

1 Revision 2: MS# 5742

2 **High-pressure compressibility and vibrational properties of (Ca,Mn)CO₃**

3 Jin Liu¹, Razvan Caracas², Dawei Fan³, Ema Bobocioiu²,

4 Dongzhou Zhang⁴, Wendy L. Mao^{1,5}

5 ¹Department of Geological Sciences, Stanford University, Stanford, CA 94305, USA

6 ²CNRS, Laboratoire de Géologie de Lyon, Université Claude Bernard Lyon 1, 69342 Lyon
7 Cedex 07, France

8 ³Key Laboratory of High-temperature and High-pressure Study of the Earth's Interior,
9 Institute of Geochemistry, Chinese Academy of Sciences, Guiyang, Guizhou 550002, China

10 ⁴Hawai'i Institute of Geophysics and Planetology, University of Hawai'i at Manoa, Honolulu,
11 HI 96822, USA

12 ⁵Stanford Institute for Materials and Energy Sciences, SLAC National Accelerator
13 Laboratory, Menlo Park, CA 94025, USA

14

15 **Abstract**

16 Knowledge of potential carbon carriers such as carbonates is critical for our understanding of
17 the deep-carbon cycle and related geological processes within the planet. Here we
18 investigated the high-pressure behavior of (Ca,Mn)CO₃ up to 75 GPa by synchrotron single-
19 crystal X-ray diffraction, laser Raman spectroscopy, and theoretical calculations. MnCO₃-
20 rich carbonate underwent a structural phase transition from the CaCO₃-I structure into the
21 CaCO₃-VI structure at 45-48 GPa, while CaCO₃-rich carbonate transformed into CaCO₃-III
22 and CaCO₃-VI at approximately 2 and 15 GPa, respectively. The equation of state and
23 vibrational properties of MnCO₃-rich and CaCO₃-rich carbonates changed dramatically

24 across the phase transition. The CaCO₃-VI-structured CaCO₃-rich and MnCO₃-rich
25 carbonates were stable at room temperature up to at least 53 and 75 GPa, respectively. The
26 addition of smaller cations (e.g., Mn²⁺, Mg²⁺, and Fe²⁺) can enlarge the stability field of the
27 CaCO₃-I phase as well as increase the pressure for the structural transition into the CaCO₃-VI
28 phase.

29 **Keywords:** Carbonate, X-ray diffraction, Raman spectroscopy, High pressure

30

31 **1. Introduction**

32 Carbonates are the major, oxidized carbon-bearing phases in the Earth's crust and represent
33 an important carrier of carbon back into the mantle. The high-pressure behavior of carbonates
34 is thus essential for understanding deep-carbon cycling and storage in the Earth's deep
35 interior (Dasgupta and Hirschmann, 2010; Hazen and Schiffries, 2013). All naturally-
36 occurring carbonates in the crust can be potentially subducted into the deep mantle at
37 subduction zones. In particular, CaCO₃ is the dominant carbonate phase in the Earth's crust
38 and has been extensively investigated in static and dynamic high-pressure studies (e.g., Liu
39 and Mernagh, 1990; Williams et al., 1992; Biellmann et al., 1993; Gillet et al., 1993; Fiquet
40 et al., 1994; Suito et al., 2001; Ono et al., 2005; Oganov et al., 2006; Ono et al., 2007;
41 Merlini et al., 2012b; Shi et al., 2012; Ishizawa et al., 2013). CaCO₃ adopts a rhombohedral
42 *R*-3*c* structure at ambient conditions (referred to as CaCO₃-I) and undergoes a series of
43 structural phase transitions upon compression at room temperature into CaCO₃-II, -III, and -
44 VI approximately at 1.7, 2-3, and 15 GPa, respectively (Suito et al., 2001; Catalli and
45 Williams, 2005; Merlini et al., 2012b). CaCO₃-I, -II, and -III transform into aragonite at
46 elevated temperatures and then become disordered calcite when approaching the liquidus

47 (Suito et al., 2001), while the phase stability of calcite-VI is not well constrained at high
48 pressure and temperature (P - T).

49

50 At ambient conditions, there is a continuous solid solution between CaCO_3 and MnCO_3 .
51 MnCO_3 occurs in the CaCO_3 -I structure at ambient conditions and its high-pressure behavior
52 has been the subject of debate. The dissociation of MnCO_3 was observed at 6-8 GPa and
53 2273 K (Liu et al., 2001), while MnCO_3 was argued to be stable up to ~47 GPa and 2000 K
54 (Santillan and Williams, 2004) and then to transform into an orthorhombic phase at 50 GPa
55 above 1500 K (Ono, 2007). Furthermore, a series of transitions in MnCO_3 at room
56 temperature were observed at 15 and 50 GPa, respectively (Farfan et al., 2013). In the most
57 recent studies, MnCO_3 was observed to transform into the CaCO_3 -VI structure at room
58 temperature upon compression, although those studies reported a different phase transition
59 pressure of either ~35 GPa (Boulard et al., 2015) or 44 GPa (Merlini et al., 2015). Hereafter
60 the low-pressure (CaCO_3 -I) and high-pressure (CaCO_3 -VI) phases of MnCO_3 will be referred
61 to as MnCO_3 -I and MnCO_3 -II, respectively, after Merlini et al. (2015). Furthermore, the
62 compressibility and vibrational properties of MnCO_3 -II have not been determined. Here we
63 combine synchrotron single-crystal X-ray diffraction (XRD) and laser Raman spectroscopy,
64 together with theoretical calculations to study the high-pressure behavior of $(\text{Ca,Mn})\text{CO}_3$ up
65 to 75 GPa.

66

67 **2. Methods**

68 Natural calcite and rhodochrosite starting samples were obtained from Dabie Mountain,
69 China. Based on electron microprobe analyses, the composition of calcite was determined to

70 be $(\text{Ca}_{0.992}\text{Mn}_{0.008})\text{CO}_3$ with very minor amounts less than 0.1% of Mg and Fe, while the
71 rhodochrosite sample was $(\text{Ca}_{0.009}\text{Mn}_{0.980}\text{Fe}_{0.009}\text{Mg}_{0.002})\text{CO}_3$. For simplicity, we will herein
72 refer to these two samples as CaCO_3 and MnCO_3 , respectively. Both CaCO_3 and MnCO_3
73 starting samples exhibit a rhombohedral shape, reflecting their $R\text{-}3c$ structure.

74

75 **2.1. Single-crystal X-ray diffraction**

76 Single-crystal XRD analyses confirmed the CaCO_3 -I structure of the MnCO_3 starting sample
77 with lattice parameters $a = 4.7762(15)$ Å and $c = 15.6477(68)$ Å at ambient conditions. A
78 small platelet of single-crystal MnCO_3 of ~ 5 μm thick and ~ 40 μm in diameter was loaded
79 into a symmetric diamond anvil cell (DAC). A tungsten (W) gasket was pre-indented by a
80 pair of diamond anvils having 300 μm flat culets to a thickness of ~ 30 μm. A hole of 170 μm
81 in diameter was drilled in the pre-indented gasket and used as a sample chamber. A small
82 piece of Pt foil was for use as a pressure calibrant. The pressure-transmitting medium, Ne,
83 was loaded into the sample chamber at GSECARS of the Advanced Photon Source (APS),
84 Argonne National Laboratory (ANL). High-pressure single-crystal XRD experiments were
85 conducted at beamline 13-IDD of the APS, ANL, using a monochromatic X-ray beam with a
86 wavelength of 0.3344 Å which was focused down to ~ 5 μm in diameter at the sample
87 position. XRD patterns were recorded on a MarCCD detector where the DAC was rotated
88 from -15° to $+15^\circ$ about the X-ray beam direction with an interval of one degree. The tilting
89 and rotation of the MarCCD detector relative to the incident X-ray beam were calibrated
90 using lanthanum hexaboride (LaB_6) powder as the X-ray diffraction standard. The calibrated
91 distance between sample and MarCCD detector was 196.79 mm. XRD images were
92 processed and integrated using the Fit2D software for deriving the lattice parameters of the

93 pressure calibrant polycrystalline Pt (Hammersley et al. 1996). The same piece of sample
94 was used for ambient and high-pressure experiments. The different phases and lattice
95 parameters of MnCO_3 were extracted by the GSE_ADA/RSV software packages (Table S1)
96 (Dera et al., 2013). Pressure was calculated from the Pt pressure calibrant using a third-order
97 Birch-Murnaghan equation of state (3rd BM EoS) (Fei et al., 2007).

98

99 **2.2. Laser Raman spectroscopy**

100 High-pressure laser Raman spectra of CaCO_3 and MnCO_3 were collected using a Renishaw
101 RM1000 Raman microscope equipped with a 250 mm spectrometer focal length in the
102 Extreme Environments Laboratory (EEL) at Stanford University. The Raman system uses a
103 514.5 nm laser excitation line and has a spectral resolution of 4 cm^{-1} with the holographic
104 diffraction grating of 1800 lines/mm. The laser with a maximum power of 25 mW was
105 focused through a Nikon L Plan EPI 20X, 0.35 objective onto the sample with the laser spot
106 of approximately $2 \mu\text{m}$ in diameter. The Raman spectra were collected between -450 to
107 $+1600 \text{ cm}^{-1}$. The single-crystal samples of $\sim 10 \mu\text{m}$ thick and $\sim 40 \mu\text{m}$ in diameter were loaded
108 into the sample chamber of a DAC with the (101) crystal plane facing the incident laser beam,
109 together with a few ruby spheres as the pressure calibrant (Mao et al., 1986). Helium was
110 loaded into the sample chamber as the pressure-transmitting medium using the high-pressure
111 gas loading system in the EEL. Pressure and uncertainties were determined by averaging
112 over multiple measurements of ruby spheres before and after each Raman spectrum. We note
113 that the Raman spectra of the sample for XRD experiments were measured during
114 decompression, indicating that the high-pressure phase transition of MnCO_3 was reversible.
115 Raman spectra were processed using the software PeakFit v4.12 with the Voigt area method.

116

117 **2.3. Theoretical Methods**

118 Theoretical Raman spectra were computed at 0 K using the density-functional perturbation
119 theory (Baroni and Resta, 1986; Baroni et al., 2001; Gonze et al., 2005) in the ABINIT
120 implementation (Veithen et al., 2005; Gonze et al., 2009) with planewaves and
121 pseudopotentials. We used experimental input for the starting configurations of the crystal
122 structures for CaCO₃-I (Graf, 1961), CaCO₃-III (Pippinger et al., 2015) and CaCO₃-VI
123 (Merlini et al., 2012b). We used Troullier-Martins-type pseudopotentials to describe the core
124 electrons (Fuchs and Scheffler, 1999). We employed 12 electronic bands per CaCO₃ formula
125 unit. The first Brillouin zone of the reciprocal space was sampled using 8x8x8, 4x4x4 and
126 6x6x6 regular grids of **k**-points (Monkhorst and Pack, 1976) for respectively the three
127 structures. Together with a kinetic energy cutoff of 40 Ha (1Ha = 27.2116 eV) this converged
128 the energy at better than 1 mHa per unit cell. The structures were first relaxed at the target
129 pressure, i.e., the atoms are displaced and the unit cell shape and volume are allowed to
130 change under symmetry constraints until the forces are minimized and the stresses have only
131 hydrostatic components. The Raman spectra were calculated on the relaxed structures: the
132 position of the peaks comes from the quasi-harmonic approximation, and the reported
133 intensity of the peaks is the average of the Raman tensors over all possible crystal
134 orientations and laser polarizations, as for ideal powders. More details can be found in the
135 WURM project original reference (Caracas and Bobocioiu, 2011) and website
136 (<http://wurm.info>), including the spectra and the corresponding atomic displacement patterns.

137

138 **3. Results and Discussion**

139 We conducted two experimental runs and obtained the single-crystal XRD patterns for
140 MnCO_3 up to 67 GPa at 300 K. The XRD patterns demonstrated that the rhombohedral
141 MnCO_3 -I was stable up to 45 GPa at ambient temperature (Figs. 1-2). At higher pressures,
142 the XRD patterns of MnCO_3 changed (Fig. 1 inserts). The X-ray diffraction spots of MnCO_3
143 at 48-67 GPa could be well indexed to the crystal structure of CaCO_3 -VI (Merlini et al.,
144 2012a). Laser Raman spectra of CaCO_3 and MnCO_3 were collected with pressure steps of 1–
145 3 GPa up to 53 and 75 GPa, respectively (Figs. 3-4). Raman spectra for CaCO_3 changed
146 around 2 and 15 GPa, respectively, while those for MnCO_3 started changes around 44 and 46
147 GPa. For CaCO_3 , Raman spectra were assigned to CaCO_3 -I below 1.9 GPa, to CaCO_3 -III
148 between 1.9-15 GPa, and to CaCO_3 -VI above 15 GPa, respectively. For MnCO_3 , Raman
149 spectra were assigned to MnCO_3 -I below 44 GPa, and to MnCO_3 -II above 47 GPa, and to a
150 mixture of these two in between. Vibrational modes of CaCO_3 in the CaCO_3 -I, -III, and -VI
151 structure were theoretically calculated in order to better understand the observed Raman
152 bands at high pressures (Figs. 5-6; see Tables S2-S4 in Supplementary Materials for more
153 details).

154

155 **3.1. Equation of state and compressibility of MnCO_3**

156 The pressure-volume (P - V) curve between 0 and 44 GPa was fit to a 3rd BM EoS to derive
157 the elastic parameters of MnCO_3 -I, yielding the zero bulk modulus $K_0 = 122(3)$ GPa and the
158 pressure derivative of bulk modulus $K' = 3.7(2)$, or $K_0 = 113(3)$ GPa if K' was fixed at 4,
159 with the zero unit-cell volume $V_0 = 309.1(1)$ Å³ from XRD measurements at ambient
160 conditions (Table S1). The derived EoS parameters in this study are comparable to most
161 previous studies (Zhang and Reeder, 1999; Boulard et al., 2015; Merlini et al., 2015), but

162 slightly lower than those of MnCO_3 reported by Ono (2007). The discrepancy may be due to
163 the lack of a pressure medium and fewer data points in the study by Ono (2007).

164

165 The relative axial compressibilities, a/a_0 and c/c_0 , and the axial ratio, c/a , of MnCO_3 -I were
166 determined as a function of pressure up to 44 GPa (Fig. 2). The compressibility of MnCO_3 -I
167 is not isotropic, with the a axis about 3–4 times less compressible than the c axis which is
168 consistent with the results reported by Boulard et al. (2015). The anisotropic change in the a
169 and c axes causes the octahedral distortion to change from trigonally elongated to trigonally
170 compressed with increasing pressure (Lavina et al., 2010). The $(\text{Mg,Fe})\text{CO}_3$ solid-solution
171 series exhibits a similar anisotropic behavior in the a/a_0 and c/c_0 ratios at high pressures
172 (Lavina et al., 2010). The a/a_0 and c/c_0 values were approximately 0.96 and 0.86,
173 respectively, for MgCO_3 , FeCO_3 , and MnCO_3 at ~40 GPa (e.g., Zhang et al., 1998; Fiquet et
174 al., 2002; Litasov et al., 2008; Lavina et al., 2010; Lin et al., 2012). The comparable axis
175 compressibilities of MgCO_3 , FeCO_3 , and MnCO_3 are likely due to them having the same
176 structure.

177

178 MnCO_3 -I started to transform into MnCO_3 -II at 45 GPa and the phase transition was
179 completed at 48 GPa. A complex diffraction pattern was observed at 47 GPa (Fig. 1 insert),
180 which cannot be indexed solely by MnCO_3 -I or MnCO_3 -II. This pattern contained both
181 phases, indicating that MnCO_3 -I and MnCO_3 -II could coexist prior to the completion of
182 phase transition. The observed transition pressure of MnCO_3 in this study is a few GPa
183 higher than that of $(\text{Mn}_{0.96}\text{Ca}_{0.04})\text{CO}_3$ reported by Merlini et al. (2015), indicating that a small
184 amount of Ca^{2+} might effectively reduce the transition pressure in the $(\text{Ca,Mn})\text{CO}_3$ solid-

185 solution system. We note that Farfan et al. (2013) and Boulard et al. (2015) reported a
186 transition of MnCO_3 between 15 and 35 GPa using NaCl or silicon oil as the pressure-
187 transmitting medium and that this transition is not consistent with the CaCO_3 -II and -VI
188 structures and may be related to a distortion of the CaCO_3 -I structure due to nonhydrostatic
189 conditions, especially when NaCl undergoes a B1 to B2 phase transition about 25-30 GPa,
190 which involves a fairly large volume change. In contrast, both Merlini et al. (2015) and this
191 study used Ne and/or He as a pressure-transmitting medium in the single-crystal XRD
192 measurements and did not observe the distorted phase of the CaCO_3 -I structure between 15
193 and 30 GPa. Similarly, Ono (2007) did not observe this distorted phase likely because the
194 MnCO_3 sample was annealed to 1000-1500 K to relax any differential stress at high pressures.
195 In particular, Merlini et al. (2015) observed MnCO_3 in the CaCO_3 -VI structure at ~36 GPa
196 during decompression, while Boulard et al. (2015) observed the presence of diffraction peaks
197 other than from MnCO_3 -I above ~35 GPa in the power XRD patterns using NaCl as the
198 pressure-transmitting medium. To date, the transition pressure of MnCO_3 into the CaCO_3 -VI
199 structure has been reported ranging approximately from 35 to 50 GPa, which seems sensitive
200 to the stress field in the sample chambers. In addition, Santillan and Williams (2004) did not
201 observe any phase transition up to 50 GPa, likely because the incorporation of 16 mol% of
202 Fe^{2+} and Mg^{2+} in their MnCO_3 sample could greatly enlarge the stability field of the calcite
203 structure of MnCO_3 . The addition of Fe and Mg into MnCO_3 could thus greatly enlarge the
204 stability field of calcite structured MnCO_3 . This is similar to the idea presented by Shi et al.
205 (2012) that an increase in Mn concentration in CaCO_3 expanded the high-pressure stability
206 range of the different CaCO_3 structures.

207

208 The P - V curve of MnCO_3 -II was fit to a second-order BM EoS from 48 to 67 GPa, resulting
209 in $K_0 = 144(13)$ GPa and $V_0 = 285(7) \text{ \AA}^3$ with K' fixed at 4. The unit cell volume per formula
210 (V/Z) of MnCO_3 -II reported by Merlini et al. (2015) is close to the P - V curve extrapolated by
211 the current EoS of MnCO_3 -II. The bulk modulus of MnCO_3 increases approximately by 12%
212 across the phase transition from I to II. The density of MnCO_3 showed a dramatic increase by
213 $\sim 5.5\%$ from MnCO_3 -I to -II, in good agreement with Merlini et al. (2015), while the density
214 increased only by $\sim 2\%$ at 35 GPa reported by Boulard et al. (2015). We note that the V/Z at
215 62 GPa measured by Boulard et al. (2015) (corresponding to the red square in Fig. 1) is
216 approximately 2% larger than that from the current EoS of MnCO_3 -II. Such a large
217 discrepancy may suggest that the unit cell of MnCO_3 -II in the triclinic system may be
218 overestimated due to the overlap and insufficient diffraction lines in the powder XRD
219 patterns by Boulard et al. (2015).

220

221 **3.2. Vibrational properties of CaCO_3 and MnCO_3**

222 Four Raman modes were observed in the CaCO_3 -I phase of CaCO_3 and MnCO_3 , including
223 the translational lattice mode E_g (T), librational lattice mode E_g (L), CO_3 in-plane bend mode
224 E_g (ν_4), and CO_3 symmetric stretch mode A_{1g} (ν_1) consistent with previous studies (e.g.,
225 White, 1974; Shi et al., 2012; Farfan et al., 2013). The frequencies of the Raman modes of
226 CaCO_3 -I at ambient conditions and their pressure dependence (dv/dP) are consistent with
227 literature results (Table 1) (Liu and Mernagh, 1990; Gillet et al., 1993). Calculations show
228 the existence of five Raman peaks with intermediate to strong intensities (Fig. 5 and Table
229 S2). The highest frequency stretching mode E_g (ν_3) overlaps with the diamond Raman peak.
230 The frequencies of the Raman modes of MnCO_3 -I at ambient conditions are consistent with

231 previous work by Shi et al. (2012). The dv/dP of $MnCO_3$ -I is comparable to Farfan et al.
232 (2013), although their frequencies are systematically higher by $\sim 10\text{ cm}^{-1}$ possibly caused by
233 the systematic error due to the incorrect setting of the zero position (Table 2).

234

235 The combined XRD and Raman results on $CaCO_3$ and $MnCO_3$ from this study and Merlini et
236 al. (2012b) were used to derive the mode Grüneisen parameters (γ_i). The vibrational
237 frequency (ν) as a function of V or P represents the mode Grüneisen parameter for a phonon
238 mode following the relation (Born and Huang, 1954):

239
$$\gamma_i = -\frac{d \ln \nu}{d \ln V} = \frac{K_T}{\nu} \left(\frac{d\nu}{dP} \right),$$

240 where K_T is the isothermal bulk modulus. The mode Grüneisen parameters describe the effect
241 that changing the volume of a crystal lattice volume have on its vibrational properties. The γ_i
242 for two external modes (T and L) of $MnCO_3$ -I are approximately 2.37 and 1.99, respectively,
243 approximately twice that of $CaCO_3$ -I (Tables 1–2). The γ_i for the internal mode ν_1 of $CaCO_3$ -
244 I is about 0.39, comparable to that for $MnCO_3$ -I, $FeCO_3$ -I, and $MgCO_3$ -I (Liu and Mernagh,
245 1990; Gillet et al., 1993; Lin et al., 2012). The γ_i for the internal mode ν_4 of $CaCO_3$ -I is
246 approximately 30% less than that of $MnCO_3$ -I. The γ_i for ν_4 of $MnCO_3$ -I of approximately
247 0.33 agrees well with infrared measurements by Boulard et al. (2015), while Santillan and
248 Williams (2004) reported a much smaller γ_i value for the ν_4 of $MnCO_3$ -I likely due to a large
249 amount of Ca, Fe, and Mg (23 mol% in total) in their rhodochrosite sample. We note that
250 Santillan and Williams (2004) suggested that ion substitution has a negligible effect on IR
251 frequencies (Dubrawski et al., 1989). However, this may be true for the room pressure
252 frequencies and not necessarily true for pressure dependency of IR modes.

253

254 The calculations showed the existence of 75 Raman-active modes in CaCO₃-III (Table S3).
255 Most of these modes may overlap in the experimental spectrum. Experimentally we observed
256 19 and 16 Raman bands in CaCO₃-III and -VI, respectively. These Raman bands can be
257 assigned to five different Raman modes, including *T*, *L*, ν_4 , ν_1 , and ν_2 (the CO₃ out-of-plane
258 bend). The reduction in frequencies and the modification in the intensity of most of the
259 Raman bands were observed when CaCO₃-III transformed into CaCO₃-VI at 14-16 GPa,
260 likely a result of the differing orientation of the CO₃²⁻ group in the phases (Fig. 3). CaCO₃-III
261 contains non-coplanar CO₃²⁻ groups with a layered structure, while CaCO₃-VI is
262 characterized by coplanar CO₃²⁻ groups with a non-layered structure (Merlini et al., 2012b).
263 Williams et al. (1992) may have missed the phase transition from CaCO₃-III to -VI due to the
264 large pressure steps for their high-pressure infrared measurements. Fiquet et al. (1994) noted
265 that there were changes in the powder diffraction patterns of CaCO₃-III starting near ~8 GPa
266 which might be explained by a new calcite structure, but they could not resolve the new
267 structure and speculated that it could also be related to normal high-pressure features of
268 CaCO₃-III. Catalli and Williams (2005) first recognized this phase transition at ~15 GPa and
269 resolved the detailed changes in the carbonate ν_4 mode using infrared spectroscopy and finer
270 pressure steps. The lowest-frequency band of CaCO₃-III experimentally decreased in
271 frequency with increasing pressure (Fig. 6; Table 1). Raman bands in the same vibrational
272 mode of CaCO₃-III seem to have similar values of dv/dP , with <2 for *T*, approximately 4-6
273 for *L*, and about 1.3 for ν_4 . The new band of 866 cm⁻¹ at 1.9 GPa was assigned to the ν_2 of
274 CaCO₃-III, having a negative γ_i and dv/dP of approximately -0.3, which is also observed in
275 calculations (Fig. 6). It is in line with infrared measurements by Williams et al. (1992), which
276 is close to a zero pressure shift of the ν_2 of CaCO₃-III reported by Catalli and Williams

277 (2005). The negative dv/dP of ν_2 could be explained by the increase in the coupling of
278 neighboring CO_3^{2-} groups, while the positive dv/dP of ν_4 could be due to an increase in the
279 Mn-O bond strength (Kraft et al., 1991). The dv/dP for the ν_1 band decreases approximately
280 by 50% from CaCO_3 -I to -III.

281

282 Theoretical calculations predict there are 30 active vibrational bands in CaCO_3 -VI (Fig. 5).
283 14 of them may not be experimentally observed due to their weak intensities at high
284 pressures (see Table S4 for more details). The vibrational modes of high frequencies at 1400-
285 1650 cm^{-1} were not experimentally observed in this study because of their superposition with
286 the diamond Raman peak. The ν_1 of CaCO_3 -VI from theoretical calculations agree perfectly
287 with experiment observations (Fig. 6). Both suggested that the dv/dP for the ν_1 band
288 decreases approximately by 30% from CaCO_3 -III to -VI. Furthermore, one ν_2 band of
289 CaCO_3 -III diverged into two ν_2 bands of CaCO_3 -VI, while theoretical calculations only
290 predict one ν_2 band with a small positive dv/dP value close to zero. For CaCO_3 -VI, the lower-
291 frequency ν_2 band of 860 cm^{-1} experimentally had a negative dv/dP of approximately -0.5,
292 but the higher-frequency ν_2 band of 867 cm^{-1} had a positive dv/dP of approximately +0.3. We
293 note that Catalli and Williams (2005) observed one ν_2 peak of CaCO_3 -VI having an
294 unchanged position of near 870 cm^{-1} throughout infrared measurements up to 52 GPa. The
295 dv/dP values of CaCO_3 -VI for the T bands from experimental measurements are around 2,
296 for the L bands at 2-3.5, and for the ν_4 bands approximately at 1.2, which are comparable to
297 those from theoretical calculations.

298

299 Although they are isostructural, the number and relative intensity of the Raman bands
300 observed in MnCO₃-II are not the same as CaCO₃-VI (Figs. 3–4). The difference in the
301 Raman spectra of the CaCO₃-VI-structured CaCO₃ and MnCO₃ suggests that cations (e.g.,
302 Mn²⁺ and Ca²⁺) can have more dramatic effects on the vibrational modes in the high-pressure
303 CaCO₃-VI phase than the rhombohedral CaCO₃-I phase. All 13 Raman bands of MnCO₃-II
304 monotonically increased in frequency with increasing pressure (Fig. 7; Table 2). The dv/dP
305 of MnCO₃-II for the *T* bands are at 1-2 and for the *L* bands at 2-2.5. All the *T* and *L* Raman
306 bands have the value of γ_i at ~ 2 . Three Raman bands for MnCO₃-II in 700–850 cm⁻¹ display a
307 comparable dv/dP of ~ 0.55 and γ_i of ~ 0.28 , and were thus assigned to ν_4 . The ν_2 band at 873
308 cm⁻¹ only appears between 44 and 47 GPa, and likely becomes too weak after the completion
309 of the phase transformation to MnCO₃-II to be observed. The Raman spectra collected at 44–
310 47 GPa exhibit the peak splitting of Raman bands within the pressure range where we
311 observed coexistence of MnCO₃-I and -II according to X-ray diffraction patterns (Figs. 1
312 insert and 4). Two ν_1 Raman bands were observed at 46 GPa, likely with one coming from
313 MnCO₃-I and the other from MnCO₃-II. Farfan et al. (2013) observed such splitting of ν_1
314 from 48.2 to 55.2 GPa, indicating that MnCO₃-I and -II coexisted at a larger pressure range
315 due to the use of silicon oil as the pressure-transmitting medium. The dv/dP of ν_1 for
316 MnCO₃-II is about 1.72, in good agreement with the infrared measurements by Boulard et al.
317 (2015). We note that Raman peak splitting in MnCO₃-I at 44-47 GPa might suggest a subtle
318 distortion of the CaCO₃-I structure between MnCO₃-I and -II, which may be similar to the
319 change from the CaCO₃-III and CaCO₃-IIIb structures (Merlini et al., 2012b), although this
320 potential transition cannot be resolved from the present diffraction patterns in this study.

321

322 **4. Implications**

323 The phase stability of divalent metal carbonates (MCO_3) is significantly affected by
324 compositional variation in the metal cations (M^{2+}). Previous studies suggested the CaCO_3 -I
325 phase stability of the $(\text{Mn,Ca})\text{CO}_3$ solid solution was greatly enhanced by adding Mn^{2+} (Shi
326 et al., 2012). For the Mn-rich $(\text{Mn,Ca})\text{CO}_3$ compositions, the solid solution may directly
327 transform into the CaCO_3 -VI structure without going through CaCO_3 -II and -III at high
328 pressures (Fig. 8) (Boulard et al., 2015; Merlini et al., 2015). The phase transition pressure
329 for MnCO_3 increased by approximately 2-3 GPa with an additional 2 mol% Mn^{2+} and 1 mol%
330 Fe^{2+} replacing Ca^{2+} in $(\text{Mn}_{0.96}\text{Ca}_{0.04})\text{CO}_3$ (Merlini et al., 2015). The presence of 7 mol% Mg^{2+}
331 and 9 mol% Fe^{2+} in $(\text{Mn,Ca})\text{CO}_3$ could stabilize the CaCO_3 -I phase at ~50 GPa and 2000 K
332 (Santillan and Williams, 2004). Furthermore, MgCO_3 can stabilize the CaCO_3 -I structure up
333 to at least 110 GPa at 2000 K (Isshiki et al., 2004), while FeCO_3 undergoes an isostructural
334 spin transition around 45 GPa and remains in the CaCO_3 -I structure to above 100 GPa at
335 ambient temperature (e.g., Lavina et al., 2010; Farfan et al., 2012; Liu et al., 2014; Liu et al.,
336 2015). In $(\text{Mn,Ca})\text{CO}_3$, the radii of Ca^{2+} (100 pm) and Mn^{2+} (83 pm) are greater than that of
337 Mg^{2+} (72pm) and Fe^{2+} (78 and 61 pm for the high-spin and low-spin states, respectively)
338 (Shannon, 1976). The incorporation of a smaller cation into CaCO_3 increases the stability
339 field of CaCO_3 -I (Fig. 9).

340

341 The phase stability and high-pressure polymorph(s) of divalent metal carbonates (MCO_3)
342 under Earth's mantle conditions have been an area of active debate. The lower-mantle
343 oxygen fugacity may be at the Ni/NiO or Fe/FeO buffers making carbonates unstable in the
344 deep mantle (see review by Frost and McCammon (2008)). However, considering the

345 uncertainty in our knowledge of lower-mantle oxygen fugacity, carbonates may be
346 transported into the lower mantle by subducting slabs if the local oxygen fugacity is
347 favorable for their survival. One piece of evidence is the observation of carbonate-bearing
348 inclusions in diamonds which were potentially brought up to the Earth's surface from the
349 deep mantle (Berg, 1986; Wang et al., 1996; Walter et al., 2011). The high-pressure
350 polymorph CaCO₃-VI has a smaller V/Z than aragonite up to at least 40 GPa and may replace
351 aragonite in the Earth's mantle (Merlini et al., 2012b). The CaCO₃-VI-structured
352 (Ca,Mn)CO₃ is stable up to at least 75 GPa at room temperature according to Raman spectra
353 in this study. The incorporation of smaller cations (e.g., Mg²⁺, Fe²⁺, and Mn²⁺) can
354 significantly reduce the V/Z of CaCO₃-I and -VI, and thus may enhance the stability field of
355 CaCO₃-rich carbonate in the CaCO₃-I and -VI structures inside cold, subducting slabs in the
356 deep mantle that carried various limestone and carbonaceous sediments.

357

358 **Acknowledgments**

359 We acknowledge R. McCarty, X. Wu, P. Dera, and R. Jones for experimental assistance and
360 helpful discussion. W.L. Mao acknowledges support from the Geophysics Program at NSF
361 (EAR 1446969) and the Deep Carbon Observatory. R. Caracas acknowledges computational
362 support from eDARI under grant stl2816 and from PSMN of ENS Lyon, and financial
363 support from the PICS program of CNRS and the Deep Carbon Observatory. D.W. Fan
364 acknowledges financial support from the National Natural Science Foundation of China
365 (41374107), and the Youth Innovative Technology Talents Program of Institute of
366 Geochemistry, Chinese Academy of Sciences. Portions of this work were performed at
367 GeoSoilEnviroCARS (Sector 13) and HPCAT (Sector 16), Advanced Photon Source (APS),
368 Argonne National Laboratory. Use of the COMPRES-GSECARS gas loading system was
369 supported by COMPRES under NSF Cooperative Agreement EAR 11-57758 and by
370 GSECARS through NSF grant EAR-1128799 and DOE grant DE-FG02-94ER14466.
371 HPCAT operations are supported by DOE-NNSA under Award No. DE-NA0001974 and
372 DOE-BES under Award No. DE-FG02-99ER45775, with partial instrumentation funding by
373 NSF. The Advanced Photon Source, a U.S. Department of Energy (DOE) Office of Science
374 User Facility is operated for the DOE Office of Science by Argonne National Laboratory
375 under Contract No. DE-AC02-06CH11357. Data used in this study are available upon
376 request from J. Liu (Email: jinliu1@stanford.edu) and R. Caracas (Email:
377 razvan.caracas@ens-lyon.fr).

378

379 **References**

- 380 Baroni, S., de Gironcoli, S., Dal Corso, A., and Giannozzi, P. (2001) Phonons and related
381 crystal properties from density-functional perturbation theory. *Review of Modern Physics*,
382 73, 515-562.
- 383 Baroni, S., and Resta, R. (1986) Ab initio calculation of the low-frequency Raman cross
384 section in silicon. *Physical Review B*, 33, 5969-5971.
- 385 Berg, G.W. (1986) Evidence for carbonate in the mantle. *Nature*, 324(6092), 50-51.
- 386 Biellmann, C., Gillet, P., Guyot, F.o., Peyronneau, J., and Reynard, B. (1993) Experimental
387 evidence for carbonate stability in the Earth's lower mantle. *Earth and Planetary Science*
388 *Letters*, 118(1-4), 31-41.
- 389 Born, M., and Huang, K. (1954) *Dynamical theory of crystal lattices*. Oxford University
390 Press, U.K.
- 391 Boulard, E., Goncharov, A.F., Blanchard, M., and L. Mao, W. (2015) Pressure-induced phase
392 transition in MnCO₃ and its implications on the deep carbon cycle. *Journal of*
393 *Geophysical Research: Solid Earth*, 120(6), 4069-4079.
- 394 Caracas, R., and Bobocoiu, E. (2011) The WURM project - a freely available web-based
395 repository of computed physical data for minerals. *American Mineralogist*, 96, 437-444.
- 396 Catalli, K., and Williams, Q. (2005) A high-pressure phase transition of calcite-III. *American*
397 *Mineralogist*, 90(10), 1679-1682.
- 398 Dasgupta, R., and Hirschmann, M.M. (2010) The deep carbon cycle and melting in Earth's
399 interior. *Earth and Planetary Science Letters*, 298(1-2), 1-13.
- 400 Dera, P., Zhuravlev, K., Prakapenka, V., Rivers, M.L., Finkelstein, G.J., Grubor-Urosevic, O.,
401 Tschauner, O., Clark, S.M., and Downs, R.T. (2013) High pressure single-crystal micro
402 X-ray diffraction analysis with GSE_ADA/RSV software. *High Pressure Research*, 33(3),
403 466-484.
- 404 Dubrawski, J.V., Channon, A.-L., and Warne, S.S. (1989) Examination of the siderite-
405 magnesite mineral series by Fourier transform infrared spectroscopy. *American*
406 *Mineralogist*, 74, 187-190.
- 407 Farfan, G., Wang, S., Ma, H., Caracas, R., and Mao, W.L. (2012) Bonding and structural
408 changes in siderite at high pressure. *American Mineralogist*, 97(8-9), 1421-1426.
- 409 Farfan, G.A., Boulard, E., Wang, S., and Mao, W.L. (2013) Bonding and electronic changes
410 in rhodochrosite at high pressure. *American Mineralogist*, 98(10), 1817-1823.
- 411 Fei, Y., Ricolleau, A., Frank, M., Mibe, K., Shen, G., and Prakapenka, V. (2007) Toward an
412 internally consistent pressure scale. *Proceedings of the National Academy of Sciences*,
413 104(22), 9182-9186.
- 414 Fiquet, G., Guyot, F., and Itie, L.P. (1994) High-pressure X-ray diffraction study of
415 carbonates: MgCO₃, CaMg(CO₃)₂, and CaCO₃. *American Mineralogist*, 79, 15-23.
- 416 Fiquet, G., Guyot, F., Kunz, M., Matas, J., Andrault, D., and Hanfland, M. (2002) Structural
417 refinements of magnesite at very high pressure. *American Mineralogist*, 87(8-9), 1261-
418 1265.
- 419 Frost, D.J., and McCammon, C.A. (2008) The redox state of Earth's mantle. *Annual Review*
420 *of Earth and Planetary Sciences*, 36(1), 389-420.
- 421 Fuchs, M., and Scheffler, M. (1999) Ab initio pseudopotentials for electronic structure
422 calculations of poly-atomic systems using density-functional theory. *Computer Physics*
423 *Communications*, 119, 67-98.

- 424
425
426 Gillet, P., Biellmann, C., Reynard, B., and McMillan, P. (1993) Raman spectroscopic studies
427 of carbonates part I: High-pressure and high-temperature behaviour of calcite, magnesite,
428 dolomite and aragonite. *Physics and Chemistry of Minerals*, 20(1), 1-18.
- 429 Gonze, X., Amadon, B., Anglade, P.M., Beuken, J.M., Bottin, F., Boulanger, P., Bruneval, F.,
430 Caliste, D., Caracas, R., Côté, M., Deutsch, T., Genovese, L., Ghosez, P., Giantomassi,
431 M., Goedecker, S., Hamann, D.R., Hermet, P., Jollet, F., Jomard, G., Leroux, S., Mancini,
432 M., Mazevet, S., Oliveira, M.J.T., Onida, G., Pouillon, Y., Rangel, T., Rignanese, G.M.,
433 Sangalli, D., Shaltaf, R., Torrent, M., Verstraete, M.J., Zerah, G., and Zwanziger, J.W.
434 (2009) ABINIT: First-principles approach to material and nanosystem properties.
435 *Computer Physics Communications*, 180(12), 2582-2615.
- 436 Gonze, X., Rignanese, G.-M., and Caracas, R. (2005) First-principles studies of the lattice
437 dynamics of crystals, and related properties. *Zeitschrift für Kristallographie*, 220, 458-
438 472.
- 439 Graf, D. L. (1961) Crystallographic tables for the rhombohedral carbonates. *American*
440 *Mineralogist*, 46, 1283-1316.
- 441 Hazen, R.M., and Schiffries, C.M. (2013) Why deep carbon? *Reviews in Mineralogy and*
442 *Geochemistry*, 75(1), 1-6.
- 443 Ishizawa, N., Setoguchi, H., and Yanagisawa, K. (2013) Structural evolution of calcite at
444 high temperatures: Phase V unveiled. *Scientific Reports*, 3, 2832.
- 445 Isshiki, M., Irifune, T., Hirose, K., Ono, S., Ohishi, Y., Watanuki, T., Nishibori, E., Takata,
446 M., and Sakata, M. (2004) Stability of magnesite and its high-pressure form in the
447 lowermost mantle. *Nature*, 427(6969), 60-63.
- 448 Kraft, S., Knittle, E., and Williams, Q. (1991) Carbonate stability in the Earth's mantle: A
449 vibrational spectroscopic study of aragonite and dolomite at high pressures and
450 temperatures. *Journal of Geophysical Research*, 96(B11), 17997-18009.
- 451 Lavina, B., Dera, P., Downs, R.T., Yang, W., Sinogeikin, S., Meng, Y., Shen, G., and
452 Schiferl, D. (2010) Structure of siderite FeCO₃ to 56 GPa and hysteresis of its spin-
453 pairing transition. *Physical Review B*, 82(6), 064110.
- 454 Lin, J.-F., Liu, J., Jacobs, C., and Prakapenka, V.B. (2012) Vibrational and elastic properties
455 of ferromagnesite across the electronic spin-pairing transition of iron. *American*
456 *Mineralogist*, 97(4), 583-591.
- 457 Litasov, K.D., Fei, Y., Ohtani, E., Kuribayashi, T., and Funakoshi, K. (2008) Thermal
458 equation of state of magnesite to 32 GPa and 2073 K. *Physics of the Earth and Planetary*
459 *Interiors*, 168(3-4), 191-203.
- 460 Liu, J., Lin, J.-F., Mao, Z., and Prakapenka, V.B. (2014) Thermal equation of state and spin
461 transition of magnesiosiderite at high pressure and temperature. *American Mineralogist*,
462 99(1), 84-93.
- 463 Liu, J., Lin, J.-F., and Prakapenka, V.B. (2015) High-pressure orthorhombic ferromagnesite
464 as a potential deep-mantle carbon carrier. *Scientific Reports*, 5, 7640.
- 465 Liu, L.-g., Lin, C.-C., and Yang, Y.-J. (2001) Formation of diamond by decarbonation of
466 MnCO₃. *Solid State Communications*, 118(4), 195-198.
- 467 Liu, L.-G., and Mernagh, T.P. (1990) Phase transitions and Raman spectra of calcite at high
468 pressures and room temperature. *American Mineralogist*, 75(7-8), 801-806.

- 469 Mao, H.-K., Xu, J., and Bell, P.M. (1986) Calibration of the ruby pressure gauge to 800 kbar
470 under quasi-hydrostatic conditions. *Journal of Geophysical Research*, 91(B5), 4673-4676.
- 471 Merlini, M., Crichton, W.A., Hanfland, M., Gemmi, M., Müller, H., Kuppenko, I., and
472 Dubrovinsky, L. (2012a) Structures of dolomite at ultrahigh pressure and their influence
473 on the deep carbon cycle. *Proceedings of the National Academy of Sciences*, 109(34),
474 13509-13514.
- 475 Merlini, M., Hanfland, M., and Crichton, W.A. (2012b) CaCO₃-III and CaCO₃-VI, high-
476 pressure polymorphs of calcite: Possible host structures for carbon in the Earth's mantle.
477 *Earth and Planetary Science Letters*, 333–334, 265-271.
- 478 Merlini, M., Hanfland, M., and Gemmi, M. (2015) The MnCO₃-II high-pressure polymorph
479 of rhodocrosite. *American Mineralogist*, 100(11-12), 2625-2629.
- 480 Monkhorst, H.J., and Pack, J.D. (1976) Special points for Brillouin-zone integrations.
481 *Physical Review B*, 13(12), 5188-5192.
- 482 Oganov, A.R., Glass, C.W., and Ono, S. (2006) High-pressure phases of CaCO₃: Crystal
483 structure prediction and experiment. *Earth and Planetary Science Letters*, 241(1–2), 95-
484 103.
- 485 Ono, S. (2007) High-pressure phase transformation in MnCO₃: a synchrotron XRD study.
486 *Mineralogical Magazine*, 71(1), 105-111.
- 487 Ono, S., Kikegawa, T., and Ohishi, Y. (2007) High-pressure transition of CaCO₃. *American*
488 *Mineralogist*, 92(7), 1246-1249.
- 489 Ono, S., Kikegawa, T., Ohishi, Y., and Tsuchiya, J. (2005) Post-aragonite phase
490 transformation in CaCO₃ at 40 GPa. *American Mineralogist*, 90(4), 667-671.
- 491 Pippinger T., Miletich R., Merlini M., Lotti P., Schouwink P., Yagi T., Crichton W.A.,
492 Hanfland M. (2015) Puzzling calcite-III dimorphism: Crystallography, high-pressure
493 behavior, and pathway of single-crystal transitions. *Physics and Chemistry of Minerals*
494 42, 29-43.
- 495 Prescher, C., and Prakapenka, V.B. (2015) DIOPTAS: a program for reduction of two-
496 dimensional X-ray diffraction data and data exploration. *High Pressure Research*, 35(3),
497 223-230.
- 498 Santillan, J., and Williams, Q. (2004) A high-pressure infrared and X-ray study of FeCO₃ and
499 MnCO₃: comparison with CaMg(CO₃)₂-dolomite. *Physics of the Earth and Planetary*
500 *Interiors*, 143–144, 291-304.
- 501 Shannon, R.D. (1976) Revised effective ionic radii and systematic studies of interatomic
502 distances in halides and chalcogenides. *Acta Crystallographica Section A*, 32(5), 751-767.
- 503 Shi, W., Fleet, M.E., and Shieh, S.R. (2012) High-pressure phase transitions in Ca-Mn
504 carbonates (Ca,Mn)CO₃ studied by Raman spectroscopy. *American Mineralogist*, 97(5-
505 6), 999-1001.
- 506 Suito, K., Namba, J., Horikawa, T., Taniguchi, Y., Sakurai, N., Kobayashi, M., Onodera, A.,
507 Shimomura, O., and Kikegawa, T. (2001) Phase relations of CaCO₃ at high pressure and
508 high temperature. *American Mineralogist*, 86(9), 997-1002.
- 509 Veithen, M., Gonze, X., and Ghosez, P. (2005) Nonlinear optical susceptibilities, Raman
510 efficiencies, and electro-optic tensors from first-principles density functional perturbation
511 theory. *Physical Review B*, 71(12), 125107.
- 512 Walter, M.J., Kohn, S.C., Araujo, D., Bulanova, G.P., Smith, C.B., Gaillou, E., Wang, J.,
513 Steele, A., and Shirey, S.B. (2011) Deep mantle cycling of oceanic crust: evidence from
514 diamonds and their mineral inclusions. *Science*, 334(6052), 54-57.

-
- 515 Wang, A., Pasteris, J.D., Meyer, H.O.A., and Dele-Duboi, M.L. (1996) Magnesite-bearing
516 inclusion assemblage in natural diamond. *Earth and Planetary Science Letters*, 141(1–4),
517 293-306.
- 518 White, W.B. (1974) The carbonate minerals. In V.C. Farmer, Ed. *Infrared spectra of minerals*,
519 p. 227-284. Mineralogical Society of Great Britain & Ireland, London.
- 520 Williams, Q., Collerson, B., and Knittle, E. (1992) Vibrational spectra of magnesite (MgCO₃)
521 and calcite-III at high pressures. *American Mineralogist*, 77(11-12), 1158-1165.
- 522 Zhang, J., Martinez, I., Guyot, F., and Reeder, R.J. (1998) Effects of Mg-Fe²⁺ substitution in
523 calcite-structure carbonates: thermoelastic properties. *American Mineralogist*, 83(3-4),
524 280-287.
- 525 Zhang, J., and Reeder, R.J. (1999) Comparative compressibilities of calcite-structure
526 carbonates: Deviations from empirical relations. *American Mineralogist*, 84(5-6), 861-
527 870.
- 528

529 **Figure Captions**

530 **FIGURE 1.** The pressure-volume relations of MnCO₃-I and -II. Insert: X-ray diffraction
531 images collected during a rotation of 30° at 32, 47, and 67 GPa, respectively, showing the
532 sharp diffraction peaks of MnCO₃ at high pressures. The diffraction rings in the XRD
533 patterns were from the pressure medium Ne while the MnCO₃ single crystal remained across
534 the phase transition. The diffraction images were illustrated by the DIOPTAS program
535 (Prescher and Prakapenka, 2015). Circles and diamonds: the unit cell volume per formula
536 (V/Z) for MnCO₃-I and -II from this study; square and triangles: V/Z of MnCO₃-II by
537 Boulard et al. (2015) and Merlini et al. (2015), respectively; lines: BM EoS fits to the data.
538 Error bars smaller than symbols are not shown for clarity. See Table S1 in Supplementary
539 Materials for more details.

540

541 **FIGURE 2.** The relative axial compressibilities of the rhombohedral phase of MnCO₃-I as a
542 function of pressure. Circles, squares, diamonds, and triangles: c/a , a/a_0 , c/c_0 , and V/V_0 ,
543 respectively. Error bars smaller than symbols are not shown for clarity.

544

545 **FIGURE 3.** Representative Raman spectra of CaCO₃ at high pressures and 300 K. The
546 metastable CaCO₃-II phase was not observed likely due to the small pressure range over
547 which this phase is stable and the relatively large pressure steps. Raman modes of CaCO₃-I
548 are labeled according to White (1974). T : translational lattice mode; L : librational lattice
549 mode; ν_4 : CO₃ in-plane bend; ν_2 : CO₃ out-of-plane bend; ν_1 : CO₃ symmetric stretch.

550

551 **FIGURE 4.** Representative Raman spectra of MnCO_3 at high pressures. MnCO_3 transformed
552 from the low-pressure phase (MnCO_3 -I) to the high-pressure phase (MnCO_3 -II) at 44-47 GPa.
553 Raman modes of MnCO_3 -I are labeled accordingly as T , L , ν_4 , and ν_1 . The splitting of Raman
554 bands of MnCO_3 at 46 GPa corresponds to the coexisting phases of I and II revealed by XRD
555 measurements at 47 GPa (Fig. 1 insert) with the ν_2 peak possibly at $\sim 875 \text{ cm}^{-1}$.

556

557 **FIGURE 5.** Vibrational modes of CaCO_3 -I and VI as a function of frequency at high
558 pressures by theoretical calculations. The intensity was normalized using the intensity of ν_1
559 as the reference. The blue diamonds indicate that the calculated intensity of modes are
560 extremely weak (see Tables S2 and S4 in Supplementary Materials for more details).

561

562 **FIGURE 6.** Raman shifts of CaCO_3 at high pressures. Black, blue, and red symbols:
563 experimental observations of CaCO_3 -I, -III, and -VI, respectively; olive diamonds, triangles,
564 and squares: theoretical calculations of CaCO_3 -I, -III, and -VI, respectively (this study). The
565 calculated modes with the extremely weak intensities (blue diamonds in Fig. 5) are not
566 shown for clarity. The calculated 75 Raman peaks of CaCO_3 -III were selectively plotted for
567 comparison. Two vertical dotted lines approximately display the phase transition pressures.
568 The solid lines represent the linear fits to the calculations. Error bars are typically smaller
569 than symbols. See Tables S2-S5 in Supplementary Materials for more details.

570

571 **FIGURE 7.** Raman shifts of MnCO_3 at high pressures. Solid, half-filled, and open symbols:
572 MnCO_3 -I, the coexisting phases of I and II, and MnCO_3 -II, respectively; blue shaded area:
573 the coexisting phases of I and II; solid lines: linear fits to experimental data. Error bars are

574 typically smaller than symbols and are not shown for clarity. See Table S6 in Supplementary
575 Materials for more details.

576
577 **FIGURE 8.** Schematic phase relations of the (Mn,Ca)CO₃ solid-solution series at 300 K.
578 Dashed lines represent the tentative phase boundaries. Solid triangles: this study; diamonds:
579 Merlini et al. (2012, 2015); squares and circles: Shi et al. (2012).

580

581 **FIGURE 9.** Schematic phase relations of the (Ca,Mn,Fe,Mg)CO₃ solid-solution system
582 under high pressures at 300 K. The effective ionic radii of Ca²⁺, Mn²⁺, Fe²⁺, and Mg²⁺ at
583 ambient conditions are illustrated accordingly from Shannon (1976). *hs*: the high-spin state
584 of Fe²⁺; *ls*: the low-spin state of Fe²⁺.

Table 1. Experimental vibrational parameters of CaCO₃ at high pressures

CaCO ₃	Raman mode	Initial frequency* (cm ⁻¹)	dv/dP (cm ⁻¹ /GPa)	Mode Grüneisen parameter (γ _i)
I	T	156	2.52(8)	1.14(4)
I	L	282	5.19(17)	1.33(6)
I	v ₄	713	2.35(7)	0.23(1)
I	v ₁	1086	5.96(15)	0.39(2)
III	T	131	-0.35(6)	-0.34(7)
III	T	140	1.98(10)	1.66(9)
III	T	161	0.76(11)	0.56(9)
III	T	168	0.36(6)	0.25(5)
III	T	202	0.56(4)	0.35(4)
III	T	206	0.89(13)	0.52(9)
III	T	219	1.91(21)	1.04(13)
III	L	242	3.78(18)	1.78(11)
III	L	256	4.41(29)	1.90(15)
III	L	266	4.32(17)	1.82(9)
III	L	275	6.31(26)	2.46(12)
III	L	296	5.76(25)	2.17(11)
III	L	312	6.25(33)	2.21(17)
III	L	330	6.28(32)	2.08(15)
III	v ₄	689	1.35(8)	0.23(2)
III	v ₄	731	1.31(17)	0.22(3)
III	v ₄	739	1.33(17)	0.22(3)
III	v ₂	866	-0.29(3)	-0.04(1)
III	v ₁	1102	2.78(14)	0.31(3)
VI	T	150	0.01(1)	0.01(1)
VI	T	172	1.37(22)	1.56(30)
VI	T	180	2.87(45)	3.03(57)
VI	T	195	2.17(49)	2.22(55)
VI	T	245	2.06(21)	1.81(20)
VI	L	284	2.04(26)	1.54(21)
VI	L	329	2.21(57)	1.48(46)
VI	L	343	1.98(62)	1.29(45)
VI	L	356	3.08(33)	1.83(24)
VI	L	382	2.82(29)	1.62(20)
VI	L	391	3.45(21)	1.87(15)
VI	v ₄	732	1.19(9)	0.39(5)
VI	v ₄	746	1.19(9)	0.39(5)
VI	v ₂	860	-0.48(7)	-0.14(3)
VI	v ₂	867	0.33(7)	0.09(2)
VI	v ₁	1134	2.04(12)	0.44(3)

*Initial frequencies of Raman modes for CaCO₃-I, -III, and -VI are at 0, 1.9, and 16 GPa, respectively.

Table 2. Experimental vibrational parameters of MnCO₃ at high pressures

MnCO ₃	Raman mode	Initial frequency* (cm ⁻¹)	dv/dP (cm ⁻¹ /GPa)	Mode Grüneisen parameter (γ_i)
I	T	184	2.91(5)	2.37(7)
I	L	290	3.68(9)	1.99(8)
I	ν_4	719	1.28(6)	0.33(3)
I	ν_1	1086	2.17(7)	0.38(3)
II	T	165	0.89(6)	1.93(17)
II	T	203	1.09(8)	1.91(19)
II	T	220	1.33(10)	2.15(18)
II	T	244	1.01(8)	1.50(14)
II	T	324	1.83(5)	2.02(8)
II	L	391	2.05(11)	1.87(11)
II	L	442	2.12(21)	1.72(19)
II	L	481	2.70(16)	2.01(19)
II	L	503	2.48(19)	1.78(22)
II	ν_4	735	0.51(4)	0.26(4)
II	ν_4	782	0.58(10)	0.28(6)
II	ν_4	826	0.63(8)	0.29(5)
II	ν_1	1202	1.72(8)	0.54(9)

*Initial frequencies of Raman modes for MnCO₃-I and -II are at 0 and 48 GPa, respectively.

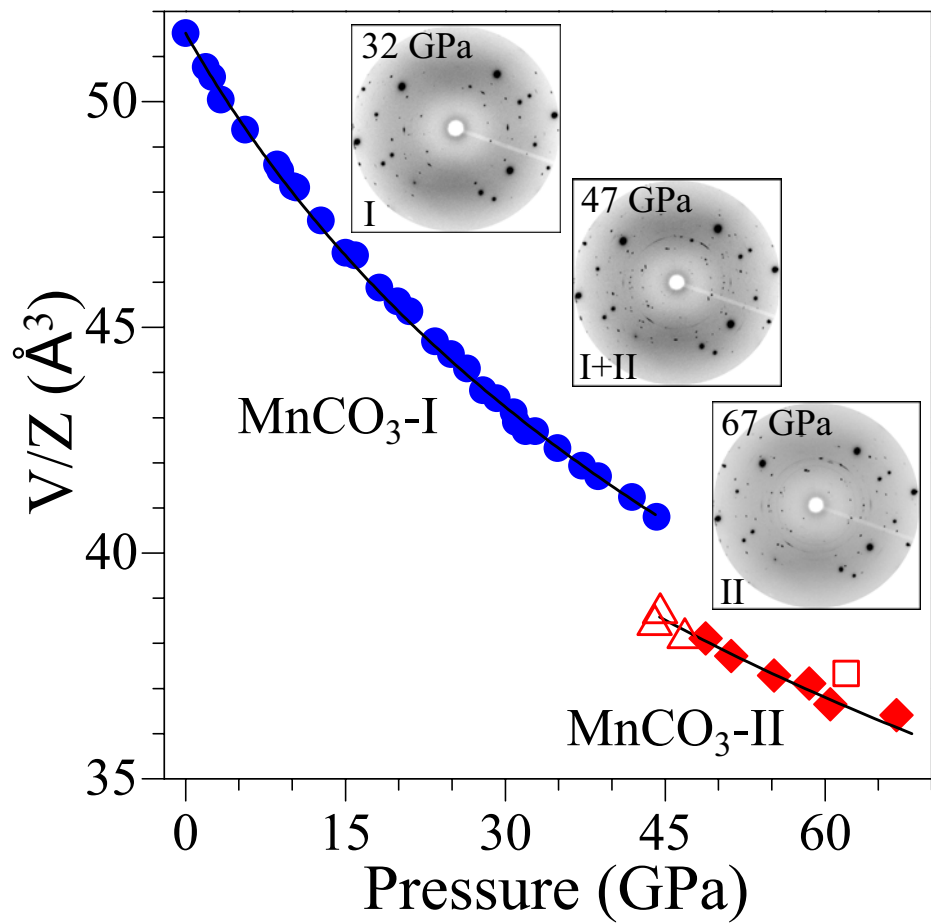


Figure 1

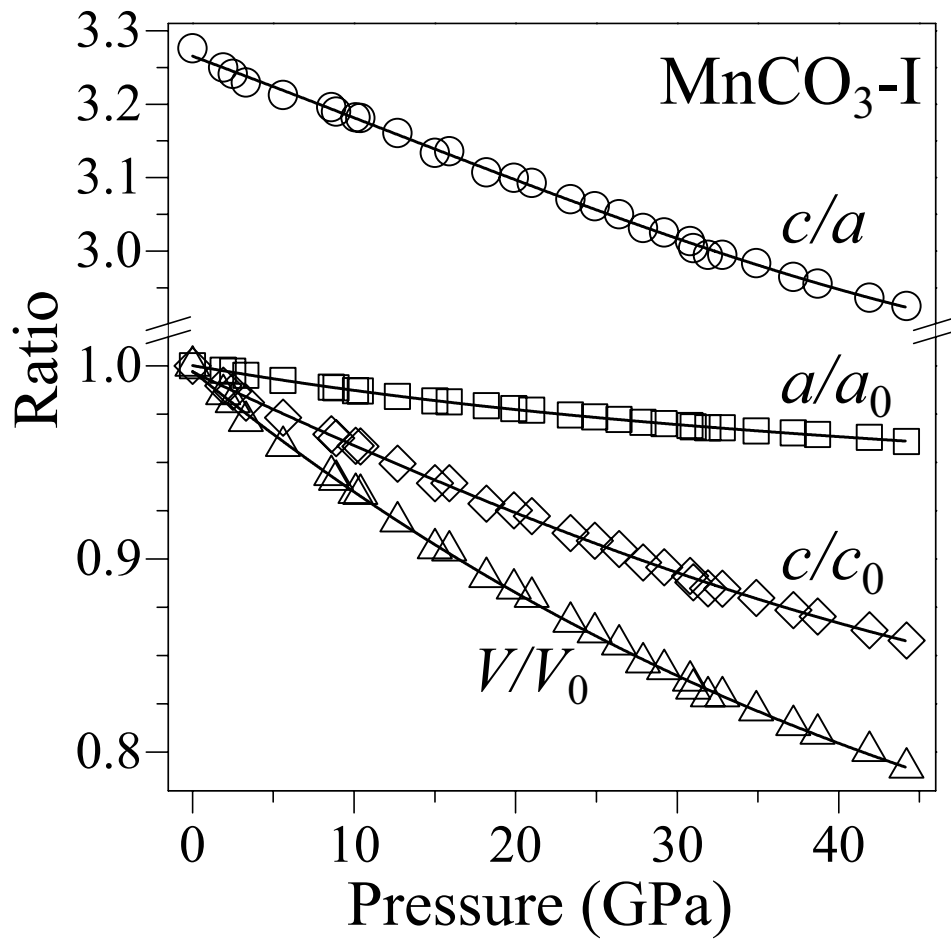


Figure 2

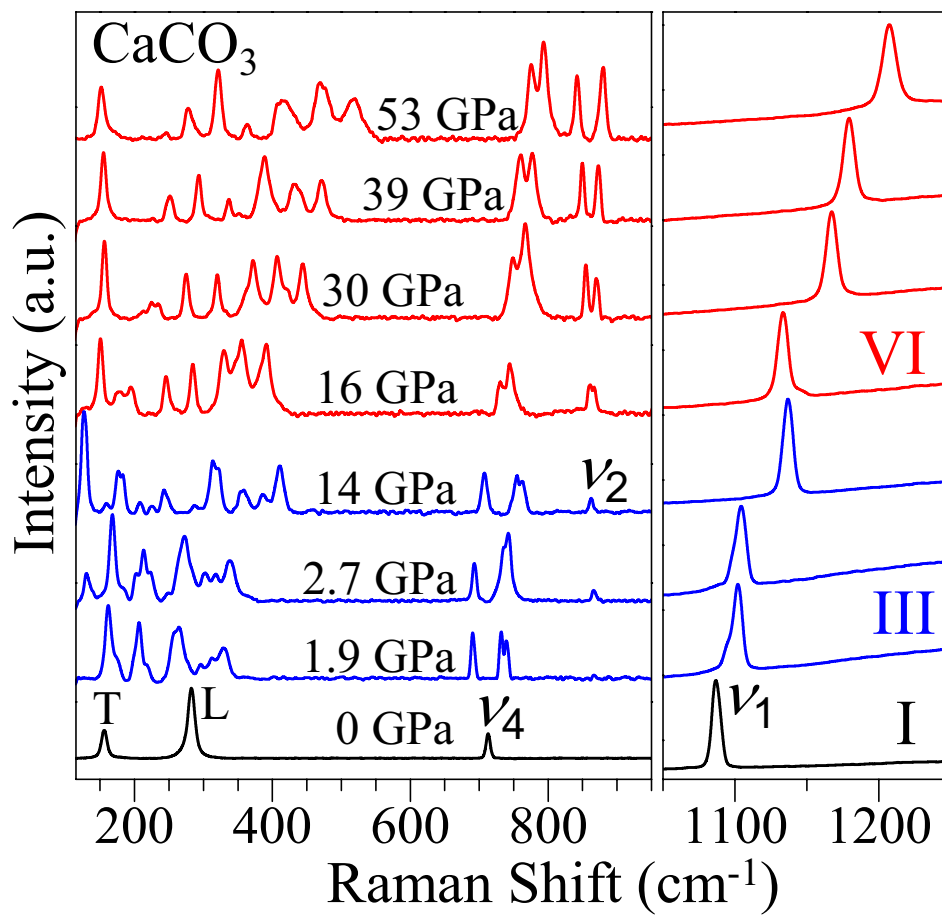


Figure 3

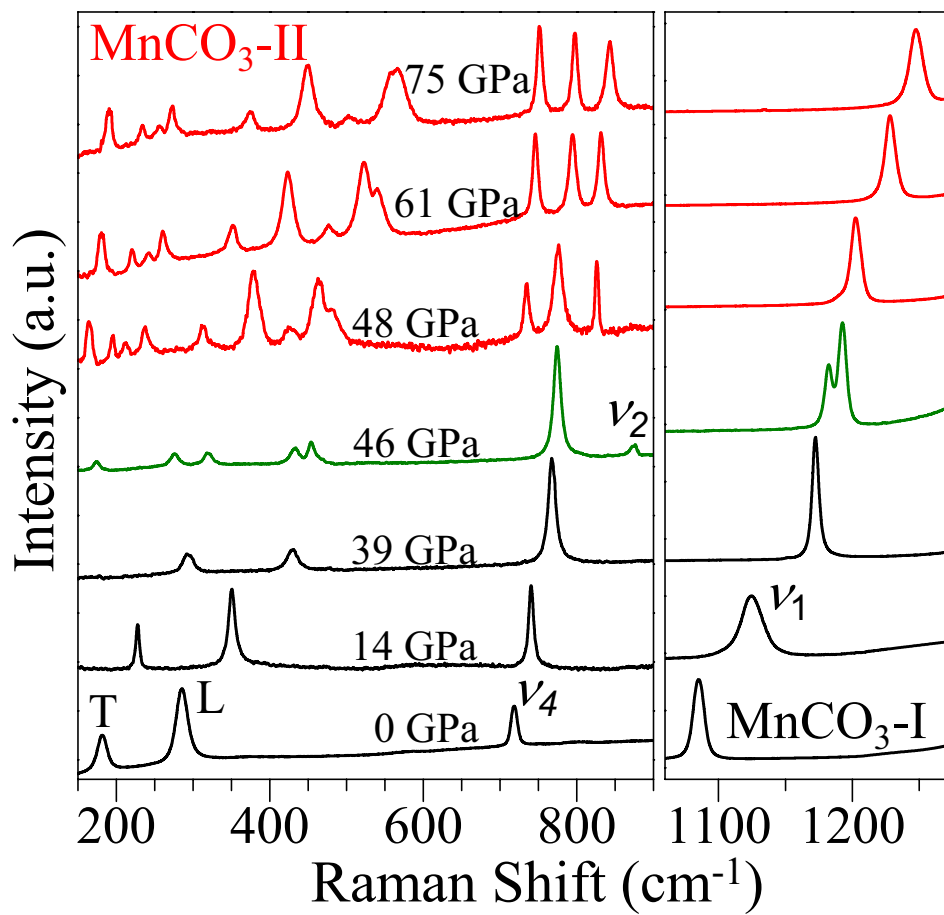


Figure 4

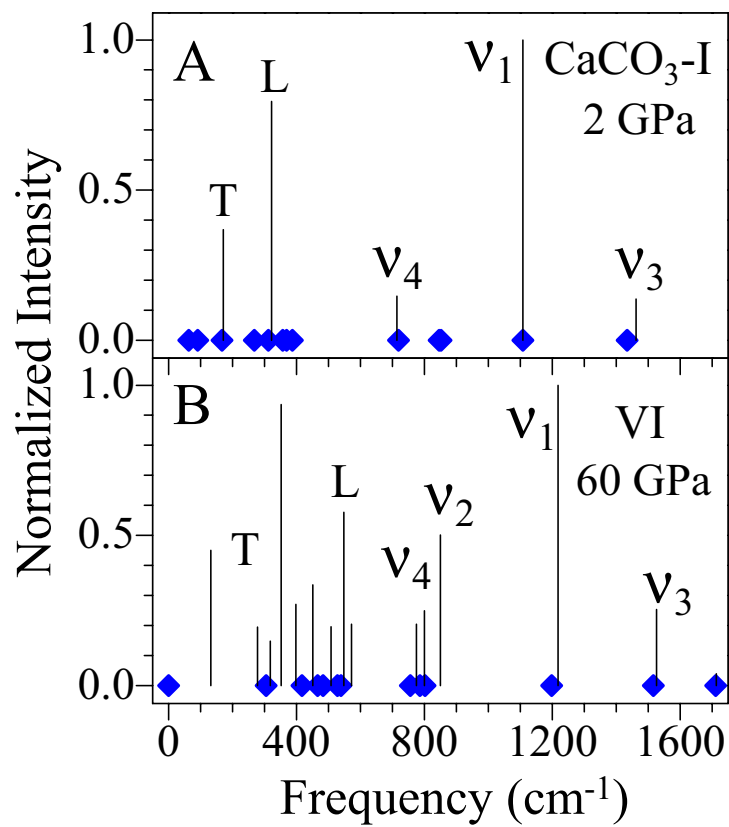


Figure 5

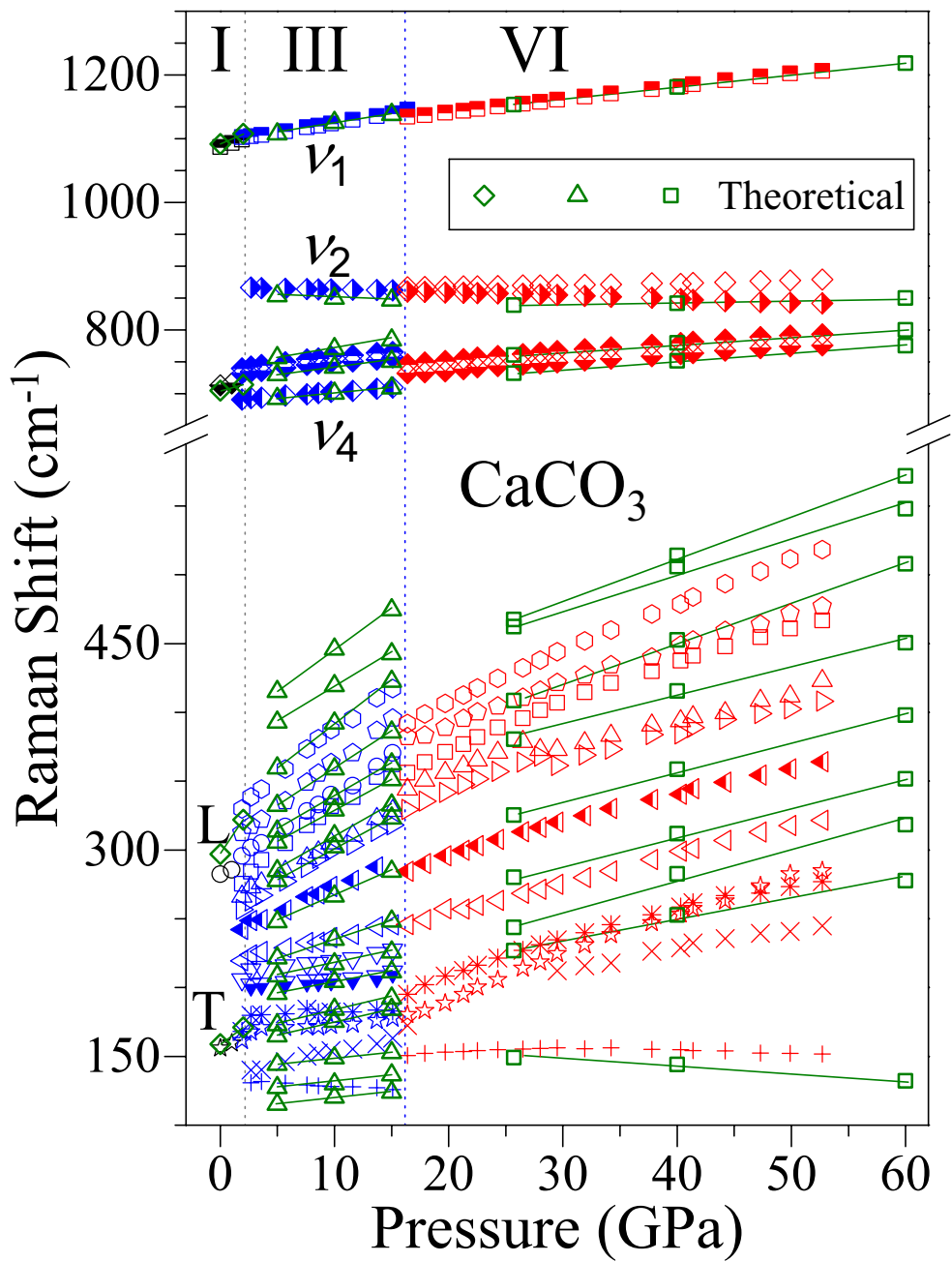


Figure 6

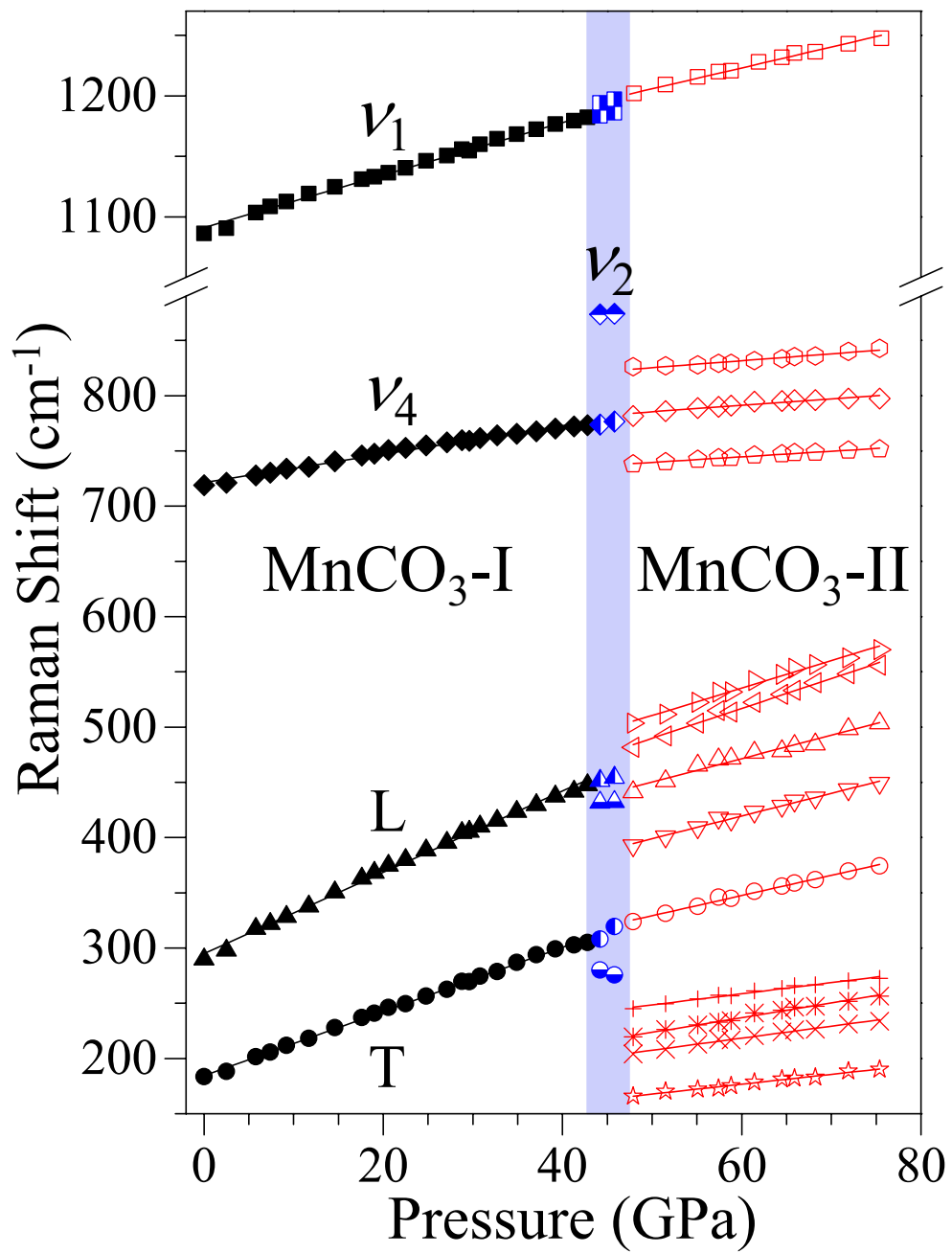


Figure 7

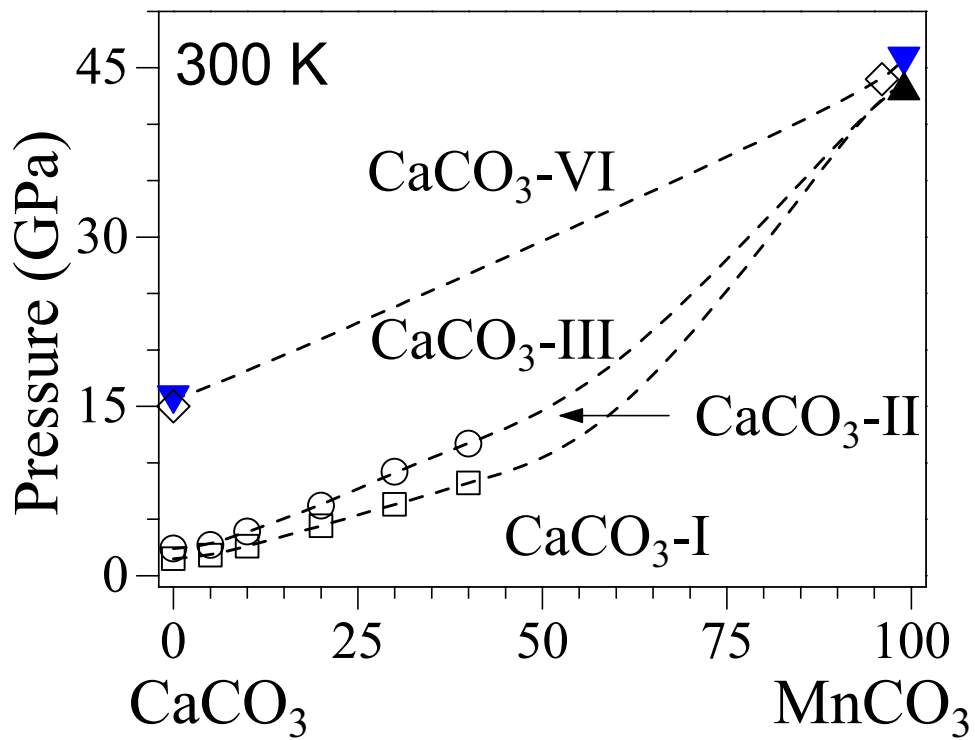


Figure 8

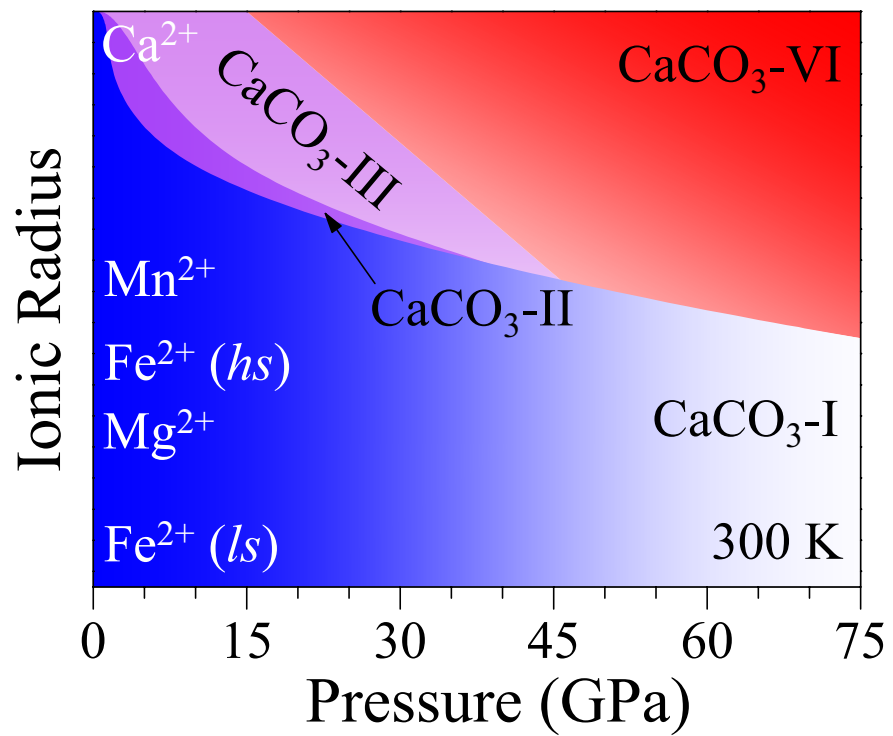


Figure 9