## Highlights and Breakthroughs

## Iron carbide in the core

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Knowledge of the high-pressure behavior of candidate inner-core components is crucial to understanding the formation, evolution, and dynamics of our planet. In the October 2018 issue of *American Mineralogist*, Lai et al. present an experimental study on the thermoelastic properties of Fe<sub>7</sub>C<sub>3</sub>, a candidate component in the inner core, by single-crystal X-ray diffraction at high pressure and high temperature. Based on the anisotropic behavior of its compressibility and thermal expansivity along the crystallographic directions, they show that that high temperatures would have a significant influence on the elastic anisotropy of Fe<sub>7</sub>C<sub>3</sub> under high pressure.

Iron is the most abundant element by mass on Earth, while carbon is the fourth most abundant element in the Solar System. The most primitive C1 chondrites, that are believed to be the building blocks of our planet, contain 3.5 wt. % C. The Fe-C system has been proposed as a candidate component in the Earth's core, largely based on its high cosmochemical abundance, the frequent occurrence of iron carbide phases (e.g., Fe<sub>3</sub>C) in meteorites, and the high solubility of carbon in Fe-Ni liquids. Recently, Nakajima et al. (2009) observed that Fe<sub>3</sub>C melted incongruently above 5–10 GPa into liquid Fe-C alloy and a more carbon-rich carbide, Fe<sub>7</sub>C<sub>3</sub>. Therefore, Fe<sub>7</sub>C<sub>3</sub> appears to be more stable than Fe<sub>3</sub>C under conditions relevant to the Earth's core and could be a more promising deep-carbon carrier in the inner core (Fei and Brosh, 2014).

Knowledge of the high-pressure behavior and nature of Fe<sub>7</sub>C<sub>3</sub> is indispensable to understanding core composition and inner-core seismic anisotropy as well as the

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core's role in the global carbon cycle (Hazen and Schiffries, 2013; Deuss, 2014). In 30 the past few years, the crystal structure, equations of state, sound velocities, melting, 31 32 spin state, and magnetism of compounds that might inhabit the core have been 33 investigated at extreme conditions (e.g., Nakajima et al., 2011; Chen et al., 2012; Chen et al., 2014; Prescher et al., 2015; Liu et al., 2016a). Lord et al. (2009) 34 determined the melting temperatures of the Fe-C system up to 70 GPa, indicating that 35 Fe<sub>7</sub>C<sub>3</sub> could form a eutectic relation with Fe at core pressures. If the solubility of 36 37 carbon in iron is limited at the core conditions, Fe<sub>7</sub>C<sub>3</sub> might crystalize out of the early Earth's molten core and could be a constituent of the innermost inner core due to its 38 melting point being higher than pure iron (Liu et al., 2016b). Most importantly, the 39 40 high Poisson's ratio and anomalously low shear velocity of Fe<sub>7</sub>C<sub>3</sub> at high pressures 41 match that of the preliminary reference Earth model (PREM) for the inner core, which favors its potential presence in the inner core (Chen et al., 2014; Prescher et al., 2015). 42 The core might thus be the largest reservoir of the Earth's carbon (Chen et al., 2014). 43 See Chen and Li (2016) for excellent reviews of deep-carbon studies. 44 Lai et al. (2018) investigated the anisotropic thermoelastic properties of 45 single-crystal Fe<sub>7</sub>C<sub>3</sub> at high-pressure and high-temperature conditions by synchrotron 46 47 X-ray diffraction. The starting material Fe<sub>7</sub>C<sub>3</sub> were synthesized in a multi-anvil apparatus at 18 GPa and 1773 K. The run products Fe<sub>7</sub>C<sub>3</sub> were further analyzed in an 48 49 externally heated diamond anvil cell. Lai et al. (2018) revealed the high-temperature effects on the anisotropy of iron 50 carbide Fe<sub>7</sub>C<sub>3</sub>. Their diffraction data indicated that Fe<sub>7</sub>C<sub>3</sub> adopts an orthorhombic 51 structure under experimentally investigated conditions. They reported the linear 52 53 compressibility anisotropy of Fe<sub>7</sub>C<sub>3</sub> along the crystallographic axes, together with the thermal equation of state of Fe<sub>7</sub>C<sub>3</sub> up to 80 GPa and 800 K. In particular, high 54 temperatures affect compressibility and the magnitude of anisotropy during thermal 55 expansion. 56 57 Lai et al. (2018) elucidated that high temperatures must be taken into account when explaining the inner-core seismic anisotropy, supporting the existence of iron 58 carbide in the inner core, with further advances in our understanding of the role of 59

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carbon in the formation and evolution of our planet.

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