

1 **Journal**

2 Invited contribution — *Special Issue* American Mineralogist on Halogens in Planetary Bodies  
3 (Editors Anita Cadoux; Keith Putirka)

4 **Title**

5 The distribution and abundance of halogens in eclogites: an in situ SIMS perspective of the  
6 Raspas Complex (Ecuador)

7 **Revision 2**

8 **Authors**

9 Benjamin M. Urann<sup>1,2\*</sup>, Véronique Le Roux<sup>2</sup>, Timm John<sup>3</sup>, Grace M. Beaudoin<sup>4</sup>, Jaime D.  
10 Barnes<sup>4</sup>

11

12 <sup>1</sup>MIT-WHOI Joint Program, Marine Geology and Geophysics, Woods Hole Oceanographic  
13 Institution, Woods Hole, MA 02543 USA

14 <sup>2</sup>Geology and Geophysics Department, Woods Hole Oceanographic Institution, Woods Hole,  
15 MA 02543 USA

16 <sup>3</sup>Institut für Geologische Wissenschaften, Freie Universität Berlin, Malteserstrasse 74-100,  
17 12249 Berlin, Germany

18 <sup>4</sup>Department of Geological Sciences, University of Texas, Austin, TX 78712 USA

19

20 \*email: burann@whoi.edu

21 **Abstract**

22 We present in situ secondary ion mass spectrometry (SIMS) and electron microprobe  
23 analyses of co-existing garnet, omphacite, phengite, amphibole, and apatite, combined with

24 pyrohydrolysis bulk-rock analyses to constrain the distribution, abundance, and behavior of  
25 halogens (F and Cl) in six MORB-like eclogites from the Raspas Complex (Southern Ecuador).  
26 In all cases concerning lattice-hosted halogens, F compatibility decreases from apatite (1.47–3.25  
27 wt %), to amphibole (563–4727 µg/g), phengite (610–1822 µg/g), omphacite (6.5–54.1 µg/g)  
28 and garnet (1.7–8.9 µg/g). The relative compatibility of Cl in the assemblage is greatest for  
29 apatite (192–515 µg/g), followed by amphibole (0.64–82.7 µg/g), phengite (1.2–2.1 µg/g),  
30 omphacite (<0.05–1.0 µg/g) and garnet (<0.05 µg/g). Congruence between SIMS-reconstructed  
31 F bulk abundances and yield-corrected bulk pyrohydrolysis analyses indicates that F is primarily  
32 hosted within the crystal lattice of eclogitic minerals. However, SIMS-reconstructed Cl  
33 abundances are a factor of five lower, on average, than pyrohydrolysis-derived bulk  
34 concentrations. This discrepancy results from the contribution of fluid inclusions, which may  
35 host at least 80% of the bulk rock Cl. The combination of SIMS and pyrohydrolysis is highly  
36 complementary. Whereas SIMS is well suited to determine bulk F abundances, pyrohydrolysis  
37 better quantifies bulk Cl concentrations, which include the contribution of fluid inclusion-hosted  
38 Cl. Raspas eclogites contain 145–258 µg/g F and at least 7–11 µg/g Cl. We estimate that ~95% of  
39 F is retained in the slab through eclogitization and returned to the upper mantle during  
40 subduction, whereas at least 95% of subducted Cl is removed from the rock by the time the slab  
41 equilibrates at eclogite facies conditions. Our calculations provide further evidence for the  
42 fractionation of F from Cl during high-pressure metamorphism in subduction zones. Although  
43 the HIMU mantle source (dehydrated oceanic crust) is often associated with enrichments in Cl/K  
44 and F/Nd, Raspas eclogites show relatively low halogen ratios identical within uncertainty to  
45 DMM. Thus, the observed halogen enrichments in HIMU ocean island basalts require either

46 further fractionation during mantle processing, or recycling of a halogen-enriched carrier  
47 lithology such as serpentinite into the mantle.

48

49 Keywords: eclogite, halogens, subduction, SIMS, nominally anhydrous minerals, HIMU

50 **1. Introduction**

51

52 Our understanding of the abundance and distribution of halogens in subducted slabs is  
53 limited. Hydrothermally altered oceanic crust (AOC) is thought to be a major halogen carrier  
54 during subduction, where F and Cl substitute into the hydroxyl sites of hydrous minerals such as  
55 amphibole and mica (Ito et al. 1983; Philippot et al. 1998; Van den Bleeken and Koga 2015;  
56 Barnes et al. 2018). Bulk estimates of pre-subduction AOC vary from 50–253 µg/g Cl and 216–  
57 400 µg/g F (Ito et al. 1983; Straub and Layne 2003; Barnes and Cisneros 2012; Van den Bleeken  
58 and Koga 2015; Chavrit et al. 2016). More recently, primary halogen measurements of altered  
59 oceanic crust from the East Pacific Rise (Penrose-type oceanic crust) and Atlantis Bank (SWIR,  
60 ultra-slow spreading) have shown extreme variability with both stratigraphic depth and lithology  
61 (Kendrick 2019a; 2019b). Bulk halogen measurements of blueschists and eclogites thought to  
62 represent high-pressure metamorphic AOC are also sparse. In a study of mélange rocks from  
63 Syros, Greece, Marschall et al. (2009) measured the bulk Cl content of eclogites (28–60 µg/g)  
64 and called into question previous estimates of eclogitized AOC Cl abundances, e.g. 100–200  
65 µg/g Cl of Philippot et al. (1998), speculating that they may be overestimated. Pagé et al. (2016)  
66 analyzed a suite of blueschists from northwest Turkey; their results (8–22 µg/g Cl) also indicate  
67 that bulk eclogitized AOC could host less Cl than previously thought. Debret et al. (2016)  
68 reconstructed bulk halogen concentrations (57–79 µg/g Cl and 10–62 µg/g F) from in situ

69 secondary ion mass spectrometry (SIMS) analyses for both blueschists and eclogites from the  
70 Western Alps, cautioning that typical bulk halogen measurements are not representative of  
71 prograde conditions due to potential retrogression during exhumation.

72 In situ halogen data in eclogitized AOC are also limited to a handful of studies, yet provide  
73 important first order observations. Debret et al. (2016) suggested that approximately 50% of F  
74 and 90% of Cl in the initial bulk-rock are lost to fluids during the first 80 km of subduction.  
75 Hughes et al. (2018) analyzed in situ halogen abundances of eclogites from the Western and  
76 Central Alps, and provided eclogitized AOC estimates of as low as 3 µg/g Cl. Finally, Pagé et al.  
77 (2016) reported low Cl and high F abundances in blueschist-hosted minerals, and concluded that  
78 Cl is primarily expelled at shallow depths, whereas F is retained to at least 80 km depth.

79 Estimates of halogen fluxes in subduction zones are typically derived from the difference  
80 between the trench inputs and extrusive products of arc magmatism (e.g. (Mather et al. 2006;  
81 Sadofsky et al. 2008; Pyle and Mather 2009; Freundt et al. 2014). Straub and Layne (2003)  
82 inferred a global arc recycling efficiency of 77–103% (Cl), and ~4–5% (F), implying near-total  
83 return of Cl to surface reservoirs whereas most F is returned to the mantle. John et al. (2011)  
84 reached similar conclusions based on Cl mass balance constraints regarding in- and out-fluxes  
85 from the subducted slab perspective. Barnes et al. (2018) also concluded that F is efficiently  
86 returned to the mantle, however their estimated Cl fluxes are more variable, indicating that some  
87 Cl could be delivered back into the mantle by serpentinites (John et al. 2011; Kendrick et al.  
88 2017). In a recent study, Pagé et al. (2018) found that Himalayan antigorite serpentinites host  
89 significant quantities of F (50–650 µg/g) and are enriched in Cl (8.8–35 µg/g) with respect to  
90 depleted mantle (11.0 µg/g F and 0.51 µg/g Cl, Salters and Stracke (2004)). Taken together,  
91 these studies illustrate the challenge of estimating halogen input fluxes. A better understanding

92 of the inherent variability within different lithologies used in flux calculations could help to  
93 refine such estimates.

94 In particular, the halogen content of eclogitized AOC remains largely unconstrained. Here  
95 we discuss the F and Cl abundances of Raspas Complex eclogites, a Cretaceous eclogite suite  
96 that we consider to be an ideal analog for subducted AOC. We combine in situ SIMS and EMP  
97 analyses of minerals with pyrohydrolysis-derived bulk halogen abundances and discuss the  
98 relative contribution of minerals, grain boundaries and fluid inclusions to the total budget of  
99 halogens in subducted eclogites. Finally, we present new constraints on slab halogen fluxes and  
100 discuss the implications for the signature of eclogitized AOC in the HIMU mantle source.

101

## 102 **2. Geologic Context**

103 An accreted parcel of oceanic lithosphere, the early Cretaceous Raspas Complex consists of  
104 blueschist- and eclogite-facies rocks (Feininger 1980; 1987; Maruyama et al. 1996). The high-  
105 pressure metamorphic suite is bound by the La Palma-El Guayabo Shear Zone to the north and  
106 lower grade greenschists and amphibolites to the south (Gabriele 2002). While no pillow  
107 structures or sheeted dikes were observed in the field, fluid-immobile REE and HFSE  
108 abundances, Nb/Zr and Hf/Yb systematics, TiO<sub>2</sub>/Yb ratios, and Nd isotope systematics provide a  
109 well-reasoned rationale for inferred protoliths (John et al. 2010; Halama et al. 2011). Based on  
110 these characteristics, the blueschists and eclogites were derived from seamount-like basalt and  
111 normal mid-ocean ridge basalt (MORB) protoliths, respectively. Both underwent similar peak  
112 metamorphic conditions (based on major element thermobarometry) of ~600°C to a maximum  
113 burial depth of approximately 60 km (John et al. 2010). Raspas eclogites show evidence for low-  
114 temperature seafloor alteration typical of basalts based on O isotopic compositions, and occur as

115 two geochemically distinct groups: a MORB-type group (this study) with LREE depleted  
116 patterns, and a zoisite eclogite group characterized by significant incompatible trace-element  
117 enrichments which are interpreted to have been derived from serpentinite-derived dehydration  
118 fluids, and which also display a fluid component indicative of some minor sediment contribution  
119 (Halama et al. 2011; Herms et al. 2012). Thus, the Raspas MORB-type group provides an  
120 excellent analog to study the in situ distribution of halogens in eclogitized AOC processed  
121 through subduction systems globally.

### 122                   3. Sample Description

123                   We chose a set of six minimally retrogressed MORB-type eclogites which have been  
124 previously described by John et al. (2010). Samples consist of omphacite + garnet + rutile +  
125 apatite  $\pm$  amphibole  $\pm$  phengite  $\pm$  quartz with minor secondary titanite. Subhedral to euhedral  
126 garnet porphyroblasts vary in size from  $<100\text{ }\mu\text{m}$  to 4mm, often hosting inclusions of rutile and  
127 quartz. Some samples exhibit weakly deformed fabrics characterized by lineated omphacite. A  
128 single sample, SEC42-06, shows a strongly foliated texture with alternating layers of nearly  
129 monomineralic garnet and omphacite domains. Modal mineralogy varies significantly between  
130 samples, from nearly bimineralic garnet + omphacite to  $\sim30\%$  modal amphibole; such variability  
131 at similar P-T conditions is likely the result of bulk chemical compositional effects (Beinlich et  
132 al. 2010; John et al. 2010; Rebay et al. 2010). Representative photomicrographs of sample  
133 textures are presented in Fig. 1 and S1. Petrographic observations reveal near-ubiquitous fluid  
134 inclusions (FI). All samples contain variable quantities of FI hosted in omphacite and garnet,  
135 typically located in mineral cores (Fig. 1; see also Herms et al. (2012)). In SEC46-02,  
136 monomineralic amphibole veinlets crosscut eclogite textures and contain secondary titanite  
137 overgrowths rimming rutile. Directly adjacent to these veinlets, FI-rich omphacite appears

138 opaque in thin section (Fig. 1), which is unique to this sample and appears to reflect fluid  
139 infiltration.

140

141 **4. Methods**

142 **4.1 SIMS**

143 SIMS analyses were conducted at the Northeast National Ion Microprobe Facility  
144 (NENIMF) at the Woods Hole Oceanographic Institution using a Cameca IMS 1280. We follow  
145 the methods outlined in Urann et al. (2017), adding OH and S measurements to the protocol.  
146 150 $\mu$ m polished thick-sections were used for all analyses. Samples were cleaned with ethanol  
147 and deionized water to remove surface contamination, then placed in a vacuum oven to dry.  
148 Samples were gold coated to ~160 nm thickness and placed into a vacuum chamber for storage  
149 until analysis. Samples were loaded into the instrument sample chamber at least 12 hours prior to  
150 analysis, to allow adequate pump down time and achieve chamber pressures of no more than  
151 ~5x10<sup>-9</sup> torr. Each spot location was chosen to avoid surface cracks, which are known to  
152 influence analyses (Urann et al. 2017). A primary Cs<sup>+</sup> beam of 5.0–7.5 nA was sputtered through  
153 the sample surface with a 30 x 30  $\mu$ m<sup>2</sup> raster and a 400  $\mu$ m field aperture, allowing only  
154 transmission of ions from the innermost 5 x 5  $\mu$ m<sup>2</sup> of the beam crater. Secondary magnet mass  
155 calibration was done before each measurement with mass resolving power of > 6000 (m/ $\Delta$  m at  
156 10 % peak height). We measured <sup>19</sup>F/<sup>30</sup>Si, <sup>16</sup>O/<sup>1</sup>H/<sup>30</sup>Si, <sup>32</sup>S/<sup>30</sup>Si, and <sup>35</sup>Cl/<sup>30</sup>Si ratios in glass  
157 reference materials D51-3, D52-5, ALV519-4-1, 46D, 1649-3, 1654-3 6001, and AII107-D20 to  
158 produce a calibration slope for each one-week analytical session (Fig. S2). Calibration slopes ( $m$ )  
159 were obtained by plotting measured isotope ratios ( $x$ ) against known reference material  
160 concentrations ( $y$ ) of the form  $y = mx$  for each element of interest. Sample unknowns were then

161 calculated by multiplying measured ratios by  $m$ . Calibration slope uncertainties were assessed  
162 utilizing a bootstrapping technique (5000 iterations) to derive confidence intervals.  
163 Measurements where the internal precision (one standard deviation) was greater than 20% were  
164 screened and excluded from the data set; these low precision measurements (less than ten in  
165 total) were associated with elevated C (not reported) and Cl concentrations, and are likely the  
166 result of analyses conducted on micro-fractures. After screening, analytical uncertainties over  
167 five counting cycles (internal precision: typical standard error ~1 % OH, <1.5 % F, <3% Cl)  
168 were combined with calibration slope uncertainties (accuracy: typical uncertainty 6% OH, 5% F,  
169 <10% Cl, 95% confidence intervals) to yield no more than 8%, 6% and 16% total uncertainty  
170 (2SE, 95% confidence intervals) for OH, F and Cl measurements, respectively. Since mineral-  
171 specific calibrations are not well constrained and/or not available for halogens in most eclogitic  
172 phases, we used a basaltic glass calibration to present a unified dataset that is self-consistent.  
173 This may introduce crystal-specific matrix effects, although preliminary calibrations of F in  
174 pyroxene and glass (Kumamoto et al. 2017) show slopes indistinguishable within uncertainty.  
175 Continuous measurements of ALV519-4-1 were used throughout the session to monitor  
176 instrument drift, which was found to be negligible. The OH, F, and Cl concentrations of Suprasil  
177 and Herasil 102, both optical-quality glasses, along with Synthetic Forsterite, were measured  
178 regularly in each session. There are no published values for the OH, F, and Cl content of  
179 Suprasil, Herasil 102 and Synthetic Forsterite, but those samples are believed to have very low  
180 OH, F, and Cl concentrations (less than 0.1  $\mu\text{g/g}$ ), respectively (E. Hauri, *pers. comm.*). To  
181 quantify our maximum backgrounds, we assume that Suprasil contains no OH, Herasil contains  
182 no F, and Synthetic Forsterite contains no Cl. Background OH, F, and Cl values were less than  
183 7.3  $\mu\text{g/g}$ , 0.19  $\mu\text{g/g}$ , and 0.15  $\mu\text{g/g}$ , respectively, for all sessions. Individual mineral analyses

184 were not corrected for background when background values were less than propagated 2SE  
185 uncertainty; otherwise, measurements are corrected to remove backgrounds (Table 1). Cl  
186 measurement uncertainties on omphacite, garnet, and phengite are near parity with Synthetic  
187 Forsterite Cl values. We therefore subtracted our background Cl values from measured values for  
188 all bulk calculations.

189

## 190 **4.2 EPMA**

191 All mineral phases were analyzed using JEOL JXA-8200 electron microprobe (EMP) at  
192 the Massachusetts Institute of Technology. A 1nA beam current and 15kV accelerating potential  
193 were used for all analyses except apatite which utilized a 10 nA beam current. Beam diameter  
194 was <1 micron for garnet and omphacite, and 10  $\mu\text{m}$  for amphibole and phengite. Data reduction  
195 was done using CITZAF software (Armstrong 1995). The counting times used for phase  
196 analysis was 40s on peak, and 20s on background. With respect to halogen abundances in  
197 amphibole and phengite, detection limits based on counting statistics were ~470  $\mu\text{g/g}$  for F and  
198 ~60  $\mu\text{g/g}$  for Cl. For apatite, care was taken to address known issues pertaining to fluorine  
199 excitation by electron microprobe analysis (Stormer and Pierson 1993). Peak searches were  
200 conducted only on the standard, and background was only measured on the first point for each  
201 grain. In addition, the LDE1 (W/Si) crystal was used for F measurements, which suppresses  
202 elemental interferences. Counting times for F were abbreviated to 10s on peak and 40 s for P, Ca,  
203 and Cl while background counts were 5s for F and 20s for Ca, P, and Cl. Analytical uncertainties  
204 including background for F and Cl in apatite based on counting statistics were 3.4–4.4% and 7.2–  
205 11.6% ( $1\sigma$ ), respectively, at concentrations of 1.47–3.25 wt % F and 192–515  $\mu\text{g/g}$  Cl.

206

207 **4.3 Bulk halogen measurements**

208        Sample powders were washed in ultra-pure ( $18\text{M}\Omega$ ) deionized water and oven-dried  
209        overnight to avoid surficial Cl contamination by adsorption. This may have removed a portion of  
210        halogens along grain boundaries or in fluid inclusions; therefore we report Cl concentrations as  
211        minimums. We note that fluid inclusion dimensions ( $\sim 1\mu\text{m}$  to  $10\mu\text{m}$ , on average) are smaller  
212        than estimated powder grain size ( $<100\mu\text{m}$ ), and were likely retained, as demonstrated in a  
213        subsequent section. Halogens were extracted from approximately 2 grams of sample powder by  
214        pyrohydrolysis, in which a stream of water vapor captures volatiles as they are released from the  
215        sample during fusion by a gas torch. The vapor is condensed to produce an aqueous solution  
216        (Schnetger and Muramatsu 1996). From this solution, F and Cl concentrations are measured via  
217        ion chromatography (IC) using a Dionex Integrion HPIC System at the University of Texas at  
218        Austin. Detection limits of the analyzed solution are  $0.05\mu\text{g/g}$  for both F and Cl. The halogen  
219        content of the whole rock is then calculated using solution concentration, solution volume, and  
220        the mass of the powder melted. Using the JB-2 basalt reference material ( $98.5\mu\text{g/g}$  F and  $281\mu\text{g/g}$  Cl  
221        (Imai et al. 1995)), F and Cl yields are 75–83% and 68–95% ( $n = 5$ ), respectively.  
222        Samples were analyzed with triple replicates in two batches. Since JB-2 yields are constrained  
223        within each contemporaneous batch, concentrations were yield-corrected by batch to the JB-2  
224        reference material to more rigorously quantify absolute halogen concentrations. The yield  
225        correction assumes that standard yields and sample yields are similar. Although not perfect, this  
226        correction likely provides more accurate values. Both uncorrected and yield-corrected values are  
227        presented in this study for comparison. Reproducibility for each sample (precision) was

228 combined with batch-specific JB-2 yield uncertainty (accuracy) to derive  $1\sigma$  uncertainties.  
229 Replicate measurements of JB-2 and uncertainties are given in Table S9.

230

231 **4.4 Modal mineralogy**

232 We calculated the modal abundances of each phase by inverting the bulk-rock major  
233 element data of John et al. (2010) with EMP analyses of individual phases (Tables S1-S6).  
234 Although EMPA indicates zonation in some phases (e.g. TiO<sub>2</sub> and MnO), changes in modal  
235 proportions were negligible (often <1% difference) when varying mineral compositions were  
236 used. We performed our calculations by taking the mean composition for each phase. SiO<sub>2</sub>,  
237 TiO<sub>2</sub>, Al<sub>2</sub>O<sub>3</sub>, FeO, MnO, MgO, CaO, Na<sub>2</sub>O, K<sub>2</sub>O, and P<sub>2</sub>O<sub>5</sub> were used in the minimization. P<sub>2</sub>O<sub>5</sub>  
238 was assumed to reside entirely within apatite. Calculated modes agree well with independently  
239 estimated mineral modes based on optical microscopy. Calculated mineral modes and  
240 uncertainties are shown in Table 2.

241 All eclogite samples contain >35 modal % garnet. Omphacite modal proportions range  
242 from 23–53 vol.%, whereas phengite, when present, makes up 2–4 vol% of the bulk-rock  
243 modally. Samples contain varying modal proportions of amphibole, from 0.5 vol% to ~30 vol%.  
244 Quartz, when present, is up to 8 vol% of the bulk rock whereas rutile (1–2 vol%) and apatite  
245 (0.25–1.6 vol%) are omnipresent.

246

247 **5. Results**

248 **5.1 Major element variability by EMPA**

249 Major element profiles show garnet zonation with respect to end member components  
250 almandine, pyrope, grossular, and spessartine (Table S1). Garnet porphyroblast cores are often

251 elevated in the spessartine component as well as TiO<sub>2</sub>, with a gradual decrease towards the rim  
252 consistent with prograde garnet growth.

253 Omphacite (Cpx) Mg# (molar Mg/(Mg+Fe)) varies from 0.68 to 0.79 (Table S2) and  
254 correlates negatively with Cpx abundances, reflecting the influence of bulk-rock chemistry on  
255 mineral modes. Omphacite EMPA profiles show some intra-mineral zonation; in particular,  
256 SEC42-06 and SEC43-01 show increasing Si concentrations from core to rim, while SEC43-03  
257 Si content decreases from core to rim. Mg# remain uniform within single grains.

258 Phengite was analyzed in samples SEC42-06, SEC43-01, and SEC47-01 (Table S5).  
259 SEC42-06 is relatively homogeneous, with slight Na enrichment approaching the rim. A full rim-  
260 to-rim profile in SEC43-01 shows strong zonation with increases in Na and Al from core to rim,  
261 and decreases in Si, K, Fe and Mg (Fig. S3). SEC47-01 shows slight Si, K, and Mg rim  
262 depletions. Both phengite profiles are consistent with continued grain growth or partial  
263 recrystallization during decompression due to decreasing Si from core to rim (Massonne and  
264 Schreyer 1987).

265 Amphibole occurs as several different species which likely reflect bulk compositional  
266 variations; analyses are presented in Table S4, along with IMA 2012 recommended  
267 nomenclature calculated using the routine of Locock (2014). SEC42-06, SEC43-01, and SEC43-  
268 03 are Na-Ca subgroup katophorite, SEC46-01 and SEC46-02 are Ca subgroup pargasite, and  
269 SEC47-01 is classified as a Na-Ca subgroup winchite. Nearly all EMP analyses show Cl  
270 abundances at or below the detection limit of ~60 µg/g. However, F concentrations show  
271 significant intra-grain variability, differing by as much as a factor of two from core to rim (e.g.  
272 SEC42-06, 869–1834 µg/g).

273 Apatite halogen contents range from 1.47–3.25% wt% F and 0.019–0.051 wt% Cl (Table  
274 S5). The hydroxyl content of the apatite are taken as the oxide sum subtracted from one hundred,  
275 which classifies apatite in samples SEC43-03, SEC46-01, SEC46-02, and SEC47-01 as  
276 hydroxylapatite and apatite in sample SEC43-01 and SEC42-06 as fluoroapatite. Apatite halogen  
277 abundances are similar to apatite from other blueschist and eclogite metamorphic terranes with  
278 low Cl abundances (Pagé et al. 2016; Hughes et al. 2018) whereas some HP terranes developed  
279 from oceanic protoliths also contain Cl-rich apatite (John and Schenk 2003). Typical MORB  
280 (0.164 wt% P<sub>2</sub>O<sub>5</sub>, Gale et al. (2013)) must undergo significant fractional crystallization (>70%)  
281 before reaching apatite saturation (~0.7 wt% P<sub>2</sub>O<sub>5</sub>); indeed, apatite-saturated MORB are few  
282 (Anderson and Greenland 1969; Watson 1979). Based on the bulk rock chemistry of Raspas  
283 provided by John et al. (2010) these conditions are unlikely to have been met. Therefore  
284 interstitial grains most likely formed as new growth during progressive metamorphism.

285

## 286 **5.2. Halogen and water variability**

287 Bulk rock pyrohydrolysis analyses are presented in Table 3. As noted previously, we  
288 report both yield-corrected and uncorrected halogen abundances. Yield-corrected F and Cl  
289 concentrations for all samples were 115–199 µg/g and 6.7–19.9 µg/g, respectively. In situ  
290 halogen measurements are presented in Tables 1 and S5. Resolvable intra- and inter-grain  
291 heterogeneity is observed for F and OH in all phases both in single-grain SIMS profiles (Table  
292 S7) and in multiple grains from the same sample.

293 Halogen variability is notable in nominally anhydrous minerals (NAMs), particularly  
294 garnet with both inter- and intra-grain heterogeneity observed ([Figure S4](#), Tables 1 and S7).  
295 Grain-averaged sample values range from 1.7 to 8.9 µg/g F and 60 to 201 µg/g OH, while Cl is

296 always at or near detection limits. Intra-sample garnet F and OH variability is up to 93% and  
297 67%, respectively ( $1\sigma$ , Table 1). Variations from core to rim were not systematic, nor can they be  
298 attributed to analytical uncertainty; in some cases cores were enriched in F or OH relative to  
299 rims, while in other grains the case was reversed. Core and rim measurements for OH and F are  
300 presented in Fig. 2, illustrating that while some intragrain variability is present, most grains plot  
301 on or near 1:1 lines within uncertainty.

302 Omphacite has highly variable F abundances (6.5–54.1  $\mu\text{g/g}$ ) while OH range from 445  
303 to 773  $\mu\text{g/g}$ . Similar to garnet, Cl abundances are less than 0.1  $\mu\text{g/g}$  for five of the six samples.  
304 However, SEC46-02 yields significantly higher mean Cl abundances of 1.05  $\mu\text{g/g}$ . F and OH  
305 variability within a single sample was less than 24% and 14%, respectively. A profile of Cpx  
306 from SEC43-01 exemplifies core to rim variations, with F and OH concentrations of 42.1–69.7  
307  $\mu\text{g/g}$  and 362–557  $\mu\text{g/g}$ , respectively (Fig. S5). While these variations are discernable,  
308 heterogeneity in this grain is 15% or less for all three species; core to rim variations are modest  
309 and typically plot along a 1:1 line, within uncertainty (Fig. 2).

310 Amphibole occurs as both interstitial grains and monomineralic veinlets, with veins only  
311 occurring in SEC46-02. Interstitial amphibole yields F and Cl abundances of 747–4727  $\mu\text{g/g}$  and  
312 0.56–82.6  $\mu\text{g/g}$ , respectively. Pargasite veins from SEC46-02 contain on average 563  $\mu\text{g/g}$  F and  
313 11.7  $\mu\text{g/g}$  Cl. Single sample averaging produces variability of no more than 18% for F and 37%  
314 for Cl ( $1\sigma$ ).

315 Phengite from three samples contains 610–1822  $\mu\text{g/g}$  F and 1.01–1.98  $\mu\text{g/g}$  Cl. Two of  
316 the three grains analyzed show no systematic core to rim halogen variations. Phengite is  
317 relatively homogeneous in terms of F and Cl, with  $1\sigma$  sample variations of less than 20%.  
318 However, in a rim to rim transect of SEC43-01 phengite (Fig. S6), intra-grain variability is

319 evident. F concentrations show a systematic decrease from core to rim, paralleling major element  
320 trends of decreasing K<sub>2</sub>O, SiO<sub>2</sub>, and MgO. Cl is enriched near grain boundaries by  
321 approximately 40% with respect to cores, following Al<sub>2</sub>O<sub>3</sub> and Na<sub>2</sub>O rim enrichments ([Fig. S3](#)  
322 [and S6](#)). Finally, a single secondary titanite in SEC46-02 contains 1394 µg/g F, <0.1 µg/g Cl,  
323 and 4010 µg/g OH.

324

## 325 **6. Discussion**

### 326 **6.1. Halogen redistribution during eclogitization**

327 Halogen distribution during prograde metamorphism depends to a large extent on bulk-  
328 rock chemistry, which dictates the equilibrium mineral assemblage and phase proportions at any  
329 given pressure and temperature. Raspas eclogites show a strong positive correlation between the  
330 proportion of omphacite in the bulk-rock and the F content of omphacite ([Fig. S7](#)), suggesting  
331 that F is repartitioned during eclogitization and progressive omphacite growth. In addition, we  
332 compare the modal proportion of nominally anhydrous minerals (omphacite + garnet + quartz) to  
333 the volatile content of analyzed phases ([Fig. S8](#)). As the proportion of NAMs in the bulk  
334 assemblage increases, the F content of each phase shows a commensurate increase. Indeed,  
335 modal abundances of amphibole are negatively correlated with F concentrations in both  
336 omphacite and apatite, again suggesting effective redistribution of F during prograde  
337 metamorphism. Further, apatite Cl abundances increase as amphibole modes decrease and the  
338 mineral assemblage becomes dominated by garnet and omphacite. In all cases concerning lattice-  
339 hosted halogens, F is accommodated primarily in apatite (1.47–3.25 wt %), amphibole (563–  
340 4727 µg/g), and phengite (610–1822 µg/g) with lesser amounts in omphacite (6.5–54.1 µg/g) and  
341 garnet (1.7–8.9 µg/g). In addition, Cl is primarily hosted in apatite (192–515 µg/g), with lesser

342 amounts in amphibole (0.64–82.7 µg/g), phengite (1.2–2.1 µg/g), omphacite (<0.05–1 µg/g) and  
343 garnet (<0.05 µg/g).

344 Our in situ measurements permit the calculation of distribution coefficients between  
345 various phases of interest ([Figs. 3 and 4](#), Table S8). Given the scarcity of data, these new values  
346 provide important constraints on the distribution of F and Cl in eclogitic phases. A number of  
347 observations can made from the above distribution coefficients. First, F partitioning between  
348 mineral pairs defines a narrow range and appears to be insensitive to concentrations, which in  
349 some cases span nearly an order of magnitude. Second, as expected apatite will preferentially  
350 host crystal lattice-bound F and Cl when present in eclogitic assemblages, similar to apatite's  
351 halogen affinity in the presence of aqueous fluids (Kusebauch et al. 2015). Finally, F and Cl are  
352 preferentially partitioned into amphibole over phengite in all samples by a factor of 2.4 and 20,  
353 respectively. Interestingly, this observation is in contrast to blueschists studied by Pagé et al.  
354 (2016) who found  $D_F^{Amp\text{-}Phen}$  was always less than unity whereas  $D_{Cl}^{Amp\text{-}Phen}$  was  
355 approximately equal to one within error. Hence, blueschist to eclogite phase transformations may  
356 elicit changes in the relative compatibilities of halogens between coexisting phases, likely as a  
357 function of changing modal proportions, mineral compositions (e.g. glaucophane to katophorite),  
358 and P-T conditions.

359

## 360 **6.2. Comparison between measured and reconstructed bulk halogen abundances**

361 Using mineral modes, in situ SIMS data, and EPMA halogen concentrations, we calculate  
362 a reconstructed bulk-rock F and Cl content for each eclogite sample, which can then be  
363 compared to bulk values measured by pyrohydrolysis (Table 3, [Fig. 5](#)). Both raw (not corrected

364 for yield) and yield-corrected pyrohydrolysis halogen abundances are given in Table 3. We use  
365 yield-corrected concentrations in subsequent discussions.

366 Reconstructed concentrations and pyrohydrolysis measurements yield F concentrations of  
367 168–415 µg/g and 115–199 µg/g, respectively. Reconstructed F abundances are broadly  
368 concordant with measured values; three of the six samples agree within uncertainty, whereas  
369 three samples display higher reconstructed values (Fig. 6; Table 3). The relative congruence  
370 between reconstructed and measured F values implies that F may be entirely hosted within the  
371 crystal lattice of eclogitic phase assemblages. The discrepancy observed for three of the six  
372 samples may be due to a combination of factors including uncertainties in pyrohydrolysis yields  
373 (Table S9) that are typically larger than SIMS analyses, inter-grain halogen heterogeneity, or  
374 uncertainties in the calculated modal proportions (Table 2). We note that calculated modal  
375 proportions as derived from whole rock analysis were assessed, but found to be minor.

376 Bulk reconstructed and measured Cl concentrations for all samples are 0.8–8.2 µg/g and  
377 6.7–19.9 µg/g, respectively. Unlike F, Cl shows very large and systematic discrepancies between  
378 reconstructed and measured values, with reconstructed values representing only 12–41% of the  
379 pyrohydrolysis-derived bulk values. We discuss these drastic differences below, in terms of  
380 uncertainties and possible means of reconciliation between the two methods. We consider three  
381 possible processes to explain the difference: retrograde alteration, grain boundary hosted Cl, and  
382 Cl-bearing fluid inclusions.

383 Bulk halogen analyses rely on the assumption that samples are perfectly preserved and  
384 have not been retrogressed along their exhumation path. Even minute amounts of alteration (e.g.  
385 chlorite or titanite) can drastically affect such bulk measurements (Debret et al. 2016). Most  
386 Raspas eclogites appear to be pristine with little to no visible alteration or retrogression, with the

exception of SEC46-02 that contains titanite-on-rutile overgrowths and monomineralic amphibole veinlets adjacent to FI-rich omphacite, indicating possible fluid assimilation. This sample also displays elevated reconstructed (8.2  $\mu\text{g/g}$ ) and bulk (19.9  $\mu\text{g/g}$ ) Cl compared to the rest of the suite. In light of this suspected fluid assimilation, and the large F incongruity (~100%) between the two methods, we exclude this sample from subsequent bulk rock calculations. Minor retrogression may have contributed to the discrepancy for SEC46-02, however all other samples still show at least a factor of three difference between reconstructed and measured bulk Cl concentrations. Therefore, a factor other than late alteration is likely at play.

To further explore this discrepancy, we conducted additional SIMS analyses on grain boundaries of FI-free amphibole and garnet (Fig. S9). Both profiles across grain boundaries display pronounced spikes in  $^{35}\text{Cl}/^{30}\text{Si}$  signals within the immediate grain boundary at near-constant  $^{30}\text{Si}$  counts. These grain boundary measurements are equivalent to 7–18  $\mu\text{g/g}$  Cl. Yet, the grain boundaries themselves are far smaller than the dimensions of our analysis spot (3.8x3.8  $\mu\text{m}$ ) and depending on their size, could contain > 1 wt% Cl assuming a grain boundary thickness of 5 nm or less. As grain boundary porosity values in eclogites are poorly constrained, we used equation 9.6 of Turcotte and Schubert (2002) to calculate grain boundary porosity values based on a simplified grain geometry over a range of grain boundary thicknesses (1–25 nm) with realistic grain sizes (500  $\mu\text{m}$ ) to yield values of  $9 \times 10^{-12}$  to  $6 \times 10^{-9}$  (Table S10). The Cl contribution from grain boundaries alone (<0.1  $\mu\text{g/g}$  Cl in all cases, or <2% contribution compared to pyrohydrolysis Cl abundances), given the above parameters, is unable to reconcile the reconstructed versus measured Cl abundances.

Finally, we consider the possible role of fluid inclusions, as they could represent an important host for Cl (Philippot and Selverstone 1991; Herms et al. 2012; Kendrick 2018). FI

410 (~1  $\mu\text{m}$  to 10  $\mu\text{m}$  on average) are found ubiquitously hosted in garnet (**Fig. S1C**) and omphacite  
411 cores, both of which appear in textural equilibrium with adjacent phases and occur in all  
412 samples. FI are far more abundant in SEC46-02, in particular where omphacite is in contact with  
413 monomineralic pargasitic amphibole veinlets (**Fig. 1B**). FI appear cloudy in thin section owing to  
414 their prolific abundance. These features likely reflect crystallization in the presence of fluids,  
415 possibly during prograde conditions (e.g. John et al (2010) and Herms et al. (2012)); such  
416 features have also been noted elsewhere in previous studies (e.g. Massonne et al. (2012), Zack et  
417 al. (2001)). Although we cannot directly analyze fluid inclusions by SIMS, heterogeneous Cl  
418 enrichments in SEC46-02 FI-rich omphacite (44–94  $\mu\text{g/g}$  Cl) indicate that each SIMS analysis  
419 may incorporate some Cl-rich FI material. FI-free (optically clear) omphacite from the same  
420 sample also contains an order of magnitude more Cl (1.05  $\mu\text{g/g}$ ) compared to omphacite from all  
421 other samples (<0.1  $\mu\text{g/g}$ ), which is consistent with assimilation of pervasive fluids during  
422 crystallization (Herms et al. 2012). FI may host significant quantities of Cl from hundreds of  
423  $\mu\text{g/g}$  to >10 wt% NaCl equivalent (Philippot and Selverstone 1991; Philippot et al. 1998;  
424 Scambelluri and Philippot 2001; Svensen et al. 2001; Kendrick et al. 2015; Kendrick 2018).  
425 Herms et al. (2012) have shown that omphacite and garnet hosted FIs from Raspas contain ~2  
426 wt% NaCl, congruent with the salinity of dehydration-related FI found in rocks that have  
427 undergone blueschist-to-eclogite conversion (Gao and Klemd 2001). Assuming that FI contain  
428 ~2 wt% NaCl and represent 0.05–0.1% of the total sample volume, one may reconcile the  
429 discrepancy between in situ and bulk Cl abundances. Thus, pyrohydrolysis appears to  
430 incorporate fluid inclusion contributions, whereas SIMS analyses do not. Our calculations imply  
431 that FI host the majority (at least 80%) of Cl in Raspas eclogites; whether these fluid inclusions

432 are all primary in nature, and therefore should be considered in the halogen budget of recycled  
433 eclogites, remains to be determined.

434 Combining SIMS and pyrohydrolysis techniques affords unique insights on the  
435 distribution and abundance of halogens in subducted slabs. SIMS provides detailed constraints  
436 on crystal lattice hosted halogen abundances, and is well suited to determine bulk F  
437 concentrations, given the uncertainty that can be associated with pyrohydrolysis yields.  
438 Minimally retrogressed Raspas eclogites analyzed by SIMS contain between 145 and 258 µg/g  
439 F; F may be entirely hosted within the crystal lattice of eclogites, whereas only ~20% of Cl (at  
440 most), on average, is crystallographically hosted when FI are present. However, SIMS cannot  
441 accurately account for Cl abundances if primary fluid inclusions are present in the sample, which  
442 dominate the Cl budget of eclogites. Assuming that fluid inclusions observed here are primary  
443 features, pyrohydrolysis is the preferred method to measure the bulk Cl content of pristine FI-  
444 bearing rocks; minimally retrogressed Raspas eclogites analyzed by pyrohydrolysis contain  
445 between 6.7 and 11.3 µg/g Cl.

446

## 447 7. Implications

448 Previous studies have suggested that amphibole is the primary carrier of halogens to the  
449 upper mantle (Debret et al. 2016; Barnes et al. 2018). Our results suggest a more nuanced view  
450 when one takes into account the contributions from coexisting phases and FI; amphibole's  
451 contribution to mineral-hosted bulk F and Cl abundances range from 6–84% and 4–85%,  
452 respectively (Fig. 6, S7). When amphibole modes are below ~5% (three of six samples), apatite  
453 is the most important carrier for mineral-hosted halogens, containing >80% of mineral-hosted Cl  
454 and 38–71% of mineral-hosted F. Apatite is thought to be stable to 7.5 GPa at 950° (Konzett and

455 Frost 2009), and will act as an important halogen carrier in eclogitized ocean crust. The  
456 remaining F is distributed between amphibole (4–53%), phengite (6–32%), omphacite (4–16%),  
457 and garnet (up to 2.2%). The above calculations consider only crystal matrix hosted halogen  
458 abundances; our results further indicate that FI host the majority of Cl in the bulk rock (at least  
459 80%, on average). The ultimate fate of Cl will be dictated by the redistribution of free fluid  
460 phases during subsequent phase transformations, diffusive re-equilibration, and dynamic  
461 recrystallization at higher pressures and temperatures.

462 Raspas reconstructed bulk F concentrations (145–258 µg/g) are similar to previous pre-  
463 subduction AOC estimates of 216 µg/g (Straub and Layne 2003), blueschist ( $457 \pm 159$  µg/g F)  
464 and eclogite-facies ( $283 \pm 146$  µg/g F) AOC (Pagé et al. 2016; Hughes et al. 2018), as well as  
465 mean N–MORB values of  $183 \pm 107$  µg/g calculated from recent work from Le Voyer et al.  
466 (2018) (Fig. 7A). Therefore, F concentrations remain relatively unchanged during the  
467 progression from nascent MORB, to AOC, and finally eclogitized AOC. Our results indicate that  
468 F is efficiently returned (~95%, derived from SIMS) to the upper mantle during subduction  
469 (Straub and Layne 2003; John et al. 2011; Pagé et al. 2016). Conversely, bulk pyrohydrolysis Cl  
470 measurements (6.7–11.3 µg/g) are at least an order of magnitude lower than pre-subduction AOC  
471 estimates (e.g. 207 µg/g (Barnes and Cisneros 2012)). Yet, observed Raspas concentrations are  
472 quite similar to recent blueschist and eclogite-facies AOC estimates by Hughes et al. (2018) of  
473 3–23 µg/g Cl and Pagé et al. (2018) of 12 µg/g Cl, and well within the range observed in MORB  
474 (excluding Cl/K >0.1 after Shimizu and Saal (2016) and references therein) (Fig. 7A).  
475 Progression from initial MORB generation to seafloor alteration leads to an increase in Cl  
476 content, while subsequent eclogitization leads to a decrease in Cl abundances, returning to values  
477 similar to MORB. Using the pre-subduction AOC Cl bulk-rock estimate of Barnes and Cisneros

478 (2012) of 207 µg/g Cl and data presented here, we calculate that the vast majority (>95%,  
479 derived from bulk pyrohydrolysis) of subducted Cl is expelled from the slab by the time the  
480 AOC parcel equilibrates at eclogite-facies conditions (Table 3), in agreement with previous  
481 workers (Straub and Layne 2003; John et al. 2011; Kendrick et al. 2014; Pagé et al. 2016). What  
482 little Cl remains will reside primarily in FI (at least 80%, on average). Thus, while most F is  
483 retained in the subducting slab at eclogite-facies conditions, most Cl previously added through  
484 hydrothermal alteration is expelled, fractionating F from Cl in eclogitized AOC. Yet, halogen  
485 abundances in Raspas are similar to unaltered MORB (Fig. 7A), while ratios of F and Cl with  
486 elements of similar compatibility during melting, Cl/K ( $0.0066 \pm 0.0043$ ), Cl/Nb ( $3.9 \pm 1.0$ ) and  
487 F/Nd ( $15.5 \pm 4.8$ ), are indistinguishable within uncertainty from DMM ratios of Salters and  
488 Stracke (2004) (Fig. 7B).

489 The HIMU (high U/Pb) mantle source is thought to contain ancient altered oceanic crust  
490 based on Pb isotope systematics (and more recently Tl, see Shu et al. (2019)), yielding  
491 enrichments in F/Nd,  $^{35}\text{Cl}/^{37}\text{Cl}$ , and often Cl/K and Cl/Nb ratios (Chase 1981; Hofmann and  
492 White 1982; Zindler et al. 1982; Stroncik and Haase 2004; John et al. 2011; Cabral et al. 2014;  
493 Jackson et al. 2015; Le Voyer et al. 2015; Kendrick et al. 2017; Rose-Koga et al. 2017; Hanyu et  
494 al. 2019). Our data indicate that halogen enrichments, and elevated F/Nd and Cl/K ratios, in the  
495 HIMU source cannot be produced by simple addition of recycled eclogitized AOC (Fig. 7B), but  
496 rather require delivery by an additional lithology, e.g. serpentinite as proposed by Kendrick et al.  
497 (2017), or an efficient fractionation mechanism to elevate halogen ratios during subsequent  
498 processing (e.g. Niu and O'Hara (2003)). The broad spectrum of F/Nd and Cl/K ratios observed  
499 in HIMU melts may then reflect variable contributions from subducted serpentinite (green line,  
500 Fig. 7B). However, antigorite serpentinite breakdown by the reaction: antigorite = olivine +

501 orthopyroxene + H<sub>2</sub>O requires geochemical transfer of halogens, likely into a fluid phase as both  
502 F and Cl are highly incompatible in olivine and orthopyroxene (Dalou et al. 2014; Joachim et al.  
503 2015; Beyer et al. 2016); their ultimate host within the upper mantle remains unknown. Direct  
504 phase transformation from antigorite serpentinite to the dense hydrous magnesium silicate Phase  
505 A may occur along cold geothermal gradients (Schmidt and Poli 1998), or halogens may become  
506 more compatible in K-rich clinopyroxene at increasing pressures as K solubility in clinopyroxene  
507 increases to ~1 wt. % K<sub>2</sub>O at 100 kbar (Schmidt and Poli 1998). Thus, subduction proceeds in  
508 returning to the upper mantle an eclogitized AOC parcel with halogen ratios quite similar to the  
509 MORB from which it was derived.

510

## 511 **Figures**

512 See attached folder

## 513 **Figure Captions**

514 Figure 1. Representative photomicrographs of Raspas eclogite textures. A. SEC43-03 contains  
515 garnet porphyroblasts within a matrix of omphacite and lesser quartz. B. Amphibole veinlet in  
516 sample SEC46-02. Adjacent omphacite shows abundant fluid inclusions (yellow arrow). C.  
517 SEC43-03 Megacrystic, inclusion-rich garnet showing sector zoning.

518

519 Figure 2. Core and rim measurements for OH (A. and B.) and F (C. and D.) in garnet and  
520 omphacite. The majority of samples show core and rim values that fall on a 1:1 line, within  
521 uncertainty, allowing a high degree of confidence in bulk rock calculations based on in situ  
522 measurements. Error bars are 2SE.

523

524 Figure 3. Inter-mineral F partitioning between A. omphacite and garnet, B. omphacite and  
525 apatite, C. omphacite and amphibole, and D. omphacite and phengite. White triangles are  
526 literature data from Debret et al. (2016). Solid lines show constant inter-mineral partition  
527 coefficients (labeled), while mean partition coefficients for each mineral pair are presented in the  
528 bottom right panel corner.

529

530 Figure 4. Inter-mineral Cl partitioning between A. phengite and apatite, B. apatite and  
531 amphibole, C. phengite and amphibole, and D. apatite Cl and garnet F concentrations. White  
532 triangles are literature data from Debret (2016), green diamonds from Hughes et al. (2018) and

533 blue upside down triangles from Pagé et al. (2016). Solid lines denote constant inter-mineral  
534 partition coefficients as in Fig. 3.

535

536 Figure 5. Bulk reconstructed and bulk measured (yield corrected) halogen abundances for F (A.)  
537 and Cl (B.). Y error bars show analytical reproducibility propagated with yield precision. Yield-  
538 uncorrected pyrohydrolysis values are shown in panels C and D, where positive Y error bars  
539 incorporate yield uncertainty. Red dots denote SEC46-02, which was not included in bulk  
540 calculations; see text for more information. Solid lines represent 1:1 lines; dashed lines plot  
541 various slopes, labeled accordingly.

542

543 Figure 6. Pie diagrams show modal abundances (A. and B.), as well as the percent contribution  
544 from each phase to mineral-hosted halogen abundances as derived from SIMS analyses for F (C.  
545 and D.) and Cl (E. and F.) in two representative, yet mineralogically distinct, samples.

546

547 Figure 7. A. F and Cl abundances from this study in comparison to literature values for MORB  
548 glasses from Le Voyer et al. (2018), HIMU OIB (Cabral et al. 2014; Jackson et al. 2015;  
549 Kendrick et al. 2017), blueschist and eclogite (Debret et al. 2016; Pagé et al. 2016; Hughes et al.  
550 2018), antigorite serpentinite (Pagé et al. 2018), and AOC estimates (Straub and Layne 2003;  
551 Van den Bleeken and Koga 2015; Barnes et al. 2018; Kendrick 2019a). MORB with Cl/K values  
552 in excess of 0.1 were screened out, due to possible seawater contamination after Shimizu and  
553 Saal (2016) and references therein. Peridotite mantle plots off the scale and would be 1.4–31  
554 µg/g for F and from 0.14 to 0.38 µg/g for Cl (Urann et al. 2017). B. F/Nd and Cl/K ratios of  
555 Raspas samples (Nd and K from John et al. (2010)), MORB glasses from Le Voyer (2018), mean  
556 antigorite serpentinite halogen concentrations from Page et al. (2018) with DMM K (60 µg/g)  
557 and Nd (0.713 µg/g) values of Salters and Stracke (2004), and OIB MIs and glasses from the  
558 literature (Cabral et al. 2014; Jackson et al. 2015; Kendrick et al. 2017). We note that our  
559 eclogitized AOC samples indeed plot below the canonical F/Nd of 21, as posited by Rose-Koga  
560 et al. (2012) for deeply subducted oceanic crust.

561 Figure S1. Representative full-section photomicrographs of Raspas samples analyzed in this  
562 study. A. Prograde assemblage in SEC43-01 contains garnet porphyroblasts within a matrix of  
563 omphacite and lesser quartz and phengite with minor apatite and rutile. The sample shows no  
564 evidence for retrogression. B. SEC43-03 megacrystic garnet showing sector zoning with minor  
565 quartz. C. Amphibole veinlet in sample SEC46-02, outlined in white dashed line. Adjacent  
566 omphacite shows abundant fluid inclusions (dark patches).

567

568 Figure S2. Representative calibration slope used in this study. Calibration slope confidence  
569 intervals were calculated with a non-linear bootstrapping technique using 5,000 iterations.

570 Figure S3. SEC43-01 phengite EMP profile from rim to rim shows Na<sub>2</sub>O and Al<sub>2</sub>O<sub>3</sub> increases  
571 from core to rim, while K<sub>2</sub>O, SiO<sub>2</sub>, and MgO all decrease toward the grain boundary.

572

573 Figure S4. SEC43-01 garnet core to rim transect. Zero values represent rim, 750 µm spot  
574 location is the garnet core. Overall, OH and F show intragrain variability likely related to  
575 partitioning and re-equilibration during garnet growth. The combination of internal and  
576 propagated calibration slope uncertainty corresponds to the positive error bar on our analyses.

577 The negative error bars are larger for Cl because they incorporate the conservative uncertainty on  
578 the maximum background.

579

580 Figure S5. SEC43-01 omphacite core (~1300  $\mu\text{m}$ ) to rim (zero  $\mu\text{m}$ ) profile for OH, F, and Cl.  
581 Significant variations occur on length scales of <100  $\mu\text{m}$ , with apparent rim depletions in F and  
582 OH.

583

584 Figure S6. SEC43-01 phengite rim-to-rim transect. F shows depletions toward the grain  
585 boundaries, while Cl shows rim enrichments of approximately 40%.

586

587 Figure S7. Fluorine concentrations (black y-axis, yellow dots) and contributions to the bulk-rock  
588 (blue y-axis, blue dots) plotted against modal proportion of each phase for A. omphacite, B.  
589 amphibole, C. apatite, and D. phengite.

590

591 Figure S8. Fluorine and chlorine concentrations in each phase as a function of the modal  
592 proportion of nominally anhydrous minerals (omphacite plus garnet). As the modal proportion of  
593 NAMs increases, so does the halogen concentration in most phases.

594

595 Figure S9. Grain boundary transects using SIMS across amphibole (A.) and garnet (B.). While  
596 there are likely unaccounted for matrix effects with this approach, both measurements on the  
597 grain boundary qualitatively show pronounced spikes in Cl/Si ratios at near-constant constant Si  
598 counts, indicating Cl enrichment within the grain boundary (C. and D.). Also shown are raw  $^{30}\text{Si}$   
599 counts for comparison (panels E and F).

600

601

## 602 Acknowledgments

603 We thank two anonymous reviewers for their insights and suggestions, which improved the  
604 manuscript. This research was supported by NSF award EAR-P&G #1524311 and #1839128 to  
605 VLR. BMU acknowledges the support of the WHOI Ocean Venture Fund. This work was also  
606 supported by NSF-PIRE grant (OIA- 1545903) to JDB.

607

## 608 References Cited

- 609 Anderson, A.T., and Greenland, L.P. (1969) Phosphorus fractionation diagram as a quantitative  
610 indicator of crystallization differentiation of basaltic liquids. *Geochimica et Cosmochimica  
611 Acta*, 33, 493–505.
- 612 Armstrong, J.T. (1995) CITZAF—a package of correction programs for the quantitative electron  
613 microbeam X-ray analysis of thick polished materials, thin lms, and particles. *Microbeam  
614 Anal*, 177–200.
- 615 Barnes, J.D., and Cisneros, M. (2012) Mineralogical control on the chlorine isotope composition  
616 of altered oceanic crust. *Chemical Geology*, 326-327, 51–60.

- 617 Barnes, J.D., Manning, C.E., Scambelluri, M., and Selverstone, J. (2018) The Behavior of  
618 Halogens During Subduction-Zone Processes. In *The Role of Halogens in Terrestrial and*  
619 *Extraterrestrial Geochemical Processes* Vol. 263, pp. 545–590. Springer International  
620 Publishing, Cham.
- 621 Beinlich, A., Klemd, R., John, T., and Gao, J. (2010) Trace-element mobilization during Ca-  
622 metasomatism along a major fluid conduit: Eclogitization of blueschist as a consequence of  
623 fluid–rock interaction. *Geochimica et Cosmochimica Acta*, 74, 1892–1922.
- 624 Beyer, C., Klemme, S., Grützner, T., Ireland, T.R., Magee, C.W., and Frost, D.J. (2016) Fluorine  
625 partitioning between eclogitic garnet, clinopyroxene, and melt at upper mantle conditions.  
626 *Chemical Geology*, 437, 88–97.
- 627 Cabral, R.A., Jackson, M.G., Koga, K.T., Rose-Koga, E.F., Hauri, E.H., Whitehouse, M.J., Price,  
628 A.A., Day, J.M.D., Shimizu, N., and Kelley, K.A. (2014) Volatile cycling of H<sub>2</sub>O, CO<sub>2</sub>, F,  
629 and Cl in the HIMU mantle: A new window provided by melt inclusions from oceanic hot  
630 spot lavas at Mangaia, Cook Islands. *Geochemistry, Geophysics, Geosystems*, 15, 4445–  
631 4467.
- 632 Chase, C.G. (1981) Oceanic island Pb: Two-stage histories and mantle evolution. *Earth and*  
633 *Planetary Science Letters*, 52, 277–284.
- 634 Chavrit, D., Burgess, R., Sumino, H., Teagle, D.A.H., Droop, G., Shimizu, A., and Ballentine,  
635 C.J. (2016) The contribution of hydrothermally altered ocean crust to the mantle halogen and  
636 noble gas cycles. *Geochimica et Cosmochimica Acta*, 183, 106–124.
- 637 Dalou, C., Koga, K.T., Le Voyer, M., and Shimizu, N. (2014) Contrasting partition behavior of F  
638 and Cl during hydrous mantle melting: implications for Cl/F signature in arc magmas.  
639 *Progress in Earth and Planetary Science*, 1, 26.
- 640 Debret, B., Koga, K.T., Cattani, F., Nicollet, C., Van den Bleeken, G., and Schwartz, S. (2016)  
641 Volatile (Li, B, F and Cl) mobility during amphibole breakdown in subduction zones.  
642 *LITHOS*, 244, 165–181.
- 643 Feininger, T. (1980) Eclogite and Related High-Pressure Regional Metamorphic Rocks from the  
644 Andes of Ecuador. *Journal of Petrology*, 21, 107–140.
- 645 Feininger, T. (1987) Allochthonous terranes in the Andes of Ecuador and northwestern Peru.  
646 *Canadian Journal of Earth Sciences*, 24, 266–278.
- 647 Freundt, A., Grevemeyer, I., Rabbel, W., Hansteen, T.H., Hensen, C., Wehrmann, H., Kutterolf,  
648 S., Halama, R., and Frische, M. (2014) Volatile (H<sub>2</sub>O, CO<sub>2</sub>, Cl, S) budget of the Central  
649 American subduction zone. *International Journal of Earth Sciences*, 103, 2101–2127.
- 650 Gabriele, P. (2002) HP terranes exhumation in an active margin setting: geology, petrology and  
651 geochemistry of the Raspas Complex in SW Ecuador.
- 652 Gale, A., Dalton, C.A., Langmuir, C.H., Su, Y., and Schilling, J.G. (2013) The mean

- 653 composition of ocean ridge basalts. *Geochemistry, Geophysics, Geosystems*, 14, 489–518.
- 654 Gao, J., and Klemd, R. (2001) Primary fluids entrapped at blueschist to eclogite transition:  
655 evidence from the Tianshan meta-subduction complex in northwestern China. *Contributions  
656 to Mineralogy and Petrology*, 142, 1–14.
- 657 Halama, R., John, T., Herms, P., Hauff, F., and Schenk, V. (2011) A stable (Li, O) and  
658 radiogenic (Sr, Nd) isotope perspective on metasomatic processes in a subducting slab.  
659 *Chemical Geology*, 281, 151–166.
- 660 Hanyu, T., Shimizu, K., Ushikubo, T., Kimura, J.-I., Chang, Q., Hamada, M., Ito, M., Iwamori,  
661 H., and Ishikawa, T. (2019) Tiny droplets of ocean island basalts unveil Earth's deep  
662 chlorine cycle. *Nature Communications*, 10, 222.
- 663 Herms, P., John, T., Bakker, R.J., and Schenk, V. (2012) Evidence for channelized external fluid  
664 flow and element transfer in subducting slabs (Raspas Complex, Ecuador). *Chemical  
665 Geology*, 310–311, 79–96.
- 666 Hofmann, A.W., and White, W.M. (1982) Mantle plumes from ancient oceanic crust. *Earth and  
667 Planetary Science Letters*, 57, 421–436.
- 668 Hughes, L., Burgess, R., Chavrit, D., Pawley, A., Tartèse, R., Droop, G., Ballentine, C.J., and  
669 Lyon, I. (2018) Halogen behaviour in subduction zones: Eclogite facies rocks from the  
670 Western and Central Alps. *Geochimica et Cosmochimica Acta*, 243, 1–23.
- 671 Imai, N., Terashima, S., Itoh, S., and Ando, A. (1995) 1994 Complilation of Analytical Data for  
672 Minor and Trace Elements in Seventeen GSJ Geochemical Reference Samples, “Igneous  
673 Rock Series.” *Geostandards and Geoanalytical Research*, 19, 135–213.
- 674 Ito, E., Harris, D.M., and Anderson, A.T. (1983) Alteration of oceanic crust and geologic cycling  
675 of chlorine and water. *Geochimica et Cosmochimica Acta*, 47, 1613–1624.
- 676 Jackson, M.G., Koga, K.T., Price, A., Konter, J.G., Koppers, A.A.P., Finlayson, V.A., Konrad,  
677 K., Hauri, E.H., Kylander-Clark, A., Kelley, K.A., and others (2015) Deeply dredged  
678 submarine HIMU glasses from the Tuvalu Islands, Polynesia: implications for volatile  
679 budgets of recycled oceanic crust. Wiley Online Library
- 680 .
- 681 Joachim, B., Pawley, A., Lyon, I.C., néé Hartmann, K.M., Henkel, T., Clay, P.L., Ruzié, L.,  
682 Burgess, R., and Ballentine, C.J. (2015) Experimental partitioning of F and Cl between  
683 olivine, orthopyroxene and silicate melt at Earth's mantle conditions. *Chemical Geology*,  
684 416, 65–78.
- 685 John, T., and Schenk, V. (2003) Partial eclogitisation of gabbroic rocks in a late Precambrian  
686 subduction zone (Zambia): prograde metamorphism triggered by fluid infiltration.  
687 *Contributions to Mineralogy and Petrology*, 146, 174–191.

- 688 John, T., Scambelluri, M., Frische, M., Barnes, J.D., and Bach, W. (2011) Dehydration of  
689 subducting serpentinite: Implications for halogen mobility in subduction zones and the deep  
690 halogen cycle. *Earth and Planetary Science Letters*, 308, 65–76.
- 691 John, T., Scherer, E.E., Schenk, V., Herms, P., Halama, R., and Garbe-Schönberg, D. (2010)  
692 Subducted seamounts in an eclogite-facies ophiolite sequence: the Andean Raspas Complex,  
693 SW Ecuador. *Contributions to Mineralogy and Petrology*, 159, 265–284.
- 694 Kendrick, M.A. (2018) Halogens in Seawater, Marine Sediments and the Altered Oceanic  
695 Lithosphere. In *The Role of Halogens in Terrestrial and Extraterrestrial Geochemical*  
696 *Processes* Vol. 87, pp. 591–648. Springer, Cham, Cham.
- 697 Kendrick, M.A. (2019a) Halogens in altered ocean crust from the East Pacific Rise (ODP/IODP  
698 Hole 1256D). *Geochimica et Cosmochimica Acta*, 261, 93–112.
- 699 Kendrick, M.A. (2019b) Halogens in Atlantis Bank gabbros, SW Indian Ridge: Implications for  
700 styles of seafloor alteration. *Earth and Planetary Science Letters*, 514, 96–107.
- 701 Kendrick, M.A., Hémond, C., Kamenetsky, V.S., Danyushevsky, L., Devey, C.W., Rodemann,  
702 T., Jackson, M.G., and Perfit, M.R. (2017) Seawater cycled throughout Earth's mantle in  
703 partially serpentinized lithosphere. *Nature Geoscience*, 10, 222–228.
- 704 Kendrick, M.A., Honda, M., and Vanko, D.A. (2015) Halogens and noble gases in  
705 Mathematician Ridge meta-gabbros, NE Pacific: implications for oceanic hydrothermal root  
706 zones and global volatile cycles. *Contributions to Mineralogy and Petrology*, 170, 173.
- 707 Kendrick, M.A., Jackson, M.G., Kent, A.J.R., Hauri, E.H., Wallace, P.J., and Woodhead, J.  
708 (2014) Contrasting behaviours of CO<sub>2</sub>, S, H<sub>2</sub>O and halogens (F, Cl, Br, and I) in enriched-  
709 mantle melts from Pitcairn and Society seamounts. *Chemical Geology*, 370, 69–81.
- 710 Konzett, J., and Frost, D.J. (2009) The High P–T Stability of Hydroxyl-apatite in Natural and  
711 Simplified MORB—an Experimental Study to 15 GPa with Implications for Transport and  
712 Storage of Phosphorus and Halogens in Subduction Zones. *Journal of Petrology*, 50, 2043–  
713 2062.
- 714 Kumamoto, K.M., Warren, J.M., and Hauri, E.H. (2017) New SIMS reference materials for  
715 measuring water in upper mantle minerals. *American Mineralogist*, 102, 537–547.
- 716 Kusebauch, C., John, T., Whitehouse, M.J., and Engvik, A.K. (2015) Apatite as probe for the  
717 halogen composition of metamorphic fluids (Bamble Sector, SE Norway). *Contributions to*  
718 *Mineralogy and Petrology*, 170, 271.
- 719 Le Voyer, M., Cottrell, E., Kelley, K.A., Bounce, M., and Hauri, E.H. (2015) The effect of  
720 primary versus secondary processes on the volatile content of MORB glasses: An example  
721 from the equatorial Mid-Atlantic Ridge (5°N–3°S). *Journal of Geophysical Research: Solid*  
722 *Earth*, 120, 125–144.
- 723 Le Voyer, M., Hauri, E.H., Cottrell, E., Kelley, K.A., Salters, V.J.M., Langmuir, C.H., Hilton,

- 724 D.R., Barry, P.H., and Füri, E. (2018) Carbon fluxes and primary magma CO<sub>2</sub>contents  
725 along the global mid-ocean ridge system. *Geochemistry, Geophysics, Geosystems*.
- 726 Locock, A.J. (2014) An Excel spreadsheet to classify chemical analyses of amphiboles following  
727 the IMA 2012 recommendations. *Computers & Geosciences*, 62, 1–11.
- 728 Marschall, H.R., Altherr, R., Gmélinc, K., and Kasztovszky, Z. (2009) Lithium, boron and  
729 chlorine as tracers for metasomatism in high-pressure metamorphic rocks: a case study from  
730 Syros (Greece). *Mineralogy and Petrology*, 95, 291–302.
- 731 Maruyama, S., Liou, J.G., and Terabayashi, M. (1996) Blueschists and Eclogites of the World  
732 and Their Exhumation. *International Geology Review*, 38, 485–594.
- 733 Massonne, H.J. (2012) Formation of amphibole and clinozoisite–epidote in eclogite owing to  
734 fluid infiltration during exhumation in a subduction channel. *Journal of Petrology*, 53, 1969–  
735 1998.
- 736 Massonne, H.J., and Schreyer, W. (1987) Phengite geobarometry based on the limiting  
737 assemblage with K-feldspar, phlogopite, and quartz. *Contributions to Mineralogy and*  
738 *Petrology*, 96, 212–224.
- 739 Mather, T.A., Pyle, D.M., Tsanev, V.I., McGonigle, A.J.S., Oppenheimer, C., and Allen, A.G.  
740 (2006) A reassessment of current volcanic emissions from the Central American arc with  
741 specific examples from Nicaragua. *Journal of Volcanology* ..., 149, 297–311.
- 742 Niu, Y., and O’Hara, M.J. (2003) Origin of ocean island basalts: A new perspective from  
743 petrology, geochemistry, and mineral physics considerations. *Journal of Geophysical*  
744 *Research: Solid Earth*, 108, 63.
- 745 Pagé, L., Hattori, K., and Guillot, S. (2018) Mantle wedge serpentinites: A transient reservoir of  
746 halogens, boron, and nitrogen for the deeper mantle. *Geology*, 46, 883–886.
- 747 Pagé, L., Hattori, K., de Hoog, J.C.M., and Okay, A.I. (2016) Halogen (F, Cl, Br, I) behaviour in  
748 subducting slabs: A study of lawsonite blueschists in western Turkey. *Earth and Planetary*  
749 *Science Letters*, 442, 133–142.
- 750 Philippot, P., Agrinier, P., and Scambelluri, M. (1998) Chlorine cycling during subduction of  
751 altered oceanic crust. *Earth and Planetary Science Letters*, 161, 33–44.
- 752 Philippot, P., and Silverstone, J. (1991) Trace-element-rich brines in eclogitic veins:  
753 implications for fluid composition and transport during subduction. *Contributions to*  
754 *Mineralogy and Petrology*, 106, 417–430.
- 755 Pyle, D.M., and Mather, T.A. (2009) Halogens in igneous processes and their fluxes to the  
756 atmosphere and oceans from volcanic activity: A review. *Chemical Geology*, 263, 110–121.
- 757 Rebay, G., Powell, R., and Diener, J.F.A. (2010) Calculated phase equilibria for a morb  
758 composition in a P-T range, 450–650 °C and 18–28 kbar: the stability of eclogite. *Journal*

- 759 of Metamorphic Geology, 28, 635–645.
- 760 Rose-Koga, E.F., Koga, K.T., Moreira, M., Vlastelic, I., Jackson, M.G., Whitehouse, M.J.,  
761 Shimizu, N., and Habib, N. (2017) Geochemical systematics of Pb isotopes, fluorine, and  
762 sulfur in melt inclusions from São Miguel, Azores. Chemical Geology.
- 763 Rose-Koga, E.F., Koga, K.T., Schiano, P., Le Voyer, M., Shimizu, N., Whitehouse, M.J., and  
764 Clochchiatti, R. (2012) Mantle source heterogeneity for South Tyrrenian magmas revealed  
765 by Pb isotopes and halogen contents of olivine-hosted melt inclusions. Chemical Geology,  
766 334, 266–279.
- 767 Sadofsky, S.J., Portnyagin, M., Hoernle, K., and van den Bogaard, P. (2008) Subduction cycling  
768 of volatiles and trace elements through the Central American volcanic arc: evidence from  
769 melt inclusions. Contributions to Mineralogy and Petrology, 155, 433–456.
- 770 Salters, V.J.M., and Stracke, A. (2004) Composition of the depleted mantle. Geochemistry,  
771 Geophysics, Geosystems, 5, n/a–n/a.
- 772 Scambelluri, M., and Philippot, P. (2001) Deep fluids in subduction zones. LITHOS, 55, 213–  
773 227.
- 774 Schmidt, M.W., and Poli, S. (1998) Experimentally based water budgets for dehydrating slabs  
775 and consequences for arc magma generation. Earth and Planetary Science Letters, 163, 361–  
776 379.
- 777 Schnetger, B., and Muramatsu, Y. (1996) Determination of halogens, with special reference to  
778 iodine, in geological and biological samples using pyrohydrolysis for preparation and  
779 inductively coupled plasma mass spectrometry and ion chromatography for measurement.  
780 The Analyst, 121, 1627.
- 781 Shimizu, K., Saal, A.E., Myers, C.E., Nagle, A.N., Hauri, E.H., Forsyth, D.W., Kamenetsky,  
782 V.S., and Niu, Y. (2016) Two-component mantle melting-mixing model for the generation  
783 of mid-ocean ridge basalts: Implications for the volatile content of the Pacific upper mantle.  
784 Geochimica et Cosmochimica Acta, 176, 44–80.
- 785 Shu, Y., Nielsen, S.G., Marschall, H.R., John, T., Blusztajn, J., and Auro, M. (2019) Closing the  
786 loop: Subducted eclogites match thallium isotope compositions of ocean island basalts.  
787 Geochimica et Cosmochimica Acta, 250, 130–148.
- 788 Stormer, J.C., and Pierson, M. (1993) Variation of F and Cl X-ray intensity due to anisotropic  
789 diffusion in apatite during electron microprobe analysis. [pubs.geoscienceworld.org](http://pubs.geoscienceworld.org)
- 790 .
- 791 Straub, S.M., and Layne, G.D. (2003) The systematics of chlorine, fluorine, and water in Izu arc  
792 front volcanic rocks: Implications for volatile recycling in subduction zones. Geochimica et  
793 Cosmochimica Acta, 67, 4179–4203.

- 794 Stroncik, N.A., and Haase, K.M. (2004) Chlorine in oceanic intraplate basalts: Constraints on  
795 mantle sources and recycling processes. *Geology*, 32, 945.
- 796 Svensen, H., Jamtveit, B., Banks, D.A., and Austrheim, H. (2001) Halogen contents of eclogite  
797 facies fluid inclusions and minerals: Caledonides, western Norway. *Journal of Metamorphic  
798 Geology*, 19, 165–178.
- 799 Turcotte, D.L., and Schubert, G. (2002) *Geodynamics* by Donald L. Turcotte, 2nd ed. Cambridge  
800 University Press, Cambridge.
- 801 Urann, B.M., Le Roux, V., Hammond, K., Marschall, H.R., Lee, C.T.A., and Monteleone, B.D.  
802 (2017) Fluorine and chlorine in mantle minerals and the halogen budget of the Earth's  
803 mantle. *Contributions to Mineralogy and Petrology*, 172, 324.
- 804 Van den Bleeken, G., and Koga, K.T. (2015) Experimentally determined distribution of fluorine  
805 and chlorine upon hydrous slab melting, and implications for F–Cl cycling through  
806 subduction zones. *Geochimica et Cosmochimica Acta*, 171, 353–373.
- 807 Watson, E. (1979) Apatite saturation in basic to intermediate magmas. Wiley Online Library  
808 .
- 809 Zack, T., Rivers, T., and Foley, S. (2001) Cs–Rb–Ba systematics in phengite and amphibole: an  
810 assessment of fluid mobility at 2.0 GPa in eclogites from Trescolmen, Central Alps.  
811 *Contributions to Mineralogy and Petrology*, 140, 651–669.
- 812 Zindler, A., Jagoutz, E., and Goldstein, S. (1982) Nd, Sr and Pb isotopic systematics in a three-  
813 component mantle: a new perspective. *Nature*, 298, 519–523.
- 814
- 815

Figure 1.

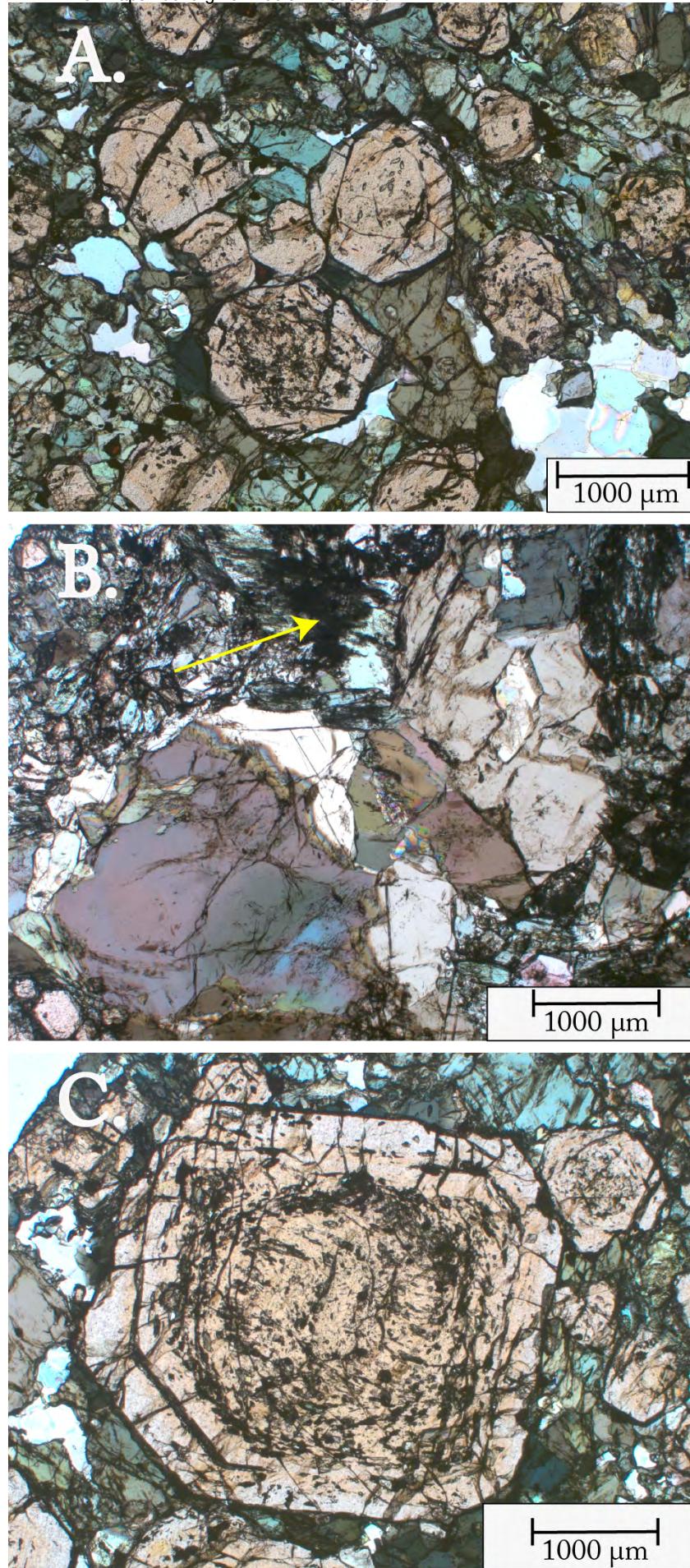


Figure 2.

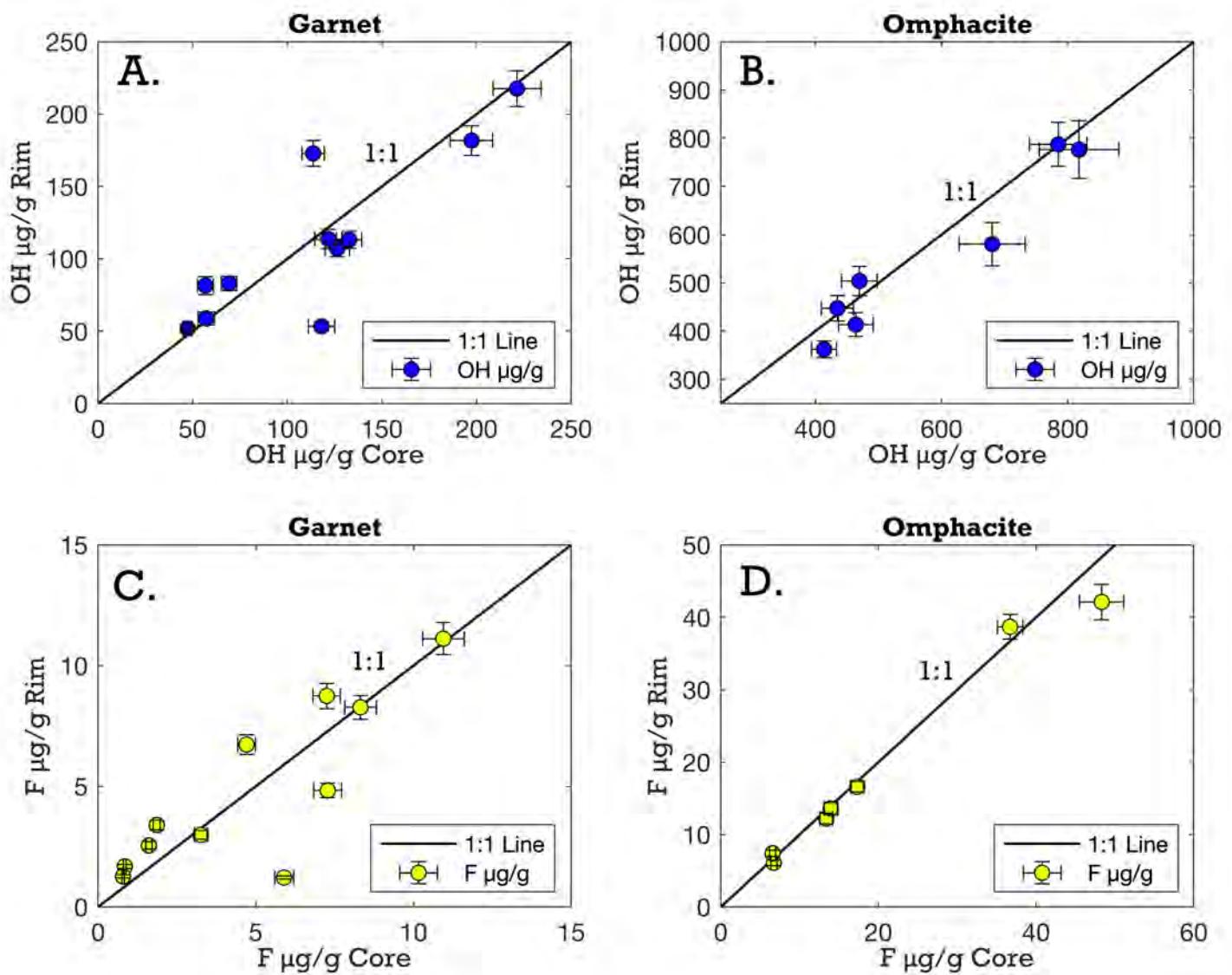


Figure 3.

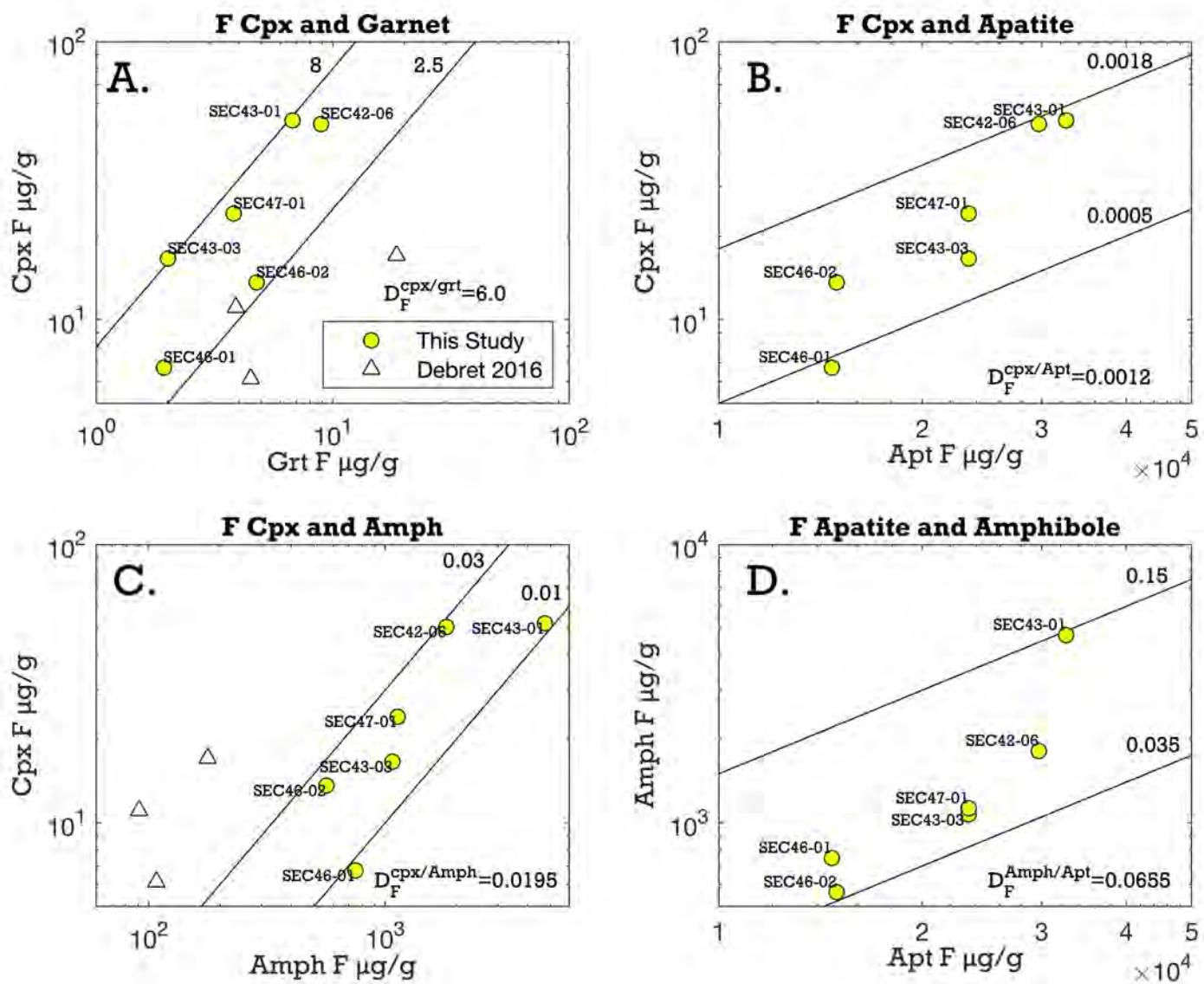


Figure 4.

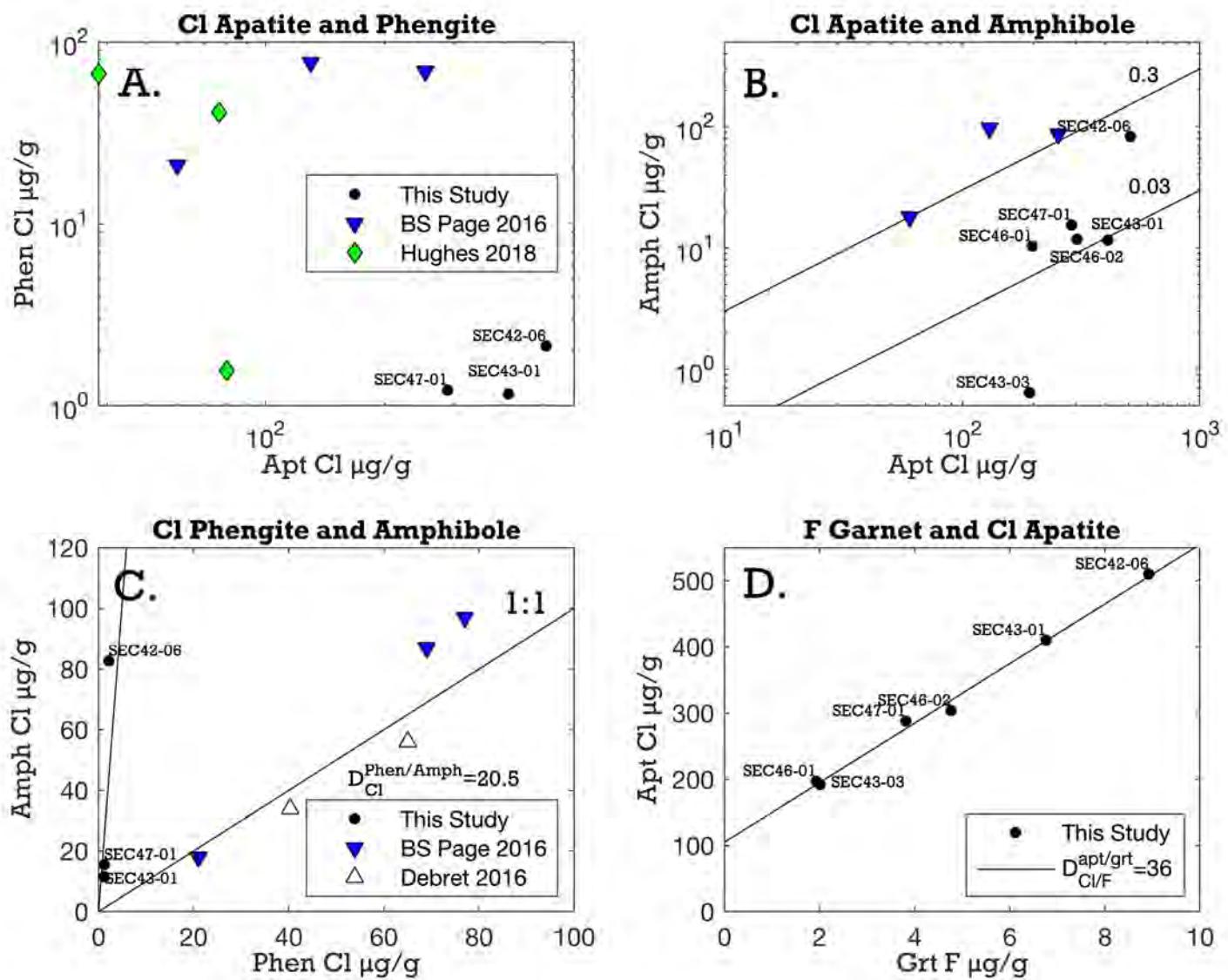


Figure 5.

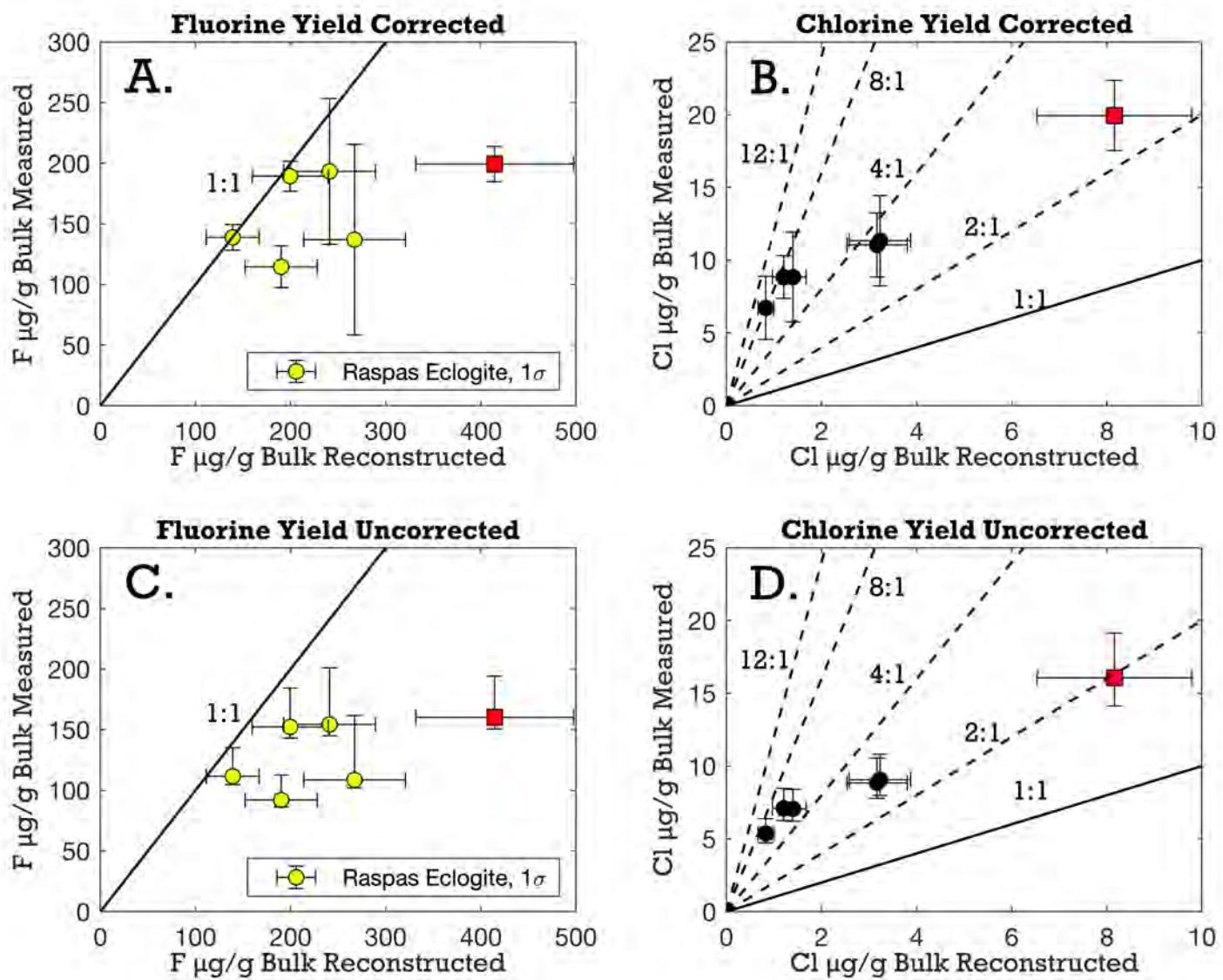


Figure 6.

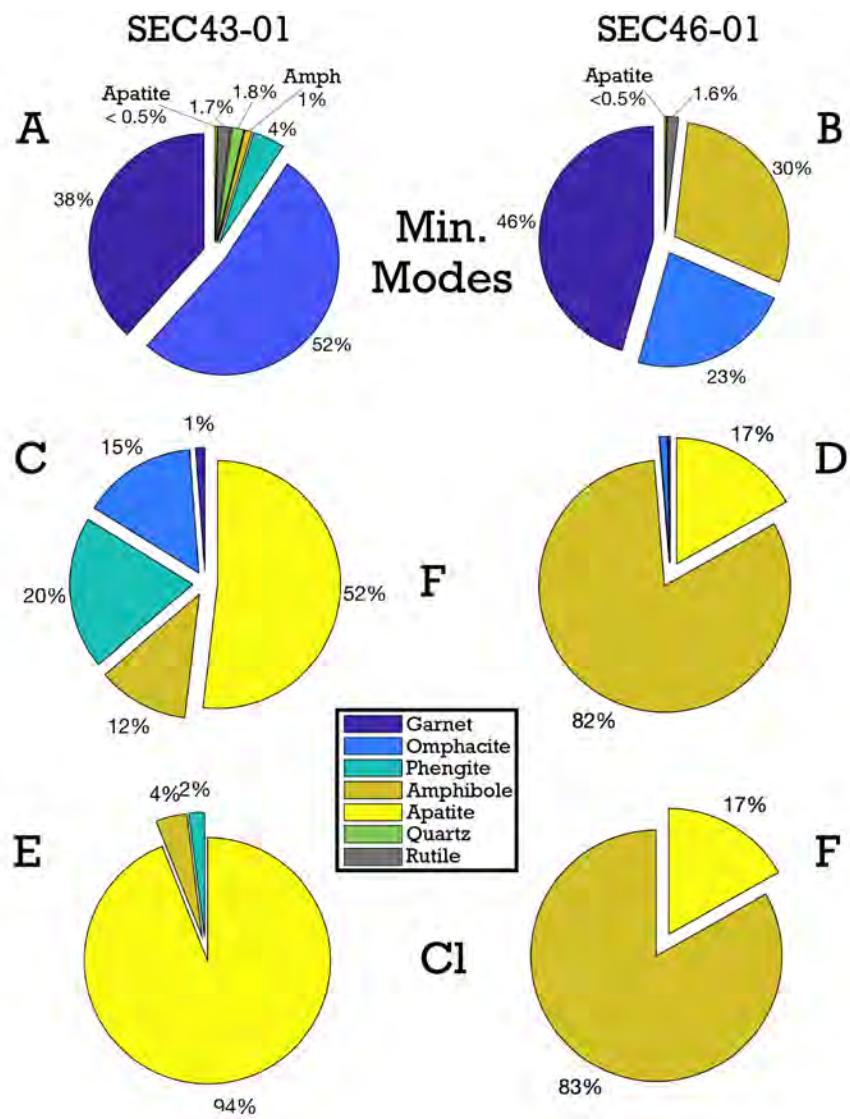


Figure 7.

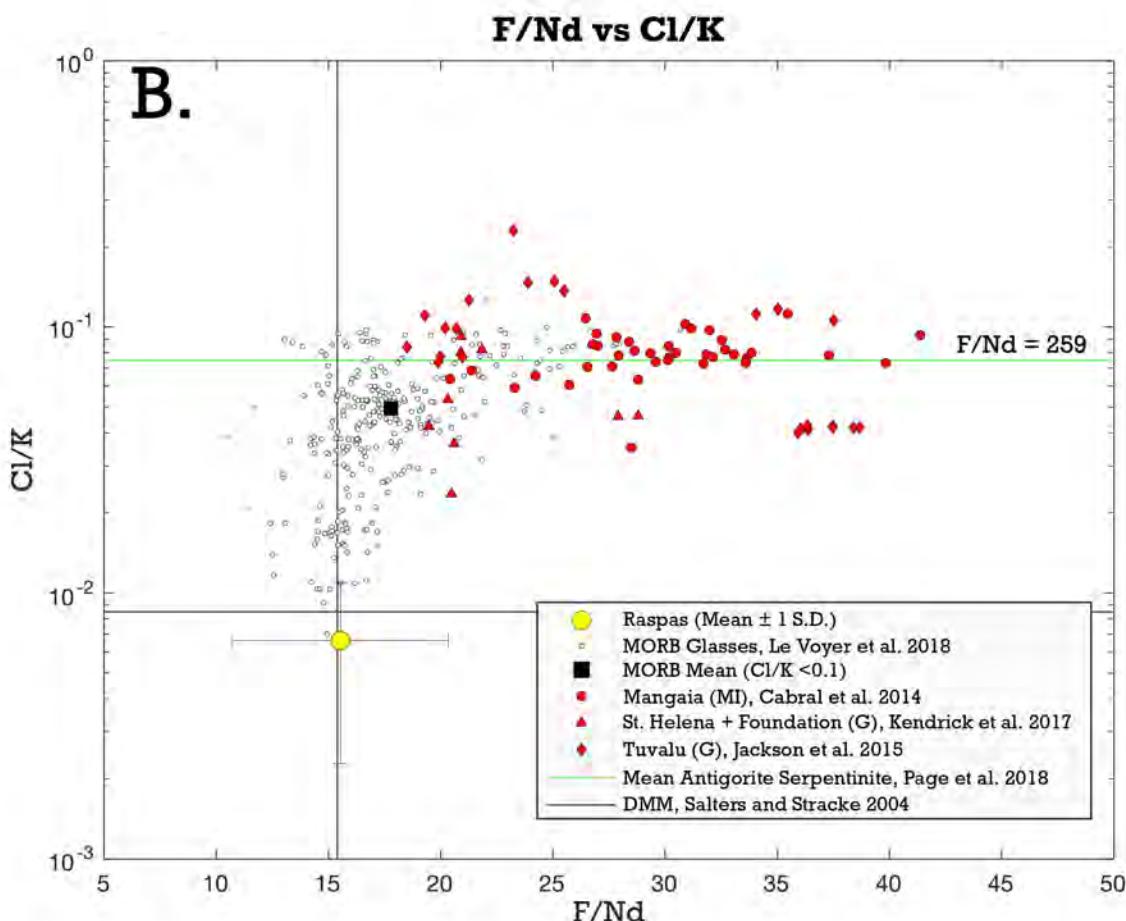
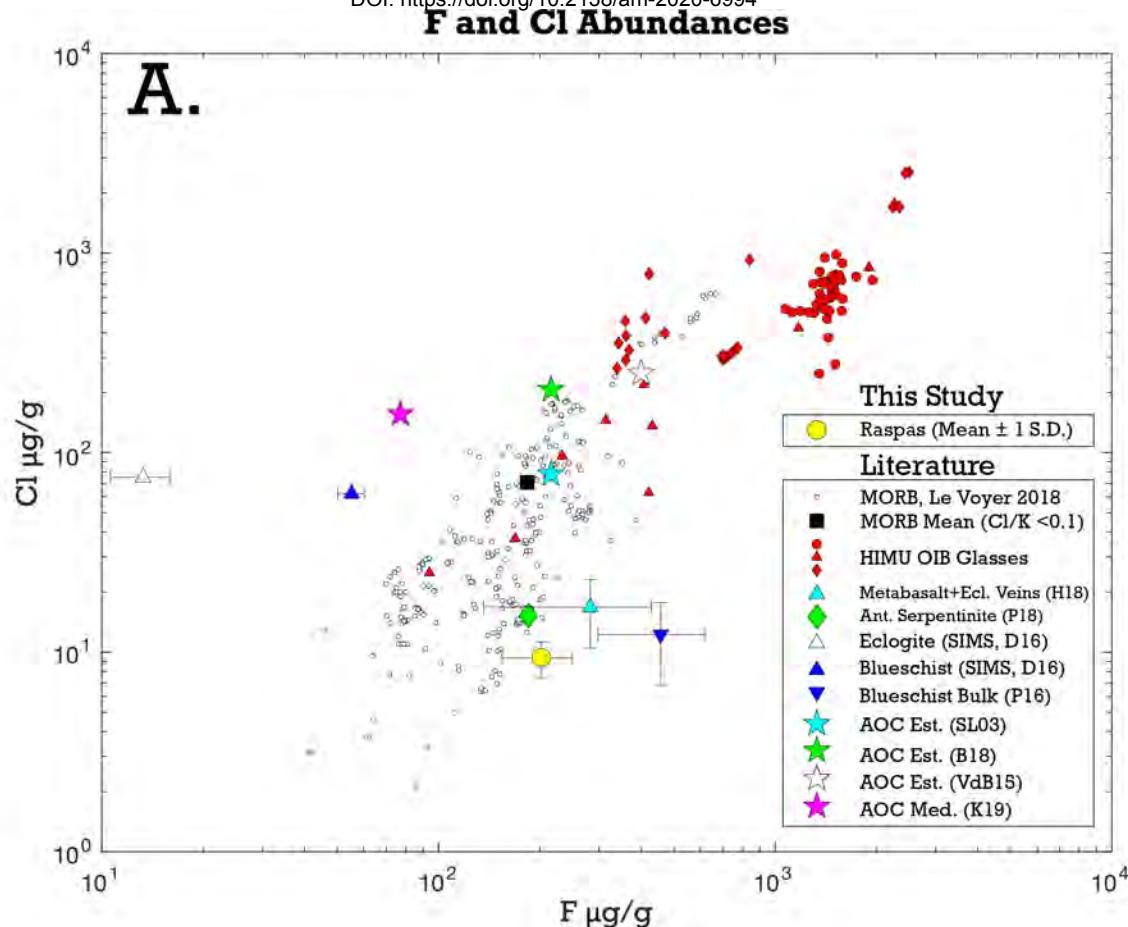


Figure S1.

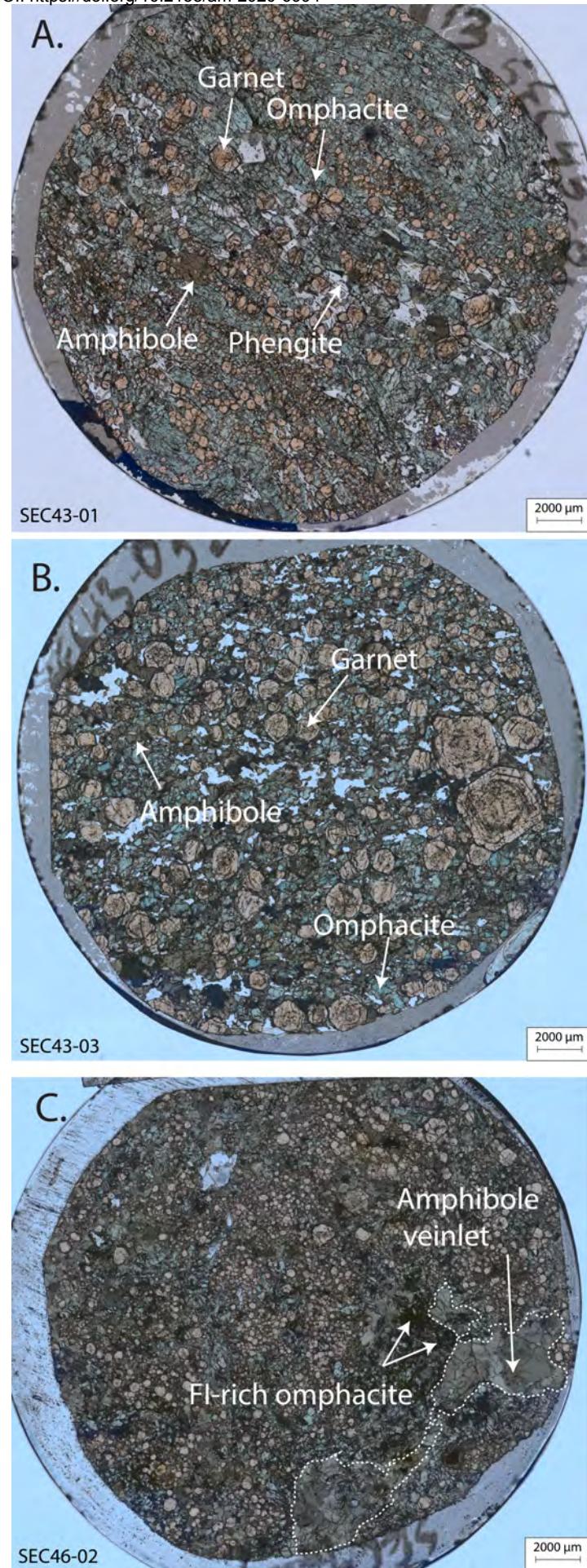


Figure S2.

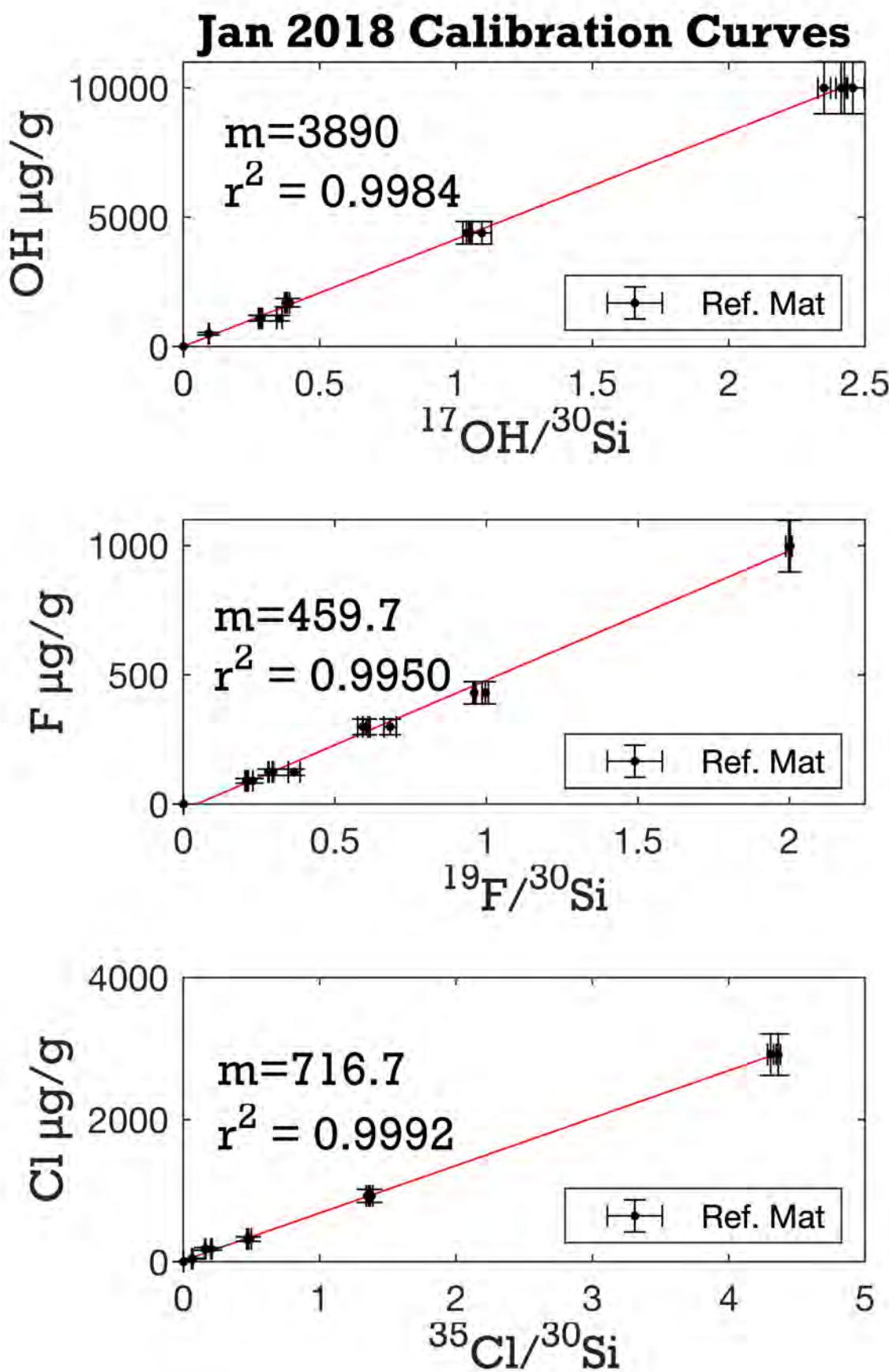


Figure S3.

## SEC43-01 Phengite Major Element Zonation

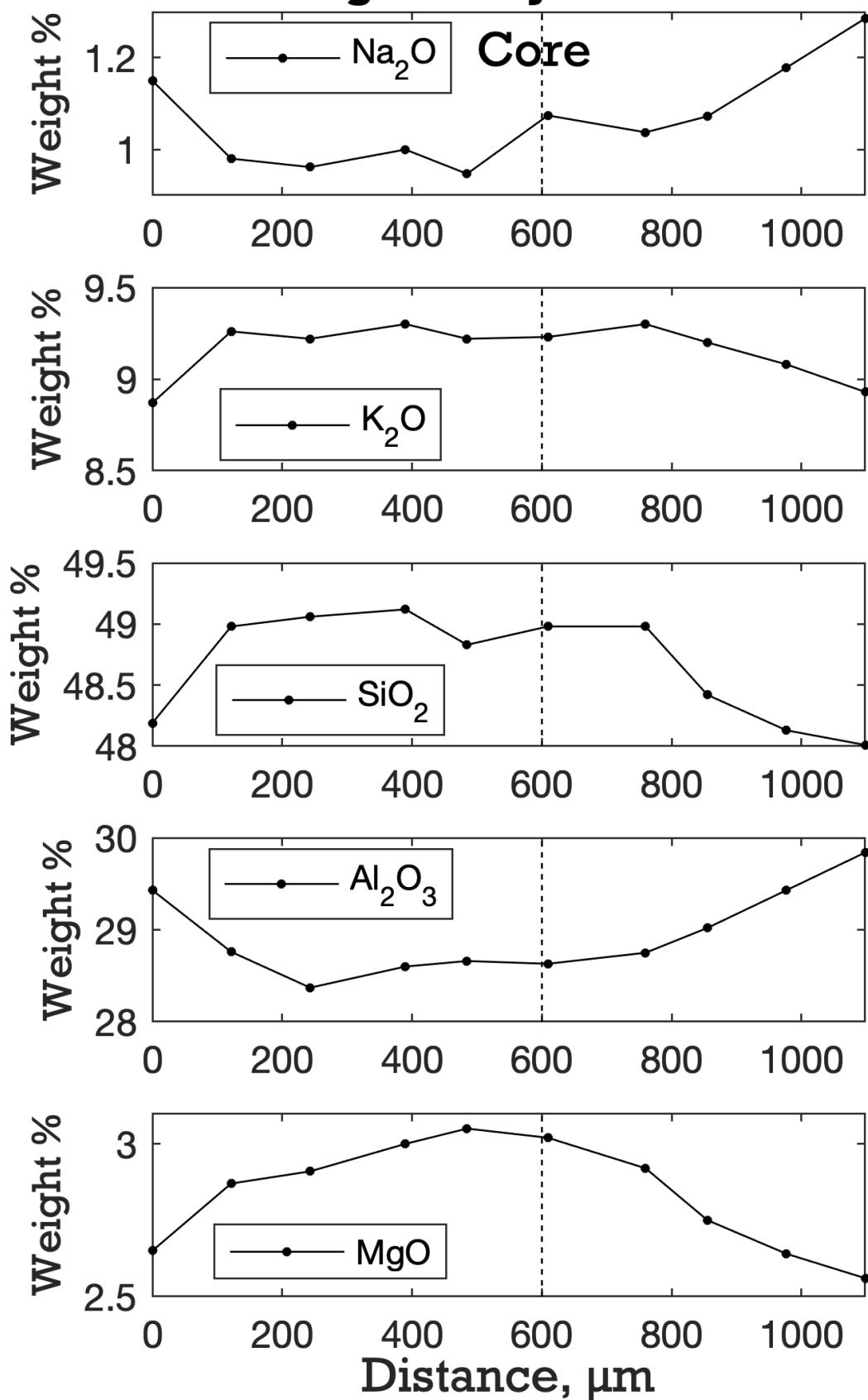


Figure S4.

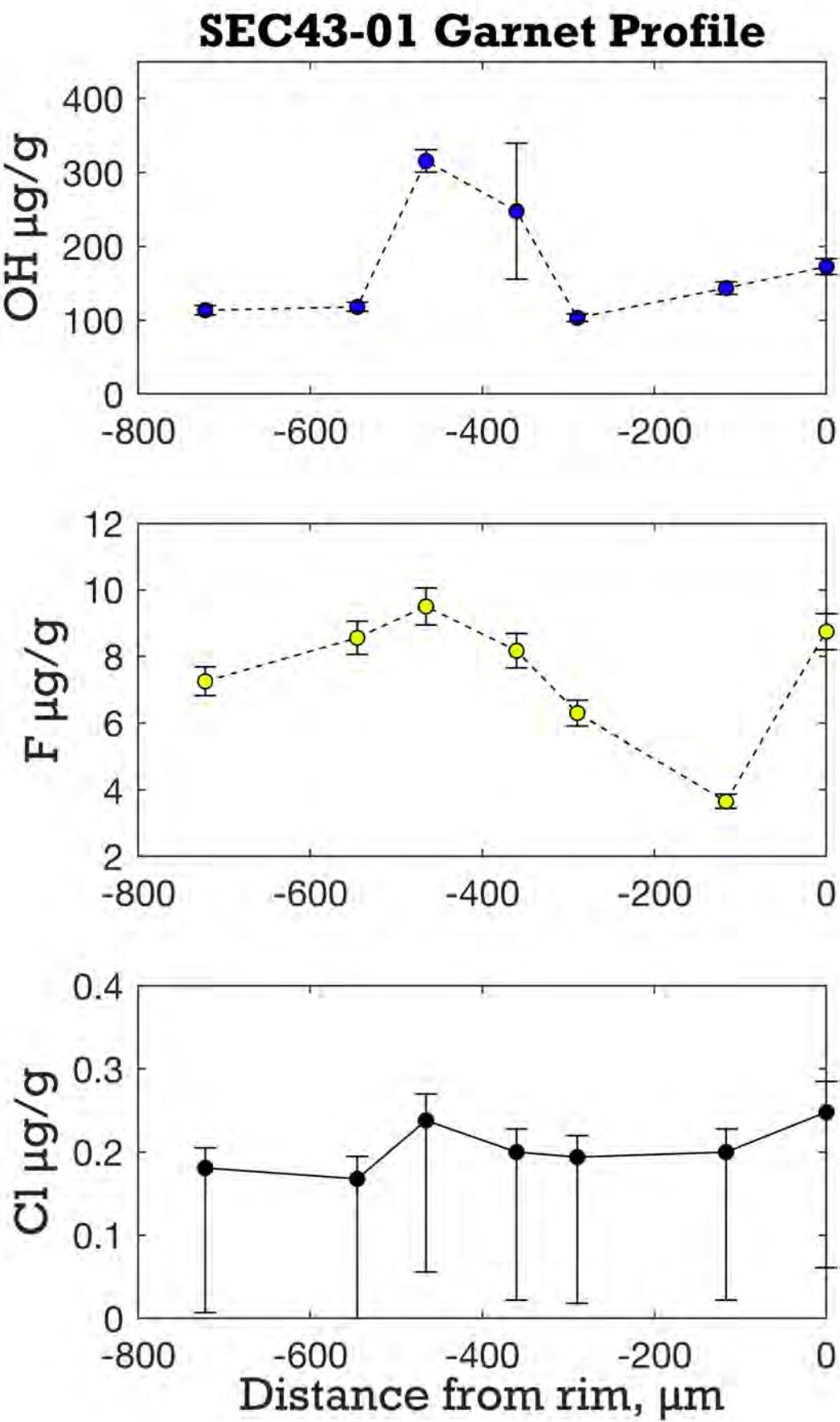


Figure S5.

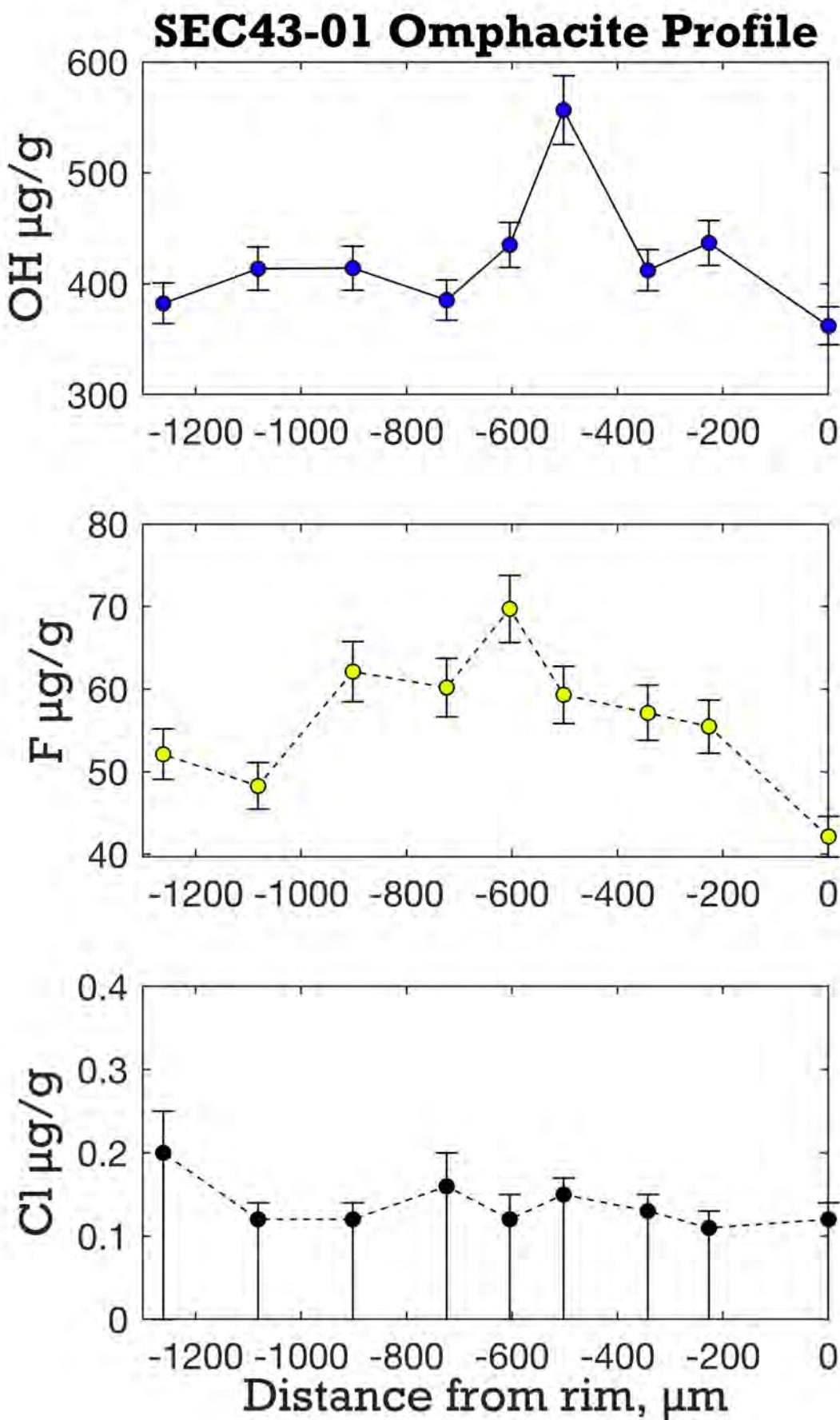


Figure S6.

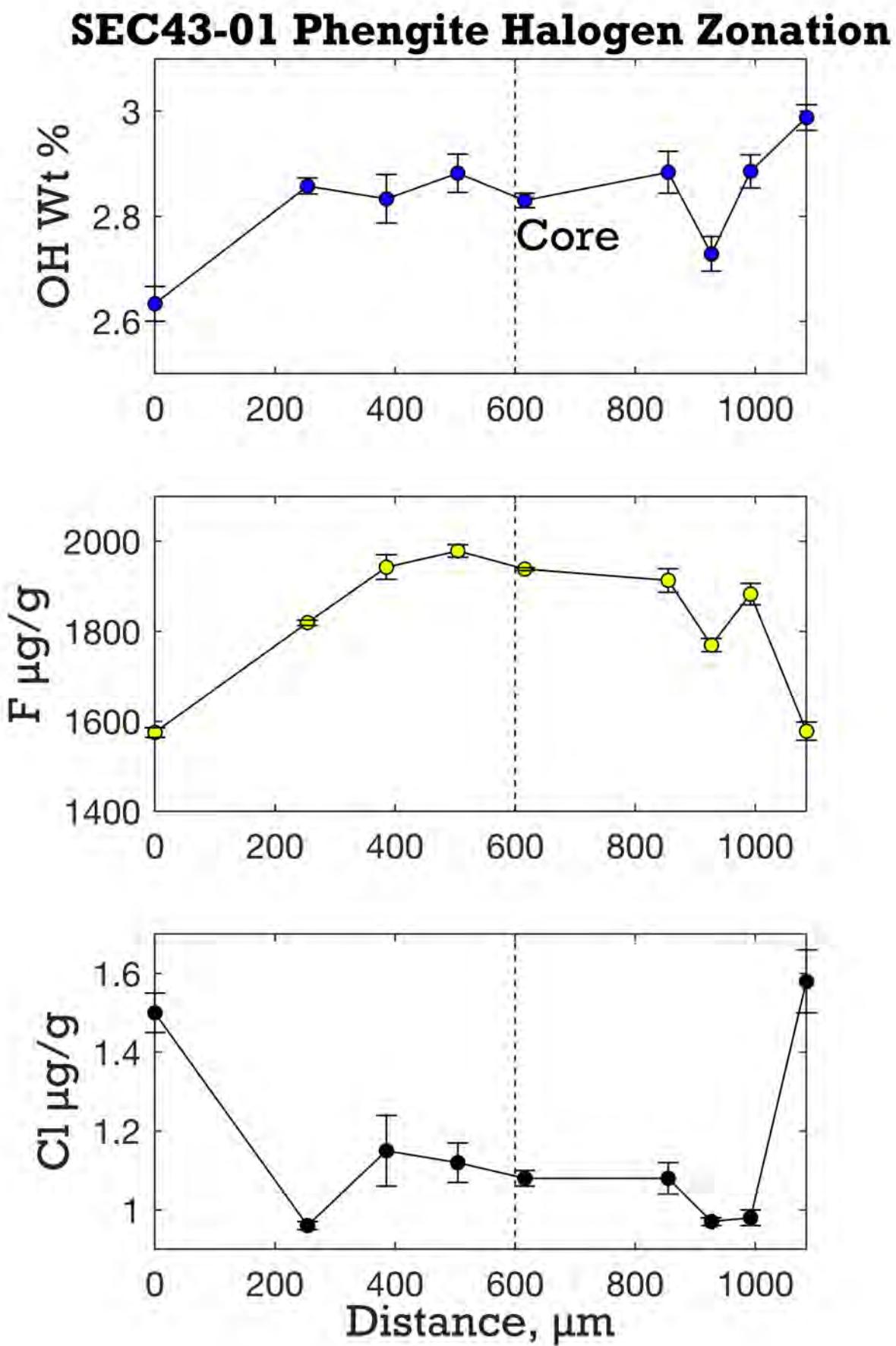


Figure S7.

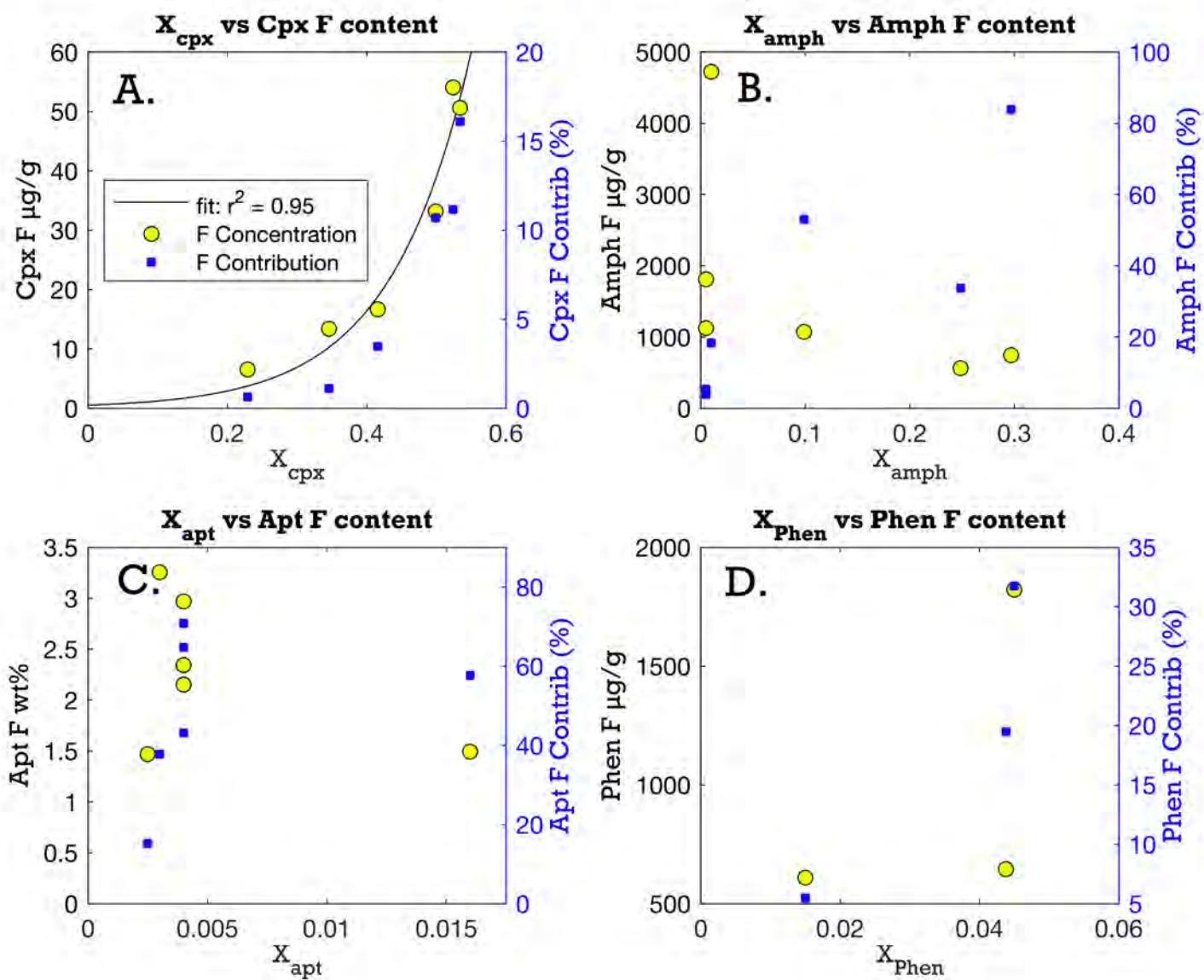


Figure S8.

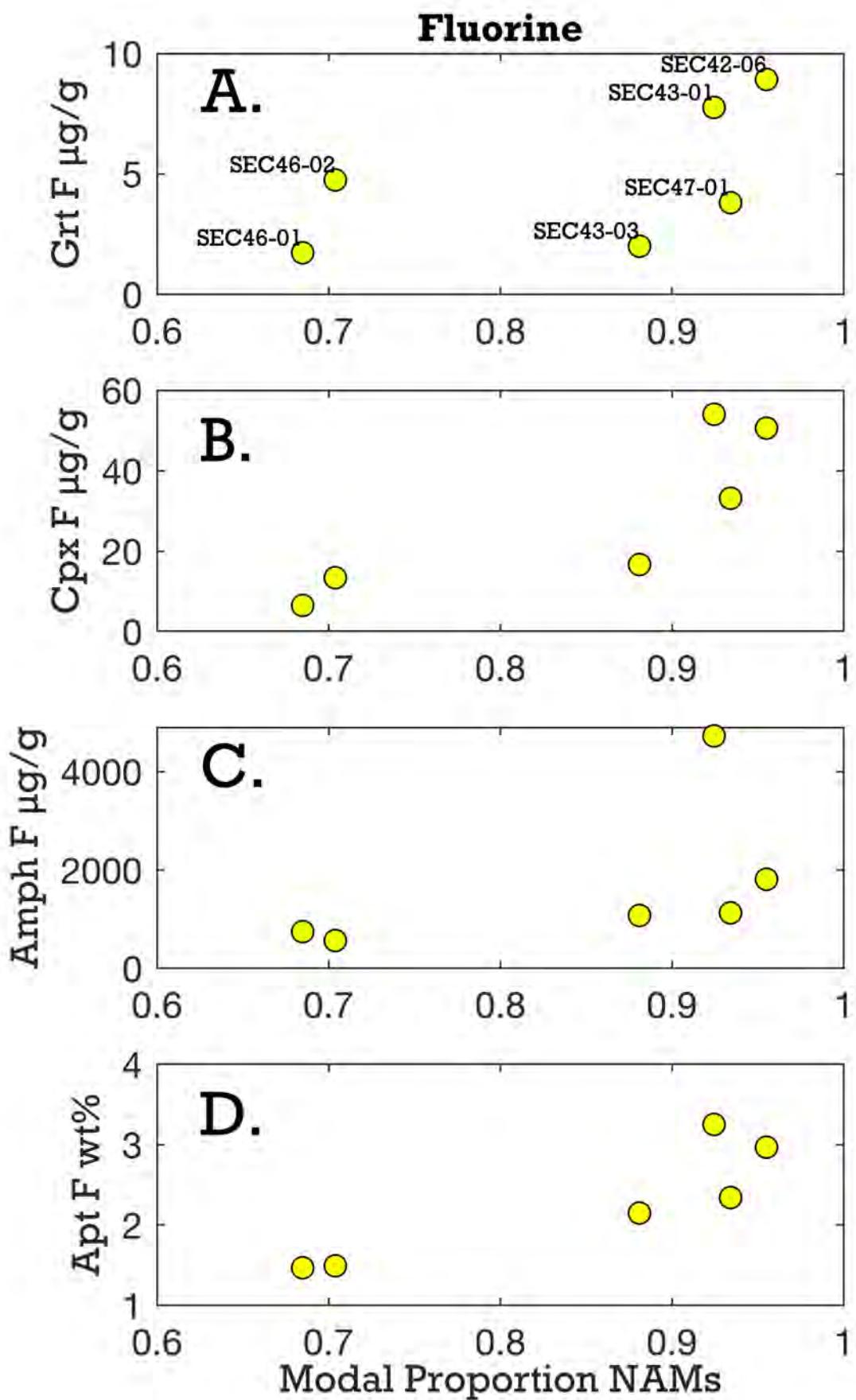


Figure S9.

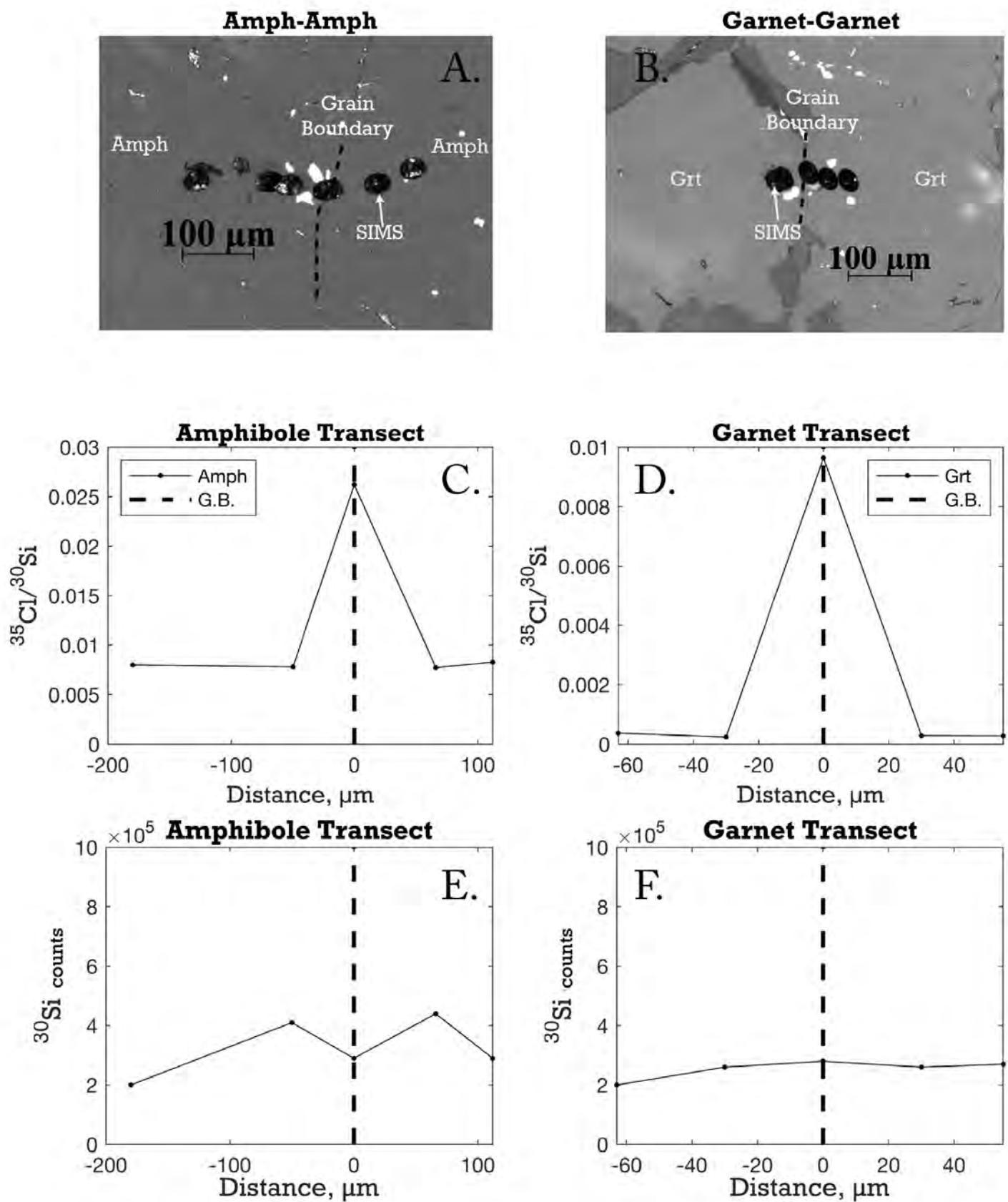


Table 1. Halogen abundances by phase (SIMS)

Sample	Description	No. grains analyzed	No. Analyses	Clinopyroxene								Intra-sample Variability				
				F (μg/g)	Internal + Calibration	Cl (μg/g)	background corrected (μg/g)	CI Error % (2SE)	OH Error % (2SE)	OH Error % (2SE Internal, Calibration, Background)	S (μg/g)	Internal + Calibration	F, 1 sigma	OH, 1 sigma	Cl, 1 sigma	
SEC 42-06	Eclogite	1	1	50.6	6.0%	0.18	0.03	15.7%	651	5.5%	<0.1	12.6%	-	-	-	
SEC 43-01	Eclogite	1	6	54.1	5.8%	0.13	0.00	14.6%	433	4.8%	<0.1	15.5%	15%	13%	-	
SEC 43-03	Eclogite	1	3	16.7	4.5%	0.08	0.00	11.2%	773	5.8%	<0.1	9.2%	3%	3%	-	
SEC46-01	Eclogite	2	5	6.53	5.8%	0.11	0.03	10.5%	446	5.9%	<0.1	8.4%	4%	5%	-	
SEC46-02	Eclogite	1	3	13.4	6.0%	1.17	1.05	9.7%	753	7.7%	<0.1	9.2%	5%	10%	25%	
SEC47-01	Eclogite	3	7	33.2	4.4%	0.07	0.00	12.4%	445	6.0%	<0.1	13.4%	24%	14%	-	
<b>Garnet</b>																
Sample	Description	No. grains analyzed	No. Analyses	F (μg/g)	Internal + Calibration	Cl (μg/g)	background corrected (μg/g)	CI Error % (2SE)	OH Error % (2SE)	OH Error % (2SE Internal, Calibration, Background)	S (μg/g)	Internal + Calibration	F, 1 sigma	OH, 1 sigma	Cl, 1 sigma	
				F (μg/g)	Internal + Calibration	Cl (μg/g)	background corrected (μg/g)	CI Error % (2SE)	OH Error % (2SE)	OH Error % (2SE Internal, Calibration, Background)	S (μg/g)	Internal + Calibration	F, 1 sigma	OH, 1 sigma	Cl, 1 sigma	
SEC 42-06	Eclogite	2	5	8.92	6.0%	0.17	0.02	14.5%	113	5.1%	10.1	<0.1	11.4%	24%	16%	-
SEC 43-01	Eclogite	2	9	7.76	6.0%	0.18	0.03	14.0%	125	5.2%	10.8	<0.1	16.1%	29%	67%	-
SEC 43-03	Eclogite	2	6	2.01	6.1%	0.10	0.02	11.5%	201	5.7%	18.8	<0.1	9.2%	42%	10%	-
SEC46-01	Eclogite	2	7	1.73	5.1%	0.11	0.03	10.0%	85	5.9%	12.3	<0.1	11.5%	93%	27%	-
SEC46-02	Eclogite	2	6	4.76	6.1%	0.11	0.00	11.8%	60	7.6%	13.8	<0.1	10.5%	35%	22%	-
SEC47-01	Eclogite	3	3	3.81	5.3%	0.08	0.00	11.7%	80	6.2%	12.2	<0.1	10.4%	12%	7%	-
<b>Phengite</b>																
Sample	Description	No. grains analyzed	No. Analyses	F (μg/g)	Internal + Calibration	Cl (μg/g)	background corrected (μg/g)	CI Error % (2SE)	OH Error % (2SE)	OH Error % (2SE Internal, Calibration, Background)	S (μg/g)	Internal + Calibration	F, 1 sigma	OH, 1 sigma	Cl, 1 sigma	
				F (μg/g)	Internal + Calibration	Cl (μg/g)	background corrected (μg/g)	CI Error % (2SE)	OH Error % (2SE)	OH Error % (2SE Internal, Calibration, Background)	S (μg/g)	Internal + Calibration	F, 1 sigma	OH, 1 sigma	Cl, 1 sigma	
SEC 42-06	Eclogite	1	2	610	5.9%	2.13	1.98	12.7%	22452	4.6%	0.65	8.4%	6%	1%	4%	
SEC 43-01	Eclogite	1	9	1822	5.9%	1.16	1.01	12.6%	28360	4.6%	0.54	7.6%	8%	4%	20%	
SEC 43-03	Eclogite															
SEC46-01	Eclogite															
SEC46-02	Eclogite															
SEC47-01	Eclogite	2	6	645	4.4%	1.22	1.14	8.7%	26410	5.7%	0.48	5.7%	7%	9%	9%	
<b>Amphibole</b>																
Sample	Description	No. grains analyzed	No. Analyses	F (μg/g)	Internal + Calibration	Cl (μg/g)	background corrected (μg/g)	CI Error % (2SE)	OH Error % (2SE)	OH Error % (2SE Internal, Calibration, Background)	S (μg/g)	Internal + Calibration	F, 1 sigma	OH, 1 sigma	Cl, 1 sigma	
				F (μg/g)	Internal + Calibration	Cl (μg/g)	background corrected (μg/g)	CI Error % (2SE)	OH Error % (2SE)	OH Error % (2SE Internal, Calibration, Background)	S (μg/g)	Internal + Calibration	F, 1 sigma	OH, 1 sigma	Cl, 1 sigma	
SEC 42-06	Eclogite	1	5	1812	4.7%	82.7	82.6	13.1%	18682	4.2%	27.9	11.0%	18%	5%	37%	
SEC 43-01	Eclogite	1	3	4727	5.9%	11.6	11.5	12.2%	20187	4.6%	18.0	5.9%	3%	3%	24%	
SEC 43-03	Eclogite	1	2	1073	4.6%	0.64	0.56	8.5%	17952	5.9%	27.0	4.4%				
SEC46-01	Eclogite	4	8	747	5.5%	10.3	10.3	8.7%	17292	5.9%	5.6	5.3%				
SEC46-02	Eclogite	2	6	563	6.2%	11.8	11.7	8.6%	23449	7.5%	5.9	6.1%				
SEC47-01	Eclogite	1	3	1127	5.0%	15.5	15.4	8.6%	15747	5.8%	3.0	5.2%				
<b>Titanite</b>																
Sample	Description	No. grains analyzed	No. Analyses	F (μg/g)	Internal + Calibration	Cl (μg/g)	background corrected (μg/g)	CI Error % (2SE)	OH Error % (2SE)	OH Error % (2SE Internal, Calibration, Background)	S (μg/g)	Internal + Calibration	F, 1 sigma	OH, 1 sigma	Cl, 1 sigma	
				F (μg/g)	Internal + Calibration	Cl (μg/g)	background corrected (μg/g)	CI Error % (2SE)	OH Error % (2SE)	OH Error % (2SE Internal, Calibration, Background)	S (μg/g)	Internal + Calibration	F, 1 sigma	OH, 1 sigma	Cl, 1 sigma	
SEC 46-02	Eclogite	1	3	1394	6.0%	0.11		9.7%	4010	7.4%	<0.1	9.3%				

Table 2. Calculated mineral modes

Est. Modal Abundances	Phase Fraction								NAMs (Grt+Cpx+Qtz)	
	Garnet	Omphacite	Phengite	Amphibole	Quartz	Rutile	Titanite	Apatite		
SEC 42-06	0.421(32)	0.534(32)	0.015(2)	0.0050(5)	0	0.021(2)	0	0.0040(4)	1	0.96
SEC 43-01	0.384(30)	0.524(31)	0.045(4)	0.010(1)	0.017(2)	0.018(2)	0	0.0030(3)	1	0.92
SEC 43-03	0.378(28)	0.416(28)	0	0.099(10)	0.087(2)	0.016(2)	0	0.0040(4)	1	0.88
SEC46-01	0.456(28)	0.229(21)	0	0.297(25)	0	0.016(2)	0	0.0025(3)	1	0.68
SEC46-02	0.358(27)	0.346(27)	0	0.248(22)	0	0.010(1)	0.02(1)	0.016(2)	1	0.70
SEC47-01	0.406(31)	0.499(31)	0.044(4)	0.0050(5)	0.024	0.018(2)	0	0.0040(4)	1	0.93
Average Mineral Modes	0.41	0.43	0.011	0.107	0.023	0.000	0.018	0.006	1.000	

Values in parenthesis represent 1 sigma  
 uncertainty

Table 3. Halogen bulk rock abundances, ratios

Sample	Calculated Bulk (WHOI)			Uncorrected for Yield		Corrected for Yield		Uncertainty (Yield Corrected)						
	OH µg/g	F µg/g	Cl µg/g	F µg/g	Cl µg/g	F µg/g	Cl µg/g	1SD Propogated	Calculated/Measured Bulk	F %	Cl %	Cl/Nb**	Cl/K**	F/Nd**
SEC42-06	901	168	2.5	92	8.9	115	11.0	15	2.2	146%	23%	3.03	0.0089	10.8
SEC43-01	1783	258	1.4	109	7.1	137	8.9	68	3.1	189%	16%	4.34	0.0020	22.5
SEC43-03	2342	200	0.8	152	5.4	189	6.7	11	2.2	105%	12%	2.80	0.0116	15.8
SEC46-01	5123	241	3.2	154	9.1	193	11.3	52	3.1	125%	28%	5.18	0.0085	17.3
SEC46-02	6836	415	8.2	160	16.1	199	19.9	12	2.4	208%	41%	4.68	0.0218	32.1
SEC47-01	1663	145	1.3	112	7.1	139	8.8	9.1	1.5	104%	15%	4.23	0.0021	11.2
Average	237.7	2.9	129.8	8.9	162.0	11.1								
Standard Deviation	96.6	2.7	29.2	3.8	36.1	4.6								
Average*	202.3	1.85	123.7	7.5	154.5	9.4						3.9	0.0066	15.5
Standard Deviation*	47.7	0.987	28.0	1.51	34.9	1.88						1.0	0.0043	4.8
AOC (Barnes et al. 2018)	216	207												
Average Recycling Efficiency (to mantle)	93.7%	0.9%	57.3%	3.6%	71.5%	4.5%								

Notes

Evidence for fluid interaction, minor retrogression  
Table 1. Halogen abundances by phase (SIMS)

\* Excludes SEC46-02 due to retrograde alteration

\* Values used in bulk flux calculations

\*\* Cl and F values from yield corrected pyrohydrolysis, K, Nb, Nd using bulk rock chemistry from John et al. 2010





SEC46-01 Average	0.02	22.50	8.09	0.87	0.03	0.06	38.55	0.00	24.61	5.99	100.72	0.00	2.04	0.67	0.06	0.00	0.00	2.97	0.00	1.58	0.69	8.01	0.53	0.23	0.22	0.02	0.70		
Line 1 SEC46-02_gr1	0	-216	0.03	22.16	10.33	0.71	0.06	0.13	38.48	0.00	24.07	4.87	100.82	0.0018	2.016	0.853	0.0464	0.0093	0.0077	2.97	0	1.554	0.56	8.0182	0.52	0.19	0.28	0.02	0.74
Line 2 SEC46-02_gr1	54	-162	0.02	22.29	10.56	0.65	0.04	0.14	38.77	0.00	23.51	5.04	101.03	0.0013	2.017	0.869	0.0424	0.0067	0.0084	2.977	0	1.51	0.576	8.0078	0.50	0.19	0.29	0.01	0.72
Line 3 SEC46-02_gr1	108	-108	0.05	22.39	10.48	0.67	0.05	0.09	38.42	0.00	23.82	4.88	100.85	0.0028	2.034	0.866	0.0436	0.0078	0.0054	2.962	0	1.536	0.561	8.0186	0.51	0.19	0.29	0.01	0.73
Line 4 SEC46-02_gr1	163	-53	0.02	22.60	9.58	0.57	0.04	0.11	38.78	0.00	24.21	5.41	101.31	0.0011	2.039	0.785	0.0367	0.0056	0.0062	2.968	0	1.55	0.617	8.0086	0.52	0.21	0.26	0.01	0.72
Line 5 SEC46-02_gr1	216	0	0.00	22.44	8.17	0.65	0.05	0.06	38.69	0.00	25.31	5.70	101.07	0	2.032	0.672	0.0425	0.0068	0.0037	2.973	0	1.626	0.653	8.009	0.54	0.22	0.22	0.01	0.71
Line 1 SEC46-02_gr2	0	-222	0.01	22.44	11.06	0.62	0.03	0.07	38.73	0.00	22.83	4.96	100.73	0.0005	2.033	0.911	0.0401	0.0039	0.0038	2.977	0	1.467	0.568	8.0043	0.49	0.19	0.31	0.01	0.72
Line 2 SEC46-02_gr2	23	-198	0.06	22.47	10.01	0.58	0.06	0.05	38.68	0.00	24.89	4.66	101.45	0.0014	2.034	0.824	0.0374	0.0089	0.0028	2.97	0	1.599	0.534	8.0135	0.53	0.18	0.28	0.01	0.75
Line 3 SEC46-02_gr2	48	-174	0.00	22.56	10.55	0.60	0.09	0.04	38.54	0.00	23.75	4.94	101.07	0.0003	2.044	0.968	0.039	0.0127	0.0022	2.961	0	1.527	0.56	8.0202	0.51	0.19	0.29	0.01	0.73
Line 4 SEC46-02_gr2	72	-150	0.01	22.43	10.37	0.60	0.00	0.04	38.51	0.00	24.63	4.57	101.16	0.0008	2.037	0.856	0.039	0	0.0021	2.967	0	1.587	0.524	8.0129	0.53	0.17	0.28	0.01	0.75
Line 5 SEC46-02_gr2	95	-127	0.03	22.33	9.73	0.70	0.03	0.01	38.31	0.00	26.11	4.21	101.47	0.0002	2.033	0.806	0.0457	0.0051	0.0008	2.96	0	1.687	0.485	8.0246	0.56	0.16	0.27	0.02	0.78
Line 6 SEC46-02_gr2	122	-100	0.03	22.48	9.50	0.66	0.06	0.10	38.84	0.00	24.54	5.40	101.62	0.0016	2.025	0.778	0.043	0.0095	0.0059	2.969	0	1.569	0.615	8.0116	0.52	0.20	0.26	0.01	0.72
Line 7 SEC46-02_gr2	146	-76	0.01	22.53	9.86	0.63	0.01	0.13	38.69	0.00	24.00	5.49	101.36	0.0007	2.033	0.808	0.0405	0.0021	0.0078	2.961	0	1.536	0.626	8.0152	0.51	0.21	0.27	0.01	0.71
Line 8 SEC46-02_gr2	171	-51	0.02	22.60	9.72	0.56	0.05	0.12	38.80	0.00	24.15	5.30	101.31	0.0001	2.039	0.797	0.0363	0.007	0.0067	2.97	0	1.546	0.604	8.007	0.52	0.20	0.27	0.01	0.72
Line 9 SEC46-02_gr2	194	-28	0.01	22.39	9.73	0.59	0.01	0.10	38.64	0.00	24.83	4.95	101.24	0.0004	2.029	0.801	0.0381	0.0018	0.0055	2.971	0	1.596	0.567	8.0098	0.53	0.19	0.27	0.01	0.74
Line 10 SEC46-02_gr2	222	0	0.06	22.63	8.32	0.59	0.04	0.02	38.98	0.00	24.64	5.98	101.26	0.0014	2.038	0.681	0.038	0.0006	0.0012	2.979	0	1.575	0.681	8.0026	0.53	0.23	0.23	0.01	0.70
SEC46-02 Average	0.02	22.45	9.86	0.62	0.04	0.08	38.66	0.00	24.35	5.09	101.18	0.00	2.03	0.81	0.04	0.01	0.00	2.97	0.00	1.56	0.58	8.01	0.52	0.19	0.27	0.01	0.73		
Line 1 SEC47-01_gr1	0	-146	0.00	22.07	9.78	1.01	0.08	0.12	38.40	0.00	25.80	4.20	101.46	0	2.011	0.81	0.0662	0.0116	0.0069	2.968	0	1.668	0.484	8.0258	0.55	0.16	0.27	0.02	0.78
Line 2 SEC47-01_gr1	37	-109	0.05	22.28	9.43	0.59	0.06	0.15	38.37	0.00	26.22	4.58	101.73	0.0032	2.022	0.778	0.0384	0.009	0.0087	2.955	0.0001	1.689	0.526	8.0294	0.56	0.17	0.26	0.01	0.76
Line 3 SEC47-01_gr1	73	-73	0.03	22.25	8.06	0.55	0.00	0.12	38.55	0.00	26.99	5.00	101.56	0.0021	2.02	0.665	0.0361	0	0.0068	2.969	0	1.739	0.574	8.0112	0.58	0.19	0.22	0.01	0.75
Line 4 SEC47-01_gr1	110	-36	0.03	22.38	7.59	0.64	0.00	0.11	38.43	0.00	27.12	5.18	101.48	0.0021	2.033	0.626	0.0416	0	0.0065	2.962	0	1.748	0.595	8.0142	0.58	0.20	0.21	0.01	0.75
Line 5 SEC47-01_gr1	146	0	0.04	22.35	8.89	0.64	0.00	0.07	38.63	0.00	26.42	4.64	101.67	0.0022	2.026	0.733	0.0418	0	0.0038	2.971	0	1.7	0.532	8.0098	0.57	0.18	0.24	0.01	0.76
Line 1 SEC47-01_gr2	0	-112	0.03	22.17	9.72	0.64	0.09	0.14	38.41	0.00	26.12	4.26	101.57	0.0004	2.017	0.804	0.0417	0.013	0.0083	2.965	0	1.686	0.49	8.0254	0.56	0.16	0.27	0.01	0.77
Line 2 SEC47-01_gr2	28	-84	0.04	22.16	9.46	0.70	0.00	0.13	38.42	0.00	26.10	4.49	101.52	0.002	2.017	0.782	0.0457	0.0004	0.0075	2.964	0	1.684	0.516	8.0188	0.56	0.17	0.26	0.02	0.77
Line 3 SEC47-01_gr2	56	-56	0.05	22.05	9.26	0.44	0.07	0.15	38.40	0.00	26.85	4.39	101.66	0.0011	2.007	0.766	0.0286	0.0098	0.009	2.965	0	1.733	0.505	8.0265	0.57	0.17	0.25	0.01	0.77
Line 4 SEC47-01_gr2	84	-28	0.02	22.42	8.13	0.41	0.07	0.08	38.47	0.00	27.12	4.85	101.57	0.0013	2.036	0.671	0.0268	0.0101	0.0045	2.964	0	1.748	0.556	8.0177	0.58	0.19	0.22	0.01	0.76
Line 5 SEC47-01_gr2	112	0	0.04	22.52	7.99	0.58	0.07	0.07	38.50	0.00	26.35	5.25	101.37	0.0027	2.045	0.659	0.0376	0.0097	0.0042	2.962	0	1.696	0.602	8.0162	0.57	0.20	0.22	0.01	0.74
Line 1 SEC47-01_gr3	0	160	0.10	22.08	0.79	0.77	0.05	0.09	38.39	0.00	26.05	4.13	101.44	0.0006	2.013	0.81	0.0502	0.0039	0.0052	2.968	0	1.605	0.476	8.0203	0.55	0.16	0.27	0.02	0.78
Line 2 SEC47-01_gr3	48	-112	0.16	22.10	0.75	0.63	0.06	0.12	38.50	0.00	25.85	4.49	101.66	0.01	2.007	0.804	0.0401	0.009	0.0071	2.966	0	1.645	0.514	8.0231	0.55	0.17	0.27	0.01	0.76
Line 3 SEC47-01_gr3	96	-74	0.09	22.10	8.86	0.66	0.01	0.16	38.25	0.00	26.91	4.35	101.40	0.0058	2.017	0.735	0.0414	0.0022	0.0091	2.962	0	1.743	0.502	8.0195	0.58	0.17	0.24	0.01	0.78
Line 4 SEC47-01_gr3	123	-36	0.03	22.40	7.98	0.57	0.01	0.11	38.39	0.00	27.23	4.80	101.52	0.0017	2.037	0.66	0.0373	0.0022	0.0063	2.962	0	1.757	0.552	8.0155	0.58	0.18	0.22	0.01	0.76
Line 5 SEC47-01_gr3	160	0	0.02	22.47	8.11	0.62	0.04	0.07	38.75	0.00	25.83	5.59	101.50	0.0014	2.03	0.666	0.0404	0.0053	0.0038	2.971	0	1.656	0.639	8.0129	0.55	0.21	0.22	0.01	0.72
SEC47-01 Average	0.05	22.25	8.85	0.63	0.04	0.11	38.46	0.00	26.46	4.68	101.54	0.00	2.02	0.73	0.04	0.01	0.01	2.96	0.00	1.71	0.54	8.02	0.57	0.18	0.24	0.01	0.76		



Supplementary Table 3. Phengite major

	Cumulative D <sub>50</sub> μm	Dist to grain boundary (μm)	TiO <sub>2</sub>	F	Cl	MnO	Na <sub>2</sub> O	Cr <sub>2</sub> O <sub>3</sub>	SiO <sub>2</sub>	K <sub>2</sub> O	FeO	MgO	Al <sub>2</sub> O <sub>3</sub>	CaO	Total	H <sub>2</sub> O (100 minus sum)	F μg/g	Cl μg/g
Line 1 SEC42-06_Phen1	0	-123	1.04	0.00	0.00	0.00	1.27	0.00	47.68	8.89	2.60	2.45	29.31	0.01	93.26	6.74	-	-
Line 2 SEC42-06_Phen1	31	-92	0.97	0.00	0.00	0.00	1.08	0.03	47.59	8.87	2.54	2.44	29.35	0.00	92.88	7.12	-	-
Line 3 SEC42-06_Phen1	61	-62	1.05	0.00	0.01	0.00	1.23	0.05	47.54	8.95	2.53	2.39	29.45	0.00	93.19	6.81	-	-
Line 4 SEC42-06_Phen1	91	-32	0.99	0.00	0.01	0.00	1.18	0.02	47.87	8.87	2.50	2.47	29.56	0.00	93.47	6.53	-	82
Line 5 SEC42-06_Phen1	123	0	0.92	0.00	0.01	0.00	1.43	0.03	47.98	8.77	2.51	2.40	29.42	0.02	93.49	6.51	-	106
SEC42-06 Average			0.99	0.00	0.01	0.00	1.24	0.03	47.73	8.87	2.54	2.43	29.42	0.01	93.26	6.74	-	-
Line 1 43-01_Phen1	0	-1098	0.92	0.00	0.01	0.01	1.15	0.06	48.19	8.87	1.72	2.65	29.43	0.01	93.02	6.98	-	71
Line 2 43-01_Phen1	121	-977	0.76	0.06	0.01	0.01	0.98	0.07	48.98	9.26	1.86	2.87	28.76	0.00	93.61	6.39	648	-
Line 3 43-01_Phen1	243	-855	0.74	0.02	0.00	0.00	0.96	0.05	49.06	9.22	1.87	2.91	28.37	0.00	93.20	6.80	-	-
Line 4 43-01_Phen1	390	-708	0.78	0.02	0.00	0.02	1.00	0.07	49.12	9.30	1.98	3.00	28.60	0.00	93.90	6.10	-	-
Line 5 43-01_Phen1	485	-613	0.68	0.07	0.01	0.00	0.95	0.11	48.83	9.22	1.87	3.05	28.66	0.00	93.45	6.55	668	102
Line 6 43-01_Phen1	609	-489	0.62	0.01	0.01	0.01	1.07	0.08	48.98	9.23	1.87	3.02	28.63	0.00	93.53	6.47	-	125
Line 7 43-01_Phen1	760	-338	0.81	0.00	0.01	0.01	1.04	0.04	48.98	9.30	1.99	2.92	28.75	0.01	93.84	6.16	-	63
Line 8 43-01_Phen1	855	-243	0.90	0.00	0.01	0.00	1.07	0.06	48.42	9.20	1.89	2.75	29.02	0.00	93.33	6.67	-	114
Line 9 43-01_Phen1	977	-121	0.96	0.04	0.00	0.00	1.18	0.05	48.13	9.08	1.90	2.64	29.43	0.00	93.41	6.59	-	-
Line 10 43-01_Phen1	1098	0	0.88	0.00	0.01	0.00	1.29	0.02	48.01	8.93	1.77	2.56	29.84	0.00	93.30	6.70	-	71
SEC43-01 Average			0.80	0.02	0.01	0.01	1.07	0.06	48.67	9.16	1.87	2.84	28.95	0.00	93.46			
Line 1 SEC47-01 phen1	0	-1108	0.75	0.05	0.00	0.00	1.20	0.06	48.42	9.09	2.32	2.83	29.1	0	93.86	6.14	519	-
Line 2 sec47-01 phen1	277	-831	0.73	0.00	0.00	0.00	1.06	0.05	48.4	9.18	2.36	2.92	28.9	0	93.58	6.42	-	-
Line 3 sec47-01 phen1	553	-555	0.69	0.04	0.00	0.01	1.01	0.03	48.77	9.13	2.28	2.85	29.0	0.0054	93.78	6.22	434	-
Line 4 sec47-01 phen1	863	-244	0.75	0.00	0.00	0.00	1.12	0.01	48.76	9.15	2.36	2.76	29.3	0.0107	94.25	5.75	-	-
Line 5 sec47-01 phen1	1108	0	0.69	0.00	0.00	0.00	1.09	0.07	48.66	9.21	2.24	2.73	29.3	0	93.96	6.04	-	-
Line 1 SEC47-01 phen2	0	-305	0.64	0.00	0.00	0.00	1.14	0.07	48.3	9.18	2.37	2.92	28.5	0	93.15	6.85	-	-
Line 2 SEC47-01 phen2	29	-275	0.58	0.00	0.00	0.01	0.97	0.06	49.13	9.28	2.42	3.05	28.2	0.0061	93.67	6.33	-	-
Line 3 SEC47-01 phen2	68	-237	0.62	0.00	0.00	0.00	1.03	0.05	48.94	9.34	2.46	3.11	28.1	0	93.59	6.41	-	-
Line 4 SEC47-01 phen2	101	-204	0.63	0.00	0.00	0.01	1.04	0.04	49.08	9.32	2.31	3.01	28.4	0.0097	93.84	6.16	-	-
Line 5 SEC47-01 phen2	136	-169	0.53	0.00	0.00	0.00	1.00	0.03	48.98	9.13	2.47	3.36	27.9	0.0286	93.44	6.56	-	-
Line 6 SEC47-01 phen2	169	-135	0.57	0.00	0.00	0.01	0.94	0.06	49.15	9.41	2.47	3.04	28.3	0	93.91	6.09	-	-
Line 7 SEC47-01 phen2	203	-102	0.59	0.00	0.00	0.00	0.89	0.06	49.15	9.38	2.43	3.05	28.3	0	93.81	6.19	-	-
Line 8 SEC47-01 phen2	237	-68	0.58	0.00	0.00	0.00	0.94	0.06	49.23	9.47	2.53	2.98	28.1	0	93.89	6.11	-	-
Line 9 SEC47-01 phen2	270	-34	0.53	0.00	0.00	0.02	1.02	0.07	48.15	8.98	3.94	2.74	28.3	0.04	93.82	6.18	-	-
Line 10 SEC47-01 phen2	305	0	0.63	0.00	0.00	0.00	1.05	0.06	48.01	9.01	2.65	2.44	29.8	0.0097	93.64	6.36	-	-
SEC47-01 Average			0.63	0.01	0.00	0.00	1.03	0.05	48.74	9.22	2.51	2.92	28.62	0.01	93.75			



Supplementary Table 5. Apatite Compositions (EMPA)

	# grains	# measurements	P2O5	1 s.d. (%)	F	1 s.d. (%)	Weight %				Total	OH
							CaO	1 s.d. (%)	Cl	1 s.d. (%)		
SEC 42-06	5	23	40.94	0.8%	2.97	7.7%	53.94	0.8%	0.051	16%	97.89	2.11
SEC 43-01	5	27	41.62	1.5%	3.25	3.2%	53.52	1.1%	0.041	75%	98.44	1.56
SEC 43-03	3	10	41.93	0.6%	2.15	4.4%	52.46	1.2%	0.019	54%	96.56	3.44
SEC46-01	3	9	42.02	0.6%	1.47	5.6%	52.19	0.5%	0.020	14%	95.70	4.30
SEC46-02	3	9	42.00	0.5%	1.49	6.6%	52.39	0.3%	0.030	13%	95.91	4.09
SEC47-01	3	9	41.92	0.7%	2.34	5.7%	51.64	0.3%	0.029	18%	95.93	4.07

Supplementary Table 6. Titanite major elem:

Comment	No.	TiO <sub>2</sub>	F	K <sub>2</sub> O	Cl	Na <sub>2</sub> O	Cr <sub>2</sub> O <sub>3</sub>	SiO <sub>2</sub>	CaO	MnO	MgO	Al <sub>2</sub> O <sub>3</sub>	FeO
Line 1 SEC46-02_titanite	46	36.97	0	0.012	0.0092	0.012	0.11	29.65	28.38	0.0454	0.01	1.84	0.20
Line 2 SEC46-02_titanite	47	36.78	0	0.014	0.0124	0.021	0.10	29.91	28.23	0.0285	0.00	1.85	0.22
Line 3 SEC46-02_titanite	48	36.48	0	0.014	0.0038	0.097	0.08	29.83	28.3	0.0310	0.00	2.07	0.20
Line 4 SEC46-02_titanite	49	36.96	0	0.012	0.0065	0.044	0.07	29.89	28.26	0.0060	0.00	1.79	0.21
Line 5 SEC46-02_titanite	50	36.91	0	0.015	0.0086	0.025	0.07	29.73	28.18	0.0320	0.00	1.78	0.28
Average		36.82	0	0.013	0.0081	0.040	0.09	29.802	28.27	0.0286	0.00	1.87	0.22

Total
97.24
97.17
97.11
97.25
97.03
97.16

Supplementary Table 7. Halogen intra-grain profiles (SIMS)

**SEC43-01**

									Grain Mean	Grain S.D. (%)
<b>Garnet</b>	SEC43-01_grt_1@44 SEC43-01_grt_07@50 SEC43-01_grt_3@46 SEC43-01_grt_06@49 SEC43-01_grt_04@47 SEC43-01_grt_5@48 SEC43-01_grt_2@45									
OH µg/g	114	118	316	248	104	144	173		174	46%
OH Error % (2SE Int+Calib)	5.4%	5.0%	4.8%	37.1%	5.1%	5.8%	6.2%			
F µg/g	7.25	8.56	9.51	8.17	6.29	3.66	8.74		7.45	26%
F Error % (2SE Int+Calib)	6.0%	5.8%	5.9%	6.3%	6.2%	5.9%	6.1%			
S µg/g	-	0.04	0.06	0.05	0.05	0.07	-			
C Error % (2SE Int+Calib)	-	13.8%	17.7%	17.1%	12.9%	16.0%	-			
Cl µg/g	0.18	0.17	0.24	0.20	0.19	0.20	0.25		0.20	14%
Cl Error % (2SE Int+Calib)	13.0%	16.1%	13.4%	13.8%	13.5%	13.8%	14.8%			
Distance (µm)	722	546	465	360	290	117	0			
									Rim	
									Core	
<b>Omphacite</b>	SEC43-01_omph_1@70 SEC43-01_omph_3@72 SEC43-01_omph_5@74 SEC43-01_omph_7@76 SEC43-01_omph_8@77 SEC43-01_omph_9@78 SEC43-01_omph_6@75 SEC43-01_omph_4@73 SEC43-01_omph_2@71									
OH µg/g	362	437	412	557	435	385	414	414	383	422
OH Error % (2SE Int+Calib)	4.7%	4.6%	4.6%	5.6%	4.7%	4.7%	4.8%	4.7%	4.8%	13%
F µg/g	42.1	55.5	57.1	59.3	69.7	60.2	62.1	48.3	52.1	56.3
F Error % (2SE Int+Calib)	5.8%	5.8%	5.8%	5.8%	5.8%	5.9%	5.9%	5.8%	5.9%	14%
S µg/g	0.03	0.03	-	0.03	0.02	-	0.04	-	-	
C Error % (2SE Int+Calib)	13.4%	16.0%	-	17.4%	13.7%	-	12.2%	-	-	
Cl µg/g	0.12	0.11	0.13	0.15	-	-	0.12	0.12	-	0.13
Cl Error % (2SE Int+Calib)	16.2%	15.1%	12.2%	14.7%	-	-	13.8%	13.5%	-	11%
Distance (µm)	0	227	343	502	604	724	902	1082	1262	
									Core	
<b>Phengite</b>	SEC43-01_phen_2@52 SEC43-01_phen_3@53 SEC43-01_phen_5@55 SEC43-01_phen_4@54 SEC43-01_phen_1@51 SEC43-01_phen_9@63 SEC43-01_phen_8@62 SEC43-01_phen_7@61 SEC43-01_phen_6@60									
OH µg/g	26335	28580	28337	28820	28304	28840	27286	28857	29884	28360
OH Error % (2SE Int+Calib)	4.6%	4.5%	4.7%	4.6%	4.5%	4.7%	4.6%	4.6%	4.6%	4%
F µg/g	1575	1819	1943	1979	1938	1913	1769	1883	1578	1822
F Error % (2SE Int+Calib)	5.8%	5.8%	6.0%	5.9%	5.8%	5.9%	5.9%	5.9%	5.9%	8%
S µg/g	0.90	0.53	0.42	0.43	0.45	0.58	0.52	0.50	0.51	
C Error % (2SE Int+Calib)	7.4%	7.4%	10.0%	6.9%	8.3%	6.7%	8.1%	7.0%	6.7%	
Cl µg/g	1.50	0.96	1.15	1.12	1.08	1.08	0.97	0.98	1.58	1.16
Cl Error % (2SE Int+Calib)	12.5%	12.1%	13.9%	12.6%	12.2%	12.6%	12.2%	12.2%	12.8%	20%
Distance (µm)	0	254	386	505	616	855	927	992	1086	
									Rim, to omphacite grain boundary	
									Core	
									Rim, to grt grain boundary	



Supplementary Table 9. Pyrohydrolysis raw data and reference material replicates

Analysis #	Uncorrected for Yield				Corrected for Yield				Batch
	F (ppm)	C1 (ppm)	F (ppm)	C1 (ppm)	F (ppm)	C1 (ppm)	F (ppm)	C1 (ppm)	
SEC42-6	2	99.74	10.00	Average	91.94	8.85	130.98	13.13	1
	3	84.72	8.35	Std. Dev.	7.53	1.00	102.35	10.09	2
	4	91.34	8.20	StdDev%	8%	11%	110.35	9.90	2
					StdDev%	12.89%	16%		
SEC43-1	2	163.91	9.27	Average	108.51	7.05	215.24	12.17	1
	3	77.69	6.40	Std. Dev.	48.08	1.97	93.86	7.73	2
	4	83.92	5.49	StdDev%	44%	28%	101.39	6.64	2
					StdDev%	49.71%	33.12%		
SEC43-3	2	151.98	6.73	Average	152.25	5.37	195.88	8.84	1
	3	148.65	5.44	Std. Dev.	3.74	1.40	179.58	6.57	2
	4	156.12	3.94	StdDev%	2%	26%	188.61	4.76	2
					StdDev%	5.3%	30.4%		
SEC46-1	2	192.26	11.03	Average	154.37	9.06	252.47	14.49	1
	3	128.56	7.56	Std. Dev.	33.52	1.78	155.31	9.14	2
	4	142.80	8.59	StdDev%	22%	20%	171.91	10.38	2
					StdDev%	26.9%	24.7%		
SEC46-2	2	162.16	15.16	Average	160.25	16.06	212.95	19.90	1
	3	156.39	16.95	Std. Dev.	3.35	0.90	188.93	20.47	2
	4	162.22	16.09	StdDev%	2%	6%	195.98	19.44	2
					StdDev%	6%	3%		
SEC47-1	2	113.07	7.62	Average	111.60	7.10	148.48	10.01	1
	3	108.15	7.03	Std. Dev.	3.00	0.49	130.65	8.49	2
	4	113.58	6.65	StdDev%	3%	7%	137.22	8.03	2
					StdDev%	6%	12%		

<sup>a</sup>postscript (') indicates a duplicated sample

Standards	
F ppm	C1 ppm
JB-2-1	76.65
	268.22
JB-2-2	73.70
	215.25
JB-2-3	74.67
	233.68
Batch 1 Mean	75.01
	239.05
Batch 1 Std	1.50
	26.89
Yield Batch 1	76.1%
	85.1%
Yield Std (%)	2.0%
	11.3%
JB-2-9	80.64
	191.80
JB-2-11	82.04
	208.73
JB-2-12	81.93
	241.47
Batch 2 Mean	81.53
	214.00
Batch 2 Std	0.78
	25.25
Yield Batch 2	82.8%
	76.2%
Yield Std (%)	1.0%
	11.8%
Average (all)	78.27
	226.53
Std. Dev. (all)	3.73
	27.07
JB-2"	98.50
	281.00
Yield (min)*	75%
	68%
Yield (max)*	83%
	95%
Yield (average)	79%
	81%
STD ALL (average)	4%
	5%

[geocas.html](#) reports the F and C1 concentration of JB-2 as 98.5ppm and 281ppm, respectively.

Sample	Uncorrected for Yield				Corrected for Yield by Batch				ISD	ISD Propagated	ISD Propagated	
	F (ppm)	C1 (ppm)	F (%)	C1 (%)	F (ppm)	C1 (ppm)	F (%)	C1 (%)				
SEC42-06	91.94	8.85	22.5%	23.0%	114.56	11.04	14.77	1.81	130.0%	19.91%	14.95	
SEC43-01	108.51	7.05	49.0%	34.4%	136.83	8.85	68.01	2.93	49.75%	34.99%	68.07	
SEC43-03	152.25	5.37	21.1%	32.8%	189.25	6.73	10.02	2.04	5.66%	32.39%	10.71	
SEC44-01	154.37	9.06	30.2%	28.1%	193.23	11.34	51.97	2.80	26.91%	27.38%	52.00	
SEC44-02	160.25	16.06	21.1%	20.8%	199.28	19.94	12.35	0.52	6.27%	12.08%	12.41	
SEC47-01	111.60	7.10	21.1%	21.2%	138.78	8.84	9.02	1.04	6.57%	16.65%	9.11	
SEC50-01												1.47

Supplementary Table 10. Grain Boundary Calculations

Grain boundary calculations with sims spot size of 3.8µm*2																										
Width	Method 1				Dolomite																					
	Est Phi, GB	Sims Spot Size (µm)	Phi	Dolomite	5.00E-04	5.00E-05	5.00E-06	5.00E-07	5.00E-08	5.00E-09	5.00E-10	5.00E-11	5.00E-12	5.00E-13												
Increments of 10x																										
3µm	0.019	0.00131579	17.65	13414	1.00E-03	0.0038	0.00026319	17.65	67070	33.54	3.35	0.34	0.03	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00				
3.8 µm	14.44		6.5	4940	5.00E-03	0.019	0.00131579	17.65	13414	6.71	0.67	0.07	0.01	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00			
5µm x 3.8µm	0.019	0.00131579	17.65	13414	2.50E-02	0.095	0.00657989	17.65	2682.8	1.34	0.13	0.01	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00			
Increments													1.25E-03	0.475	0.0328947	17.65	538.56	0.27	0.03	0.00	0.00	0.00	0.00	0.00	0.00	
of 5x													6.25E-03	2.375	0.164474768	17.65	107.312	0.05	0.01	0.00	0.00	0.00	0.00	0.00	0.00	
of 5x													3.80E+00	14.44		1	17.65	17.65	0.01	0.00	0.00	0.00	0.00	0.00	0.00	0.00
of 5x													1.90E+01	72.2		5	17.65	3.53	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
of 5x													9.50E+01	361	25	17.65	0.706	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	
of 5x													4.75E+02	1805	125	17.65	0.1412	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	
of 5x													Grain Size µm	500	CL GB meas	17.65										
of 5x													CL GB contribution (%)	#REF!												
of 5x													Delta µm	Phi	GB Cl contribution µg/g											
of 5x													Method 2	1mm												
of 5x													Calc Phi	5mm												
of 5x													from GB	25mm												
of 5x													125mm	0.125	1.47E-07	0.00008										
of 5x													625mm	0.625	3.68E-06	0.00039										
of 5x													3.125mm	3.125	9.20E-05	0.00162										
of 5x													15.625mm	15.625	2.30E-03	0.00812										
of 5x													78.125mm	78.125	5.75E-02	0.04059										
of 5x													390.625mm	390.625	1.44E-00	0.20296										

696

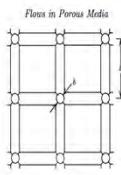


Figure 9.1 An idealized model of a porous medium. Circular tubes of diameter  $\delta$  and length  $b$  lie within the cube. The porosity is therefore given by

$$\phi = 3\left(\frac{\delta}{2}\right)^2 \frac{b}{\delta^3} = \frac{3\pi \delta^2}{4 b^3} \quad (9.6)$$

Turcotte and Schubert 2002