

1 **Revision 1**

2 **Melting in the Fe-FeO system to 204 GPa: Implications for oxygen in Earth's core**

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18 **ABSTRACT**

19 We performed melting experiments on Fe-O alloys up to 204 GPa and 3500 K in a diamond-
20 anvil cell (DAC) and determined the liquidus phase relations in the Fe-FeO system based on
21 textural and chemical characterizations of recovered samples. Liquid-liquid immiscibility was
22 observed up to 29 GPa. Oxygen concentration in eutectic liquid increased from >8 wt% O at 44
23 GPa to 13 wt% at 204 GPa, and is extrapolated to be about 15 wt% at the inner core boundary
24 (ICB) conditions. These results support O-rich liquid core, although oxygen cannot be a single

25 core light element. We estimated the range of possible liquid core compositions in Fe-O-Si-C-S
26 and found that the upper bounds for silicon and carbon concentrations are constrained by the
27 crystallization of dense inner core at the ICB.

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INTRODUCTION

30 Oxygen is one of the likely candidates for light elements in the Earth's core (Hirose et al. 2013;
31 Morard et al. 2017). This hypothesis is supported by recent first-principles calculations showing
32 that both the density and sound velocity observed in the liquid core are compatible with the
33 presence of oxygen (Badro et al. 2014). The earlier calculations by Alfè et al. (2002) also stated
34 that oxygen is an important alloying element in the core; it is mostly partitioned into liquid iron
35 and can thus account for the large density contrast observed across the ICB.

36 Liquidus phase relations, in particular eutectic composition, provide another clue to constrain
37 the outer core composition. Even if oxygen is likely to be an abundant core light element, the liquid
38 core should not crystallize as FeO at the ICB; FeO is indeed lighter and thus cannot form the denser
39 solid inner core. Previous experiments clearly demonstrated that oxygen concentration in the Fe-
40 FeO eutectic liquid increases with increasing pressure (Ohtani et al. 1984; Ringwood and
41 Hibberson 1990; Seagle et al. 2008). Komabayashi (2014) developed a thermodynamic model that
42 predicts ~12 wt% O in the eutectic liquid composition at 330 GPa corresponding to the pressure
43 at the ICB. The most recent study by Morard et al. (2017) carried out melting experiments on Fe-
44 FeO to 102 GPa. They reported a sharp and large increase in oxygen concentration in the eutectic
45 liquid around 80 GPa, which contradicts earlier experimental results and the thermodynamic model
46 (Seagle et al. 2008; Komabayashi 2014).

47 Previous experiments showed immiscible Fe-rich metallic liquid and FeO-rich ionic liquid

48 below 25 GPa (Ringwood and Hibberson 1990; Tsuno et al. 2007). Frost et al. (2010) found a
49 small immiscible two-liquid region at 25 GPa and 2473 K, and the immiscibility is no longer
50 observed at 50 GPa (Seagle et al. 2008). The disappearance of the two-liquid region from the Fe-
51 FeO liquidus diagram should change the eutectic liquid composition as illustrated in [Figure 1](#).

52 Here we report melting experiments on Fe-O alloys up to pressures of 204 GPa, which extends
53 the pressure range examined in earlier studies by a factor 2. On the basis of the eutectic liquid
54 composition determined in the core pressure range, we discuss the presence of oxygen and other
55 light elements in the outer core.

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EXPERIMENTAL METHODS

58 High-pressure and -temperature (P - T) conditions were generated in a laser-heated DAC using
59 diamonds with culet size of 120, 150, or 300 μm ([Table 1](#)). We employed homogeneous Fe-
60 4.1wt%O and Fe-13.0wt%O starting materials, prepared by magnetron sputtering of Fe in a
61 inductively coupled RF (radio-frequency) reactive oxygen-rich argon plasma (Morard et al. 2017).
62 A sample flake was loaded with an Al_2O_3 pressure medium, into a 40–140 μm hole at the center
63 of a pre-indented rhenium gasket. After loading the sample, the whole DAC was dried in a vacuum
64 oven at 423 K for at least 12 hrs to eliminate moisture from the sample. It was then compressed in
65 an argon atmosphere.

66 At each pressure of interest, we heated the sample from both sides with a couple of 100 W
67 single-mode Yb fiber lasers (YLR-100, IPG Photonics) with flat-top beam-shaping optics. A
68 visible laser-heated spot was approximately 20 μm in diameter. The heating duration was limited
69 to less than 3 sec, in order to avoid temperature fluctuation that leads to a complex melting texture.
70 It is long enough for oxygen to diffuse over an entire melt pocket (Helffrich 2014) ([Figs. 2a, b](#)),

71 which assures chemical equilibrium between liquid and solid since melting/crystallization occurs
72 almost instantaneously. Indeed, the time-series experiments performed by Mori et al. (2017)
73 demonstrated that the compositions of coexisting liquid and solid in the Fe-S system did not change
74 after 1 sec in a DAC. Note that the diffusivity of oxygen is slightly higher than that of sulfur in
75 liquid Fe (Helffrich 2014). The temperature profile was obtained using a spectro-radiometric
76 method (Mori et al. 2017) (Fig. 2c). Following Mori et al. (2017) and Tateno et al. (2018), we
77 determined the temperature at the liquid/solid boundary by combining the temperature profile with
78 a sample cross section. Pressure was measured at 300 K after heating based on a Raman shift of
79 diamond (anvil) (Akahama and Kawamura 2004) and then corrected for thermal pressure based
80 on Andrault et al. (1998). We considered 60% of theoretical thermal pressure for purely isochoric
81 heating $\Delta P = \alpha K_T \times (T - 300)$, $\alpha K_T = 4$ MPa/K below 100 GPa, and adopted 90% of that with
82 $\alpha K_T = 9$ MPa/K for runs at higher pressures (Hirose et al. 2017). These estimates differ from the
83 empirical one (5% pressure increase per temperature increase for every 1000 K) (Fiquet et al.
84 2010) by less than 7 GPa in all runs. The uncertainties in both P and T should therefore be less
85 than ± 5 –10 % (Table 1).

86 Textural and chemical characterizations were carried out on recovered DAC samples. A cross
87 section of the laser-heated portion of a sample was obtained parallel to the compression axis by
88 milling with a focused Ga ion beam (FIB, *FEI VersaTM 3D DualBeamTM*). It was then examined
89 by a field-emission (FE)-type scanning electron microscope (SEM) and energy dispersive X-ray
90 spectrometry (EDS) for elemental mapping (Fig. 2a). We also used an FE-type electron probe
91 micro-analyzer (FE-EPMA, JEOL JXA-8530F) with wavelength dispersive spectrometry (WDS)
92 for quantitative chemical analyses; an accelerating voltage of 12 kV, a current of 15 nA, and the
93 X-ray counting time of 20s/10s for peak/background. Fe, Fe₃C, and corundum were used as

94 standards. LDE1 (O), LIF (Fe), LDE2H (C), and TAP (Al) were analyzing crystals. The FE-EPMA
95 analyses of quenched liquid iron sometimes included a small amount of Al, which is likely a signal
96 from the surrounding pressure medium or Al₂O₃ grains that mechanically intruded into the liquid
97 during the experiment (Figs. 2a, b). Al was therefore subtracted as Al₂O₃ from raw data. Carbon
98 was detected not only in the quenched liquid, but also inside the rhenium gasket; the latter could
99 be due to contamination during FIB processing or FE-EPMA analysis. Thus, we subtracted 0.2–
100 0.4 wt% C from the raw analyses of liquids, which was found on the rhenium gasket in the same
101 cross section (or in an unheated portion of a sample when the gasket was not included). The liquids
102 should have included the remaining 0.8–3.5 wt% C (Table 1).

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RESULTS

105 We performed eight separate runs in a pressure range from 23 to 204 GPa (Table 1). In sample
106 cross sections, except for runs performed below 44 GPa, we found a chemically homogeneous area
107 with non-stoichiometric Fe-O composition at the center of a laser-heated area (Figs. 2a, b), which
108 likely represents a quenched liquid pool. There are layers of single-phase Fe or FeO (or both) next
109 to the liquid, which indicates which side of the eutectic the liquid was. A subsolidus part is found
110 outside such a liquidus phase layer.

111 In run #1 performed at 23 GPa and 2760 K, two large liquid pools were observed (Fig. 2d);
112 each one exhibits heterogeneous textures, similar to the ones reported in previous multi-anvil
113 experiments (e.g., Ringwood and Hibberson 1990; Tsuno et al. 2007). The two-liquid
114 immiscibility was also found in run #2 at 29 GPa, but not at pressures above 44 GPa in this study.

115 In runs #3–#6, the composition of a liquid represents the lower and upper bounds for oxygen
116 concentration in the eutectic liquid when coexisting with Fe and FeO, respectively (Table 1, Fig.

117 3). In runs #4 and #6, both Fe and FeO layers were found. As commonly observed in both multi-
118 anvil (e.g., Hirose and Fei 2002) and DAC experiments (e.g., Mori et al. 2017), a temperature
119 gradient in a sample induced phase segregation of each phase; in the present case, the layers of
120 liquid, FeO, and Fe were formed separately. Nevertheless, we estimate the difference in
121 temperature between the liquid/FeO boundary and the FeO/Fe boundary to be less than 50–60 K
122 considering the temperature gradient (Fig. 2c) (note that the FeO layer was about 1 μm wide as
123 shown in Fig. 2a). Since 50–60 K is less than the uncertainty in the present experimental
124 temperature determination, we considered that liquid, FeO, and Fe coexisted practically at the
125 same temperature. The liquid composition may therefore be the eutectic liquid when both Fe and
126 FeO layers were observed. The liquids obtained in runs #7 and #8 included remarkably less oxygen
127 than in the other runs; alternatively, they were enriched in carbon (Table 1). Except for these two
128 runs, the carbon amount was limited to ~ 1 wt%, with probably a limited effect on the Fe-FeO
129 diagram.

130 The experiments with limited carbon contamination show a change in the Fe-FeO eutectic
131 liquid composition with increasing pressure (Fig. 3). Above 44 ± 4 GPa, the eutectic liquid includes
132 >8 wt% O, which is a remarkably high amount compared to that found in previous multi-anvil
133 experiments below 16 GPa (Ringwood and Hibberson 1990; Tsuno et al. 2007), where liquid-
134 liquid immiscibility is observed. The oxygen concentration increases to about 13 wt% at the
135 maximum pressure of 204 ± 20 GPa explored here.

136 Since the composition of liquid obtained in the present study is close (or identical) to the
137 eutectic liquid composition at each pressure, the liquidus temperature determined as temperature
138 at the liquid/solid boundary should be close to (or represent) the eutectic temperature. Note that
139 the compositions of the present Fe-O-(C) liquids with limited carbon contamination do not

140 represent the Fe-O-C ternary eutectic compositions. Indeed iron carbides did not crystallize;
141 therefore their liquidus temperatures give the upper bounds for the ternary eutectic temperatures.
142 All data are in good agreement with the Fe-O eutectic melting curve reported by Morard et al.
143 (2017) on the basis of X-ray diffraction (XRD) observations (Fig. 4).

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DISCUSSION

146 The present results show that the eutectic liquid composition in the Fe-FeO binary evolves
147 from Fe + >8 wt% O at 44 ± 4 GPa to +13 wt% O at 204 ± 20 GPa (Fig. 3). This observation is in
148 reasonable agreement with the previous DAC experiments by Seagle et al. (2008) based on the
149 disappearance of Fe or FeO from the XRD pattern, and with the thermodynamic model proposed
150 by Komabayashi (2014) considering a non-ideal mixing based on Frost et al. (2010). Extrapolation
151 of our data to 330 GPa suggests that the Fe-FeO eutectic liquid contains about 15 wt% O at the
152 ICB.

153 We observed the liquid immiscibility to 29 ± 3 GPa and 2900 ± 290 K, while Frost et al. (2010)
154 found a narrow Fe-FeO immiscible two-liquids region at 25 GPa and 2473 K. These two conditions
155 are similar when considering errors in P and T for both the experiments. In addition, since the
156 textures of quenched immiscible liquids are quite heterogeneous (see Fig. 9 in Frost et al. 2010
157 and Fig. 2d in this study), a precise determination of their chemical compositions is difficult.
158 Indeed, the compositions of the immiscible two liquids reported by Frost and others include large
159 uncertainties, suggesting that the immiscible two-liquids region may exist to P - T conditions higher
160 than 25 GPa at 2473 K to a certain extent.

161 The vanishing of the liquid immiscibility changes the Fe-FeO liquidus curves. According to
162 previous experiments (Ohtani et al. 1984; Ringwood and Hibberson 1990), the Fe-FeO eutectic

163 liquid includes minor amounts of oxygen at low pressures (~2 wt% at 16 GPa). The immiscible
164 two-liquids region is found around Fe+11wt%O to >25 GPa at 2473 K according to Frost et al.
165 (2010). Using the Fe-13.0wt%O starting material, we observed two immiscible liquids at 23 and
166 29 GPa (Fig. 2c). A single liquid was formed at 44 GPa, suggesting the disappearance of the
167 immiscible two-liquids region from the phase diagram. As illustrated in Fig. 1, a reduction in the
168 immiscible two-liquids region with increasing pressure causes the Fe-FeO eutectic point to shift
169 to a more O-rich region. It explains the high (>8 wt%) oxygen concentration in the eutectic liquid
170 at 44±4 GPa (Fig. 3).

171 The closure of immiscible two-liquids region also indicates that mixing between liquids Fe and
172 FeO is non-ideal at low pressures and becomes closer to ideal above 44 GPa. It is supported by the
173 eutectic melting curve shown in Fig. 4 that is in broad agreement with the thermodynamic
174 prediction by Komabayashi (2014) considering the ideal mixing. Furthermore, the pressure–
175 eutectic liquid composition relation observed in the present experiments is consistent with the
176 behavior shown by the non-ideal mixing model below ~50 GPa, while it could be explained by the
177 ideal solution model at higher pressures (Komabayashi 2014) (Fig. 3). The ideal solution model
178 by Komabayashi shows the lower oxygen concentration in the Fe-FeO eutectic liquid than the
179 present experimental results at >50 GPa, but his thermodynamic prediction depends on the
180 uncertain melting temperature of FeO, which can be much lower than he assumed.

181 While these results suggest that the oxygen content in the eutectic liquid increases with
182 increasing pressure to ~40 GPa, Morard et al. (2017) demonstrated it rises sharply around 80 GPa
183 (Fig. 3). The liquid coexisting with both Fe and FeO at 39 GPa in run #7 was contaminated by 3.1
184 wt% C and included only 0.4 wt% O, much less than that in the Fe-FeO eutectic liquid at equivalent
185 pressure. It is possible that the incorporation of carbon expands the immiscible two-liquid region

186 to higher pressures, which could explain the difference between these two studies. Morard et al.
187 (2017) estimated 0.8–2.7 wt% C with ± 1.0 wt% uncertainty for liquids obtained below 80 GPa,
188 using Rietveld analyses of XRD patterns to find the amount of iron carbide formed in a quenched
189 liquid portion (after moderate re-heating after melting experiments). Since solid Fe can include 1
190 to 2 wt% C below 100 GPa (Walker et al. 2013; Fei and Brosh 2014), such estimates may be
191 underestimated. The observed low O content may be more relevant to the ternary Fe-O-C system
192 than for the binary Fe-FeO system.

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IMPLICATIONS

195 In order to form the dense inner core, the liquid core should crystallize Fe that is depleted in
196 light elements compared to the outer core liquid. It was once suggested that the inner core is
197 composed mainly of Fe₇C₃ (Chen et al. 2014; Prescher et al. 2015), but recent calculations
198 demonstrated that the density of Fe₇C₃ is too small for the inner core (Li et al. 2016). The present
199 experiments demonstrate that oxygen concentration in the Fe-FeO eutectic liquid increases (in
200 other words, the liquidus field of Fe expands) with increasing pressure (Fig. 3). It contrasts the
201 other binary iron alloy systems. The amounts of silicon, carbon, and sulfur in corresponding Fe-Si
202 (Ozawa et al. 2016), Fe-C (Lord et al. 2009; Mashino et al. 2019), and Fe-S binary eutectic liquids
203 (Kamada et al. 2012; Mori et al. 2017) decrease at higher pressures. At 330 GPa, these amounts
204 are less than those required to explain the outer core density deficit when assuming a single light
205 element. It indicates that none of silicon, carbon, and sulfur can be a predominant light component
206 in the core.

207 The present experiments allow us to determine the Fe-FeO eutectic liquid composition to 204
208 GPa, which is a much higher pressure compared to earlier experiments (Seagle et al. 2008; Morard

209 et al. 2017). Our extrapolation to 330 GPa indicates ~15 wt% O in the eutectic liquid at the ICB
210 conditions. If oxygen is assumed to be the sole light element in the core, the calculations by Badro
211 et al. (2014) argued that 5.4 wt% O is required to account for both the density and bulk sound
212 velocity in the outer core. If this amount (Fe + 5.4 wt% O) is retained, such a system should sit on
213 the Fe-rich side of the eutectic and thus crystallize solid Fe at the ICB, which is consistent with a
214 dense inner core. However, oxygen is hardly soluble in solid Fe (<0.1 wt% at ICB conditions - see
215 Alfè et al. 2002; Ozawa et al. 2010). And the inner core is less dense than pure iron to some extent
216 (Antonangeli and Ohtani, 2015), thus suggesting that oxygen cannot be the sole light element in
217 the core.

218 Even if oxygen is not a predominant light element, the core should include (or should have
219 included) a certain amount of oxygen (Hirose et al. 2017), according to recent high-temperature
220 core formation models (e.g., Siebert et al. 2013). The present estimate of the Fe-FeO eutectic liquid
221 composition at 330 GPa is thus important for modeling the compositional evolution of the Earth's
222 core and for constraining the present-day core composition from the phase diagrams of iron alloys.
223 Recently Tateno et al. (2018) examined the liquidus phase relations in the Fe-Si-S ternary system
224 and found that the liquidus field of solid Fe is approximated by a tie-line connecting the two
225 eutectic points in the Fe-Si and Fe-S binary systems. If this is also true in Fe alloy systems
226 containing more light elements, we can estimate the liquidus field of solid Fe in Fe-O-Si-C-S at
227 the ICB, considering each binary eutectic liquid composition that has been already reported in the
228 literature: 15 wt% O in Fe-O (this study), 1.5 wt% Si in Fe-Si (Ozawa et al. 2016), 3 wt% C in Fe-
229 C (Mashino et al. 2019), and 5 wt% S in Fe-S (Mori et al. 2017).

230 The present-day outer core composition should be within the liquidus field of solid Fe at 330
231 GPa. It provides an additional constraint to narrow down the possible range for the outer core

232 composition (Fig. 5). The first-principles calculations by Badro et al. (2014) showed the range of
233 liquid Fe-O-Si-C-S compositions that explain both the density and sound velocity of the outer core.
234 That range is reduced when we consider such additional conditions required to form a dense inner
235 core. While Badro and others suggested that the silicon content could be as high as 4.5 wt%, the
236 amount of silicon in the possible liquid core composition is found to be small, and constrained by
237 the Fe-Si eutectic liquid composition. The estimate of sulfur concentration is similar. The minor
238 silicon content in the present-day outer core is consistent with the limited Si+O solubility in molten
239 iron under the core *P-T* conditions (Hirose et al. 2017). In addition, Badro et al. (2014)
240 demonstrated that the carbon content can be as high as 3.9 wt%, but our new constraint limits it to
241 less than 2.5 wt%.

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REFERENCES CITED

251 Akahama, Y., and Kawamura, H. (2004) High-pressure Raman spectroscopy of diamond anvils to
252 250GPa: Method for pressure determination in the multimegabar pressure range. *Journal of*
253 *Applied Physics*, 96, 3748–3751.
254 Alfè, D., Gillan, M.J., and Price, G.D. (2002) Composition and temperature of the Earth's core

- 255 constrained by combining ab initio calculations and seismic data. *Earth and Planetary Science*
256 *Letters*, 195, 91–98.
- 257 Andrault, D., Fiquet, G., Itié, J.-P., Richet, P., Gillet, P., Häusermann, D., and Hanfland, M. (1998)
258 Thermal pressure in the laser-heated diamond-anvil cell: An X-ray diffraction study.
259 *European Journal of Mineralogy*, 10, 931–940.
- 260 Antonangeli, D., and Ohtani, E. (2015) Sound velocity of hcp-Fe at high pressure: Experimental
261 constraints, extrapolations and comparison with seismic models. *Progress in Earth and*
262 *Planetary Science*, 2:3.
- 263 Badro, J., Côte, A.S., and Brodholt, J.P. (2014) A seismologically consistent compositional model
264 of Earth's core. *Proceedings of the National Academy of Sciences USA*, 111, 7542–7545.
- 265 Chen, B., Li, Z., Zhang, D., Liu, J., Hu, M.Y., Zhao, J., Bi, W., Alp, E.E., Xiao, Y., Chow, P., and
266 others (2014) Hidden carbon in Earth's inner core revealed by shear softening in dense Fe₇C₃.
267 *Proceedings of the National Academy of Sciences*, 111, 17755–17758.
- 268 Fei, Y., and Brosh, E. (2014) Experimental study and thermodynamic calculations of phase
269 relations in the Fe–C system at high pressure. *Earth and Planetary Science Letters*, 408, 155–
270 162.
- 271 Fiquet, G., Auzende, A.L., Siebert, J., Corgne, A., Bureau, H., Ozawa, H., and Garbarino, G.
272 (2010) Melting of Peridotite to 140 Gigapascals. *Science*, 329, 1516–1518.
- 273 Frost, D.J., Asahara, Y., Rubie, D.C., Miyajima, N., Dubrovinsky, L.S., Holzapfel, C., Ohtani, E.,
274 Miyahara, M., and Sakai, T. (2010) Partitioning of oxygen between the Earth's mantle and
275 core. *Journal of Geophysical Research*, 115, B02202.
- 276 Helffrich, G. (2014) Outer core compositional layering and constraints on core liquid transport
277 properties. *Earth and Planetary Science Letters*, 391, 256–262.

- 278 Hirose, K., and Fei, Y. (2002) Subsolidus and melting phase relations of basaltic composition in
279 the uppermost lower mantle. *Geochimica et Cosmochimica Acta*, 66, 2099–2108.
- 280 Hirose, K., Labrosse, S., and Hernlund, J. (2013) Composition and state of the core. *Annual*
281 *Review of Earth and Planetary Sciences*, 41, 657–691.
- 282 Hirose, K., Morard, G., Sinmyo, R., Umemoto, K., Hernlund, J., Helffrich, G., and Labrosse, S.
283 (2017) Crystallization of silicon dioxide and compositional evolution of the Earth's core.
284 *Nature*, 543, 99–102.
- 285 Kamada, S., Ohtani, E., Terasaki, H., Sakai, T., Miyahara, M., Ohishi, Y., and Hirao, N. (2012)
286 Melting relationships in the Fe–Fe₃S system up to the outer core conditions. *Earth and*
287 *Planetary Science Letters*, 359/360, 26–33.
- 288 Komabayashi, T. (2014) Thermodynamics of melting relations in the system Fe–FeO at high
289 pressure: Implications for oxygen in the Earth's core. *Journal of Geophysical Research: Solid*
290 *Earth*, 119, 4164–4177.
- 291 Li, Y., Vočadlo, L., Brodholt, J., and Wood, I.G. (2016) Thermoelasticity of Fe₇C₃ under inner
292 core conditions. *Journal of Geophysical Research: Solid Earth*, 121, 5828–5837.
- 293 Lord, O.T., Walter, M.J., Dasgupta, R., Walker, D., and Clark, S.M. (2009) Melting in the Fe–C
294 system to 70 GPa. *Earth and Planetary Science Letters*, 284, 157–167.
- 295 Mashino, I., Miozzi, F., Hirose, K., Morard, G., and Sinmyo, R. (2019) Melting experiments on
296 the Fe–C binary system up to 255 GPa: Constraints on the carbon content in the Earth's core.
297 *Earth and Planetary Science Letters*, 515, 135–144.
- 298 Morard, G., Andrault, D., Antonangeli, D., Nakajima, Y., Auzende, A.L., Boulard, E., Cervera, S.,
299 Clark, A., Lord, O.T., Siebert, J., and others (2017) Fe–FeO and Fe–Fe₃C melting relations
300 at Earth's core–mantle boundary conditions: Implications for a volatile-rich or oxygen-rich

- 301 core. *Earth and Planetary Science Letters*, 473, 94–103.
- 302 Mori, Y., Ozawa, H., Hirose, K., Sinmyo, R., Tateno, S., Morard, G., and Ohishi, Y. (2017)
303 Melting experiments on Fe–Fe₃S system to 254 GPa. *Earth and Planetary Science Letters*,
304 464, 135–141.
- 305 Ohtani, E., Ringwood, A.E., and Hibberson, W. (1984) Composition of the core II. Effect of high
306 pressure on solubility of FeO in molten iron, *Earth and Planetary Science Letters*, 71, 94–103.
- 307 Ozawa, H., Hirose, K., Tateno, S., Sata, N., and Ohishi, Y. (2010) Phase transition boundary
308 between B1 and B8 structures of FeO up to 210 GPa. *Physics of the Earth and Planetary*
309 *Interiors*, 179, 157–163.
- 310 Ozawa, H., Hirose, K., Yonemitsu, K., and Ohishi, Y. (2016) High-pressure melting experiments
311 on Fe–Si alloys and implications for silicon as a light element in the core. *Earth and Planetary*
312 *Science Letters*, 456, 47–54.
- 313 Prescher, C., Dubrovinsky, L., Bykova, E., Kuppenko, I., Glazyrin, K., Kantor, A., McCammon, C.,
314 Mookherjee, M., Nakajima, Y., Miyajima, N., and others (2015) High Poisson's ratio of
315 Earth's inner core explained by carbon alloying. *Nature Geoscience*, 8, 220–223.
- 316 Ringwood, A.E., and Hibberson, W. (1990) The system Fe-FeO revisited. *Physics and Chemistry*
317 *of Minerals*, 17, 313–319.
- 318 Seagle, C.T., Heinz, D.L., Campbell, A.J., Prakapenka, V.B., and Wanless, S.T. (2008) Melting
319 and thermal expansion in the Fe-FeO system at high pressure. *Earth and Planetary Science*
320 *Letters*, 265, 655–665.
- 321 Siebert, J., Badro, J., Antonangeli, D., and Ryerson, F.J. (2013) Terrestrial accretion under
322 oxidizing conditions. *Science*, 339, 1194–1197.
- 323 Tanaka, T., and Nagasaka, T. (2006) Fundamentals of thermodynamics of phase diagrams. *Ferrum*,

- 324 11, 15–22 (in Japanese).
- 325 Tateno, S., Hirose, K., Sinmyo, R., Morard, G., Hirao, N., and Ohishi, Y. (2018) Melting
326 experiments on Fe–Si–S alloys to core pressures: Silicon in the core? *American Mineralogist*,
327 103, 742–748.
- 328 Tsuno, K., Ohtani, E., and Terasaki, H. (2007) Immiscible two-liquid regions in the Fe-O-S system
329 at high pressure: Implications for planetary cores. *Physics of the Earth and Planetary Interiors*,
330 160, 75–85.
- 331 Walker, D., Dasgupta, R., Li, J., Buono, A. (2013) Nonstoichiometry and growth of some Fe
332 carbides. *Contributions to Mineralogy and Petrology*, 166, 935–957.
- 333

334 **Figure captions**

335 **FIGURE 1.** Schematic illustration of the change in liquidus curve and eutectic point with
336 changing interaction energy in liquid after Tanaka and Nagasaka (2006).
337

338 **FIGURE 2.** Sample cross sections and temperature profile. **(a)** X-ray maps combining oxygen
339 (red) and aluminum (cyan) for a sample recovered from run #6 performed at 204 GPa and 3500 K.
340 Back-scattered electron image shows the liquid/solid boundary. **(b, c)** SEM image and the
341 corresponding temperature profile for run #5 at 132 GPa. The boundary between quenched liquid
342 and solid Fe is found on the basis of the difference in microtexture and the presence of Al₂O₃ grains
343 (black) that mechanically intruded from surroundings into a liquid. The temperature at the liquid-
344 solid boundary is estimated to be 3050 K, considering that the boundary was isothermal (the same
345 temperature for both sides). **(d)** The SEM image from run #1 showing immiscible two liquids at
346 23 GPa.
347

348 **FIGURE 3.** Oxygen concentration in liquid (red symbols). Normal and reverse triangles
349 indicate liquids coexisting with either one of solid Fe or FeO, respectively, giving the lower and
350 upper bounds for eutectic liquid composition. Diamonds show liquids coexisting with both Fe and
351 FeO that may represent the eutectic liquids. Closed circles are carbon-contaminated samples (>3
352 wt% C). Black diamonds represent eutectic compositions at 3 GPa (Ohtani et al. 1984) and 16 GPa
353 (Ringwood and Hibberson 1990). Red, blue, and black curves show the change in the Fe-FeO
354 eutectic liquid composition from this study (only with ~1 wt% carbon contamination), Morard et
355 al. (2017), and Seagle et al. (2008), respectively. Green and cyan curves are the thermodynamic
356 models assuming non-ideal and ideal solutions, respectively, by Komabayashi (2014).
357

358 **FIGURE 4.** Fe-FeO eutectic melting curve. Diamonds represent eutectic temperature obtained
359 at 3 GPa (Ohtani et al. 1984), 16 GPa (Ringwood and Hibberson 1990), and 48 and 204 GPa (this
360 study). Reverse triangles indicate the upper bounds for the eutectic temperature. Red and black
361 lines are the Fe-FeO eutectic melting curves obtained in this study and by Morard et al. (2017),
362 respectively. Small circles and squares are the previous experimental data by Seagle et al. (2008)
363 and Morard et al. (2017), respectively, showing the upper (filled) and lower bounds (open symbols)
364 for eutectic temperatures.
365

366 **FIGURE 5.** The range of possible Fe-O-Si-C-S liquid core compositions that crystallize solid
367 Fe at the ICB and account for the density and velocity of the outer core (Badro et al. 2014).
368

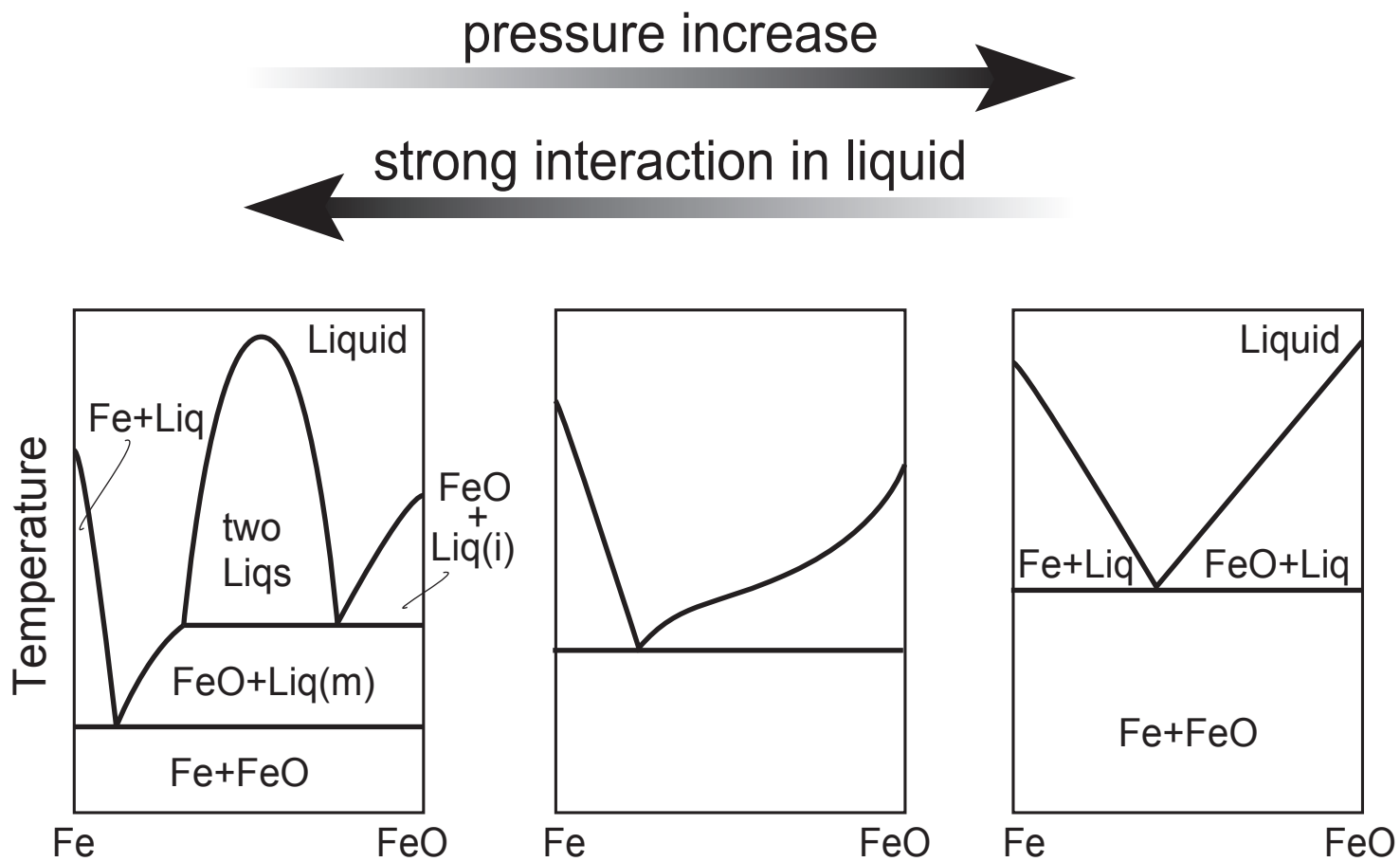


Fig. 1

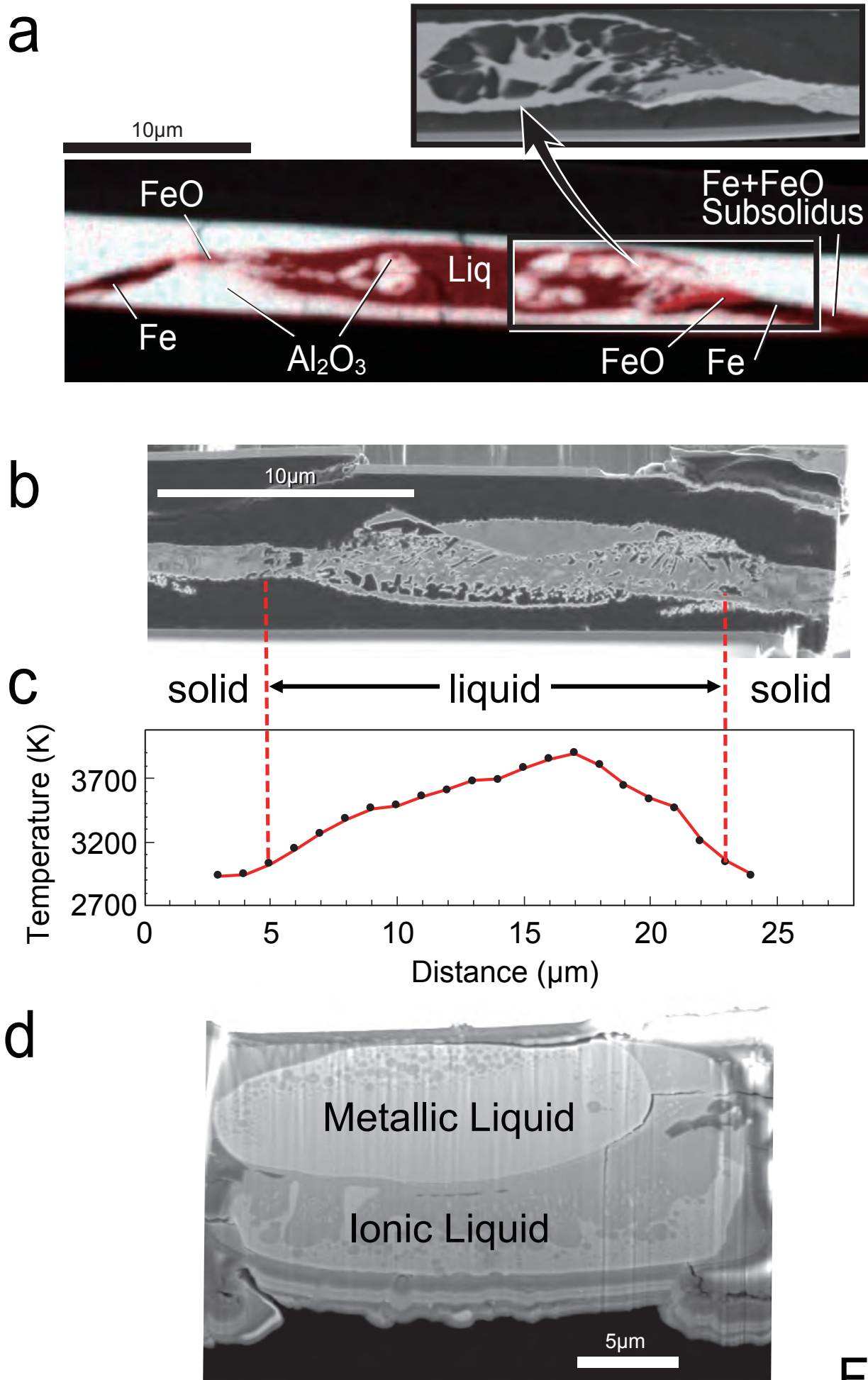


Fig. 2

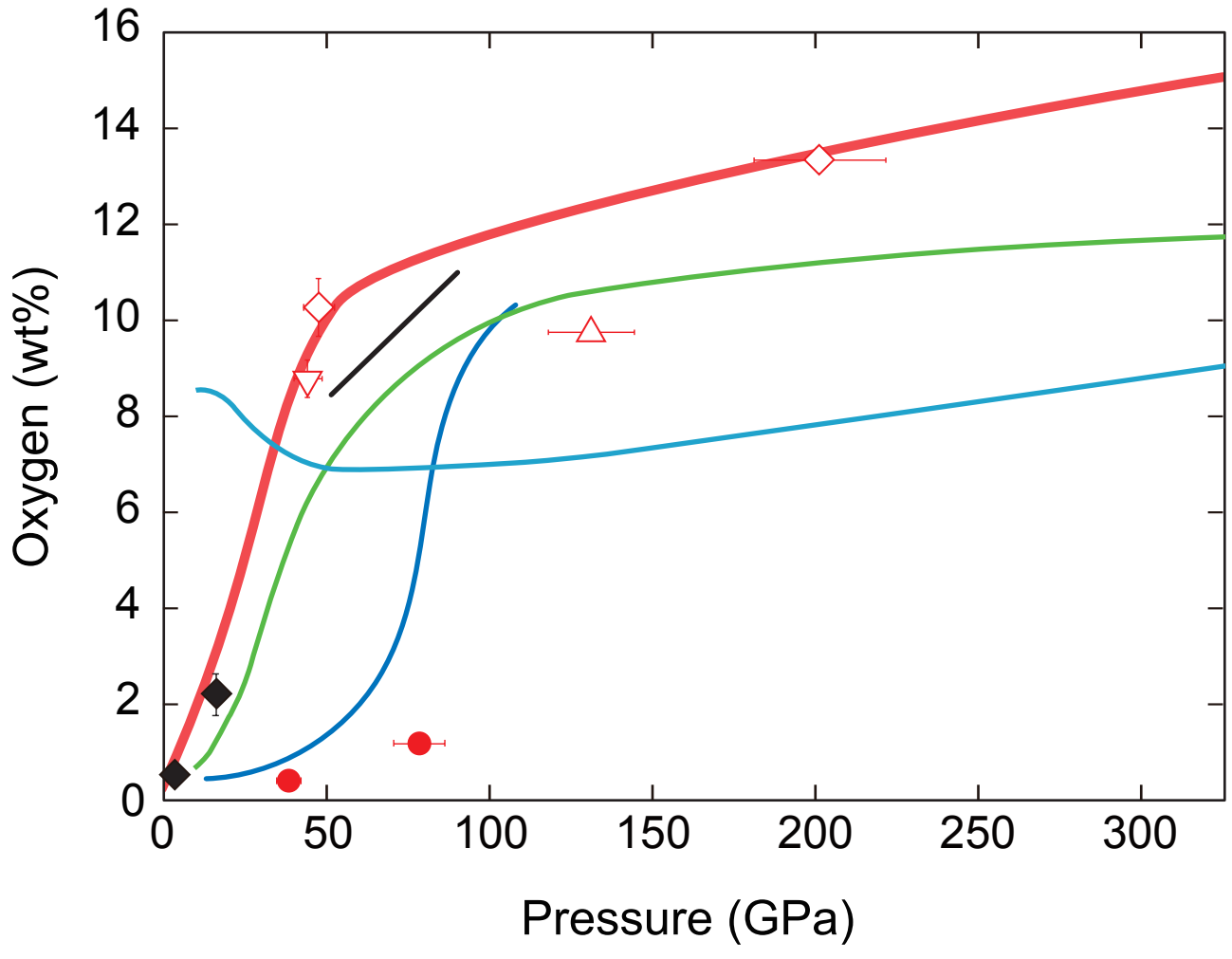


Fig. 2

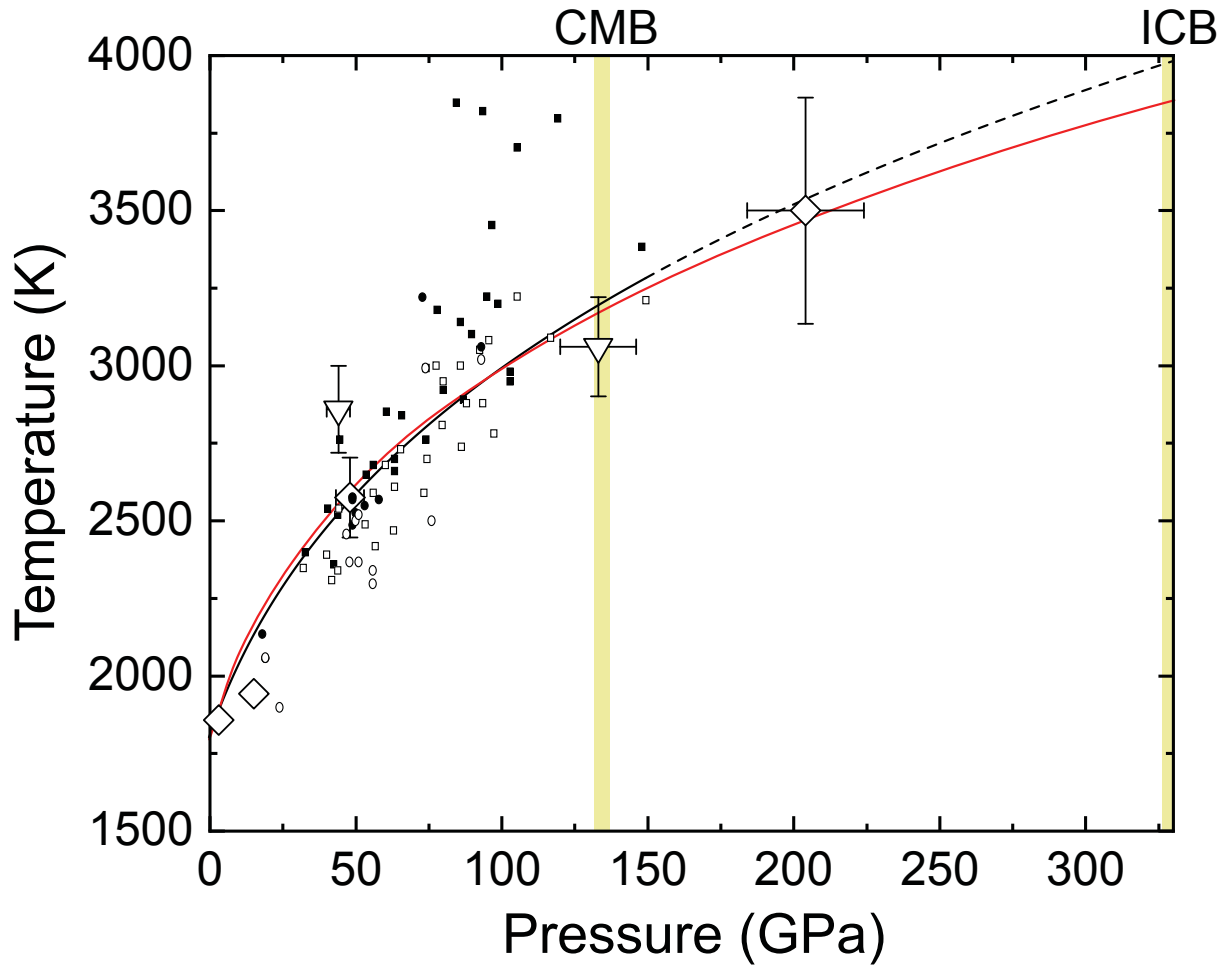


Fig. 4

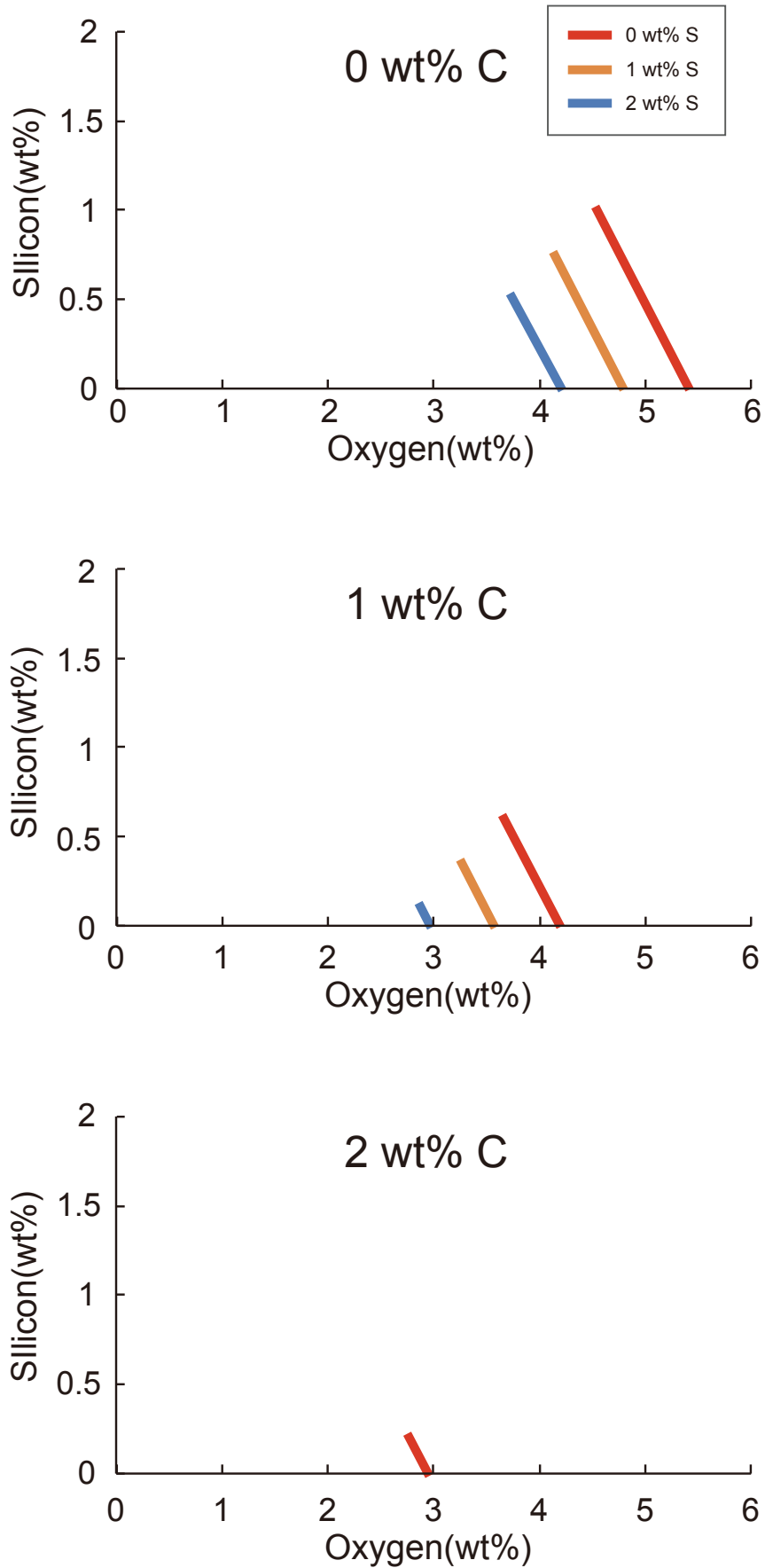


Fig. 5

TABLE 1. Experimental conditions and results

Run #	Starting material	P (GPa) at 300 K	P (GPa)	T (K)	O (wt%) in liquid	C (wt%) in liquid	Coexisting solid phase
#1	Fe-13.0wt%O	17	23(2)	2760(280)	two liquids		none
#2	Fe-13.0wt%O	23	29(3)	2900(290)	two liquids		FeO
#3	Fe-13.0wt%O	38	44(4)	2860(140)	8.8(4)	1.3(1)	FeO
#4	Fe-13.0wt%O	43	48(5)	2580(130)	10.3(6)	1.4(1)	Fe + FeO
#5	Fe-4.1wt%O	110	132(13)	3050(150)	9.8(1)	0.8(0)	Fe
#6	Fe-13.0wt%O	178	204(20)	3500(350)	13.3(0)	1.1(0)	Fe + FeO
#7	Fe-4.1wt%O	34	39(4)	2260(110)	0.4(0)	3.1(1)	Fe + FeO
#8	Fe-4.1wt%O	74	79(8)	2310(120)	1.0(0)	3.5(1)	FeO