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1	Revision 1
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3	Negevite, the pyrite-type NiP ₂ , a new terrestrial phosphide
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14	
15	Abstract
16	Negevite, ideally NiP ₂ , is a new phosphide mineral from pyrometamorphic complex of the Hatrurim
17	Formation (the Mottled Zone), Southern Levant. It is found in phosphide assemblages of the
18	Hatrurim Basin, south Negev Desert, Israel, and Daba-Siwaqa complex, Jordan. The mineral occurs
19	as tiny isometric grains reaching $15\mu m$ in size, and forms intimate intergrowths with other
20	phosphides related to the Fe-Ni-P system. In reflected light, negevite is white with yellowish tint.
21	Isotropic. Reflectance values for COM recommended wavelengths [R (%), λ (nm)] are as follows:
22	54.6 (470), 55.0 (546), 55.3 (589), 55.6 (650). Chemical composition of the holotype specimen
23	(electron microprobe, wt%): Ni 42.57, Co 3.40, Fe 2.87, P 42.93, S 8.33, total 100.10,
24	corresponding to the empirical formula $(Ni_{0.88}Co_{0.07}Fe_{0.06})_{\Sigma 1.01}(P_{1.68}S_{0.31})_{\Sigma 1.99}$. The crystal structure of
25	negevite was solved and refined to $R_1 = 1.73\%$ based on 52 independent observed $[I > 2\sigma(I)]$

26	reflections. The mineral is cubic, space group $Pa\overline{3}$, a 5.4816(5) Å, V 164.71(3) Å ³ , and Z = 4. D_x =
27	4.881(1) g/cm ³ calculated on the basis of empirical formula. Negevite is a first natural phosphide
28	belonging to the pyrite structure type. It is a chemical and structural analog of vaesite, NiS ₂ ,
29	krutovite, NiAs ₂ , and penroseite, NiSe ₂ . The well explored catalytic and photocatalytic properties of
30	a synthetic counterpart of negevite could provide new insights into the possible role of higher
31	phosphides as a source of low-valent phosphorus in prebiotic phosphorylation processes.
32	
33	Keywords: NiP ₂ , negevite, nickel phosphide, Fe-Ni-P system, pyrite, crystal structure,
34	pyrometamorphism, prebiotic phosphorylation, meteorite
35	
36	Introduction

37 Since the discovery of first phosphide mineral, schreibersite (Fe,Ni)₃P (Berzelius 1832), natural 38 iron-nickel phosphides were recognized as mandatory accessory constituents of different meteorite 39 groups (Buchwald 1975; Papike 1998). This mineral family is suggested to play an important role in 40 the highly reduced assemblages of deep planetary interiors (Scott et al. 2007; Dera et al. 2008; Gu et 41 al. 2011, 2012, 2014). Meanwhile, occurrences of terrestrial phosphides of non-anthropogenic origin 42 are quite rare and confined to a few localities worldwide (e.g., Britvin et al. 2015); the most notable 43 one is schreibersite in native iron from basalts of Disko Island, Greenland (Pauly 1969). The 44 practical absence of phosphides in the present-day lithosphere can not be accounted just for highly 45 reducing conditions required for their formation. A possible explanation might imply oxidative 46 decomposition of phosphides at the early stages of Earth evolution, leading to a release of low-47 valent phosphorus required for initiation of prebiotic phosphorylation processes (Pasek et al. 2017; 48 Kitadai and Maruyama 2018; Gibard et al. 2019). The recent discovery of rich phosphide 49 assemblages preserved in geologically juvenile pyrometamorphic complex of the Mottled Zone 50 (Hatrurim Formation) supports this point of view (Britvin et al. 2015).

51 Of twelve phosphide minerals related to the Fe-Ni-P system, nine are currently reported from 52 the Hatrurim Formation (Table 1). We herein present the results of a study of a new mineral negevite, NiP₂ - the first natural phosphide which crystallizes in a pyrite structure type. Negevite is 53 54 named for its type locality in the Negev Desert, Israel; both the mineral and the name have been 55 approved by the Commission on New Minerals, Nomenclature and Classification of International 56 Mineralogical Association (IMA 2013-104). The holotype specimen of negevite is deposited in the 57 collections of the Mineralogical Museum of the Department of Mineralogy, St. Petersburg State 58 University, St. Petersburg, Russia, catalogue number 19604.

59

60 Occurrence and general appearance

61 The Mottled Zone is the world's widest area of sedimentary rocks affected by combustion 62 metamorphism - the suite of processes leading to high-temperature annealing and melting of 63 surficial sediments (Sokol et al. 2005). Outcrops of the Mottled Zone are scattered across the territory of 150×200 km² in Southern Levant in the surroundings of the Dead Sea. The detailed 64 65 description of geological setting along with the hypotheses explaining origin of the Mottled Zone 66 are given in previous works (Gross 1977; Burg et al. 1991; Vapnik et al. 2007; Geller et al. 2012; 67 Novikov et al. 2013). Metamorphic processes were followed by pronounced hydrothermal alteration 68 and weathering stage (Gross 1977; Kolodny et al. 2014), that resulted in emergence of unique 69 mineral assemblages, combining ultrahigh-temperature minerals (Weber and Bischoff 1994; 70 Murashko et al. 2010; Sharygin et al. 2013; Galuskina et al. 2014; Khoury et al. 2016) and 71 secondary phases corresponding to Earth's extreme oxidative and/or alkaline environment (Hauff et 72 al. 1983; Sokol et al. 2011; Galuskin et al. 2013, 2014). 73 Negevite was discovered in phosphide assemblages found along the upper stream of the 74 Halamish Wadi, Hatrurim Basin, Southern Negev desert, Israel (Britvin et al. 2015). The mineral

75 forms irregular minute grains up to 15μm in size intimately intergrown with murashkoite,

76	zuktamrurite, transjordanite or halamishite (Fig. 1). Phosphide assemblages often occur together
77	with magnetite, pyrrhotite and andradite in diopside microbreccia which is severely altered by
78	hydrothermal and weathering processes (Fig. 1). Secondary minerals are comprised by calcite and
79	unidentified X-ray amorphous Ca-Fe hydrous silicates and phosphates. The Hatrurim Basin is a type
80	locality for negevite. Later on, the mineral was identified in one sample of weathered
81	pyrometamorphic paralavas of the Daba-Siwaqa complex (Um Al-Rasas Sub-District, 80 km SSE of
82	Amman, Jordan, 31° 21' 52" N, 36° 10' 55" E) where it forms micrometer-sized aggregates of
83	irregularly fractured grains often intergrown with transjordanite and zuktamrurite (Fig. 2).
84	

85 Physical properties and chemical composition

86 In reflected light, negevite has a white color with bluish tint (Fig. 1A). It is optically isotropic and 87 has no internal reflections. Reflectance values (Table S1) were measured in air by means of a MSF-88 21 spectrophotometer (LOMO, St. Petersburg, Russia) using monochromator slit of 0.4mm and 89 beam diameter of 0.1mm, against WTiC standard. The mineral has no observable cleavage. Due to 90 small size of the grains available, microindentation hardness could not be measured. The calculated 91 density of the holotype specimen calculated for an empirical formula is 4.881(1) g/cm³. Negevite is 92 insoluble in cool 10% HCl. Chemical composition of negevite and associated phosphides was 93 studied in carbon-coated polished thick sections by means of a Hitachi S-3400N scanning electron 94 microscope equipped with (1) an Oxford Instruments AzTec Energy X-Max 20 energy dispersive 95 (EDX) spectrometer and (2) an INCA WAVE 500 wavelength-dispersive (WDX) spectrometer. 96 Preliminary screening of chemical composition and elemental mapping (Fig. 1C, 1D) was carried 97 out in EDX mode whereas quantitative data (Table 1) were obtained with WDX spectrometer using 98 the following analytical standards: GaP (PK α), pyrite (FeK α , SK α), PbSe (SeL α), metallic Co 99 $(CoK\alpha)$, Ni $(NiK\alpha)$, Mo $(MoL\alpha)$, Ag $(AgL\alpha)$. The measurement conditions were: 20kV accelerating 100 voltage, 15nA beam current, peak counting time, 20s peak counting time; 10s background counting

101time. Chemical composition of the holotype negevite (Table 2) corresponds to the empirical formula102 $(Ni_{0.88}Co_{0.07}Fe_{0.06})_{\Sigma 1.01}(P_{1.68}S_{0.31})_{\Sigma 1.99}$ leading to the ideal formula NiP2.

103

104 X-ray single crystal study

105 In order to establish and refine the crystal structure of negevite, a $\sim 10 \mu m$ single-crystal grain of the 106 mineral was extracted from the polished section, mounted onto the glass fiber and subjected to a 107 conventional X-ray single-crystal data collection by means of a Bruker APEX Kappa DUO CCD 108 diffractometer. Data processing and integration routines were performed using a Bruker AXS 109 instrument built-up software (Bruker 2003). The crystal structure of negevite was solved by direct 110 methods and refined assuming NiP₂ formula, by means of a SHELX-2018 software (Sheldrick 2015) 111 incorporated into Olex2 program environment (Dolomanov et al. 2009). The essential parameters of 112 data collection and structure refinement are summarized in Table S2; the complete structural 113 information can be retrieved from the CIF file in the Supplementary Data (see also CSD entry 114 1936174). The insufficient amount of substance had precluded obtaining experimental X-ray 115 powder diffraction data for the mineral. Therefore, powder diffraction pattern of negevite was 116 calculated for CuKal radiation on the basis of refined atomic coordinates and unit cell metrics, 117 using ATOMS v.5.0 software (Dowty 2006) (Table S3).

118

119 **Discussion**

120 Negevite is the first natural phosphide related to a pyrite structure type (Fig. 3). From the crystal-

121 chemical point of view, the mineral can be regarded as phosphide analog of either vaesite, NiS₂,

122 krutaite, NiAs₂, and penroseite, NiSe₂ (Table 3). Though the ideal formula of negevite is NiP₂, its

123 real composition shows substantial (up to 0.48 atoms per MX_2 formula) contents of sulfur (Table 2),

124 suggesting the occurrence of at least partial isomorphism along the join $NiP_2 - NiS_2$. Substitution of

P for S correlates with an increase of the unit cell parameter of negevite as compared to pure NiP₂
(Table 3).

127 Contrary to the synthetic marcasite-type FeP_2 and its natural analog, zuktamurite (Table 1), 128 NiP₂ can exist in two polymorphic modifications, both of which are stable under ambient conditions: 129 (1) cubic, pyrite-type one (Donohue et al. 1968) and (2) monoclinic polymorph (space group $C^{2/c}$) 130 which has no sulfide or phosphide structural analogs (Larsson 1965; Orishchin et al. 2000). Synthetic pyrite-type NiP₂ was first synthesized under high-pressure-temperature environment and 131 132 thus was considered to be metastable, high-pressure polymorph of NiP₂ (Donohue et al. 1968). 133 However, further studies revealed that this modification can be obtained using different methods at 134 atmospheric pressure: via solid-state synthesis (Barry and Gillan 2009) and even by "soft chemistry" 135 solvothermal techniques (Barry and Gillan 2008). Because NiP₂ was found to be promising material 136 in modern electrochemical and catalytic applications (Gillot et al. 2005; Jiang et al. 2014), its 137 chemical and physical characteristics were studied in detail. In particular, it was determined that 138 pyrite-type NiP₂ is a low-temperature modification stable below 600°C (Owens-Baird et al. 2019). 139 Moreover, it was shown that transformation of cubic NiP₂ into the monoclinic modification is 140 irreversible, implying that the presence of cubic NiP₂ in a given assemblage evidences that the 141 temperature has never passed the level of cubic-to-monoclinic transition (Owens-Baird et al. 2019). 142 These results could be helpful in understanding the formation conditions of natural phosphide 143 assemblages of the Mottled Zone (Britvin et al. 2015). However, the presence of sulfur in the 144 composition of natural negevite (Table 2) might stabilize its pyrite-type structure towards higher 145 transition temperature. 146 An interesting feature of negevite from the Daba-Siwaga complex in Jordan is its relative

146 An interesting feature of negevite from the Daba-Siwaqa complex in Jordan is its relative 147 enrichment in Ag (Table 2). It becomes even more attractive taking into account the absence of Ag 148 in closely associated zuktamrurite, FeP₂, and transjordanite, Ni₂P (Fig. 2, Table 2). Ag is an element 149 which is structurally incompatible with Ni. Synthetic phosphide AgP₂ is known though it is not

isostructural with negevite (Möller and Jeitschko 1982). Recent investigation of Ag-doped synthetic
pyrite demonstrated that it may contain up to 0.4 wt% (3820ppm) Ag incorporated via mechanism
of lattice-scale structural defects (Li and Ghahreman 2018). Therefore, one can assume that minor
(0.01 atoms per formula unit) incorporation of Ag into negevite is also permissible.

154

155 Implications

156 Discovery of negevite, naturally occurring NiP₂ might have promising implications 157 considering the general role of phosphides as a source of phosphorus required for phosphorylation 158 processes. It was shown that gentle aquatic oxidation of schreibersite, (Fe,Ni)₃P, results in a release 159 of diverse water-soluble phosphorus compounds which could serve as building blocks during 160 prebiotic phosphorylation at the early stages of Earth evolution (Pasek et al. 2017; Kitadai and 161 Maruyama 2018; Gibard et al. 2019). Other phosphides related to the Fe-Ni-P system (Table 1) were 162 not considered for that role. However, our recent findings demonstrate that these minerals could be 163 formed in a reducing environment of Archean era as well (Britvin et al. 2015). In that respect, 164 unique catalytic and electrochemical properties of the synthetic counterpart of negevite, including its 165 ability to photoinduced water splitting under soft conditions and hydrogenation activity (Gillot et al. 166 2005; Jiang et al. 2014), could provide new routes for further exploration of possible natural 167 phosphorylation pathways. The discovery of negevite-bearing phosphide assemblages of the 168 Mottled Zone undoubtedly imply the extremely reducing ("super-reducing") environment occurred 169 during their crystallization. This enigma of the Hatrurim Formation is not yet resolved and requires 170 gathering of further evidences.

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321 List of figure captions

322	Figure 1. Negevite in phosphide assemblages of the Halamish wadi, Hatrurim Basin, Southern
323	Negev Desert, Israel. (A) Photomicrograph in reflected light. (B) The same area, SEM BSE image.
324	(C) Element distribution map for phosphorus. (D) Element distribution map for iron and nickel.
325	Abbreviations: Ng, negevite (NiP2); Zk, zuktamrurite (FeP2); Tj, transjordanite (Ni2P). (Color
326	online.)
327	
328	Figure 2. Fractured grain of Ag-bearing negevite among secondary Ca-Fe-silicates in altered
329	pyroxene-anorthite paralava. Note grains of Ag-free transjordanite and zuktamrurite associated with
330	negevite. Daba-Siwaqa complex, Transjordan plateau, Jordan. SEM BSE image. Abbreviations: Ng,
331	negevite (NiP2); Zk, zuktamrurite (FeP2); Tj, transjordanite (Ni2P).
332	
333	Figure 3. Crystal structure of negevite (pyrite structure type). (A) General view: corner-sharing
334	[NiP ₆] octahedra (green) linked by P–P "dumbbell" bonds (yellow) characteristic of the pyrite
335	structure. (B) Skeletal view of [NiP ₆] octahedron connected to [P(Ni ₃ P)] tetrahedra. (C) Slice of
336	substructure framework composed of corner-sharing [NiP6] octahedra. (D) Slice composed of
337	[P(Ni ₃ P)] tetrahedra. Legend: green, Ni [4 <i>a</i> ; 0, 0, 0]; yellow, P [8 <i>c</i> ; 0.3844(1), <i>x</i> , <i>x</i>]. (Color online.)
338	

340 Tables

341

342

343 Table 1. Natural phosphides related to the Fe-Ni-P system

344

Mineral	End member	Structure type	Occurrence ^a	Reference
Schreibersite	Fe ₃ P	$Fe_3P(\overline{I4})$	M T	Berzelius (1832)
Barringerite	Fe ₂ P	$Fe_2P(\overline{P62m})$	M T	Buseck (1969); Britvin et al. (2017)
Allabogdanite	Fe ₂ P	Co ₂ Si (Pnma)	М	Britvin et al. (2002)
Murashkoite	FeP	MnP (Pnma)	Т	Britvin et al. (2019b)
Zuktamrurite	FeP ₂	Marcasite (Pnnm)	Т	Britvin et al. (2019a)
Melliniite	(Ni,Fe) ₄ P ^b	$\operatorname{Au}_{4}\operatorname{Al}(P2_{1}3)$	М	Pratesi et al. (2006)
Nickelphosphide	Ni ₃ P	$Fe_3P(\overline{I4})$	М	Britvin et al. (1999)
Nazarovite	Ni ₁₂ P ₅	Ni ₁₂ P ₅ (<i>I</i> 4/ <i>m</i>)	M T	Britvin et al. (2019c) ^c
Transjordanite	Ni ₂ P	$Fe_2P(\overline{P62m})$	Т	Britvin et al. (2015)
Orishchinite	Ni ₂ P	Co ₂ Si (Pnma)	Т	Britvin et al. (2019d) ^c
Halamishite	Ni ₅ P ₄	$Ni_5P_4(P6_3mc)$	Т	Britvin et al. (2015)
Negevite	NiP ₂	Pyrite ($Pa\overline{3}$)	Т	This work

345 *a* Occurrence: M – meteoritic; T – terrestrial (the Hatrurim Formation). *b* End-member is not known.

^c Approved by the Commission on New Minerals, Nomenclature and Classification of International

347 Mineralogical Association, IMA 2019-013 (nazarovite) and 2019-039 (orishchinite).

349	Table 2. Chemical	composition (of negevite and	associated phosphides
		1		1 1

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Mineral ^a	Ng	Ng	Ng	Tj	Zk	Ng	Zk	Tj
Locality ^b	HB	HB	HB	HB	HB	DS	DS	DS
Notes	Holotype		Fig. 1	Fig. 1	Fig. 1	Fig. 2	Fig. 2	Fig. 2
				Wt%				
Ni	42.57	38.25	37.77	71.95	8.83	38.79	15.66	74.80
Co	3.40	2.92	3.44	0.30	1.24	3.40	0.37	0.15
Fe	2.87	6.15	6.41	6.26	36.74	5.12	30.87	3.30
Ag	-	-	-	-	-	1.01	-	-
Р	42.93	40.05	40.07	21.14	52.25	39.51	51.43	21.09
S	8.33	12.12	12.78	0.00	1.51	10.82	0.00	0.00
Se	-	-	0.24	-	-	-	-	-
Total	100.10	99.49	100.71	99.65	100.57	98.65	98.33	99.34
			Form	ula amou	nts (3 <i>apfu</i>	ı)		
Ni	0.88	0.79	0.77	1.82	0.18	0.82	0.32	1.90
Со	0.07	0.06	0.07	0.01	0.02	0.07	0.01	0.00
Fe	0.06	0.13	0.14	0.17	0.77	0.11	0.67	0.09
Ag						0.01		
Σ	1.01	0.98	0.98	2.00	0.97	1.01	1.00	1.99
Р	1.68	1.56	1.55	1.01	1.97	1.57	2.00	1.01
S	0.31	0.46	0.48		0.06	0.42		
Se			0.004					
Σ	1.99	2.02	2.03	1.01	2.03	1.99	2.00	1.01

351

352 ^{*a*}Abbreviations: Ng, negevite; Tj, transjordanite; Zk, zuktamrurite. ^{*b*}Localities: HB, Hatrurim Basin,

353 Israel; DS, Daba-Siwaqa complex, Jordan. ^{*c*}The bar "-" denotes below detection limit (<0.05 wt%).

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	Mineral	Negevite	Synthetic	Vaesite	Krutovite	Penroseit		
357 358	Table 3. Unit cell parameters of negevite and related minerals and compounds							

Mineral	Negevite	Synthetic	Vaesite	Krutovite	Penroseite
Ideal formula	NiP ₂	NiP ₂	NiS ₂	NiAs ₂	NiSe ₂
a (Å)	5.4816	5.4475	5.6679	5.7634	5.988
Reference ^a	[1]	[2]	[3]	[4]	[5]

a [1] This work; [2] Owens-Baird et al. (2019); [3] Kerr (1945); [4] Donohue et al. (1968);

361 Vinogradova et al. (1976); [5] Bindi et al. (2008).

365 Figures

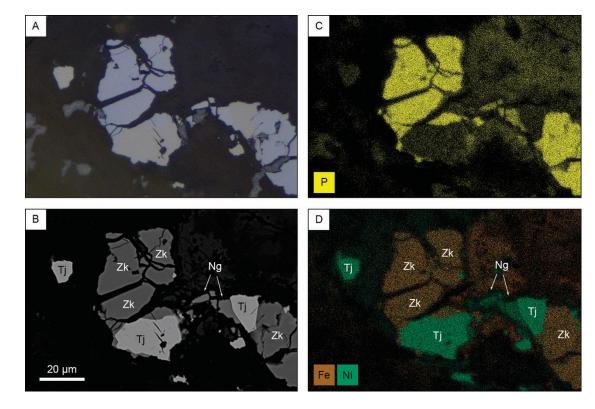


Figure 1

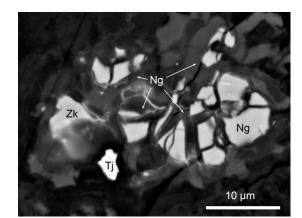
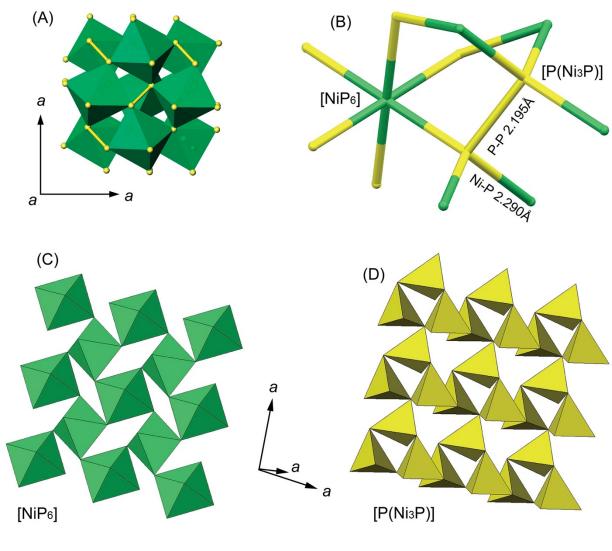


Figure 2



374 375 376

Figure 3