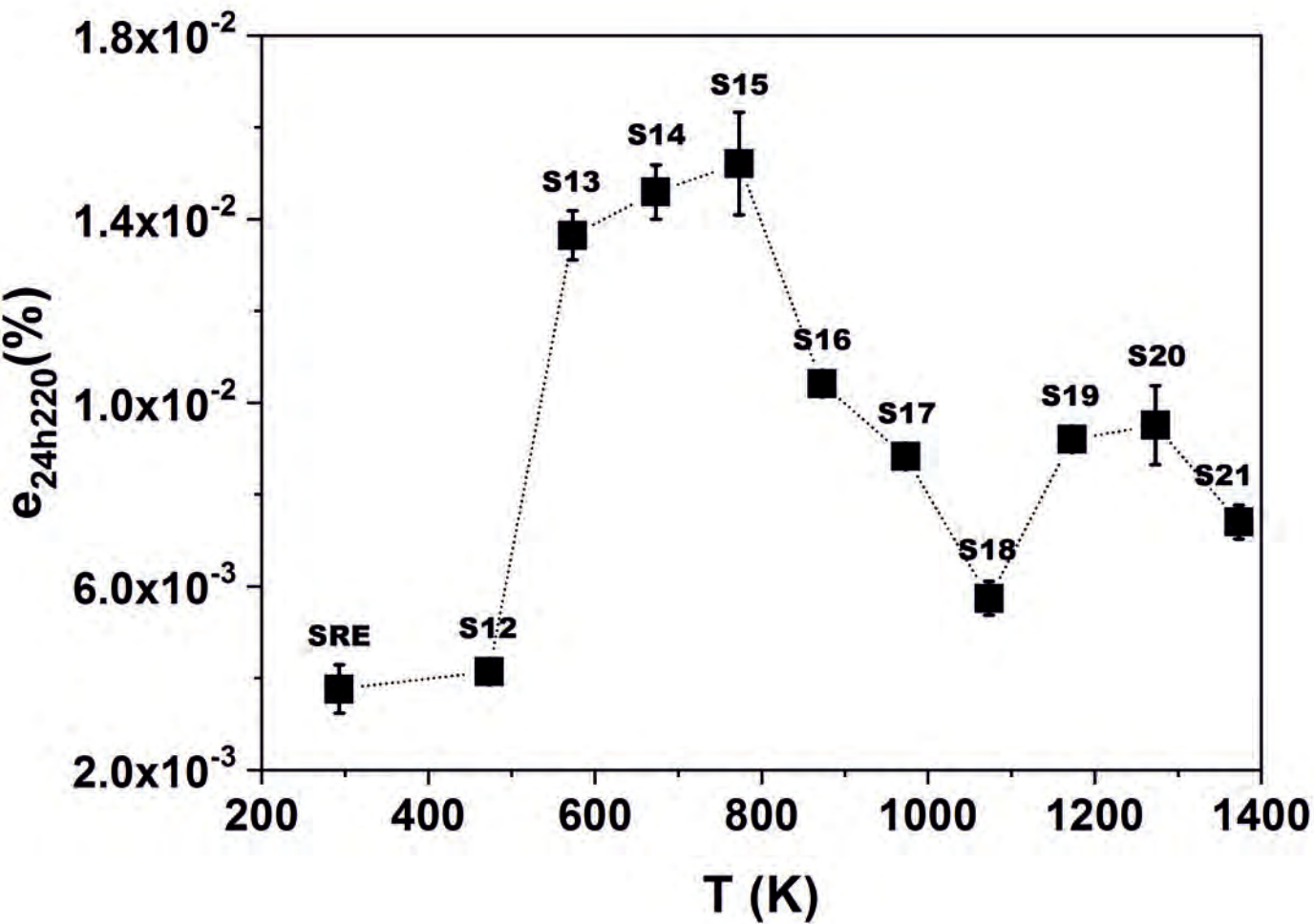


FIGURE 8

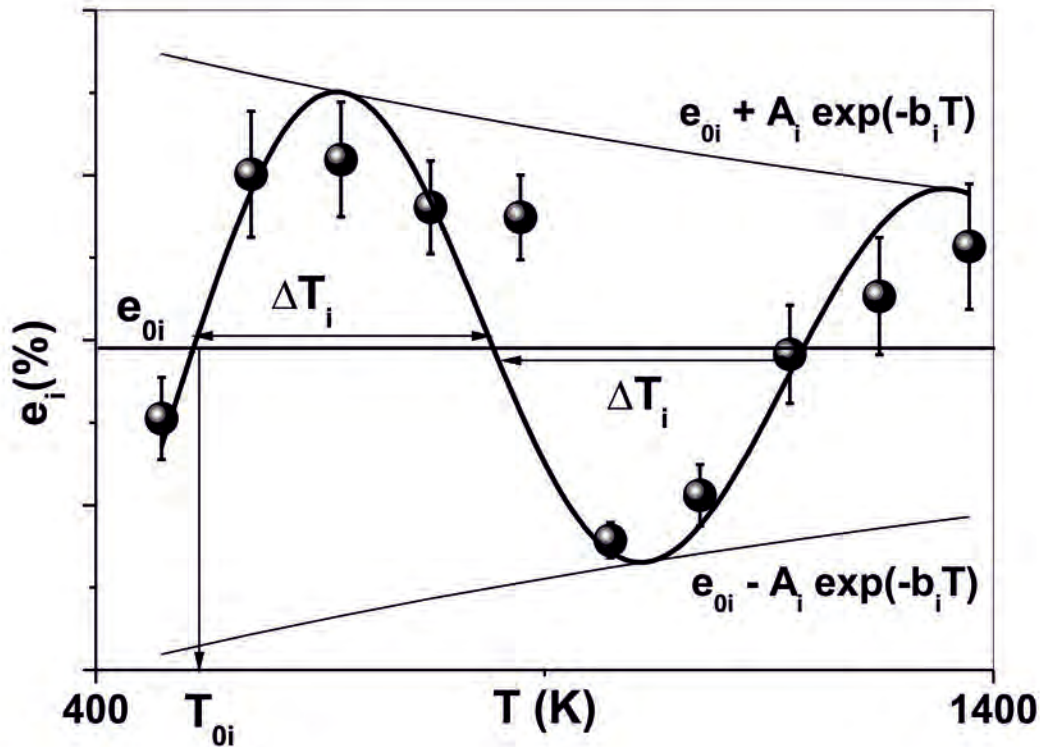


FIGURE 9

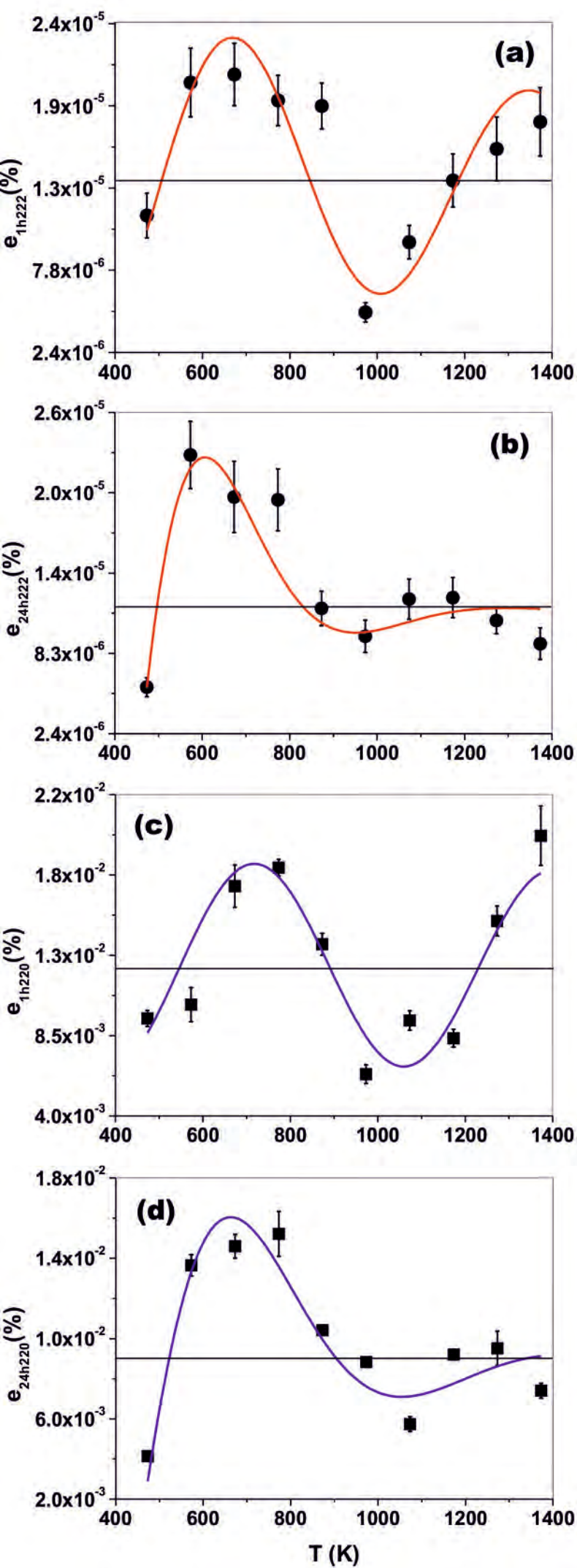


FIGURE 10

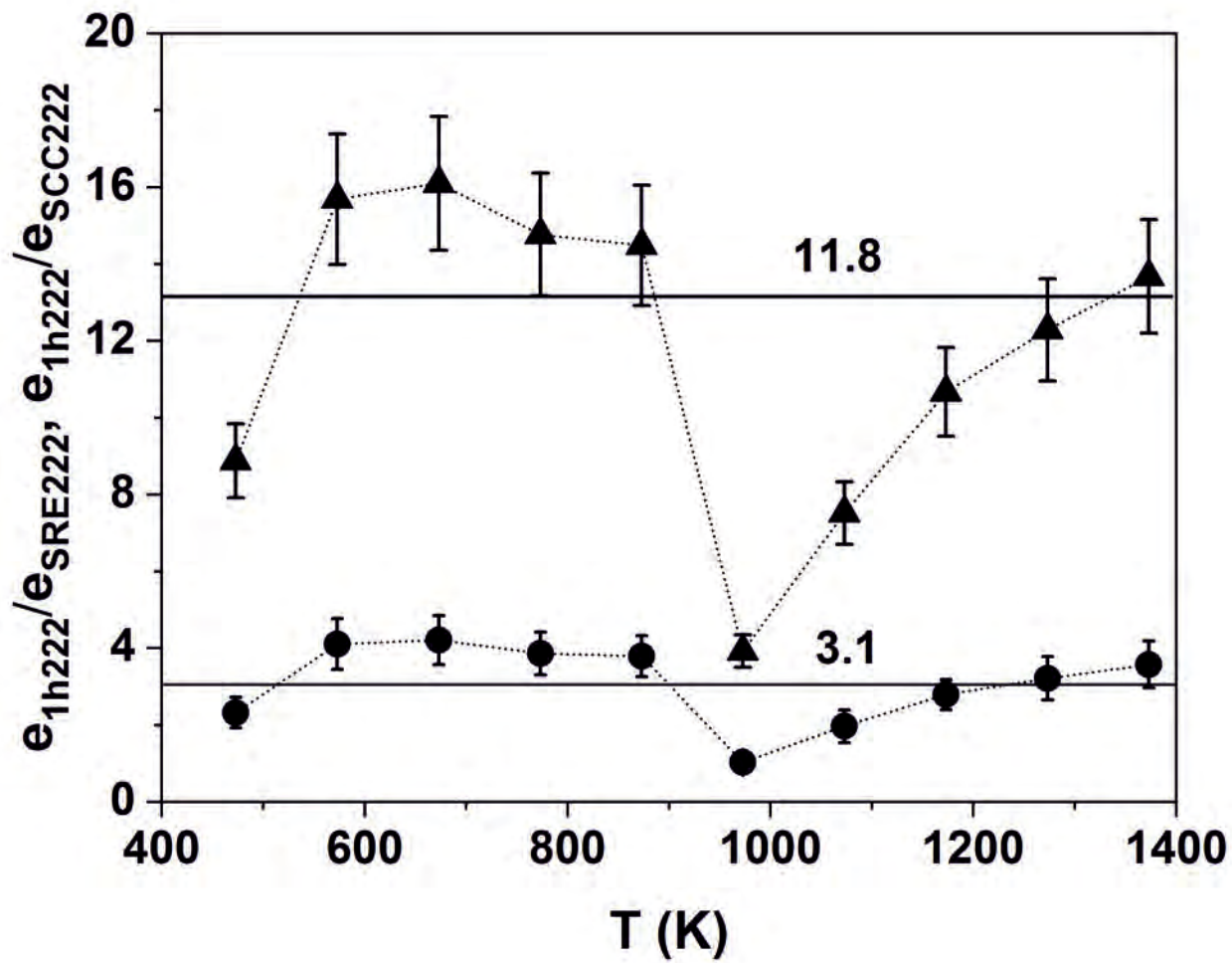


FIGURE 11

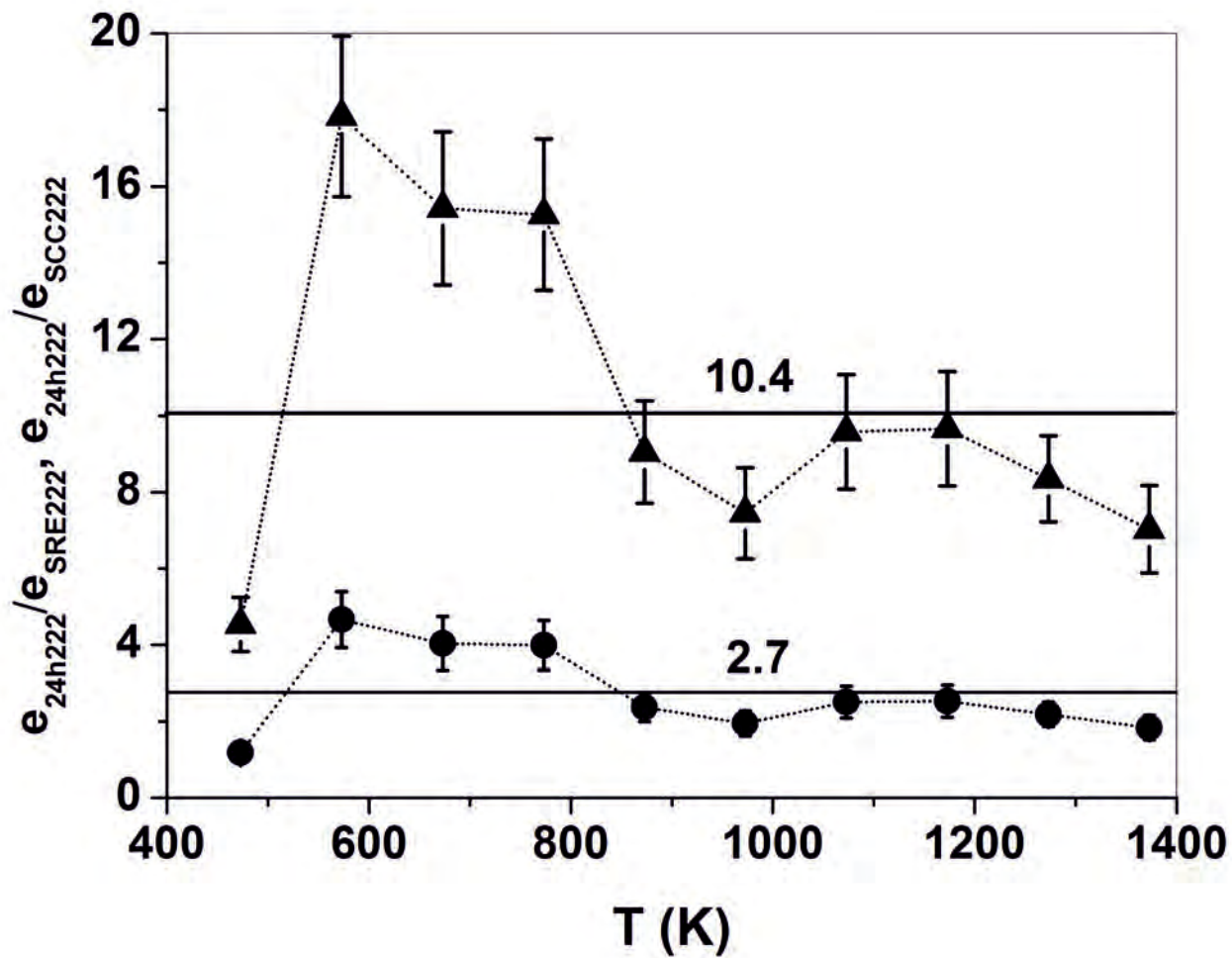


FIGURE 12

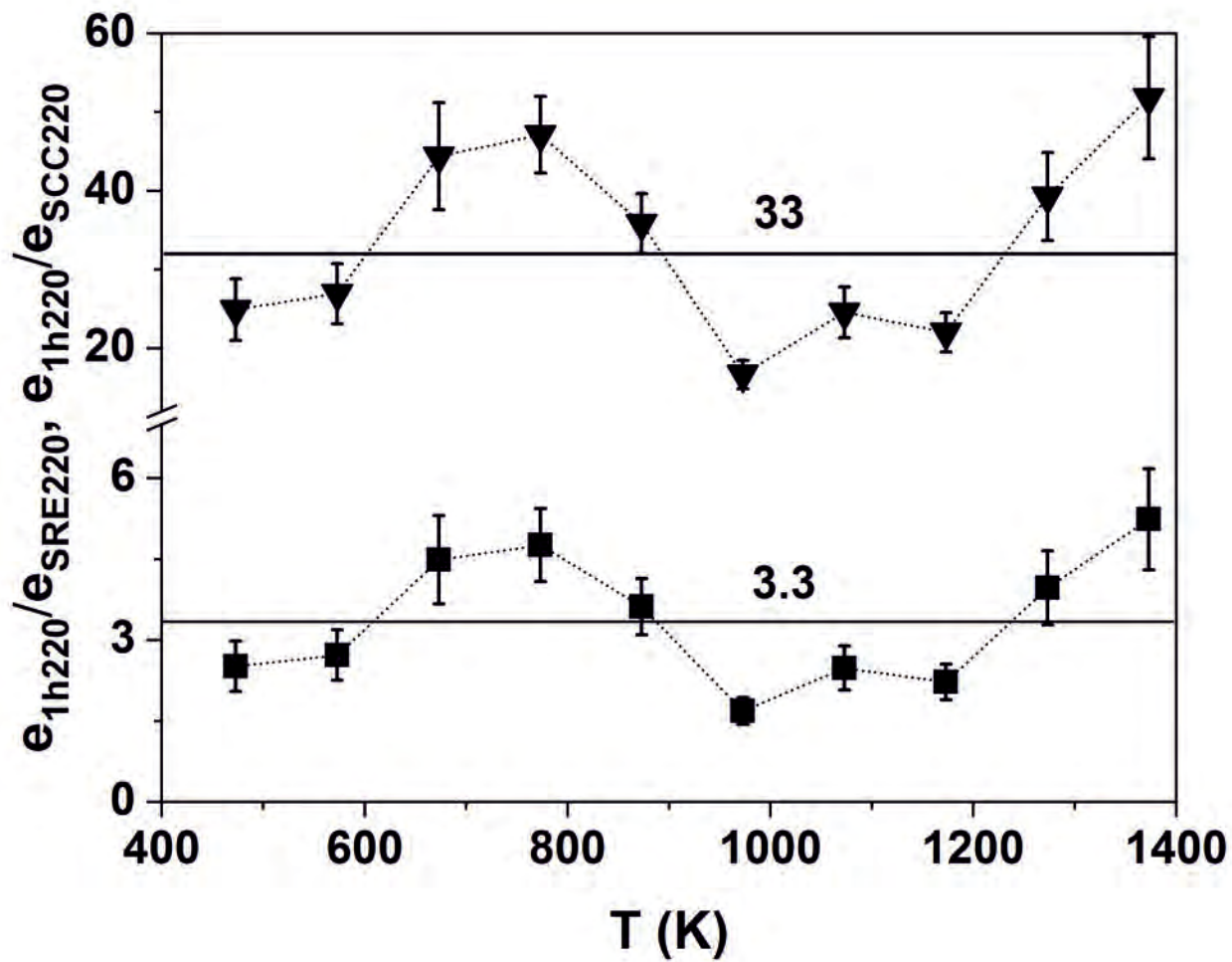


FIGURE 13

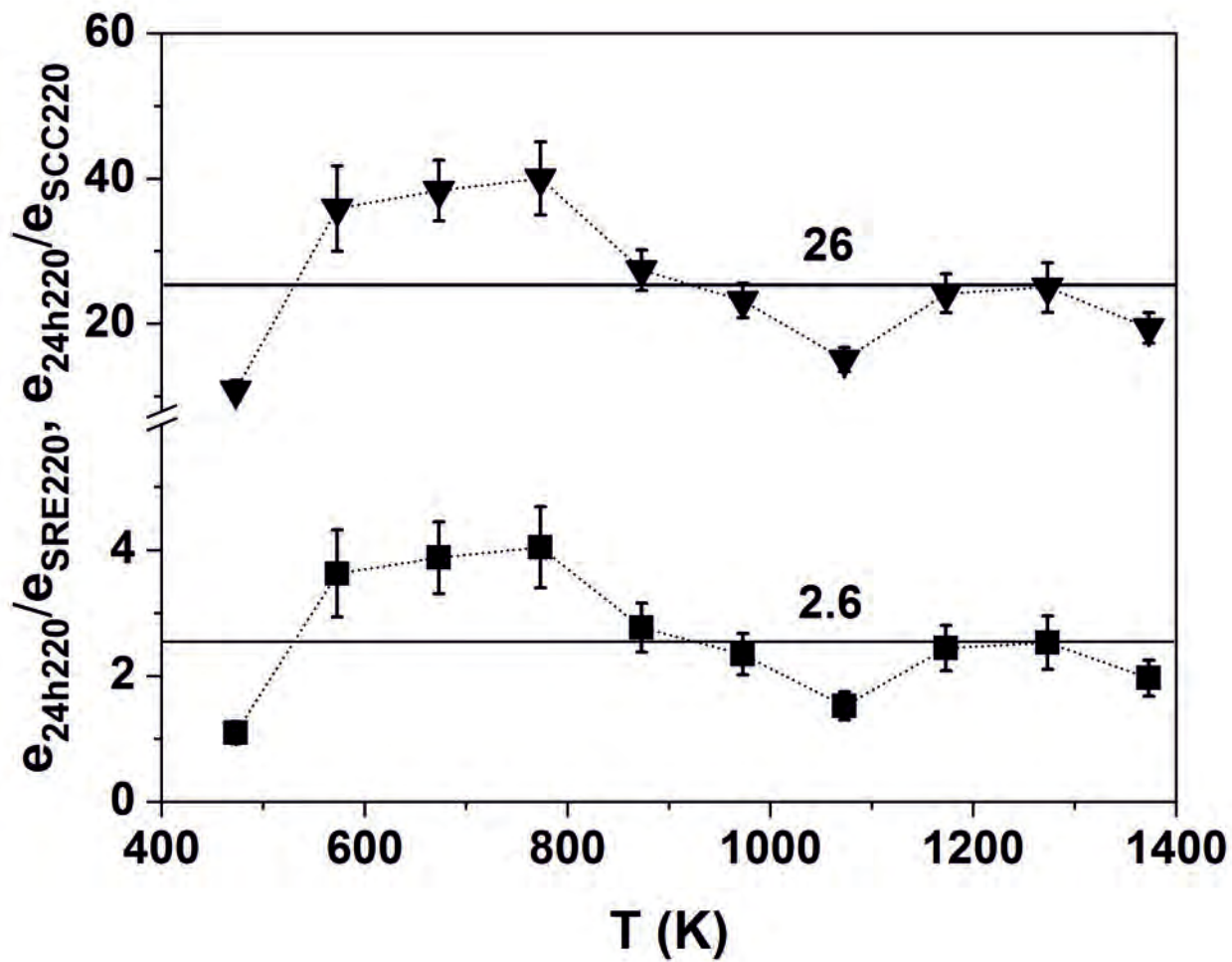


FIGURE 14